

## The Diel Cycle of Modeled and Measured Latent Heat Flux in a Colorado Subalpine Forest: Measurements vs CLM

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Community Land Model (CLM) Tutorial

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## Niwot Ridge AmeriFlux Tower (US-NR1)

(photo credit: Bill Bowman)

Influence of Precipitation on Warm-Season Fluxes

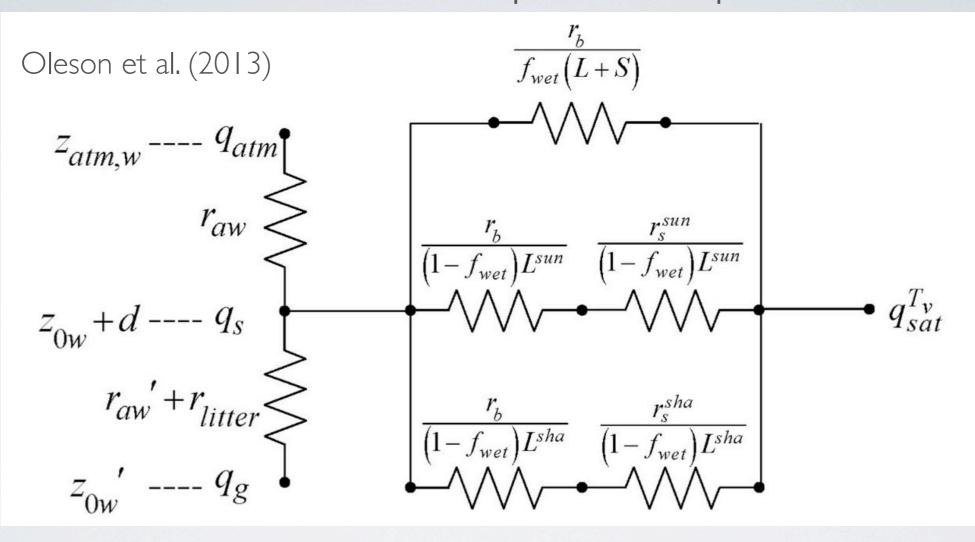
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  - In CLM, fluxes calculated using M-O Similarity Theory + additional resistance and turb trans coefficient to represent the subcanopy transport

CLM Water Vapor Transport



The aerodynamic resistances to heat (moisture) transfer between the ground at

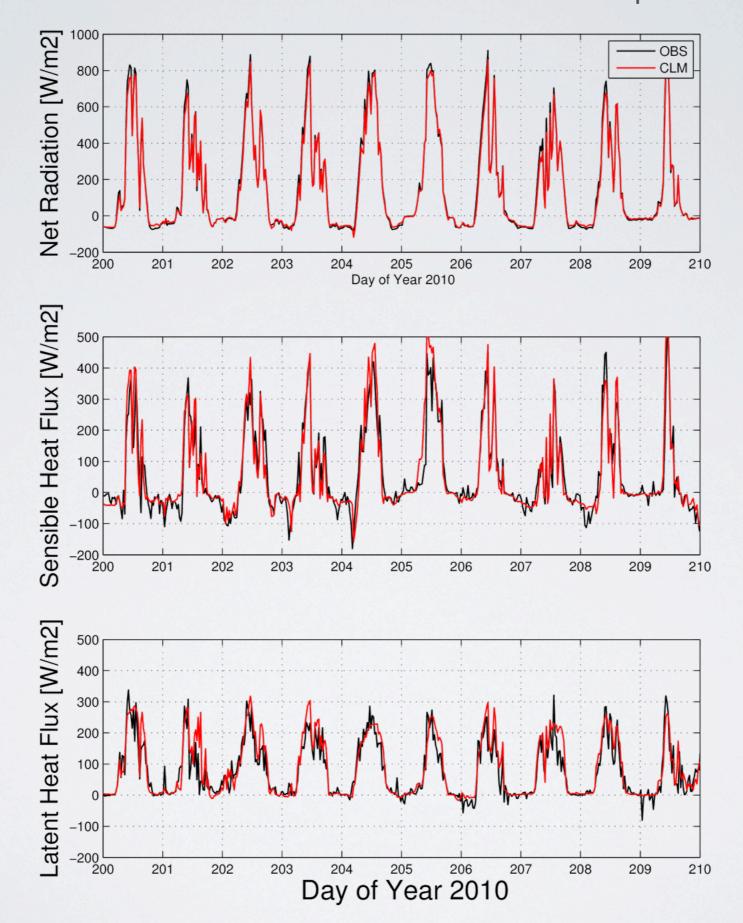
height  $z_{0h}'(z_{0w}')$  and the canopy air at height  $z_{0h} + d(z_{0w} + d)$  are

$$r_{ah}' = r_{aw}' = \frac{1}{C_s U_{av}}$$
(5.113)

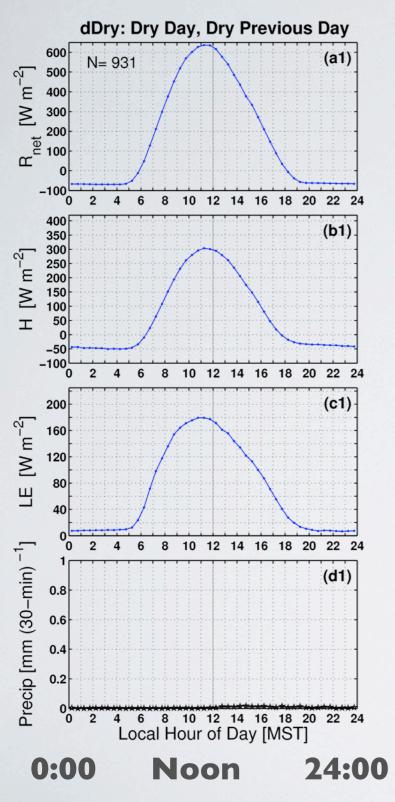
where

$$U_{av} = V_a \sqrt{\frac{1}{r_{am}V_a}} = u_*$$
 (5.114)

#### CLM vs OBS Time Series Comparison

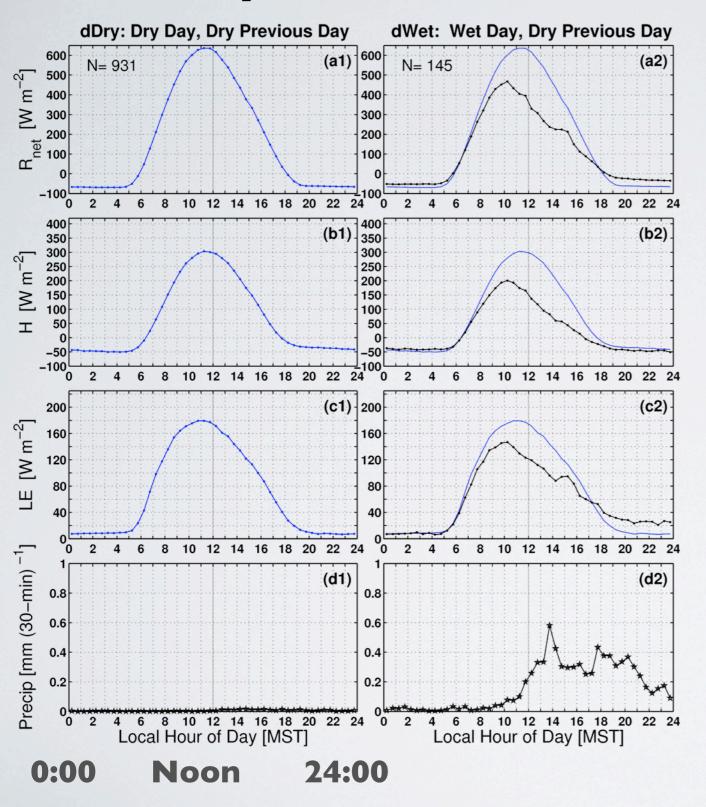


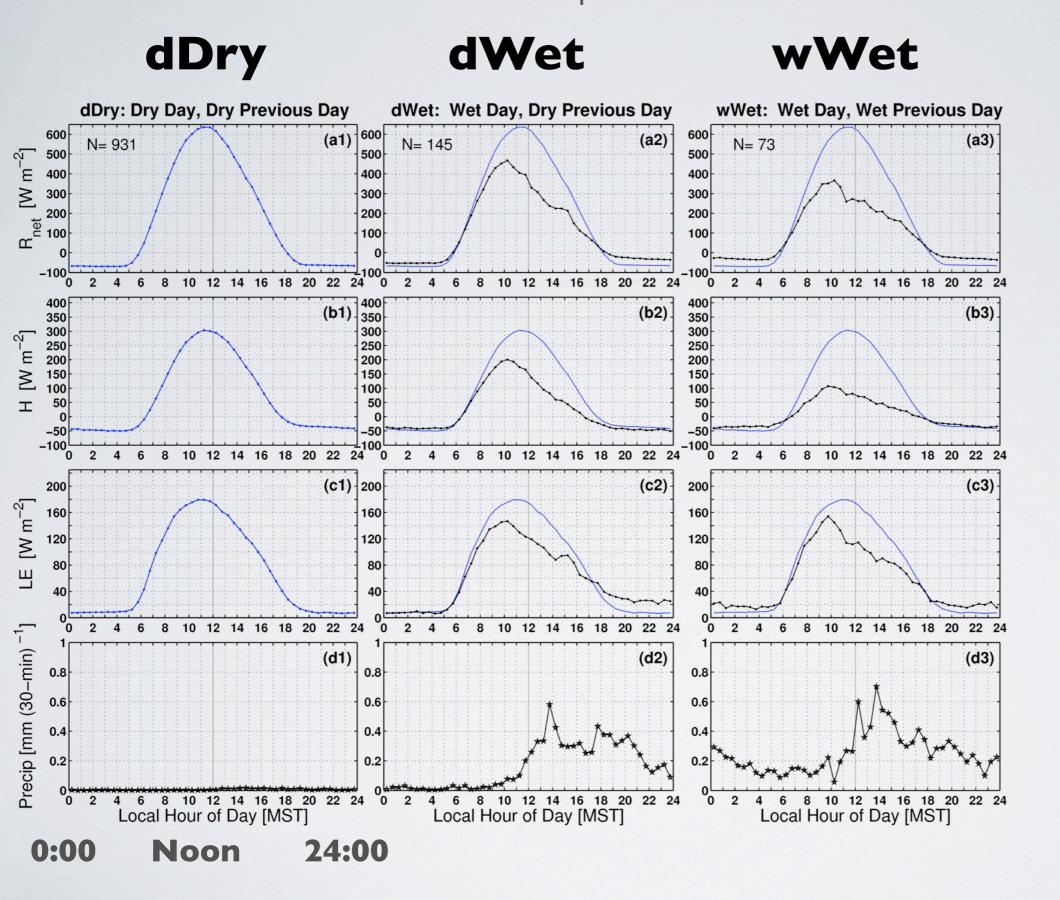
#### dDry

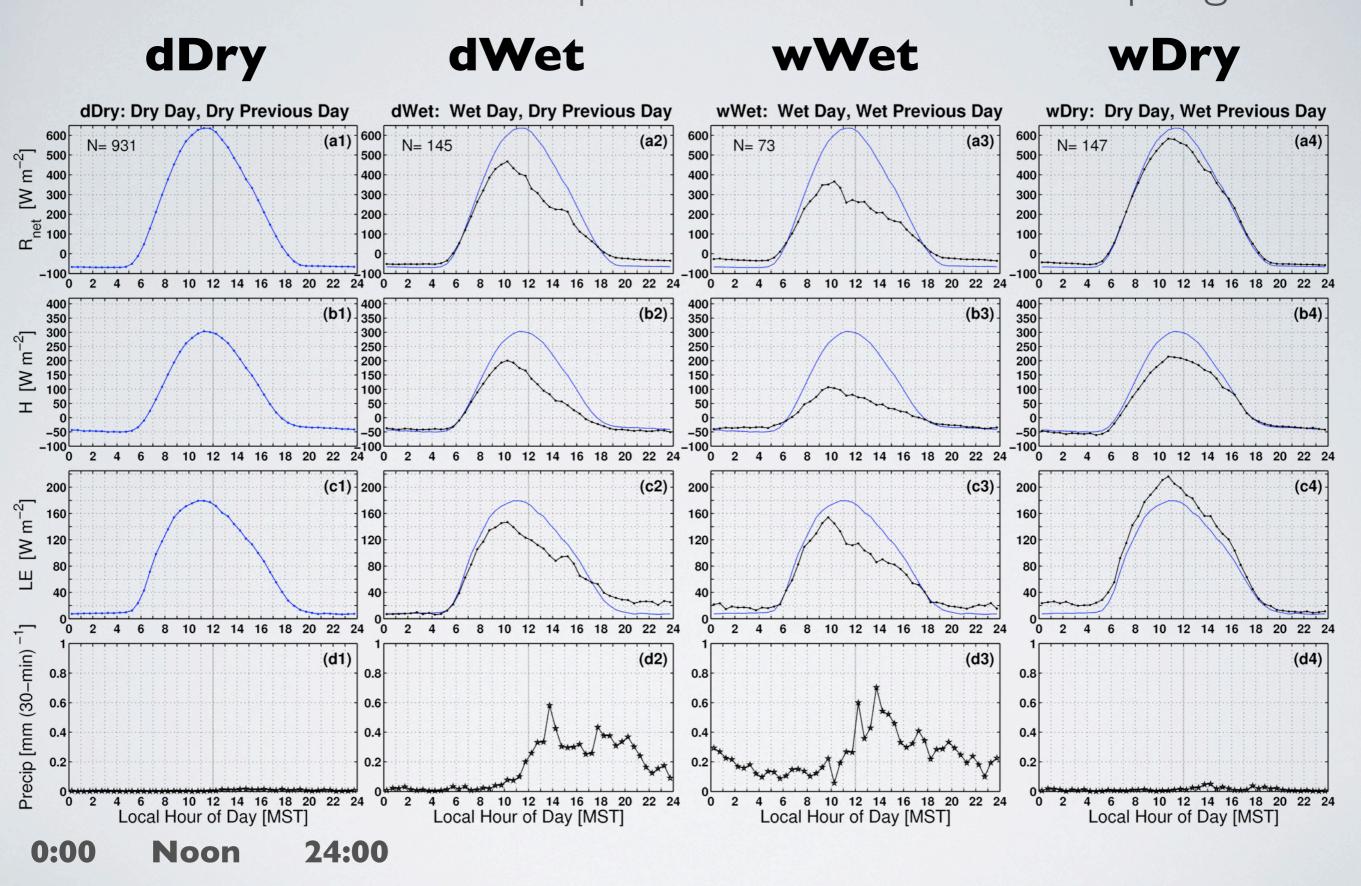


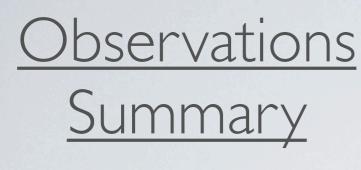
dDry

dWet

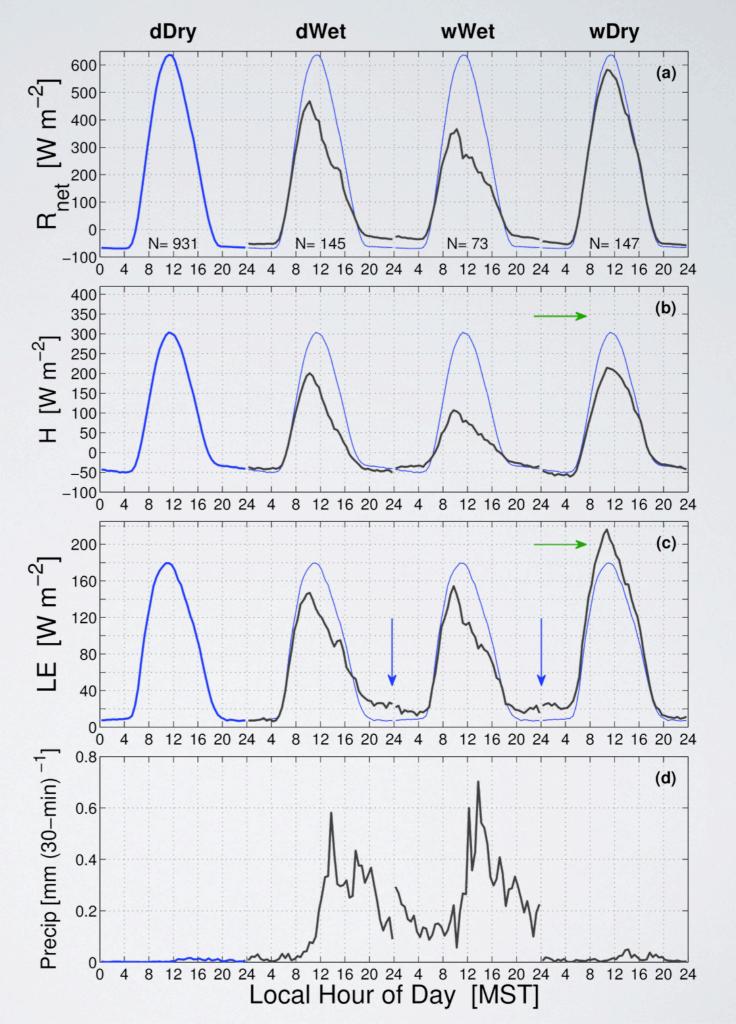


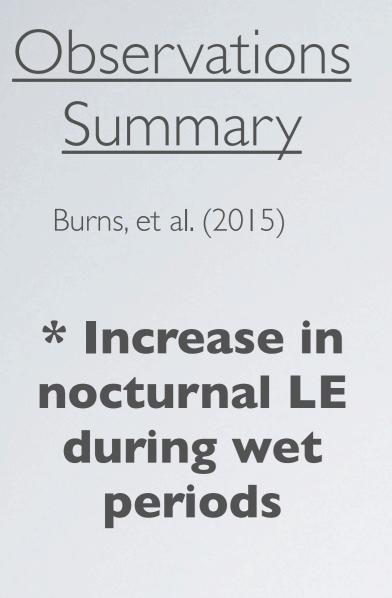


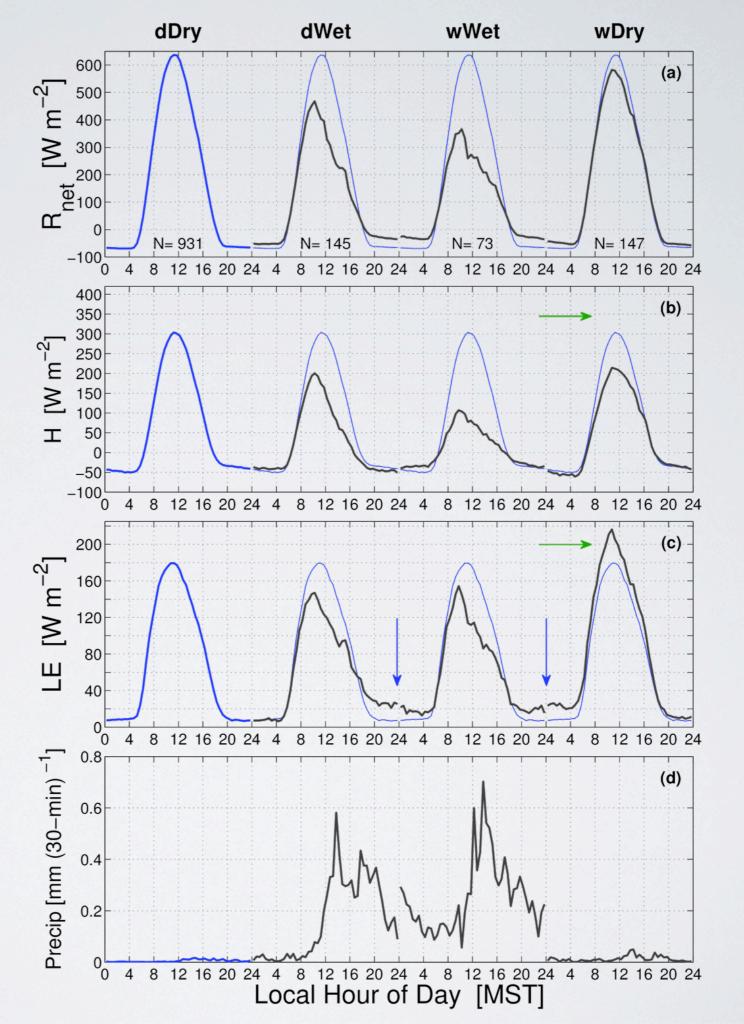


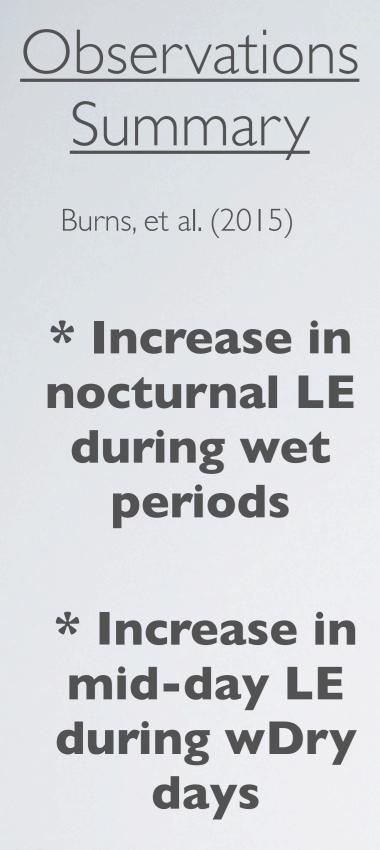


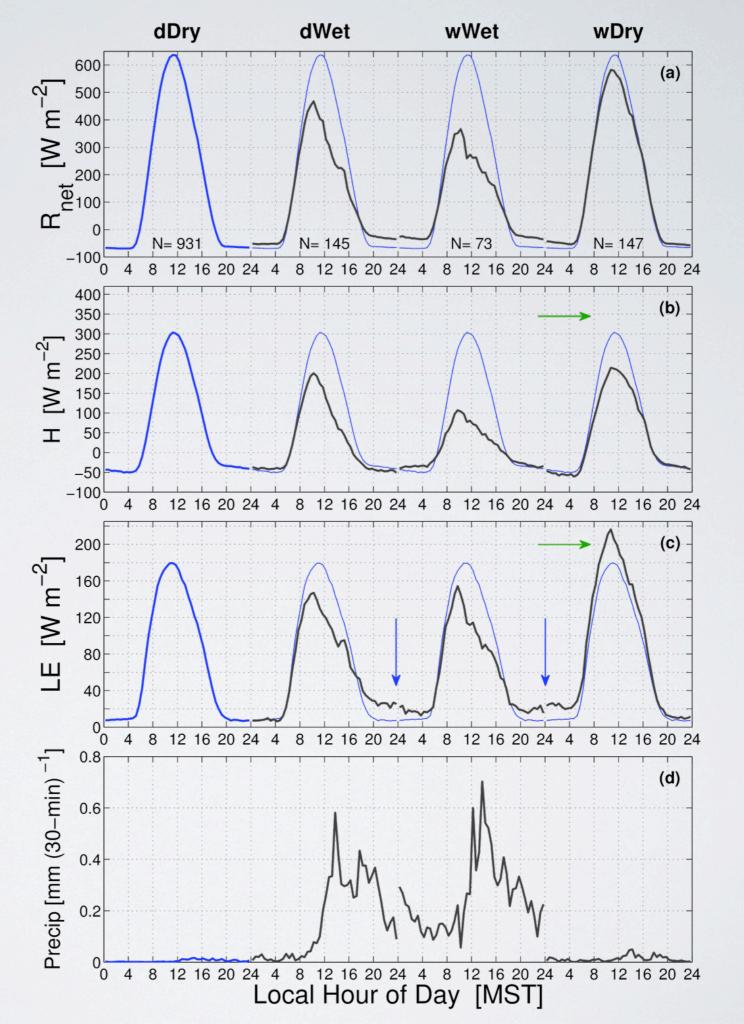
Burns, et al. (2015)



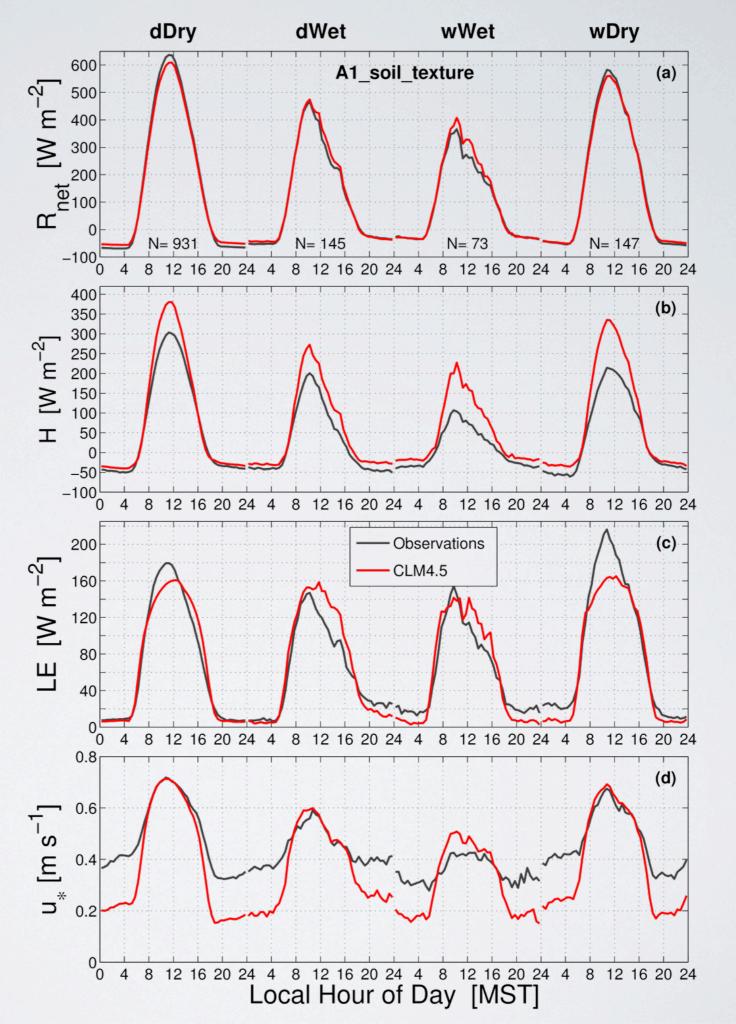


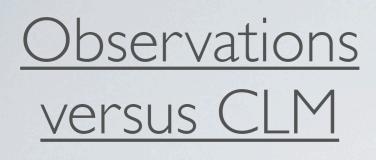




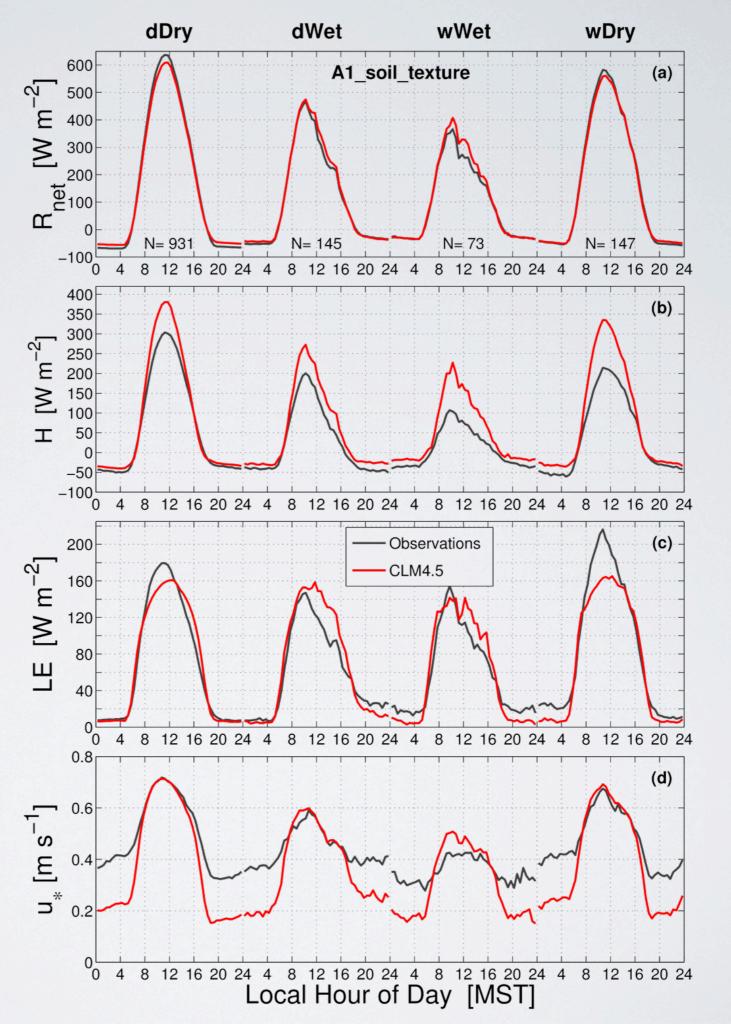


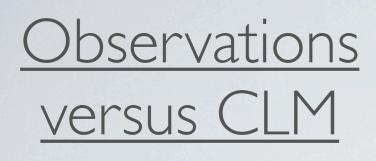






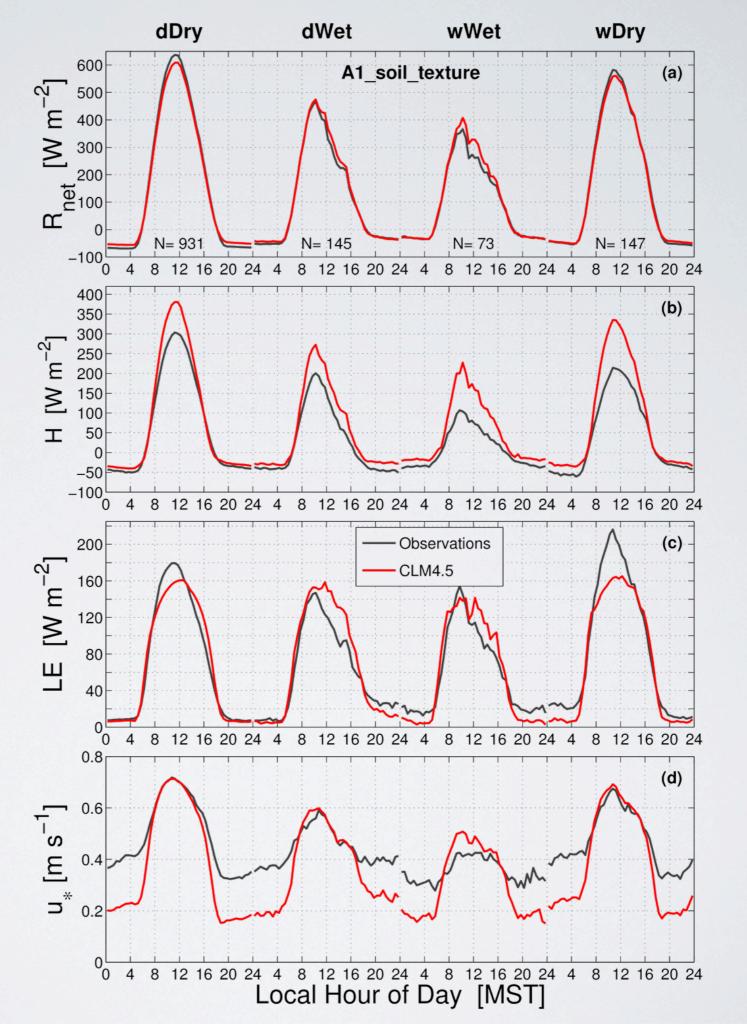
#### \* CLM LE at night too low (in wet conditions)

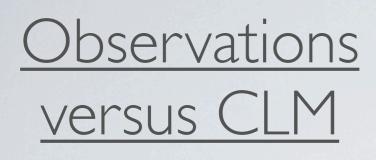




\* CLM LE at night too low (in wet conditions)

\* No increase in CLM LE on wDry days

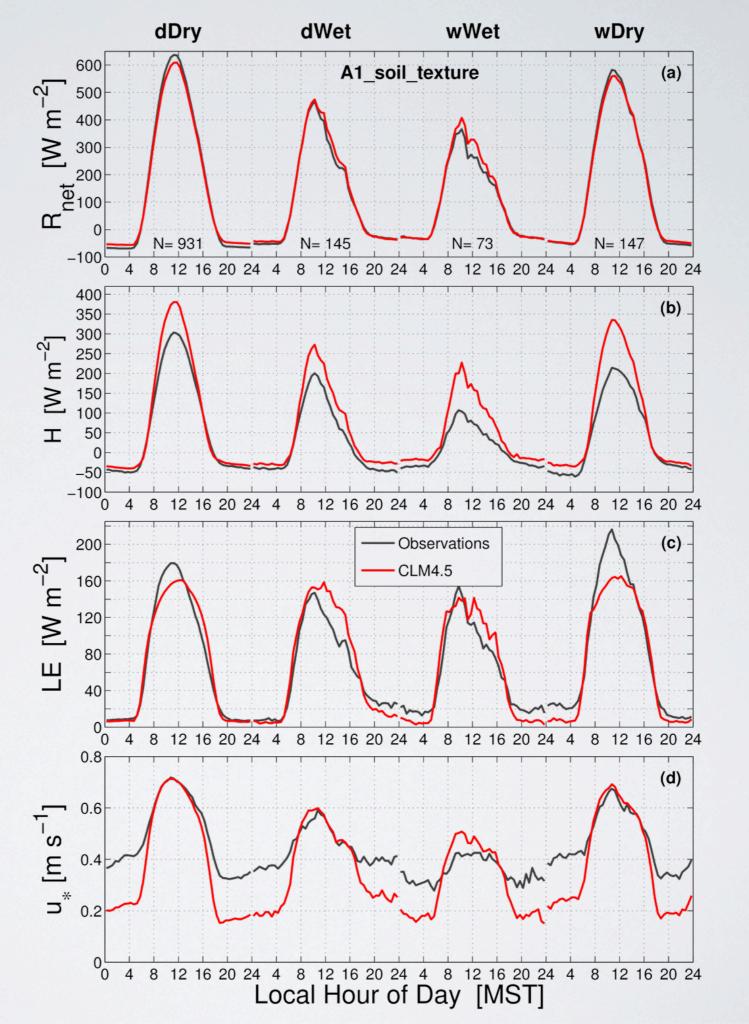




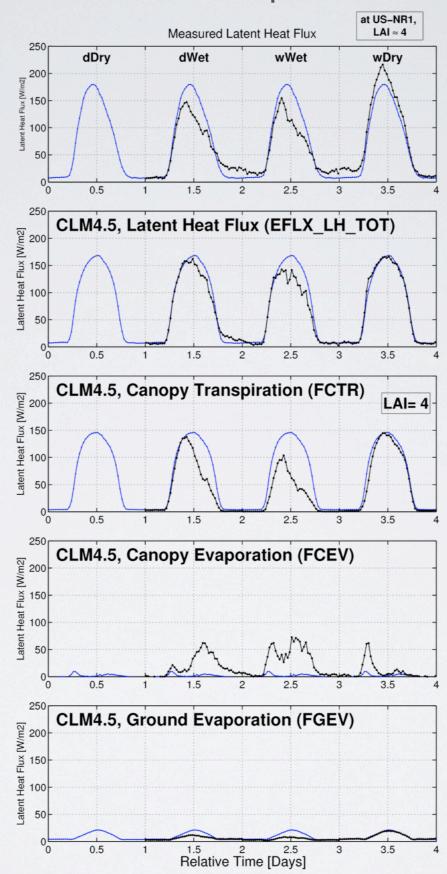
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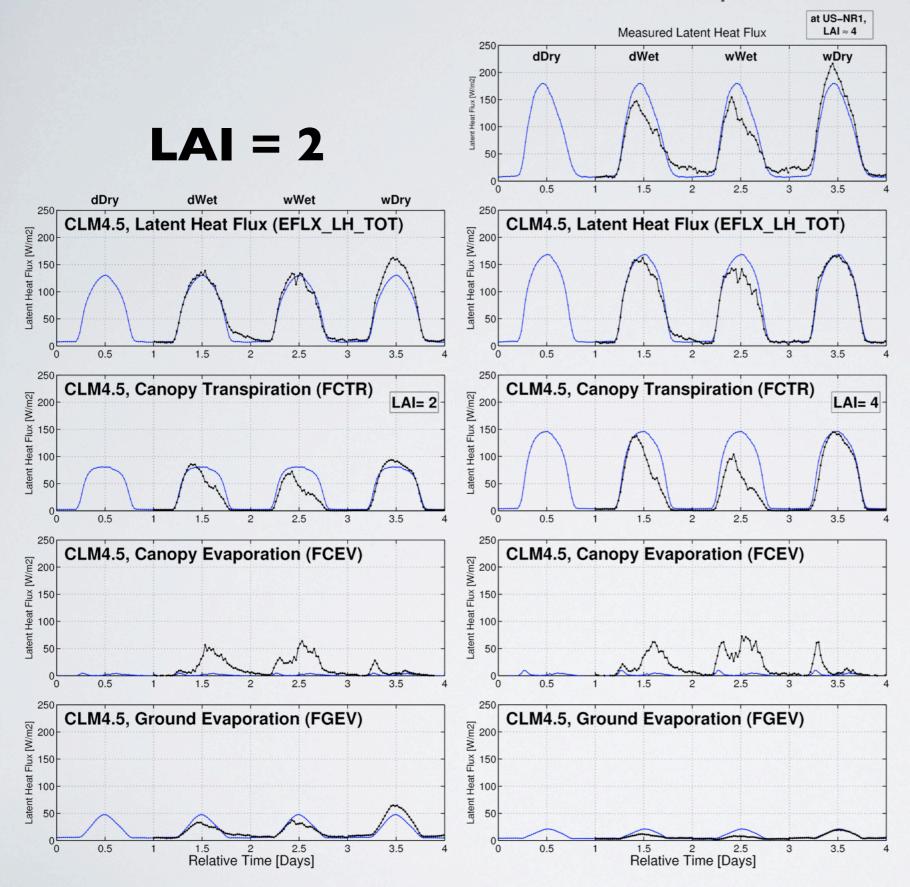
\* ustar at night is too low (e.g., Bonan, Patton)



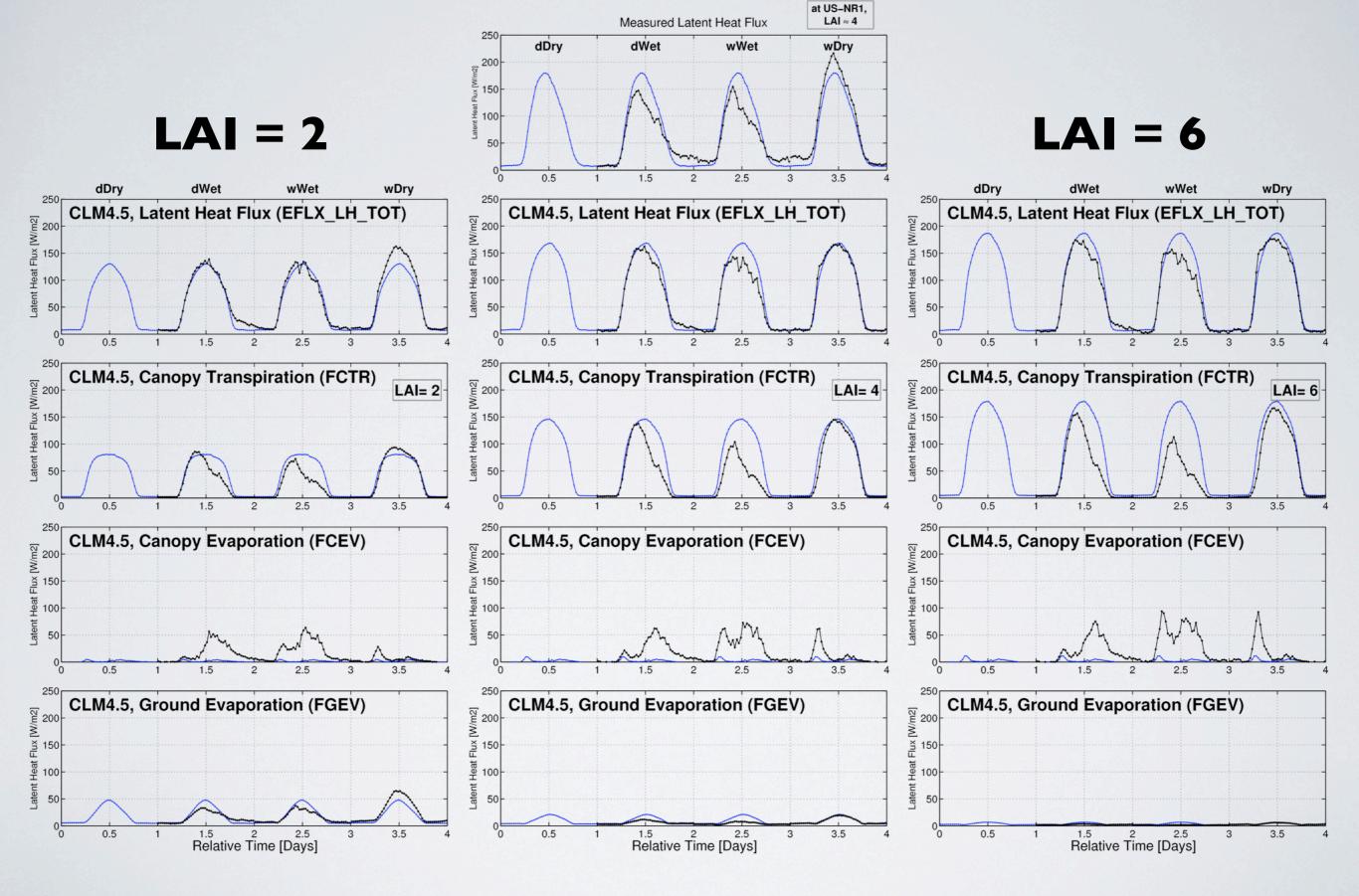
## Latent Heat Flux Components of CLM4.5



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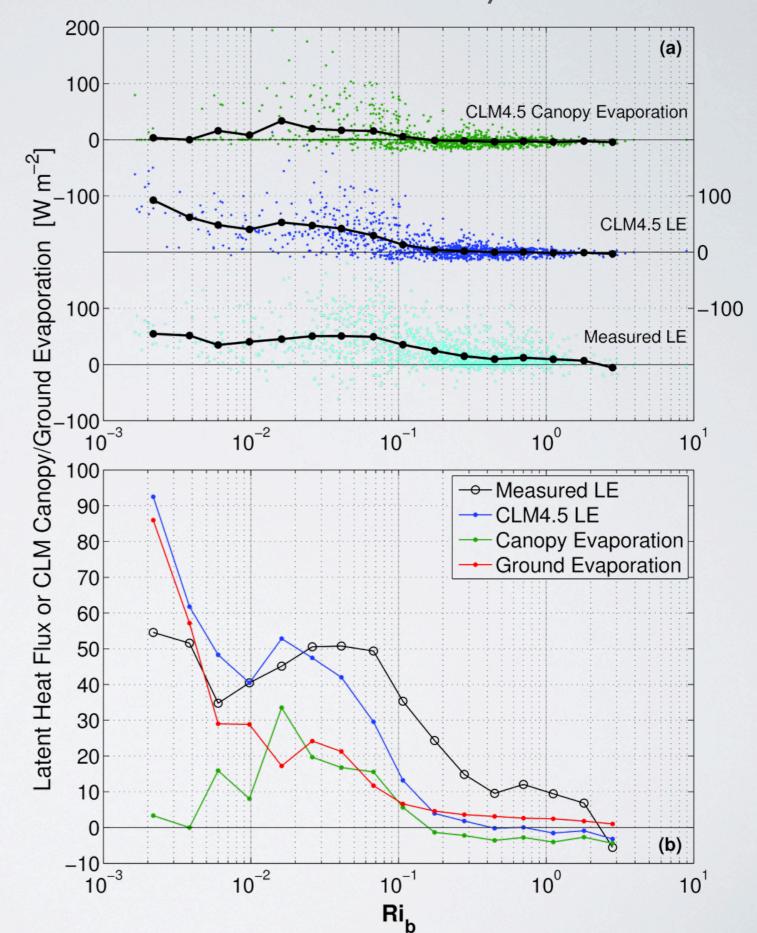
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## Latent Heat Flux vs Stability

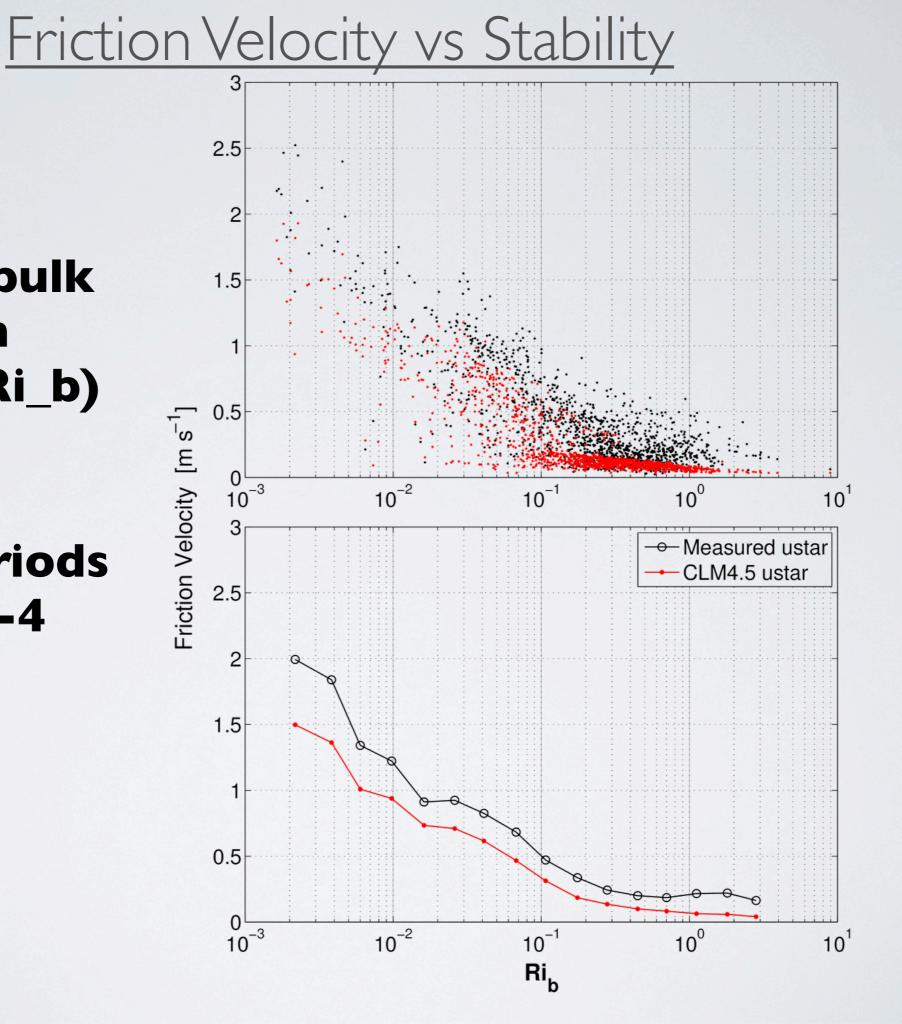
\* Latent Heat Flux vs bulk Richardson Number (Ri\_b)

\* wDry periods between 0-4 LST



## \* ustar vs bulk Richardson Number (Ri\_b)

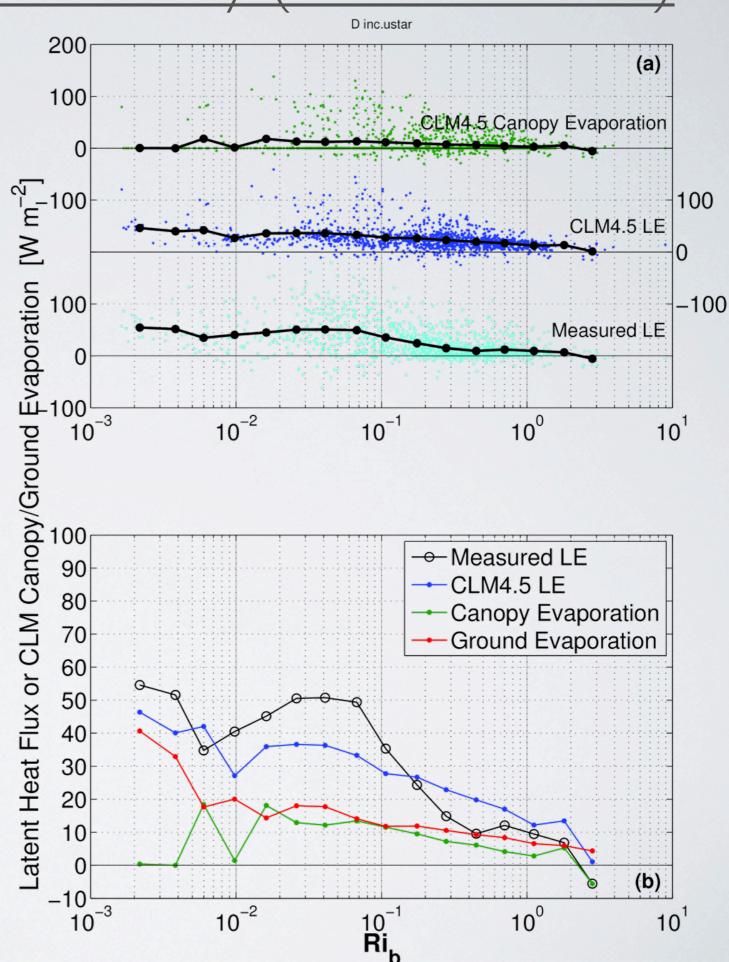
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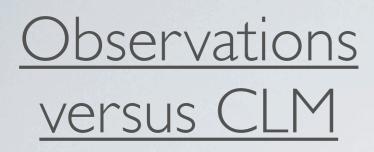


# Latent Heat Flux vs Stability (w/ u\* increase)

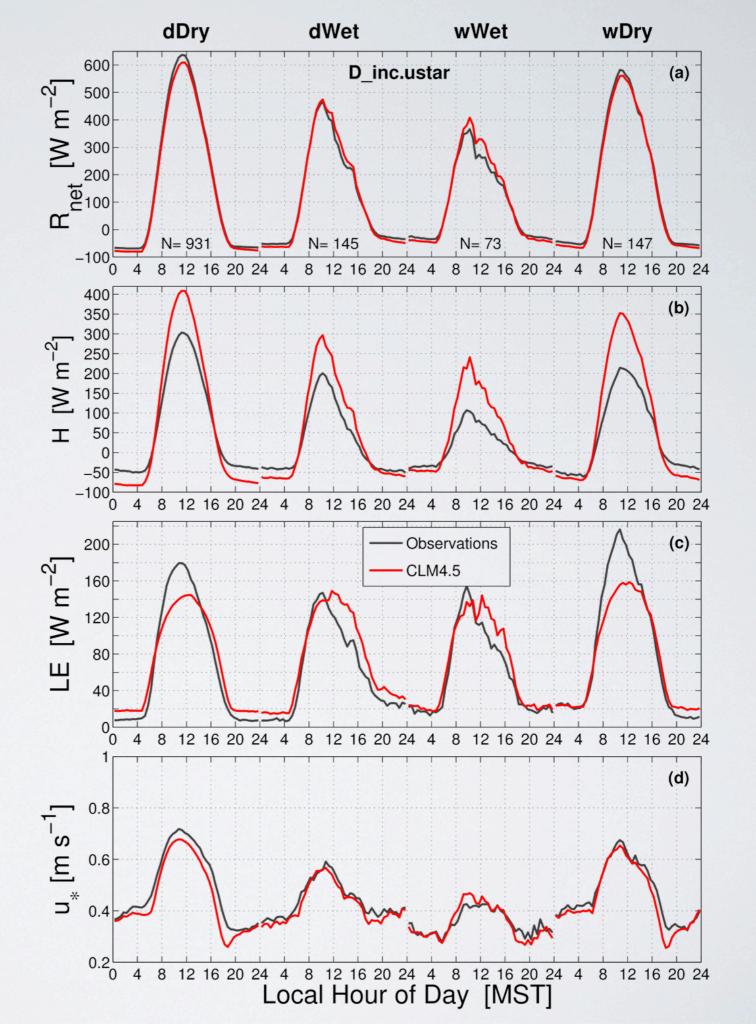
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#### \* CLM results with an ad-hoc increase to u\* of 0.2 m/s



#### **Conclusions and Future Plans**

- In partially-wet (wDry) conditions, observations show an increase in mid-day latent heat flux that is not reproduced in CLM LE (for  $LAI \approx 4$ ).
- At night: (1) CLM ground and canopy evaporation are smaller than observed LE, (2) CLM friction velocity is smaller than the observations, (3) Suspect that low  $u_*$  is at least the partial cause of too small CLM soil and/or canopy evaporation.
- Ad-hoc increase to CLM ustar increases nocturnal LE, but for both wet and dry conditions.
- Are results site specific? Repeat this analysis at another "less complicated" flux site (looked at Howland and Metolius)?
- M-O Similarity inappropriate in strongly stable conditions—what to use instead?
- What results would be achieved using a continuous, multi-level canopy model (e.g., ACASA)?

# COMMENTS/QUESTIONS?

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# Latent Heat Flux Comparison: CLM vs Obs

