The Community Land Model

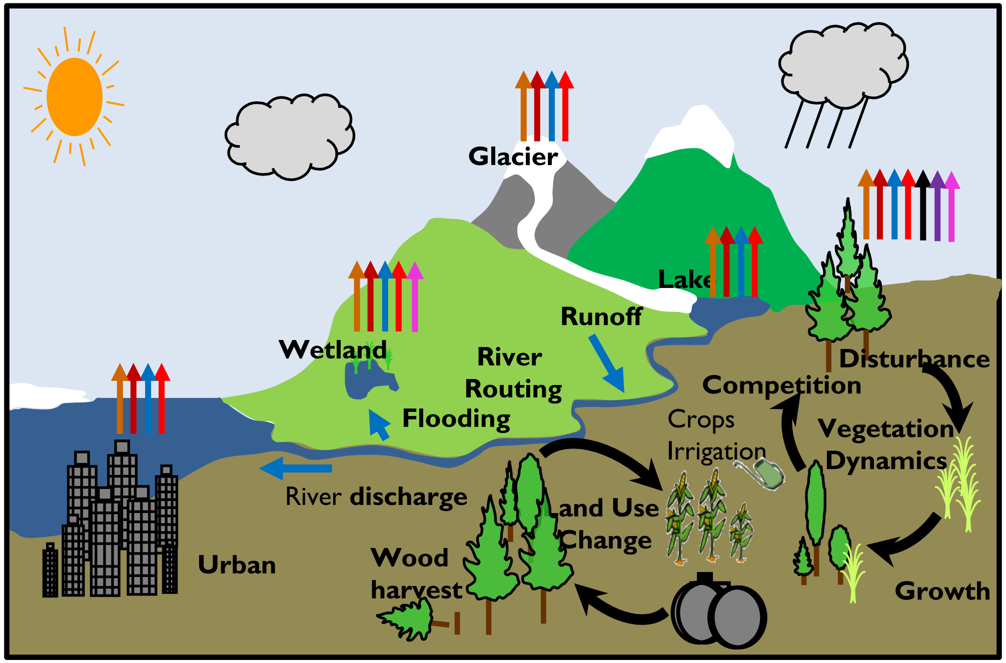
Representing terrestrial processes in the Earth System

David Lawrence

Co-chair of Land Model Working Group
Climate and Global Dynamics Lab
Terrestrial Sciences Section
dlawren@ucar.edu

NCAR is sponsored by the National Science Foundation

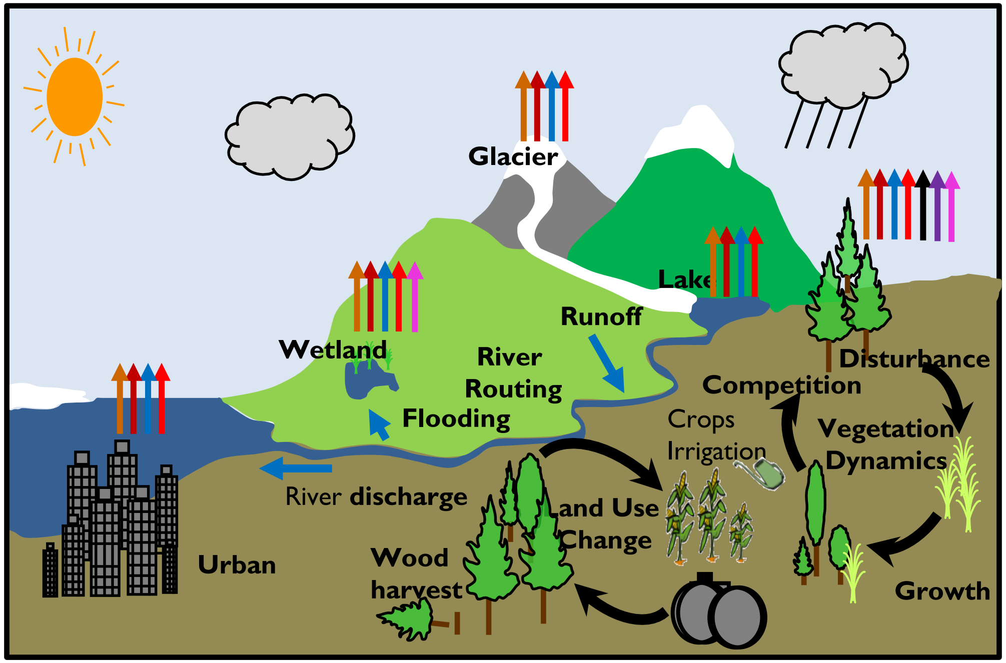
Land Modeling



“Why?”

“Are you sure this is necessary?”

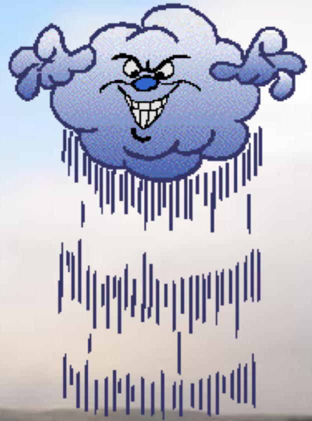
Land Modeling

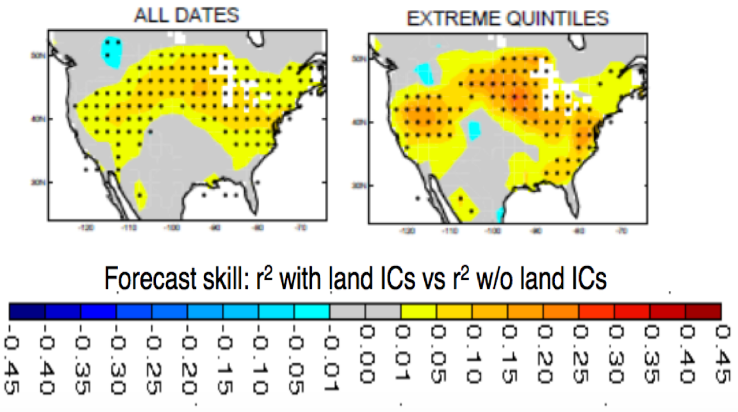


Yes!

Land is the critical interface through which humanity affects and is affected by, adapts to, and mitigates global environmental change

Land modelling, why?   Land-atmosphere interactions

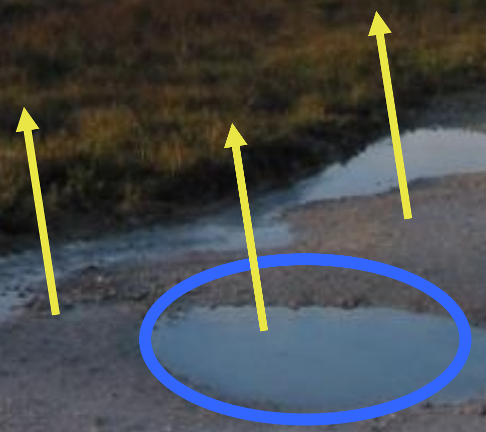




30-45 day forecast conditioned on SM

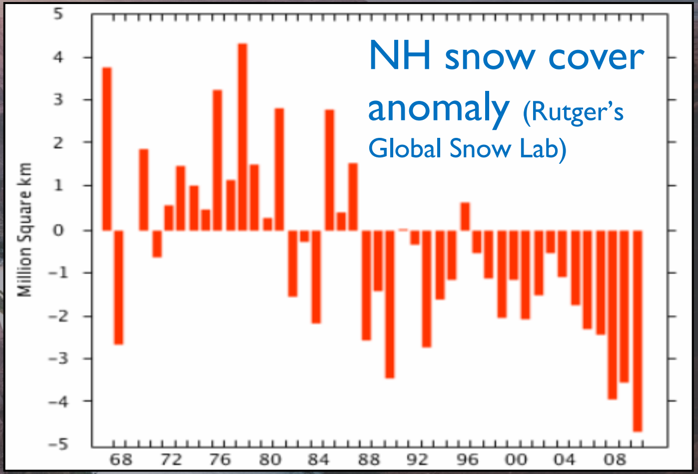
Koster et al., 2010

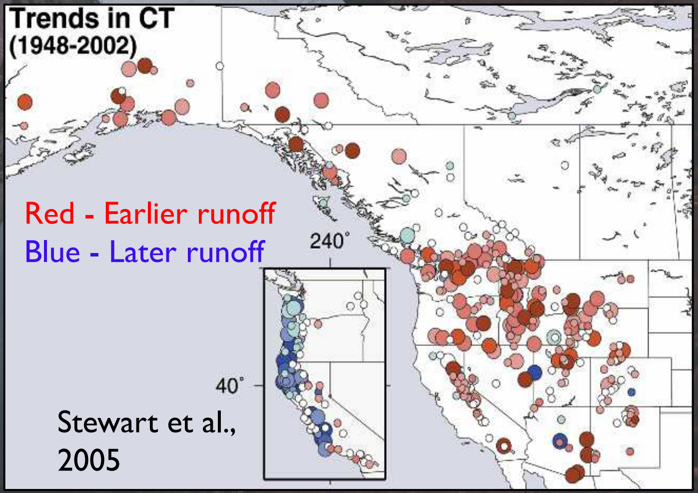
- When, where, and by how much do land fluxes influence atmosphere, surface temperature, clouds, precipitation, etc.?
- Land-driven predictability
  - Significant skill, especially when conditioned on amplitude of initial soil moisture anomaly
  - Increased land-atmosphere coupling in future warmer climate, increased land-driven skill?
- Land influence on extremes



Land modeling, why? Water

- Land feedbacks on droughts and floods
- Snow-albedo and snow-soil T feedbacks
- Water and food security
  - >1/6th world population dependent on water from seasonal snowpacks
- Water – plant interactions
  - Plant water use efficiency likely to increase with CO_2
- Streamflow prediction


NH snow cover anomaly (Rutger's Global Snow Lab)


Trends in CT (1948-2002)

Red - Earlier runoff

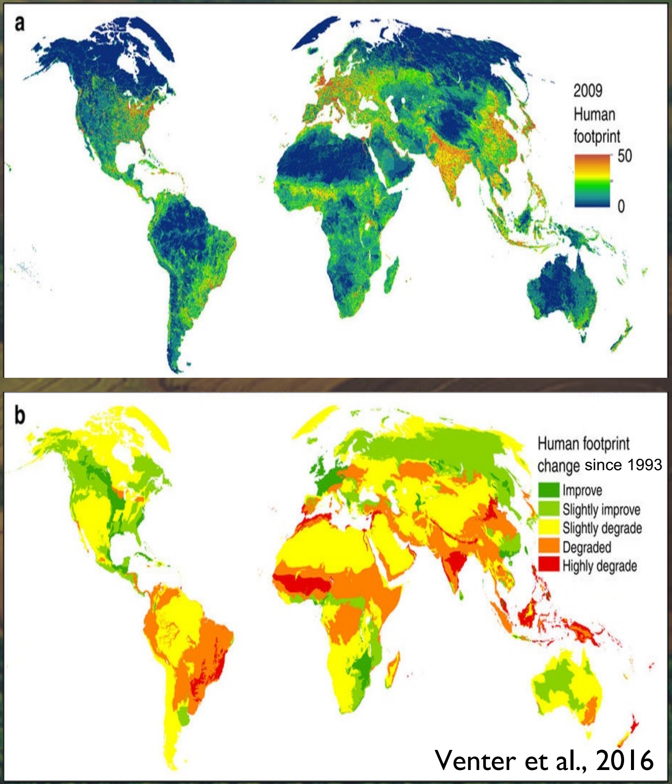
Blue - Later runoff

Stewart et al., 2005

Image: Kimon Maritz

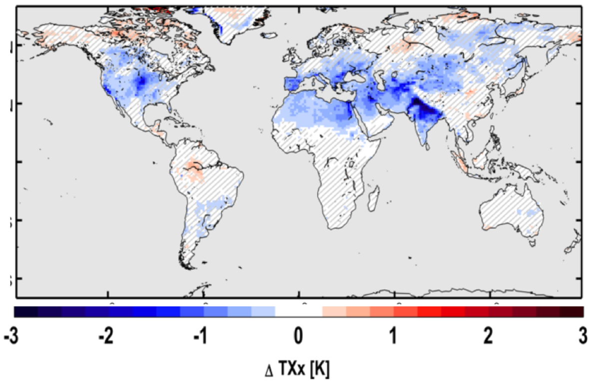
Land modeling why? Land-use and land-cover change

- ~25% non-ice land area undergone anthropogenic land-cover change
- ~80% non-ice land area under some form of land management
- Regionally, LULCC as impactful on surface climate as greenhouse gases
- ~1/3 of direct historic carbon emissions (180 ± 80PgC from land use, ~400 PgC from fossil fuel and cement),
- Deforestation: loss of Additional Sink Capacity yields indirect C impact
- Effectiveness of afforestation and biofuels for CO_2 mitigation
- Urban-rural differences in climate change impacts, e.g. ,heat stress



Venter et al., 2016

Irrigation mitigates heat extremes



Thierry, Lawrence, et al., 2017

Image: Frans Lanting/Robert Harding Picture Library

Land modeling, why? Carbon and ecology

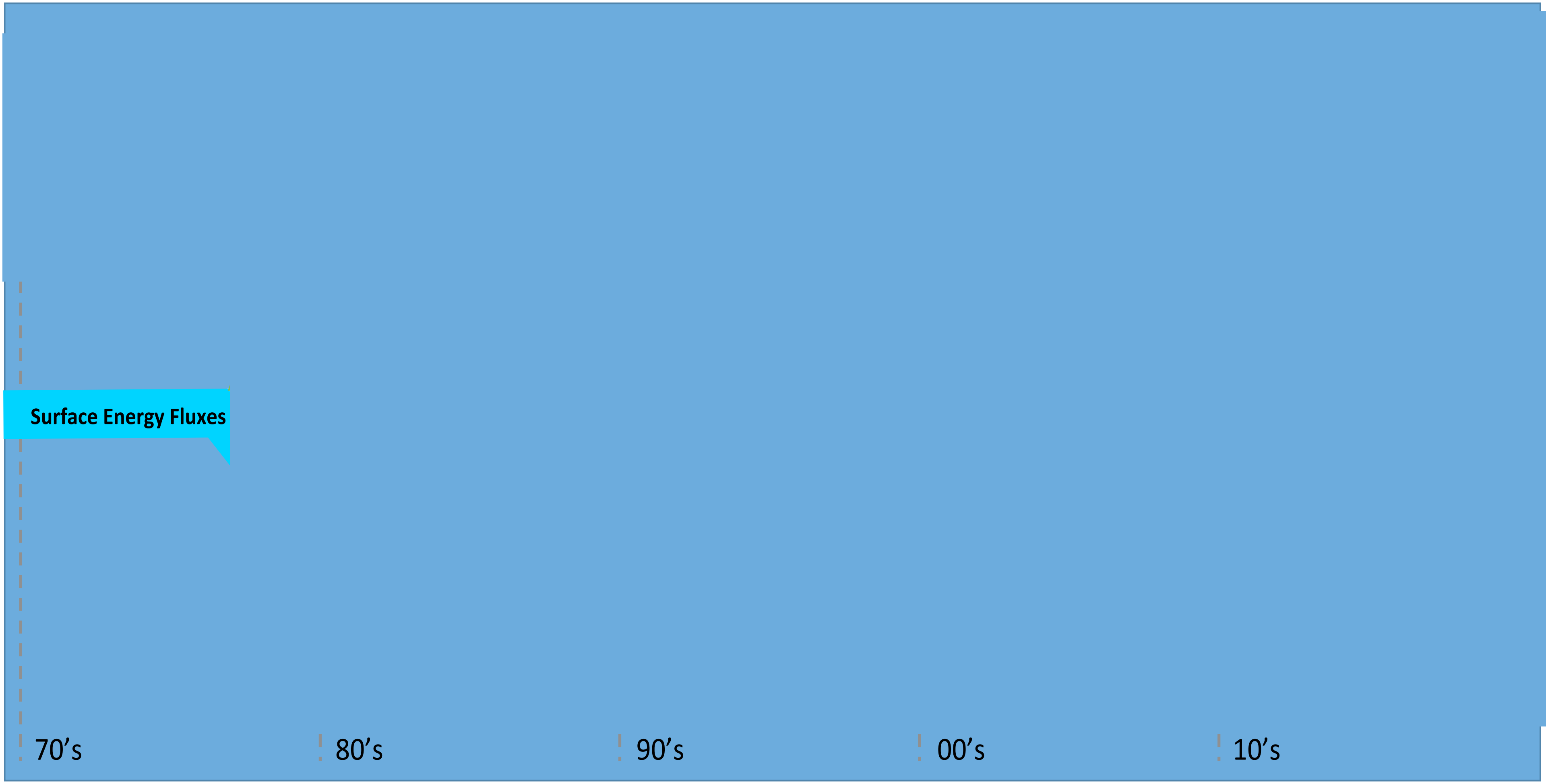
- Carbon and nitrogen cycle interactions and their impact on long term trajectory of terrestrial carbon sink
- High uncertainty in projected land C sink
  - Emissions driven RCP8.5:
    795 to 1140 ppm (source of ±1.2C uncertainty on top of 3.7C projected change)
- Vulnerability of ecosystems to climate change as well as natural and human disturbances
- Ecosystem services
- Ecosystem management to mitigate climate change

Image: Joel Vodell

The interdisciplinary evolution of land models

The interdisciplinary evolution of land models

Land as a lower boundary to the atmosphere → **Land as an integral component of the Earth System**



Surface Energy Fluxes

70's

80's

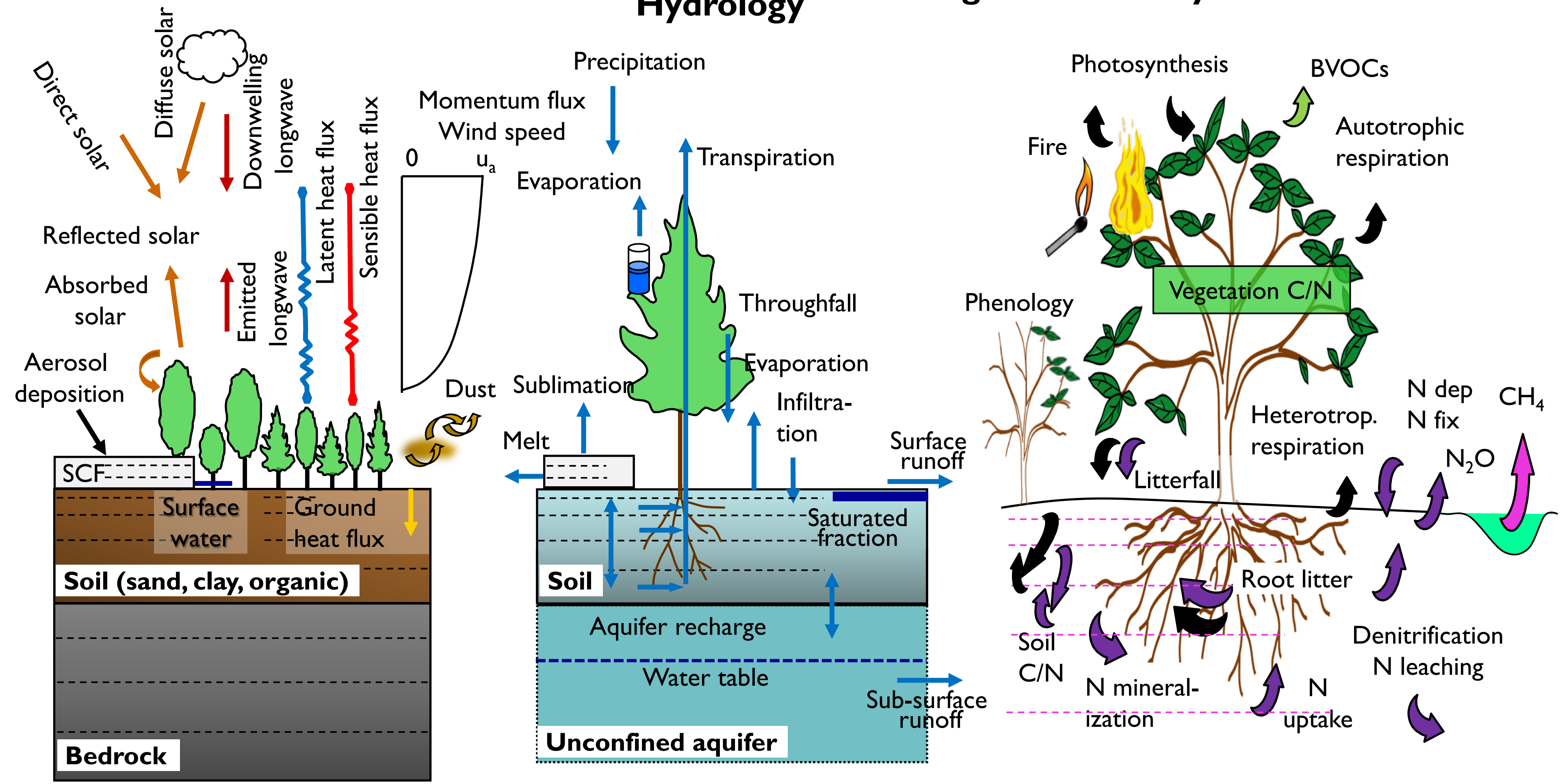
90's

00's

10's

Figure: Fisher, Lawrence, Bonan, Clark, unpublished

Community Land Model Key Processes



Surface energy fluxes

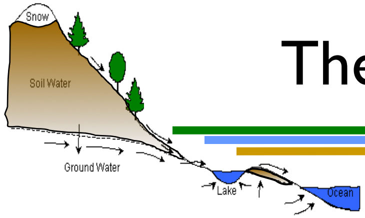
Direct solar
Diffuse solar
Downwelling longwave
Reflected solar
Absorbed solar
Emitted longwave
Latent heat flux
Sensible heat flux
Aerosol deposition
SCF
Surface water
Ground heat flux
Soil (sand, clay, organic)
Bedrock

Hydrology

Momentum flux
Wind speed
0
u_a
Precipitation
Evaporation
Transpiration
Throughfall
Evaporation
Dust
Sublimation
Melt
Infiltration
Surface runoff
Saturated fraction
Soil
Aquifer recharge
Water table
Sub-surface runoff
Unconfined aquifer

Biogeochemical cycles

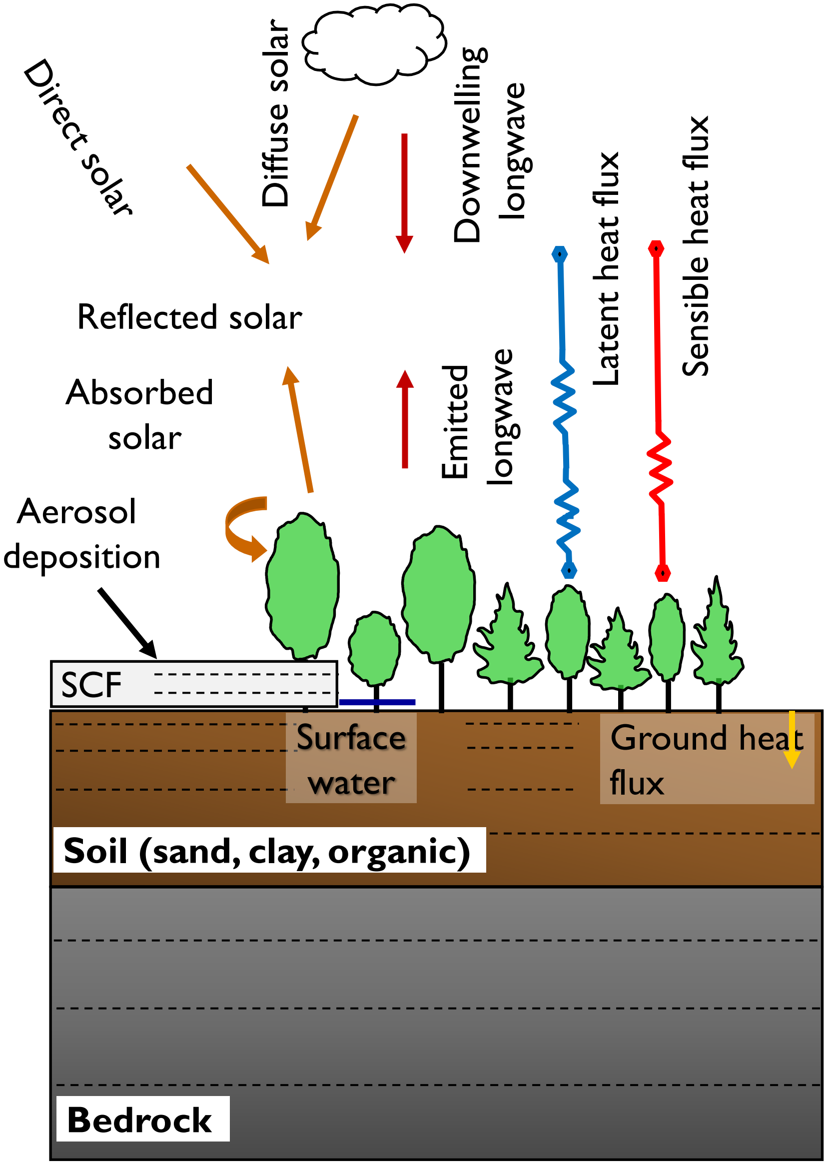
Photosynthesis
BVOCs
Fire
Autotrophic respiration
Vegetation C/N
Phenology
Heterotrop. respiration
N dep
N fix
CH_4
N_2O
Litterfall
Root litter
Soil C/N
N mineralization
N uptake
Denitrification
N leaching



The role of a land model within an Earth System Model

- exchanges of momentum, energy, water vapor, CO_2, dust, and other trace gases/materials between land surface and the overlying atmosphere (and routing of runoff to the ocean)
- states of land surface (e.g., soil moisture, soil temperature, canopy temperature, snow water equivalent, C and N stocks in vegetation and soil)
- characteristics of land surface (e.g., soil texture, surface roughness, albedo, emissivity, vegetation type, cover extent, leaf area index, and seasonality)

At each time step the land model solves Surface Energy Balance



Direct solar

Diffuse solar

Downwelling longwave

Latent heat flux

Sensible heat flux

Reflected solar

Absorbed solar

Emitted longwave

Aerosol deposition

SCF

Surface water

Ground heat flux

Soil (sand, clay, organic)

Bedrock

$$S^{\uparrow} - S^{\downarrow} + L^{\uparrow} - L^{\downarrow} = \lambda E + H + G$$

$S^{\uparrow}$, $S^{\downarrow}$ are down(up)welling solar radiation,

$L^{\uparrow}$, $L^{\downarrow}$ are up(down)welling longwave rad,

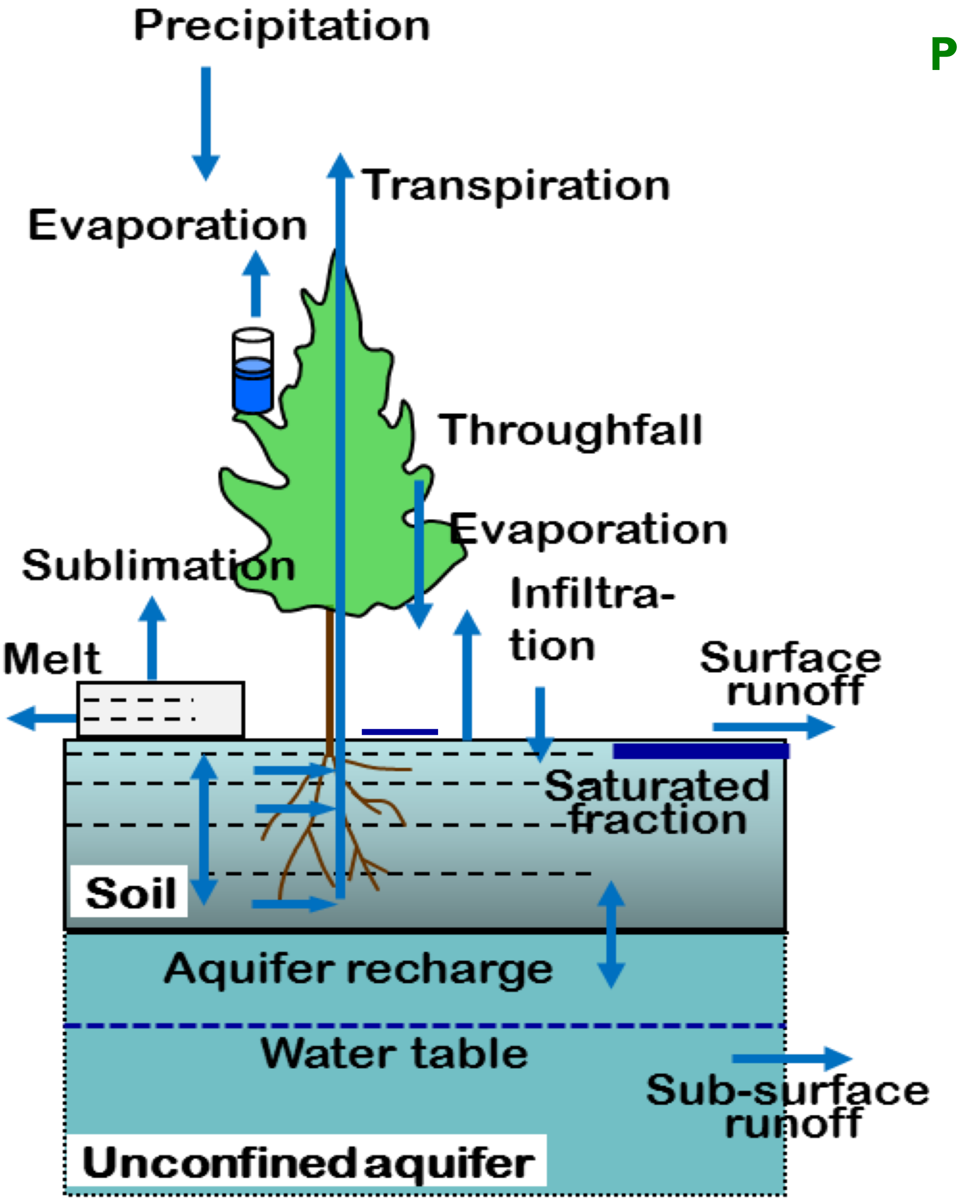
λ is latent heat of vaporization,

E is evaporation,

H is sensible heat flux

G is ground heat flux

… and the Surface Water Balance



$$P = E_S + E_T + E_C + R + (\Delta W_{soi} + \Delta W_{snw} + \Delta W_{sfcw} + \Delta W_{can}) / \Delta t$$

P is rainfall/snowfall,

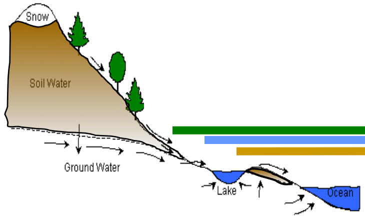
E_S is soil evaporation,

E_T is transpiration,

E_C is canopy evaporation,

R is runoff (surf + sub-surface),

$\Delta W_{soi} / \Delta t$, $\Delta W_{snw} / \Delta t$, $\Delta W_{sfcw} / \Delta t$, $\Delta W_{can} / \Delta t$, are the changes in soil moisture, surface water, snow, and canopy water over a timestep

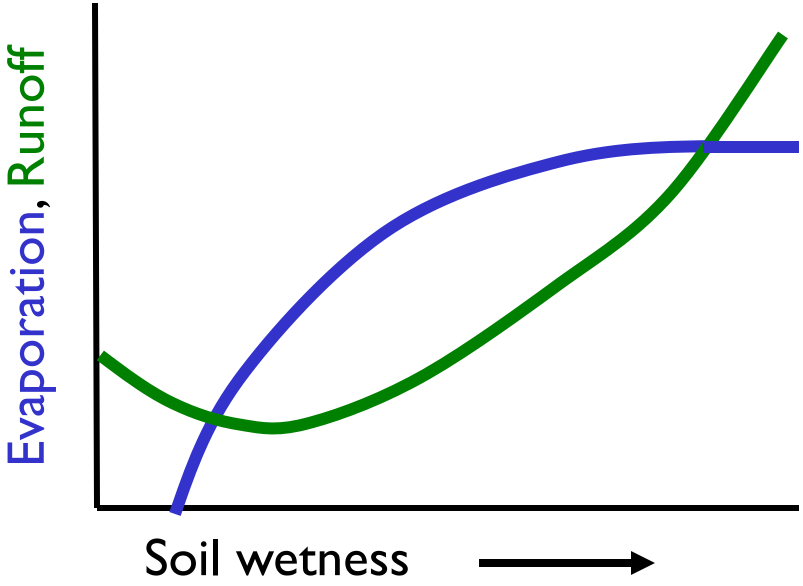


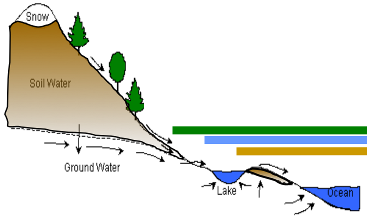
Terrestrial water and energy cycles intricately linked

> *"The ability of a land-surface scheme to model evaporation correctly depends crucially on its ability to model runoff correctly. The two fluxes are intricately related through soil moisture."*

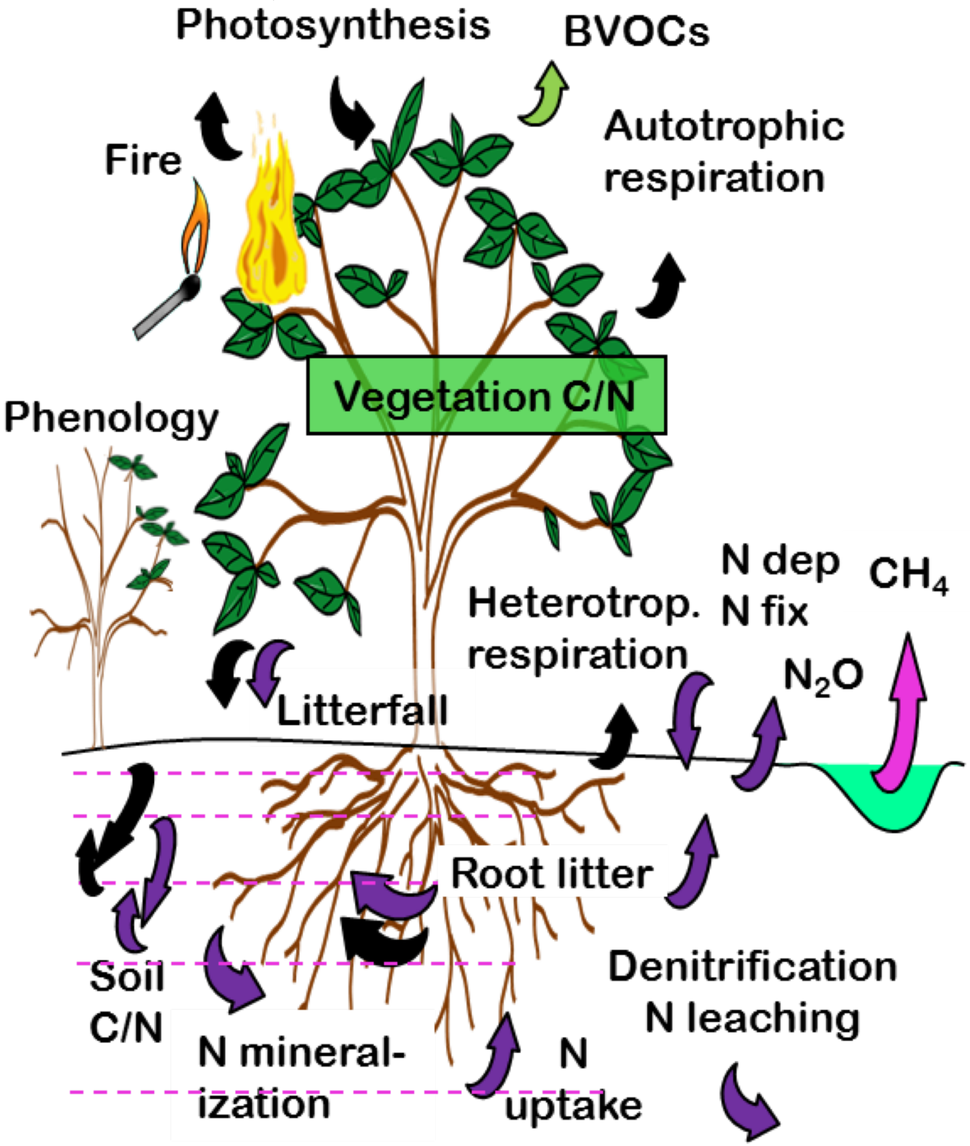
(Koster and Milly, 1997).

Runoff and evaporation both vary non-linearly with soil moisture





… and Surface Carbon Exchange



NEE = GPP – HR – AR – Fire – LUC

NEE is net ecosystem exchange

GPP is gross primary productivity

HR is heterotrophic respiration

AR is autotrophic respiration

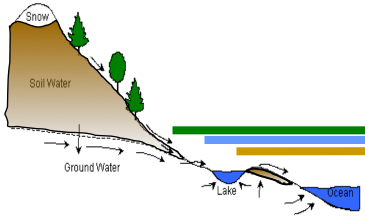
Fire is carbon flux due to fire

LUC is C flux due to land use change

Land complexity: Submodels of CLM

- Biogeophysics
  - Photosynthesis and stomatal resistance
  - Hydrology
  - Snow
  - Soil thermodynamics
  - Surface albedo and radiative fluxes
- Biogeochemistry
  - Carbon / nitrogen pools, allocation, respiration
  - Vegetation phenology
  - Decomposition
  - Plant mortality
  - External nitrogen cycle
  - Methane production and emission

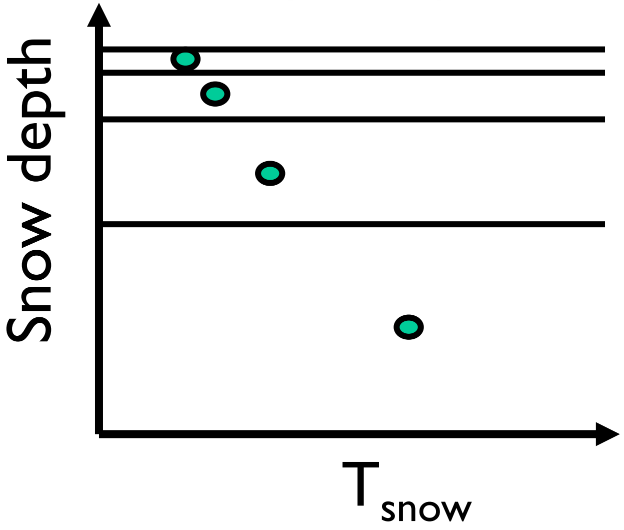
- Vegetation dynamics
- Urban
- Crop and irrigation
- Lakes
- Glaciers and ice sheets
- Fire and fire emissions
- Dust emissions
- River flow
- Biogenic Volatile Organic Compound emissions



Land model complexity: Snow model example

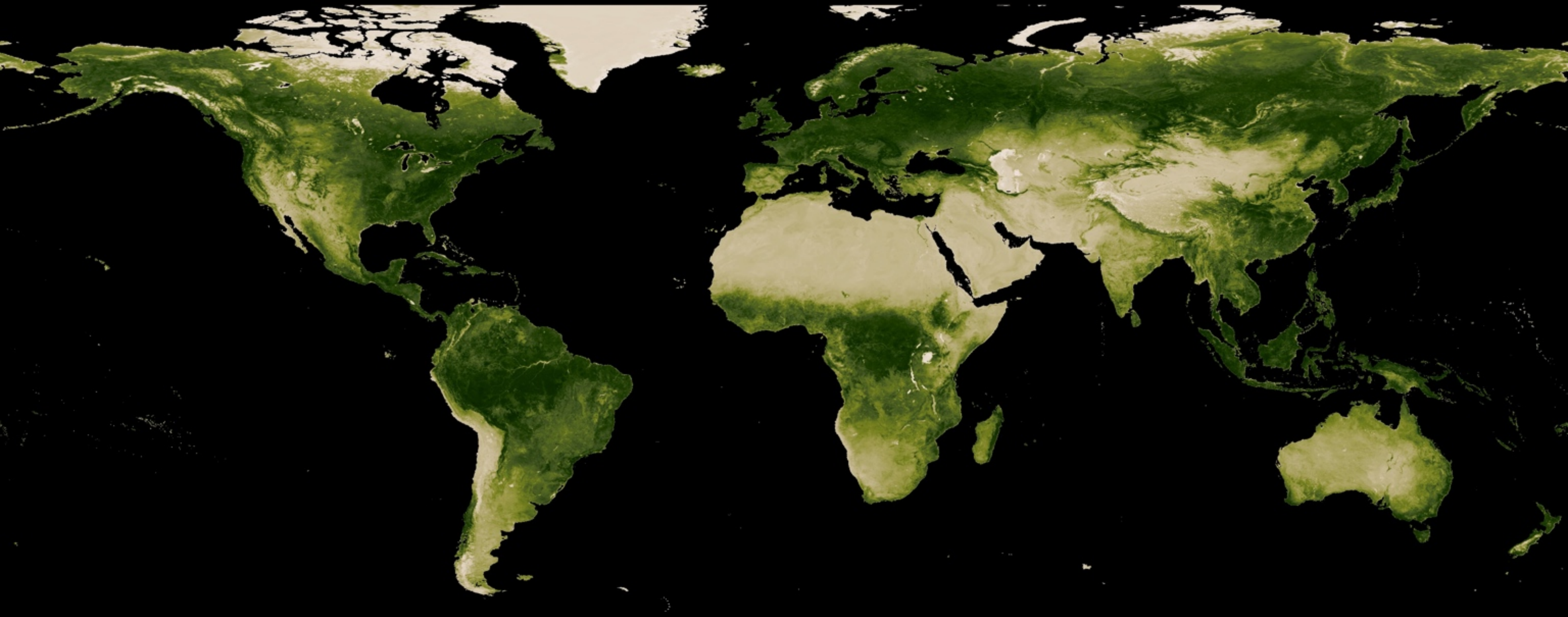
State Variables

$$N, w_{liq,i}, w_{ice,i}, \Delta z_i, T_i$$

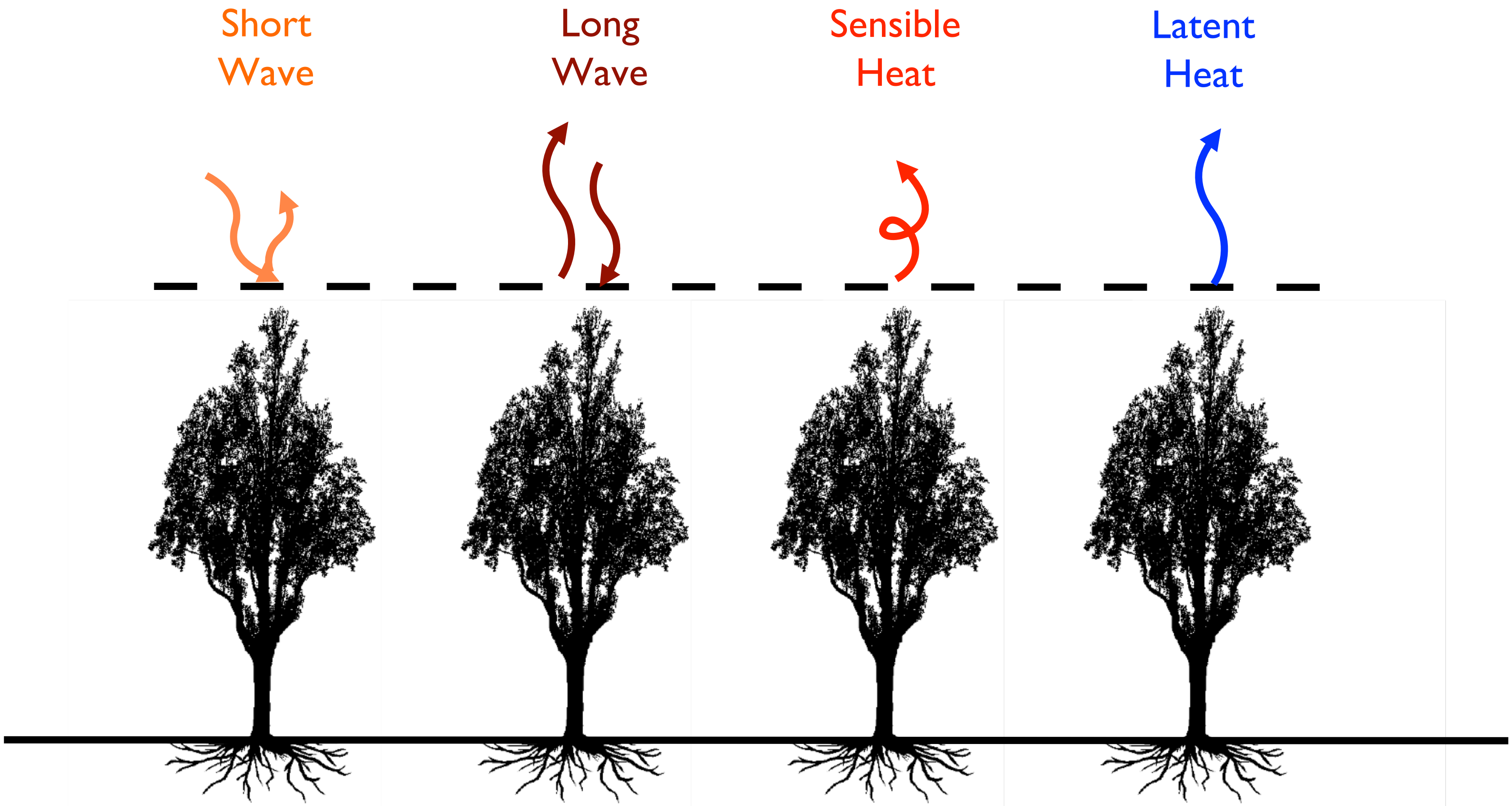


- Up to 10-layers of varying thickness
- Represented processes
  - Accumulation and fresh snow density *f* (T, wind)
  - Snow melt and refreezing
  - Snow aging
  - Water and energy transfer across snow layers
  - Snow compaction
    - destructive metamorphism due to temperature and wind
    - overburden
    - melt-freeze cycles
  - Sublimation
  - Aerosol (black carbon, dust) deposition
  - Canopy snow storage and unloading
  - Canopy snow radiation
  - Snow burial of vegetation
  - Snow cover fraction
- Missing processes
  - Blowing snow
  - Subgrid variations in snow depths
  - Depth hoar

Plants ⟷ Climate

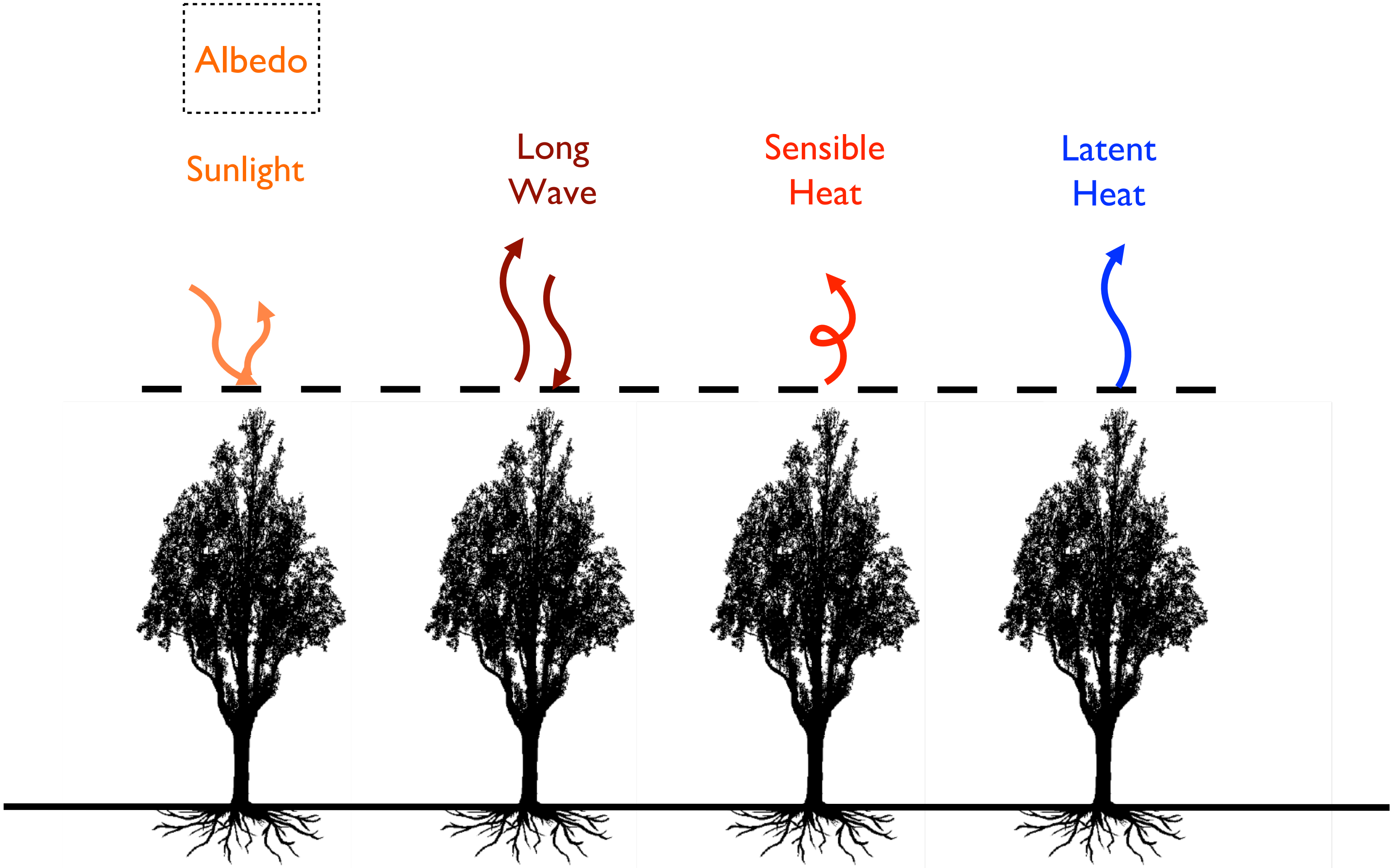


How do Plants and Climate Interact?



Terrestrial Surface Energy Budget

How do Plants and Climate Interact?

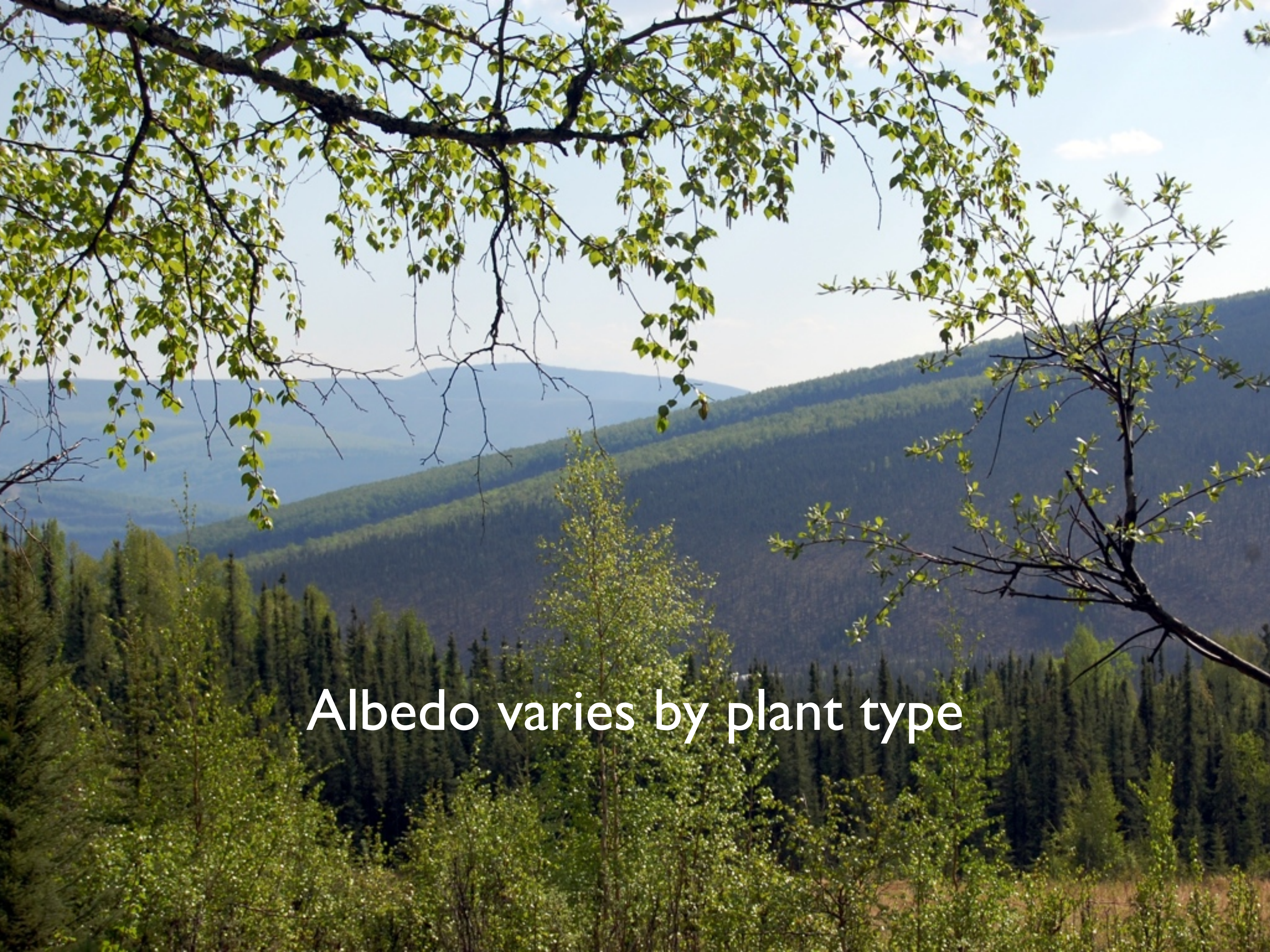

Albedo

Sunlight

Long Wave

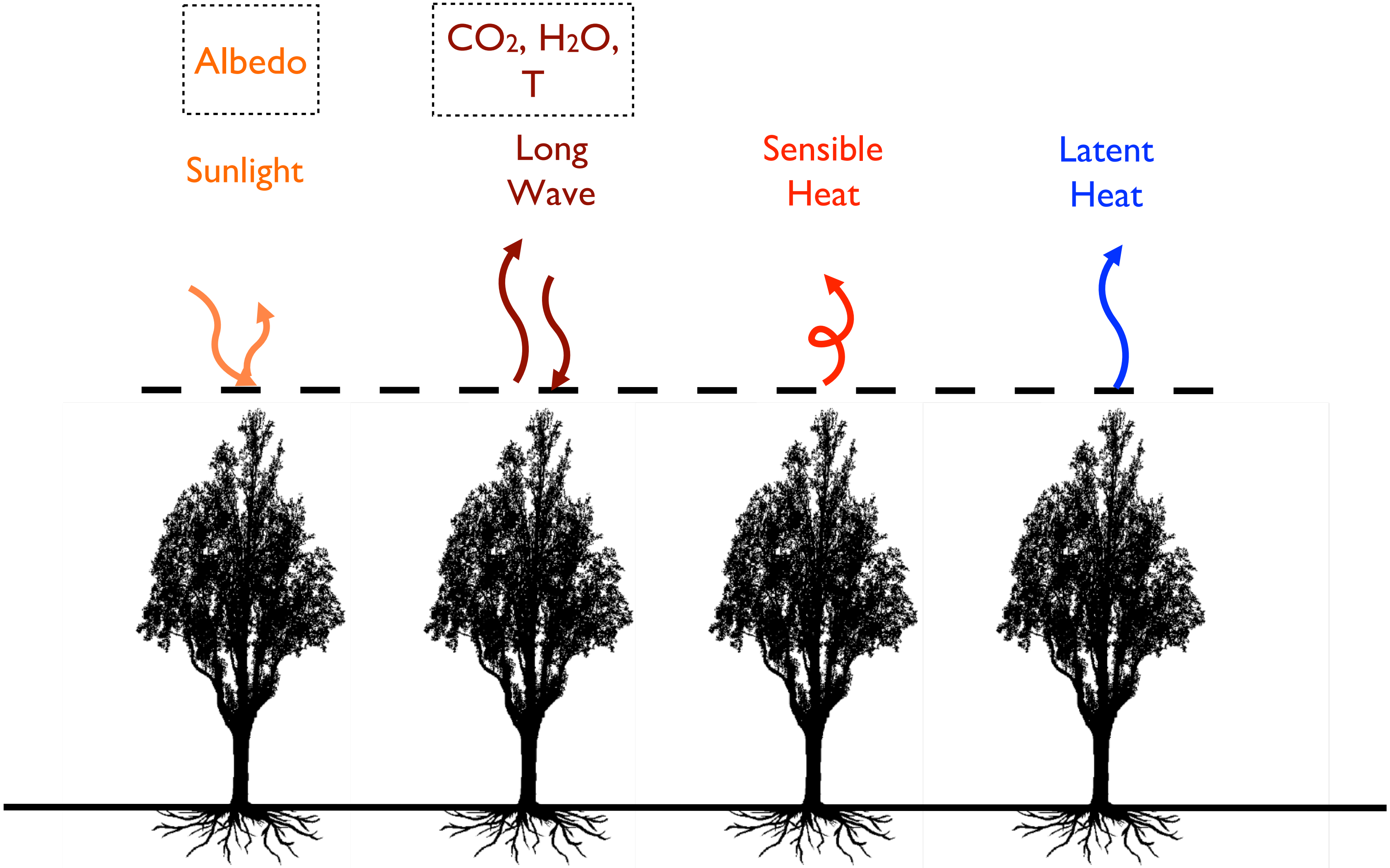
Sensible Heat

Latent Heat



Albedo varies by plant type

How do Plants and Climate Interact?



Albedo

CO_2, H_2O, T

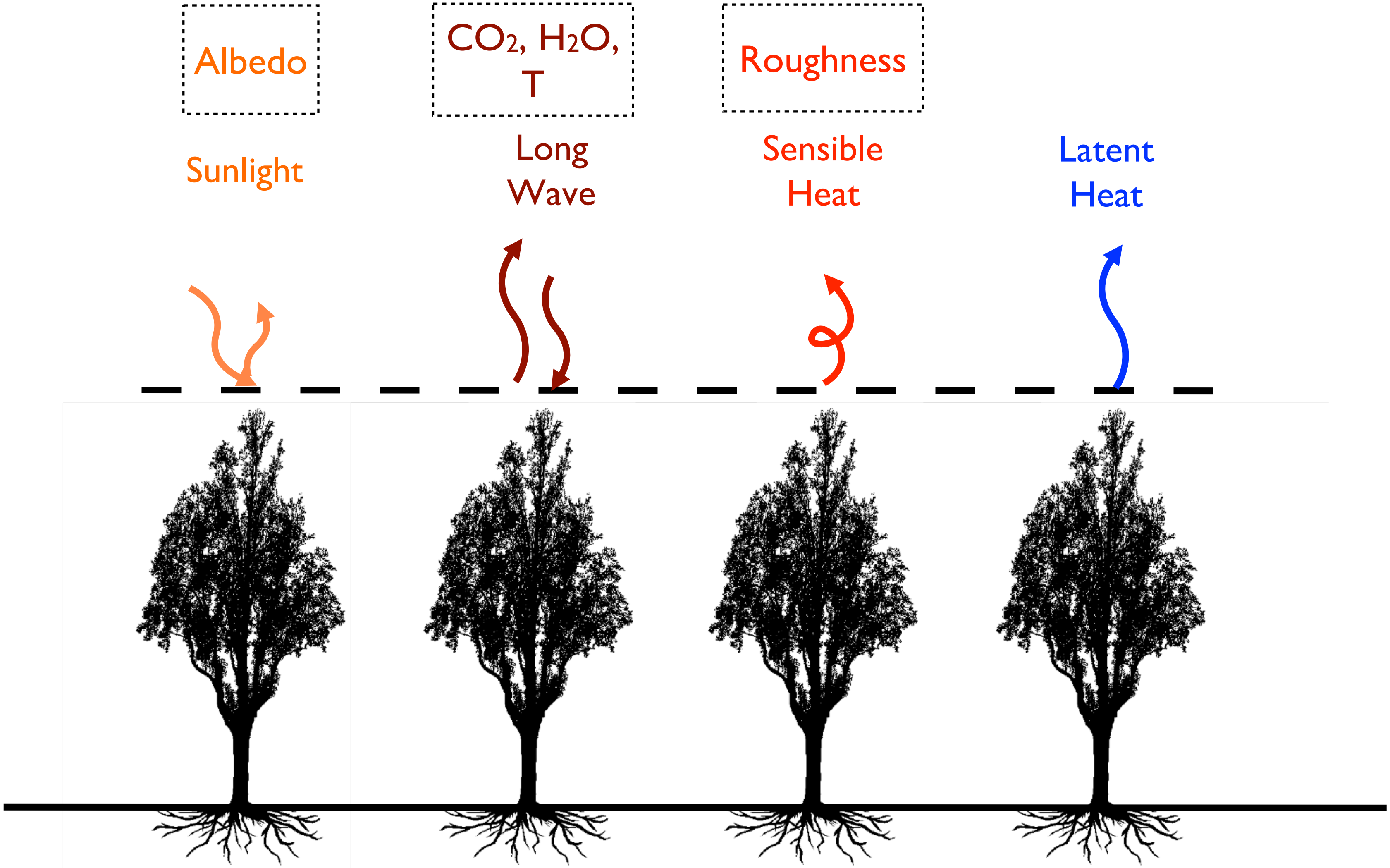
Sunlight

Long Wave

Sensible Heat

Latent Heat

How do Plants and Climate Interact?



Albedo

CO_2, H_2O, T

Roughness

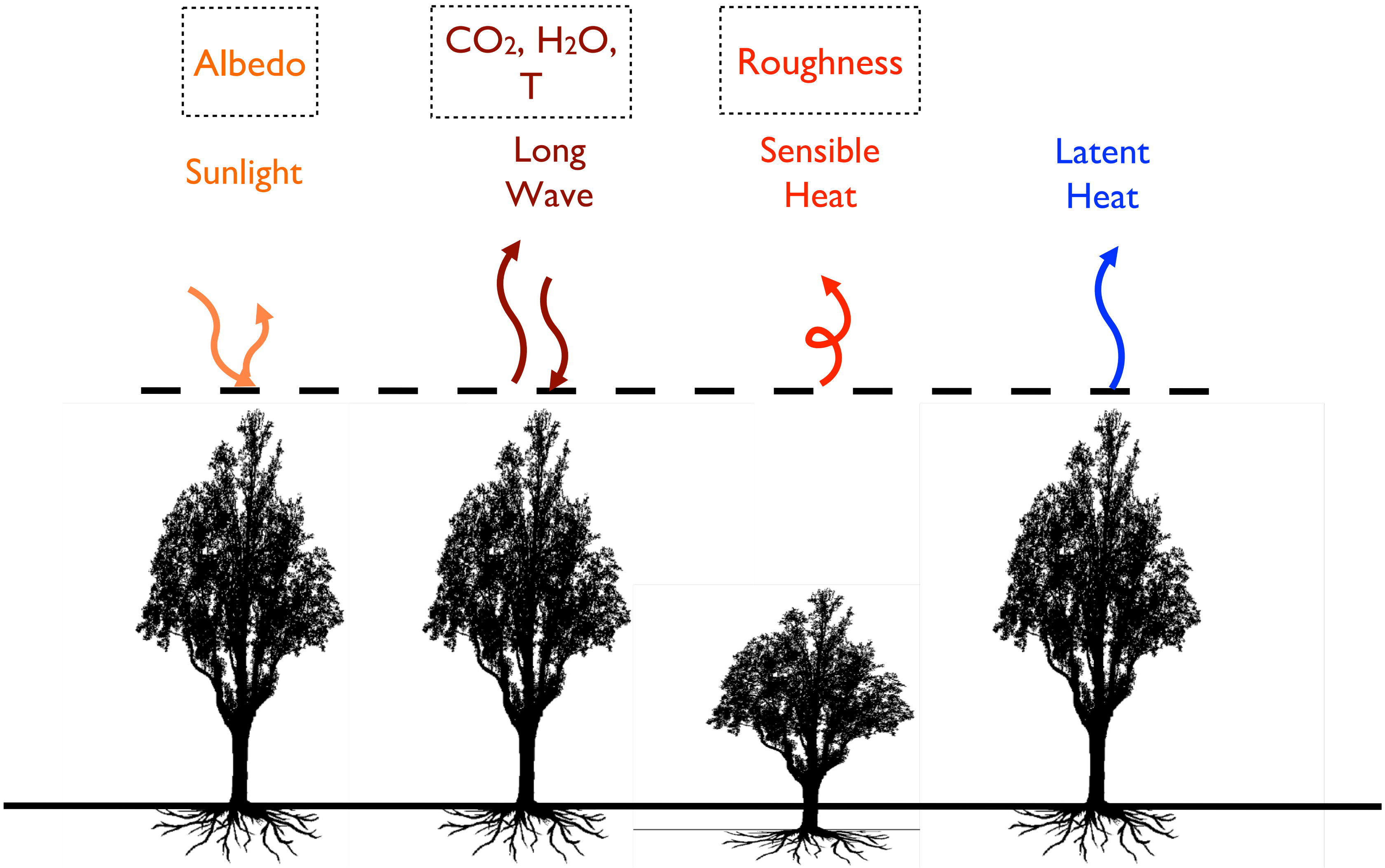
Sunlight

Long Wave

Sensible Heat

Latent Heat

How do Plants and Climate Interact?



Albedo

CO_2, H_2O, T

Roughness

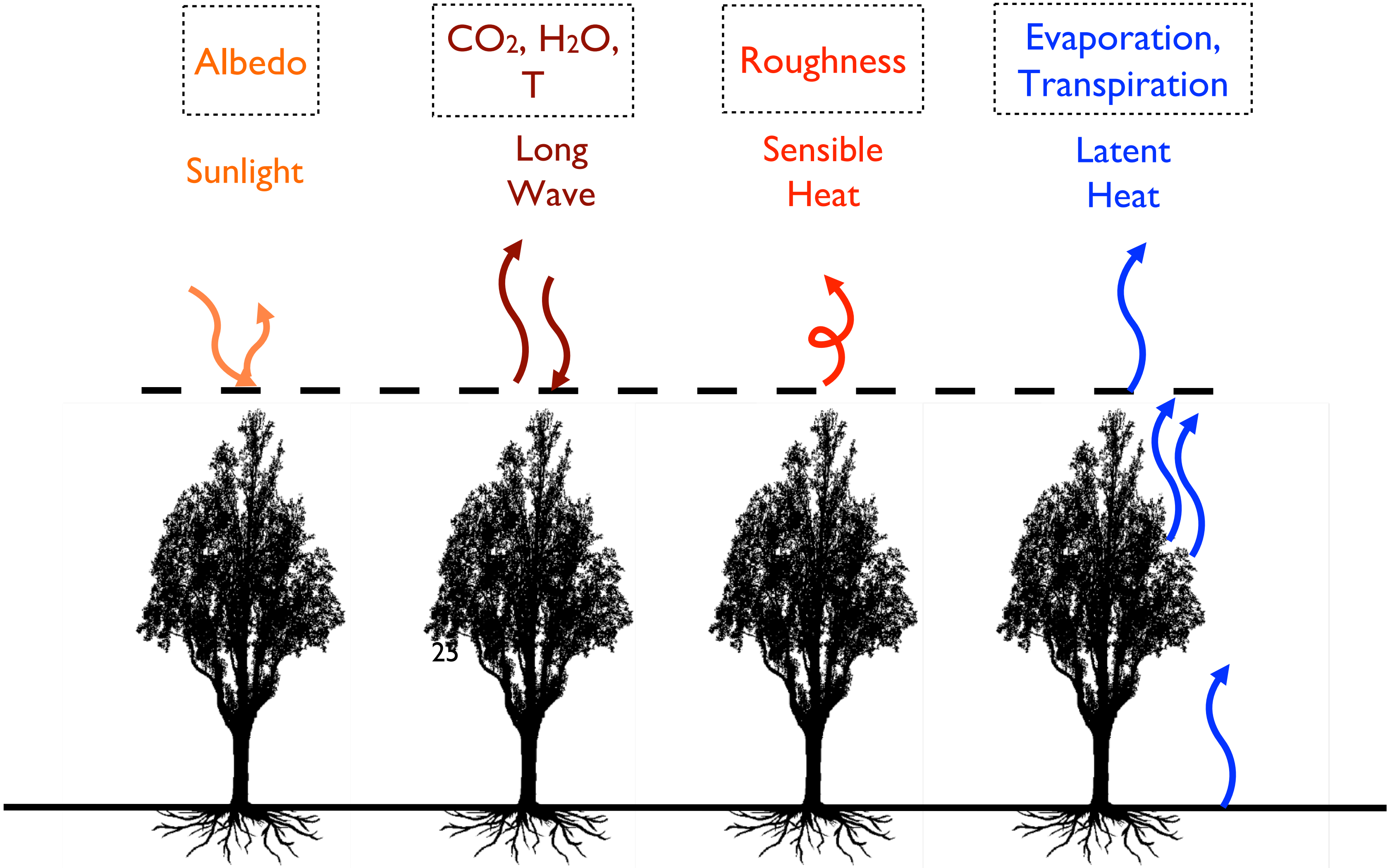
Sunlight

Long Wave

Sensible Heat

Latent Heat

How do Plants and Climate Interact?



Albedo

Sunlight

CO_2, H_2O, T

Long Wave

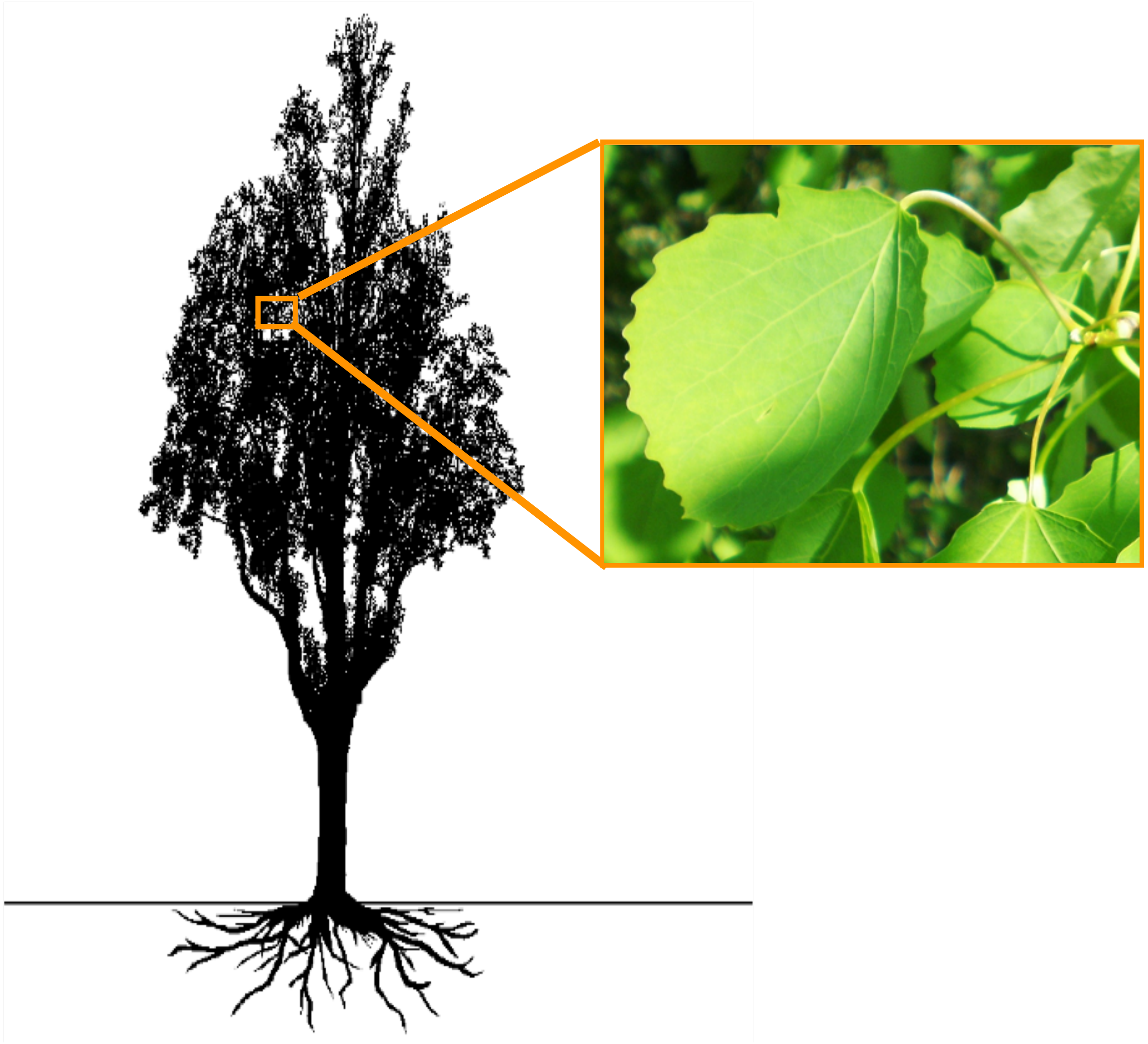
Roughness

Sensible Heat

Evaporation, Transpiration

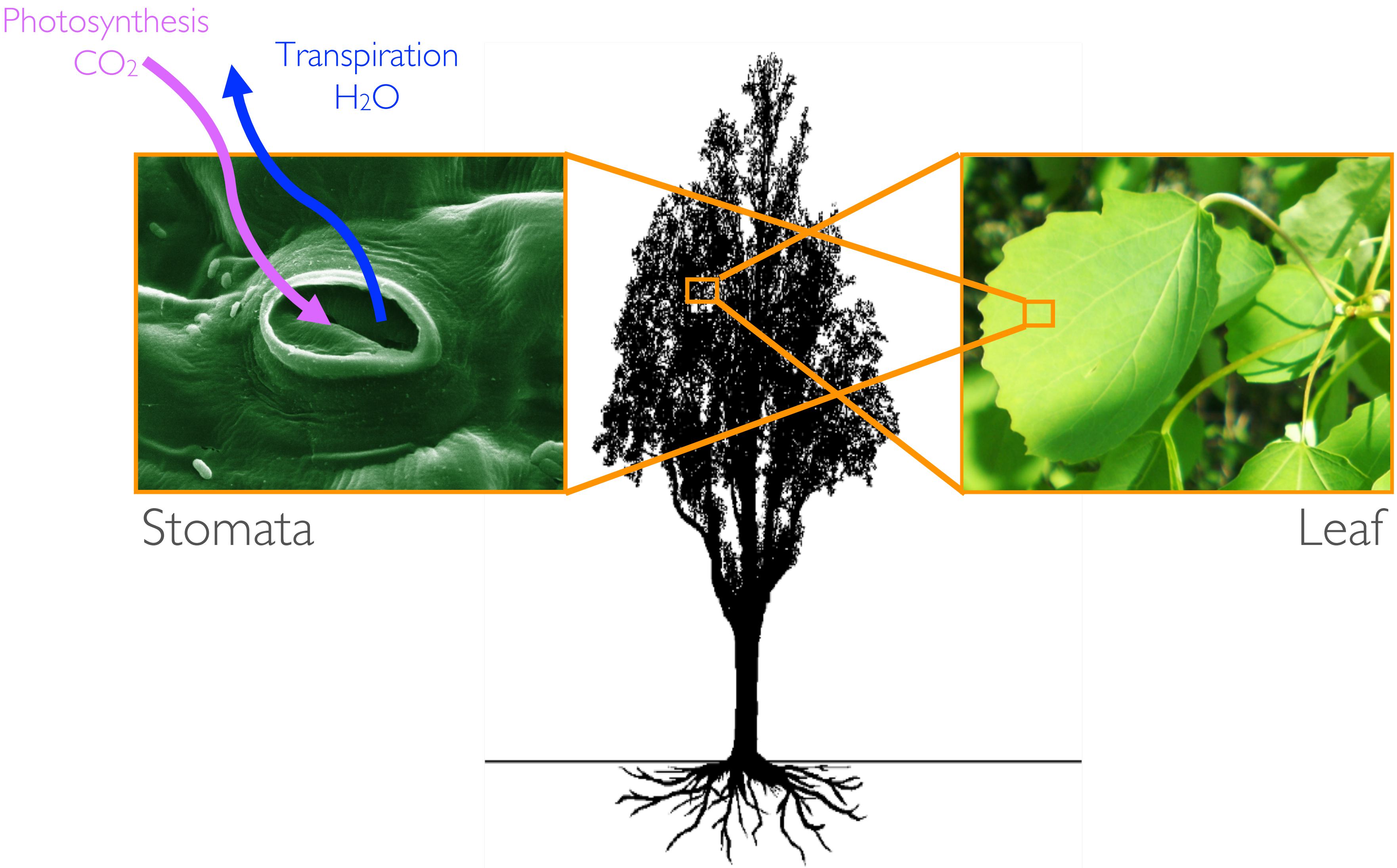
Latent Heat

Transpiration flux of water



Photos: Wikimedia Commons

Carbon in, water out

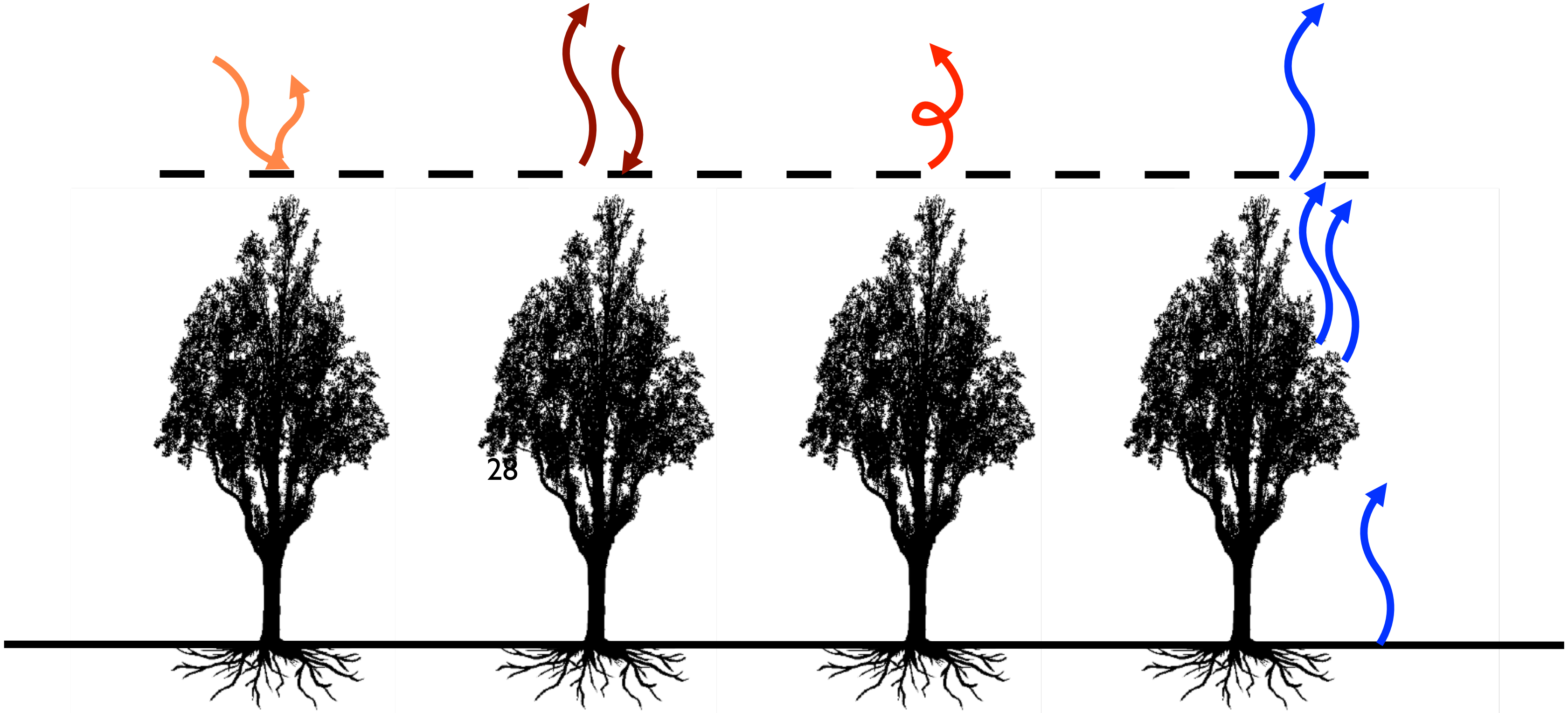


Plant physiological controls on CO_2 exchange and transpiration

Function of solar radiation, humidity deficit, soil moisture, [CO2], temperature, leaf N content

Photos: Wikimedia Commons

ΔPlants => ΔSurface Energy Budget



Albedo

Sunlight

CO_2, H_2O, T

Long Wave

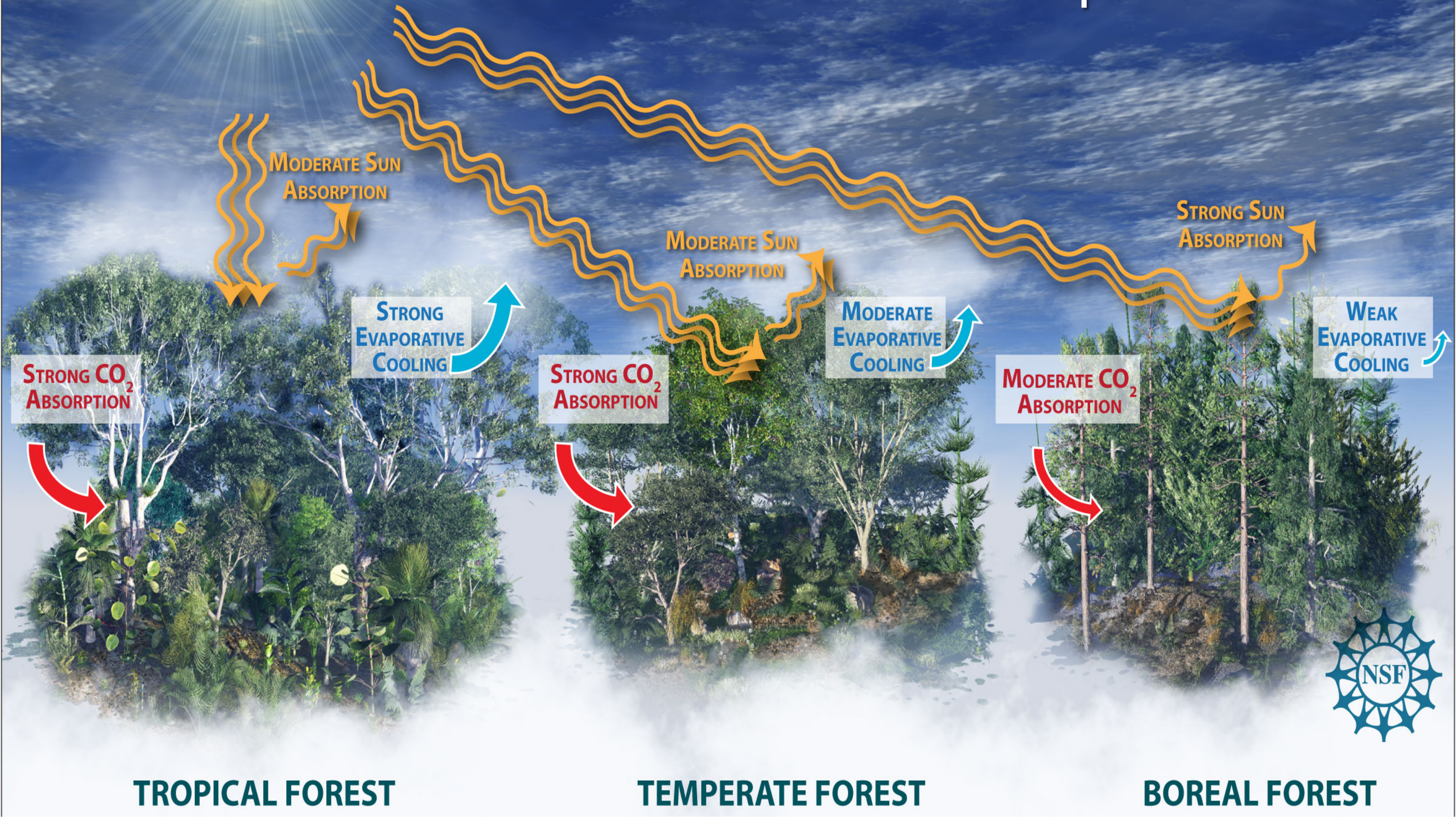
Roughness

Sensible Heat

Evaporation, Transpiration

Latent Heat

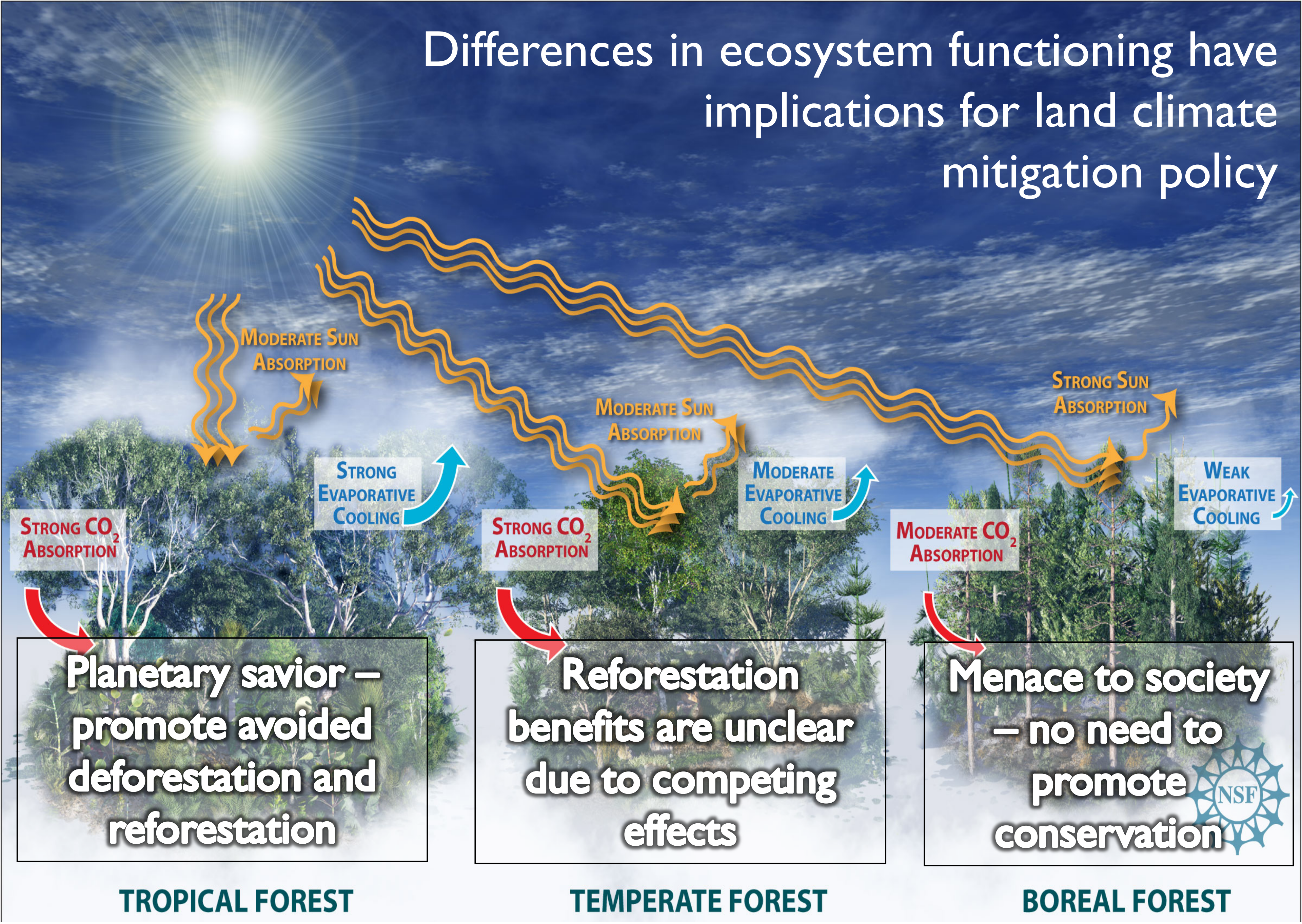
Land in Earth System:
Not all forest ecosystems have the same impact on climate



Bonan (2008) Science 320:1444-1449

Credit: *Nicolle Rager Fuller, National Science Foundation*

Differences in ecosystem functioning have implications for land climate mitigation policy



Planetary savior – promote avoided deforestation and reforestation

Reforestation benefits are unclear due to competing effects

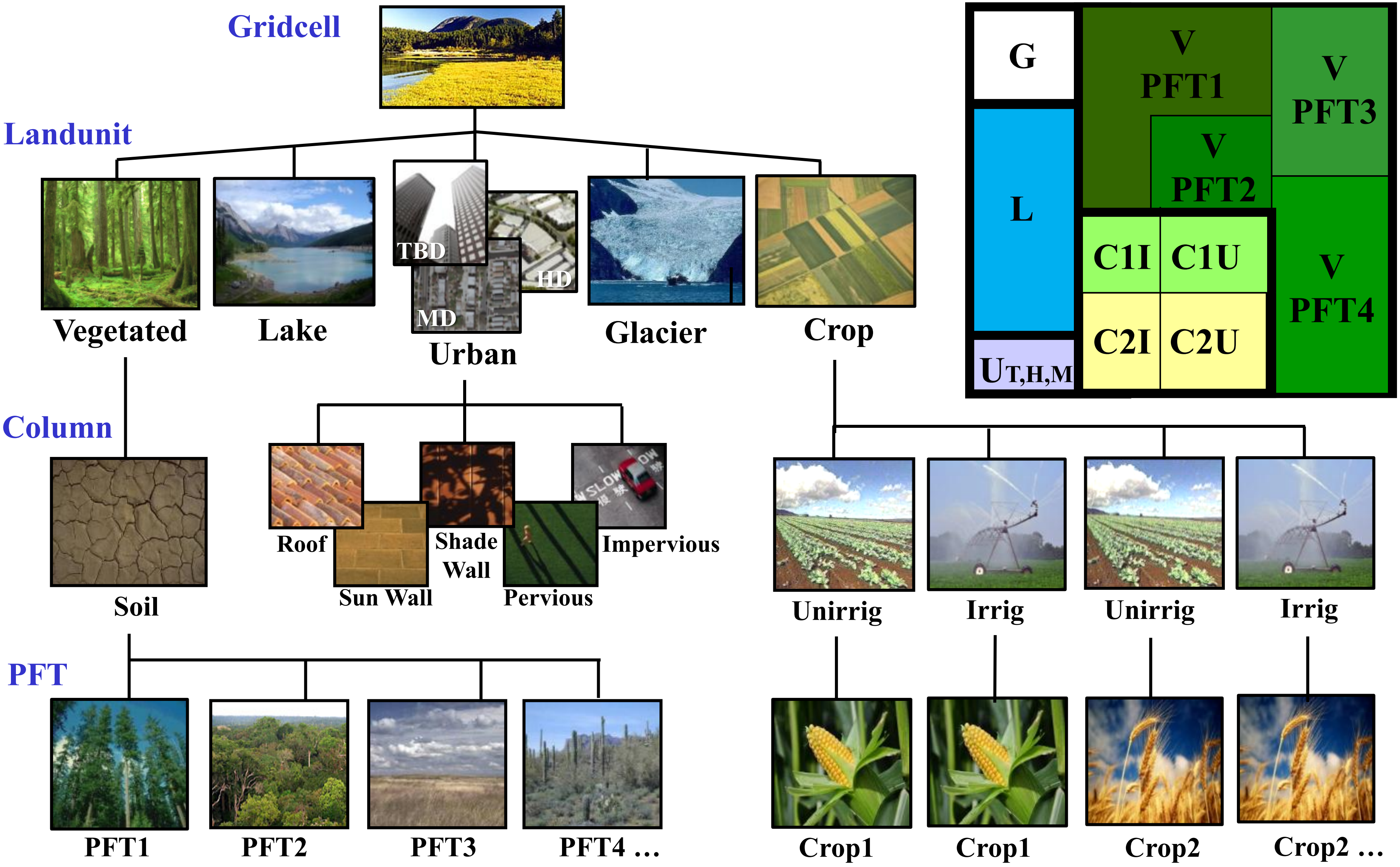
Menace to society – no need to promote conservation

Bonan (2008) Science 320:1444-1449

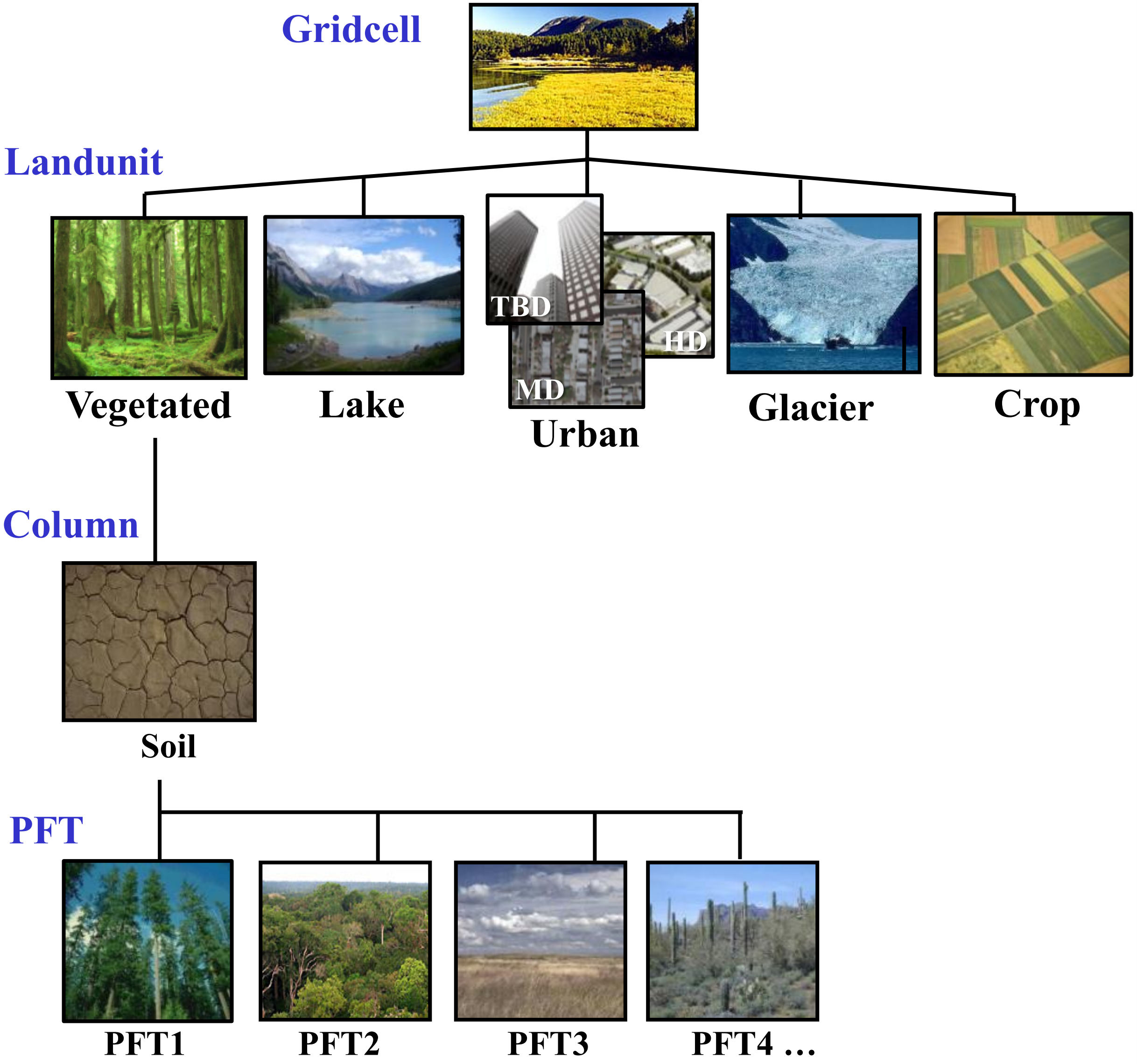
Credit: *Nicolle Rager Fuller, National Science Foundation*

Land Modeling Challenges: Land surface heterogeneity

Land surface heterogeneity
CLM subgrid tiling structure



Land surface heterogeneity
CLM subgrid tiling structure



Gridcell

Landunit

Vegetated

Lake

TBD

HD

MD

Urban

Glacier

Crop

Column

Soil

PFT

PFT1

PFT2

PFT3

PFT4 …

Plant Functional Types:

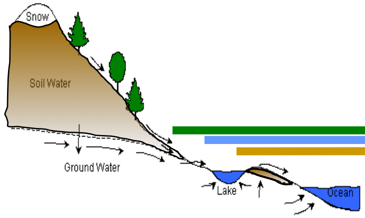
0. Bare

Tree:

1. Needleleaf Evergreen, Temperate
2. Needleleaf Evergreen, Boreal
3. Needleleaf Deciduous, Boreal
4. Broadleaf Evergreen, Tropical
5. Broadleaf Evergreen, Temperate
6. Broadleaf Deciduous, Tropical
7. Broadleaf Deciduous, Temperate
8. Broadleaf Deciduous, Boreal

Herbaceous / Understorey:

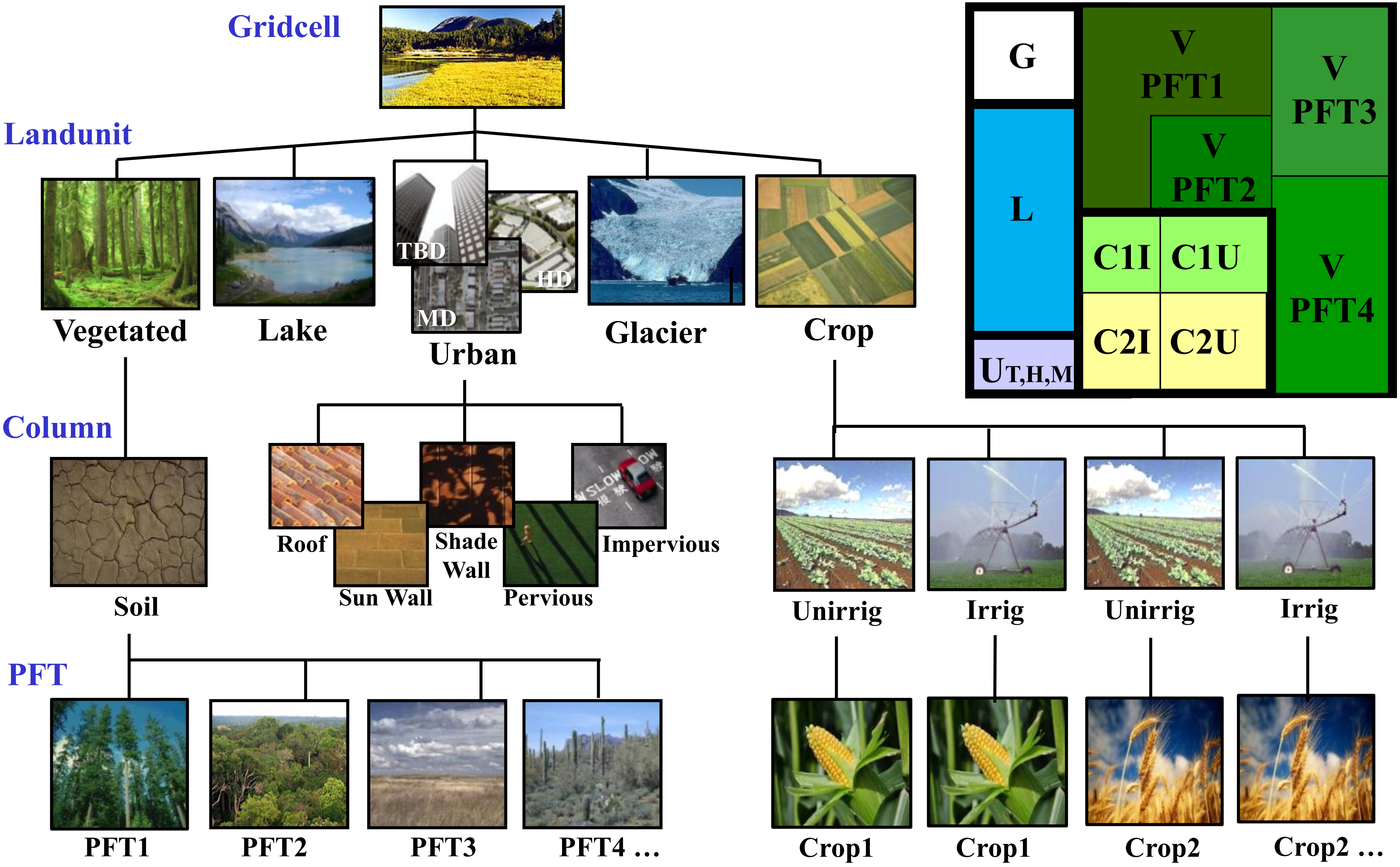
9. Broadleaf Evergreen Shrub, Temperate
10. Broadleaf Deciduous Shrub, Temperate
11. Broadleaf Deciduous Shrub, Boreal
12. C3 Arctic Grass
13. C3 non-Arctic Grass
14. C4 Grass
15. Crop

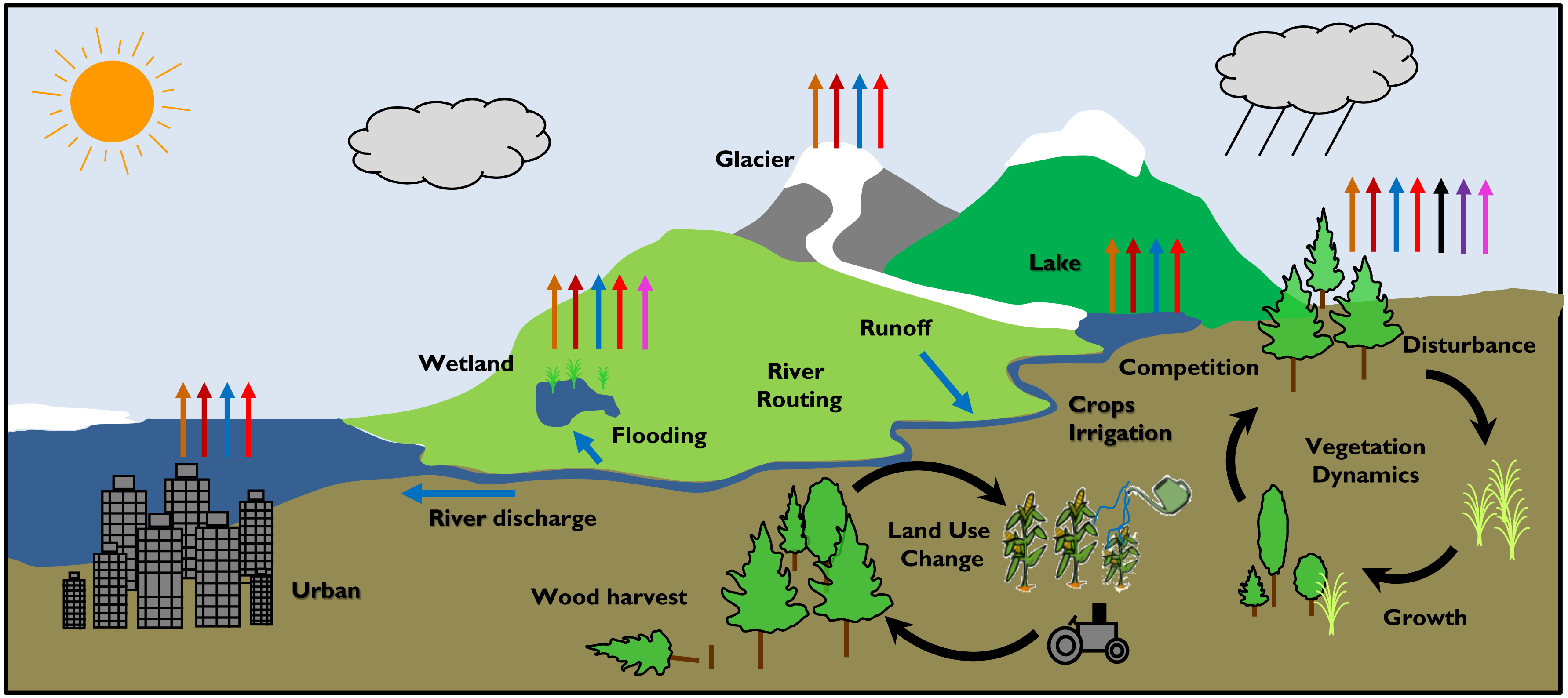


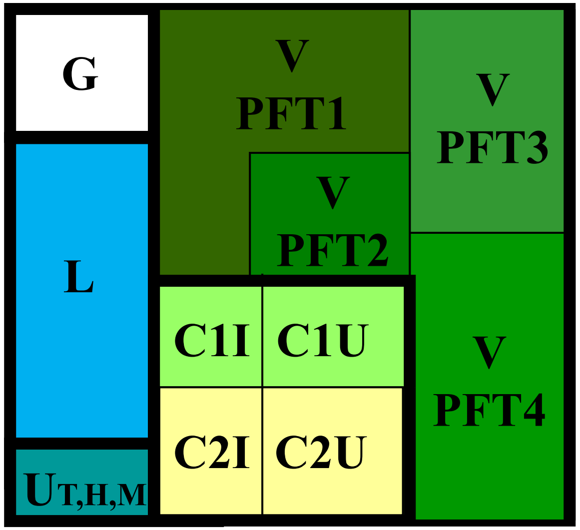
Plant Functional Type Parameters

- Optical properties (visible and near-infrared):
  - Leaf angle
  - Leaf reflectance
  - Stem reflectance
  - Leaf transmittance
  - Stem transmittance
- Fire:
  - Combustion completeness
  - Fire mortality
- Land models are parameter heavy!!!
- Morphological properties:
  - Leaf area index (annual cycle)
  - Stem area index (annual cycle)
  - Leaf dimension, leaf orientation
  - Roughness length/displacement height
  - Canopy top and bottom height
  - Root depth and distribution
- Photosynthetic parameters:
  - Specific leaf area
  - m (slope of conductance-photosynthesis relationship)
  - Vcmax (maximum rate of carboxylation)
  - Leaf carbon to nitrogen ratio
  - Fraction of leaf nitrogen in Rubisco
  - Root conductivity, plant conductivity

Land surface heterogeneity
CLM subgrid tiling structure



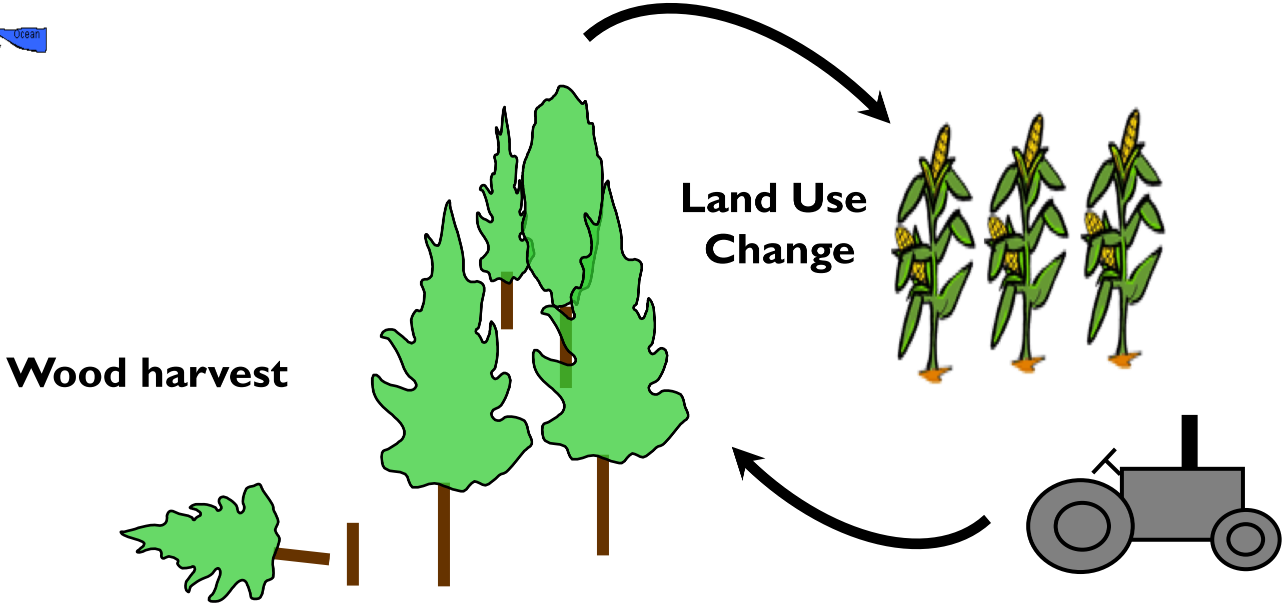


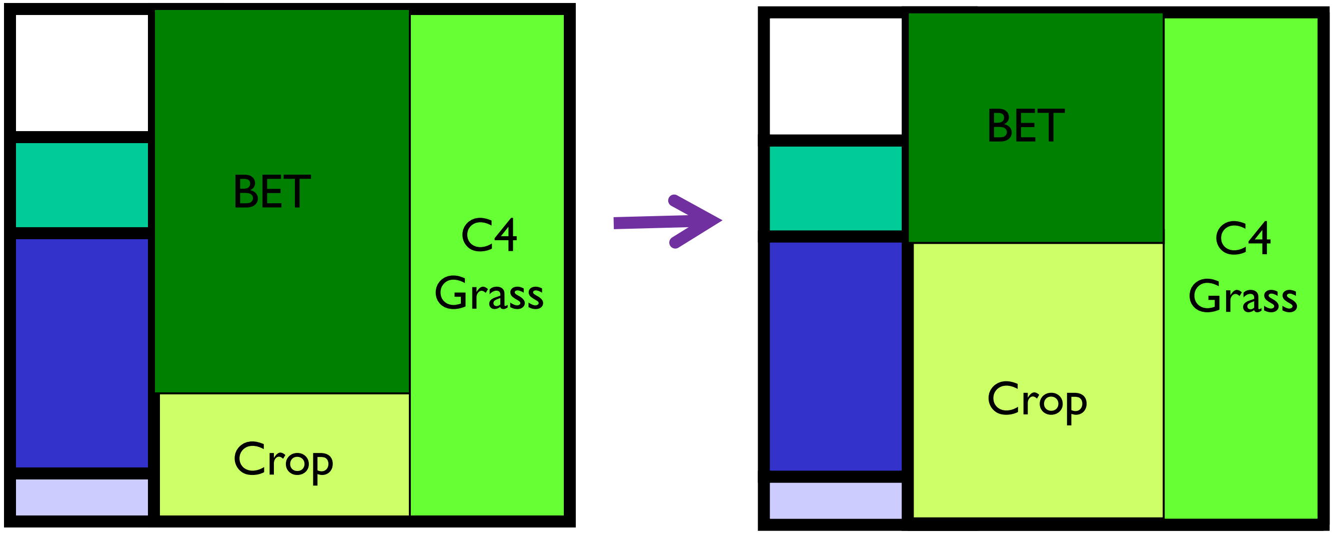


Landscape-scale dynamics

Long-term dynamical processes that affect fluxes in a changing environment (disturbance, land use, succession)

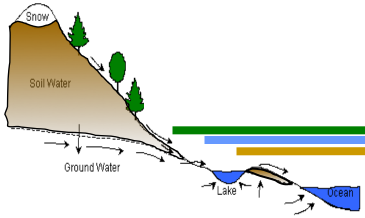
Land-cover / land-use change (prescribed)





Land Modeling Challenges: Land surface heterogeneity

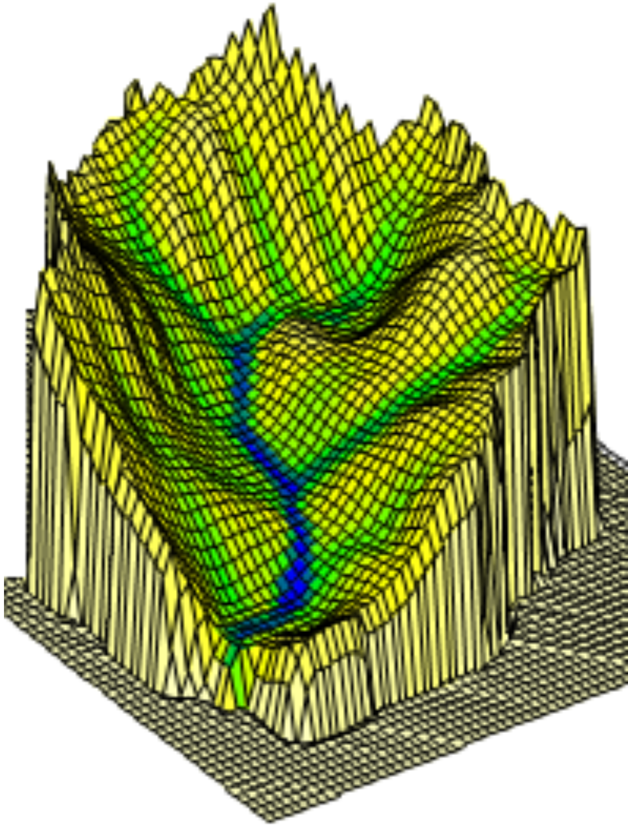
Natural vegetation patterns imply controls from soil moisture convergence, slope, and aspect



Parameterize impact of subgrid-scale soil moisture heterogeneity

A major control on soil moisture heterogeneity and thus runoff is topography.

Lowland soils tend to be zones of high soil moisture content, while upland soils tend to be progressively drier.

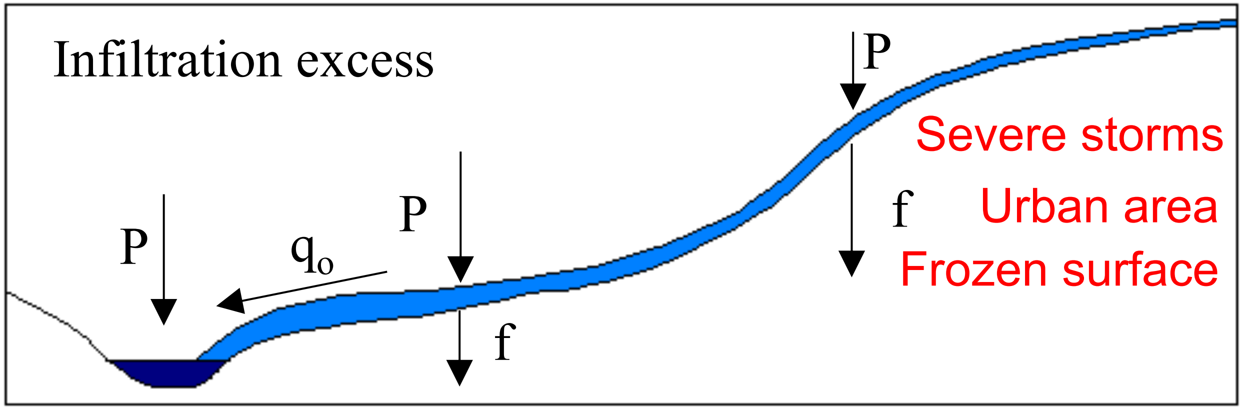


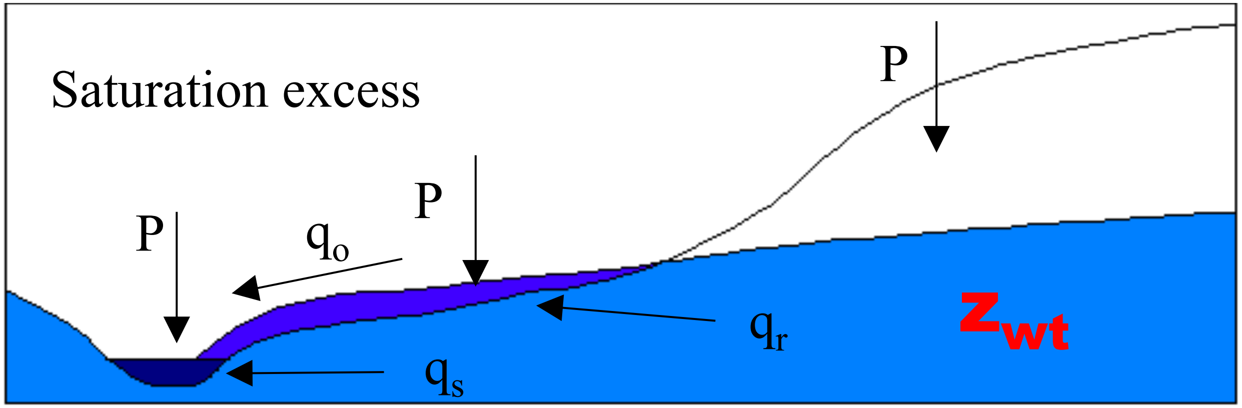
Three main sources of runoff:

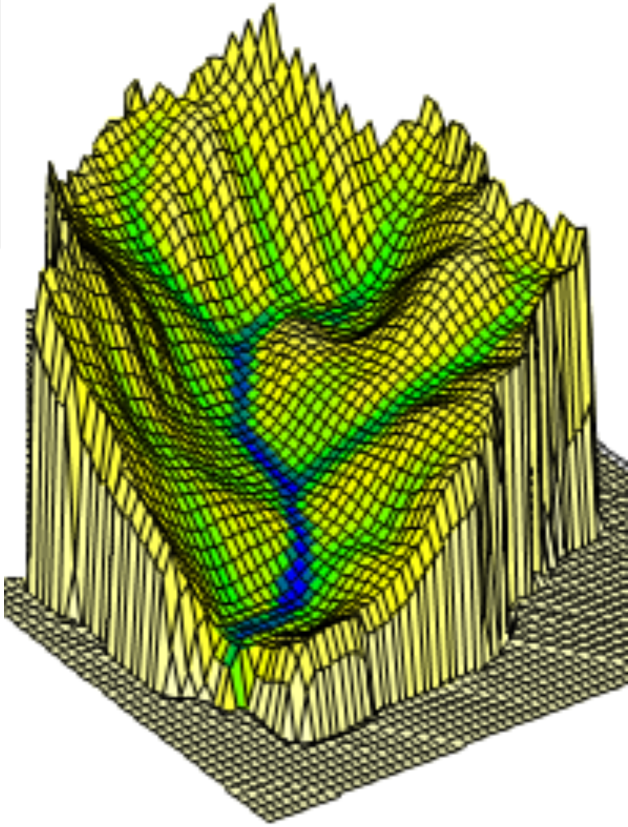
- Infiltration excess
- Saturation excess
- Baseflow (drainage)

Subgrid-scale soil moisture heterogeneity

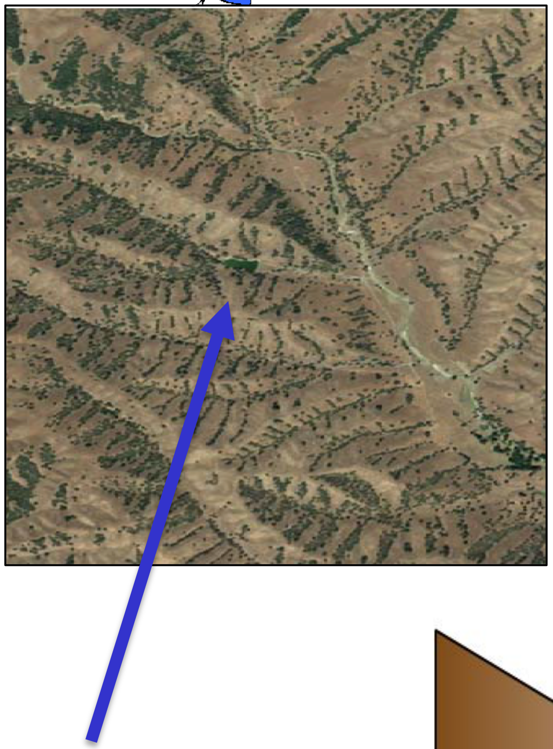
SIMTOP: TOPMODEL-based runoff

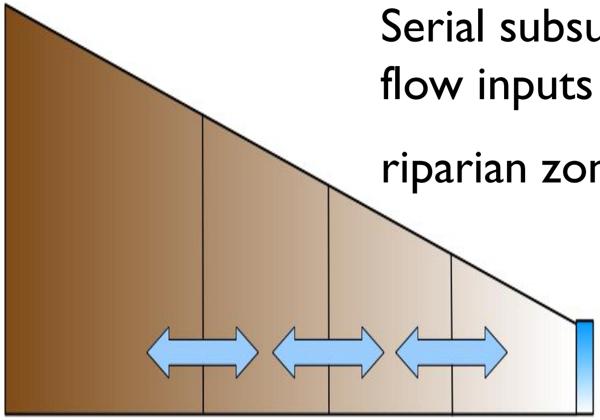




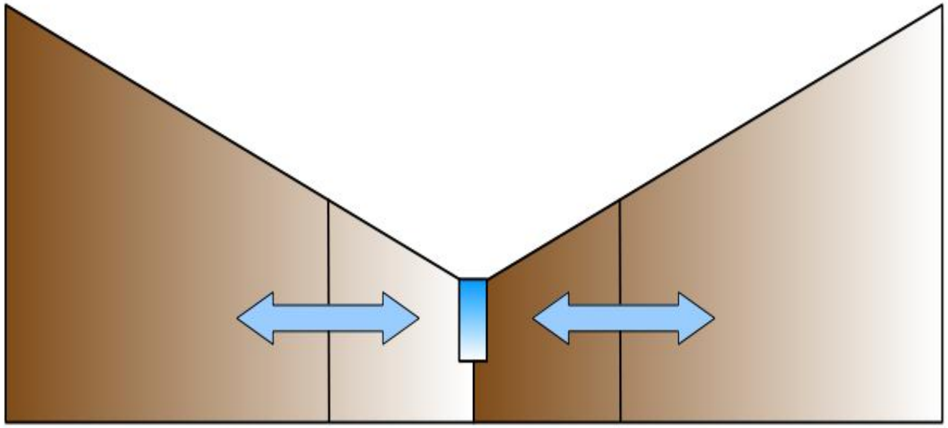


Representative hillslopes (CLM5 option)





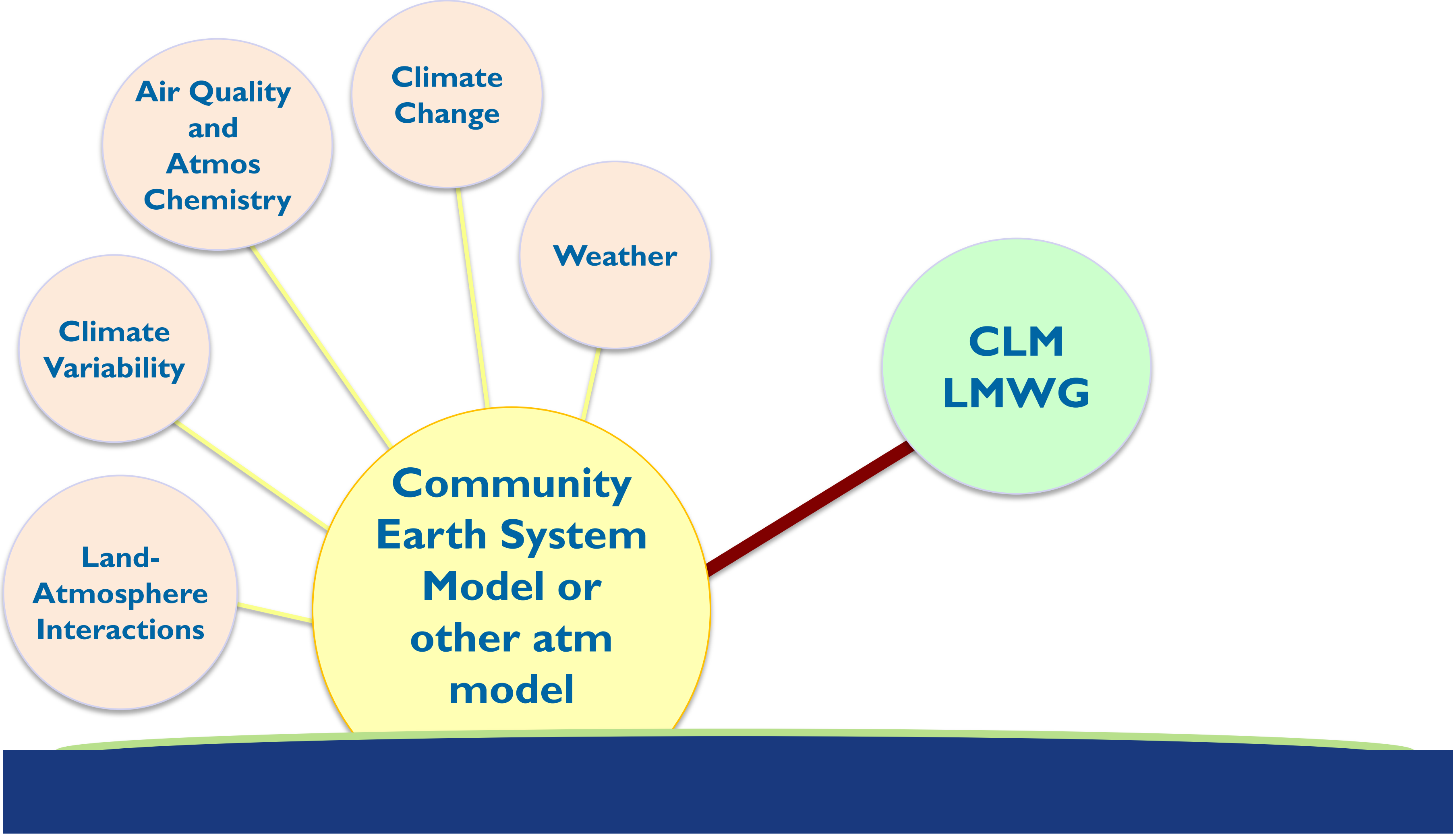
Serial subsurface flow inputs to riparian zone



Independent (parallel) subsurface flow inputs to riparian zone

Model development and assessment

CLM as a community modeling tool



Air Quality and Atmos Chemistry

Climate Change

Weather

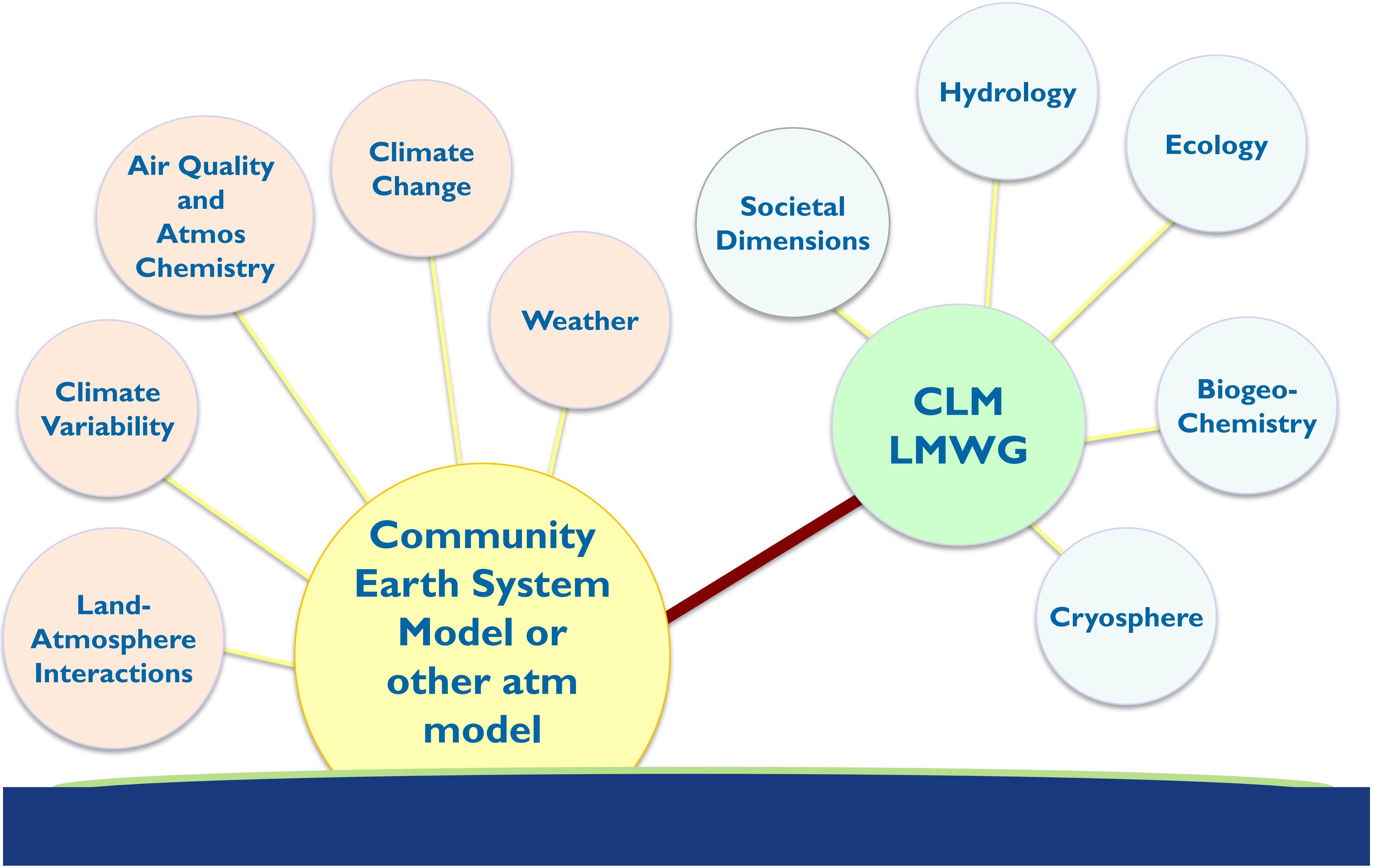
Climate Variability

Land-Atmosphere Interactions

Community Earth System Model or other atm model

CLM LMWG

CLM as a community modeling tool



Community Earth System Model or other atm model

Land-Atmosphere Interactions

Climate Variability

Air Quality and Atmos Chemistry

Climate Change

Weather

CLM LMWG

Societal Dimensions

Hydrology

Ecology

Biogeo-Chemistry

Cryosphere

CLM4 (June 2010)

- Carbon and nitrog
  prognostic vegetat
- Transient land cove
  wood harvest
- ‘Permafrost-enable
  deep ground
- Aerosol depositior
- Simple groundwate
- Urban model

CLM4.5 (June 2013)

- Vertically-resolved soil C/N
- Co-limitation and acclimation of photosynthesis
- Variable river flow rates
- Natural CH_4 emissions
- Human triggering and suppression of fire
- Cold region hydrology
- Revised lake model
- Multiple urban density classes

What's New for CLM5

A LOT!

More than 50 researchers from 15 different institutions have been involved in development of CLM5

Parallel focus on mechanistic improvements and expansion of capabilities

- hydrology more consistent with state-of-art understanding
- more ecologically relevant plant nutrient, carbon, and water dynamics
- land management including global crop model, wood harvest, urban environments
- prognostic Greenland ice sheet model
- …

CLM4 (June 2010)

- Carbon and nitrog
  prognostic vegetat
- Transient land cove
  wood harvest
- ‘Permafrost-enable
  deep ground
- Aerosol depositior
- Simple groundwate
- Urban model

CLM4.5 (June 2013)

- Vertically-
- Co-limitat
  of photosy
- Variable ri
- Natural C
- Human tri
  suppressic
- Cold regic
- Revised la
- Multiple u
  classes

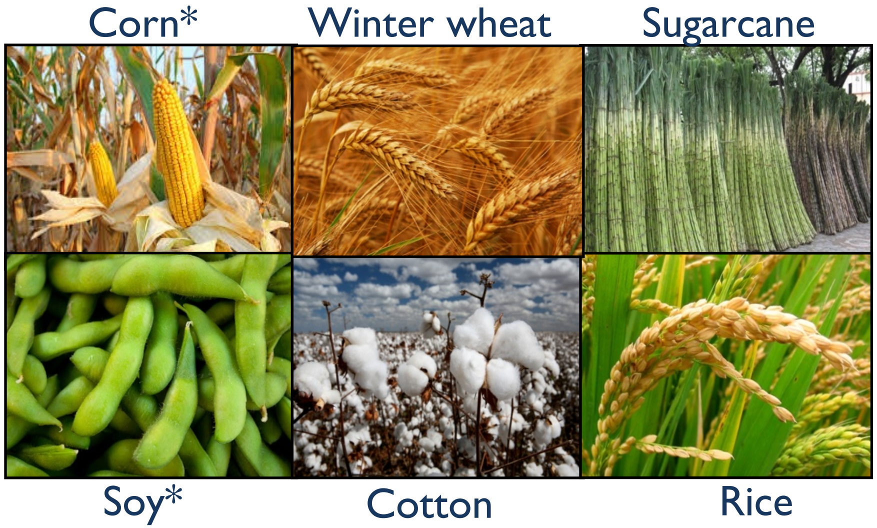
CLM5 (Feb 2018)

- Flexible leaf stoichiometry, leaf N optimize for photosynthesis
- Carbon costs for plant N uptake
- Plant hydraulics w/ hydraulic redistribution, *Ecosystem demography (FATES), ozone damage*
- Spatially explicit soil depth (0.4 – 8.5m), dry surface layer, revised GW, canopy interception, *representative hillslopes*
- MOSART river model (hillslope → tributary → main channel)
- Canopy snow, snow dens (T, wind), simple firn model
- Global crop model (8 crop types), transient irrigation and fertilization, *shifting cultivation*
- Dynamic landunits (nat veg ←→ crop, *glacier ←→ nat veg*,)
- Urban heating and AC, heat stress indices
- Carbon isotopes
- *Coupled fire trace gas emissions*

Land management in CLM5

<u>Included in default CLM5</u>

- Global crop model with 8 basic crop types; planting, grain fill, harvest
- Crop irrigation
- Crop industrial fertilization
- Wood harvest
- Urban environments
- Anthropogenic fire ignition and suppression



** Temperate and tropical varieties*

Fertilization

Irrigation

CLM4 (June 2010)

CLM4.5 (June 2013)

CLM5 (Feb 2018)

<u>A central challenge: Model assessment</u>

Are land models getting better or just more complex?

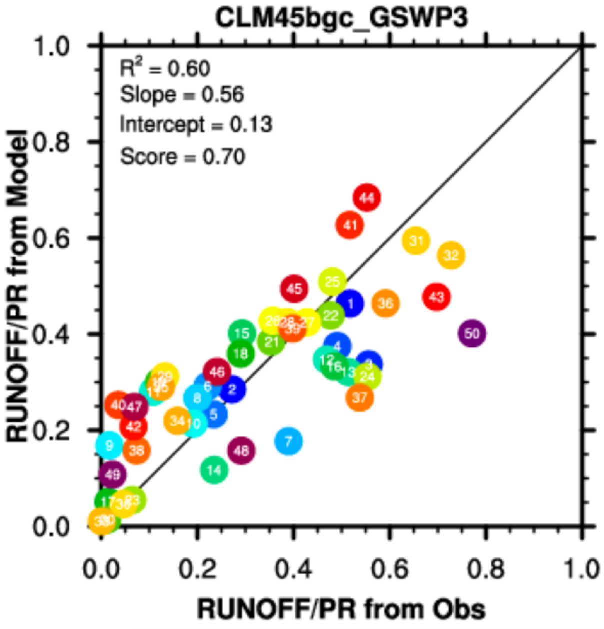
Do land models need to be more complex to be better?

How do we interpret results from disparate set of models with varying degrees of comprehensiveness and complexity?

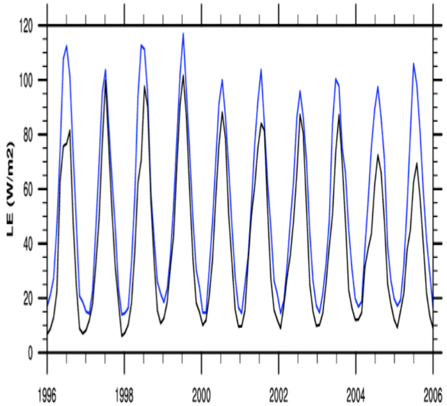
CMIP5 models, TRENDY models

International Land Model Benchmarking (ILAMB) Package

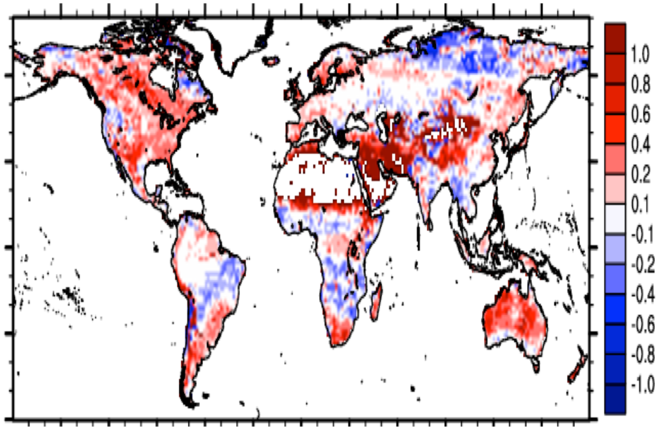
Land diagnostics package (25+ variables, 60+ datasets) with metrics for
RMSE, bias, spatial pattern corr, interannual variability, functional relationships



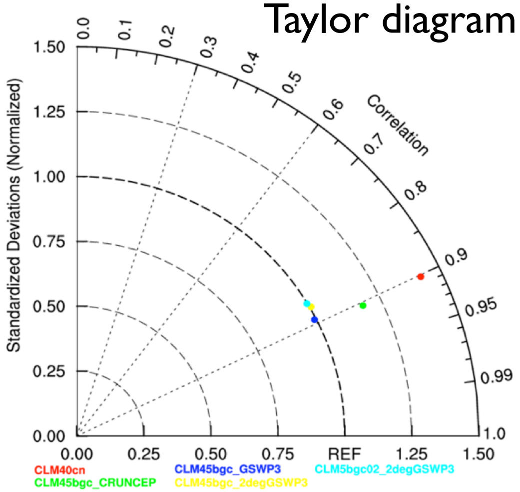
Tower Site



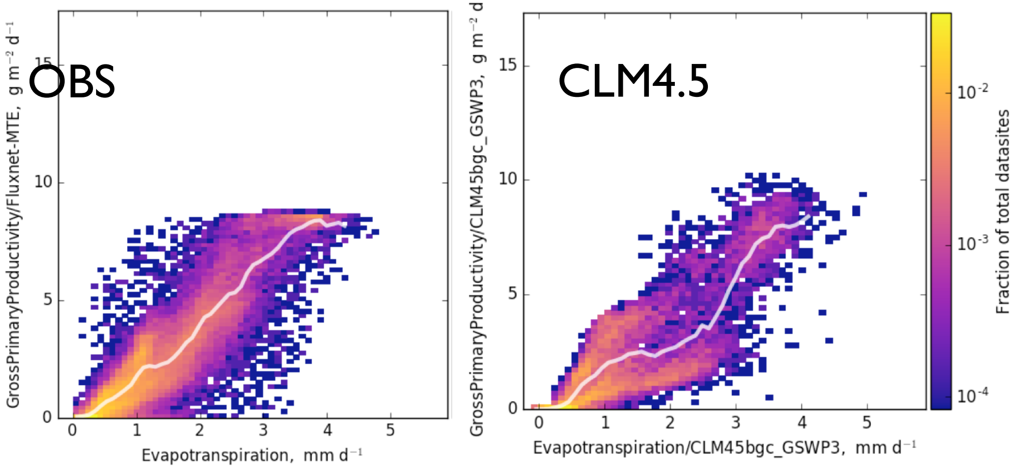
Global bias, relative bias, RMSE



Taylor diagram

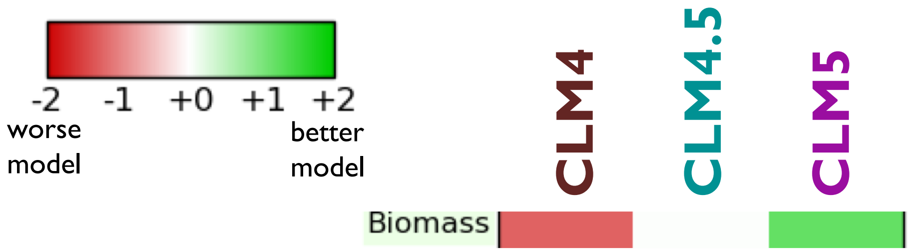


Functional relationships



Collier, Hoffman, Randerson, Lawrence, Mu, Koven et al., 2018

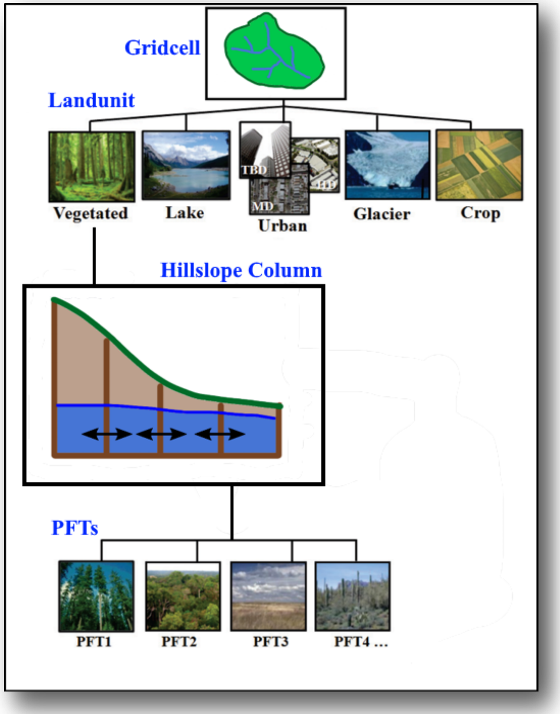
CLM land-only forced with GSWP3 prognostic vegetation and carbon configuration



- For majority of variables, progression in simulation quality from CLM4 to CLM5
- Why?
  - Improvements in mechanistic treatment of processes (e.g., hydrology, plant N processes, land use)
  - But, at same time, many more moving parts, additional unconstrained parameters

Some priorities and plans for next generation CLM

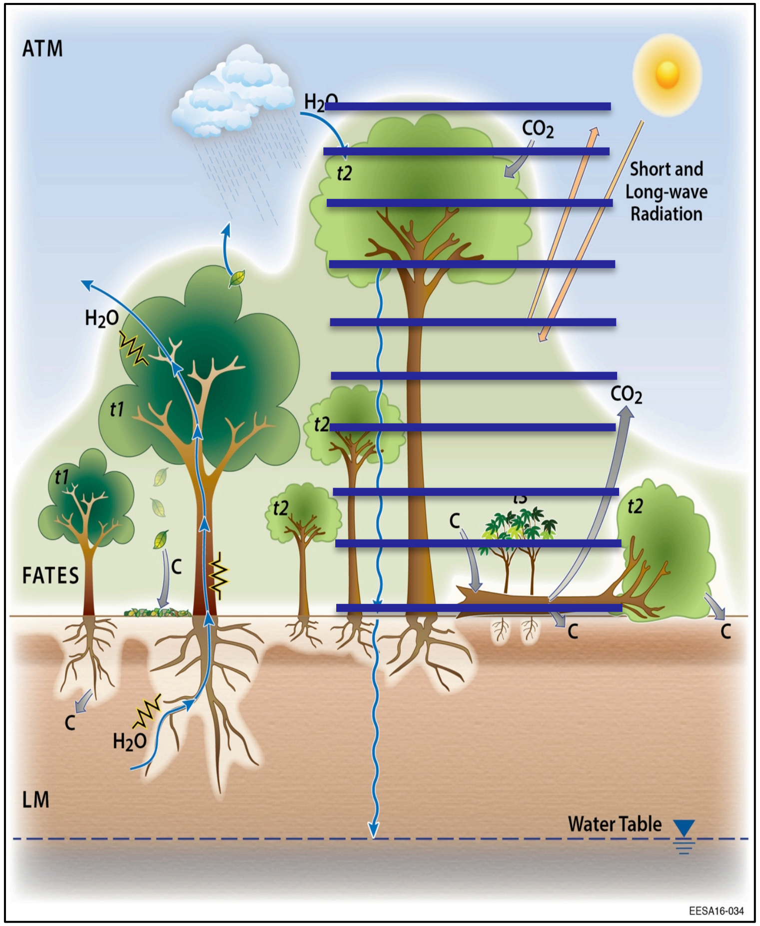
- Water and food security in context of climate variability, change, and extreme weather
- Ecosystem vulnerability and impacts on carbon cycle and ecosystem services
- Sources of predictability from land processes
- Impacts of land use and land-use change on climate, carbon, water, and extremes



Lateral fluxes of water

Water and land management

Ecosystem Demography / Multi-layer canopy



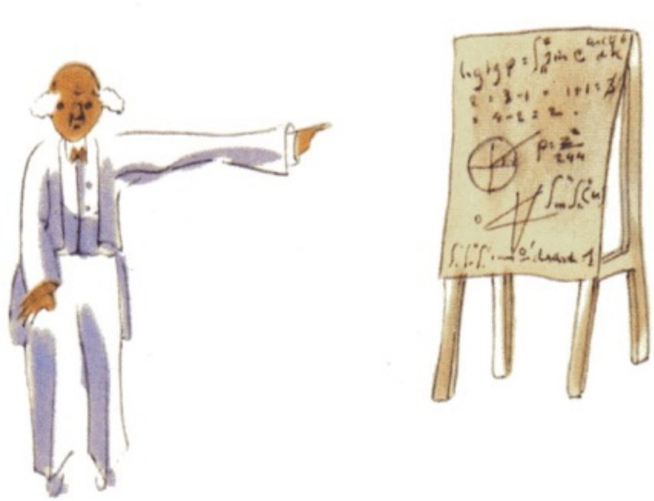
Questions?

Extra slides

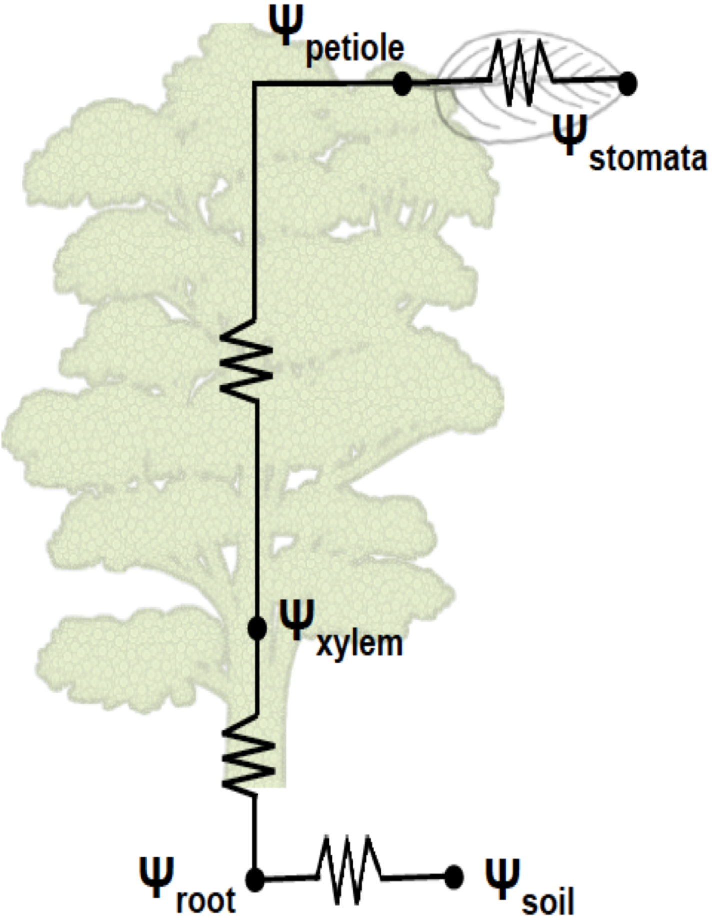
Modeling caveats

CLM (CESM) is just a starting point for the science. It is not the science itself

- Easy to run the model and get an answer
- Much harder to understand why you got that answer
- CLM is a very complex, multidisciplinary model

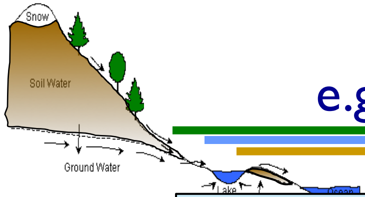


CLM5: Plant Hydrodynamics

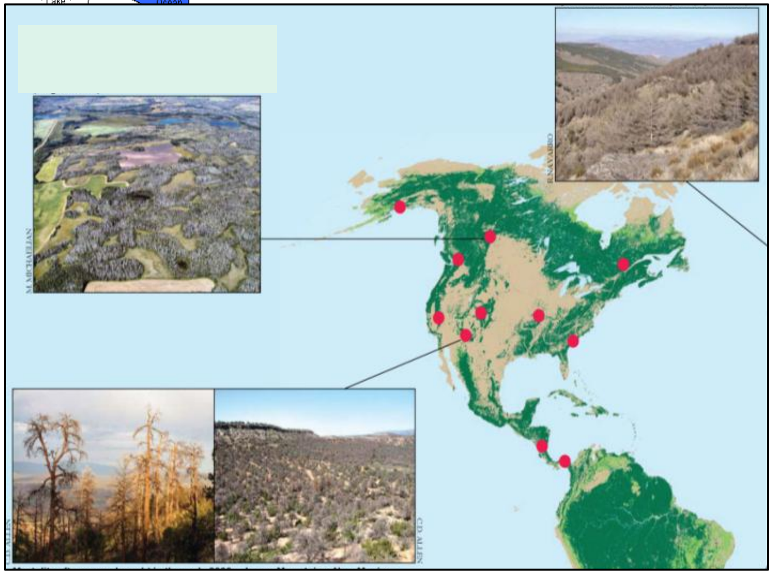


Why plant hydrodynamics

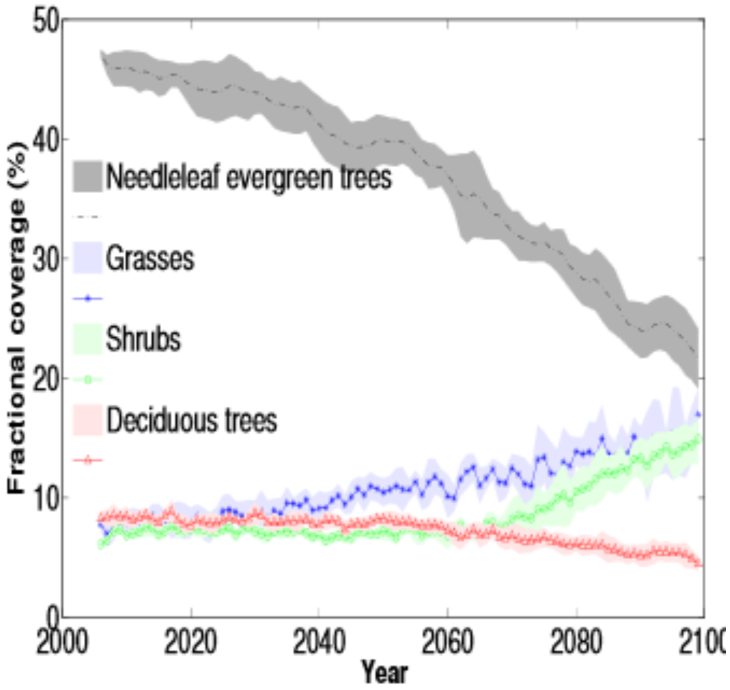
- BTRAN (soil moisture stress), and it’s parameters, θ_{crit} and θ_{wilt}, have no physical meaning and cannot be measured.
- Flux tower ET convolutes transpiration with canopy and soil evap making it difficult to use for process-level assessment. With plant hydrodynamics, sap flow measurements could be utilized.
- Satellites increasingly observe properties related to canopy or leaf water content.



Ecosystem vulnerability to climate change
e.g., how vulnerable are western US forests to climate change?



CLM4(DGVM), suggests widespread die-off of forests by 2100, but simple representation of hydrology, plant water use, mortality, ecosystem dynamics



Jiang et al. 2013

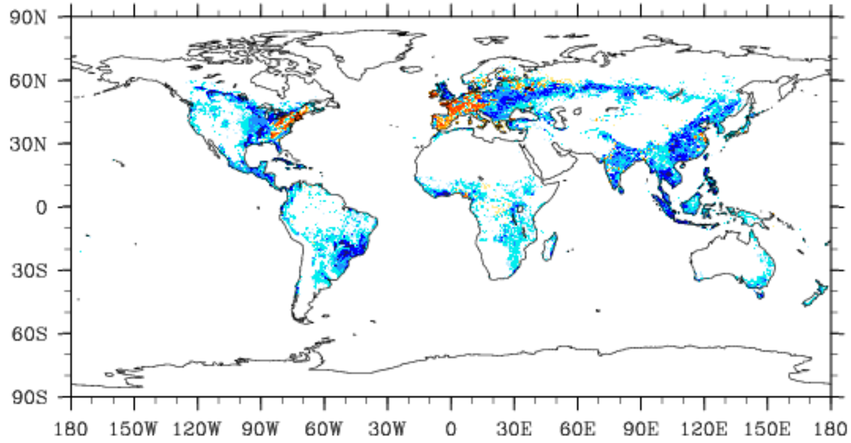
But … these results are likely unreliable; tree response to soil moisture deficits represented in ad hoc way in land models. Forest loss is complex problem that requires combined consideration of climate, hydrology, ecology, and plant physiology and diversity

Historical land use & land cover change, 1850-2005

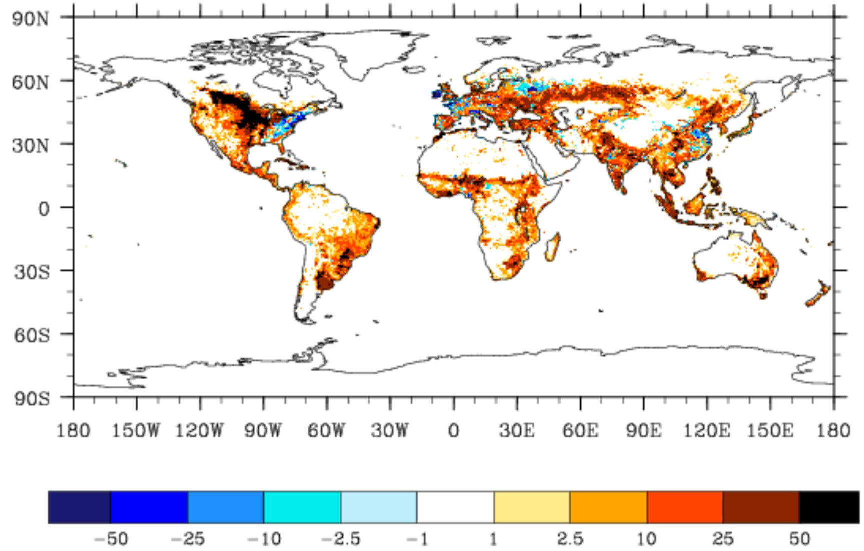
Change in tree and crop cover (% of grid cell)

Cumulative percent of grid cell harvested

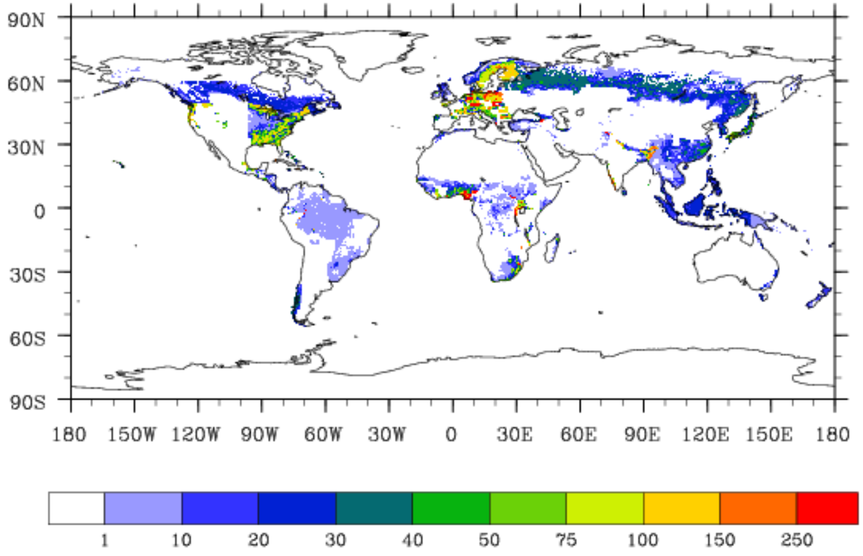
(a) Historical (2005-1850) Tree PFTs %



(a) Historical (2005-1850) Crop PFTs %



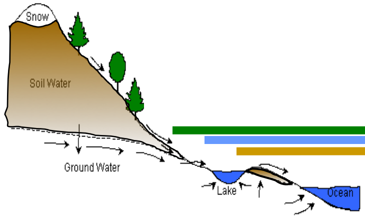
(b) Historical (2005-1850) Tree PFT Harvest %



Historical LULCC

- ❑ Loss of tree cover and increase in cropland
- ❑ Farm abandonment and reforestation in eastern U.S. and Europe
- ❑ Extensive wood harvest

P. Lawrence et al. (2012) J Climate 25:3071-3095



Many paths to improve models and reduce model uncertainty

Model intercomparisons (MIPs)
- CMIP6: carbon cycle, land use, land-atmosphere coupling, …
- Range of plausible outcomes, but more models ≠ better results

Model benchmarking
- Comprehensive model evaluation against observations

Real-world experiments and models
- FACE, N addition

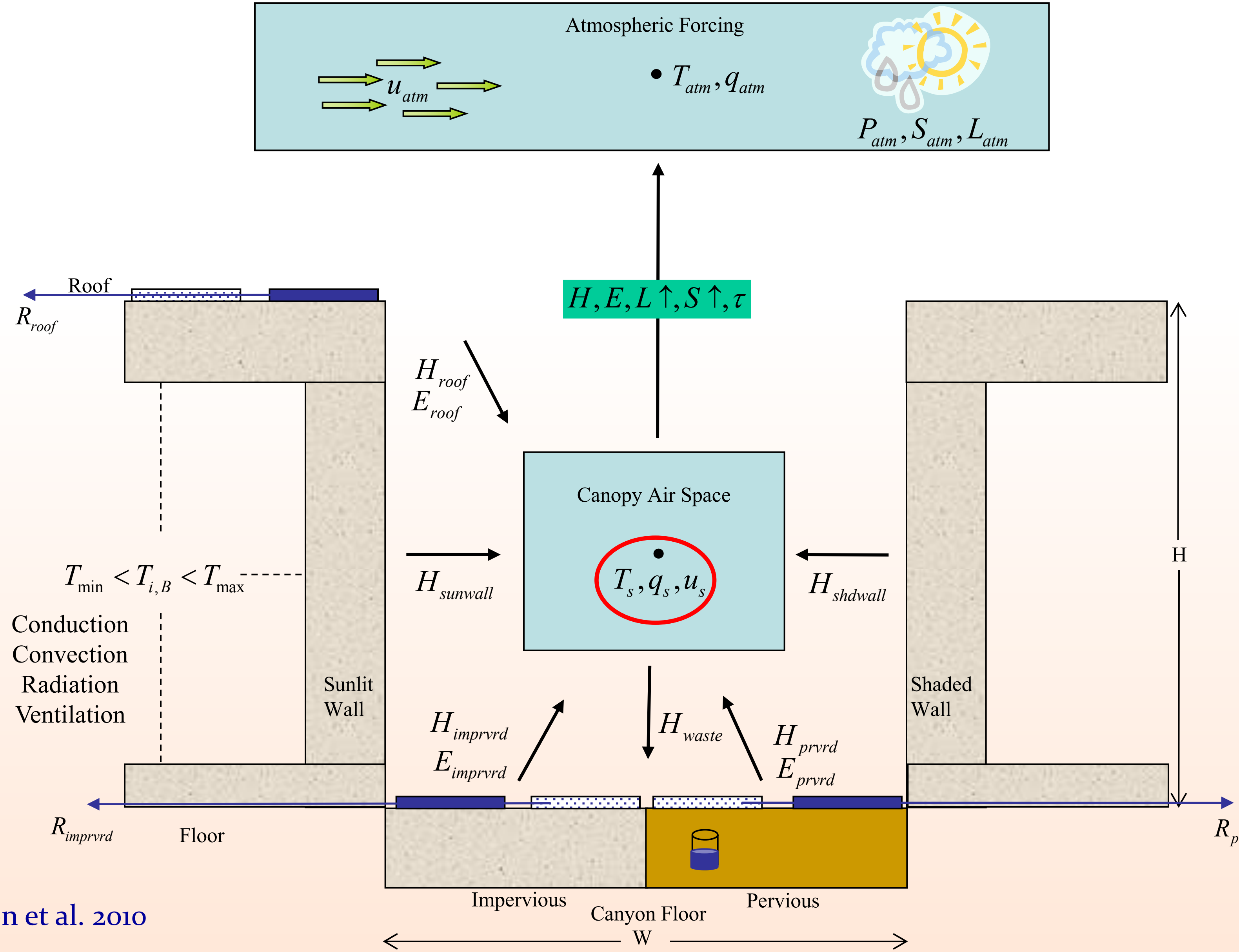
Model-data fusion
- Data assimilation, parameter estimation

“Discover” critical missing process
- Add another process that is ecologically or hydrologically important but poorly known at the global scale. Tune a key parameter to get a good simulation.

Model intracomparison
- Focus on model structural uncertainty to identify processes contributing to uncertainty

Model hierarchy
- CLM
- Process models (multilayer canopy, MIMICS)
- Simple land models (Marysa Lague)

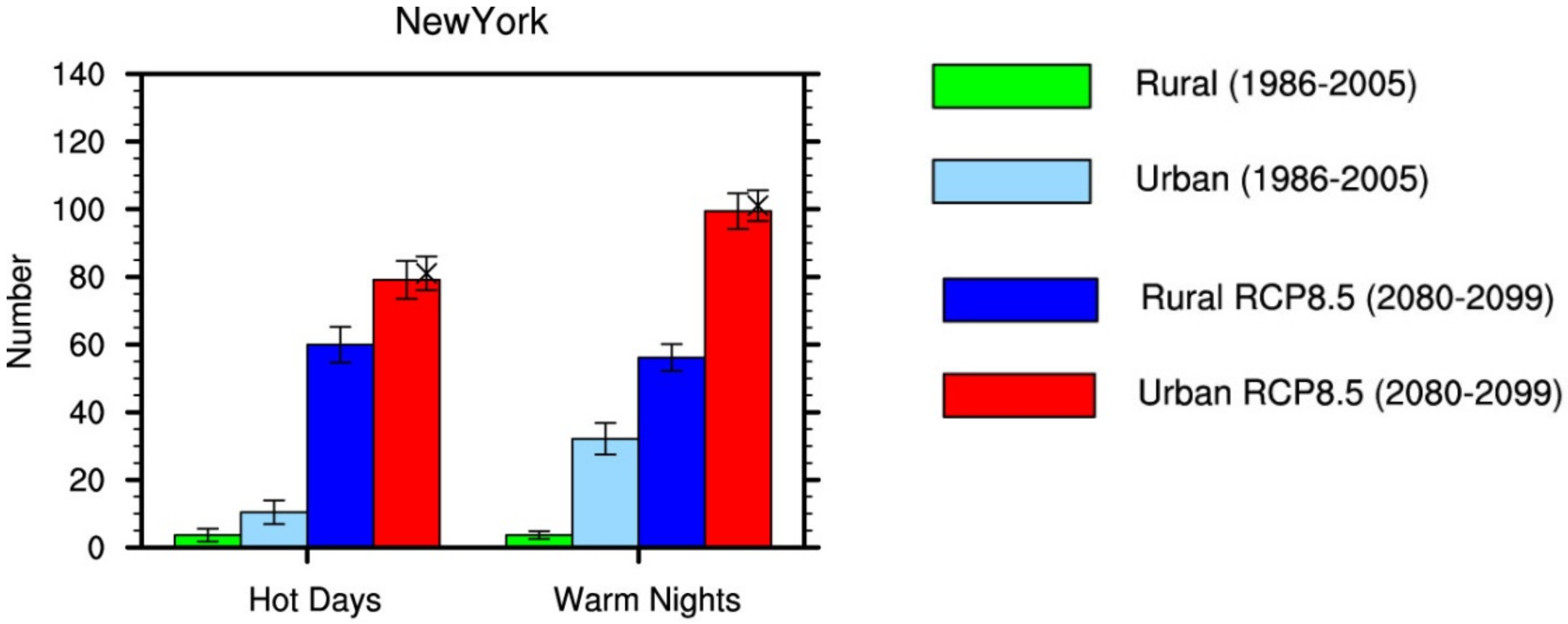
Urban Model



Oleson et al. 2010

Changes in hot days and warm nights – RCP8.5

Hot days (warm nights) – Number of days per year that daily TMAX (TMIN) exceeds 99th percentile of present day Rural daily TMAX (TMIN)



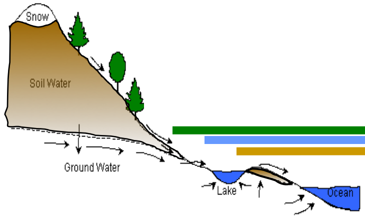
Present-day climate

Cities have more hot days and warm nights than rural land

21st century climate change

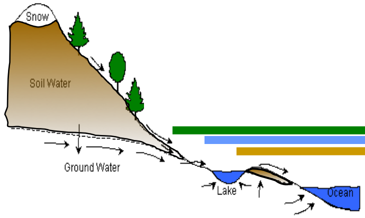
Cities increase more in hot days and warm nights than does rural land

Slide courtesy K. Oleson



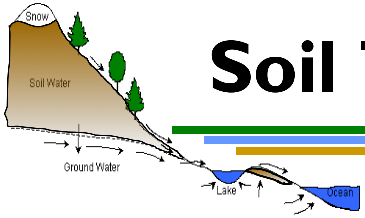
The role of CLM in CESM: Land to Atmosphere

[1]Latent heat flux	$\lambda_{vap}E_v + \lambda E_g$	W m^{-2}
Sensible heat flux	$H_v + H_g$	W m^{-2}
Water vapor flux	$E_v + E_g$	mm s^{-1}
Zonal momentum flux	τ_x	kg m^{-1} s^{-2}
Meridional momentum flux	τ_y	kg m^{-1} s^{-2}
Emitted longwave radiation	$L\uparrow$	W m^{-2}
Direct beam visible albedo	$I\uparrow_{vis}^{\mu}$	-
Direct beam near-infrared albedo	$I\uparrow_{nir}^{\mu}$	-
Diffuse visible albedo	$I\uparrow_{vis}$	-
Diffuse near-infrared albedo	$I\uparrow_{nir}$	-
Absorbed solar radiation	$\vec{S}$	W m^{-2}
Radiative temperature	T_{rad}	K
Temperature at 2 meter height	T_{2m}	K
Specific humidity at 2 meter height	q_{2m}	kg kg^{-1}
Snow water equivalent	W_{sno}	m
Aerodynamic resistance	r_{am}	s m^{-1}
Friction velocity	u_*	m s^{-1}
[2]Dust flux	F_j	kg m^{-2} s^{-1}
Net ecosystem exchange	NEE	$kgCO_2$ m^{-2} s^{-1}



The role of CLM in CESM: Atmosphere to Land

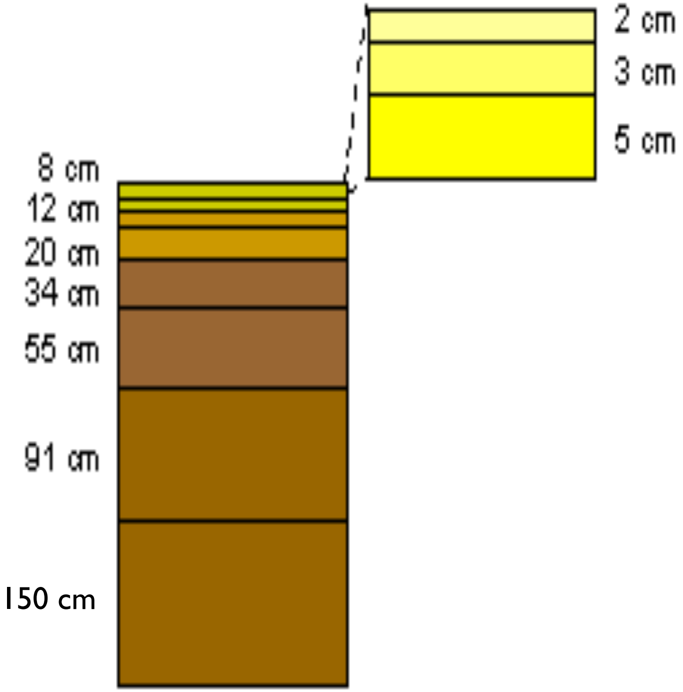
[1]Reference height	z'_{atm}	m
Zonal wind at z_{atm}	u_{atm}	m s^{-1}
Meridional wind at z_{atm}	v_{atm}	m s^{-1}
Potential temperature	$\overline{\theta_{atm}}$	K
Specific humidity at z_{atm}	q_{atm}	kg kg^{-1}
Pressure at z_{atm}	P_{atm}	Pa
Temperature at z_{atm}	T_{atm}	K
Incident longwave radiation	$L_{atm}\downarrow$	W m^{-2}
[2]Liquid precipitation	q_{rain}	mm s^{-1}
[2]Solid precipitation	q_{sno}	mm s^{-1}
Incident direct beam visible solar radiation	$S_{atm}\downarrow^{\mu}_{vis}$	W m^{-2}
Incident direct beam near-infrared solar radiation	$S_{atm}\downarrow^{\mu}_{nir}$	W m^{-2}
Incident diffuse visible solar radiation	$S_{atm}\downarrow_{vis}$	W m^{-2}
Incident diffuse near-infrared solar radiation	$S_{atm}\downarrow_{nir}$	W m^{-2}
Carbon dioxide (CO_2) concentration	c_a	ppmv
[3]Aerosol deposition rate	D_{sp}	kg m^{-2} s^{-1}
[4]Nitrogen deposition rate	NF_{ndep_sminn}	g (N) m^{-2} yr^{-1}
[5]Lightning frequency	I_l	flash km^{2} hr^{-1}



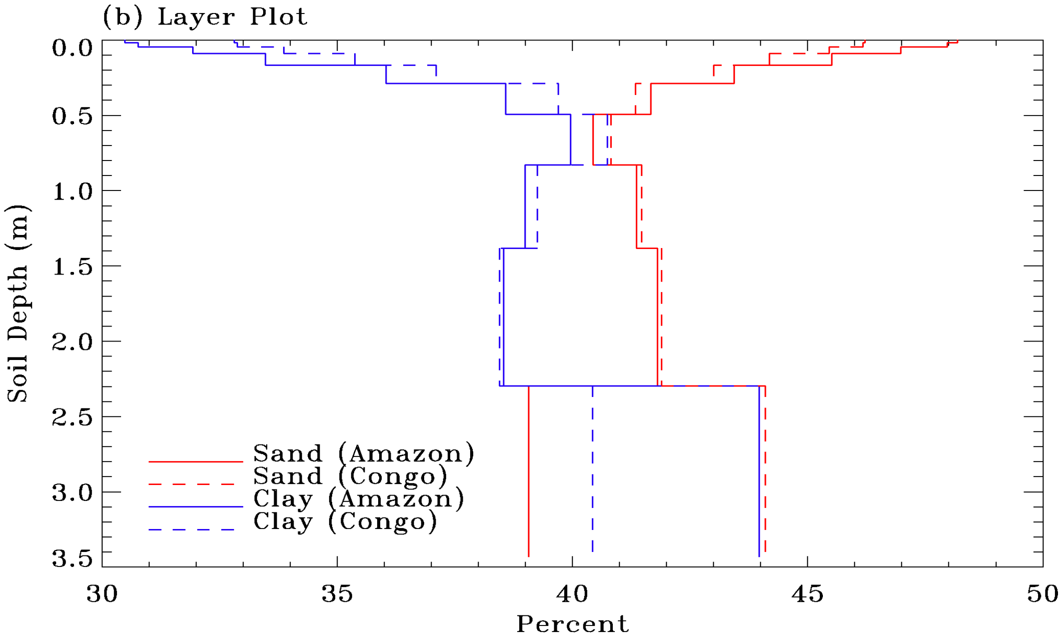
Soil Texture – thermal/hydrologic parameters

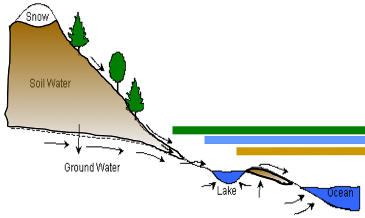
Soil parameters are derived from sand / clay percentage and soil organic matter content which is specified geographically and by soil level

- Soil moisture concentration at saturation
- Soil moisture concentration at wilting point
- Hydraulic conductivity at saturation
- Saturated soil suction
- Thermal conductivity
- Thermal capacity

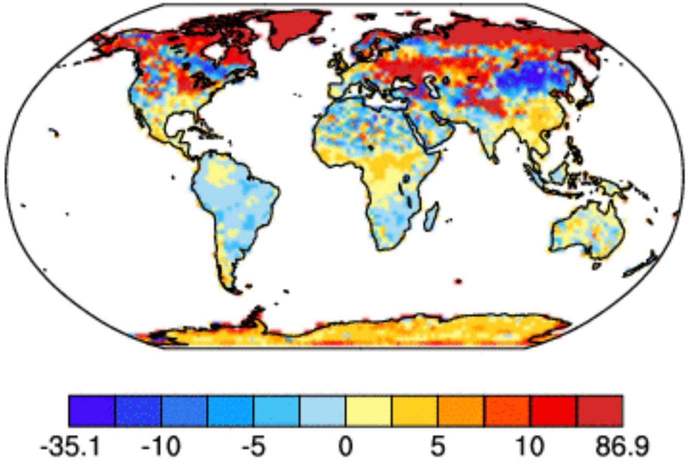


Soil profile
10 soil levels (~3.8m)
5 bedrock levels (~42m)

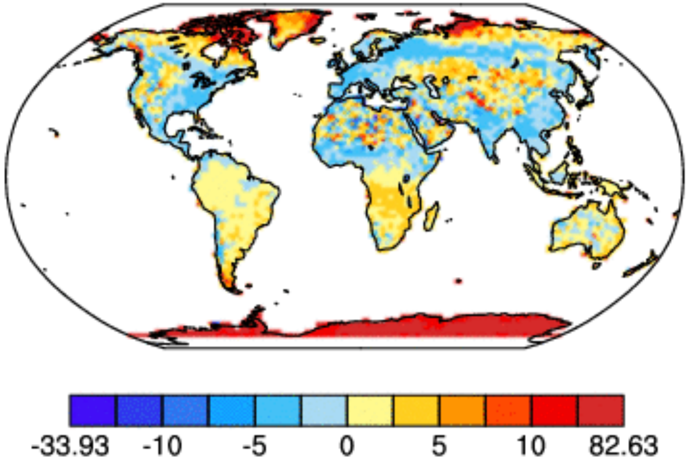




Modeling surface albedo


DJF ASA (% reflected)

CLM45SP_CRUNCEP
- Observations


JJA ASA (% reflected)

CLM45SP_CRUNCEP
- Observations

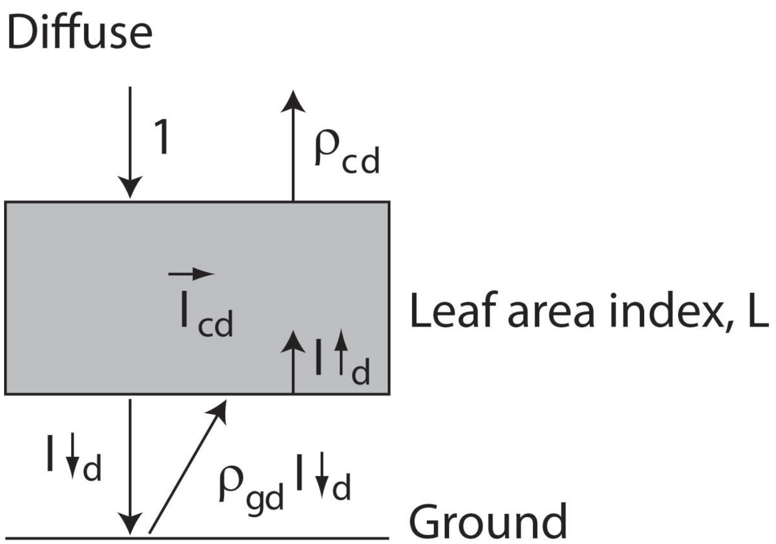
Surface albedo a function of

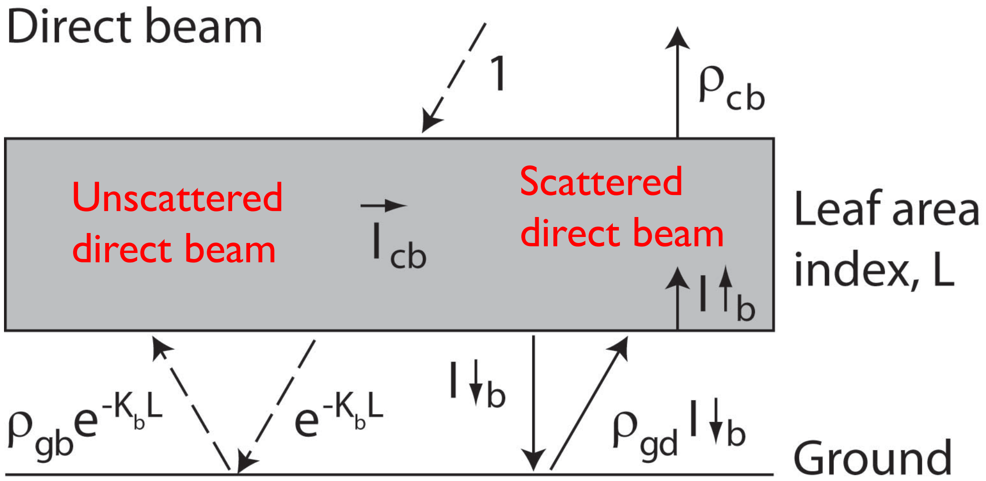
- Vegetation cover and type
- Snow cover
- Snow age
- Soil moisture
- Soil color
- Solar zenith angle
- Amount of direct vs diffuse solar radiation
- Amount of visible vs IR solar radiation

Note: MODIS albedo biased low for snow at high zenith angle

(Wang and Zender, 2010)

Two-stream radiative transfer





Radiative transfer uses the two-stream approximation (Dickinson, Sellers) to determine reflected and absorbed solar radiation

Slide courtesy G. Bonan

Momentum, and sensible heat and evaporation fluxes

Momentum flux

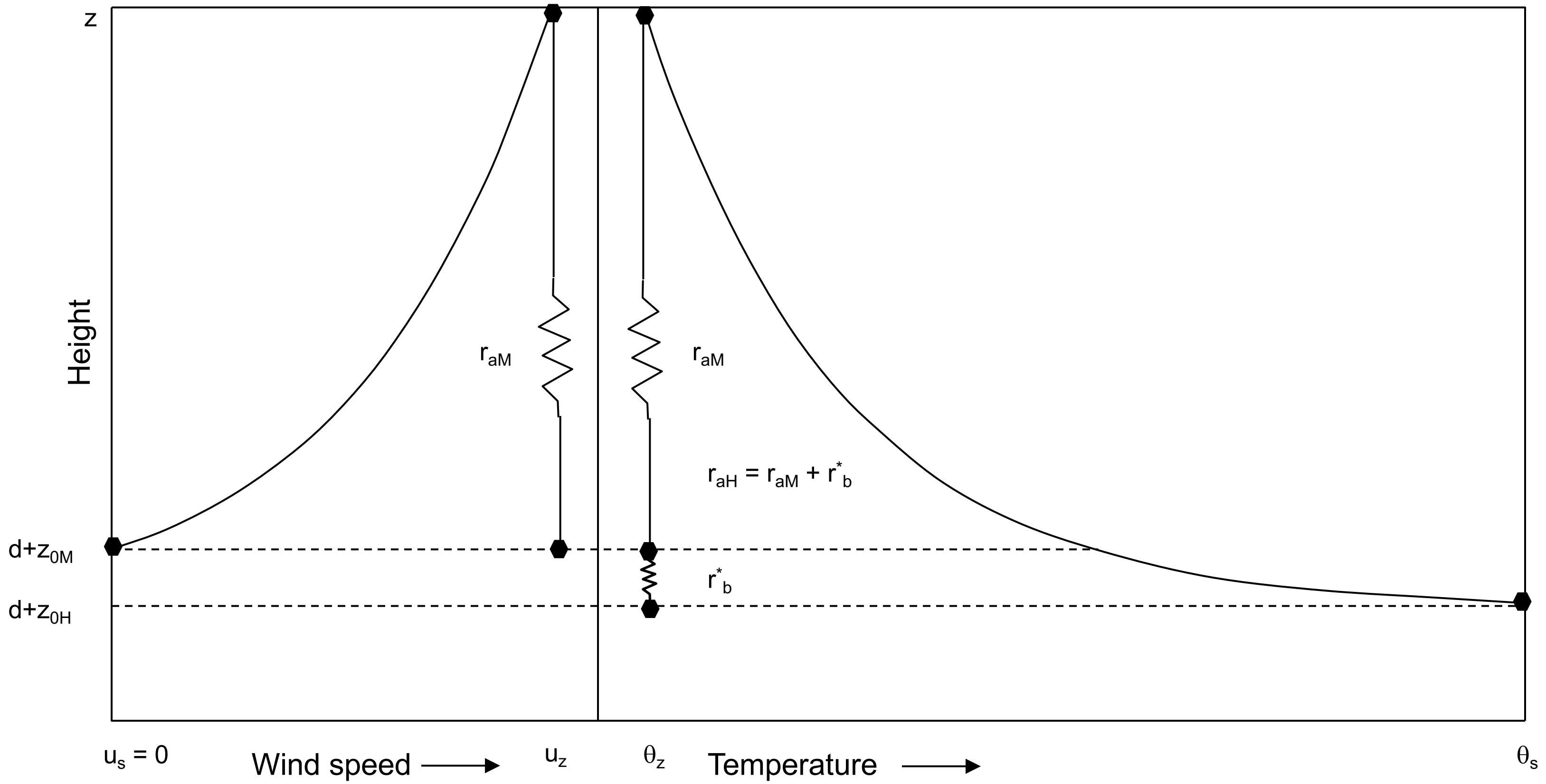
$u_* u_* = \tau / \rho$ and $\tau = \rho(u_a - u_s)/r_{aM} = \rho u / r_{aM}$ → $r_{aM} = \frac{1}{k^2 u}\left[\ln\left(\frac{z-d}{z_{0M}}\right) - \psi_m(\zeta)\right]^2$

Sensible heat flux

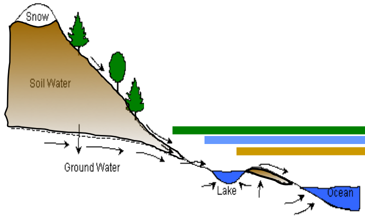
$\theta_* u_* = -H/(\rho c_p)$ and $H = -\rho c_p(\theta_a - T_s)/r_{aH}$ → $r_{aH} = \frac{1}{k^2 u}\left[\ln\left(\frac{z-d}{z_{0M}}\right) - \psi_m(\zeta)\right]\left[\ln\left(\frac{z-d}{z_{0H}}\right) - \psi_h(\zeta)\right]$

Evaporation

$q_* u_* = -E/\rho$ and $E = -\rho(q_a - q_s)/r_{aW}$ → $r_{aW} = \frac{1}{k^2 u}\left[\ln\left(\frac{z-d}{z_{0M}}\right) - \psi_m(\zeta)\right]\left[\ln\left(\frac{z-d}{z_{0W}}\right) - \psi_w(\zeta)\right]$



Slide courtesy G. Bonan



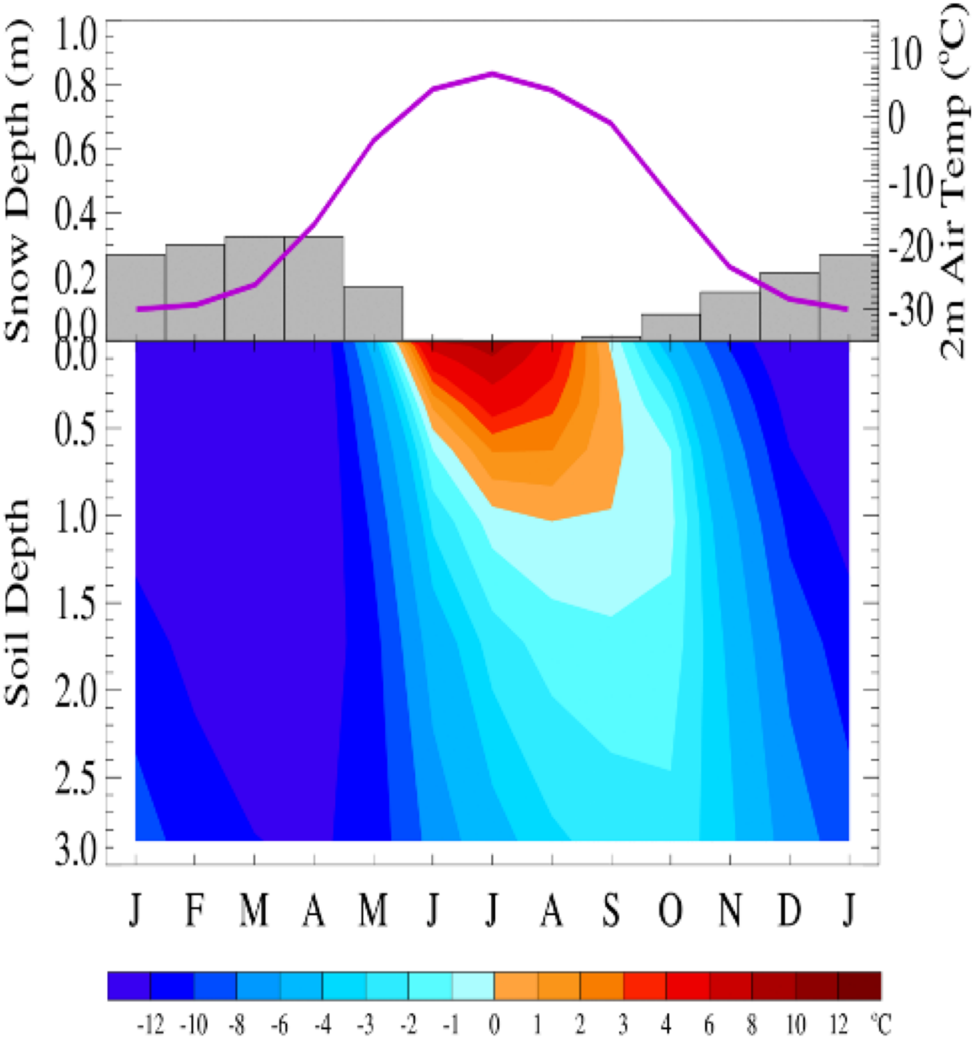
Snow/Soil thermodynamics

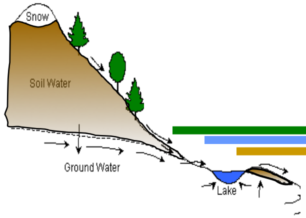
Solve the heat diffusion equation for multi-layer snow and soil model

$$C_p \frac{\partial T}{\partial t} = \frac{\partial}{\partial z}\left(K \frac{\partial T}{\partial z} \right)$$

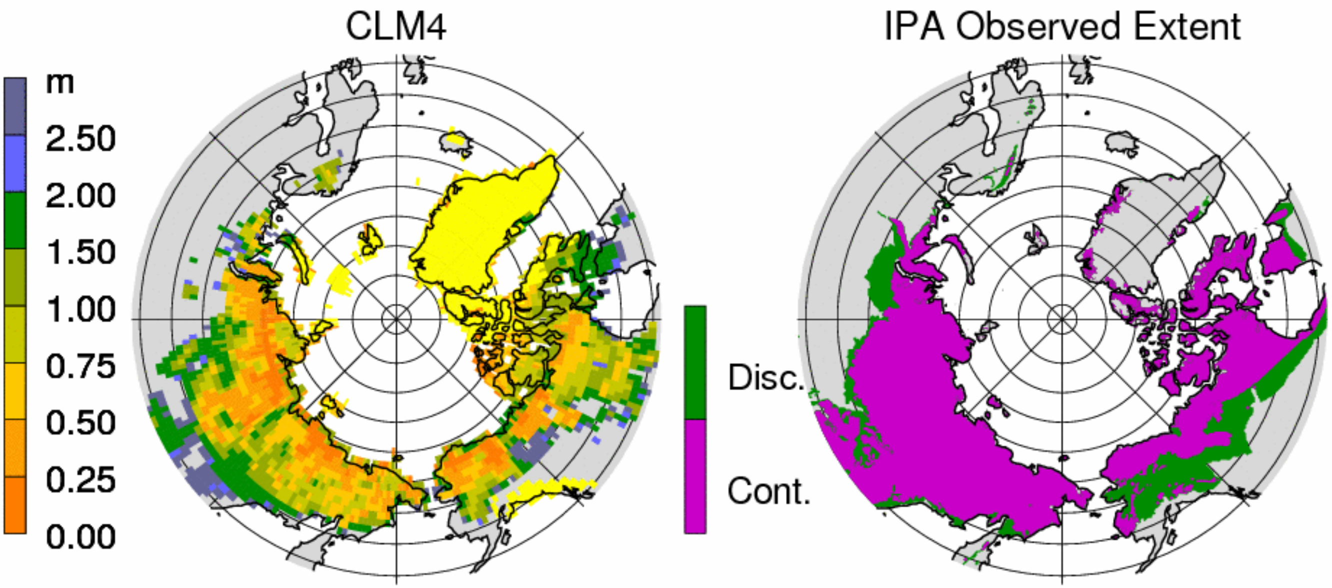
where C_p (heat capacity) and K (thermal conductivity) are functions of:

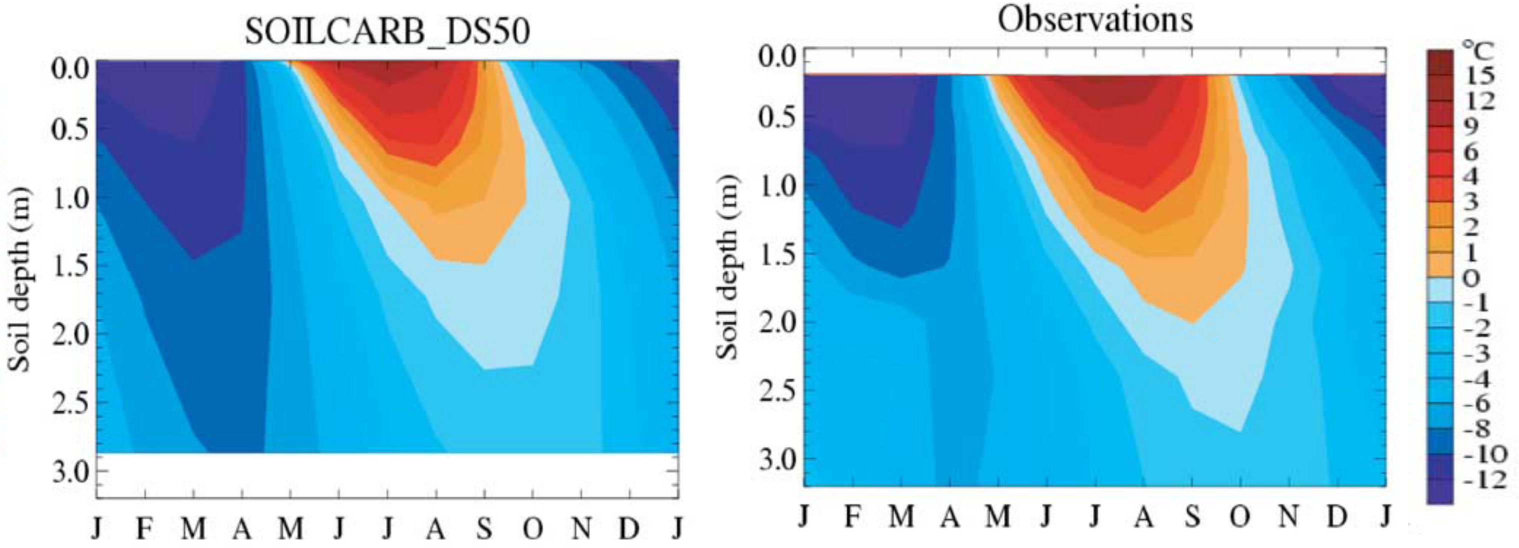
- temperature
- total soil moisture
- soil texture
- ice/liquid content





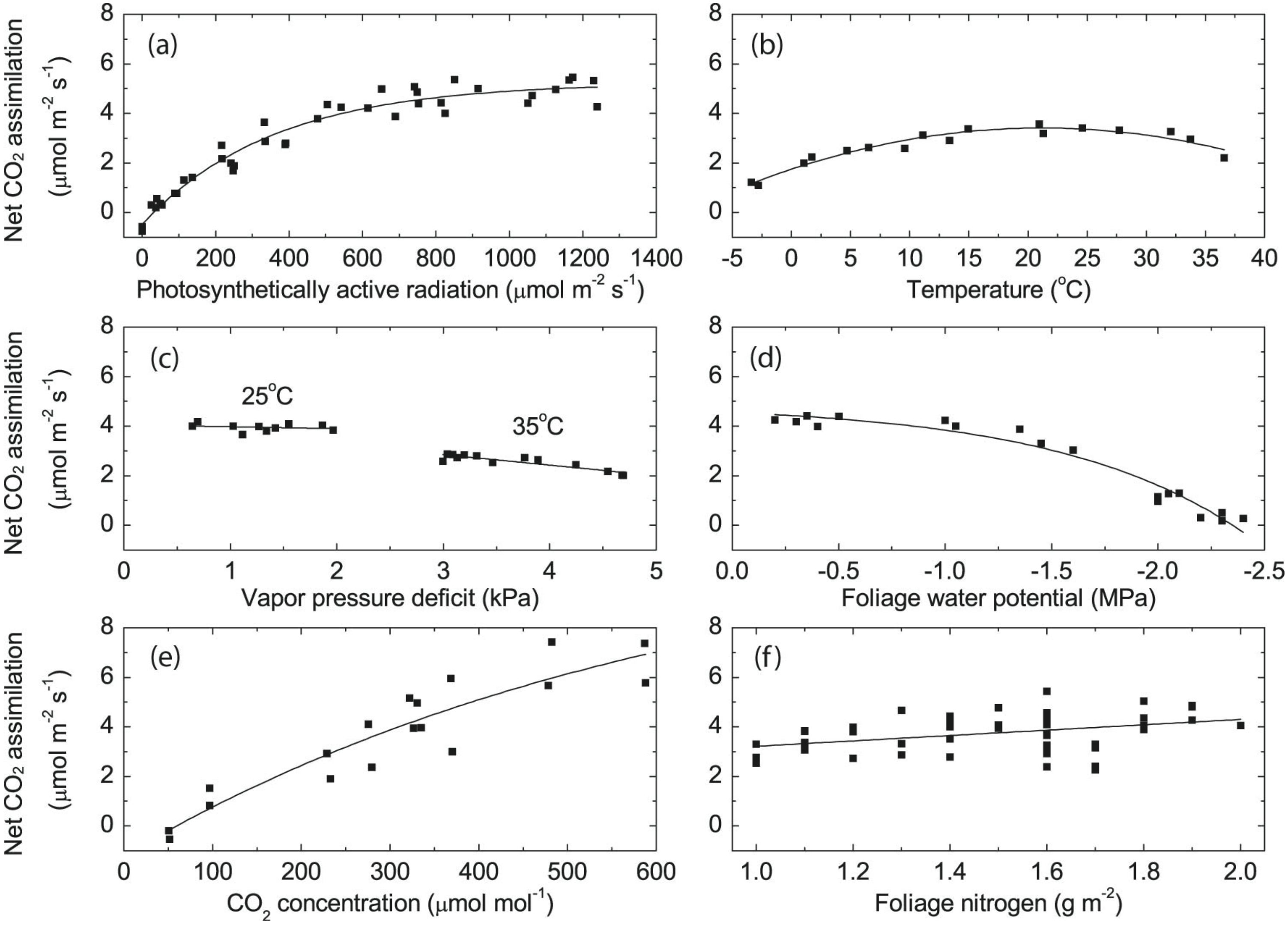
Modeling Permafrost in CLM





Lawrence et al., J. Climate, 2011

Leaf photosynthesis



Slide courtesy G. Bonan

Leaf photosynthesis and stomatal conductance

Farquhar photosynthesis model

$$A_n = \min(w_c, w_j, w_p) - R_d$$

w_c is the rubisco-limited rate of photosynthesis, w_j is light-limited rate allowed by RuBP regeneration, w_p is product limited rate of carboxylation

rubisco-limited rate is

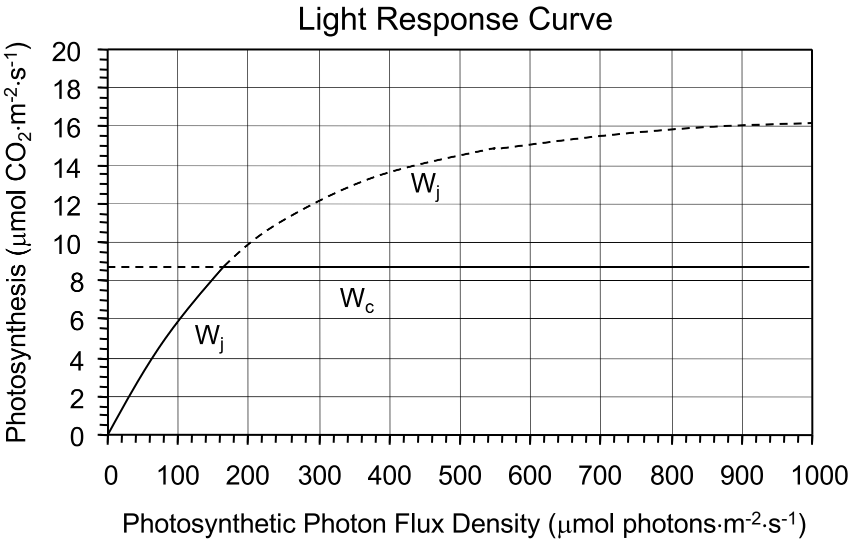
$$w_c = \frac{V_{c\max}(c_i - \Gamma_*)}{c_i + K_c(1 + O_i/K_o)}$$

RuBP regeneration-limited rate is

$$w_j = \frac{J(c_i - \Gamma_*)}{4(c_i + 2\Gamma_*)}$$

product-limited rate is

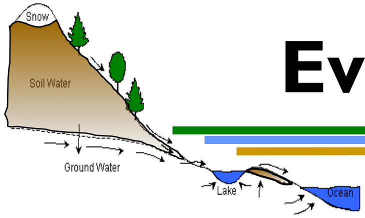
$$w_p = 3T_p$$



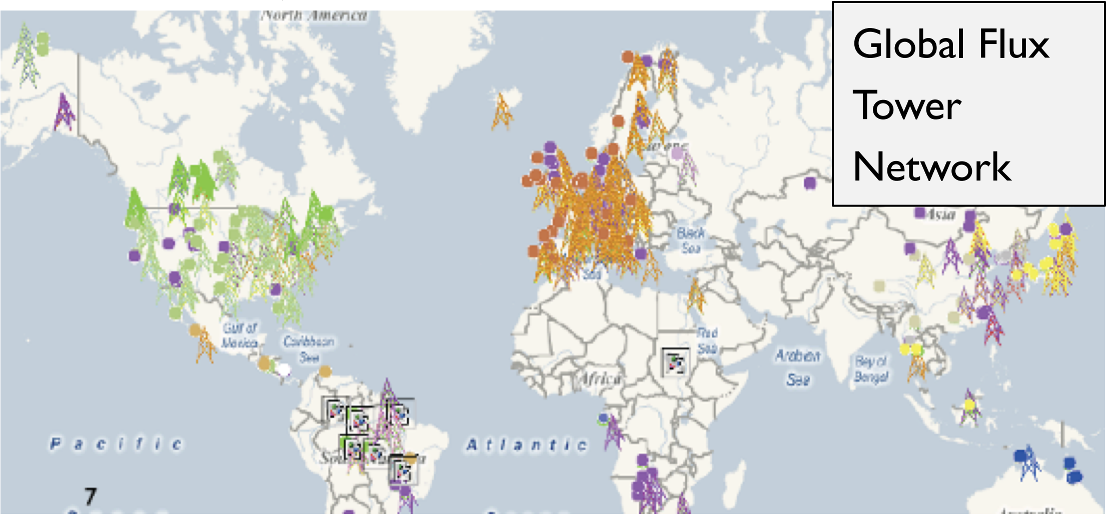
Ball-Berry stomatal conductance

$$\frac{1}{r_s} = g_s = g_1 \frac{A_n h_s}{c_s/P_{atm}} + g_0 \beta_t$$

Slide courtesy G. Bonan



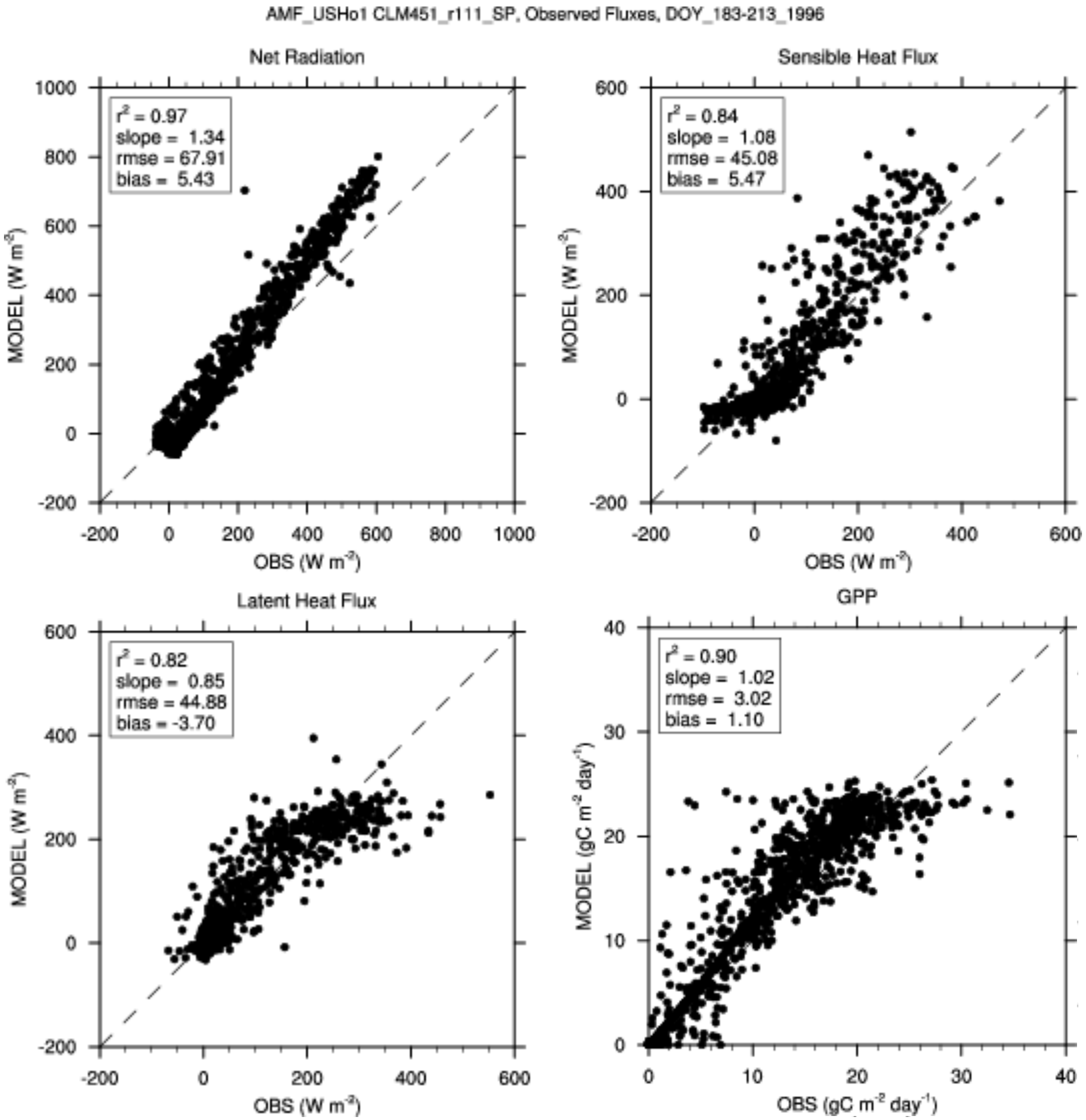
Evaluating the model with tower flux data



Global Flux Tower Network

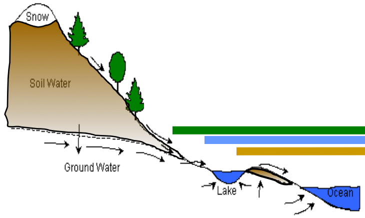
Evaluating CLM4.5 with tower flux data

Howland Forest, Maine, July, 1996

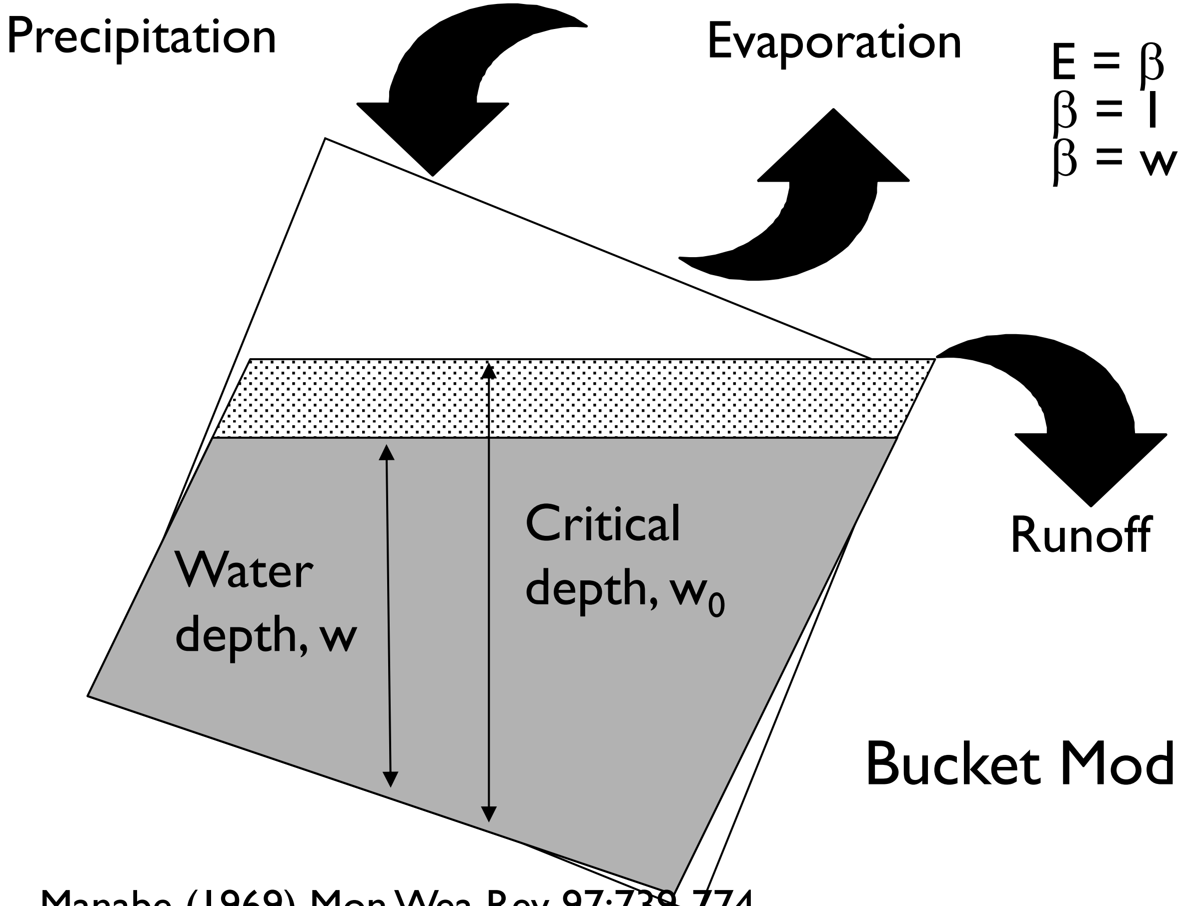


Manhattan-Mannahatta: on right is a reconstruction of Manhattan Island circa 1609 (called "Mannahatta" by the Lenape native Americans), as compared to today, based on historical landscape ecology and map data.

Markley Boyer / The Mannahatta Project / Wildlife Conservation Society and the aerial view of modern Manhattan, Amiaga Photographers. In: Climate Change and Cities: First Assessment Report of the Urban Climate Change Research Network (ARC3) (2011)



Land models have come a long way
1st Generation: Bucket Model

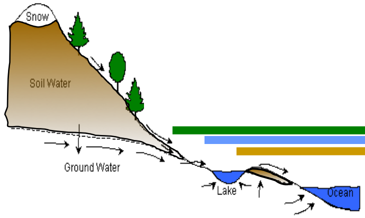


$E = \beta E_P$
$\beta = 1 \quad \text{for } w \geq w_0$
$\beta = w/w_0 \text{ for } w < w_0$

Bucket Model

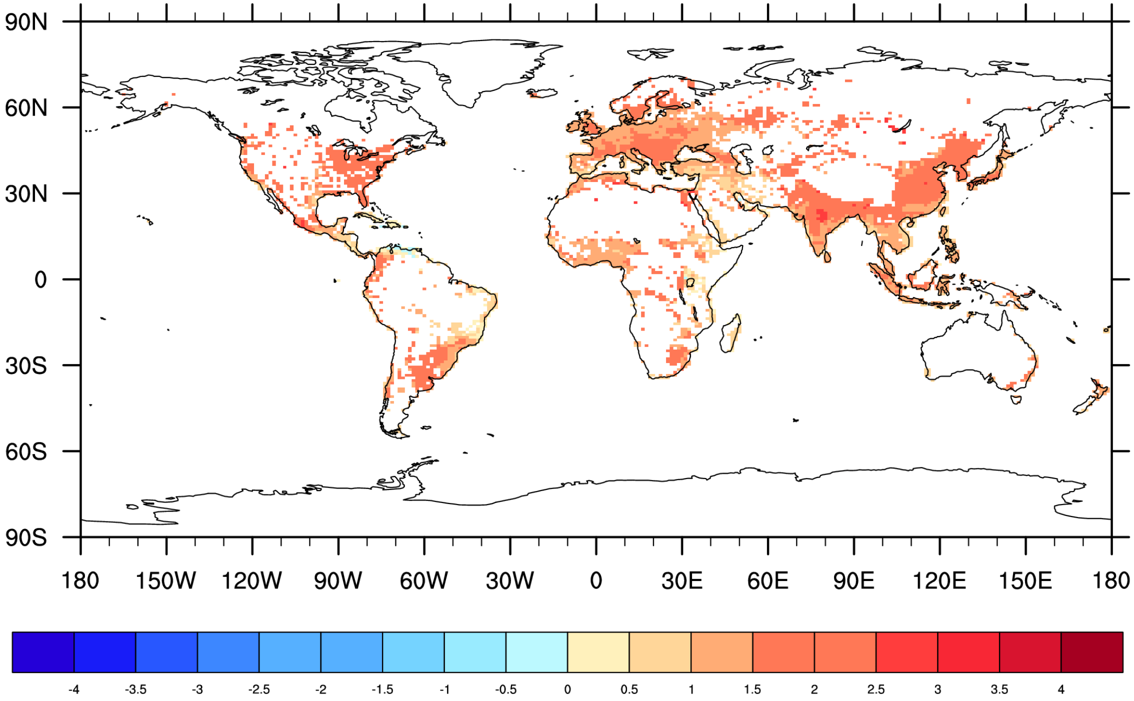
Manabe (1969) Mon Wea Rev 97:739-774
Williamson et al. (1987) NCAR/TN-285+STR

Figure courtesy G. Bonan

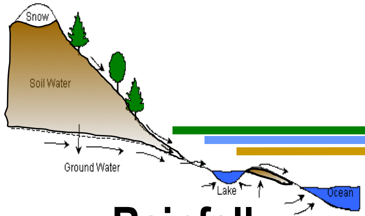


Urban Heat Island in CCSM4

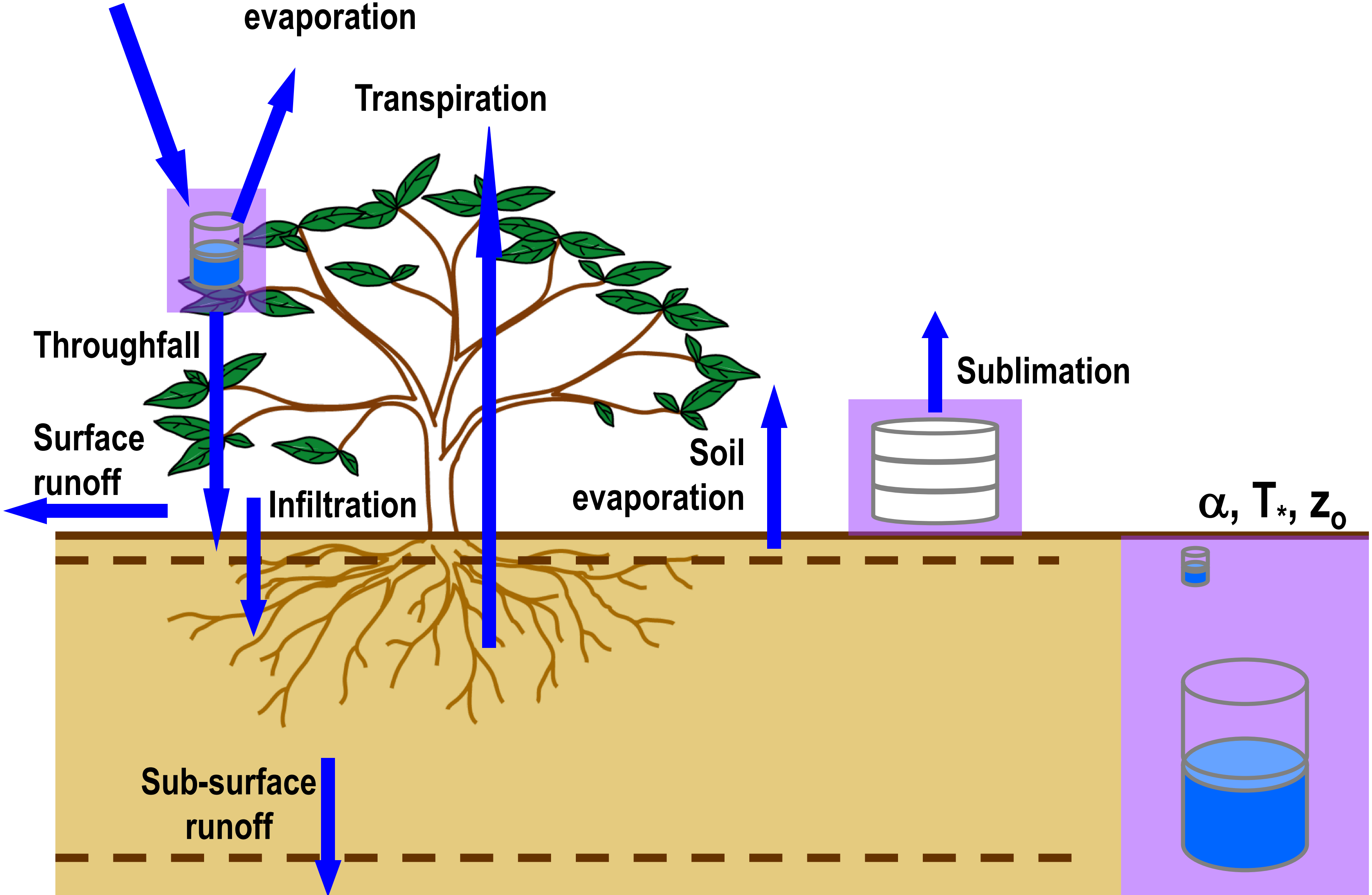
Present day Urban Heat Island (UHI) simulated by CLM Urban (°C)



Modeled UHI ranges from near-zero up to 4°C with spatial and seasonal variability controlled by urban to rural contrasts in energy balance.



Land model processes - Hydrology



Rainfall

Canopy evaporation

Transpiration

Throughfall

Surface runoff

Infiltration

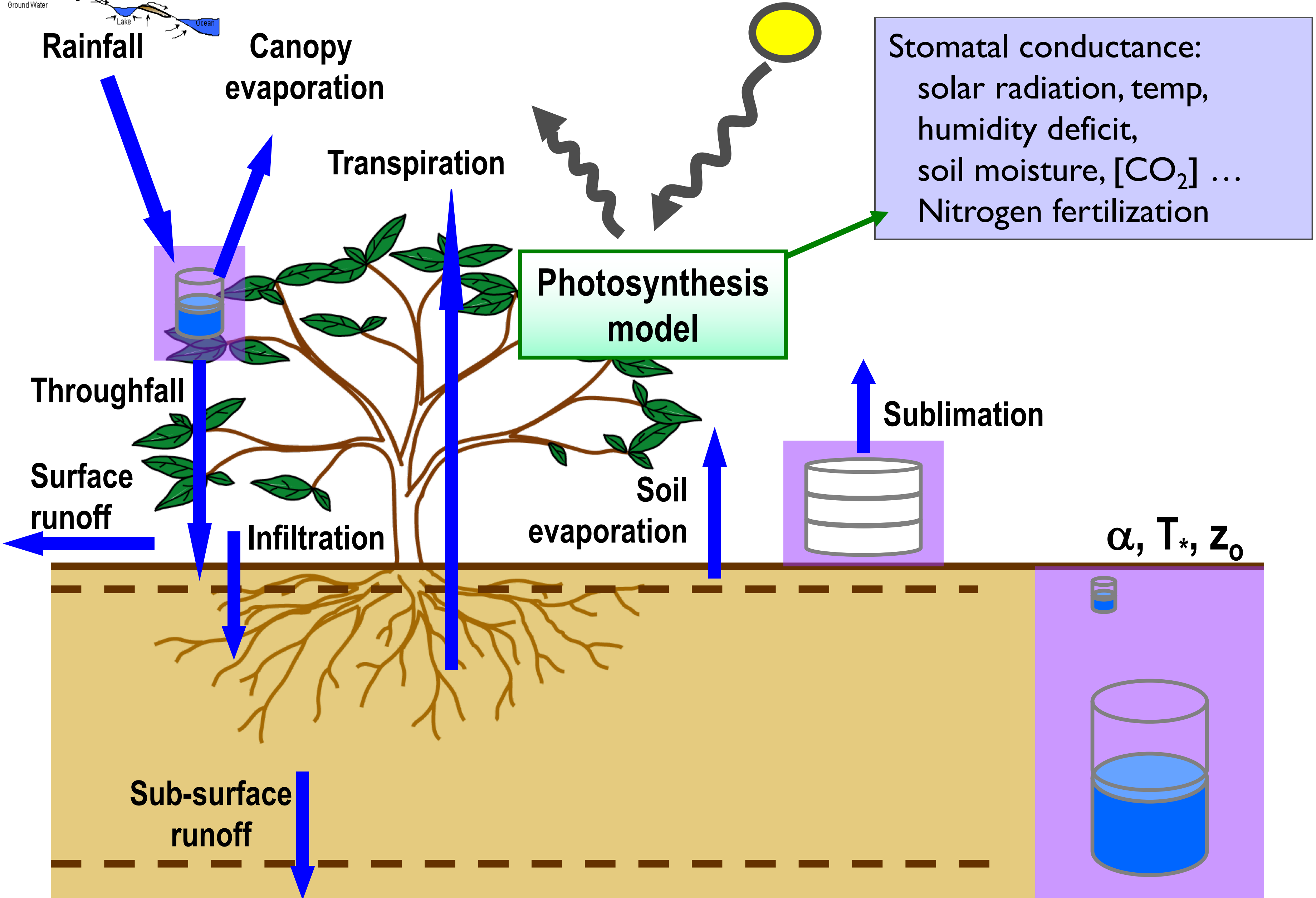
Soil evaporation

Sublimation

α, T_*, z_o

Sub-surface runoff

Land model processes - Hydrology



Rainfall

Canopy evaporation

Transpiration

Stomatal conductance:
- solar radiation, temp,
- humidity deficit,
- soil moisture, [CO_2] …
- Nitrogen fertilization

Photosynthesis model

Throughfall

Surface runoff

Infiltration

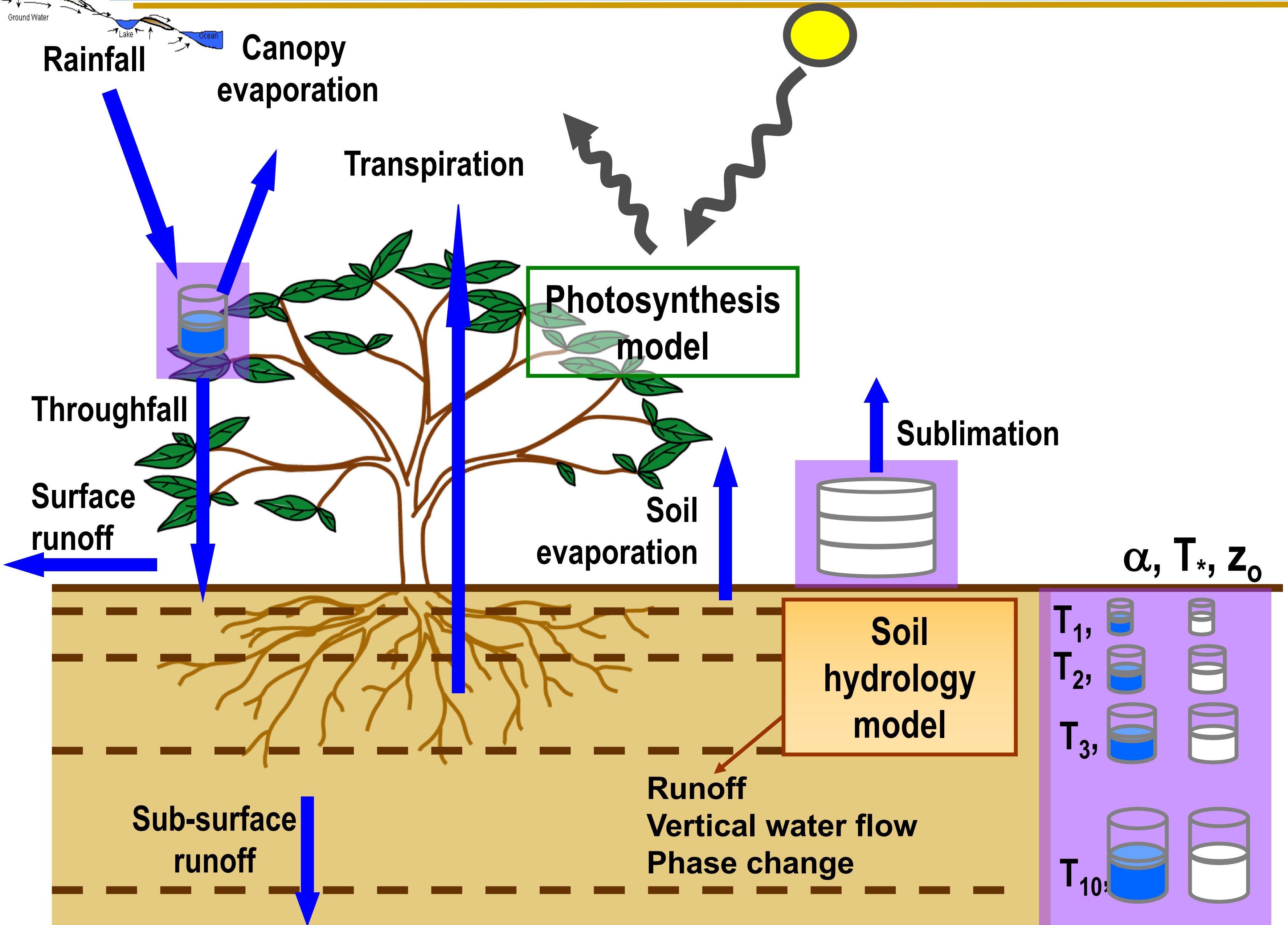
Soil evaporation

Sublimation

α, T_*, z_o

Sub-surface runoff

Land model processes - Hydrology



Rainfall

Canopy evaporation

Transpiration

Photosynthesis model

Throughfall

Sublimation

Surface runoff

Soil evaporation

α, T_*, z_o

Soil hydrology model

T_1,

T_2,

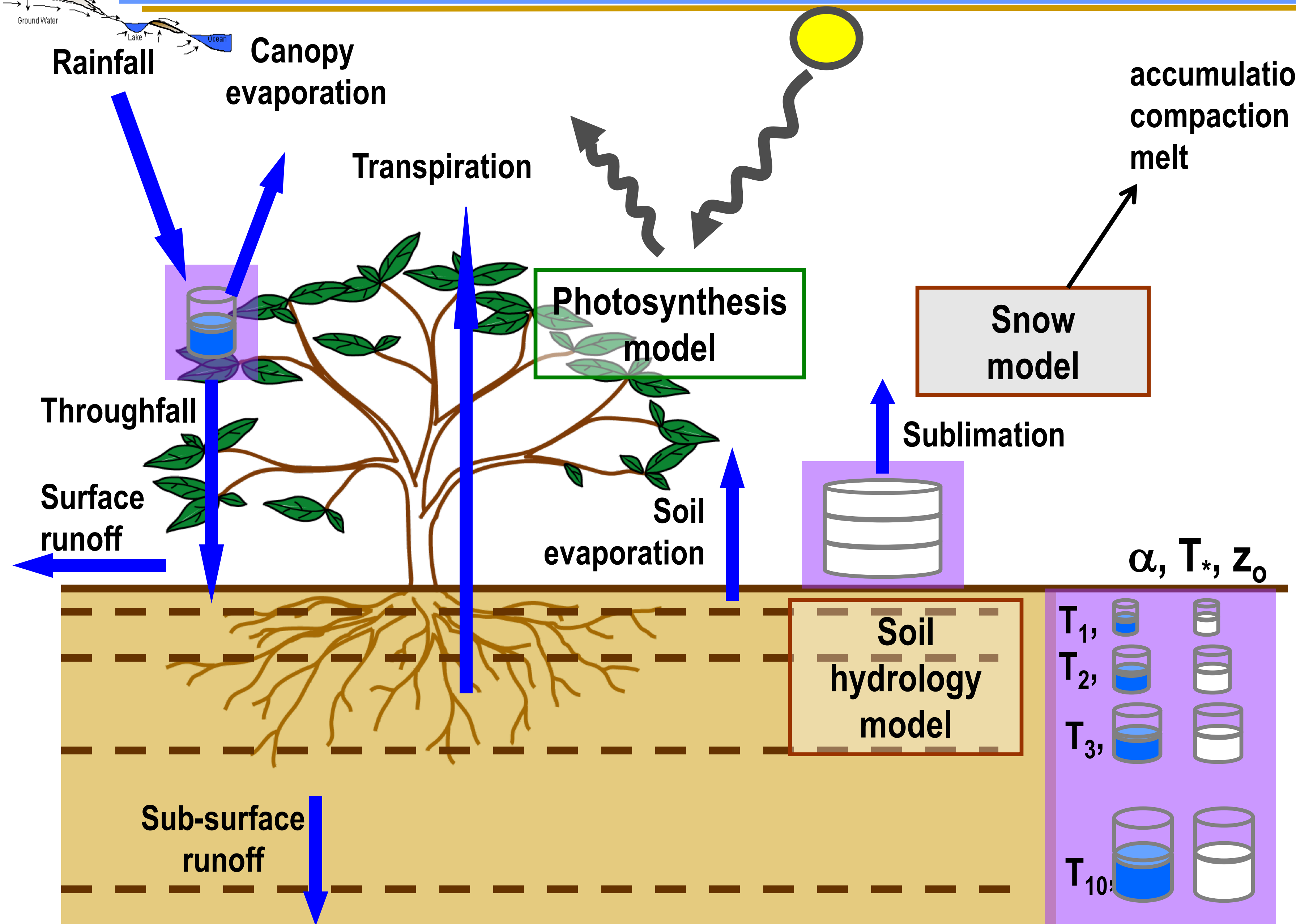
T_3,

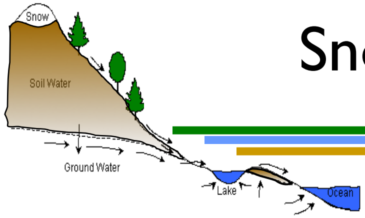
T_{10};

Runoff
Vertical water flow
Phase change

Sub-surface runoff

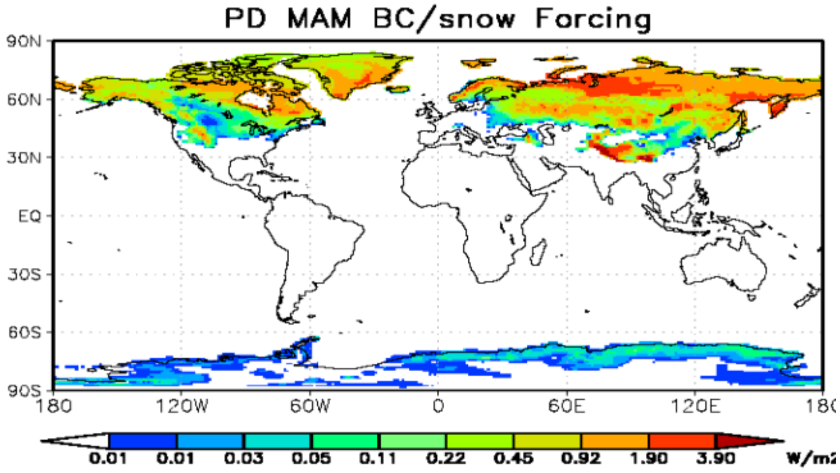
Land model processes - Hydrology

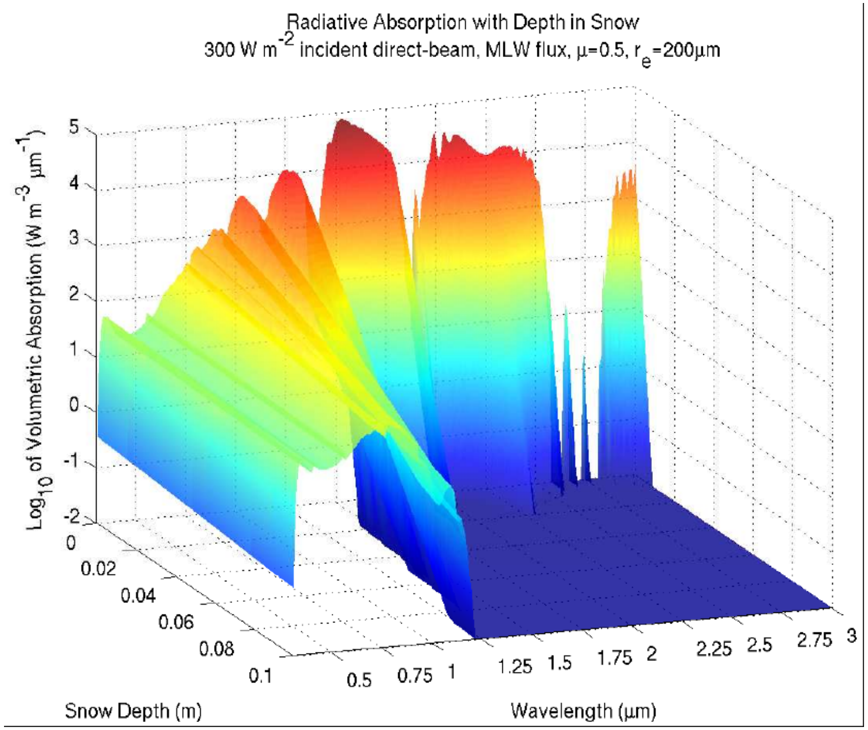




Snow, Ice, and Aerosol Radiative Model (SNICAR)

- Snow darkening from deposited black carbon, mineral dust, and organic matter
- Vertically-resolved solar heating in the snowpack
- Snow aging (evolution of effective grain size) based on:
  - Snow temperature and temperature gradient
  - Snow density
  - Liquid water content and
  - Melt/freeze cycling


PD MAM BC/snow Forcing


Radiative Absorption with Depth in Snow

$300\ W\ m^{-2}$ incident direct-beam, MLW flux, $\mu=0.5$, $r_e=200\mu m$

Flanner et al (2007), *JGR*
Flanner and Zender (2006), *JGR*
Flanner and Zender (2005), *GRL*