



Creation of an Eddy Parameterization Challenge Suite

Scott Bachman

**with Frank Bryan, John Dennis, and Baylor
Fox-Kemper**

NSF OCE 0825614

Outline

- What is the motivation behind the challenge suite?
- What will it be used for and what is the justification for creating it?
- What the challenge suite *REALLY* is
- Idealized examples
- Future Work

Tracer Flux-Gradient Relationship

$$\overline{\mathbf{u}'\tau'} = -\mathbf{M}\nabla\bar{\tau}$$

- Virtually all subgridscale eddy closures may be written as: GM, Redi, FFH Submesoscale
- Relates the eddy flux to the coarse-grain gradients
- May have a flow/property dependent \mathbf{M}

$$\overline{\mathbf{u}'\tau'} = -\mathbf{M}\nabla\bar{\tau}$$

Fox-Kemper, Ferrari, & Hallberg (2008) form
(a mixed layer (submeso) eddy param.):

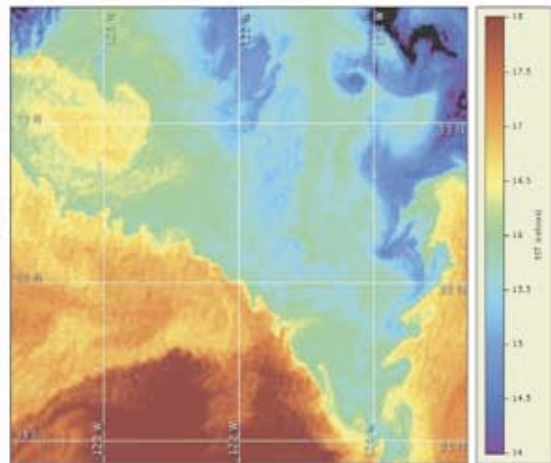
$$\begin{bmatrix} \overline{u'\tau'} \\ \overline{v'\tau'} \\ \overline{w'\tau'} \end{bmatrix} = - \begin{bmatrix} 0 & 0 & -\Psi_y \\ 0 & 0 & \Psi_x \\ \Psi_y & -\Psi_x & 0 \end{bmatrix} \begin{bmatrix} \bar{\tau}_x \\ \bar{\tau}_y \\ \bar{\tau}_z \end{bmatrix}$$

Antisymmetric Elements in Fox-Kemper, Ferrari, & Hallberg (2008) are scaled to **overtake fronts**, making vertical fluxes **extract PE**, and **restratify the fluid**,

At a rate validated against eddy simulations!

The Character of the Submesoscale

(Capet et al., 2008)



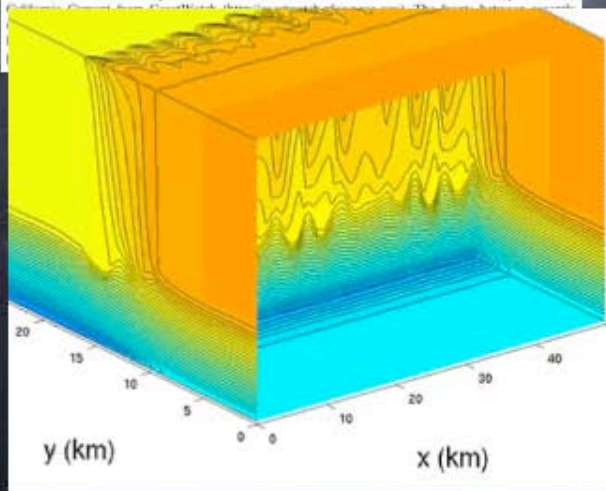
Longitude

FIG. 16. Sea surface temperature measured at 1832 UTC 3 Jun 2006 off Point Conception in the

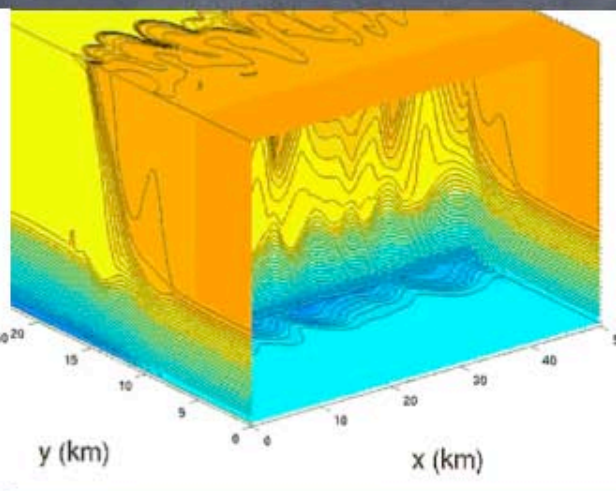
- Fronts
- Eddies
- $Ro=O(1)$
- $Ri=O(1)$
- near-surface



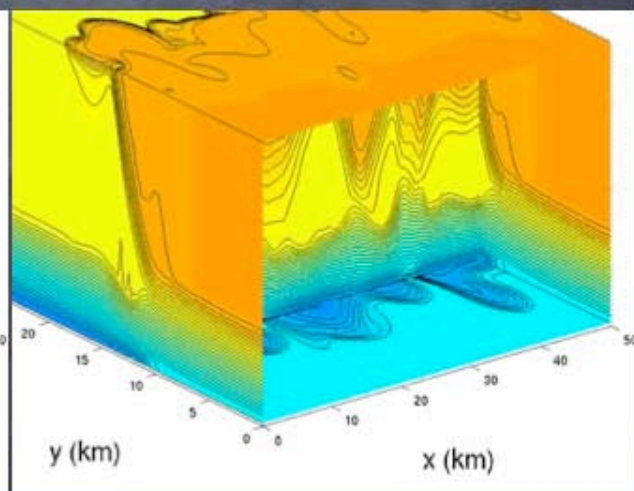
100 km



Simple Spindown



Simple Spindown



Simple Spindown

The Problem is:

The mesoscale equivalent isn't rEady

- ① FFH param. doesn't do interior stratification/
PV gradients
- ① PV jumps are OK, e.g, surface & mixed layer
base
- ① But, Mesoscale==Full Depth, so PV Varies
- ① Smith (07) shows interior PV gradients
dominate mesoscale energy extraction

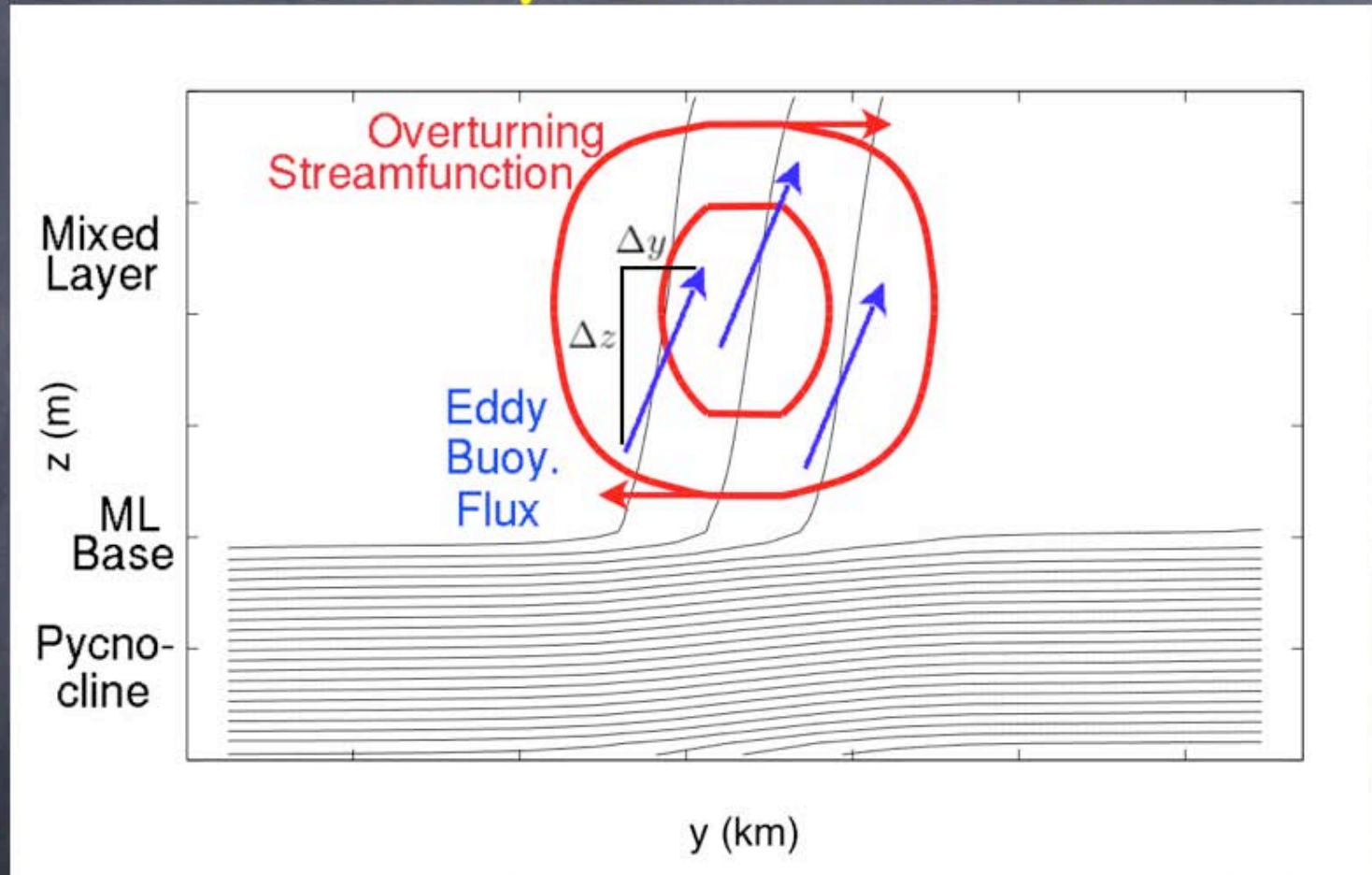
What to do?

The Parameterization Challenge Suite

- Uses a model (MITgcm) that includes shear, strain, stratification, etc.
- Will simulate individually: $O(2000)$ simulations
- Simultaneous runs exploring parameter space
- A suite of several different challenges, using both idealized and reality-based problems
- Will extract “typical” eddy configurations in the global model by EOF or SOM

Examples?

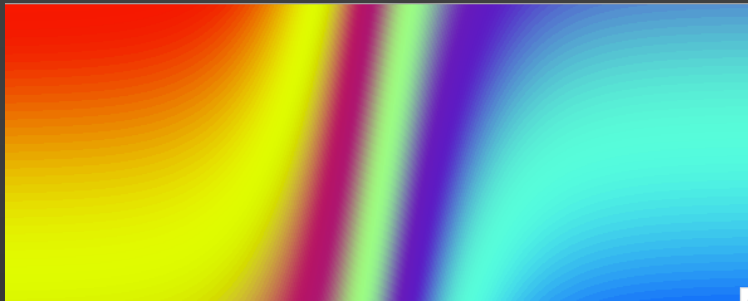
Overturning Schematic: An Eady-like Problem



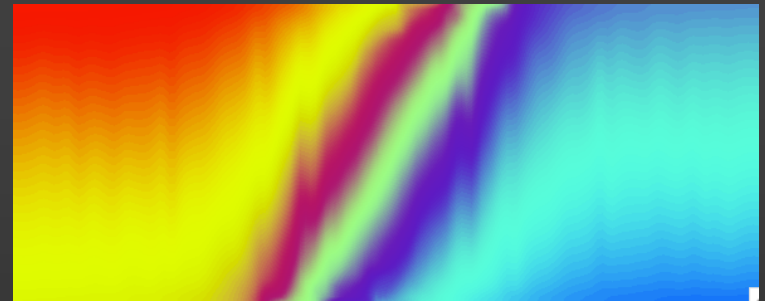
Horizontal scale of overturning = scale of front
Vertical structure of overturning = ?

Eady Snapshots

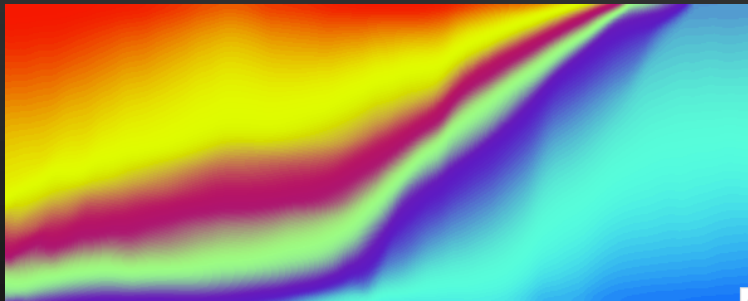
Density surfaces



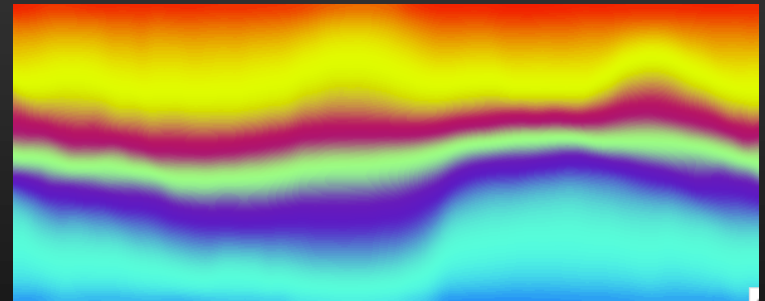
Initial configuration



10 days



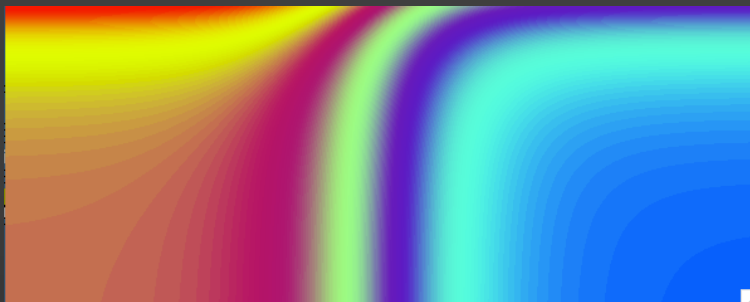
20 days



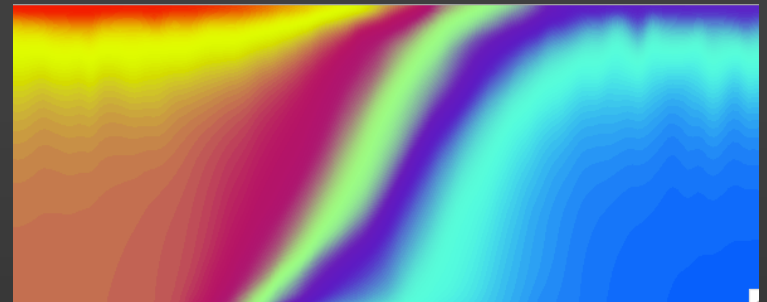
30 days

Charney Snapshots

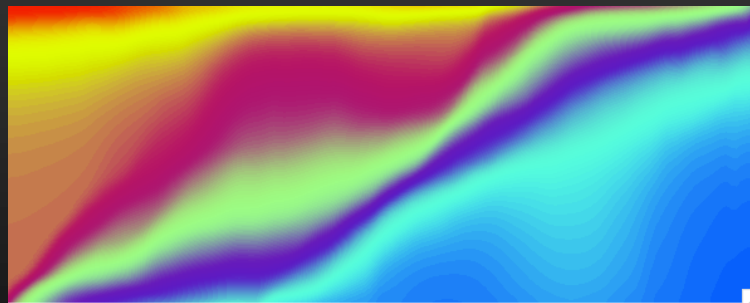
Density surfaces



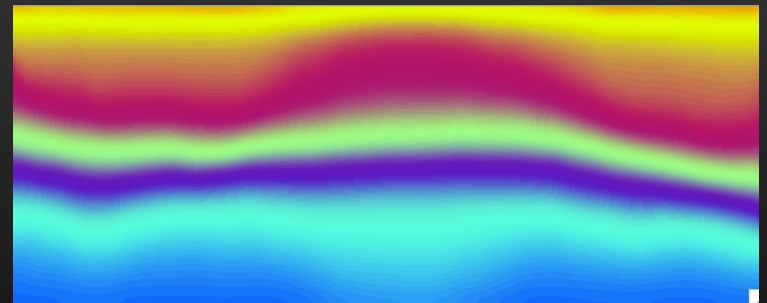
Initial configuration



10 days



20 days



30 days

Evaluating the Results

- How fast is the slumping?
- Multiple tracer approach to diagnose M (Bratseth, 1998)

The Future!

- Investigate smoothing operator in FVG

$$\left(c^2 \frac{d^2}{dz^2} - N^2 \right) \tilde{\Psi} = -\kappa \nabla \bar{b}$$
$$\tilde{\Psi} = 0, \quad z = 0, -H$$

- Selection of more challenges
 - PV jumps
 - Sinusoids
- Suggestions from the audience?

The Character of the Mesoscale

(Capet et al., 2008)

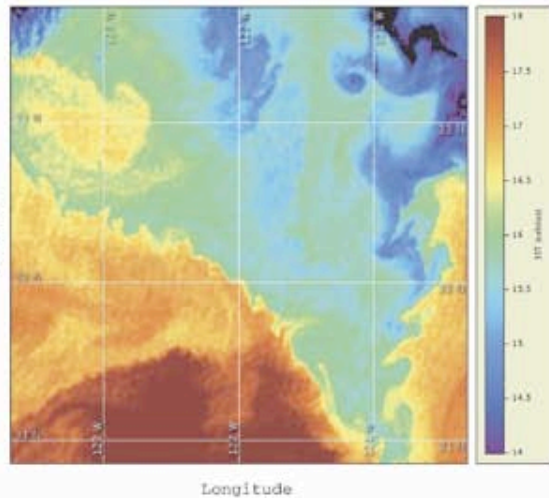


FIG. 16. Sea surface temperature measured at 1800 UTC 3 Jun 2006 off Point Conception in the California Current from CoastWatch (<http://coastwatch.pfeg.noaa.gov>). The fronts between recently upwelled water (i.e., 15°–16°C) and offshore water ($\geq 17^\circ\text{C}$) show submesoscale instabilities with wavelengths around 30 km (right front) or 15 km (left front). Images for 1 day earlier and 4 days later show persistence of the instability events.

- Boundary Currents
- Eddies
- $Ro=O(0.1)$
- $Ri=O(1000)$
- Full Depth



But, dominant eddy process **still baroclinic instability.**

Why can't we just use submesoscale scaling of **FFH**? for a flow-dependent **GM-like param**?

MORAL OF THIS SLIDE:

Scaling Laws for streamfunction
can be deduced empirically from
the models

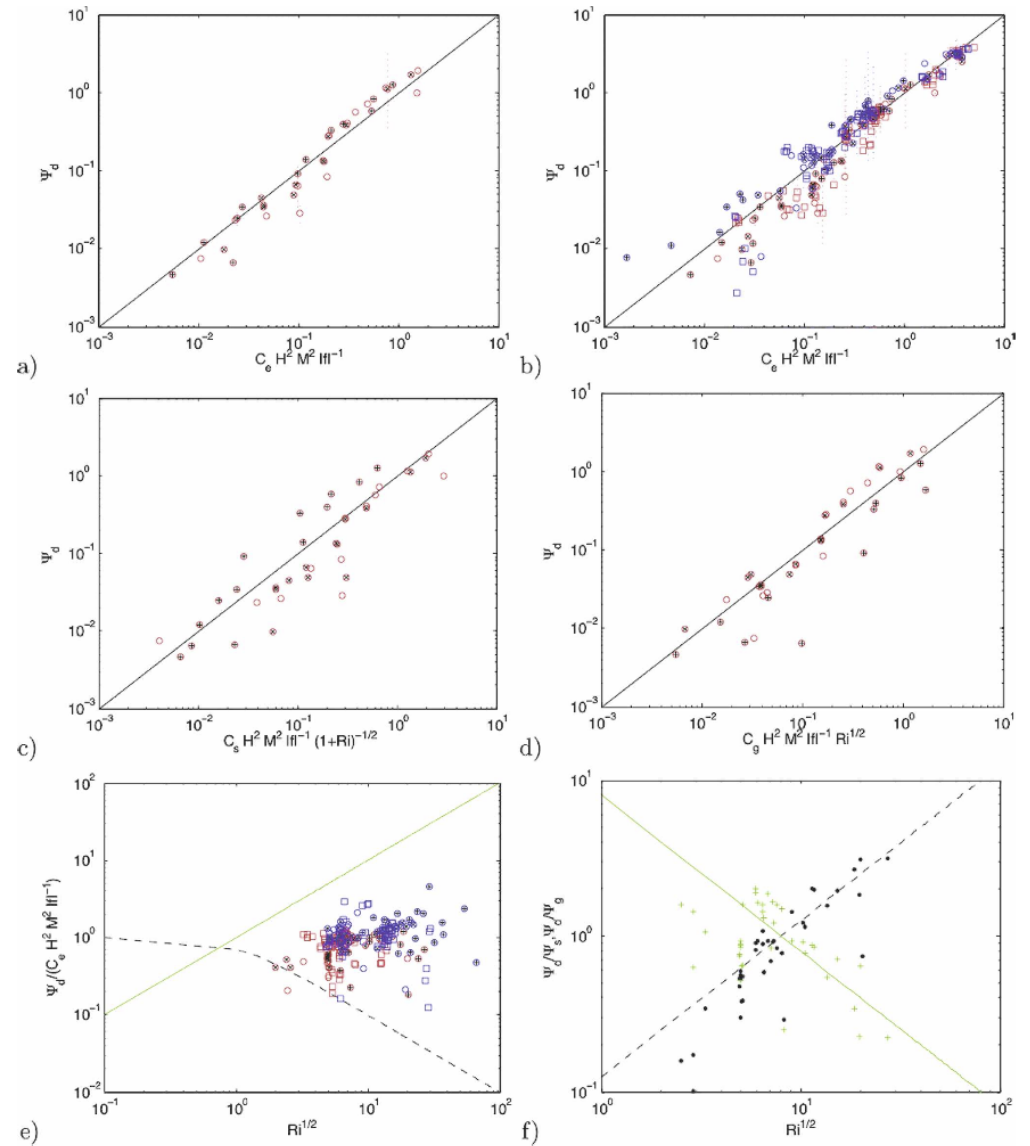


FIG. 14. Magnitude of Ψ_d vs theories for magnitude of Ψ for simulations with diurnal cycle (blue) and without (red) starting from balanced (circles) or unbalanced (squares) initial conditions. Plus signs and crosses indicate balanced simulations where $Ri_0 > 1$ or $Ri_0 < 1$ initially: (a) Ψ_d in the balanced, no diurnal cycle simulations vs $C_e b_y^{xz} H^2 |f|^{-1}$, $C_e = 0.06$, and (b) unbalanced and diurnal cycle simulations, $C_e = 0.08$; (c) Stone's theory, (36), $C_s = 0.53$; (d) Green's theory, (37), $C_g = 0.0085$; (e) $\Psi_d / C_e b_y^{xz} H^2 |f|^{-1}$ vs $Ri^{1/2}$. Also shown are lines parallel to $Ri^{1/2}$ and $(1 + Ri)^{-1/2}$; (f) Ψ_d / Ψ_g (black dots) and Ψ_d / Ψ_g (green crosses) vs $Ri^{1/2}$. Also shown are lines parallel to $Ri^{\pm 1/2}$; Ψ_d , Ψ_s , and Ψ_g are defined in (35), (36), and (37).