

Permafrost-climate feedbacks in CESM/CLM

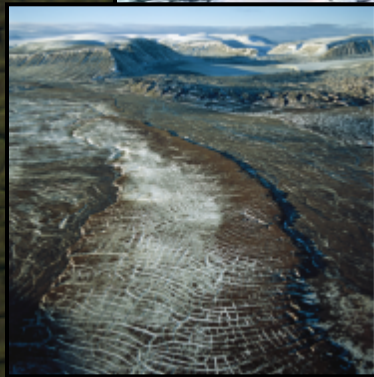
David Lawrence

**Andrew Slater², Sean Swenson¹,
Charlie Koven³, Bill Riley³,
Zack Subin³, Hanna Lee¹ and
the CESM LMWG**

¹NCAR Earth System Lab, Boulder, CO

²NSIDC, Boulder, CO

³LBL, Berkeley, CA



NCAR is sponsored by the National Science Foundation



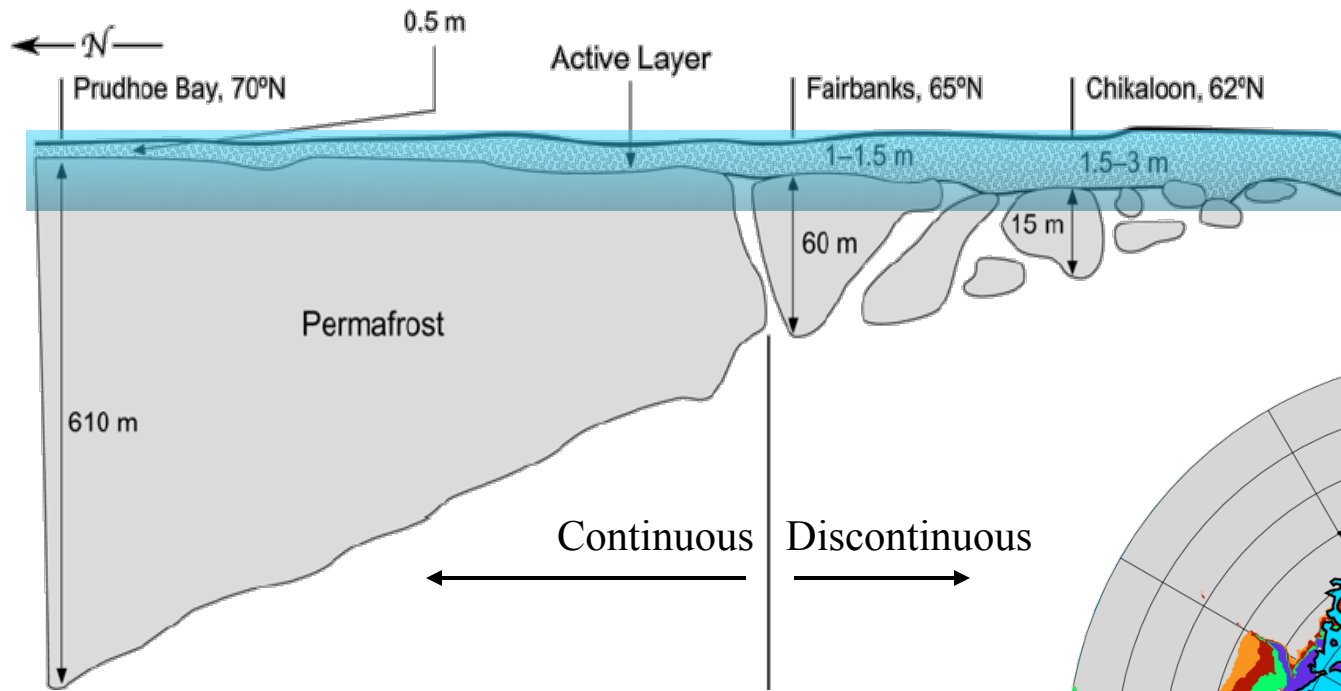
A photograph of two men standing in a vast, flat, green field under a clear blue sky. Both men are wearing backpacks and have their hands to their chins in a thinking pose. The man on the left is wearing a grey long-sleeved shirt, khaki pants, and a dark baseball cap. The man on the right is wearing a light blue long-sleeved shirt, khaki pants, a purple baseball cap, and sunglasses. A yellow and white plaid shirt is tied around his waist. In the top right corner, there is a white thought bubble containing the text 'What is permafrost?'.

What is permafrost?

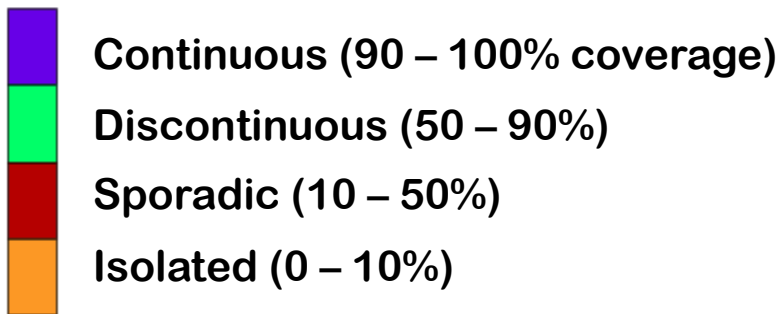
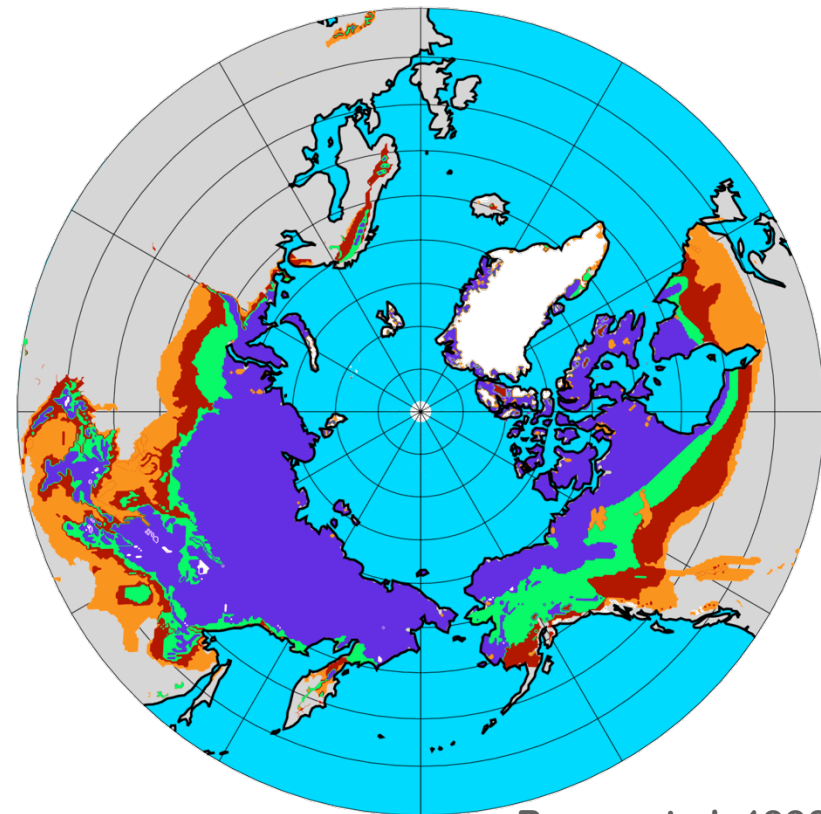
Definition: Soil or rock that remains below 0°C for two or more consecutive years



Global Permafrost Distribution

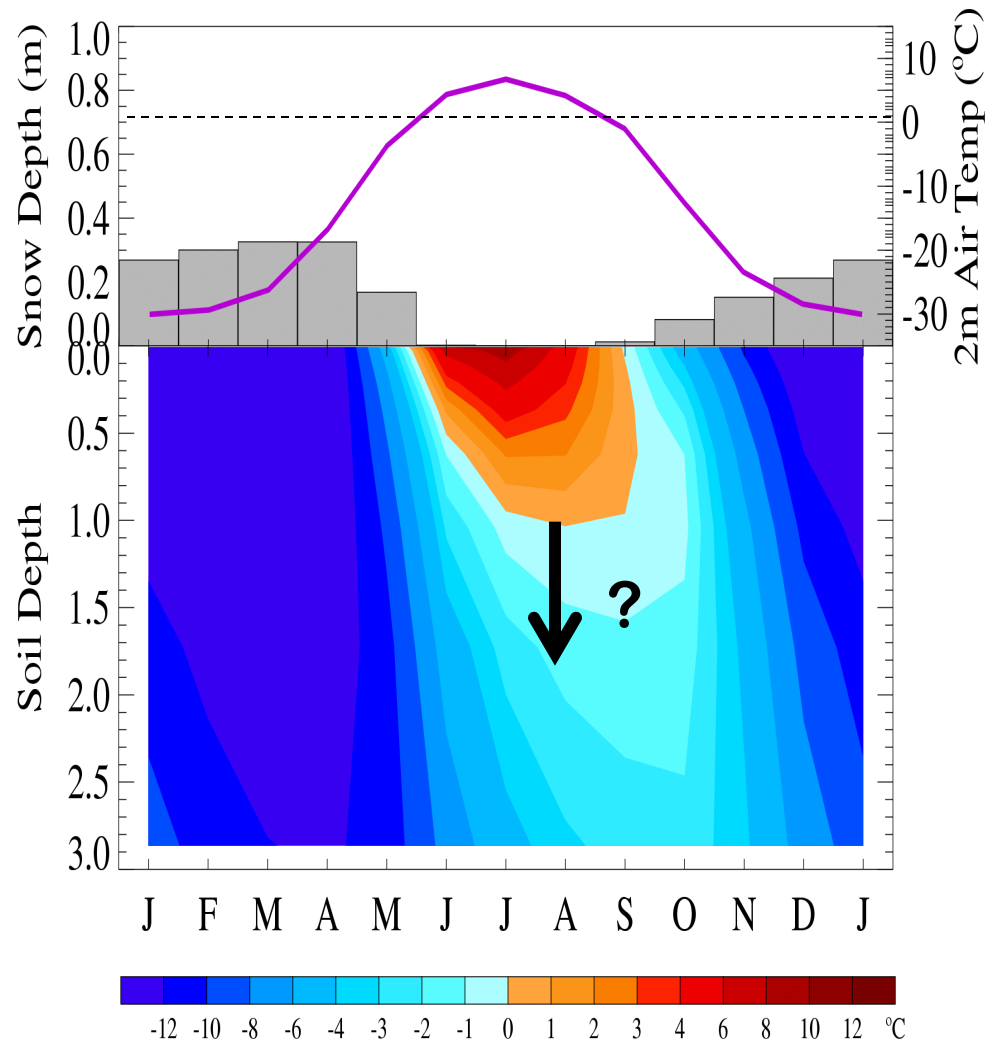


IPA Permafrost Distribution Map



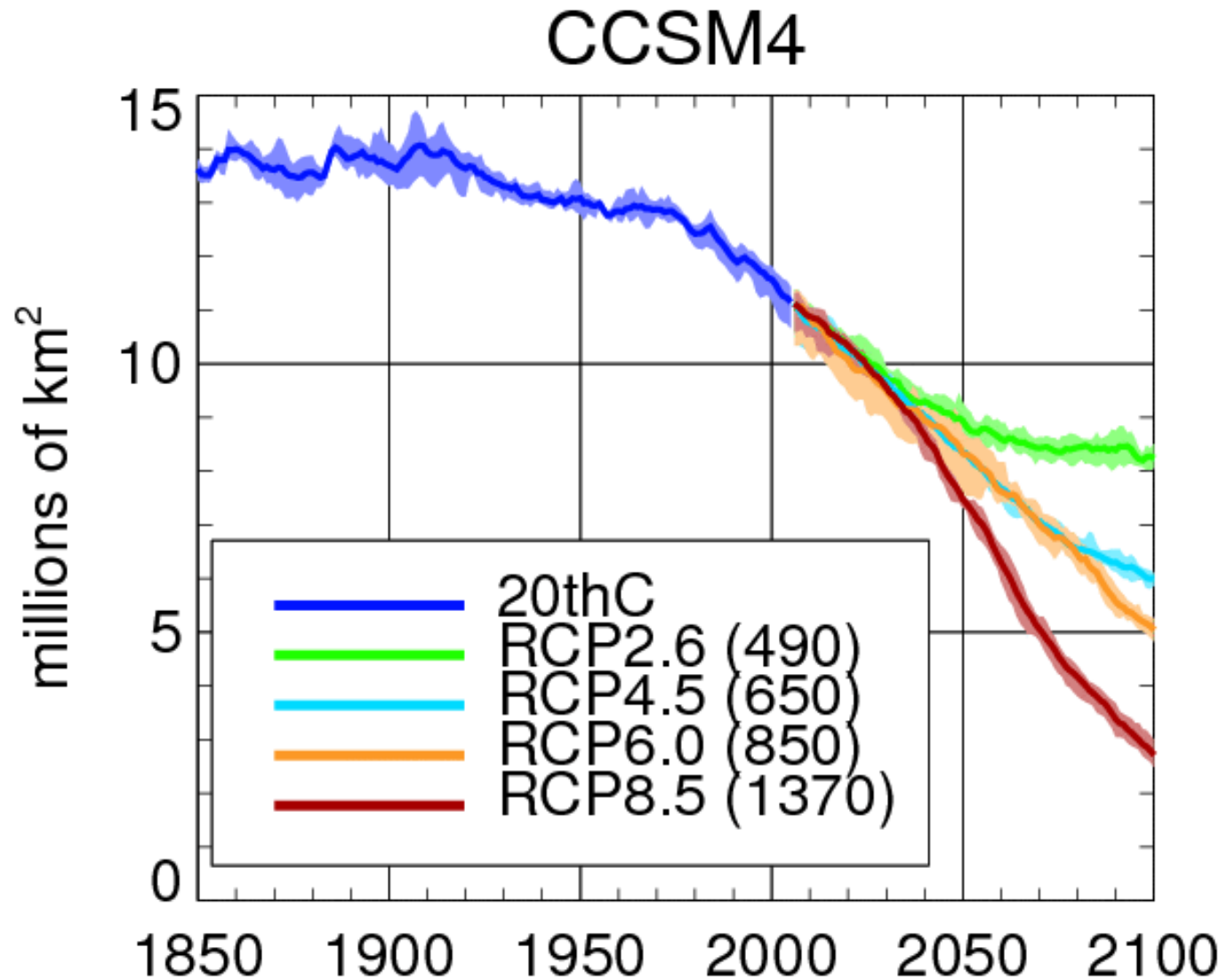


Active Layer Thickness (ALT)



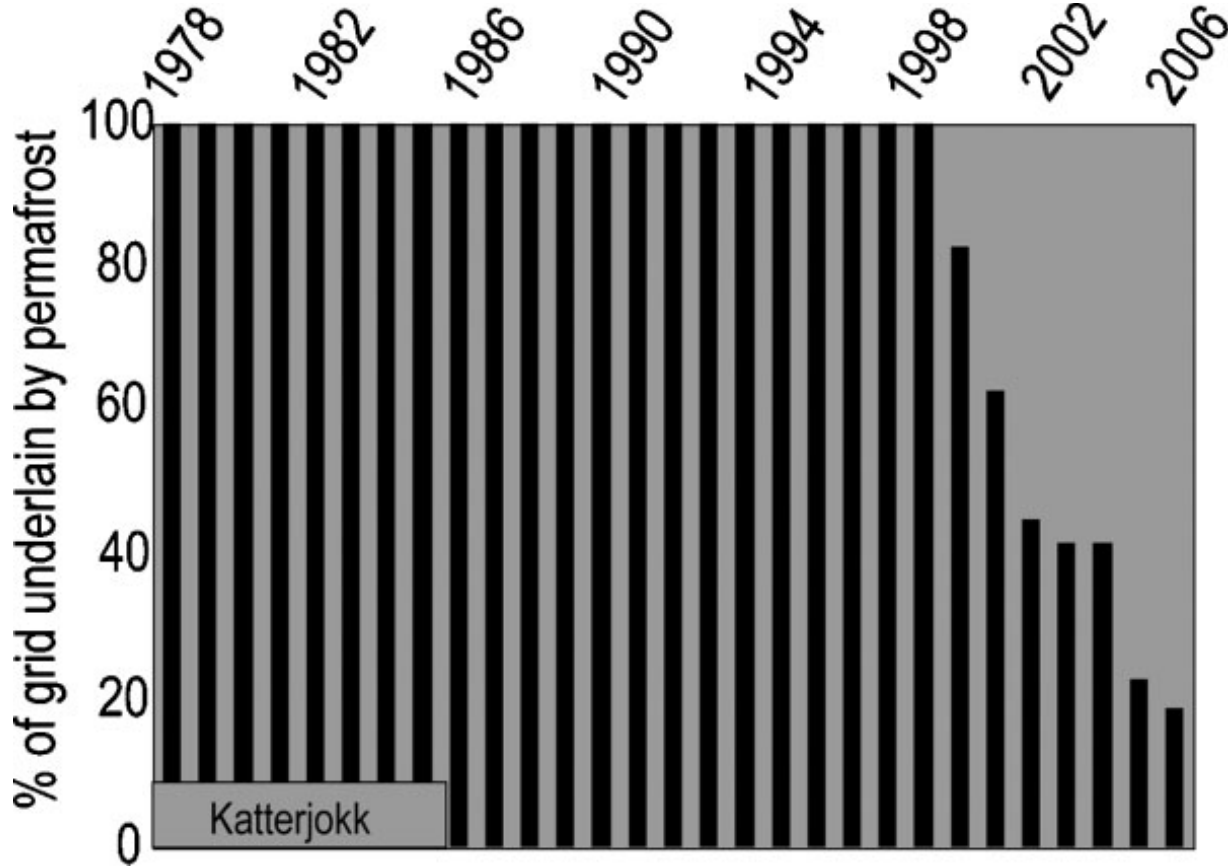


Projections of near-surface permafrost thaw





Observed rapid permafrost degradation

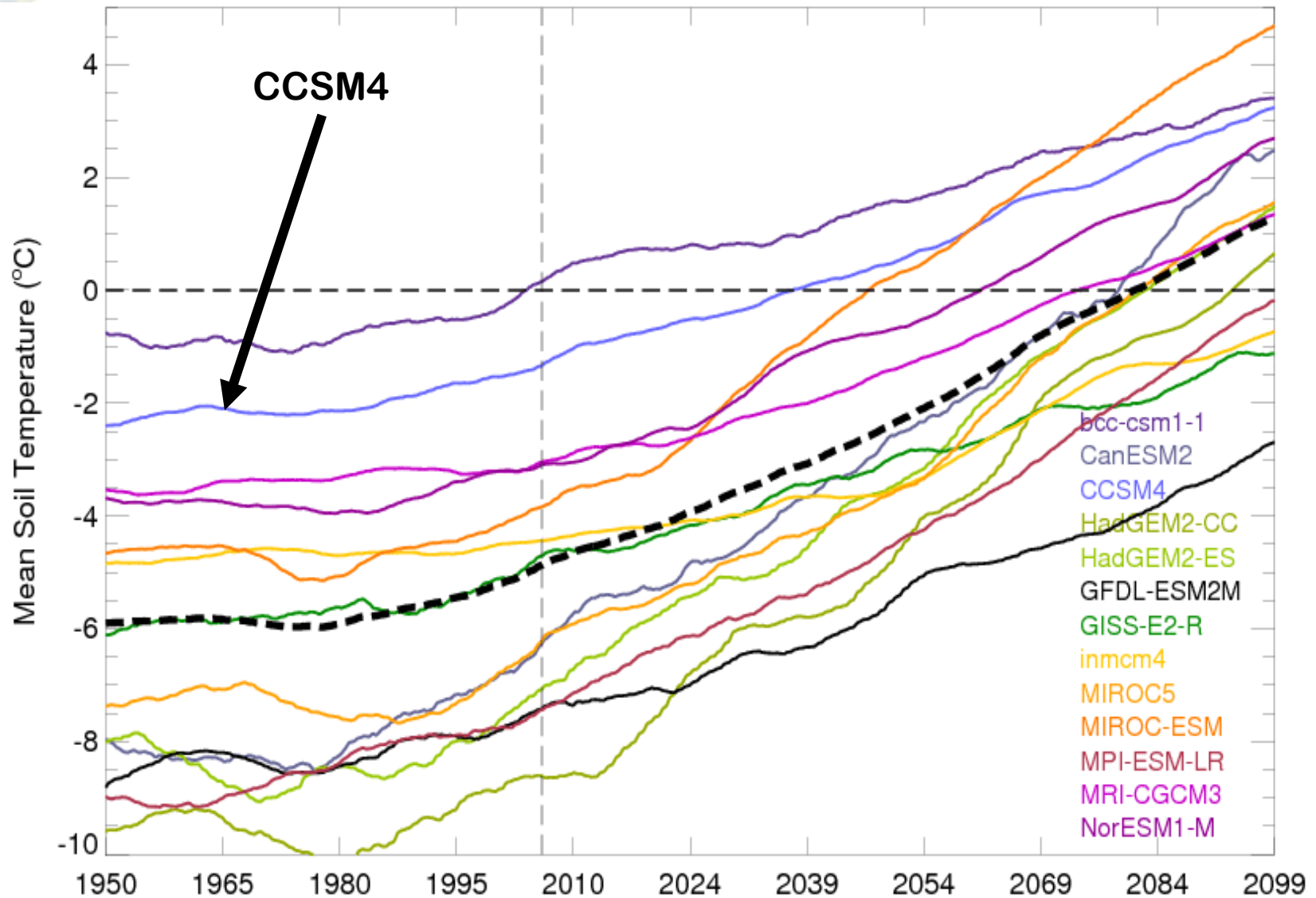


**IPY synthesis:
Widespread warming
and thawing**
(Romanovsky et al. 2010)

Akerman and Johansson, 2008

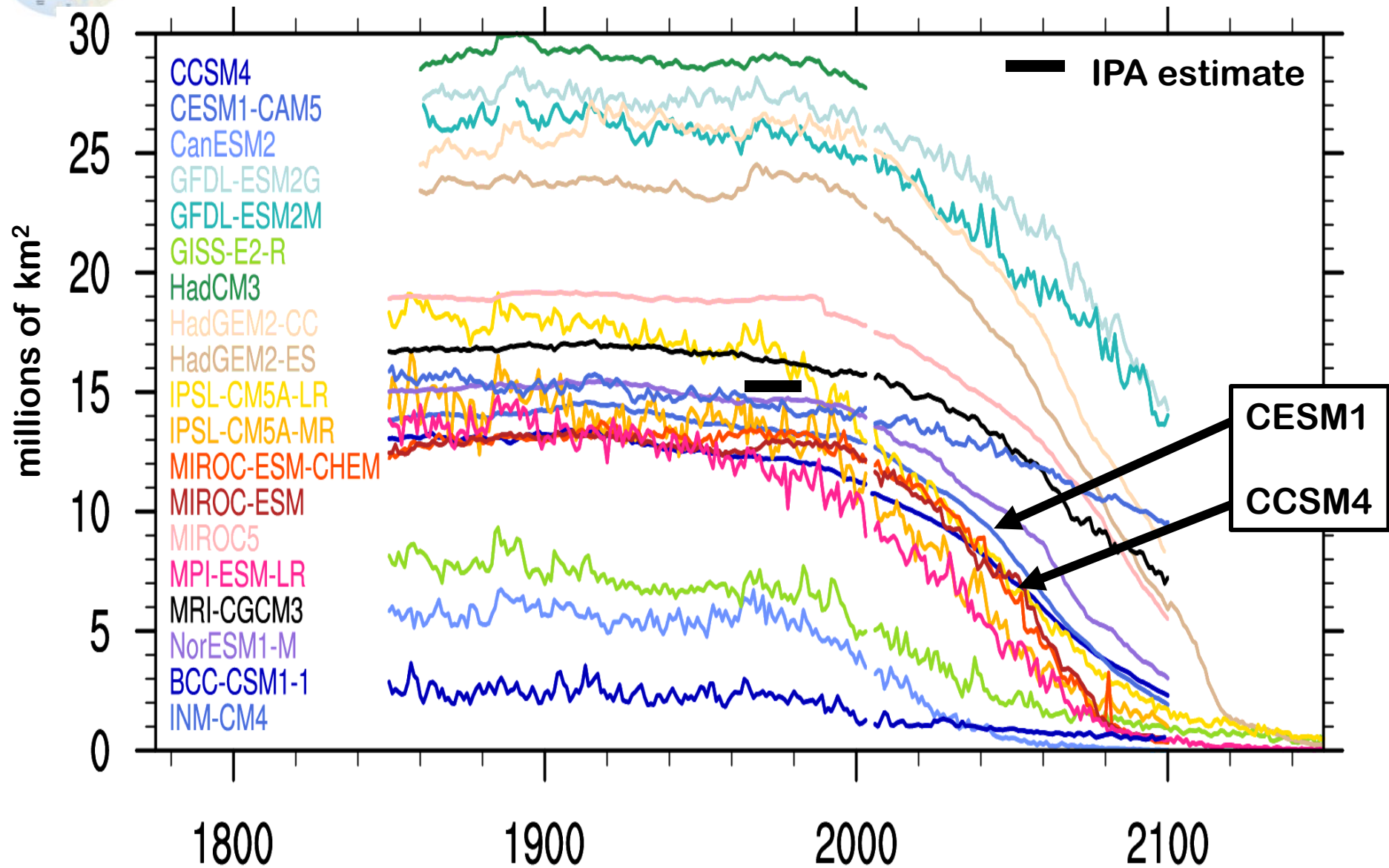


CMIP5 Models: Mean Soil Temperature across permafrost domain @ 3.3m (RCP 8.5)



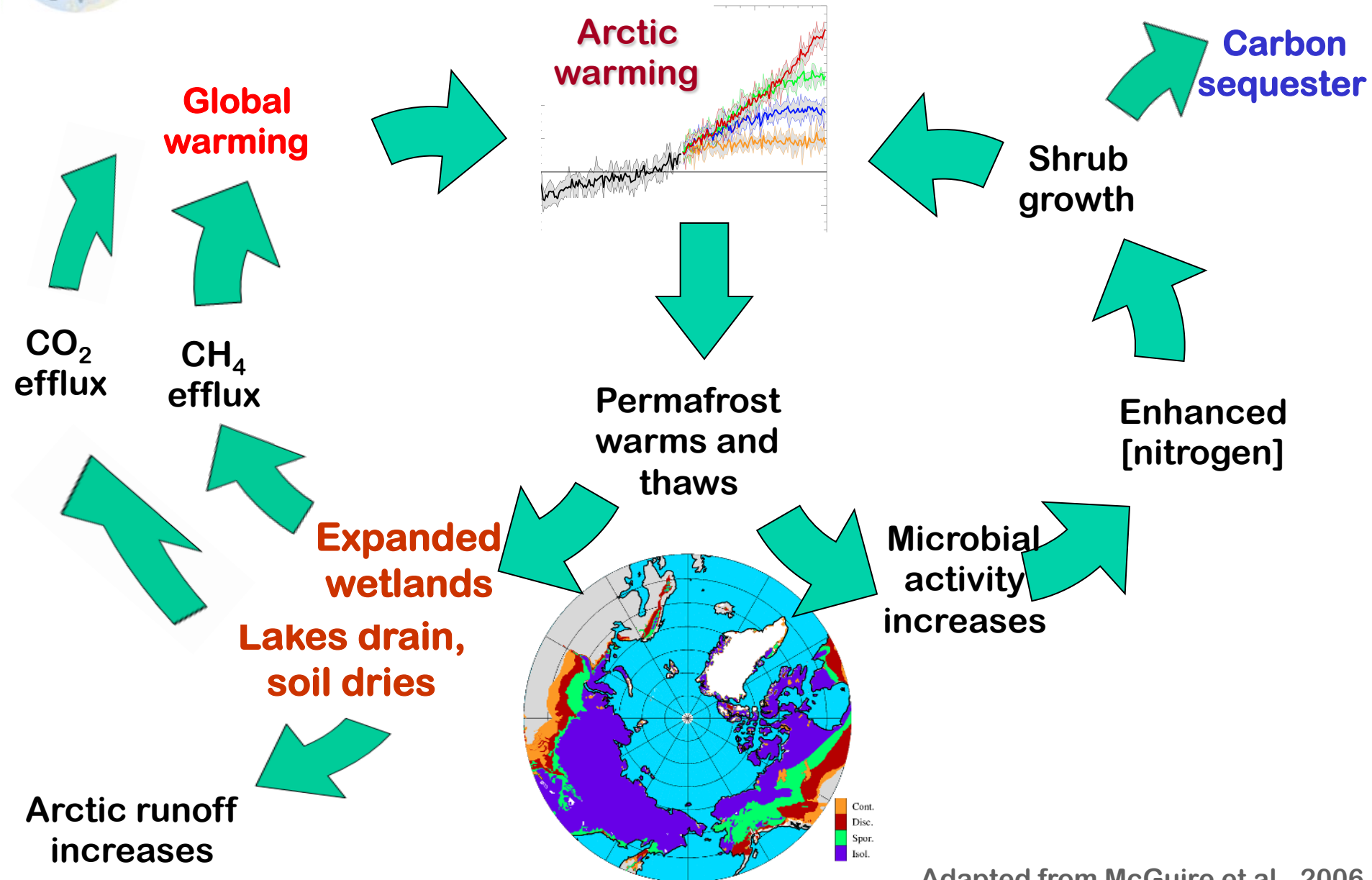


CMIP5 Models: Near-surface permafrost extent (RCP 8.5)



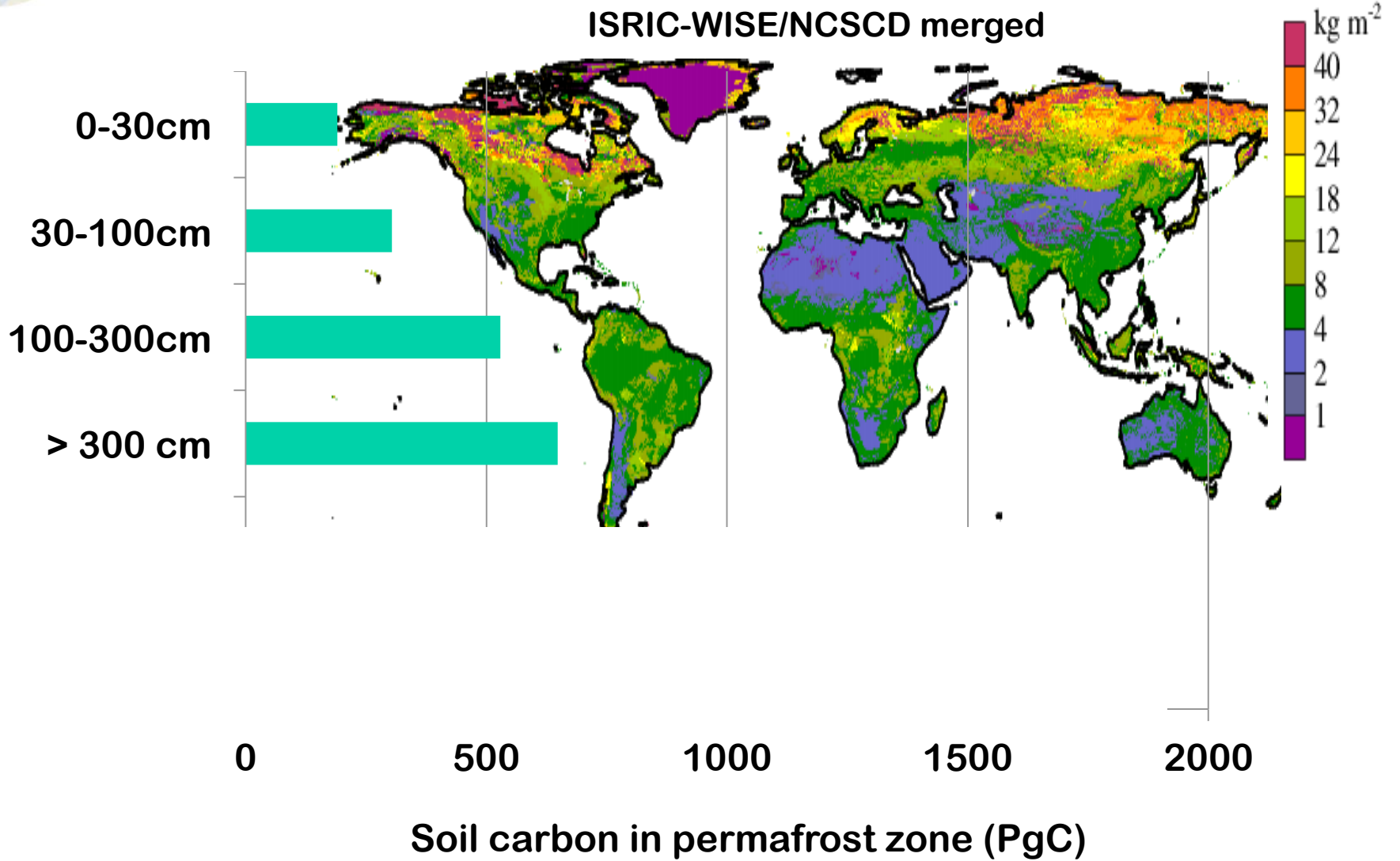


Potential Arctic terrestrial climate-change feedbacks



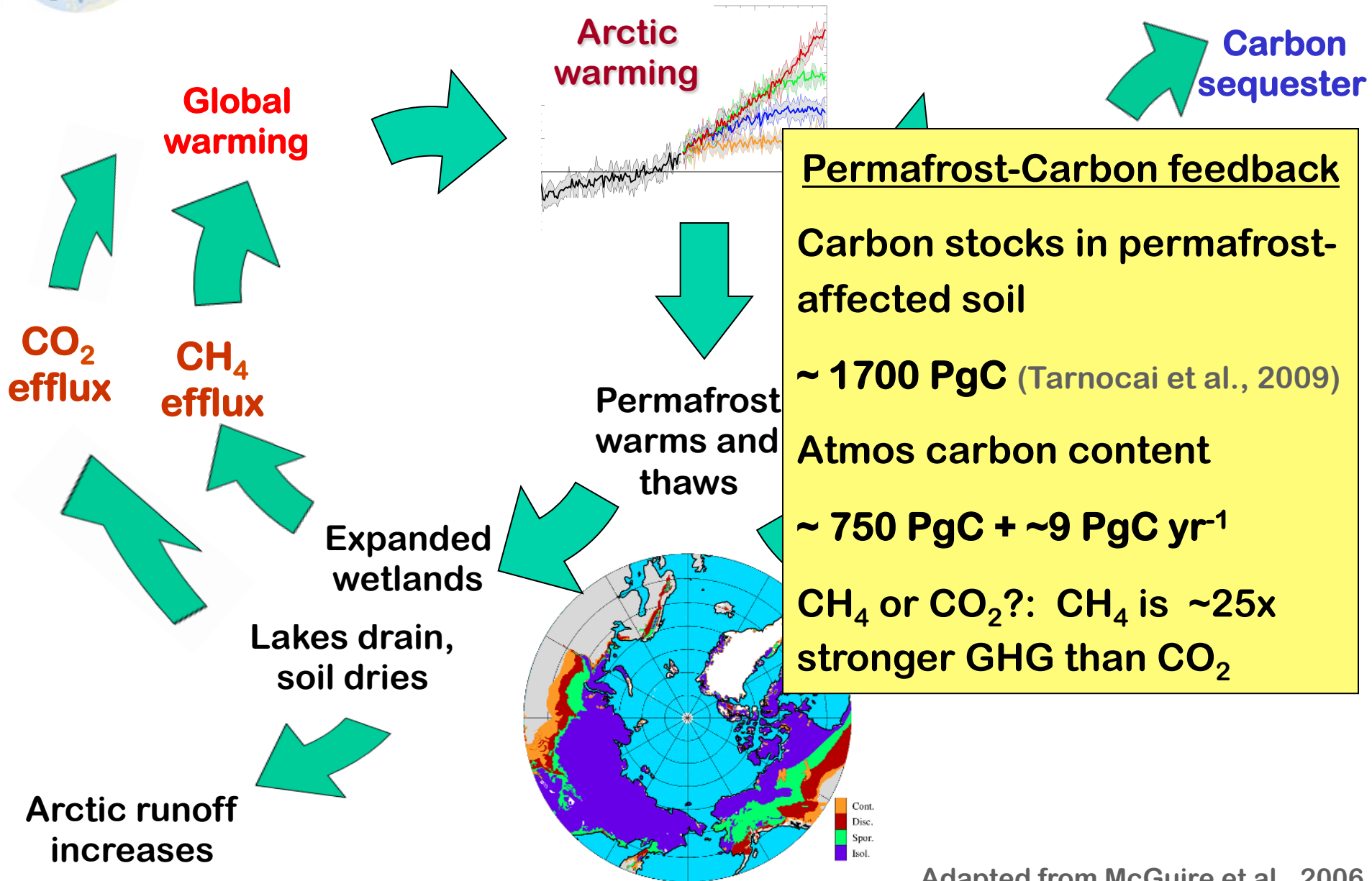


Soil carbon in permafrost zone





Potential Arctic terrestrial climate-change feedbacks





What happens to soil carbon as soil warms and permafrost thaws?

dry, well-drained soil
aerobic decomposition
→ CO₂ emissions

increased wetlands and warmer soil
anaerobic decomposition
→ CH₄ production (25x GWP)

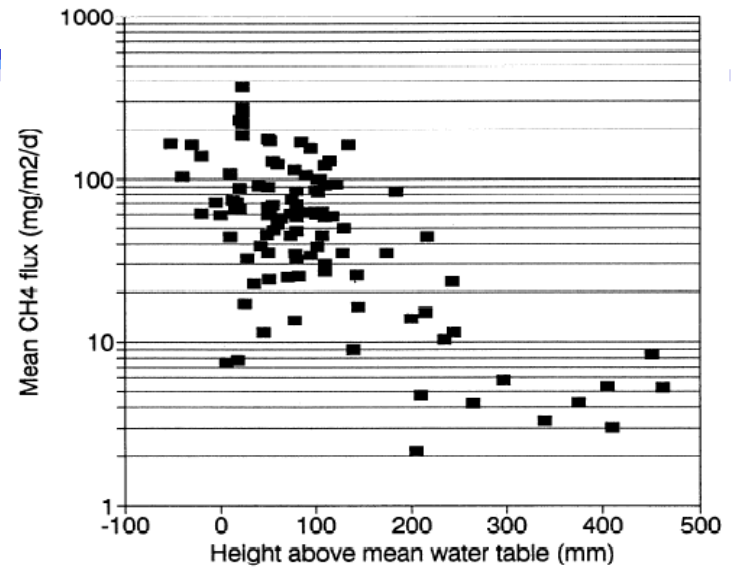


1978

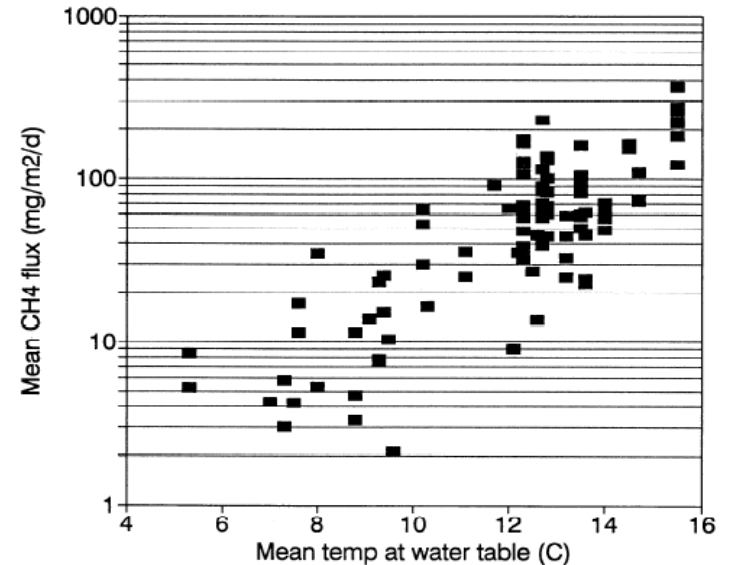


1998

Methane vs. water table



Methane vs. temperature

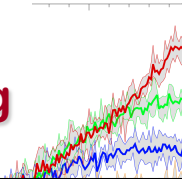




Potential Arctic terrestrial climate-change feedbacks

Global

Arctic warming



Carbon sequester

What is the integrated effect of Arctic land feedbacks?

Is it **+** or **-**?

The hydrology and permafrost-carbon feedbacks are not represented in CMIP3 or CMIP5 era Earth System models

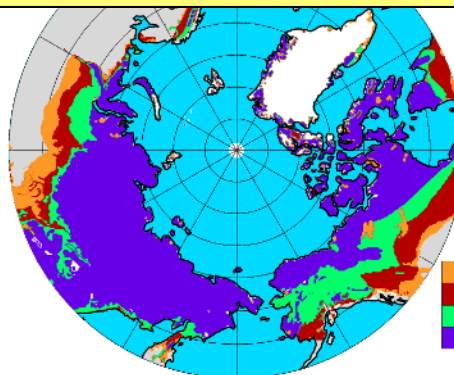
Limits our capacity to provide quantitative analysis on a key vulnerability in Earth system

CO₂ efflux

Lakes drain, soil dries

Increases

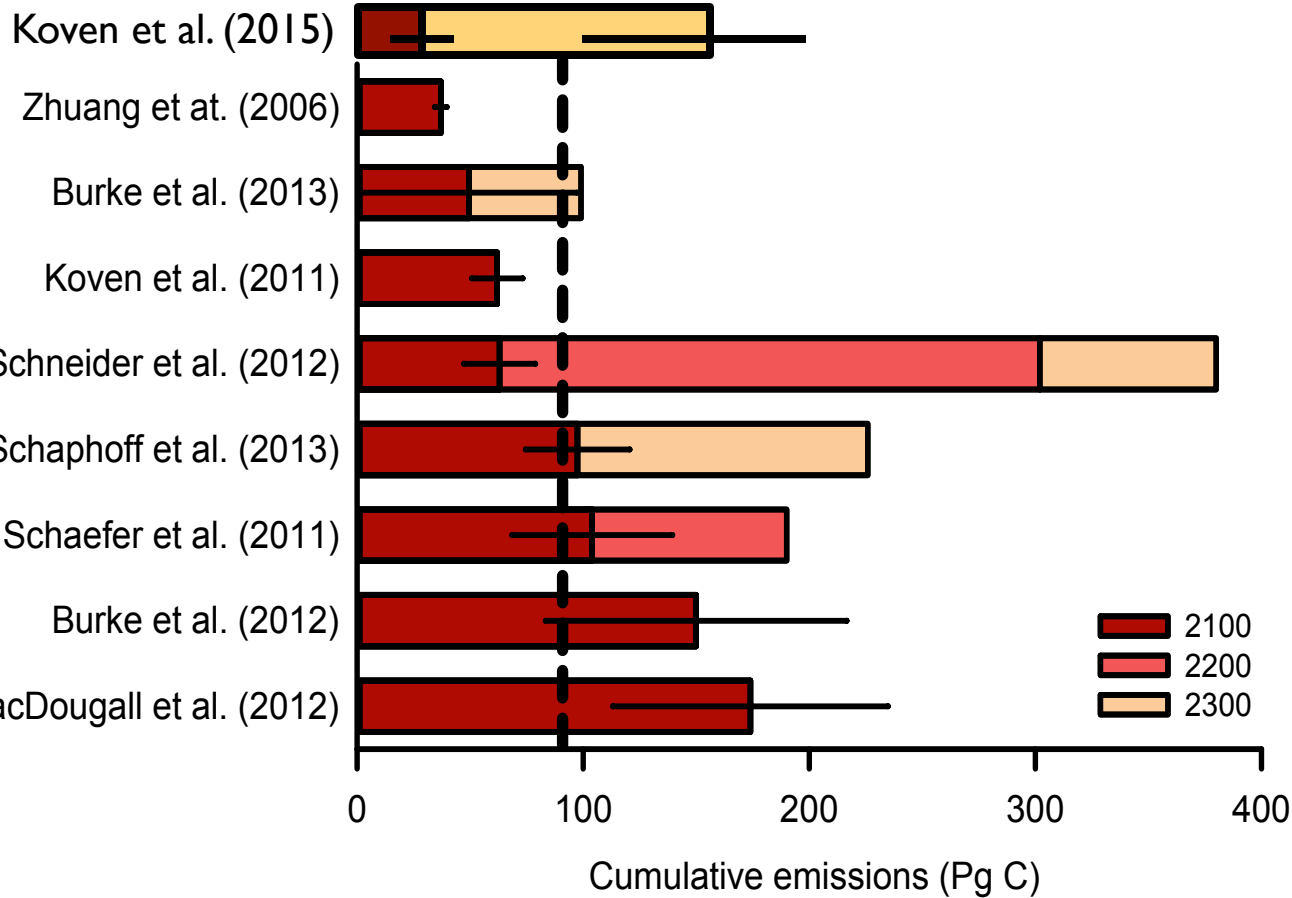
Arctic runoff increases



Cont.
Disc.
Spor.
Isol.



Model projections of permafrost carbon emissions



By 2100 (RCP8.5)

Ensemble Mean

92 ± 17 Pg C

5-15% of initial pool

Timescale

60% of permafrost carbon emissions after 2100

Context

Similar magnitude to land use change



LMWG Progress towards goal of representing permafrost feedbacks in CLM4.5

Soil biogeochemistry: vertically resolved soil carbon model; accounts for limitations on decomposition in cold/saturated conditions

Glo
warm

Carbon
sequester

Shrub
growth

CLM-CNDV (dynamic vegetation): added shrub PFT

CH₄ emission model:
- moisture, T,
vegetation controls
on CH₄ emissions

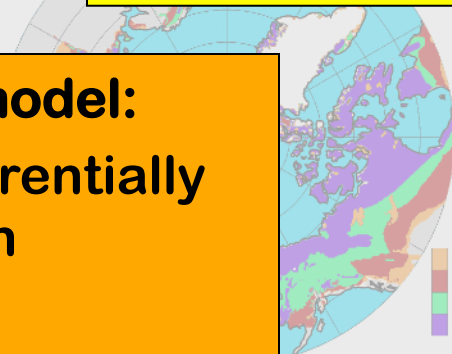
Cold region hydrology/snow:
- more realistic active layer hydrology
- new snow cover fraction

Expanded
wetlands

Lakes drain,

Prognostic wetland model:
- wetlands form preferentially in low gradient terrain
- flooding

Arctic
incr

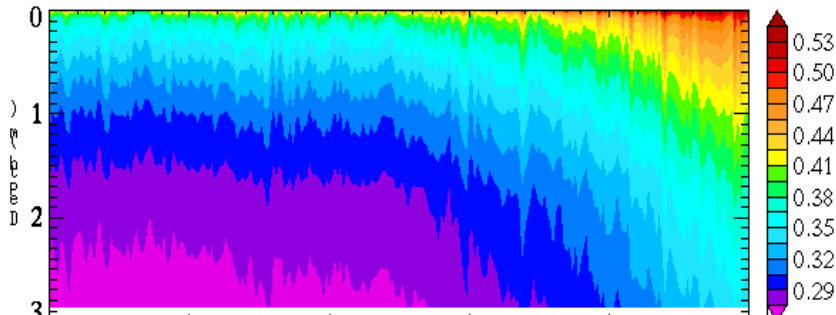




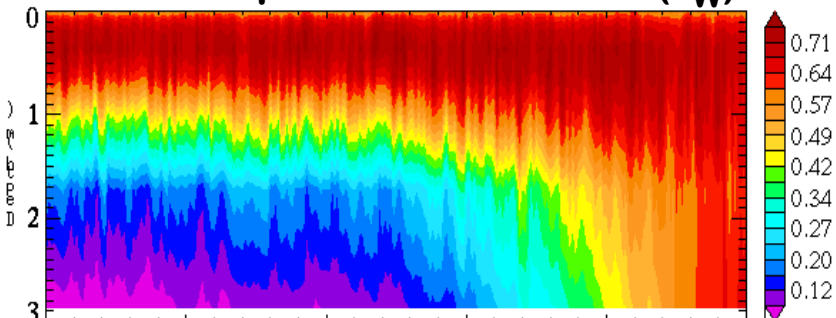
Soil carbon decomposition in CLM4.5

Permafrost zone

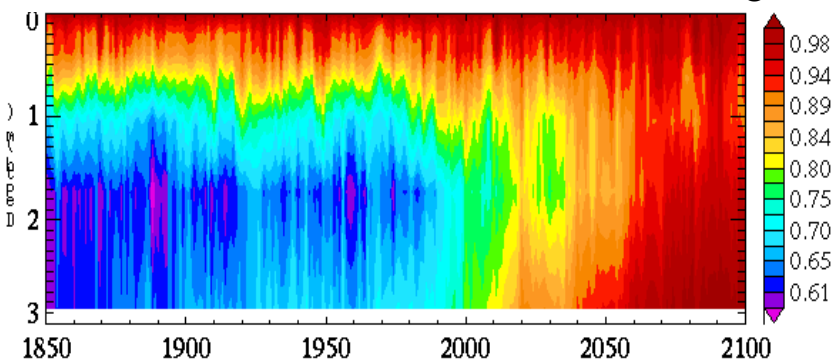
Temperature scalar (r_T)



Soil liquid water scalar (r_W)

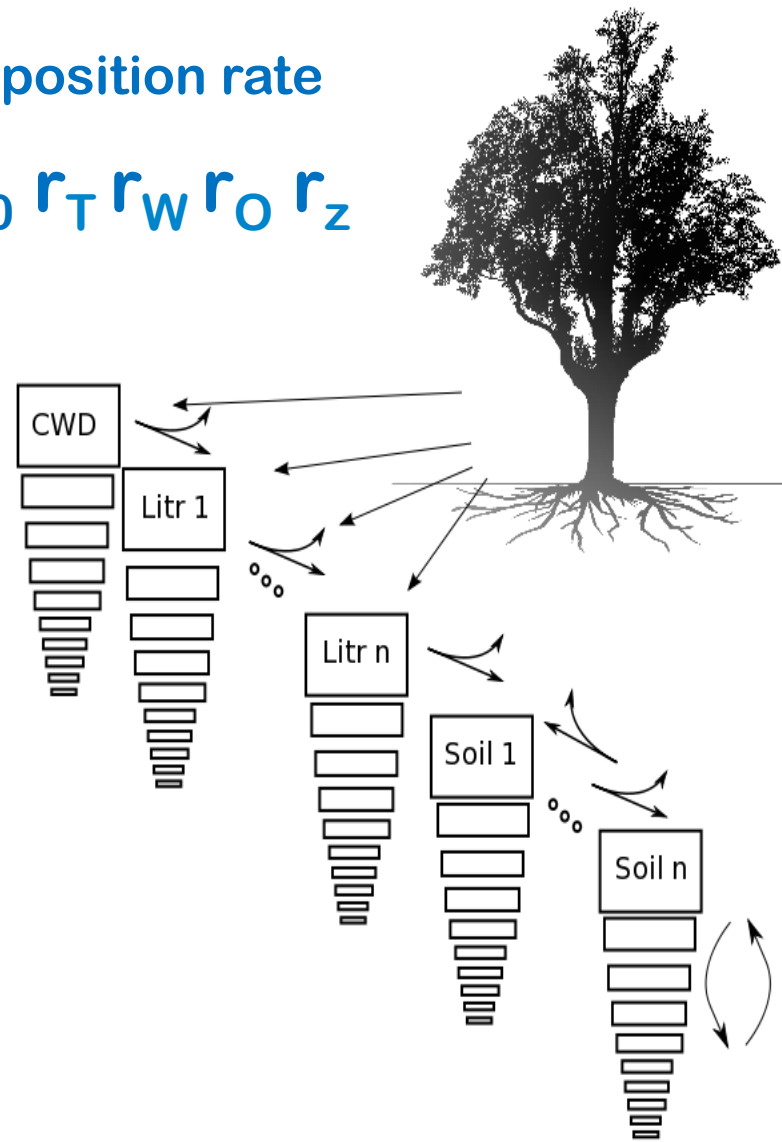


Oxygen availability scalar (r_O)



Decomposition rate

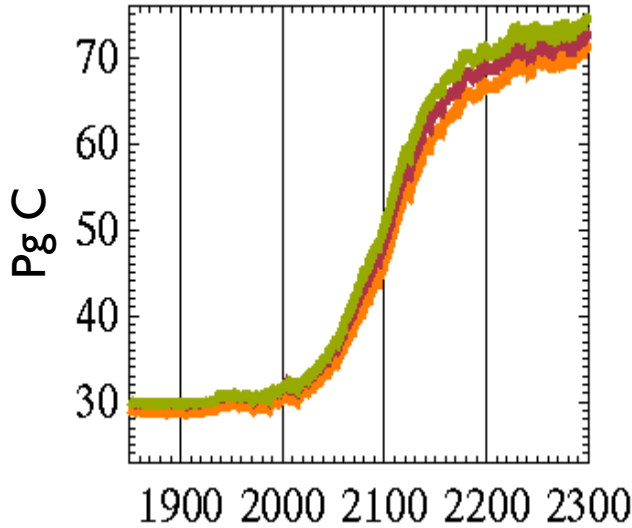
$$k = k_0 r_T r_W r_O r_z$$



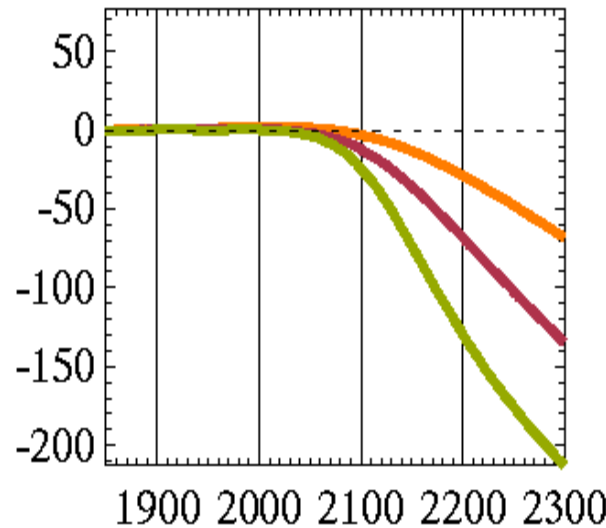


Permafrost Carbon Feedback (CLM4.5BGC)

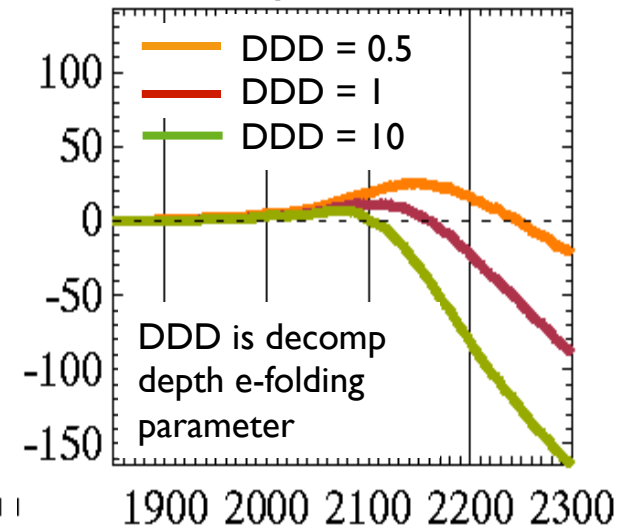
Δ Veg Carbon



Δ Soil Carbon

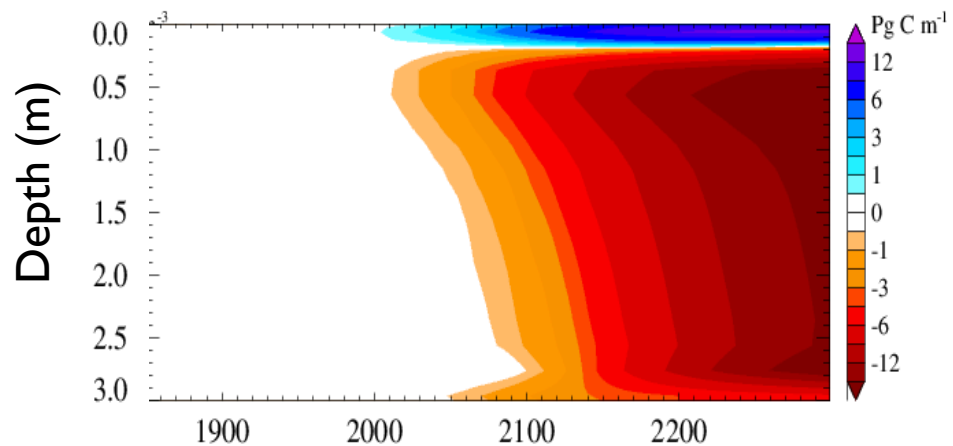


Δ Ecosys Carbon



17 – 42 Pg of 'deep' carbon lost by 2100
103 – 252 Pg by 2300

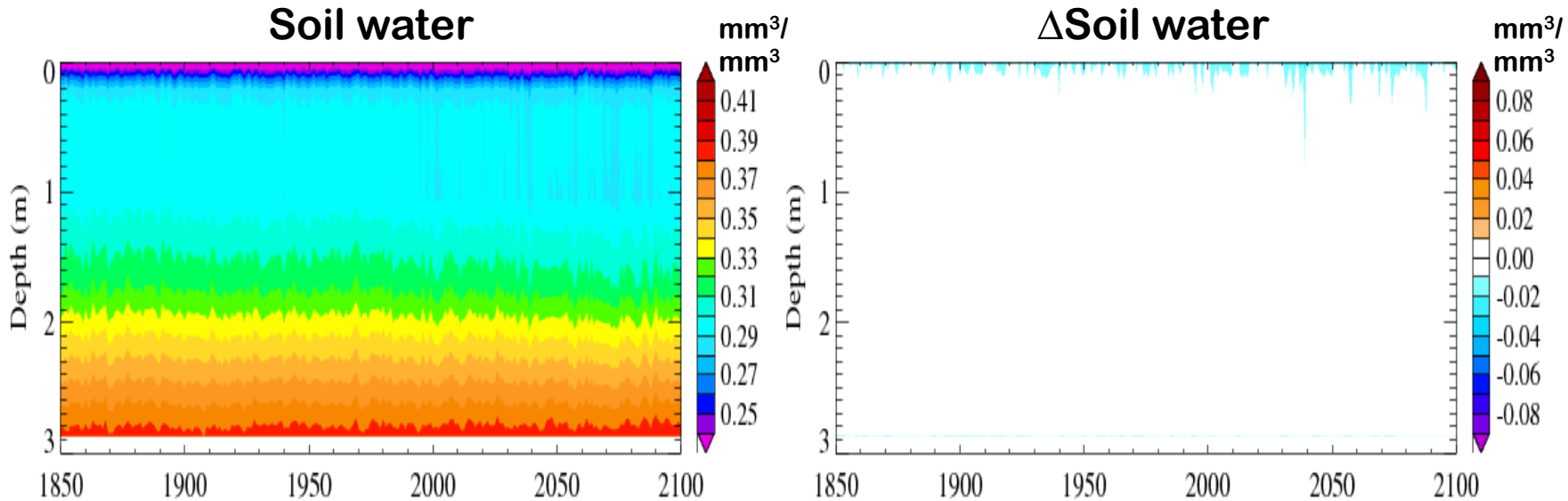
Δ Soil carbon since 1850





Soil hydrologic response to permafrost thaw (RCP8.5)

CLM4



Problems with CLM4 active layer hydrology

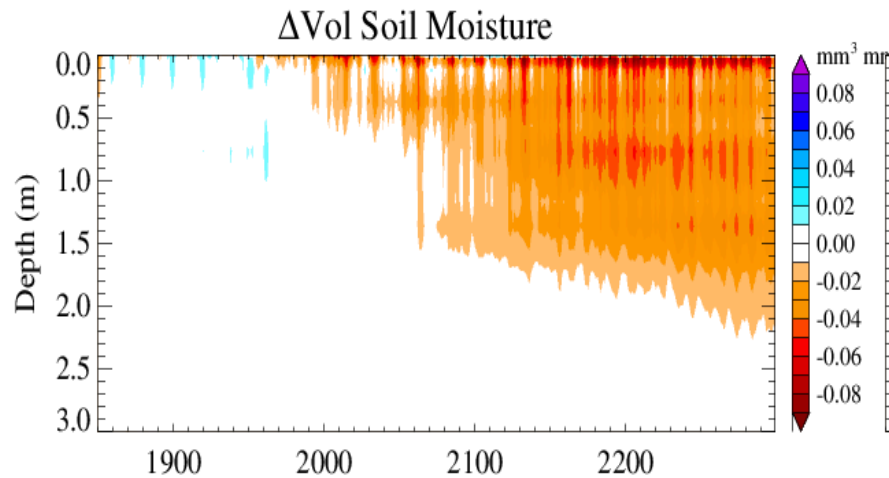
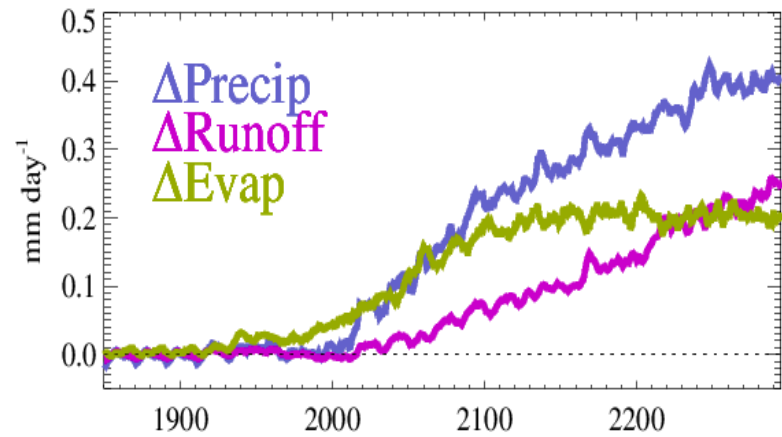
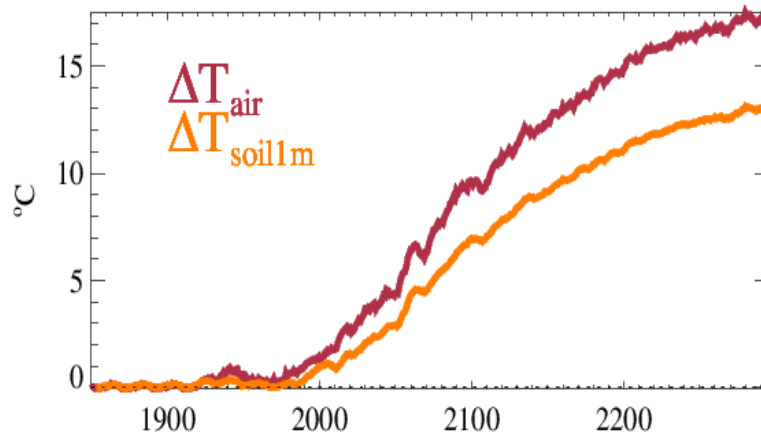
Surface soils are very dry

(some locations are too dry to support vegetation)

No soil moisture response to climate change or permafrost thaw



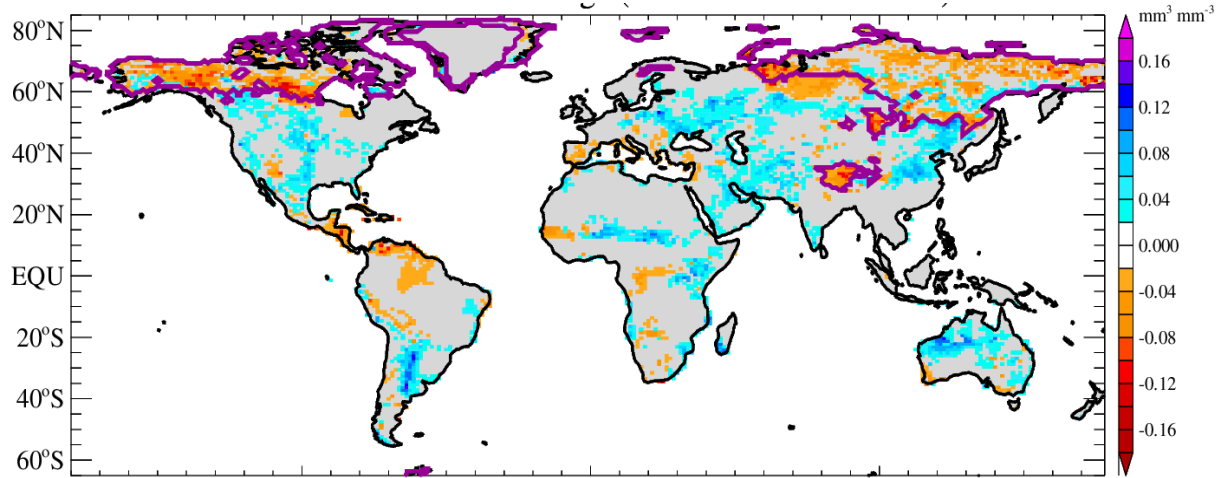
CESM Projections of temperature and water balance for permafrost domain (RCP8.5)



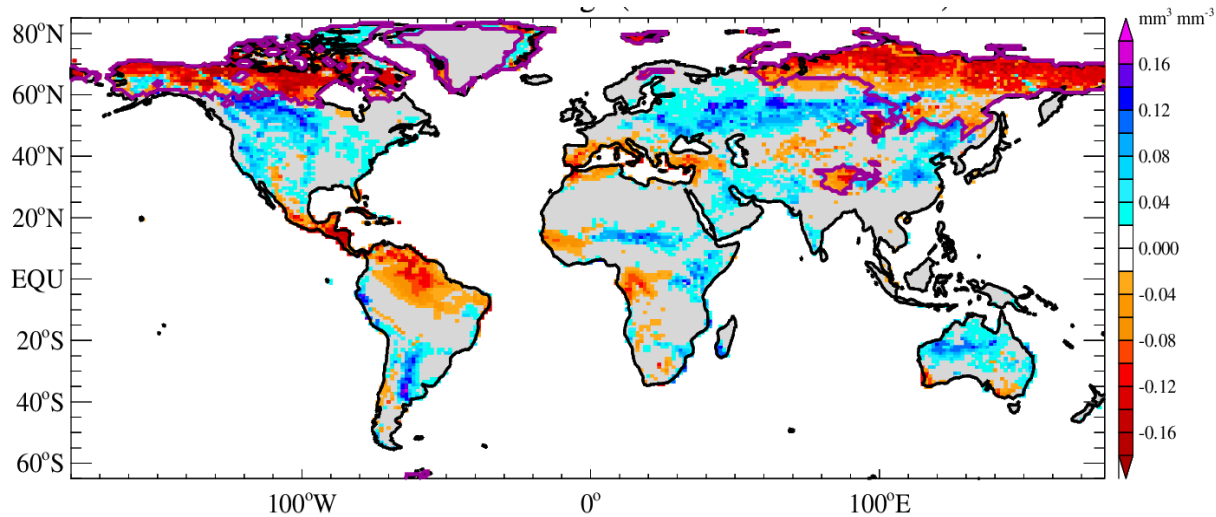


Projected soil moisture change (RCP8.5) CLM4.5

Column soil moisture change by 2100

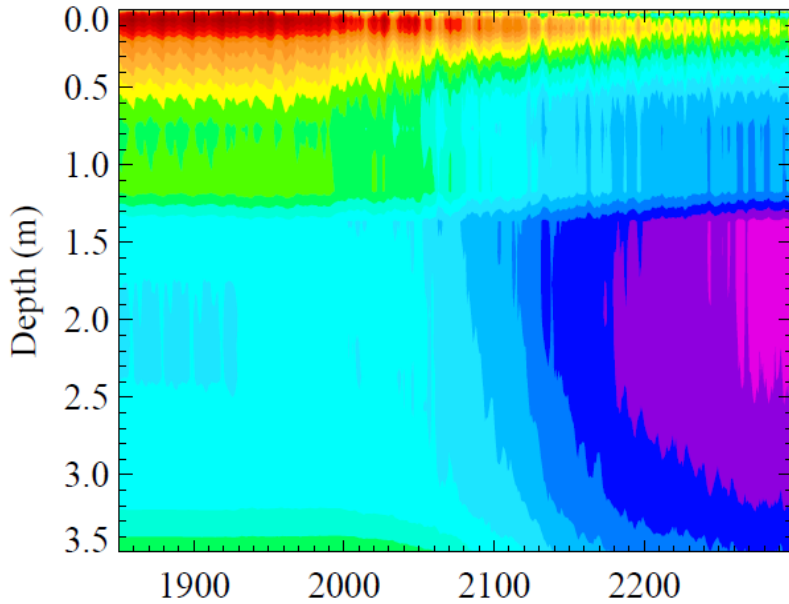


Column soil moisture change by 2300

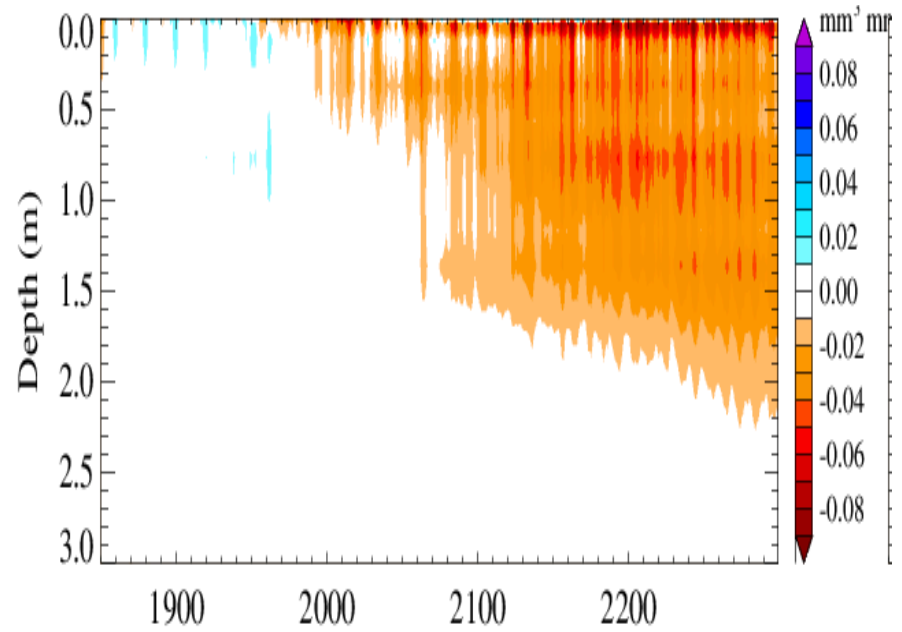


DRY SOIL

Soil Moisture

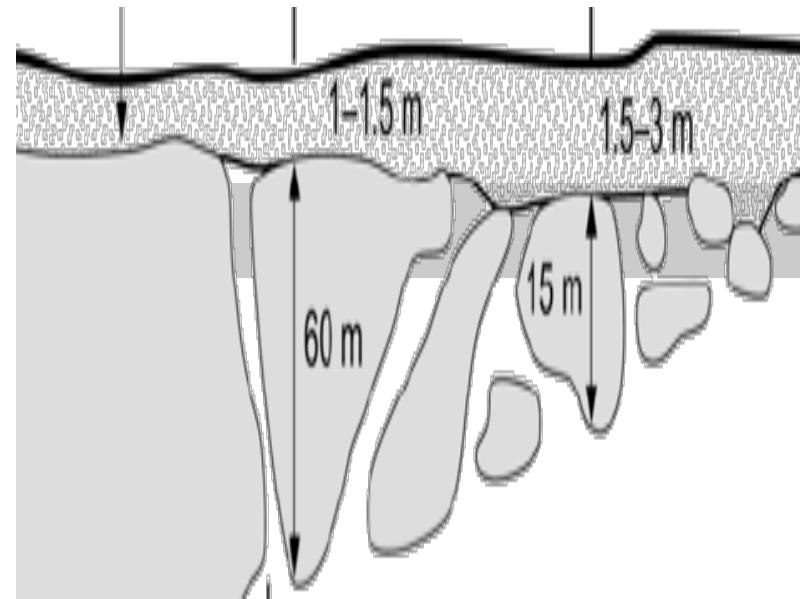


Δ Soil Moisture



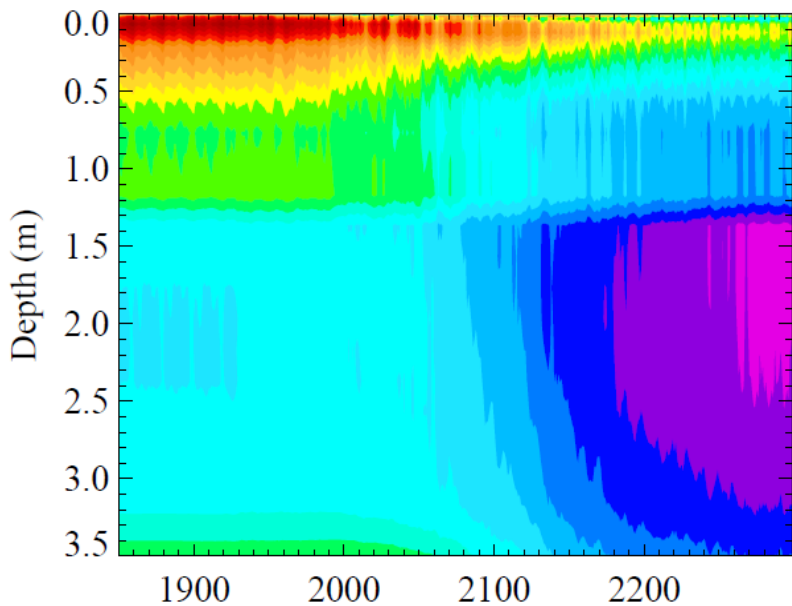
“WETSOIL” experiment

‘Maintain’ initial soil moisture conditions by maintaining the vertical profile of impedance to liquid water drainage

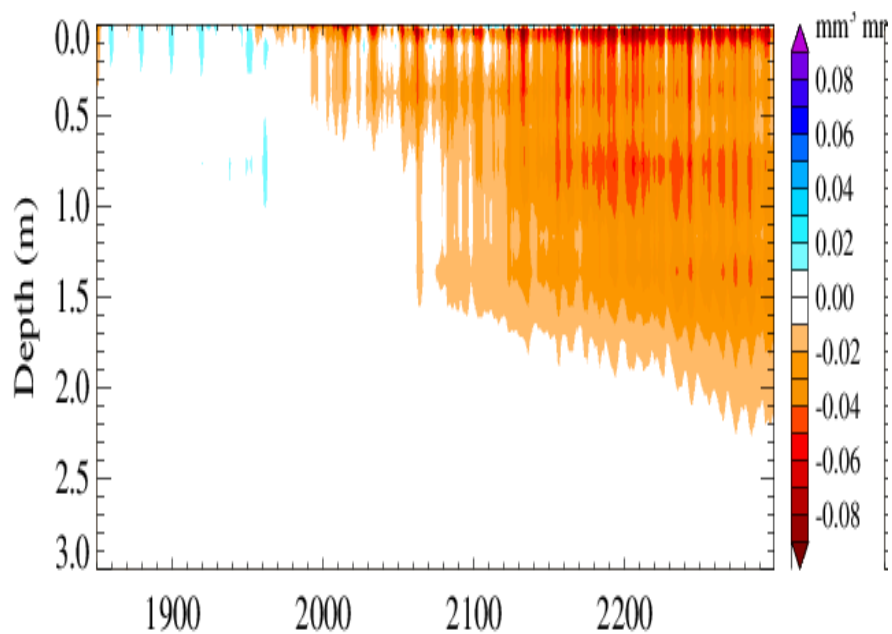


Soil Moisture

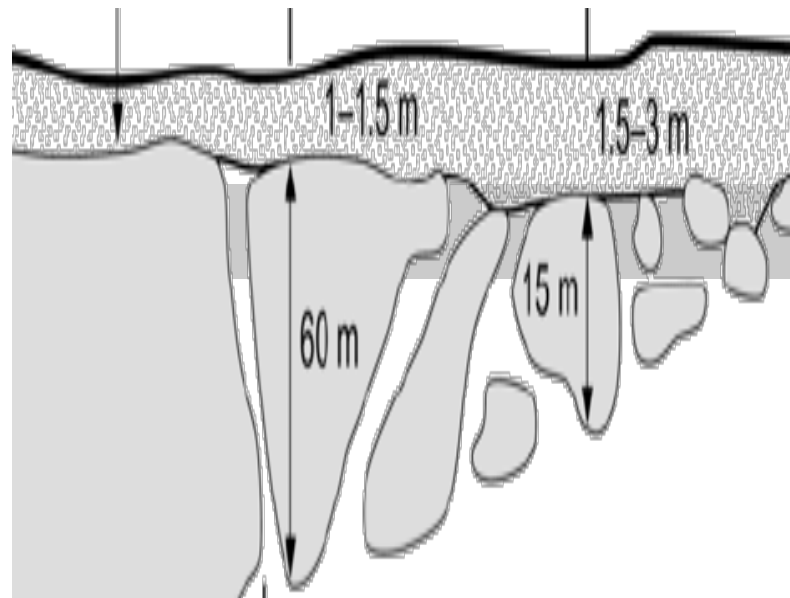
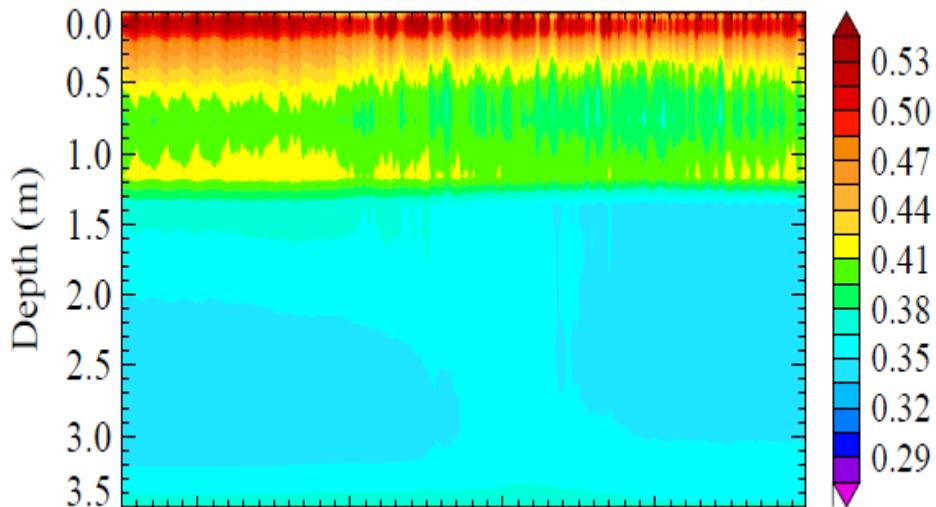
DRY SOIL



Δ Soil Moisture



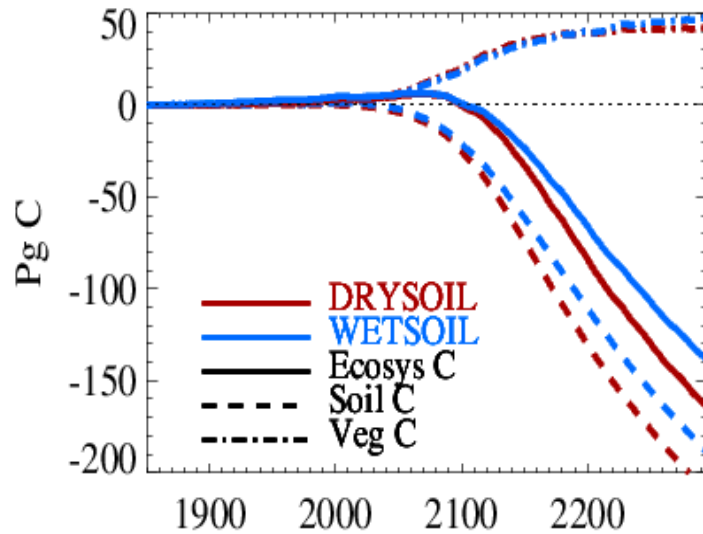
WET SOIL





Global Warming Potential for DRYSOIL and WETSOIL expts

Δ Carbon Stock

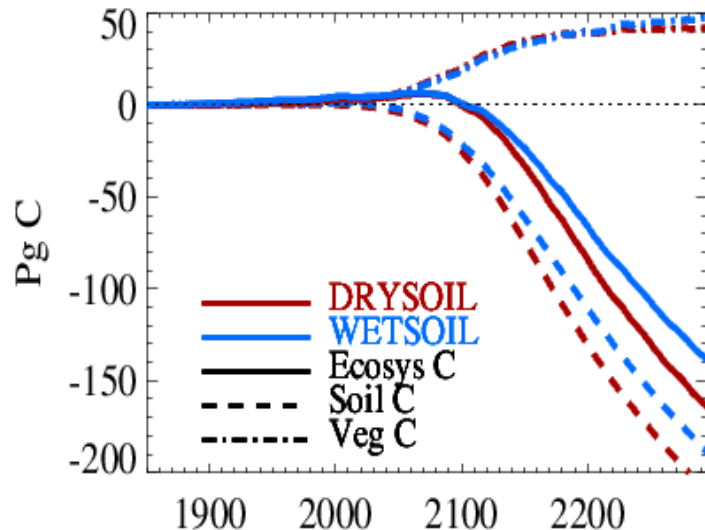


18% more permafrost
soil carbon lost in
DRYSOIL case

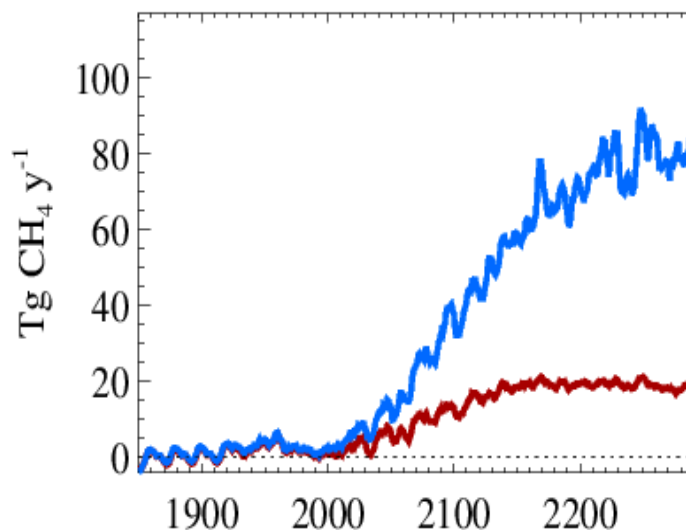


Global Warming Potential for DRYSOIL and WETSOIL expts

Δ Carbon Stock



Δ CH₄ emissions



DRYSOIL

$$\text{CO}_2 = 0.75$$

$$\text{CH}_4 = 0.19$$

$$\mathbf{0.94} \text{ Pg CO}_2\text{e-C y}^{-1}$$

WETSOIL

$$\text{CO}_2 = 0.67$$

$$\text{CH}_4 = 0.82$$

$$\mathbf{1.49} \text{ Tg CO}_2\text{e-C y}^{-1}$$

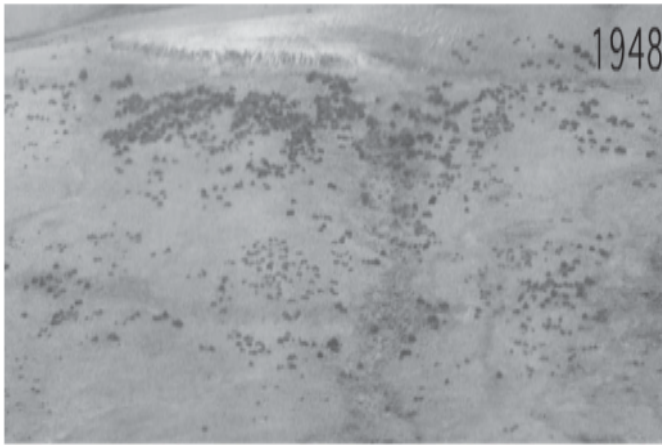
50% higher
GWP in
WETSOIL
case



Photos: Bernhard Edmaier , National Geographic



Shrub – permafrost interactions



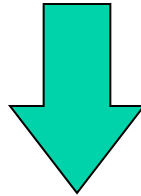
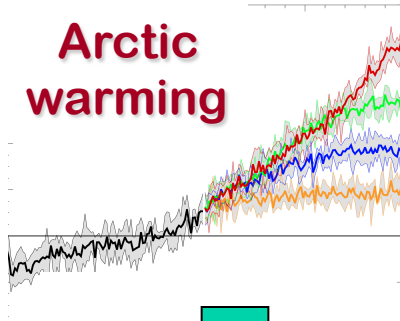
1948



2002

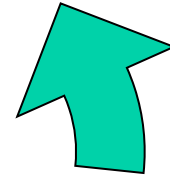
+7% increase in shrubs in Alaska, 1950 to 2005

Arctic warming



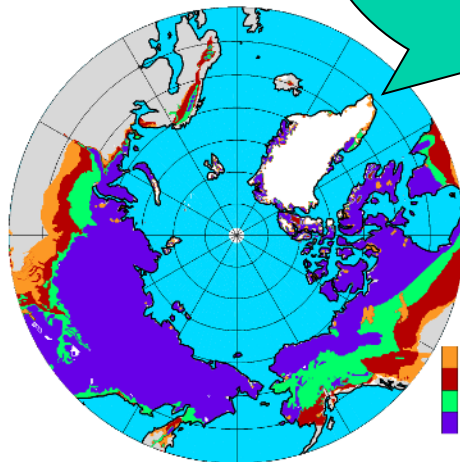
Permafrost warms and thaws

Shrub growth



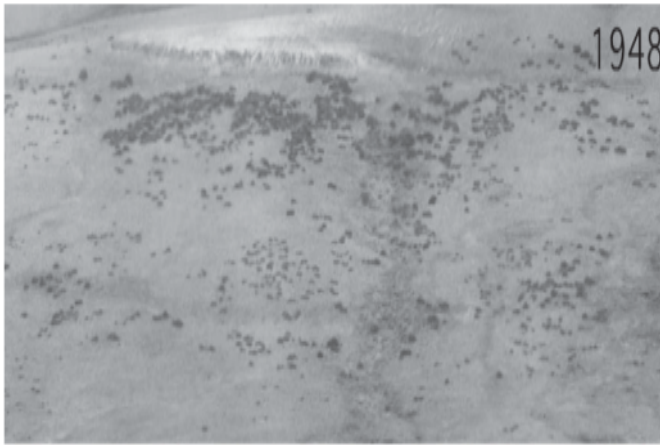
Enhanced [nitrogen]

Microbial activity increases





Shrub – permafrost interactions

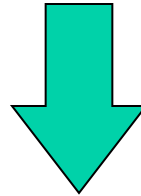
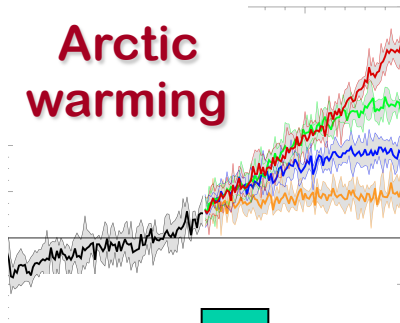


1948

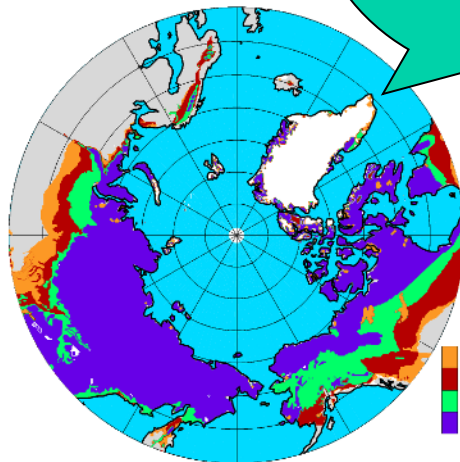


2002

+7% increase in shrubs in Alaska, 1950 to 2005



Permafrost warms and thaws

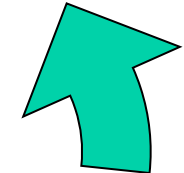


Microbial activity increases

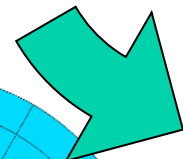
+20PgC for +20% shrub

Shrub growth

Carbon Sequester

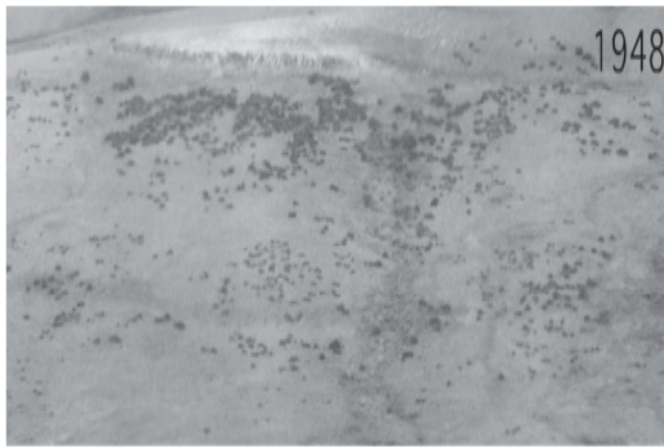


Enhanced [nitrogen]





Shrub – permafrost interactions

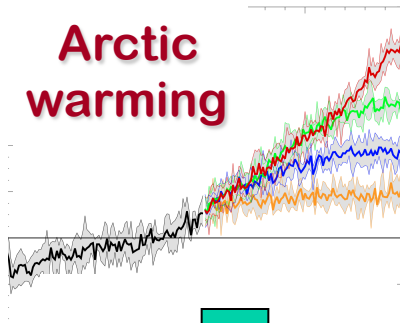


1948



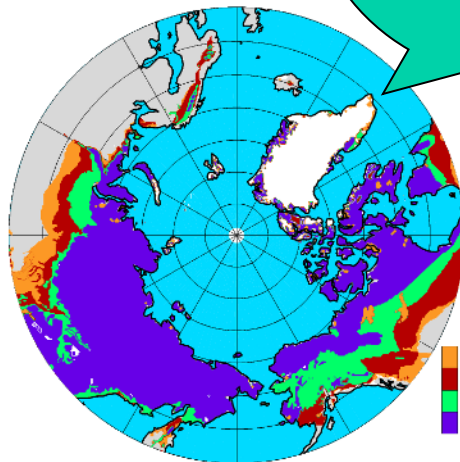
2002

+7% increase in shrubs in Alaska, 1950 to 2005



Arctic warming

Permafrost warms and thaws



Cont.
Disc.
Spor.
Isol.

Microbial activity increases

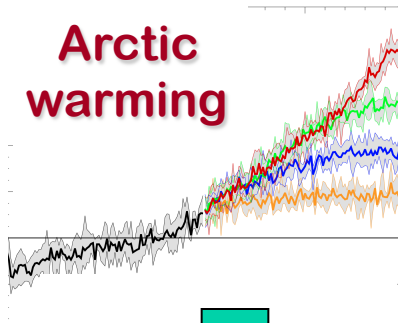
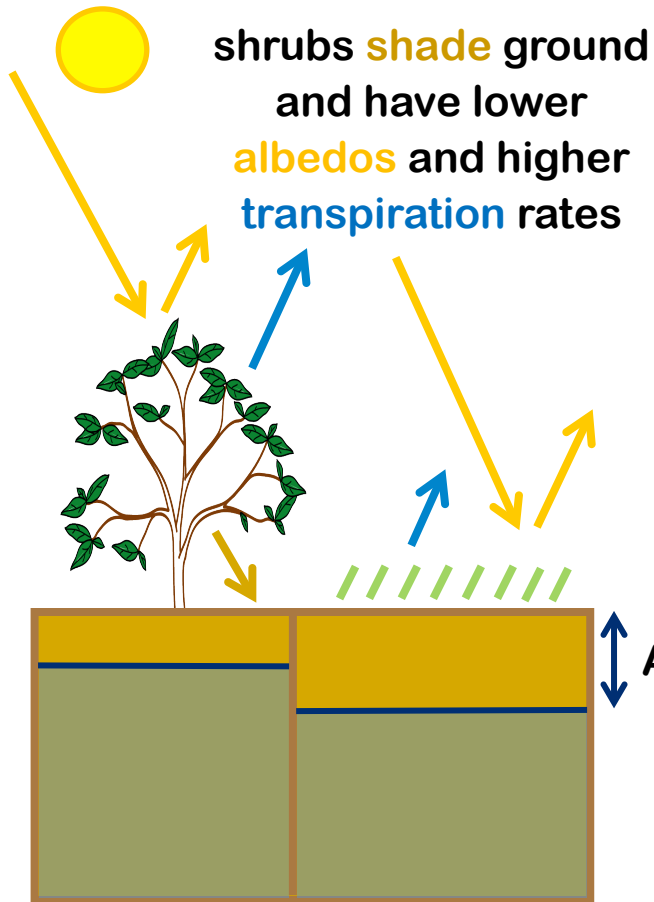
Shrub growth

Carbon Sequester

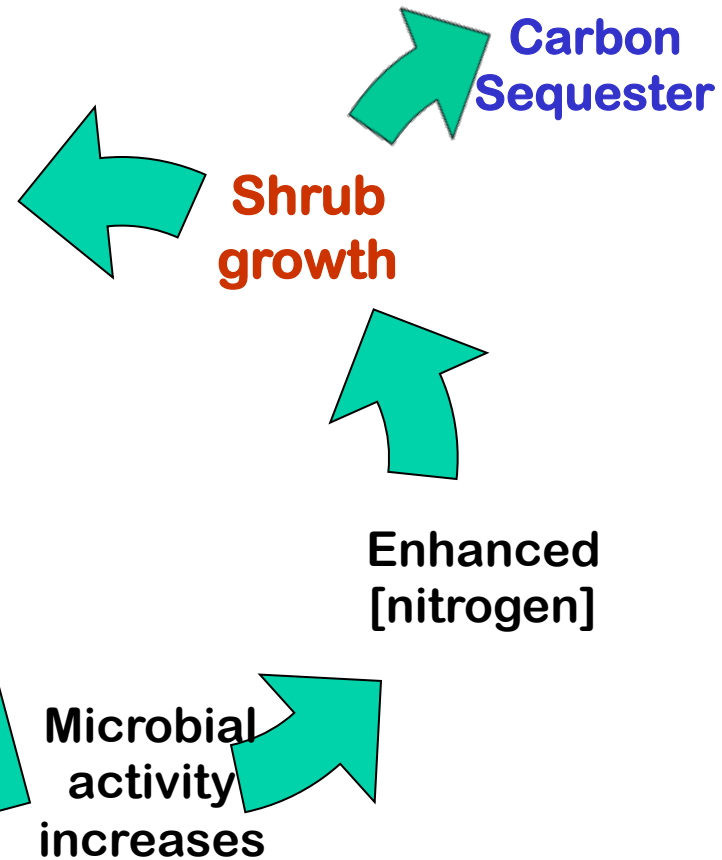
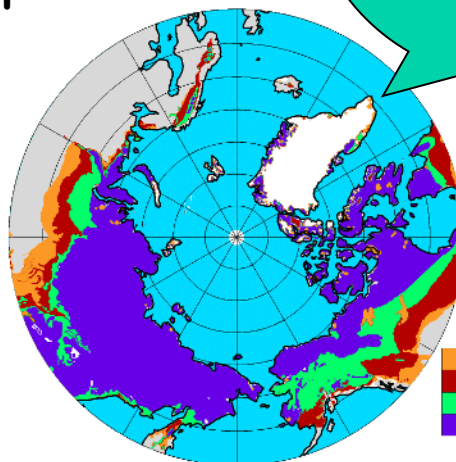
Enhanced [nitrogen]



Potential Arctic terrestrial climate-change feedbacks

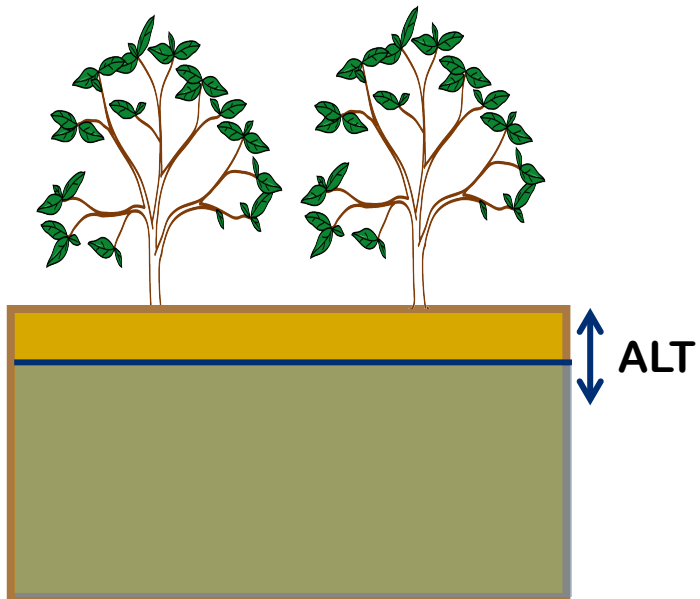


Permafrost warms and thaws



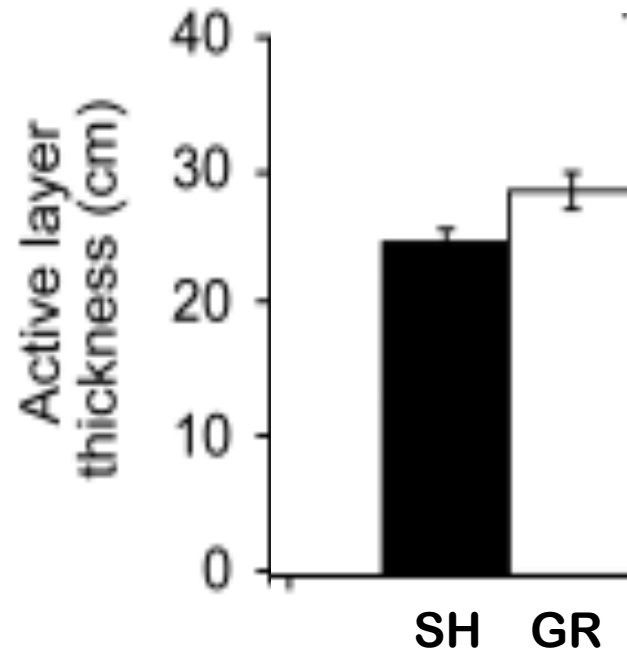
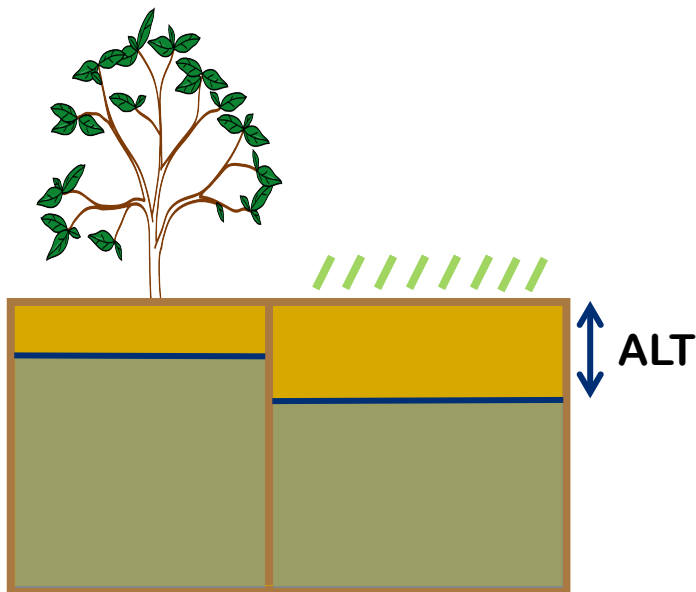
Shrub expansion may reduce summer permafrost thaw in Siberian tundra

D. BLOK*, M. M. P. D. HEIJMANS*, G. SCHAEPMAN-STRUB*†, A. V. KONONOV‡,
T. C. MAXIMOV‡ and F. BERENDSE*



Shrub expansion may reduce summer permafrost thaw in Siberian tundra

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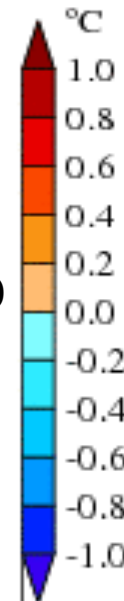
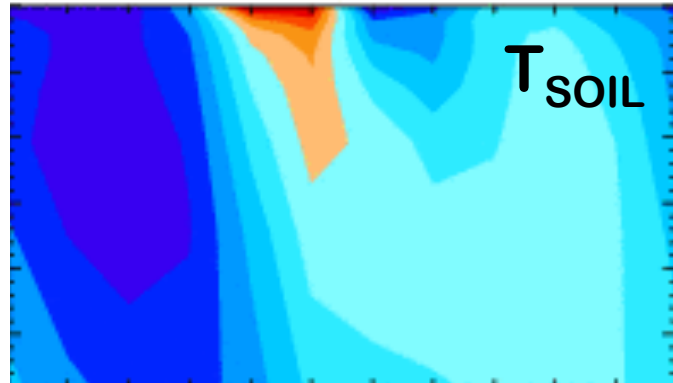
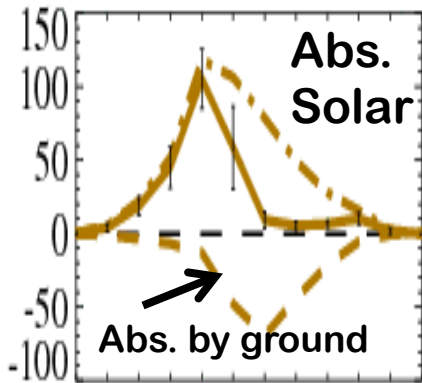
“These results suggest that the expected expansion of deciduous shrubs in the Arctic region, triggered by climate warming, may reduce summer permafrost thaw.”

Evaluate this hypothesis using CCSM4

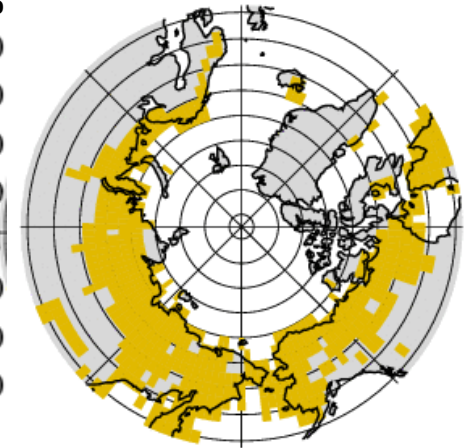


Examining impact of shrubs on permafrost using CESM

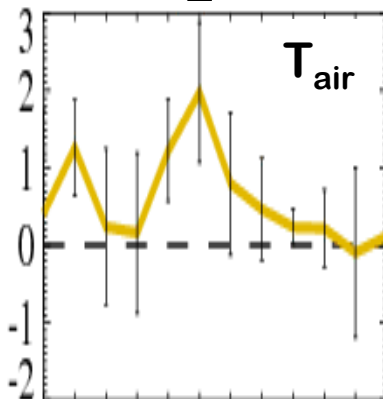
SB_LOW: Shrub – Grass



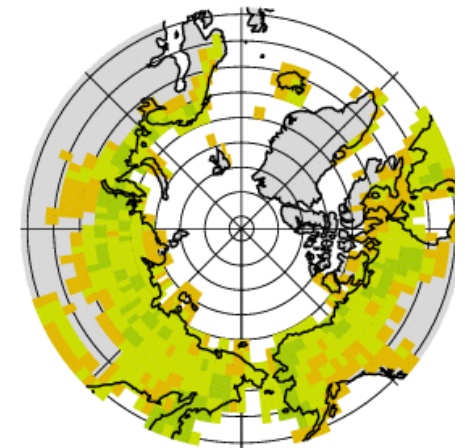
SB_LOW Boreal Shrub



SB_HIGH – SB_LOW: Grid cell mean



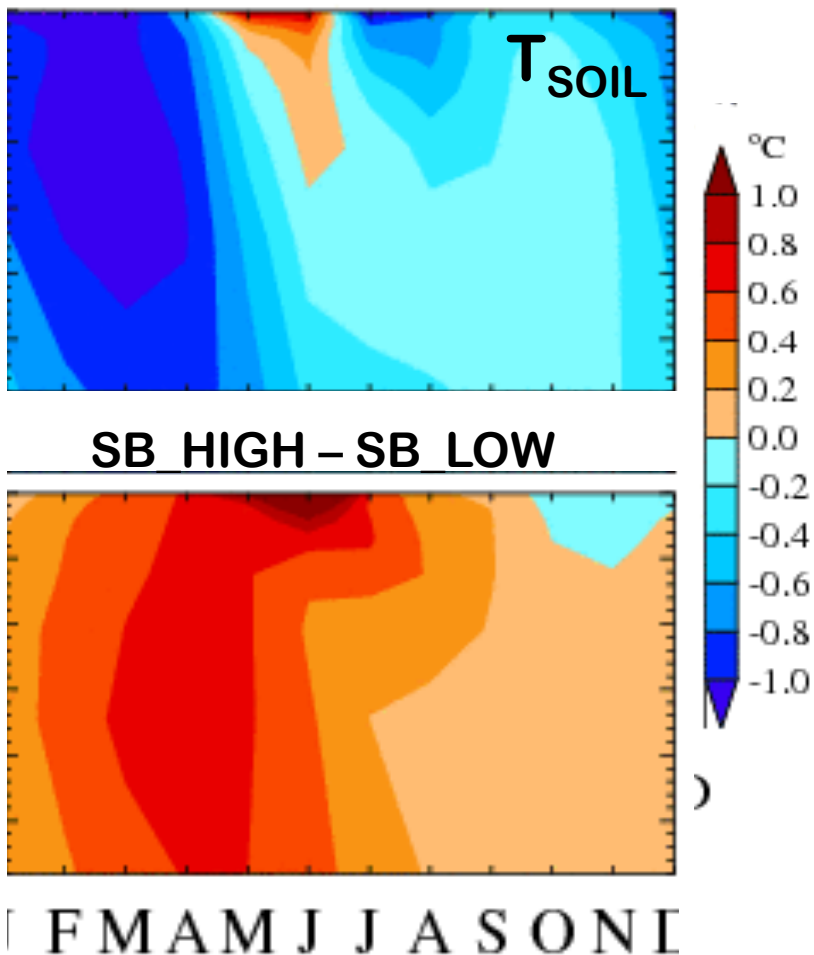
SB_HIGH Boreal Shrub





Impact of shrubs on permafrost

Shrub - Grass



Will expanding Arctic shrub cover decrease permafrost vulnerability to climate change?

A. Not necessarily. Depends on whether direct local cooling or indirect climate warming dominates.

CAM/CLM results indicate that shrub expansion **may actually increase rather than decrease permafrost vulnerability to climate change.**



Summary

- Substantial near-surface permafrost degradation is projected for 21st century
- Process-rich enhancements to CLM (soil thermodynamics and hydrology, soil biogeochemistry, CH₄ emissions, prognostic wetlands) are enabling study of permafrost dynamics and feedbacks
- Initial results suggest that feedbacks will amplify climate change, though magnitude is highly uncertain
 - Warming feedbacks related to shrub encroachment may dominate in 21st century
 - Permafrost-carbon feedback might be relatively small in 21st century but likely to amplify and extend into 22nd century and beyond as soils warm and dry
 - Permafrost soils may dry after permafrost thaws, which enhances aerobic SOM decomposition and limits increase in CH₄ emissions

Summary

- Substantial near-surface permafrost degradation is projected for 21st century and beyond, likely initiating several feedbacks that could amplify Arctic and global climate change

- Using CLM4.5BGC, we identify several potential controls on the strength permafrost-carbon feedback

- 1: Deeper (0.5-3m) soil carbon decomposability is first order determinant of amplitude of permafrost-carbon feedback

2. Deep soil N mineralization doesn't strongly fertilize plants due to asynchrony of plant N demand (summer) and additional deep N supply (fall)

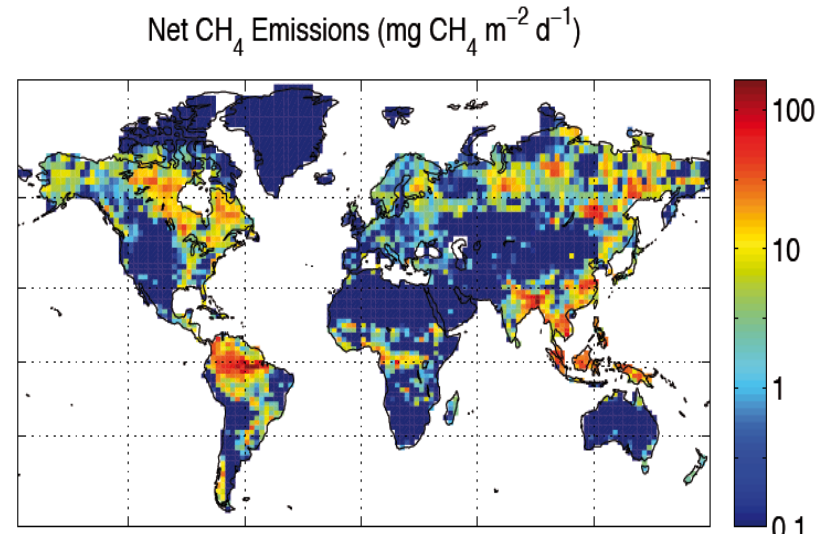
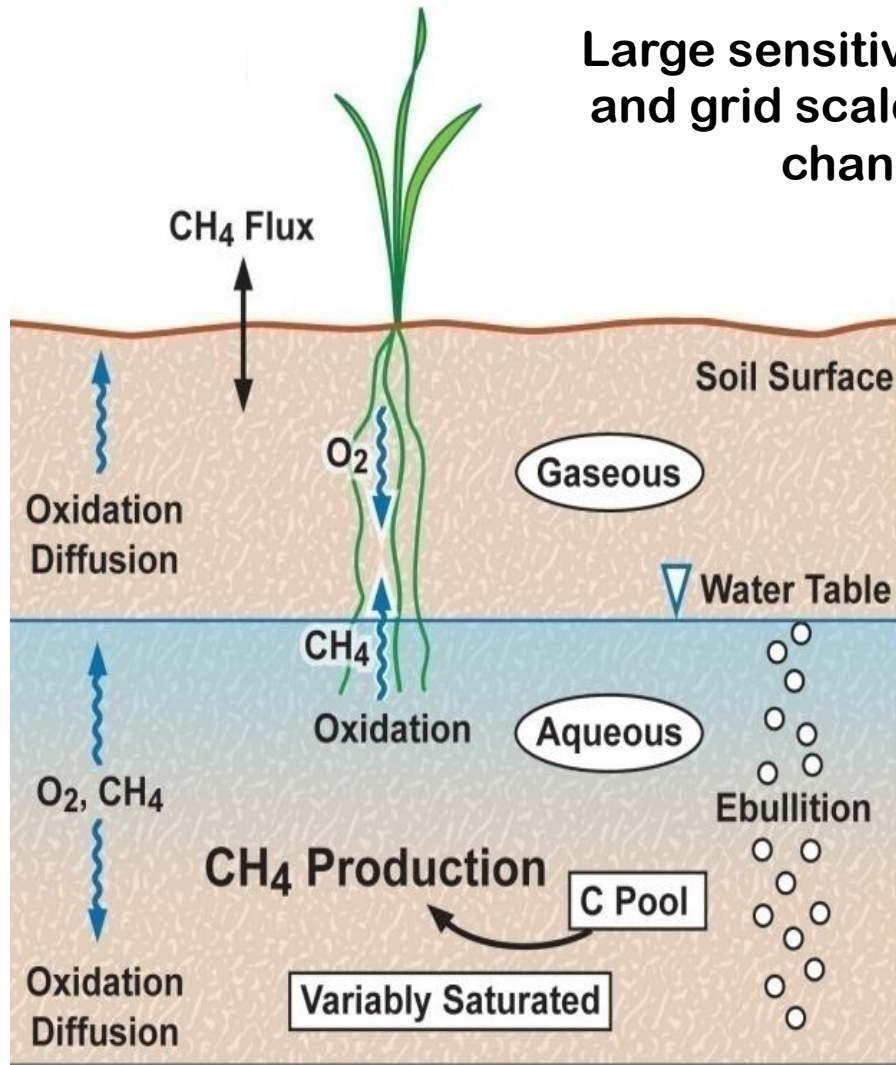
- 3: Permafrost soils will dry after permafrost thaws, which enhances aerobic SOM decomposition and limits increase in CH₄ emissions



Process based methane emissions model

“Barriers to predicting changes in global terrestrial methane fluxes”

Large sensitivities (up to 4x and 10x at regional and grid scales) in CH₄ fluxes from reasonable changes in model parameters

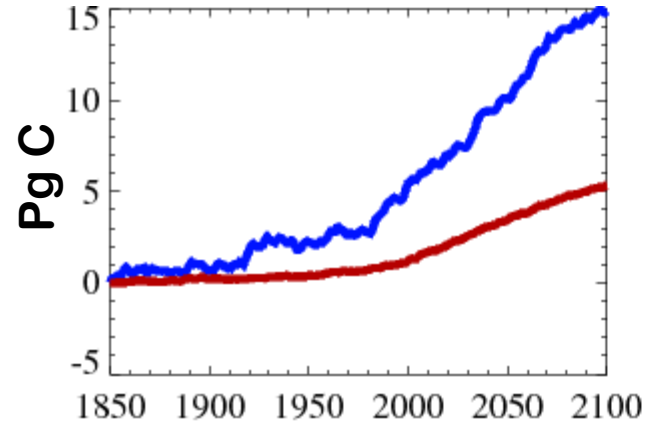


Projections highly uncertain, but with default parameters ~ +20% increase in high-lat CH₄ emissions (A1B)

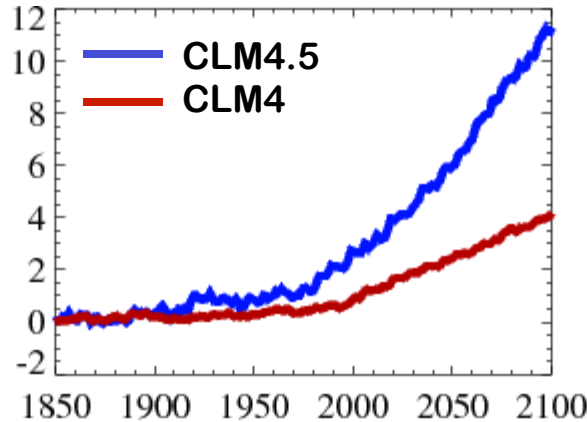


Carbon stock trends in permafrost zone

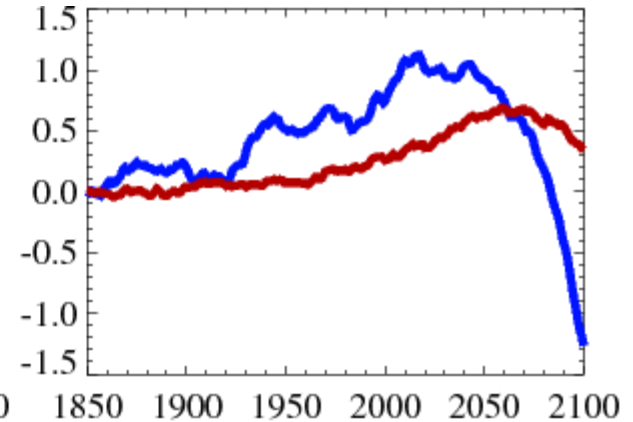
Ecosystem Carbon



Vegetation Carbon



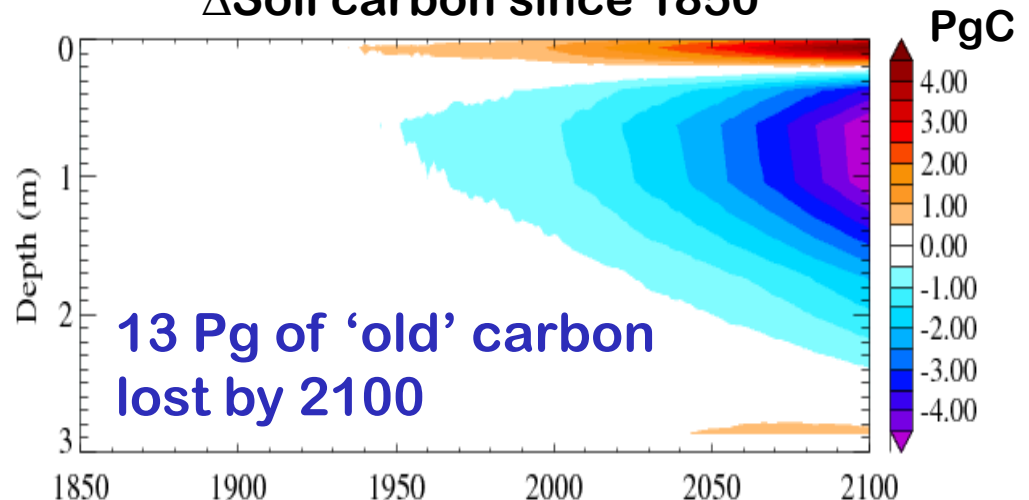
Soil Carbon



Prior estimates of carbon loss (PgC)

- 62 ± 6 ORCHIDEE (Koven et al., 2011)
- 100 ± 40 SibCASA (Schaefer et al. 2011)
- 72 ± 40 MAGICC (Deimling et al., 2011)
- 12 ± 6 TEM (Zhuang et al. 2006)

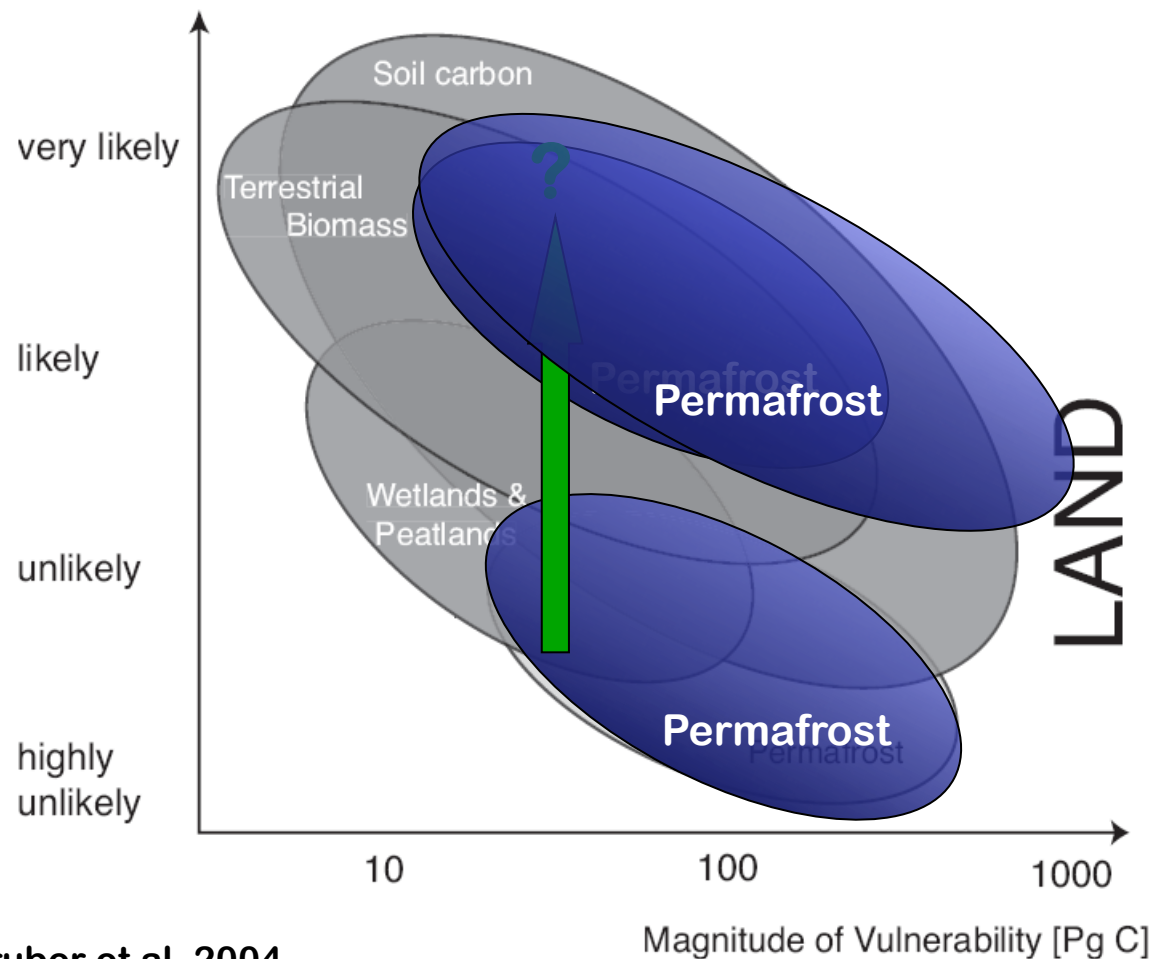
ΔSoil carbon since 1850





Release of Soil Carbon Frozen in Permafrost

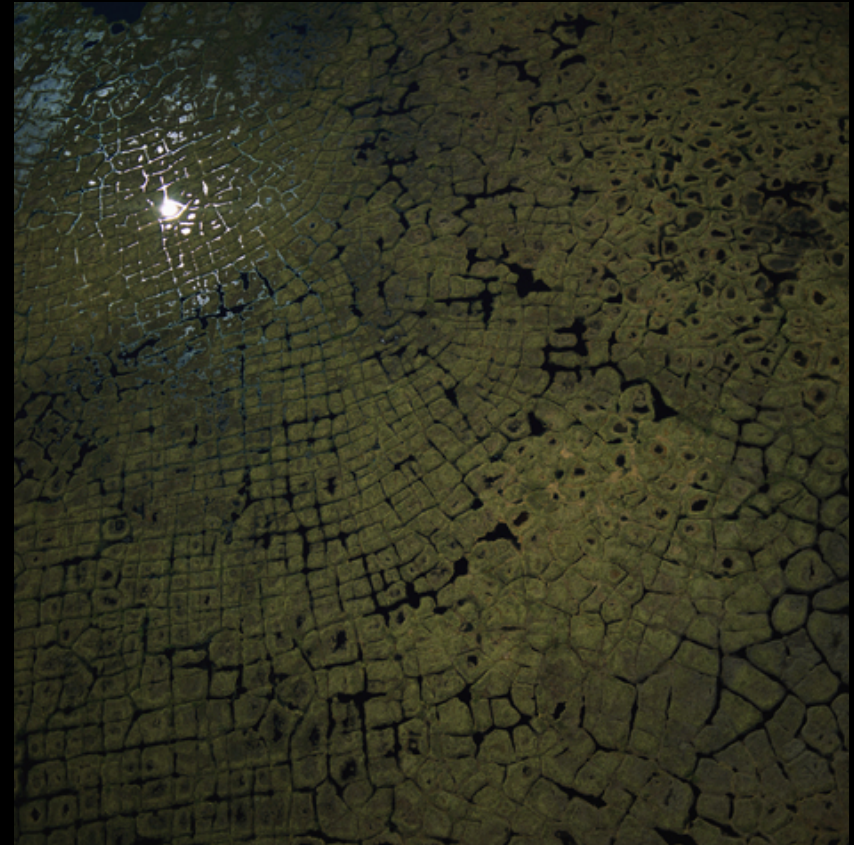
Global Carbon Project C-POOLS AT RISK IN THE 21st CENTURY



Gruber et al. 2004

Magnitude of Vulnerability [Pg C]

Extra Slides

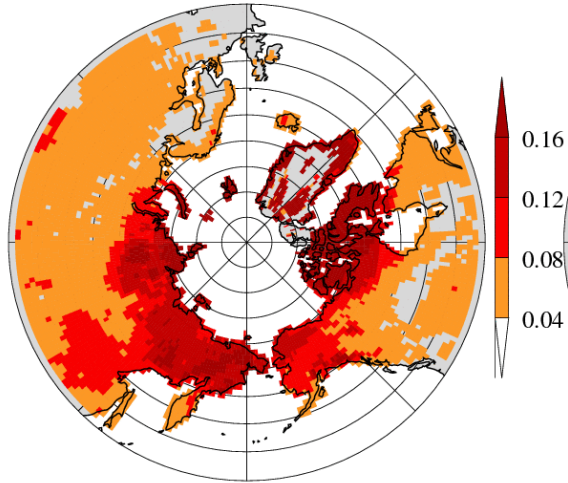


Bernhard Edmaier
National Geographic

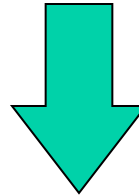
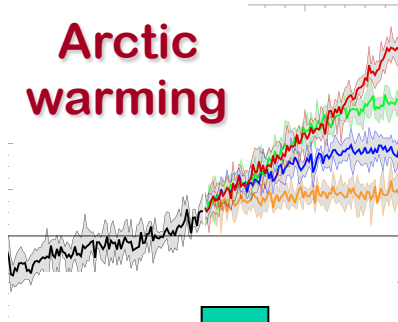


Potential Arctic terrestrial climate change feedbacks

GH / R_{NET}
1980-1999 (JA)

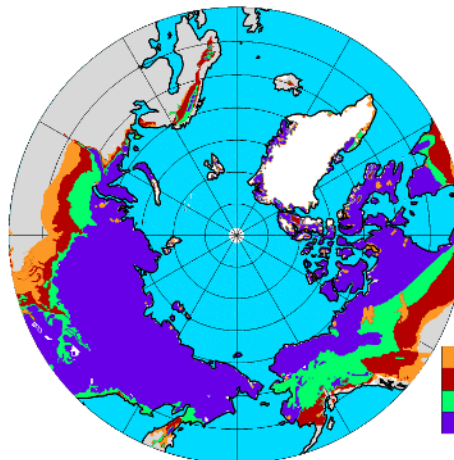
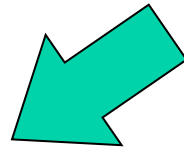
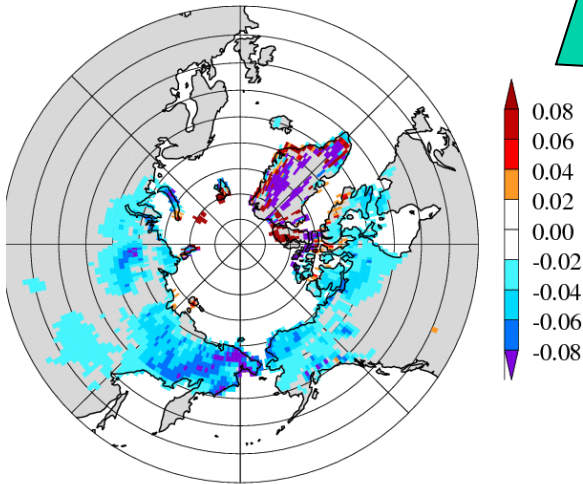


Arctic
warming



Permafrost
warms and
thaws

$\Delta(\text{GH} / \text{R}_{\text{NET}})$
2080-2099 minus 1980-1999 (JA)



Cont.
Disc.
Spor.
Isol.

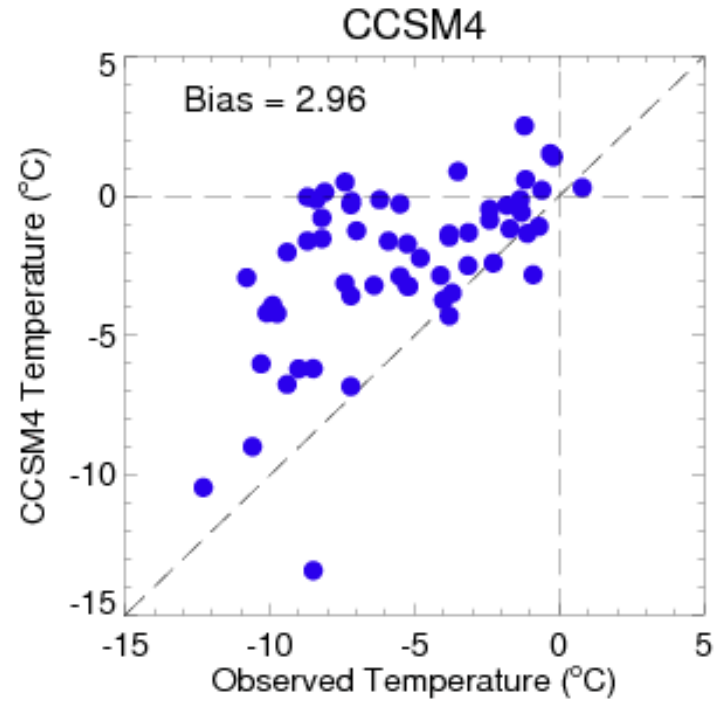
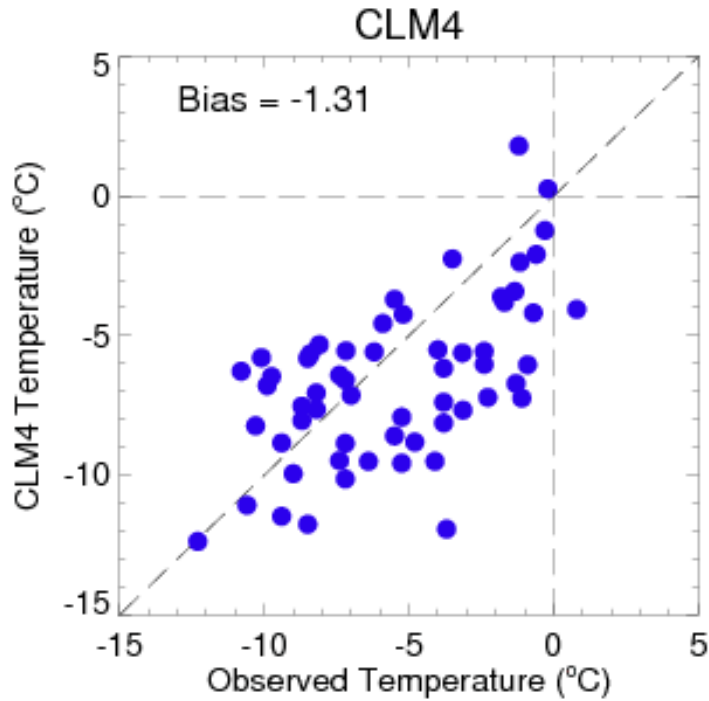
Direct feedback

Surface energy partitioning

Permafrost state (especially presence or absence of soil ice) affects partitioning of net radiation into ground, latent, and sensible heat fluxes

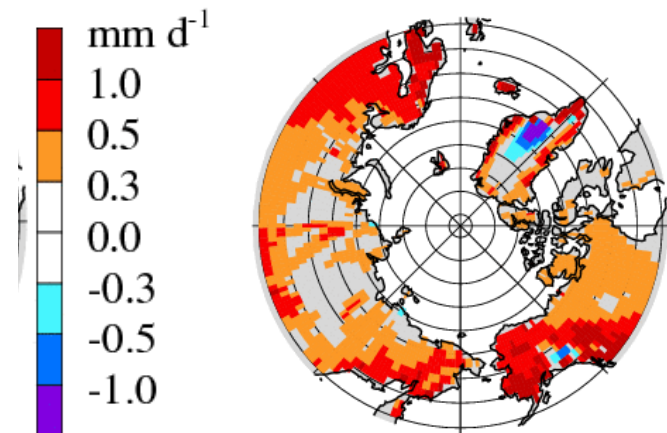


Offline (CLM) vs coupled (CCSM) model deep (> 15m) ground temperatures



Cold bias because soils too dry?

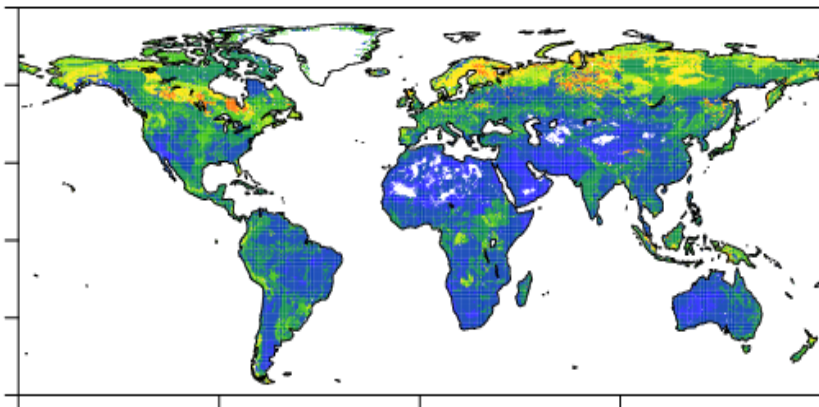
CCSM4 Snowfall bias



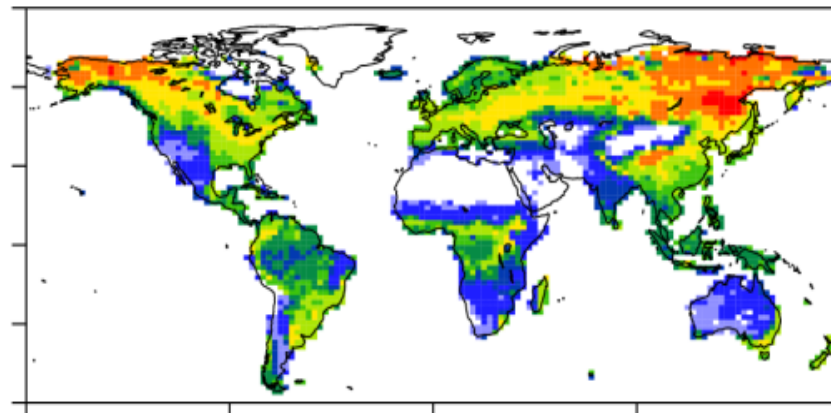


Soil carbon in CLM

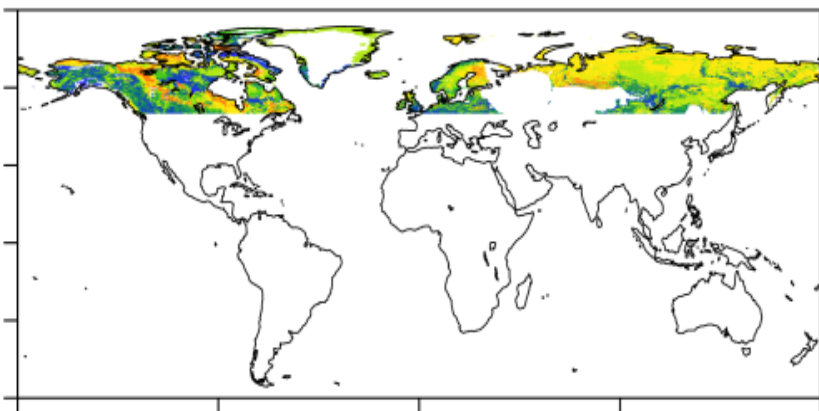
IGBP (900-1650 PgC, to 1m)



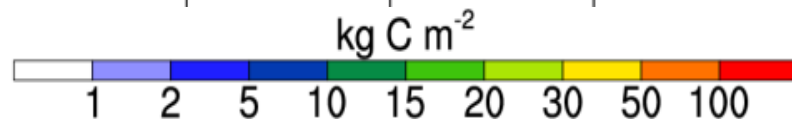
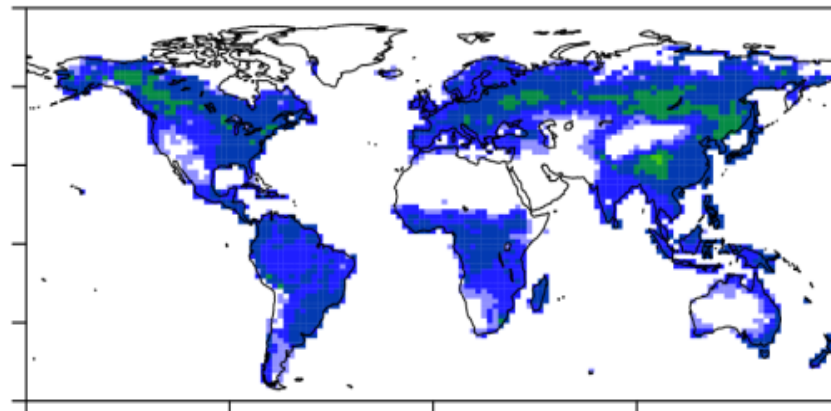
CLM4.5BGC (to 1m; 1900 PgC)



NCSCD (to 1m)



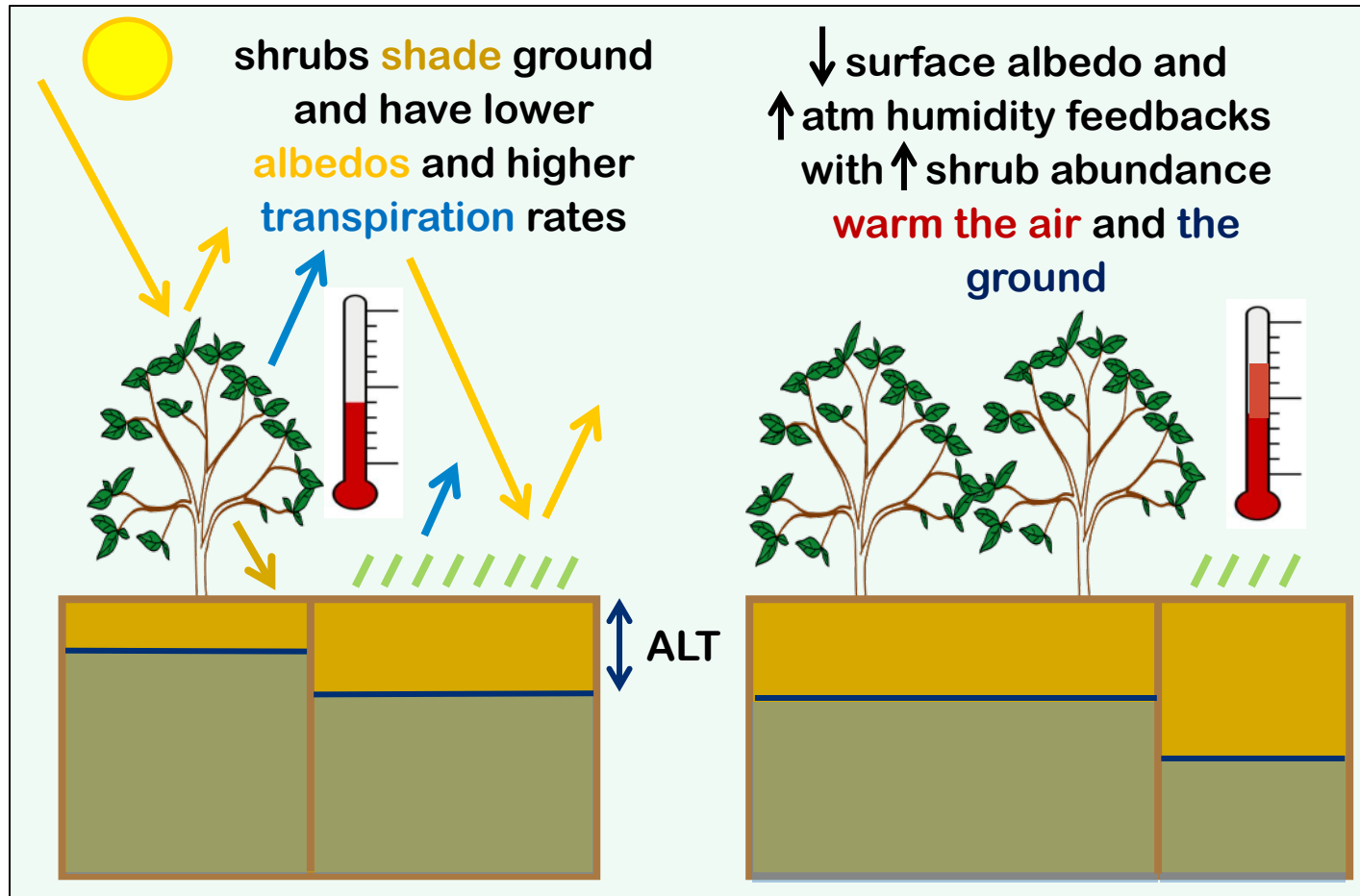
CLM4CN (650 PgC)





Summary (Lawrence and Swenson, ERL, 2011)

Will expanding Arctic shrub cover decrease permafrost vulnerability to climate change?



A. Not necessarily. Depends on whether the direct local cooling or the indirect climate warming dominates. Our results indicate that **shrub expansion may increase rather than decrease permafrost vulnerability** to climate change.



What is permafrost?

Definition: Soil or rock that remains below 0°C for two or more consecutive years