



# Isotope modeling with CESM

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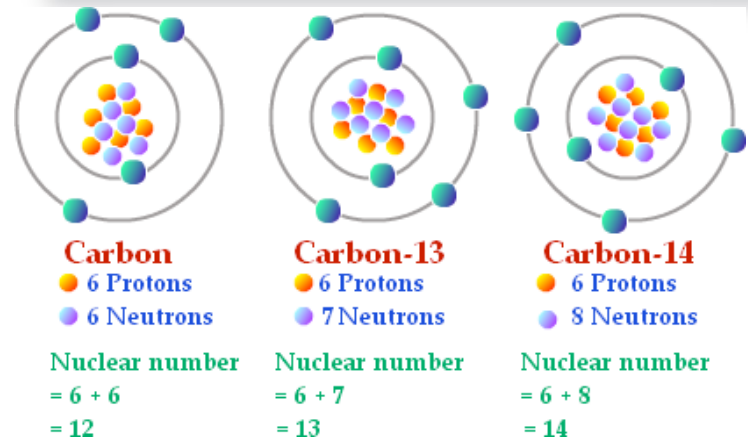
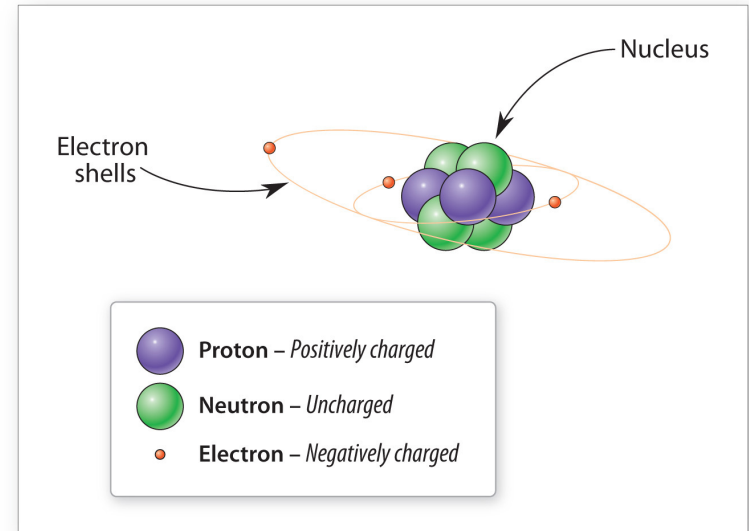
*With contributions from B. Otto-Bliesner and E. Brady (NCAR)*

*and the CESM Isotope Development team: Z. Liu<sup>2</sup>, D. Noone<sup>4</sup>, K. Lindsay<sup>1</sup>, M. Vertenstein<sup>1</sup>, D. Bailey<sup>1</sup>, F. Joos<sup>3</sup>, A. Bozbiyik<sup>3</sup>, A. Gettelman<sup>1</sup>, S. Gu<sup>2</sup>, C. Koven<sup>5</sup>, E. Kluzek<sup>1</sup>, J. Nusbaumer<sup>4,7</sup>, B. Riley<sup>5</sup>, J. Tang<sup>5</sup>, P. Thornton<sup>6</sup>, X. Wen<sup>2</sup>, T. Wong<sup>4</sup>, J. Zhang<sup>2</sup>, J. Zhu<sup>2</sup>*

- 1. National Center for Atmospheric Research, 2. University of Wisconsin, Madison,*
- 3. University of Bern, 4. Oregon State University, 5. Lawrence Berkeley Laboratory,*
- 6. Oak Ridge National Laboratory, 7. University of Colorado, Boulder*

# What are isotopes?

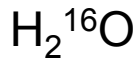
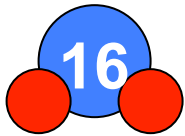
- Isotopes are variants of an element with additional neutrons → these lead to small differences in the physical properties of the element
- Radioactive or unstable isotopes break down by radioactive decay at a constant rate to produce a “daughter” element
- The half life of a radioactive isotope is the time it takes for 50% of the original nuclei in a sample to decay to the daughter element



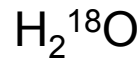
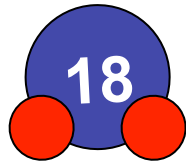
# Fractionation

- The small differences in the physical properties of each isotope affect how isotopes are partitioned between different parts of the climate system → this is called “fractionation”
- Due to fractionation, isotopes can be used to measure environmental changes
  - Oxygen and hydrogen Isotopes help us measure temperature and ice volume
  - Carbon isotopes help us infer changes in the ocean circulation and in the biological productivity, as well as use radiocarbon as “clock”

# Isotope Delta Notation

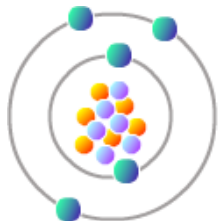


Light Oxygen



Heavy Oxygen  
(0.2% of all oxygen, 11% heavier than light oxygen)

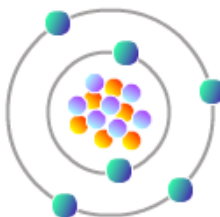
$$\delta^{18}\text{O} = \left( \frac{\left( \frac{^{18}\text{O}}{^{16}\text{O}} \right)_{\text{sample}}}{\left( \frac{^{18}\text{O}}{^{16}\text{O}} \right)_{\text{standard}}} - 1 \right) * 1000 \text{ ‰}$$



**Carbon**

- 6 Protons
- 6 Neutrons

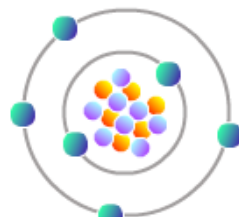
Nuclear number  
= 6 + 6  
= 12



**Carbon-13**

- 6 Protons
- 7 Neutrons

Nuclear number  
= 6 + 7  
= 13



**Carbon-14**

- 6 Protons
- 8 Neutrons

Nuclear number  
= 6 + 8  
= 14

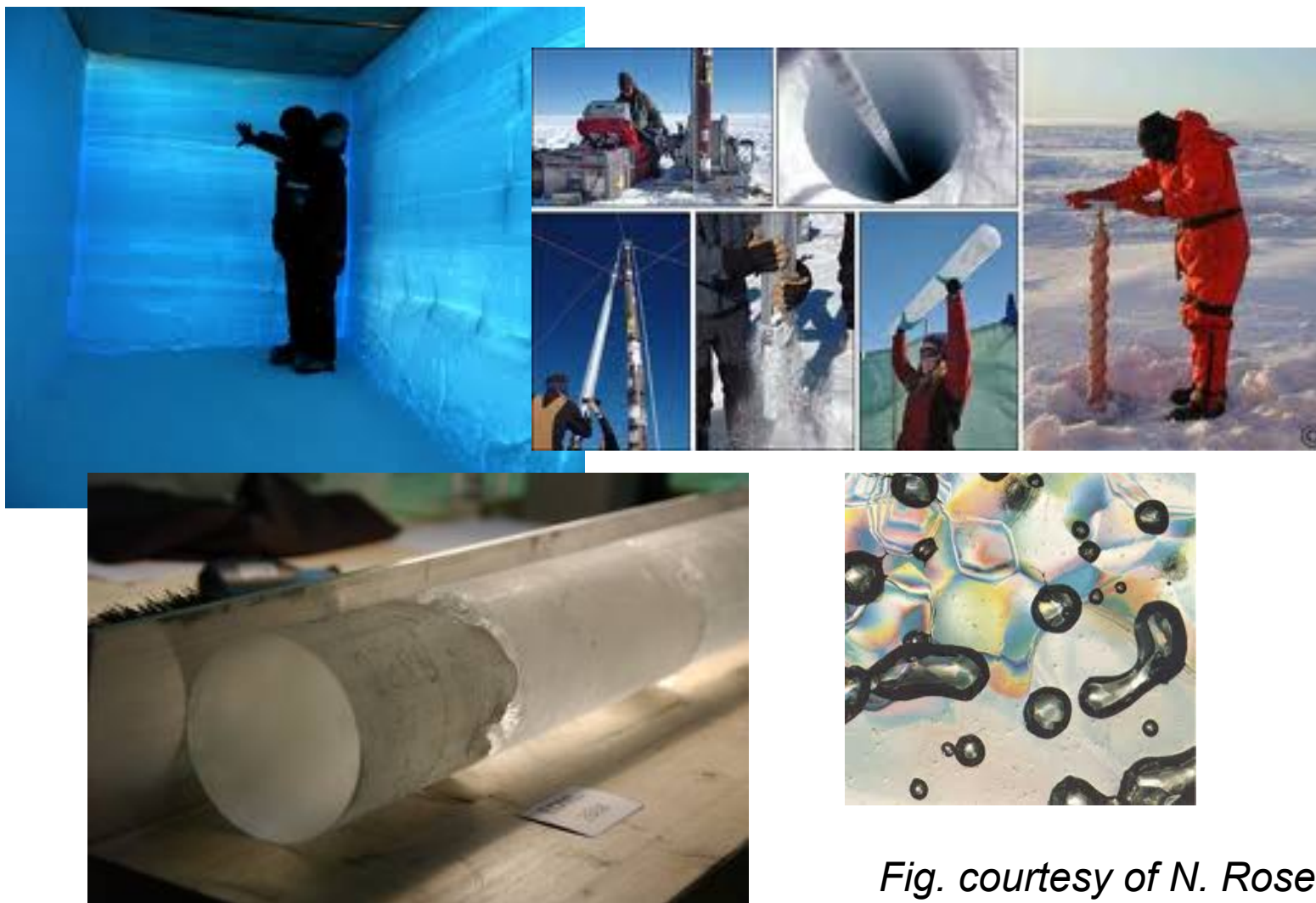
98.89%

1.11%

<0.01%

$$\delta^{13}\text{C} = \left( \frac{\left( \frac{^{13}\text{C}}{^{12}\text{C}} \right)_{\text{sample}}}{\left( \frac{^{13}\text{C}}{^{12}\text{C}} \right)_{\text{standard}}} - 1 \right) * 1000 \text{ ‰}$$

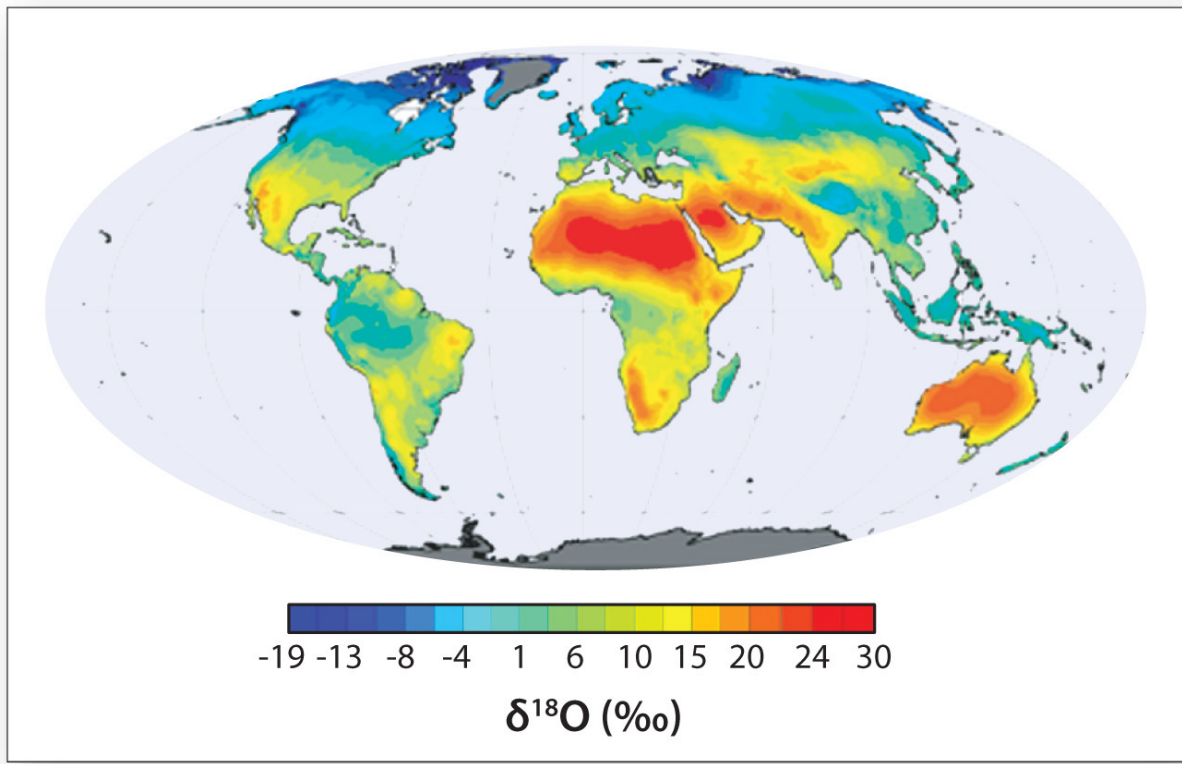
Many climate proxies are inferences of past climate change from isotopic records in archives like sediment and ice cores



*Fig. courtesy of N. Rosenbloom,  
Deep Time Liaison*

# Observed Oxygen isotope distribution of precipitation over land

- Lowest values of  $\delta^{18}\text{O}$  towards the poles, meaning more  $^{16}\text{O}$  reaches these latitudes than  $^{18}\text{O}$
- Ice sheets are located at high latitudes, so depleted in  $^{18}\text{O}$



# Temperature Effect for $\delta^{18}\text{O}$ in Precipitation

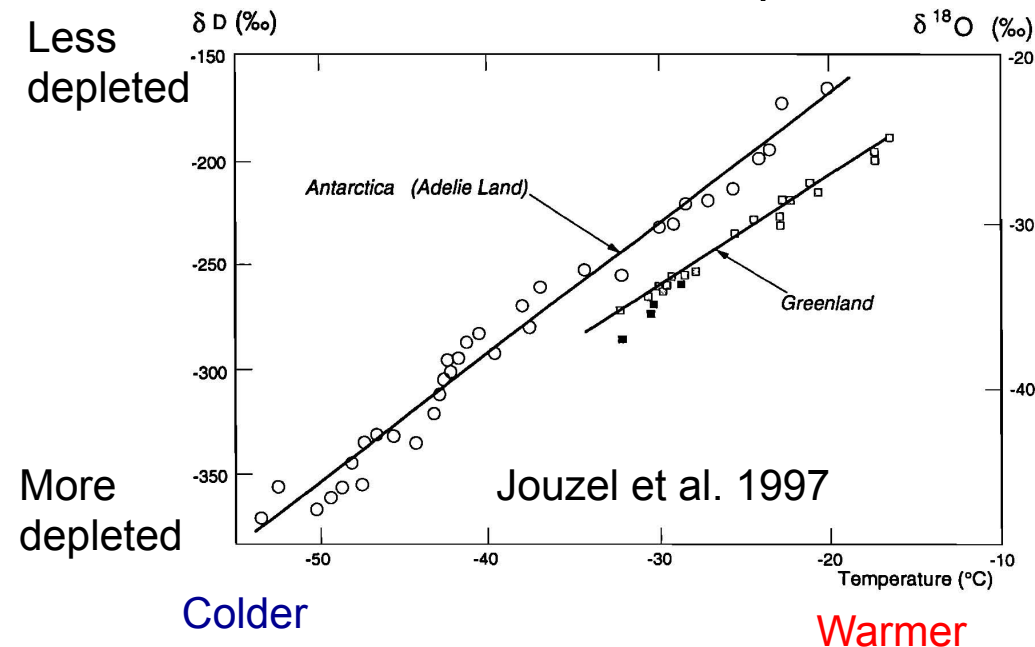
“Paleo-Thermometer”

$$\Delta(\delta^{18}\text{O}) \sim 0.7 [‰/K] \Delta T$$

Spatial relationship –  
Observed to hold temporally  
over seasons and  
interannually in ice cores.

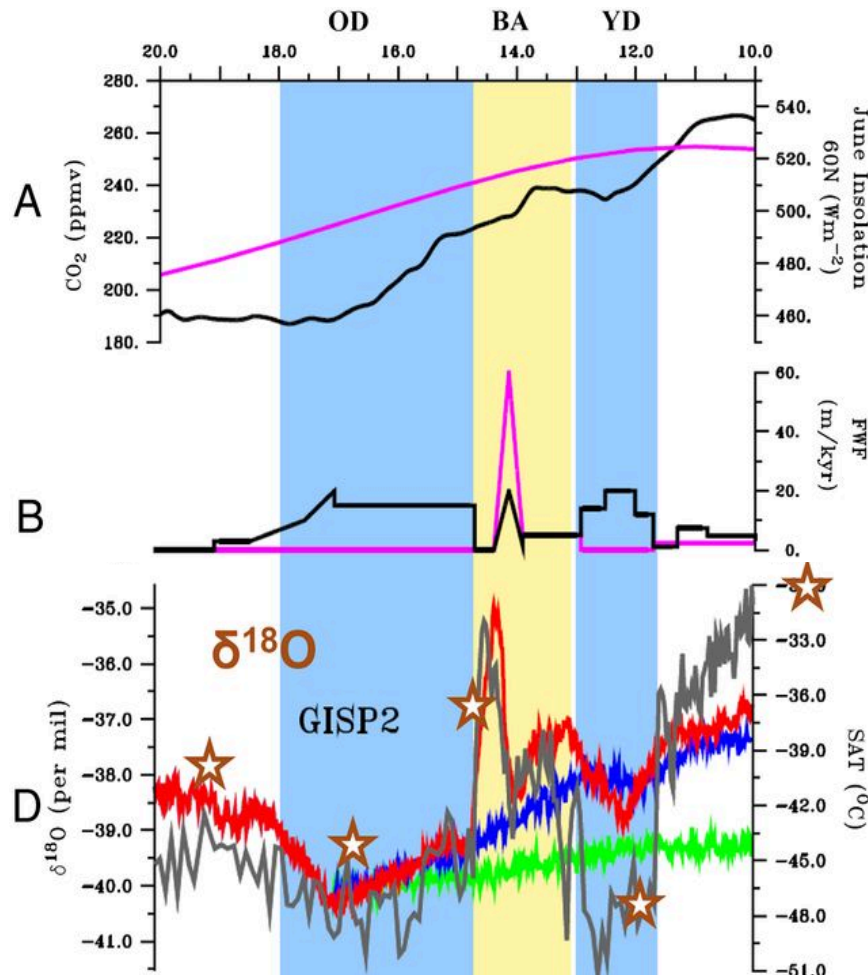
But can the modern  
relationship be used to  
reconstruct high-latitude  
temperature change over the  
long time periods of Glacial-  
Interglacial change?

$\delta^{18}\text{O}$  in Snow vs. Temp.





# CCSM3 TraCE, Transient Simulation of the Last Deglaciation



## Forcing:

60°N June Insolation (Orbital)  
 Atm CO<sub>2</sub> concentration  
 (+Ice Sheet orography)

## Meltwater input

NH  
 SH

## Proxy Comparison

★  $\delta^{18}O_p^*$  over Greenland  
 From IsoCAM3 'slices'

Greenland SAT CCSM3-Full  
 CO<sub>2</sub>+IS, Orbital+IS  
 $\delta^{18}O$  GISP2 Record

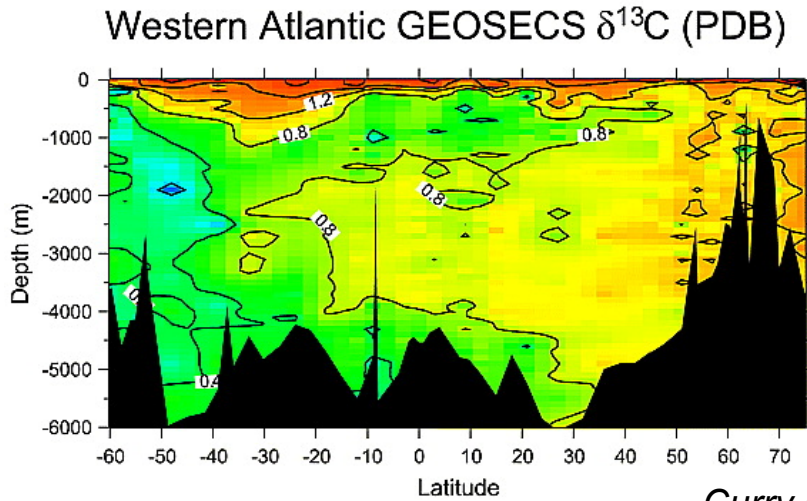
\* $\delta^{18}O_p$  ~ w/offset for bias

Liu Z et al. PNAS 2012;109:11101-11104

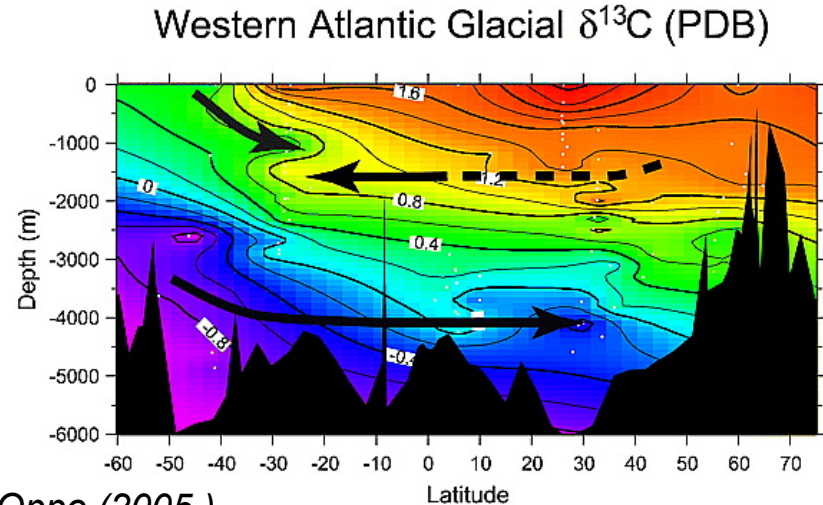
*CCSM3 suggests weaker YD cooling  
 But, Iso-CAM3 agrees better with  $\Delta\delta^{18}O_p$*



# Carbon isotopes as ocean tracers



*Curry and Oppo (2005)*



- The  $\delta^{13}\text{C}$  ratio of shells is a measure of how much photosynthesis is taking place in the oceans and/or how much organic material is removed from the surface to the deep ocean via circulation, as both processes enrich the water in  $^{13}\text{C}$
- paleo proxies mainly use  $\delta^{13}\text{C}$  as a water mass tracer, neglecting biological productivity changes
- $\delta^{13}\text{C}$  can be used as tracers of carbon cycle processes → e.g., used to diagnose the oceanic uptake of anthropogenic  $\text{CO}_2$
- $\Delta^{14}\text{C}$  is used as ocean reservoir age tracer

# Isotope-enabled CESM for studying abrupt change

*Bette Otto-Bliesner, Zhengyu Liu<sup>2</sup>, and iCESM Team<sup>\*\*\*</sup>*



*\*\*\*E. Brady, A. Jahn<sup>7</sup>, D. Noone<sup>4</sup>, K. Lindsay<sup>1</sup>, M. Vertenstein<sup>1</sup>, D. Bailey<sup>1</sup>, F. Joos<sup>3</sup>, A. Bozbiyik<sup>3</sup>, A. Gettelman<sup>1</sup>, S. Gu<sup>2</sup>, C. Koven<sup>5</sup>, E. Kluzek<sup>1</sup>, J. Nusbaumer<sup>4</sup>, B. Riley<sup>5</sup>, J. Tang<sup>5</sup>, P. Thornton<sup>6</sup>, X. Wen<sup>2</sup>, T. Wong<sup>4</sup>, J. Zhang<sup>2</sup>, J. Zhu<sup>2</sup>*

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# iCESM: Water Isotopic and Carbon isotope tracers development

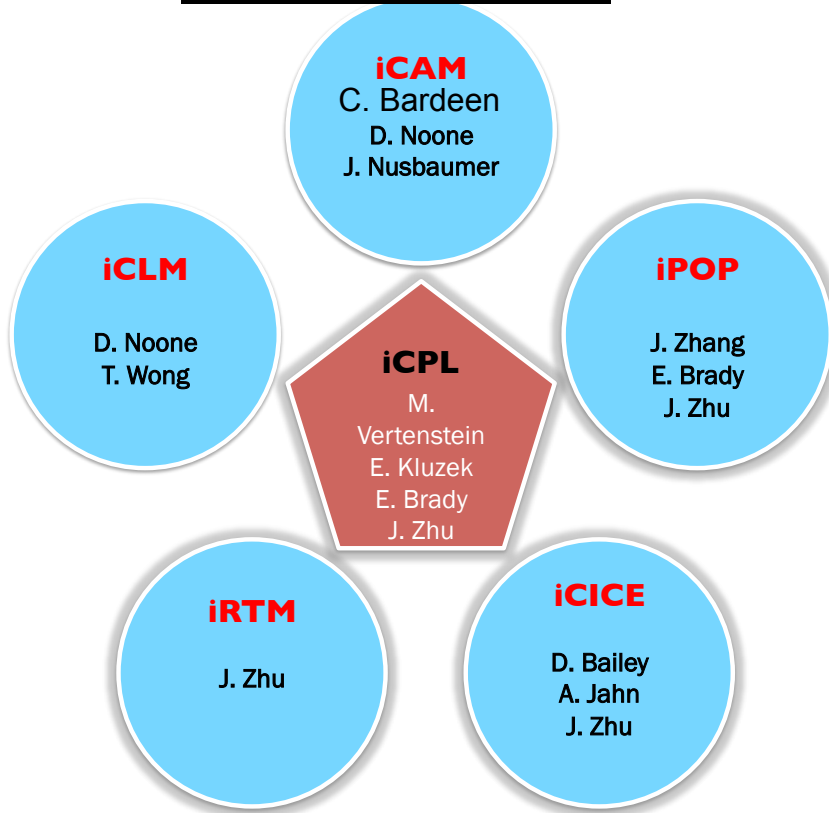


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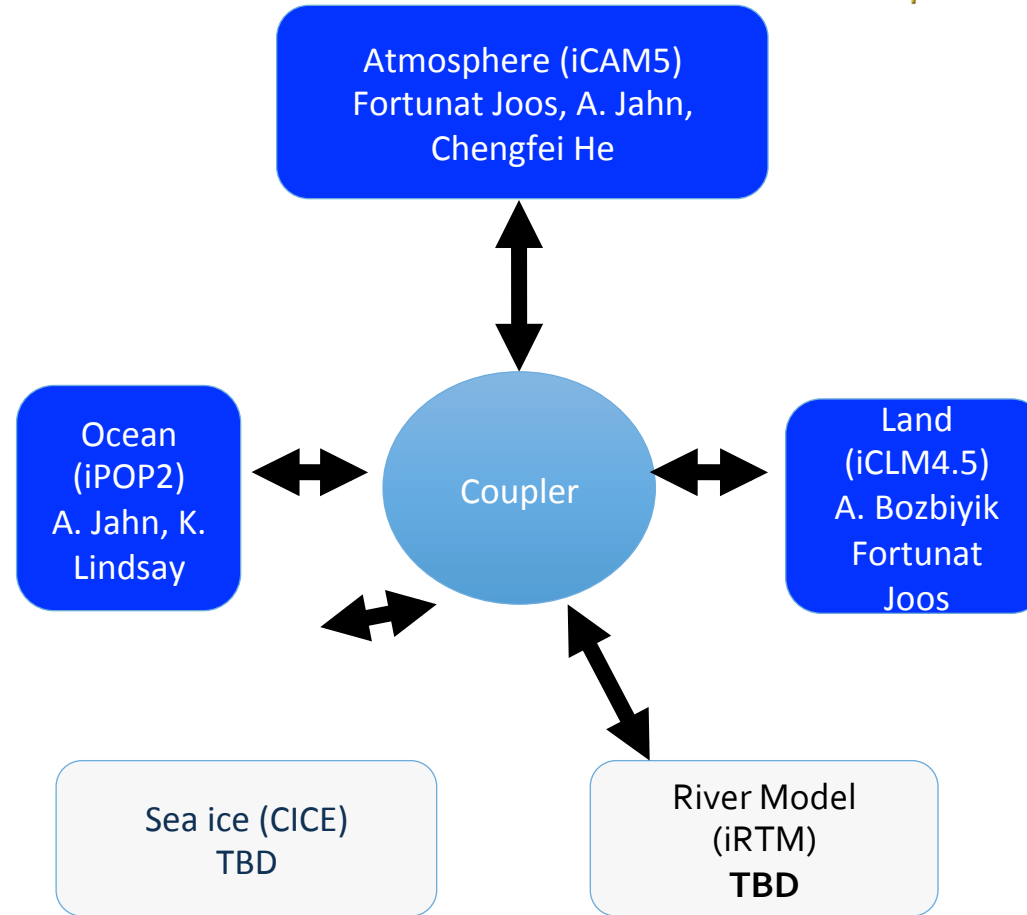
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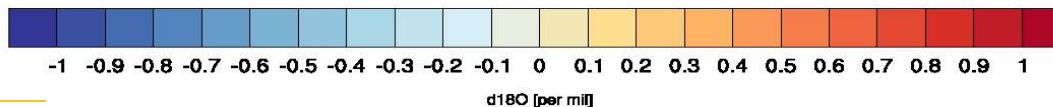
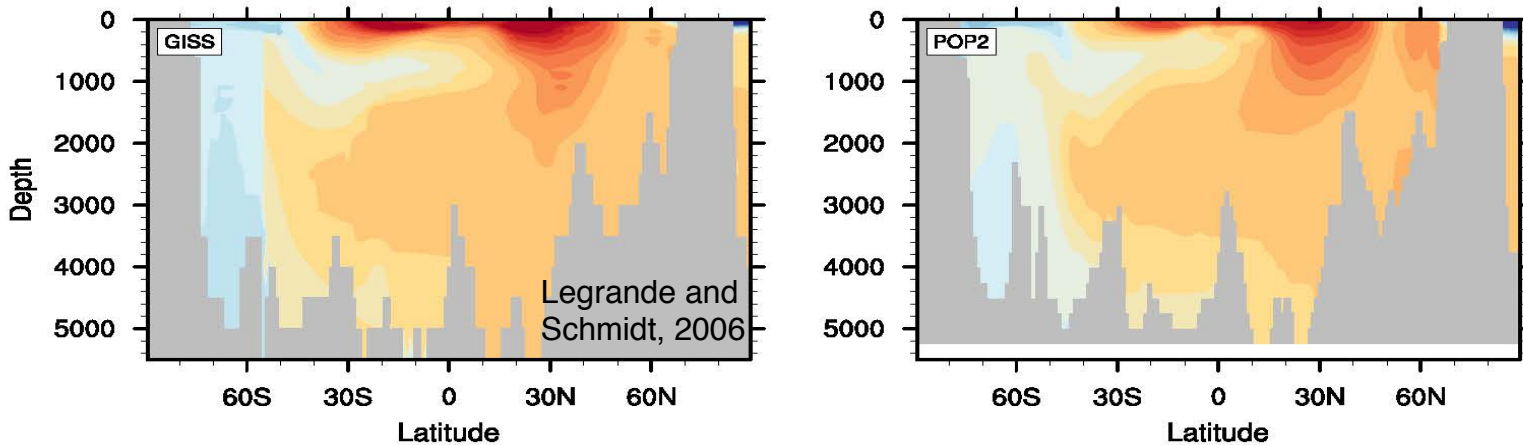
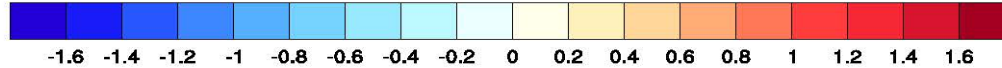
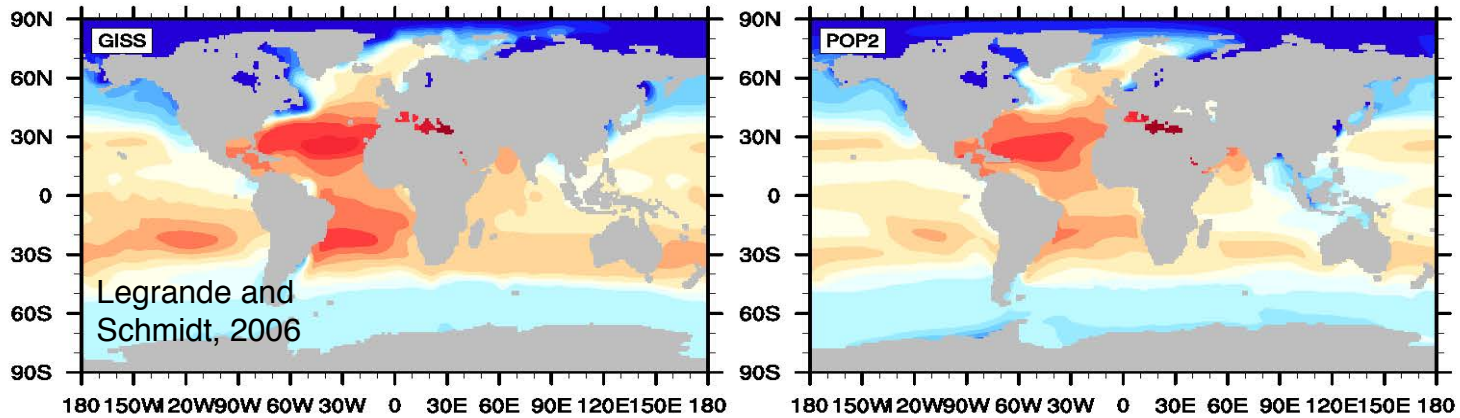
## Water Isotopes



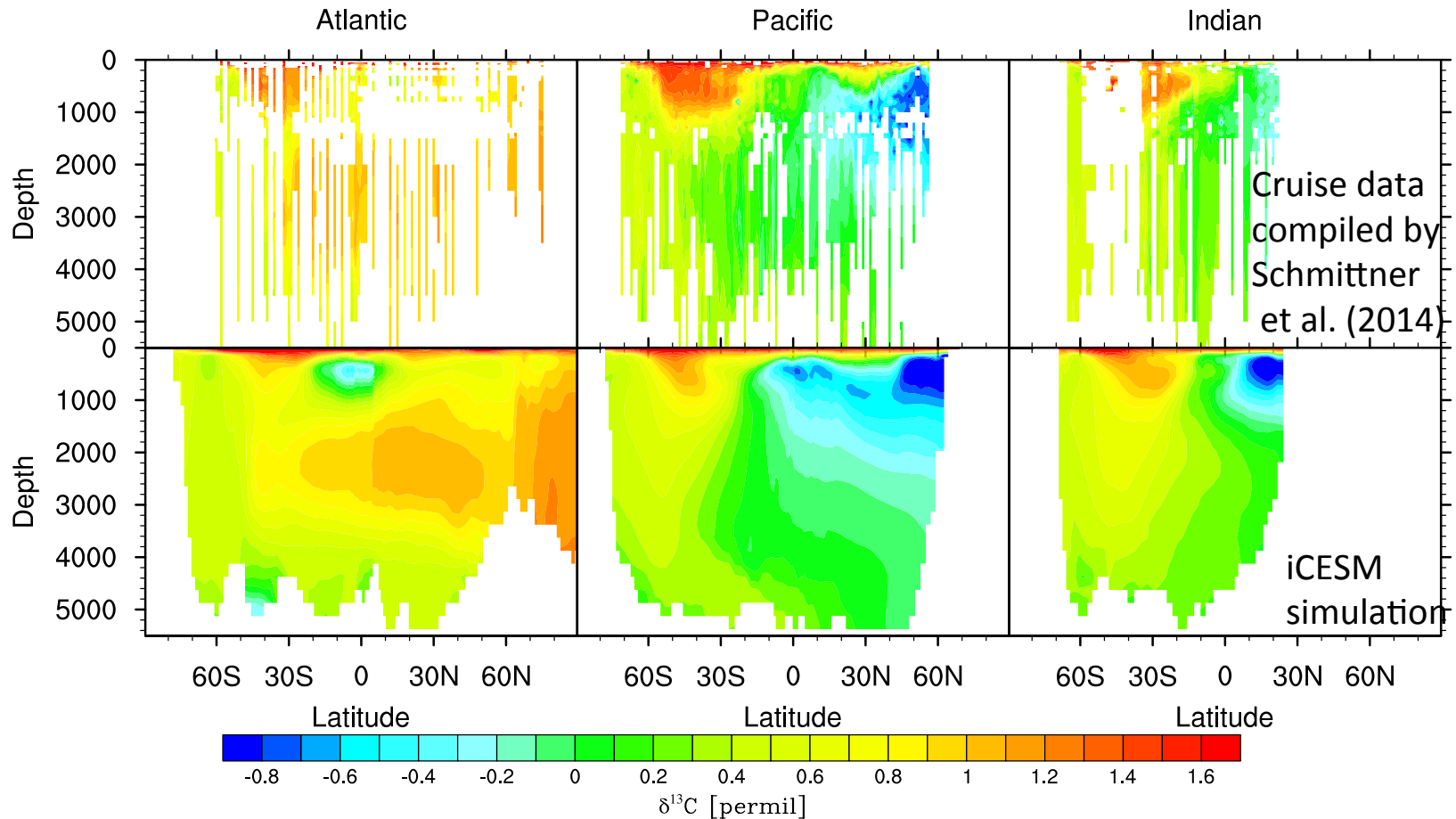
## Carbon Isotopes



# Coupled iCESM water isotopes

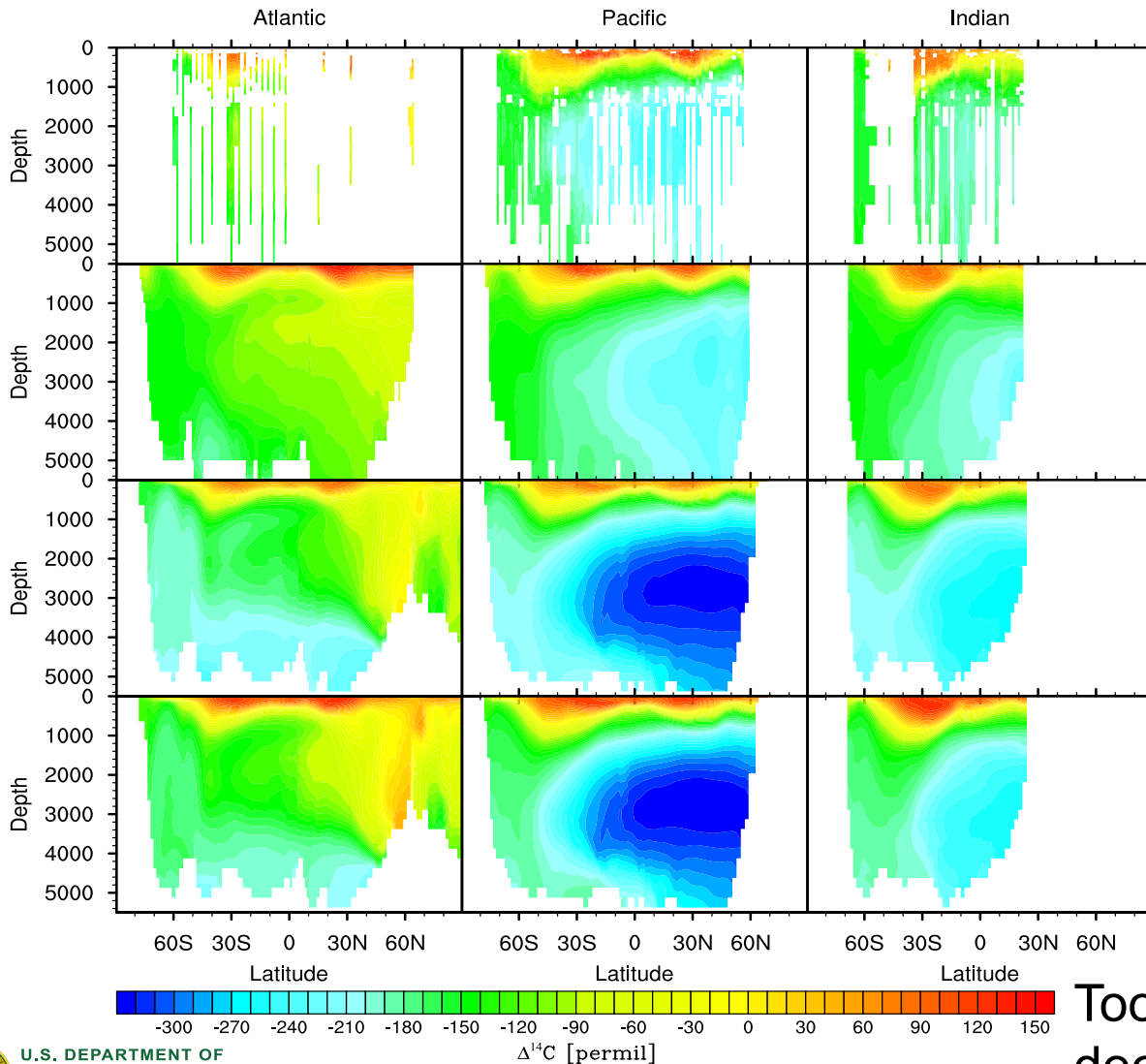


# Cross sections of oceanic $\delta^{13}\text{C}$ (1990s)





# Cross sections of Radiocarbon



Cruise data  
compiled by  
Schmittner  
et al. (2014)

GLODAP  
(Kay et al.  
2004)

Biotic Radiocarbon  
Simulation

Abiotic Radiocarbon  
Simulation

Too old radiocarbon ages in  
deep Pacific → shows model  
biases in that region



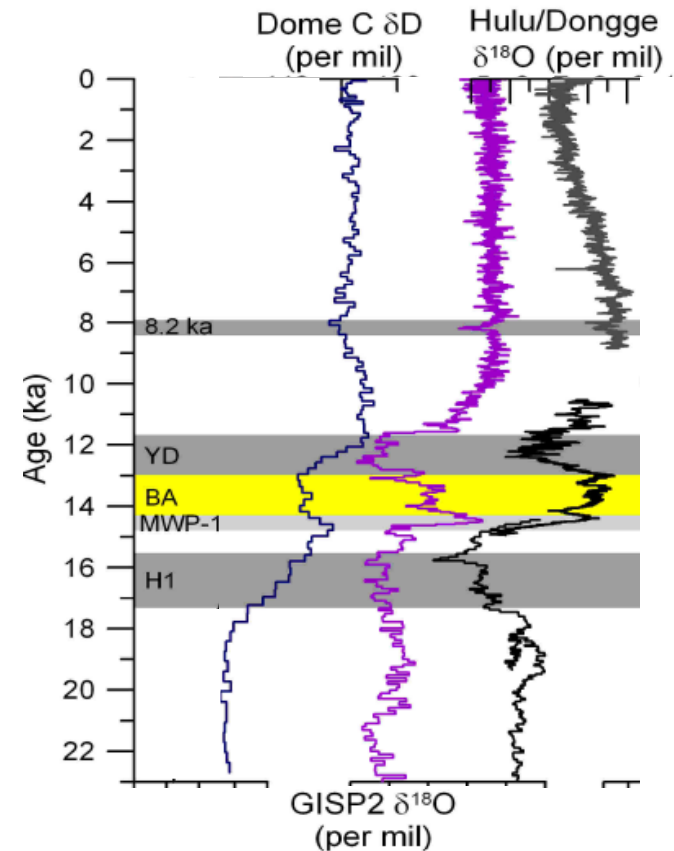
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Office of Science

Jahn et al. (2015), GMD

# Paleo Application: *iTraCE* simulations

- *iTraCE* will allow us to evaluate the skill of CESM, stability through time of the interpretations of the proxies, and the mechanisms associated with abrupt changes of the last 21,000 years
- Most of the isotope capabilities will be available to the community in CESM2 (to be released in Dec. 2016)



**Water isotope *iTraCE*:** Funded by NSF-P2C2: PIs: B. Otto-Bliesner (NCAR), Z. Liu (U. Wisc.), P. Clark (OSU)

**Carbon Isotope *O-iTraCE*:** Funded by NSF-P2C2: PIs: A. Jahn (CU) and Z. Liu (U. Wisc.)



# Thank You!

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