

# Atmospheric Modeling I: Physics in the Community Atmosphere Model (CAM)

CESM Tutorial

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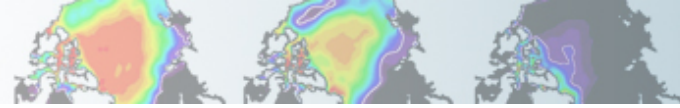
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NCAR is sponsored by the National Science Foundation



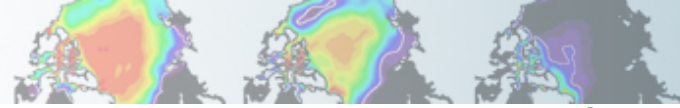
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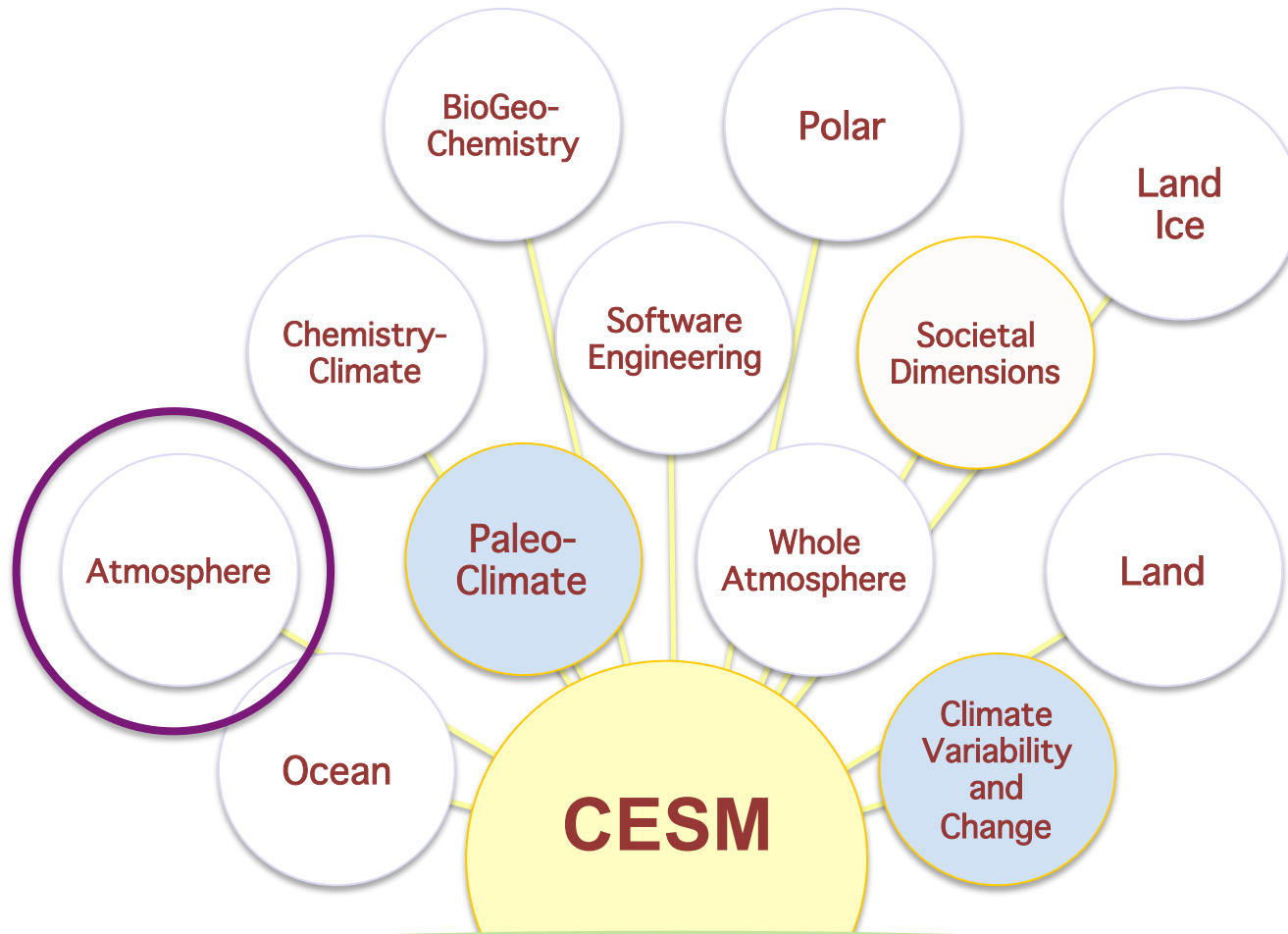
# Outline

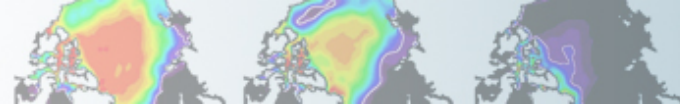
- **Physical processes** in an atmosphere GCM
- Distinguishing GCMs from other models (scales)
- Concept of **'Parameterization'**
- Physics representations (CESM)
  - **Clouds** (different types), cloud fraction and microphysics
  - Radiation
  - Boundary layers, surface fluxes and gravity waves
- **Process interactions**
- Model **complexity, sensitivity** and **climate feedbacks**



# Community Atmosphere Model (CAM)

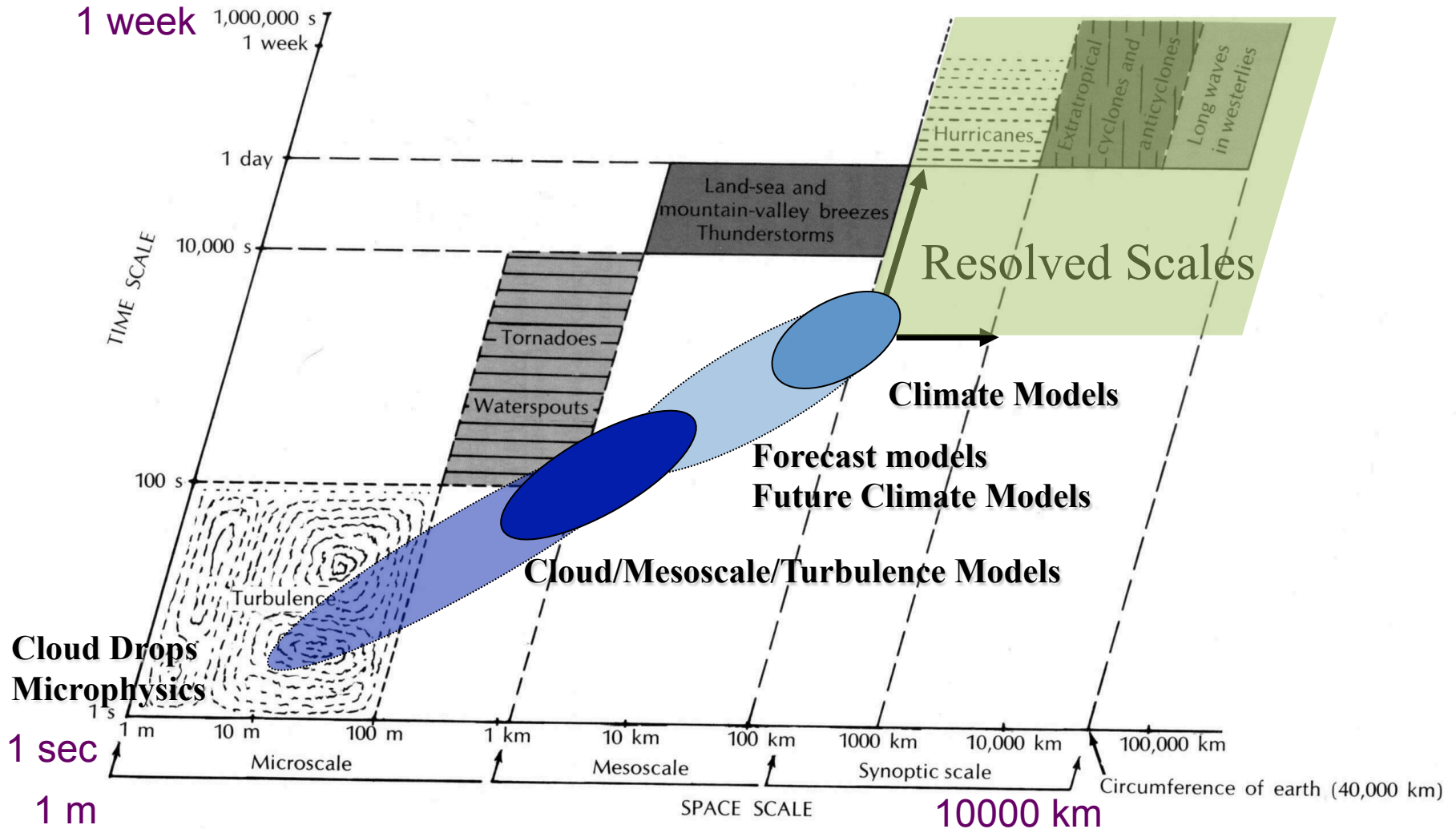
## Atmosphere Model Working Group (AMWG)



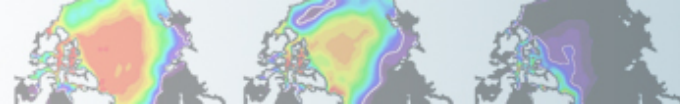


# Scales of Atmospheric Processes

Determines the formulation of the model







# Hydrostatic Primitive Equations

Where do we put the physics?

Horizontal scales  $\gg$  vertical scales

Vertical acceleration  $\ll$  gravity

$$d\bar{\mathbf{V}}/dt + f\mathbf{k} \times \bar{\mathbf{V}} + \nabla\bar{\phi} = \mathbf{F}, \quad \mathbf{F}_V \quad (\text{horizontal momentum})$$

$$d\bar{T}/dt - \kappa\bar{T}\omega/p = Q/c_p, \quad F_T \quad (\text{thermodynamic energy})$$

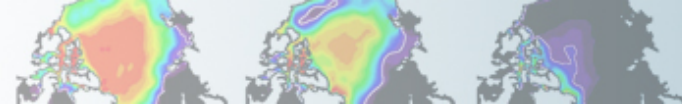
$$\nabla \cdot \bar{\mathbf{V}} + \partial\bar{\omega}/\partial p = 0, \quad (\text{mass continuity})$$

$$\partial\bar{\phi}/\partial p + R\bar{T}/p = 0, \quad (\text{hydrostatic equilibrium})$$

$$d\bar{q}/dt = S_q. \quad \mathbf{F}_{QV}, \mathbf{F}_{QL}, \mathbf{F}_{QI} \quad (\text{water vapor mass continuity})$$

+transport

Harmless looking terms  $\mathbf{F}$ ,  $Q$ , and  $S_q \implies$  “physics”



# What is a 'Parameterization'?

- Usually based on
  - Basic physics (conservation laws of thermodynamics)
  - Empirical formulations from observations
- In many cases: no explicit formulation based on first principles is possible at the level of detail desired. Why?
  - Non-linearities & interactions at 'sub-grid' scale
  - Often coupled with observational uncertainty
  - Insufficient information in the grid-scale parameters



Vertical eddy transport of  $\chi$

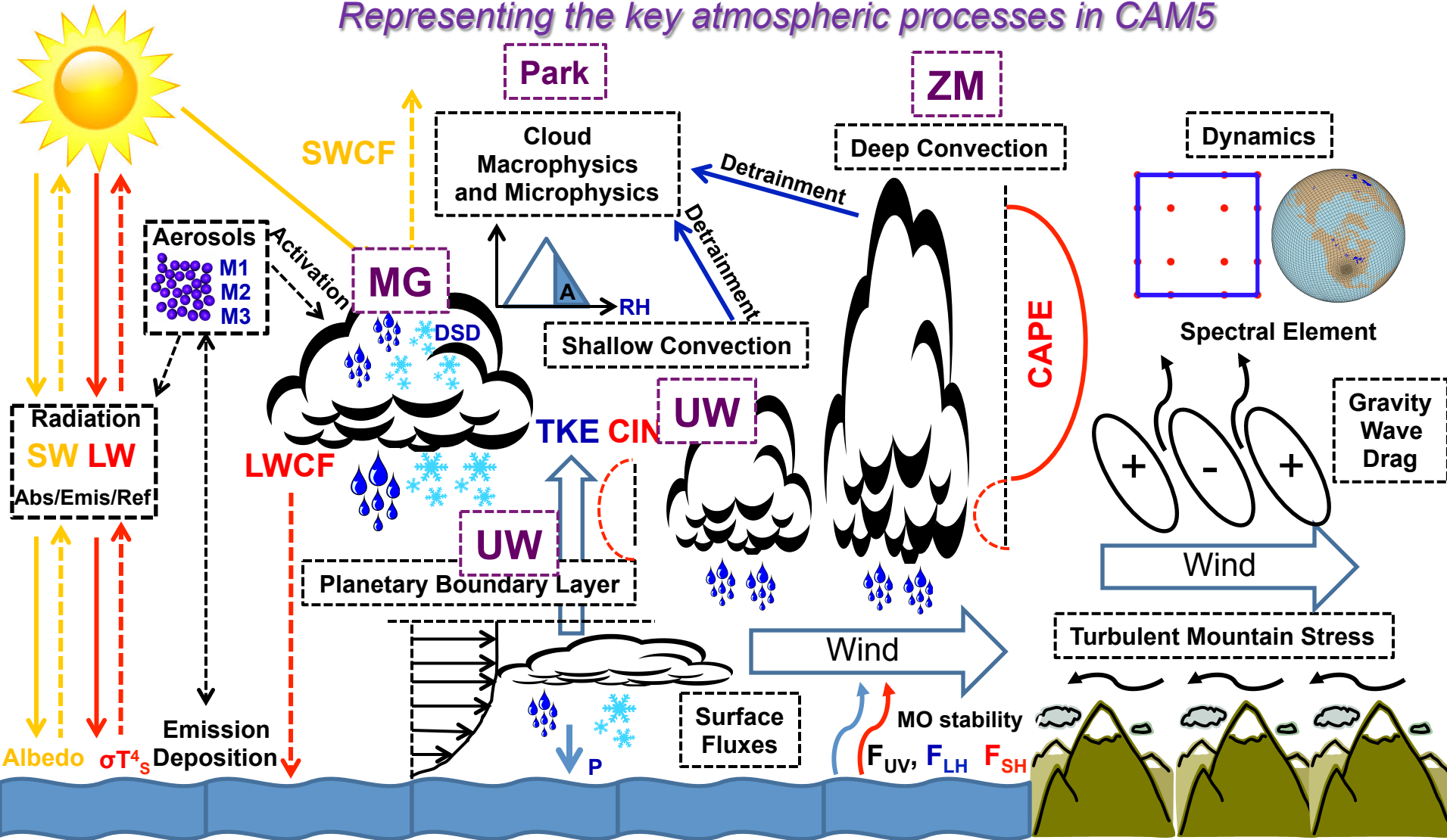
$$\overline{w'\chi'} = -K_x \frac{\partial \chi}{\partial z}$$

↑ **Unresolved 'sub-grid'**
↙ **'Diffusivity'**
↘ **Resolved 'grid-scale'**

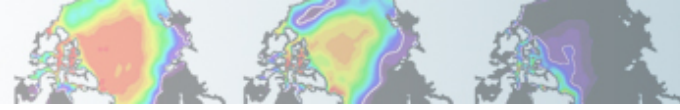


# Community Atmosphere Model

Representing the key atmospheric processes in CAM5

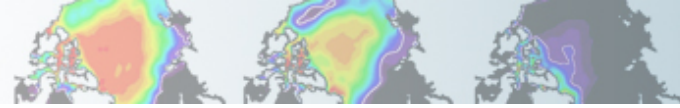






# Clouds





# Clouds

## Multiple Categories

- **Stratiform (large-scale) clouds**

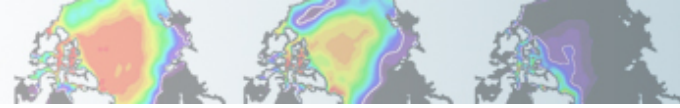
- Responds to large-scale saturation fraction, RH (parameterized)
- Coupled to presence of condensate (microphysics, advection)

- **Shallow convection clouds**

- Symmetric turbulence in lower troposphere
- Non precipitating (mostly)
- Responds to surface forcing

- **Deep convection clouds**

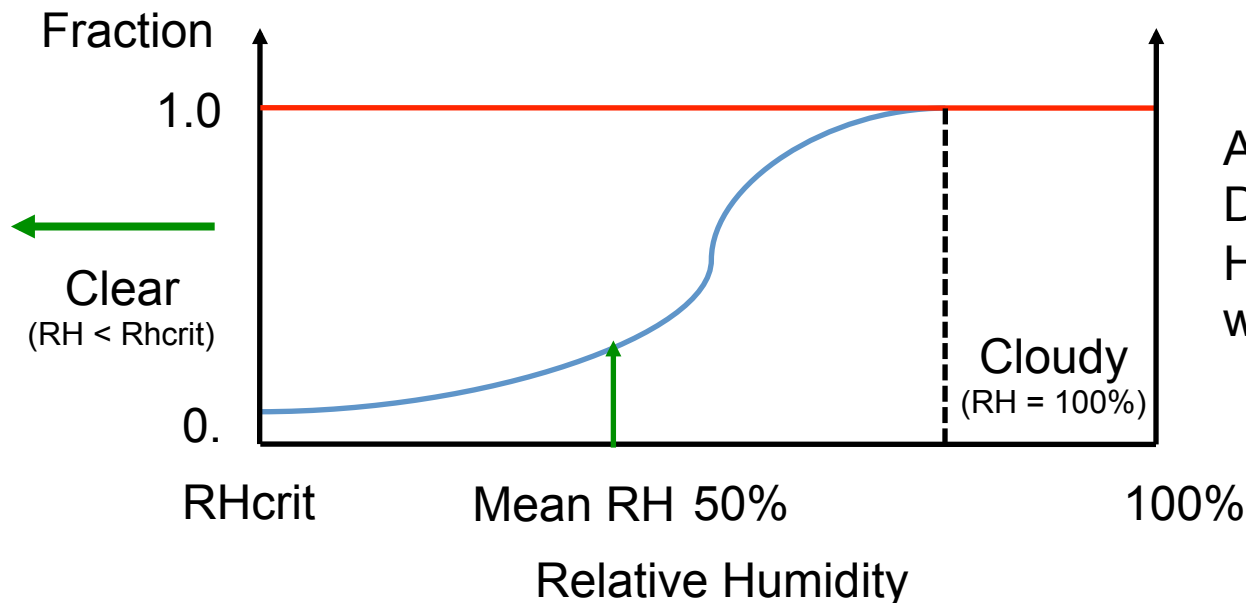
- Asymmetric turbulence
- Penetrating convection (surface -> tropopause)
- Precipitating
- Responds to surface forcing and conditional instability



# Stratiform Clouds (macrophysics)

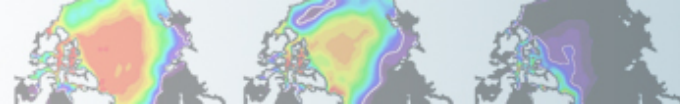
## Sub-Grid Humidity and Clouds

- ✓ Liquid clouds form when  $RH = 100\%$  ( $q = q_{sat}$ )
- ✓ But if there is variation in RH in space, some clouds will form before *mean*  $RH = 100\%$
- ✓  $RH_{crit}$  determines cloud fraction  $> 0$ ; Value is lower over land due to higher humidity variance



Assumed Cumulative Distribution function of Humidity in a grid box with sub-grid variation





# Shallow and Deep Convection

## Exploiting conservation properties

### Common properties

Parameterize consequences of vertical displacements of air parcels

**Unsaturated:** Parcels follow a dry adiabat (conserve **dry static energy**)

**Saturated:** Parcels follow a moist adiabat (conserve **moist static energy**)

### Shallow (10s-100s m)

Parcels remain stable (buoyancy $<0$ )

Shallow cooling mainly

Some latent heating and precipitation

Generally a source of water vapor

Small cloud radius large entrainment



### Deep (100s m-10s km)

Parcels become unstable (buoyancy $>0$ )

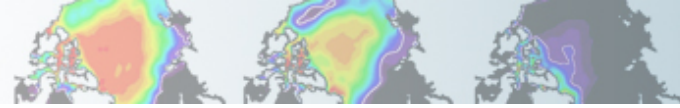
Deep heating

Latent heating and precipitation

Generally a sink of water vapor

Large cloud radius small entrainment



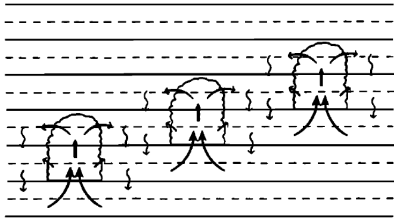


# Shallow and Deep Convection

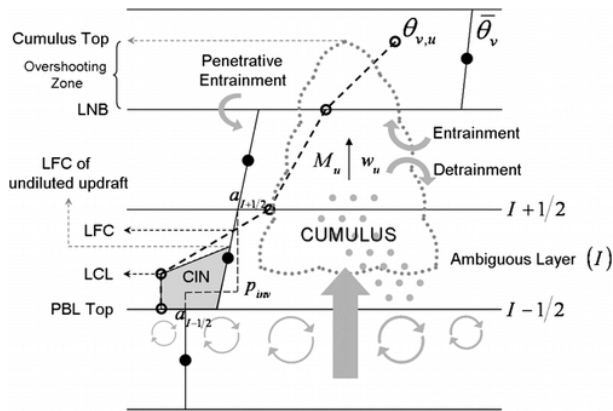
## Closure: How much and when?

### Shallow

Local conditional instability **CAM4**



Convective inhibition and turbulent kinetic energy (TKE) **CAM5**

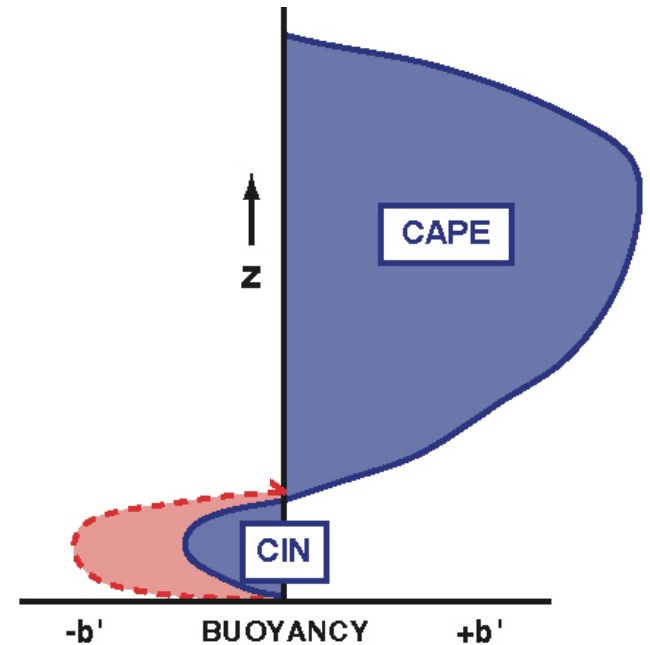


### Deep

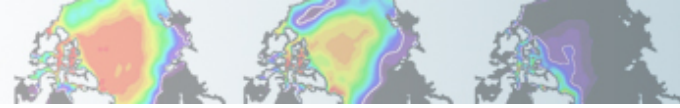
Convective Available Potential Energy (CAPE) **CAM4 and CAM5**

$$CAPE > CAPE_{\text{trigger}}$$

Timescale=1 hour



Shallow and deep convection and stratiform cloud fractions combined for radiation



# Cloud Microphysics

- **Condensed phase water processes**

- Properties of condensed species (=liquid, ice)
  - size distributions, shapes
- Distribution/transformation of condensed species
  - Precipitation, phase conversion, sedimentation

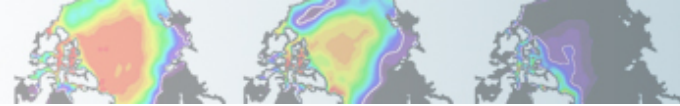
- **Important for other processes:**

- Aerosol scavenging
- Radiation



- **In CAM = 'stratiform' cloud microphysics**

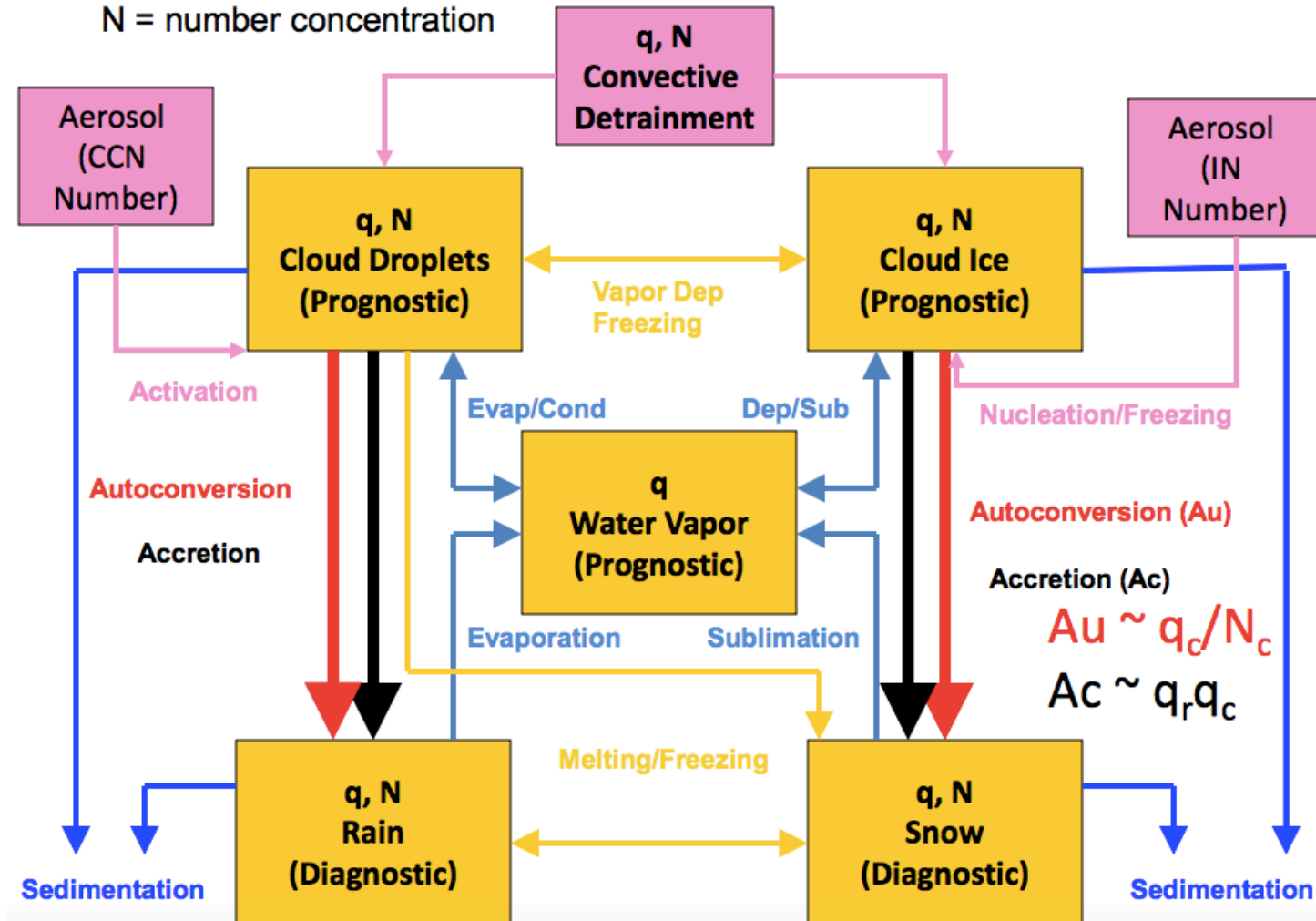
- Convective microphysics simplified
- Formulations currently being implemented into convection

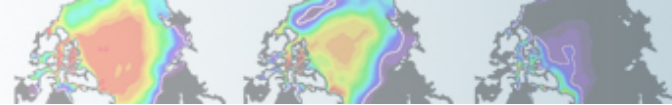


# CAM5 Microphysics

q = mixing ratio  
N = number concentration

Morrison & Gettelman 2008



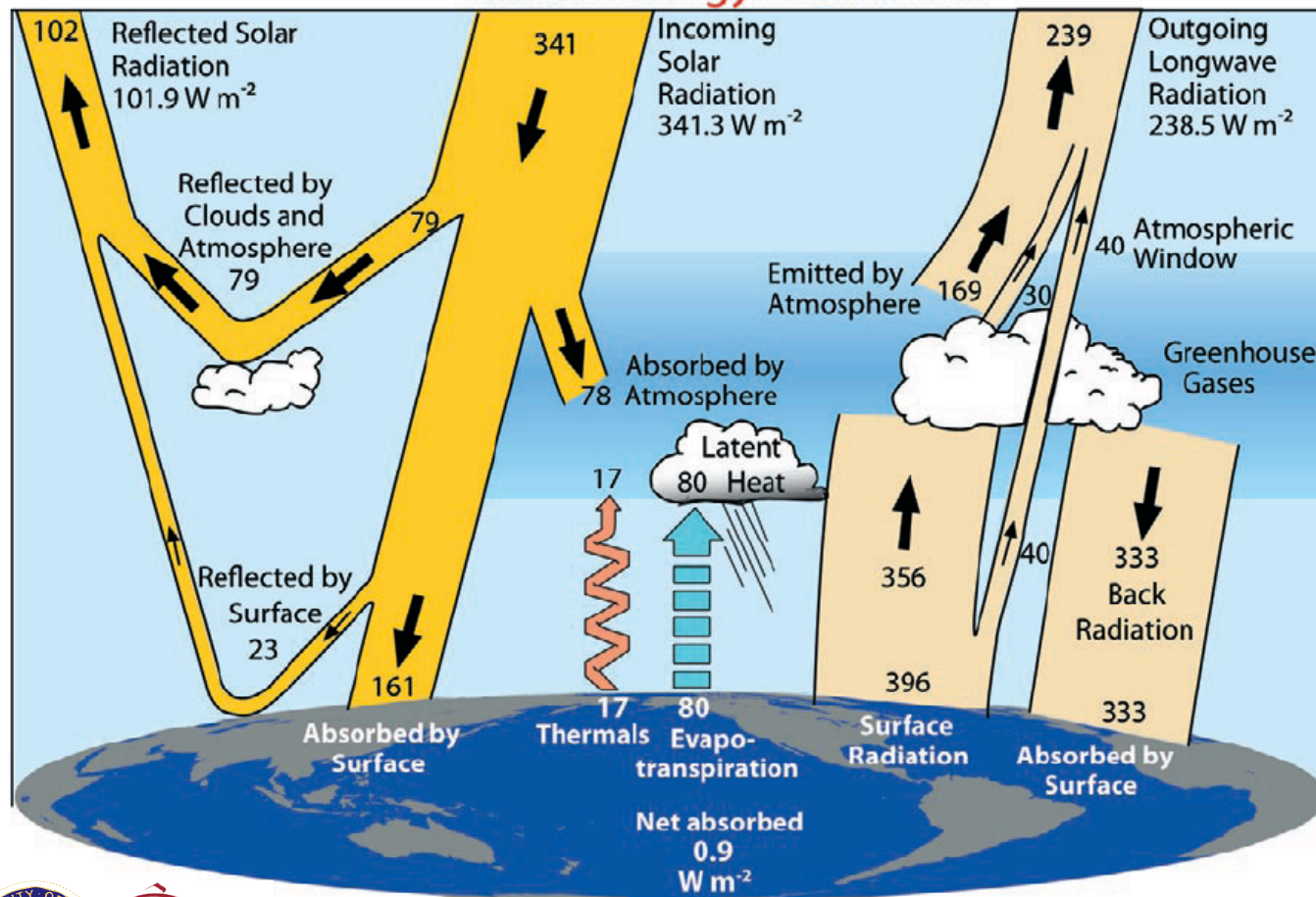


# Radiation

## The Earth's Energy Budget

Trenberth & Fasullo, 2008

Global Energy Flows  $W m^{-2}$

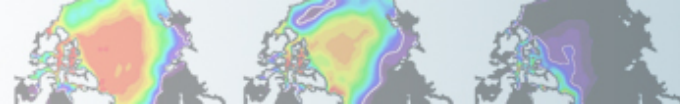


Gas	SW Absorption ( $W m^{-2}$ )
CO <sub>2</sub>	1
O <sub>2</sub>	2
O <sub>3</sub>	14
H <sub>2</sub> O	43



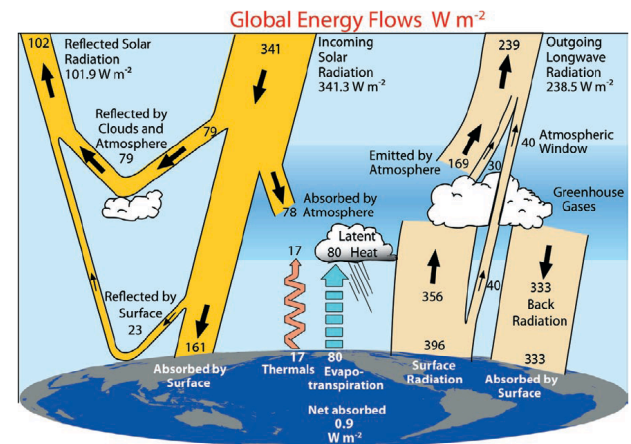
Thanks to: Bill Collins, Berkeley & LBL



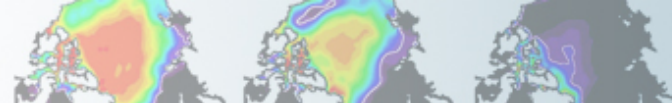


# Goals of GCM Radiation Codes

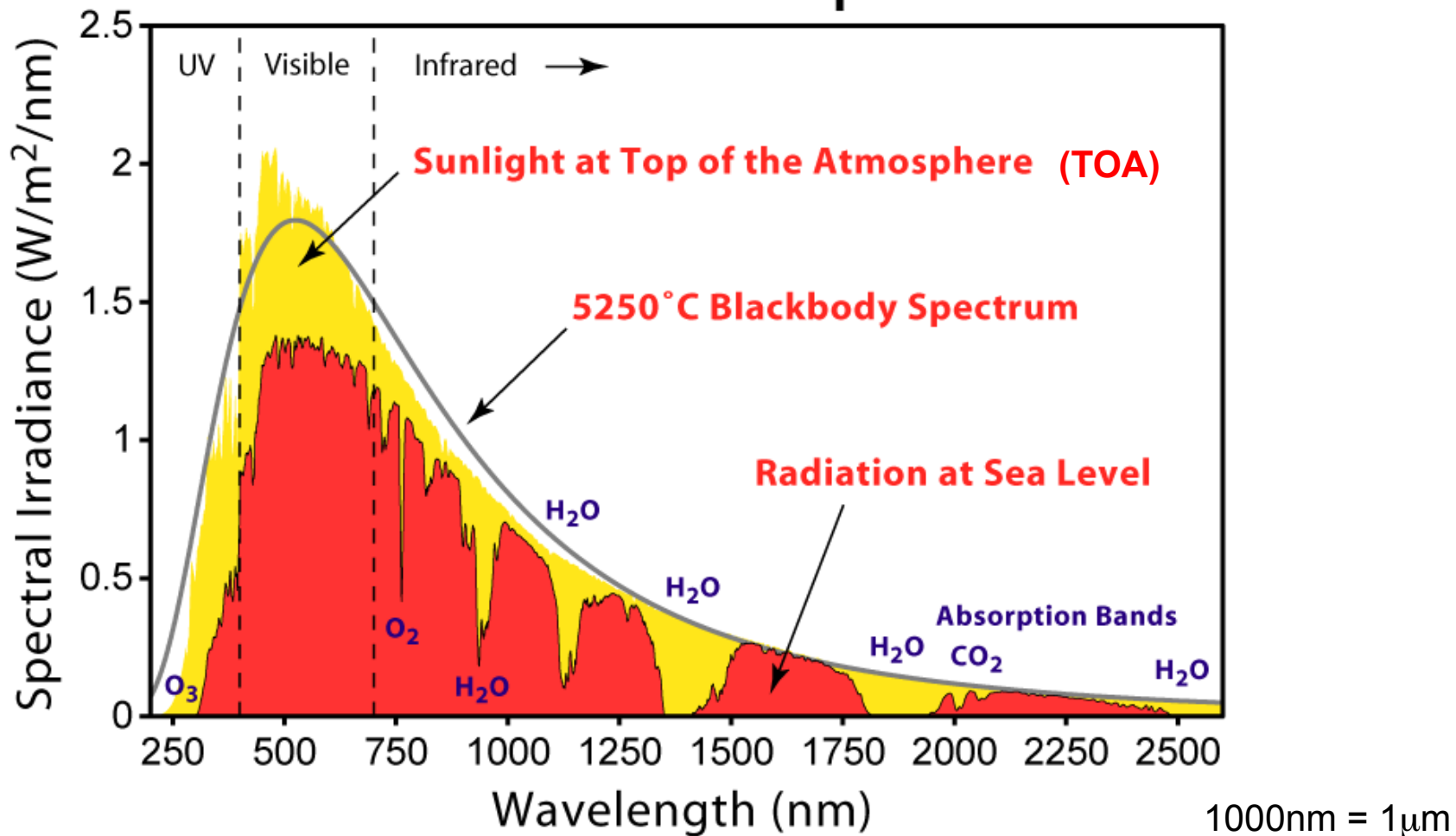
- Accurately represent the input and output of energy in the climate system and how it moves around
  - Solar Energy
  - Thermal Emission
  - Gases
  - Condensed species: Clouds & Aerosols



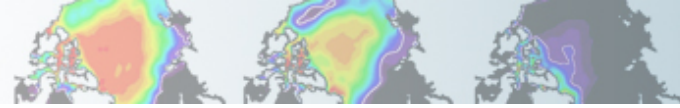




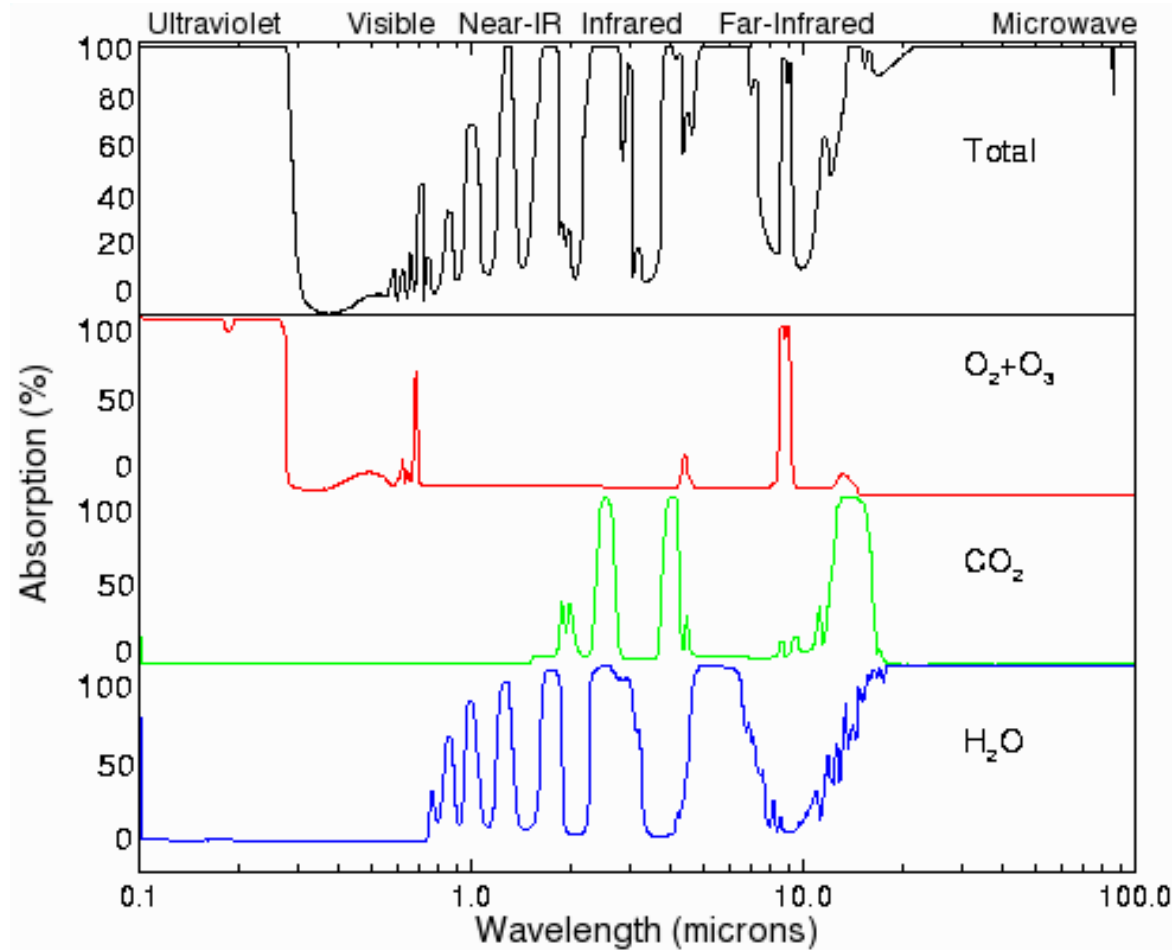
# Solar Radiation Spectrum



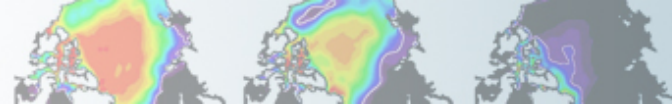
Input at TOA, Radiation at surface



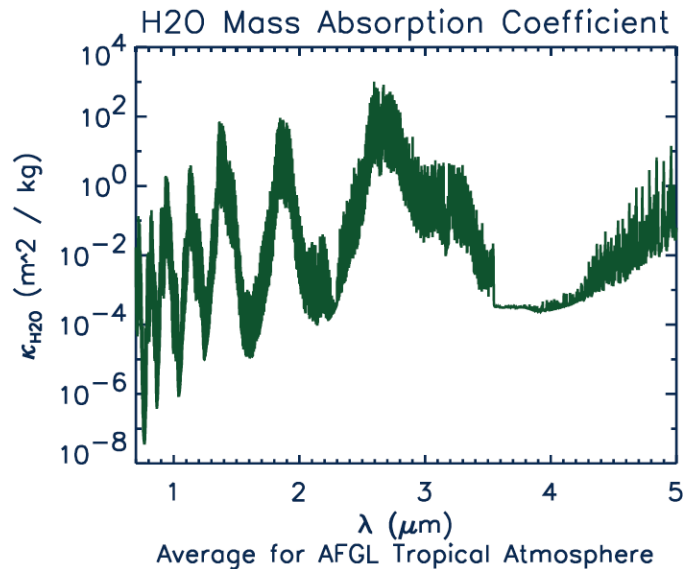
# IR absorption



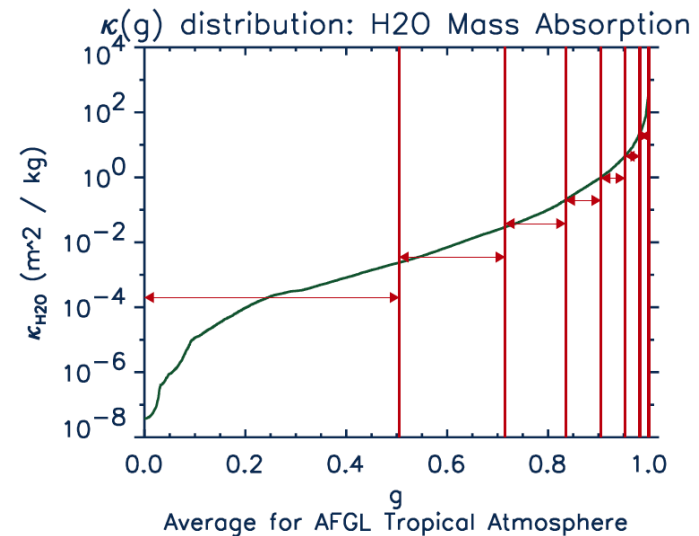
1000nm = 1 $\mu$ m



# k-distribution Band Models

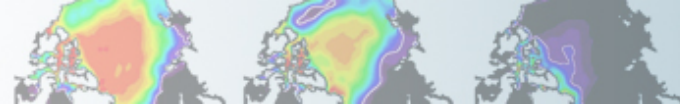


Sort



- Line-by-line calculations
- **Very expensive/slow, accurate**

- k-distribution band model, sort absorption coefficients by magnitude
- **Cheaper/fast, less accurate**



# Planetary Boundary Layer (PBL)

## Regime dependent representations

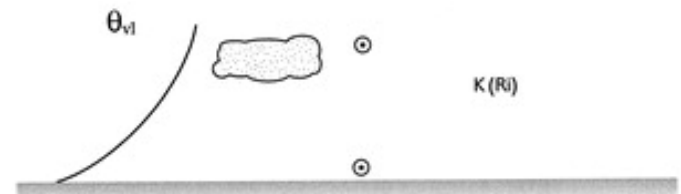
- Vital for near-surface environment (humidity, temperature, chemistry)
- Exploit **thermodynamic conservation** (liquid virtual potential temperature  $\theta_{vl}$ )
- **Conserved** for rapidly well mixed PBL
- **Not conserved** for stable PBL
- Critical determinant is the presence of turbulence

- **Richardson number** 
$$Ri = \frac{g\beta}{(\partial u / \partial z)^2},$$
- $\ll 1$ , flow becomes turbulent

- **CAM4**: Gradient Ri # + non-local transport (Holtslag and Boville, 1993)
- **CAM5**: TKE-based Moist turbulence (Park and Bretherton, 2009)

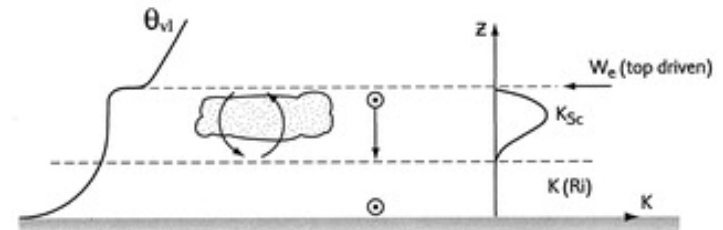
(a)

I. Stable boundary layer, possibly with non-turbulent cloud (no cumulus, no decoupled Sc, stable surface layer)



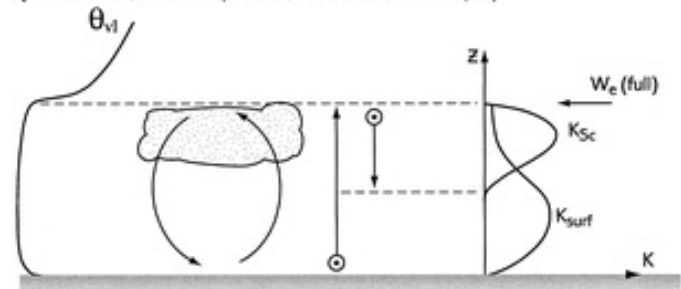
(b)

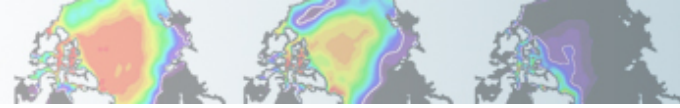
II. Stratocumulus over a stable surface layer (no cumulus, decoupled Sc, stable surface layer)



(c)

III. Single mixed layer, possibly cloud-topped (no cumulus, no decoupled Sc, unstable surface layer)



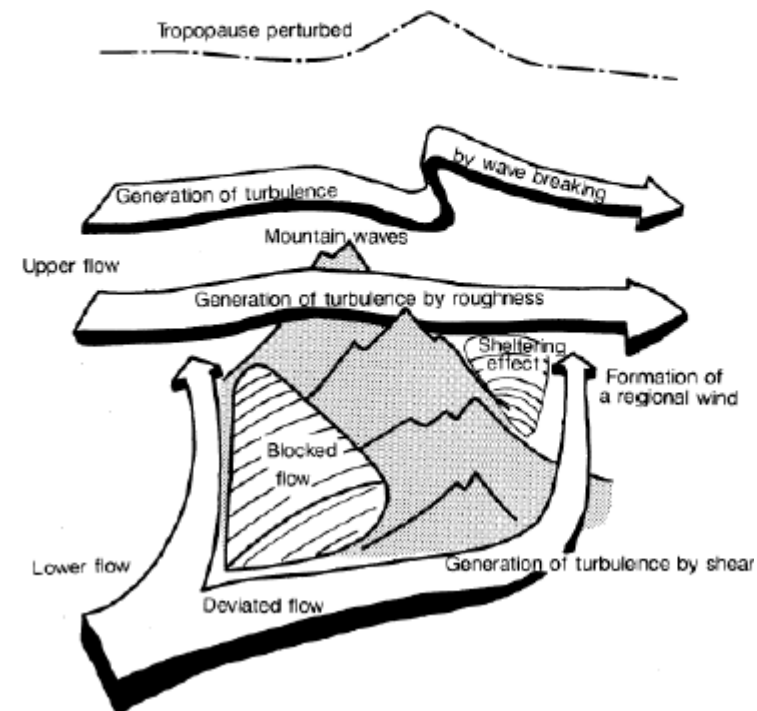


# Gravity Waves and Mountain Stresses

## Sub-grid scale dynamical forcings

- **Gravity Wave Drag**

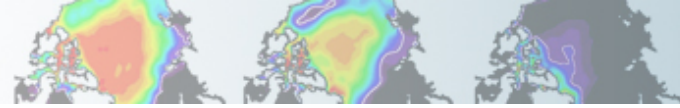
- Determines flow effect of upward propagating (sub-grid scale) gravity waves that break and dump momentum
- Generated by surface orography (mountains) and deep convection
- Important for closing off jet cores in the upper troposphere (strat/mesosphere)



- **Turbulent mountain stress**

- Local near-surface stress on flow
- Roughness length < scales < grid-scale
- Impacts mid/high-latitude flow (CAM5)

- **More difficult to parameterize than thermodynamic impacts (conservation?)**



# Surface Exchange

- Surface fluxes (bulk formulations)

## Stresses

$$\tau_x = -\rho_1 \overline{(u'w')} = -\rho_1 u_*^2 (u_1/V_a) = \rho_1 \frac{u_s - u_1}{r_{am}}$$

$$\tau_y = -\rho_1 \overline{(v'w')} = -\rho_1 u_*^2 (v_1/V_a) = \rho_1 \frac{v_s - v_1}{r_{am}}$$

## Specific Heat

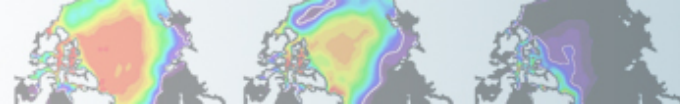
$$H = \rho_1 c_p \overline{(w'\theta')} = -\rho_1 c_p u_* \theta_* = \rho_1 c_p \frac{\theta_s - \theta_1}{r_{ah}}$$

## Latent heat (evaporation)

$$E = \rho_1 \overline{(w'q')} = -\rho_1 u_* q_* = \rho_1 \frac{q_s - q_1}{r_{aw}}$$

- Resistances  $r_{ax}$  based on
  - Monin-Obhukov similarity theory

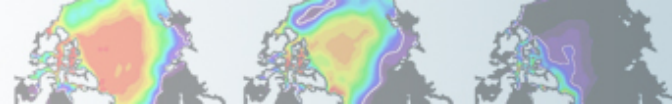




# Parameterization Interactions

## Direct and Indirect Process Communication

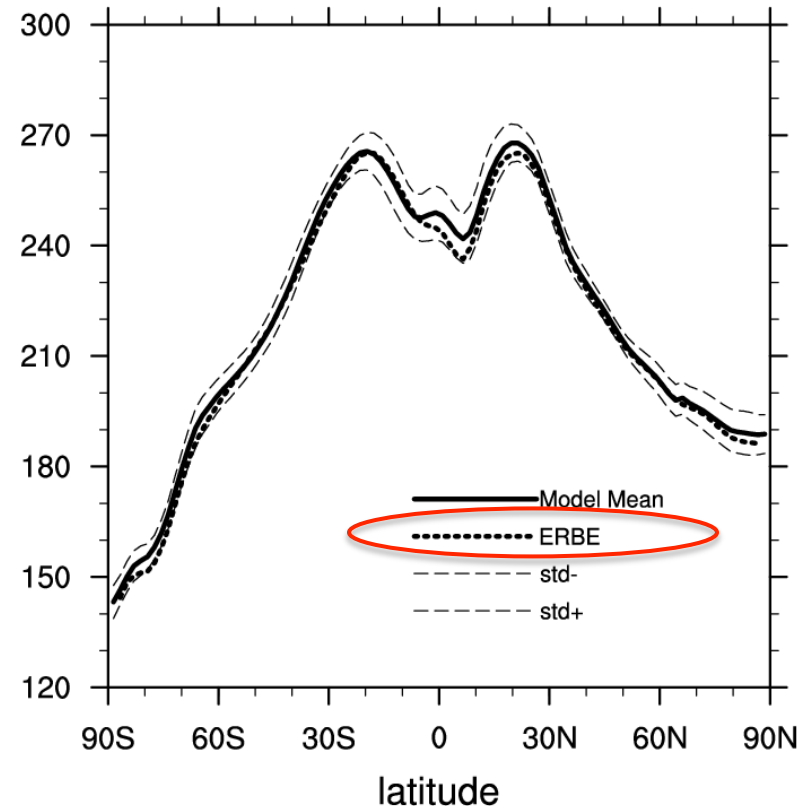
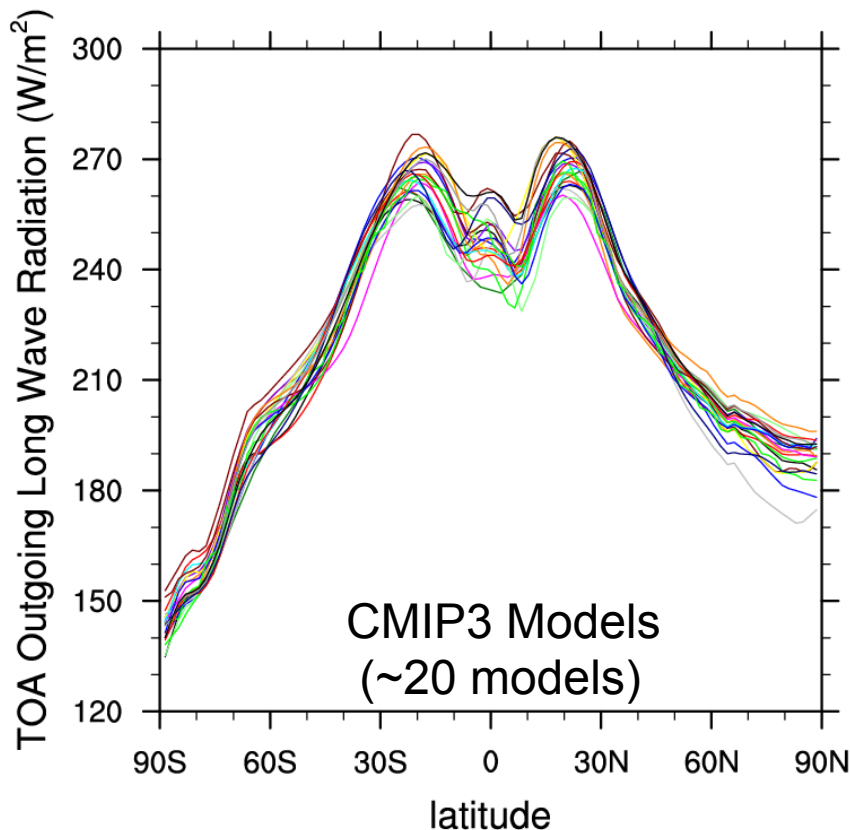
- Cloud Processes & Radiation
  - Feedbacks
- Boundary Layer / Cumulus & Dynamics
- Precipitation & Scavenging
  - Chemical (gas phase) constituents
  - Aerosols (condensed phase constituents)
- Microphysics and Aerosols
- Physics and surface components (ice, land, ocean)
- Resolved scales and unresolved scales

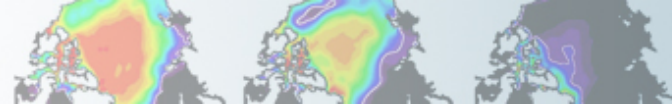


# Clouds in GCMs

## State of the Art from CMIP3

Outgoing Long-wave Radiation  
(Annual, 1990-1999)

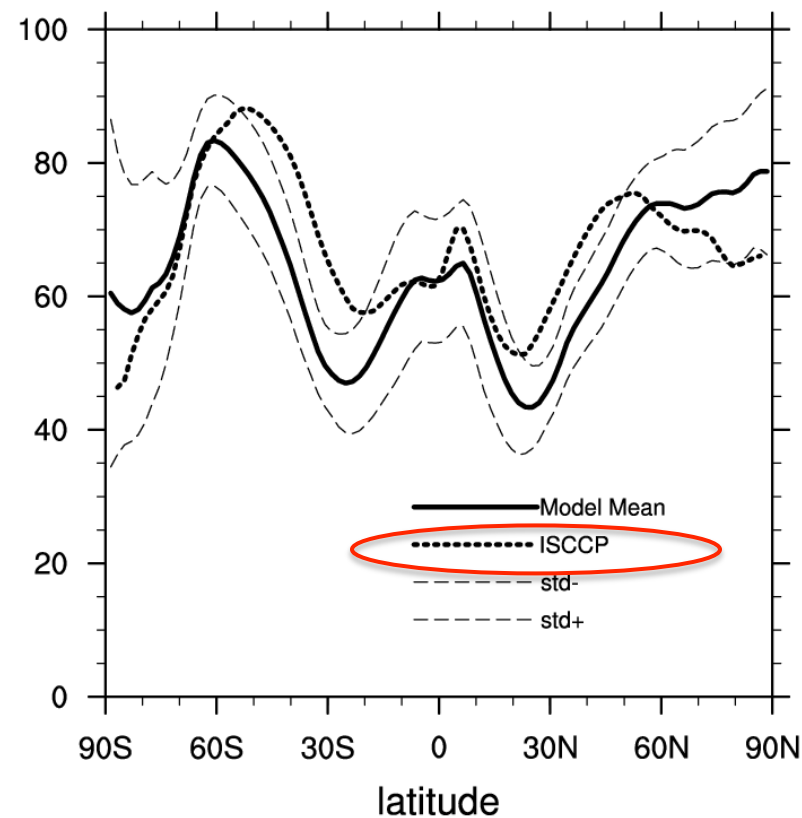
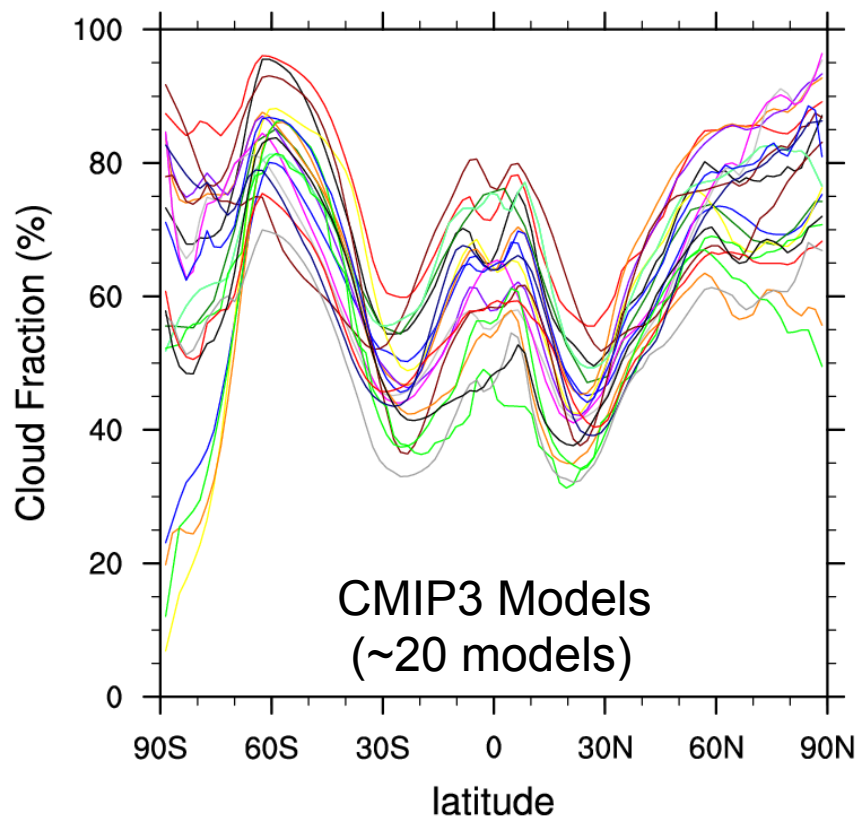


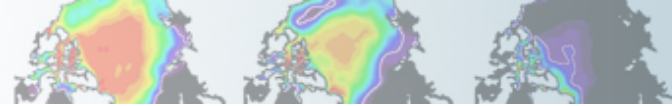


# Clouds in GCMs

## State of the Art from CMIP3

Total Cloud Fraction  
(Annual, 1990-1999)

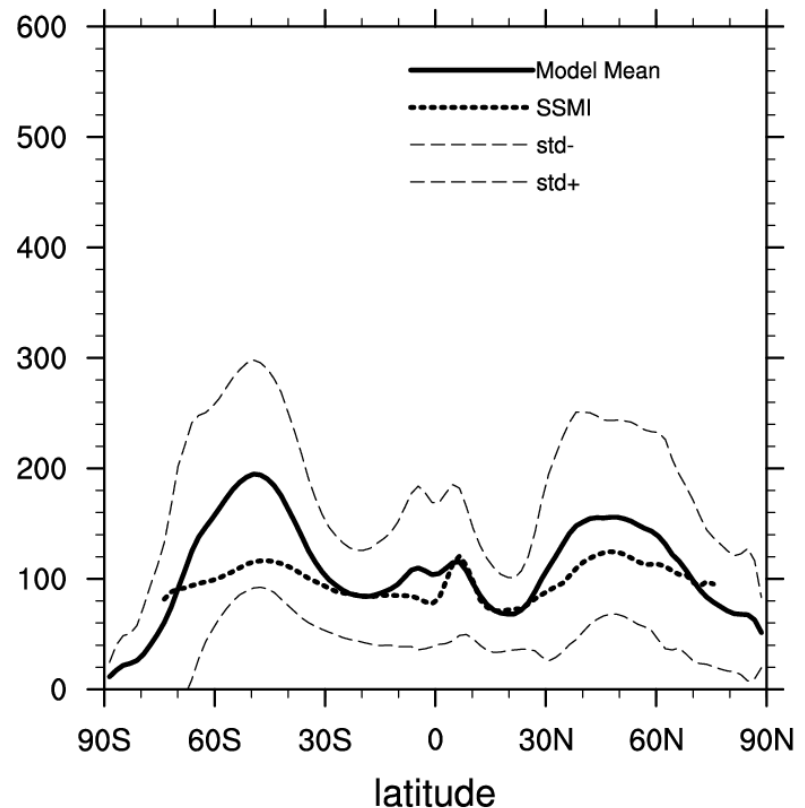
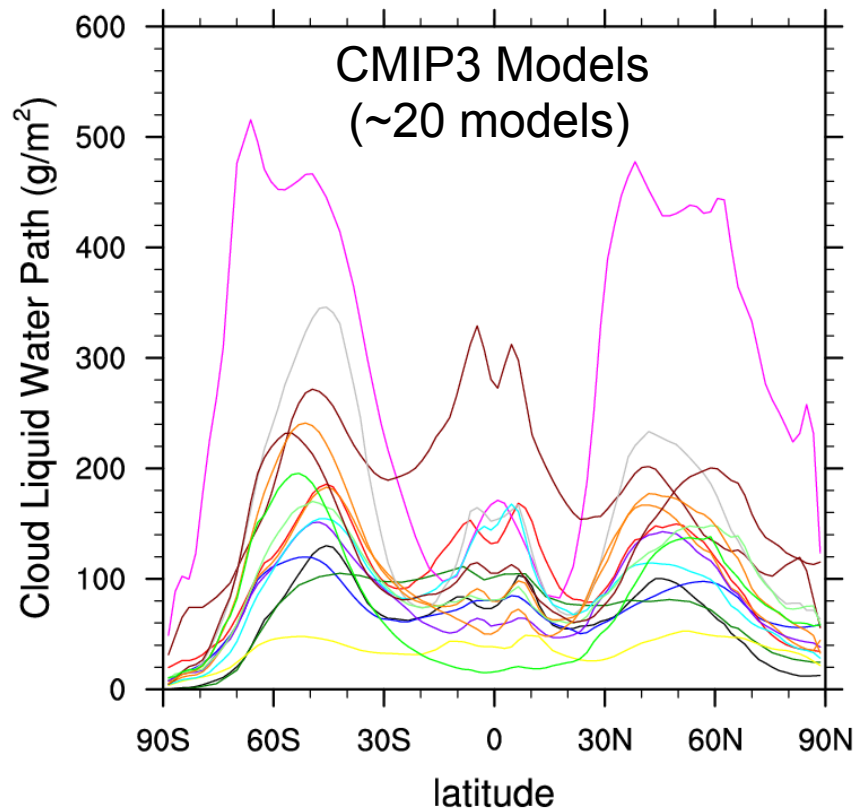




# Structural Clouds in GCMs

## State of the Art from CMIP3

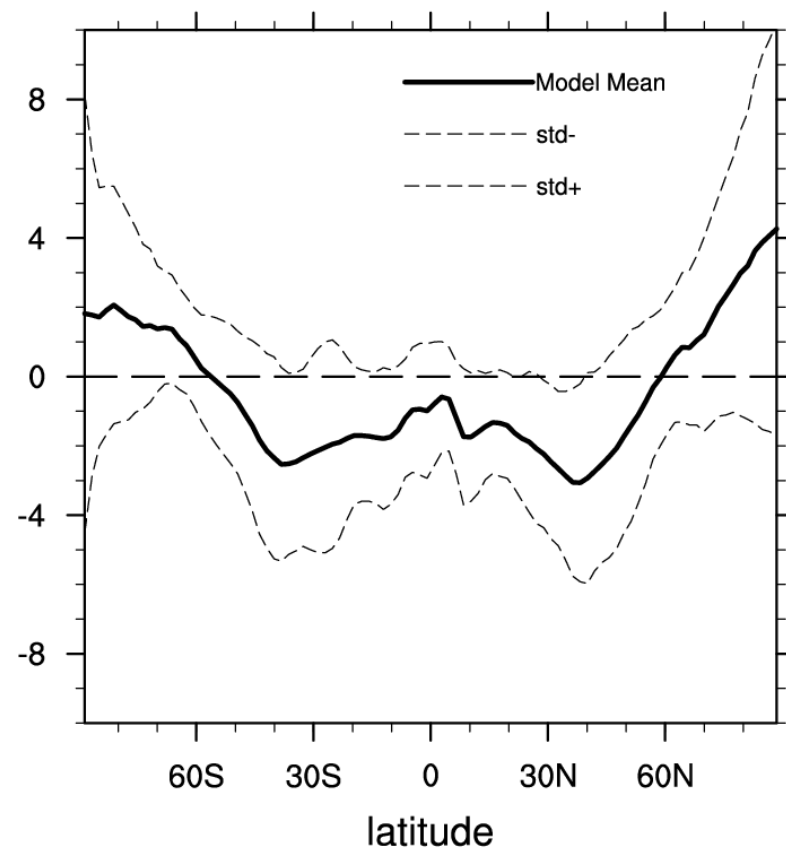
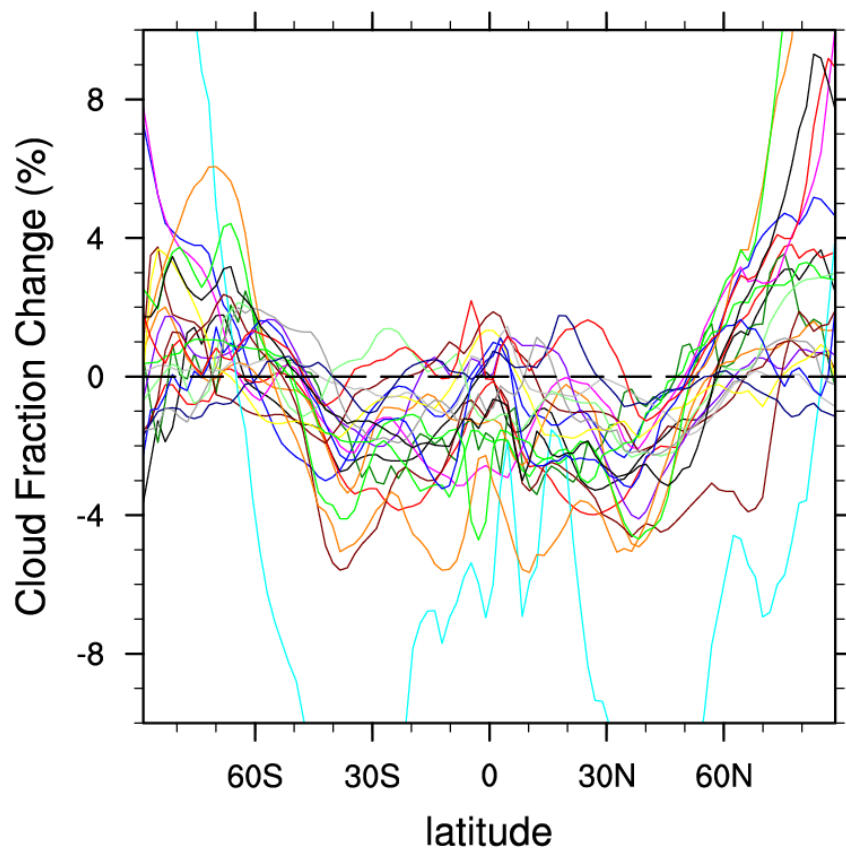
Liquid Water Path  
(Annual, 1990-1999)

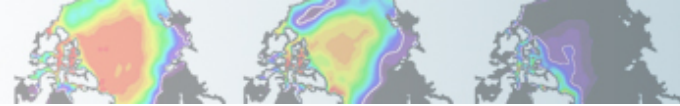


# Future Clouds in GCMs

State of the Art from CMIP3 – response to climate change

Total Cloud Fraction Change  
(Annual, SRES A1B: 2090-2099 minus 1990-1999)

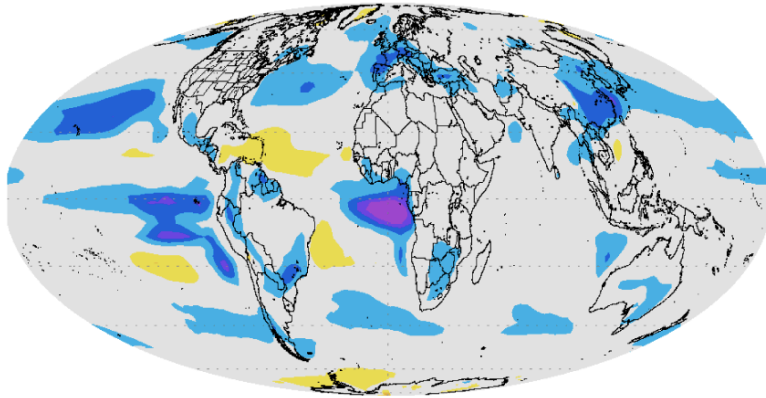




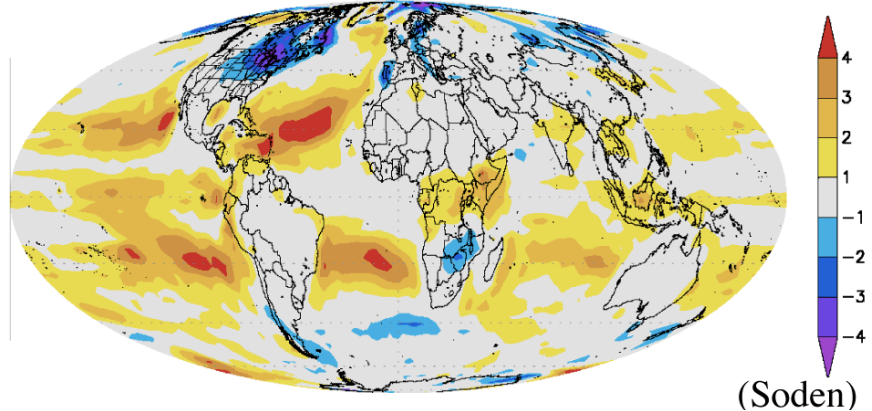
# Climate Sensitivity

What happens to clouds when we double CO<sub>2</sub>?

GFDL Model **+4.2K**



NCAR Model **+1.8K**



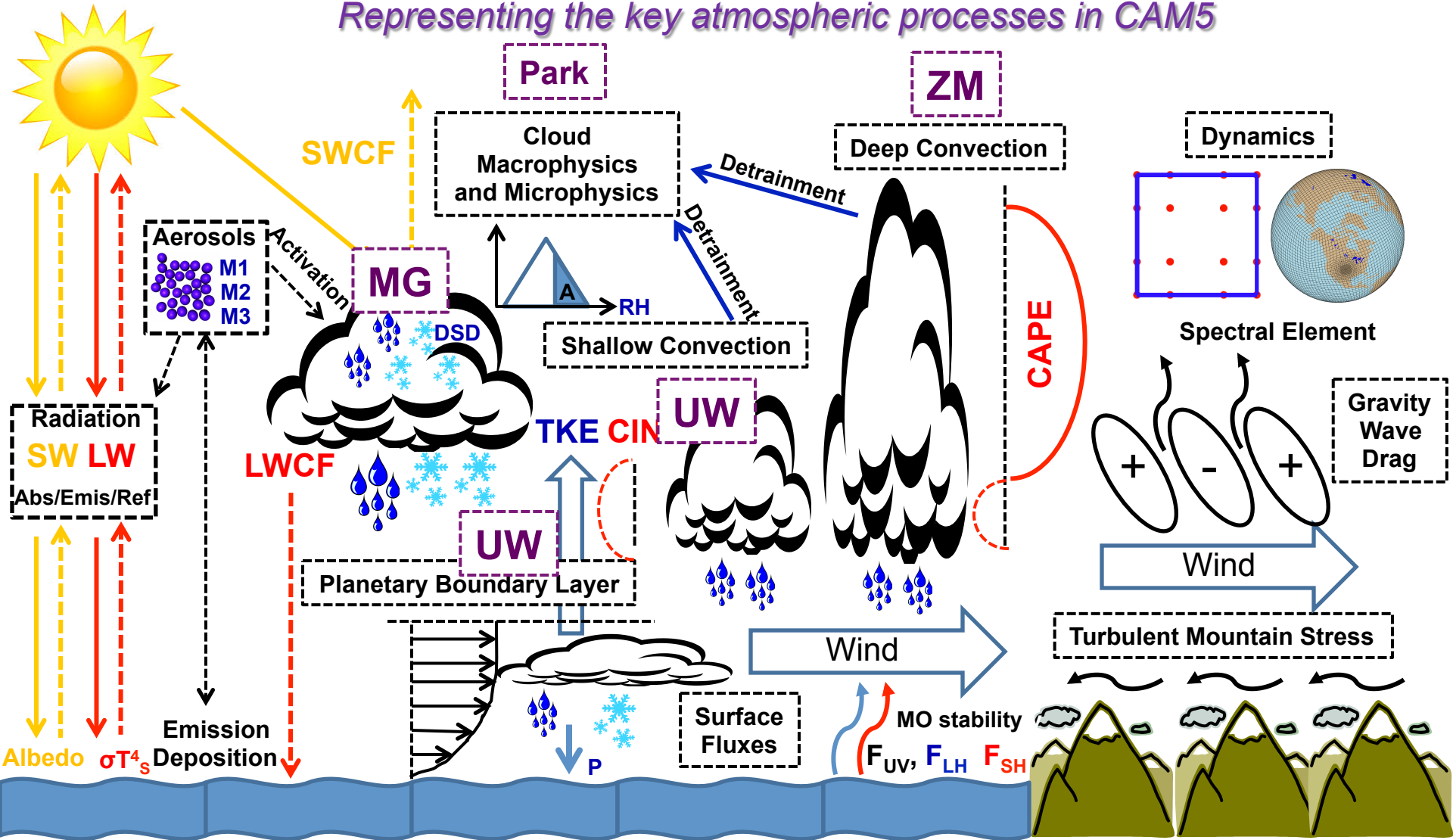
Change in low cloud amount (%)

- Significant range in **low-cloud sensitivity** (low and high end of models)
- Cloud regimens are largely **oceanic stratocumulus** (difficult to model)
- Implied temperatures change is due to (higher/lower) solar radiation reaching the ground because of **clouds feedbacks**.



# Community Atmosphere Model

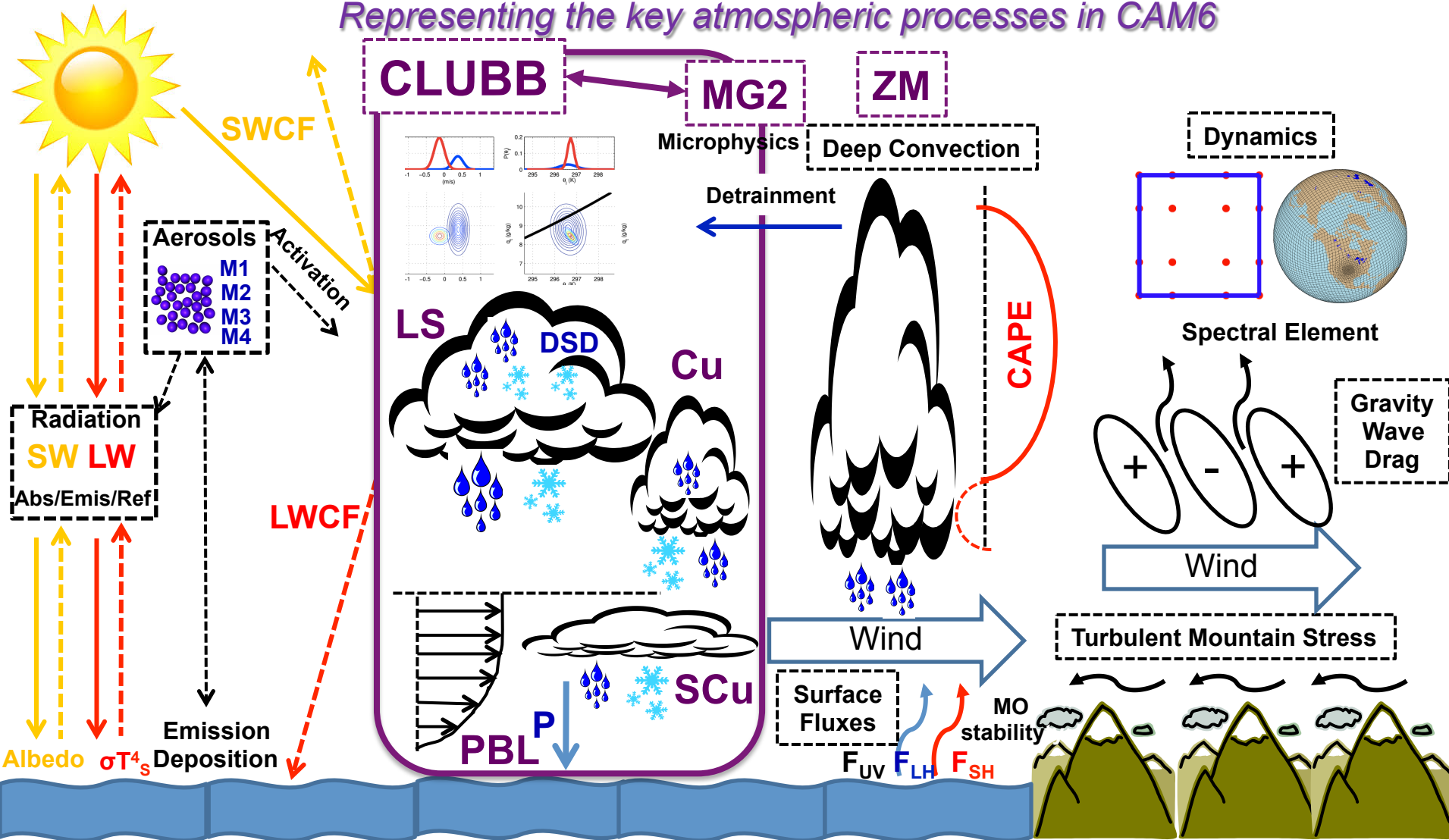
Representing the key atmospheric processes in CAM5

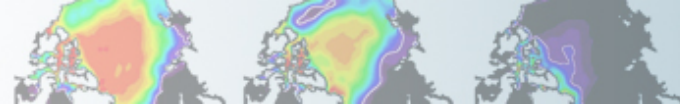




# Community Atmosphere Model

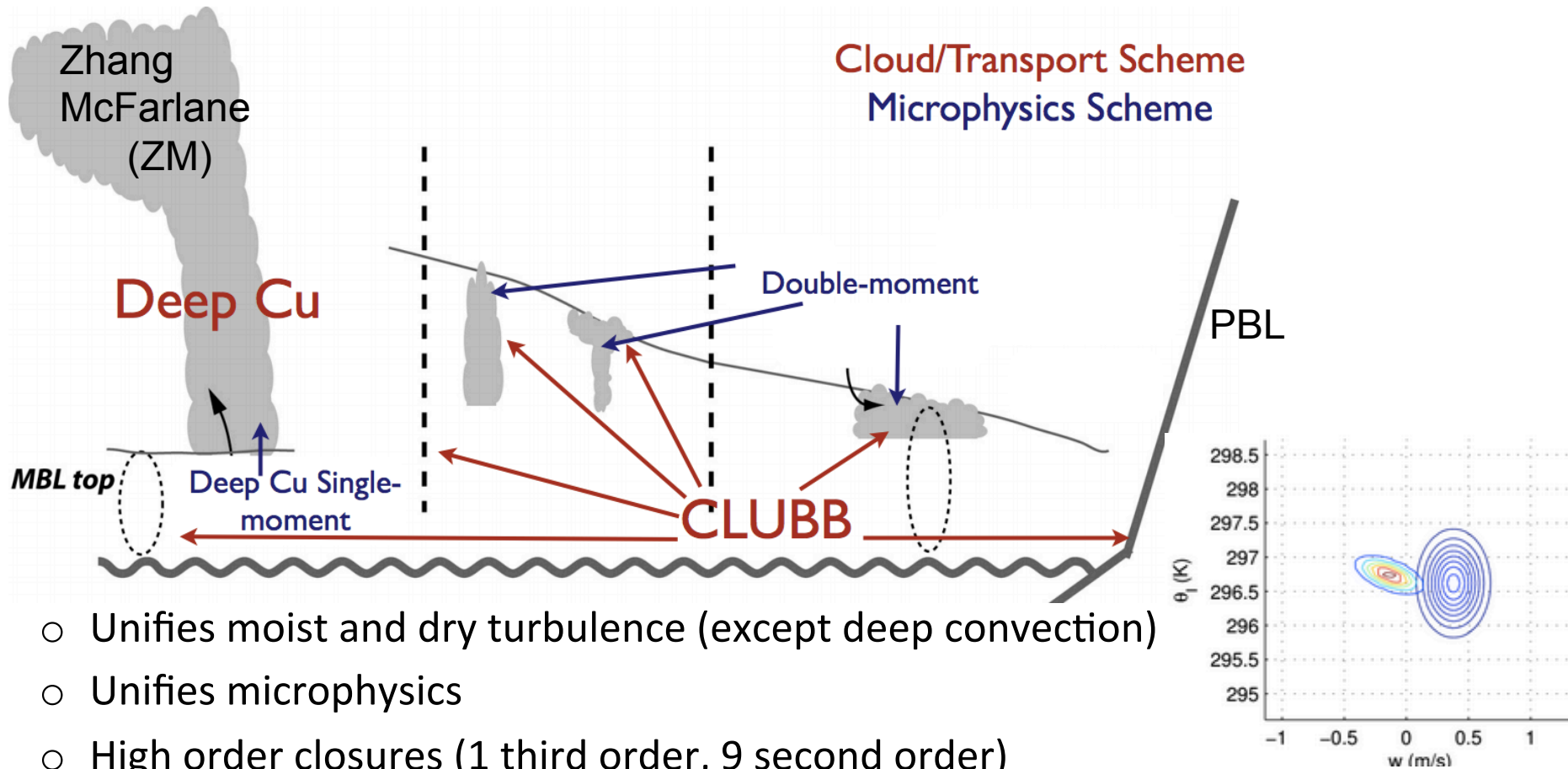
Representing the key atmospheric processes in CAM6



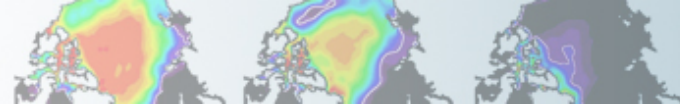


## CLUBB: Cloud Layers Unified By Binormals

Golaz 2002b, J. Atmos. Sci.



- Unifies moist and dry turbulence (except deep convection)
- Unifies microphysics
- High order closures (1 third order, 9 second order)
- Use two Gaussians to described the sub-grid multivariate PDF:  $P=P(w, q_t, \theta_L)$

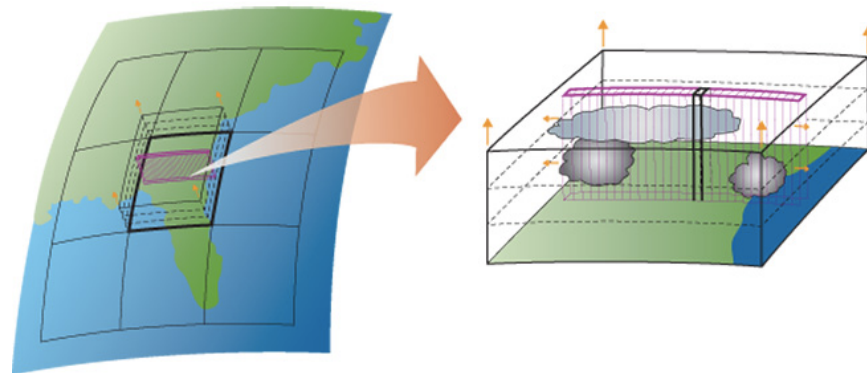


# Model physics: The future

- 1. How to operate in varying grid scale environments
- 2. Advanced representation of processes.

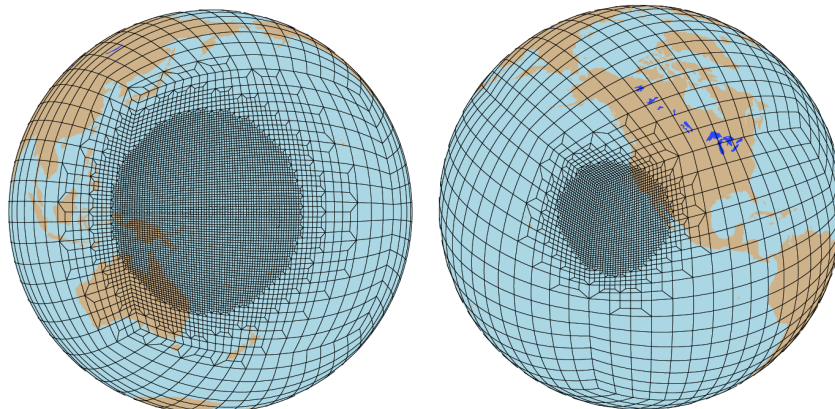
**Microphysics Species' Characteristics**

<p><b>CLOUD WATER</b></p> <ul style="list-style-type: none"> <li>• Gamma distribution with shape factor dependent on droplet concentration</li> <li>• Does not sediment</li> <li>• "Autoconverts" to rain using Berry &amp; Reinhardt with dependence on droplet concentration</li> </ul>	<p><b>CLOUD ICE</b></p> <ul style="list-style-type: none"> <li>• Gamma distribution</li> <li>• "Pristine" ice (<math>D &lt; 125</math> microns)</li> <li>• Initiation <math>T</math>-dependent (Cooper)</li> <li>• Prognosed <math>N_i</math></li> <li>• Slowly sediments at <math>\sim 10</math>-30 cm/s</li> </ul>
<p><b>RAIN</b></p> <ul style="list-style-type: none"> <li>• Gamma distribution</li> <li>• Variable equiv <math>y</math>-intercept: <math>2 \times 10^6 \text{m}^{-3}</math> (drizzle) <math>8 \times 10^6</math> (melted snow)</li> <li>• Accurate fallspeed relation</li> </ul>	<p><b>SNOW</b></p> <ul style="list-style-type: none"> <li>• Sum of two gamma distributions (Field et al. 2005)</li> <li>• Variable <math>y</math>-intercept depends on ice content and temperature</li> <li>• Non-spherical geometry (<math>m = ud^3</math>)</li> <li>• Variable snow density (<math>U/D</math>)</li> </ul>
<p><b>GRAUPEL</b></p> <ul style="list-style-type: none"> <li>• Gamma distribution</li> <li>• Variable equiv <math>y</math>-intercept depends on mixing ratio (simulate hail and snow-like graupel) <math>2 \times 10^6 \text{m}^{-4}</math> (graupel) <math>1 \times 10^4</math> (hail)</li> </ul>	

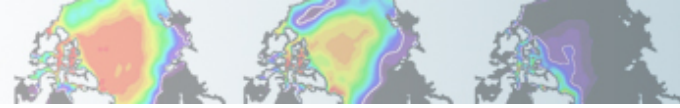


**New and more complex** processes

Cloud **super-parameterization**



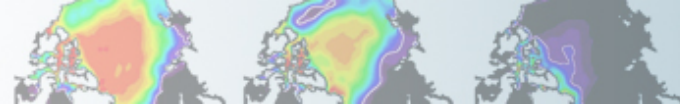
**High Resolution, Regional grid refinement and scale-aware** physics



# Summary

- GCMs physics=**unresolved processes**=**parameterization**
- Parameterization (CESM) = **approximating reality**
  - Starts from and maintains **physical constraints**
  - Tries to represent effects of smaller ‘sub-grid’ scales
- Fundamental constraints, **mass & energy conservation**
- Clouds are **fiendishly hard**: lots of **scales**, lots of **phase changes**, lots of **variability**
- **Clouds** are **coupled to radiation** (also hard) = biggest uncertainties (in future climate); largest dependencies
- CESM physics increasingly **complex** and **comprehensive**
- Future parameterizations aim to be process **scale-aware** and model **grid-scale independent**





Questions?