

# **Coupling between the lower and upper atmosphere as simulated by the Whole Atmosphere Model (WAM)**

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# **Integrated Dynamics in the Earth's Atmosphere (IDEA) project**

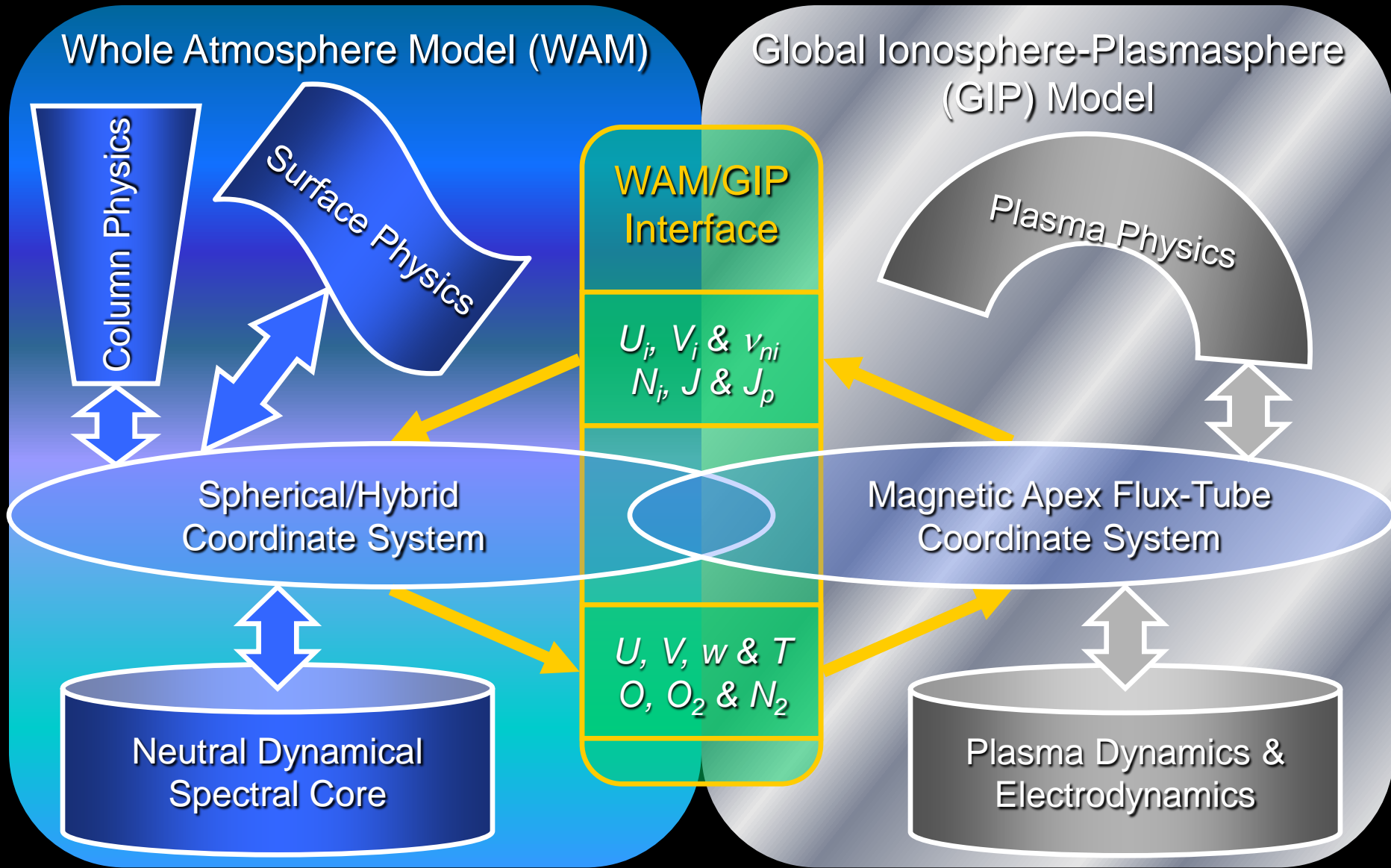
## **Collaborators**

- **NCEP Environmental Modeling Center (EMC)**
- **Naval Research Laboratory (NRL)**
- **National Center for Atmospheric Research (NCAR)**
- **Others**

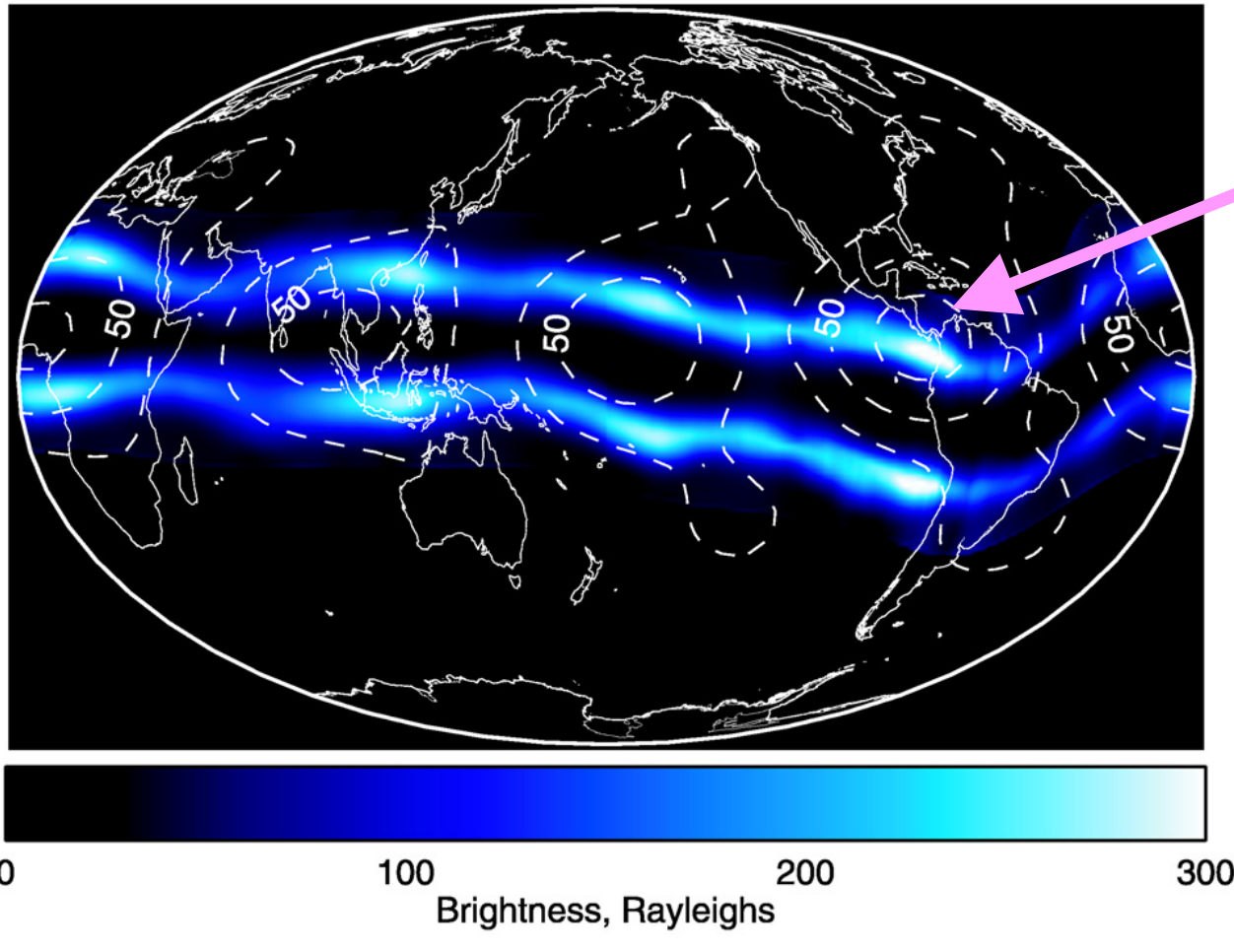
## **Sponsored by**

- **NASA LWS and Heliophysics Theory programs**
- **AFOSR Multidisciplinary University Research Initiative (MURI) program**

# Coupled IDEA model



# Motivation: Tidal signatures in nightside Equatorial Ionospheric Anomaly (EIA)



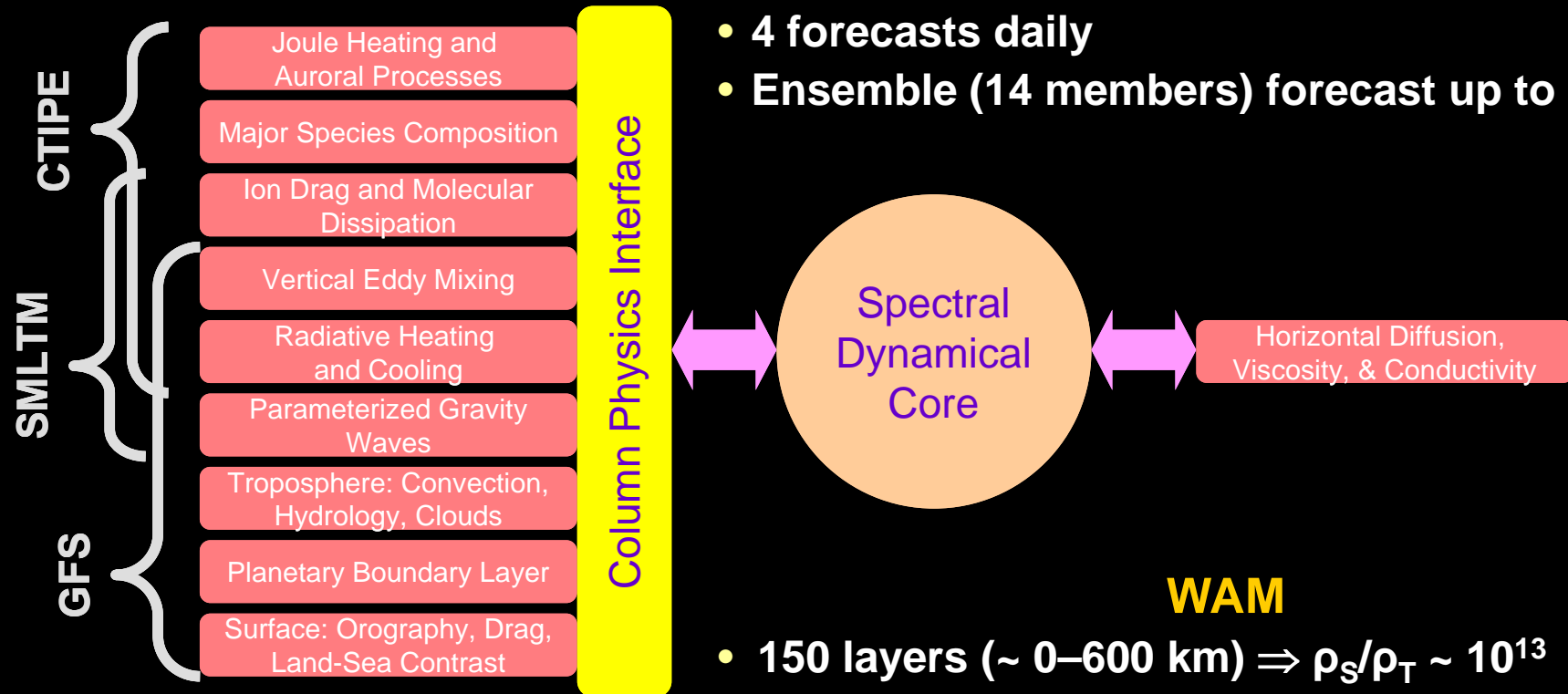
The four peaks in diurnal temperature amplitude result from superposition of the migrating (to the west) tide (DW1) and nonmigrating eastward mode with zonal wavenumber 3 (DE3).

**IMAGE composite of 135.6-nm O airglow (350–400 km) in March–April 2002 for 20:00 LT and amplitude of modeled diurnal temperature oscillation @ 115 km (Immel et al., 2006).**

# WAM = Extended GFS + Physics

## Operational **Global Forecast System (GFS)**

- T382L64 (~ 0.3°×0.3°, ~ 0–62 km)
- Hydrology, surface exchange processes, ozone transport, radiation, cloud physics, etc.
- 4 forecasts daily
- Ensemble (14 members) forecast up to 16 days



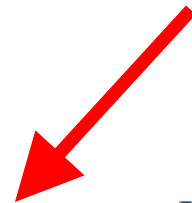
## WAM

- 150 layers (~ 0–600 km)  $\Rightarrow \rho_s/\rho_T \sim 10^{13}$
- Numerical stability issues
- Timing: T62L150 ~ 8 min/day on 32 P6 nodes
- **Column + horizontal-surface physics**
- **Variable composition  $\Rightarrow$  thermodynamics**

# Composition effects on thermodynamics

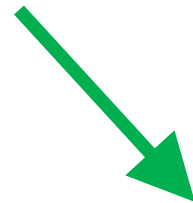
Traditional approximation

$$\frac{dc_p T}{dt} - \frac{1}{\rho} \frac{dp}{dt} = Q$$


$$c_p \frac{dT}{dt} - RT \frac{dp}{p dt} = Q$$

$$\frac{dT}{dt} - \kappa T \frac{dp}{p dt} = \frac{Q}{c_p}$$

Consistent approach

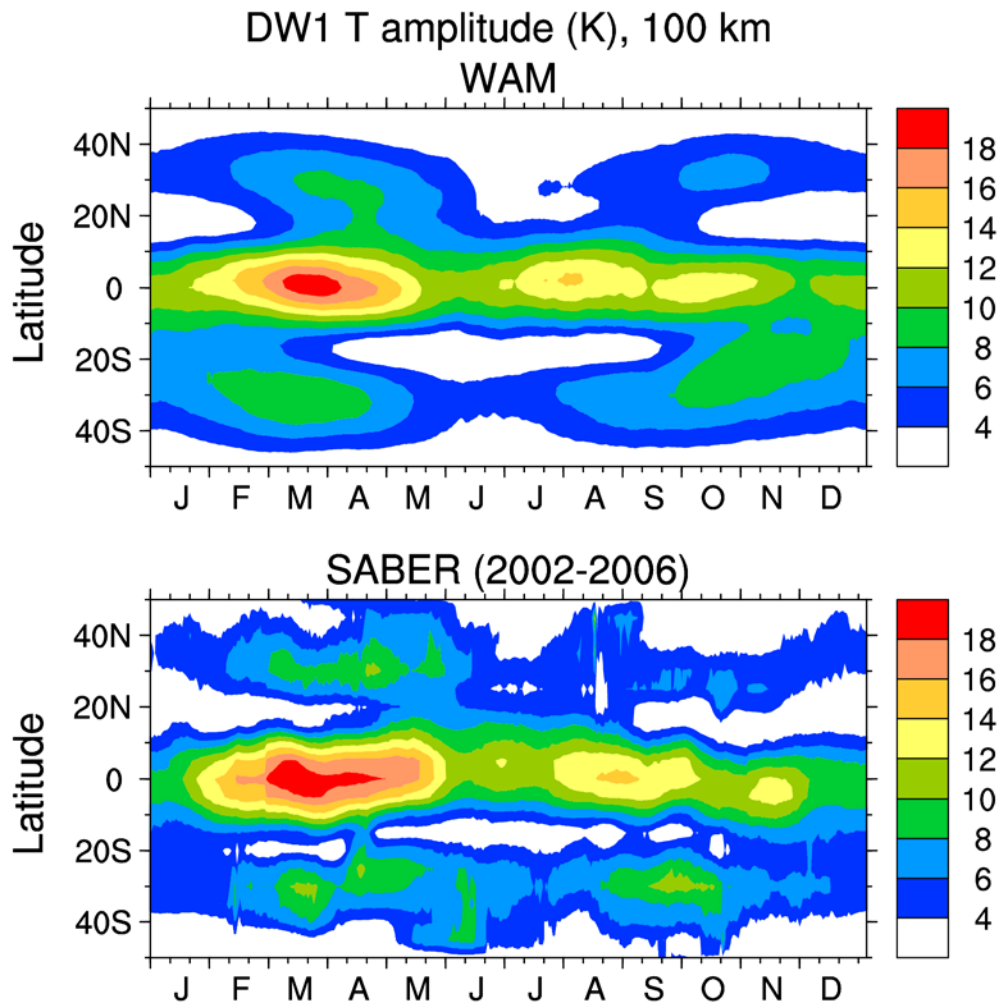

$$\frac{dh}{dt} - \kappa h \frac{dp}{p dt} = Q$$

$$h = c_p T$$

$$\kappa = R/c_p$$

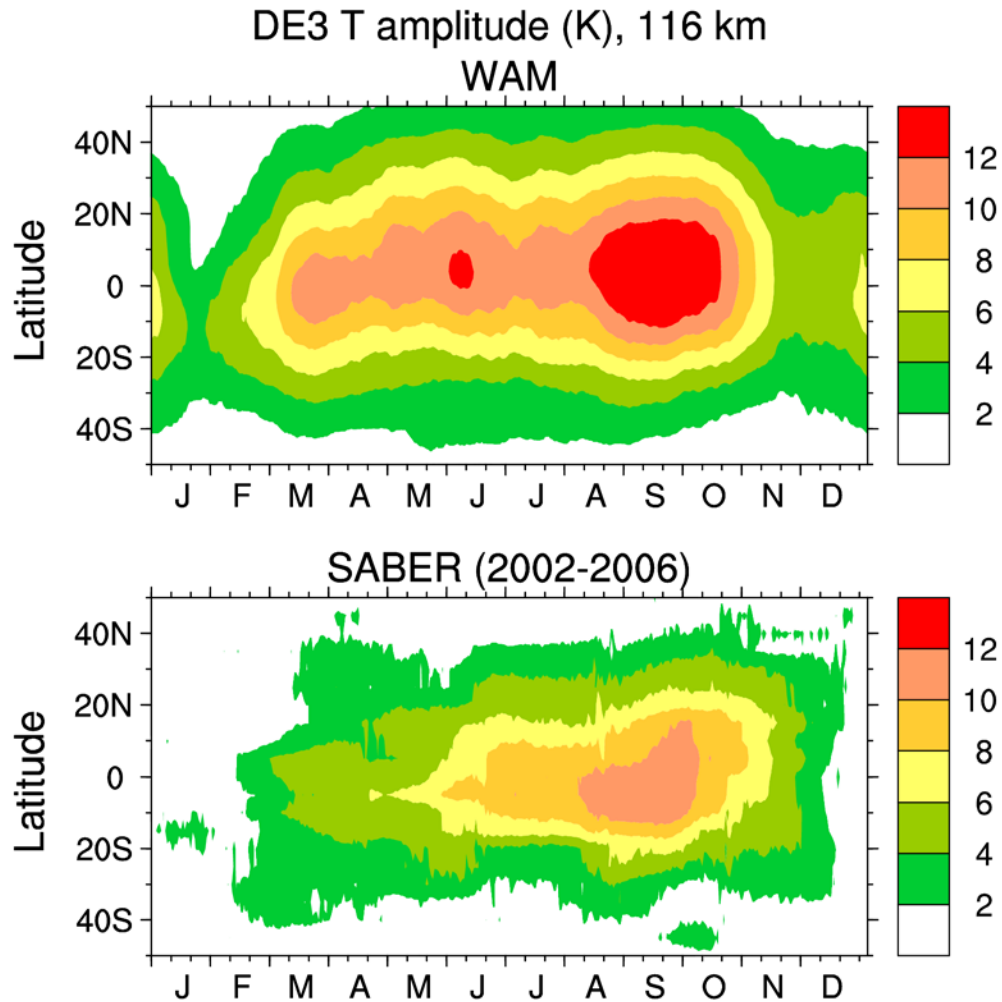
Richmond & Matsushita (JGR, 1975), Fuller-Rowell & Rees (JAS, 1980),  
Akmaev and Juang (QJRMS, 2008)

# Validation: DW1 tide @ 100 km



**Example: DW1 migrating tidal temperature amplitude compared to TIMED/SABER data analysis (Forbes et al., 2008) @ 100 km.**

# Validation: DE3 tide in the E-layer

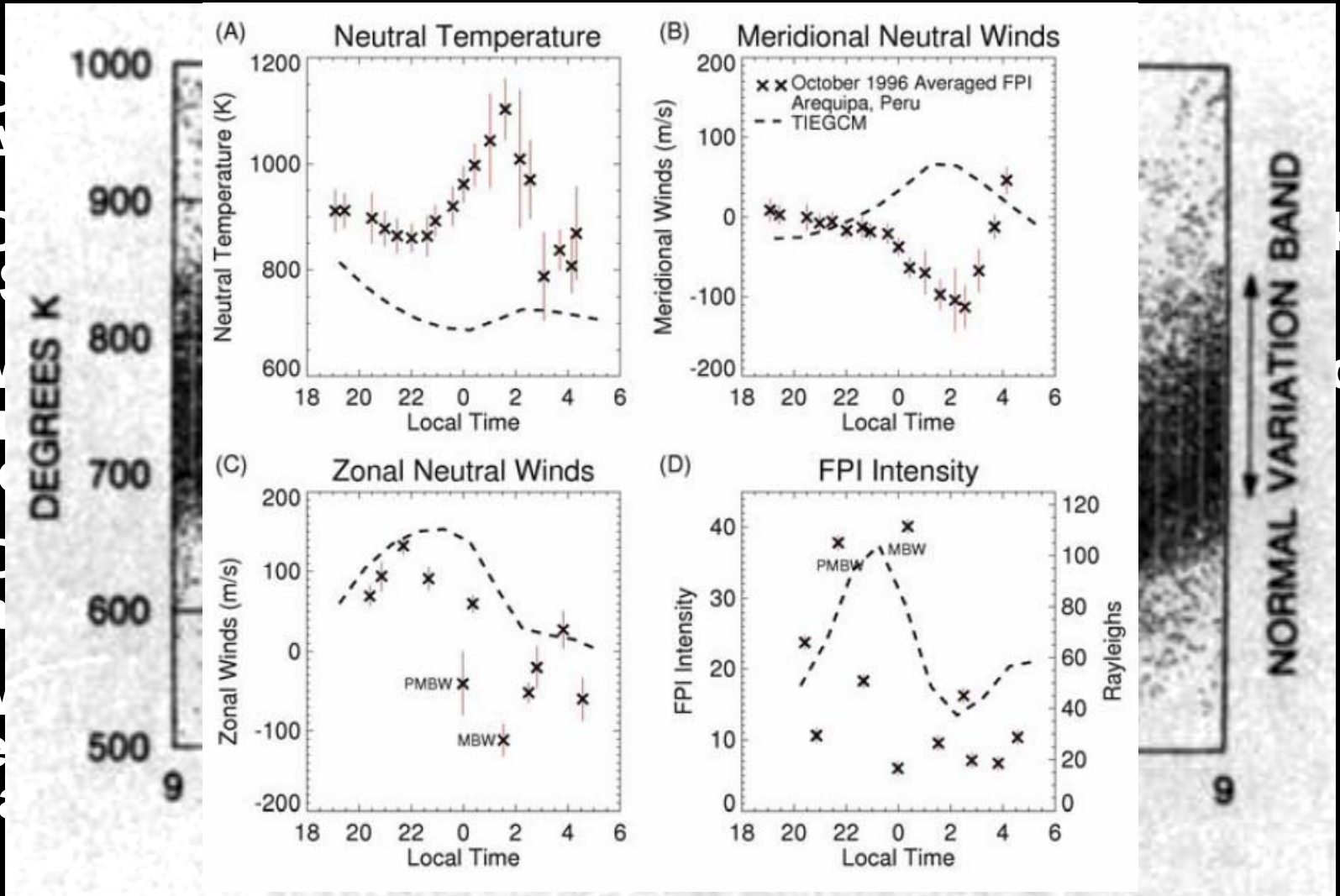


**Example: DE3 nonmigrating tidal temperature amplitude compared to TIMED/SABER data analysis (Forbes et al., 2008) in the E-layer dynamo region.**

**Other tides (e.g., SW2) and variables (winds) validated as well (Akmaev et al., GRL, 2008; Fuller-Rowell et al., GRL, 2008).**

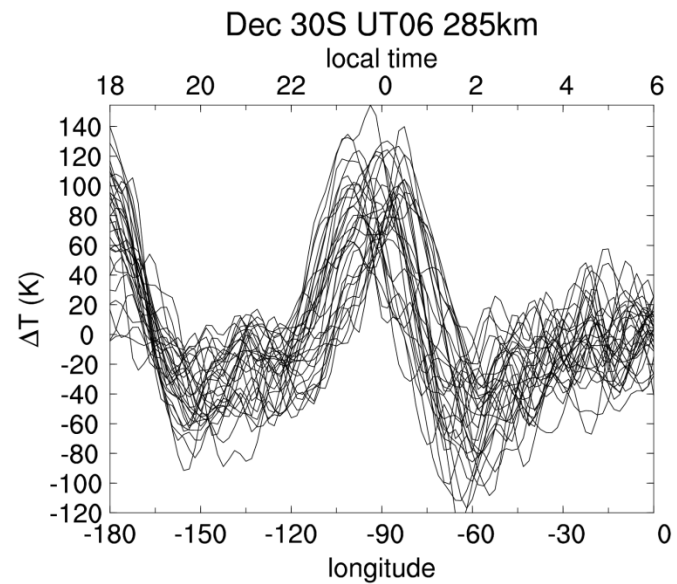
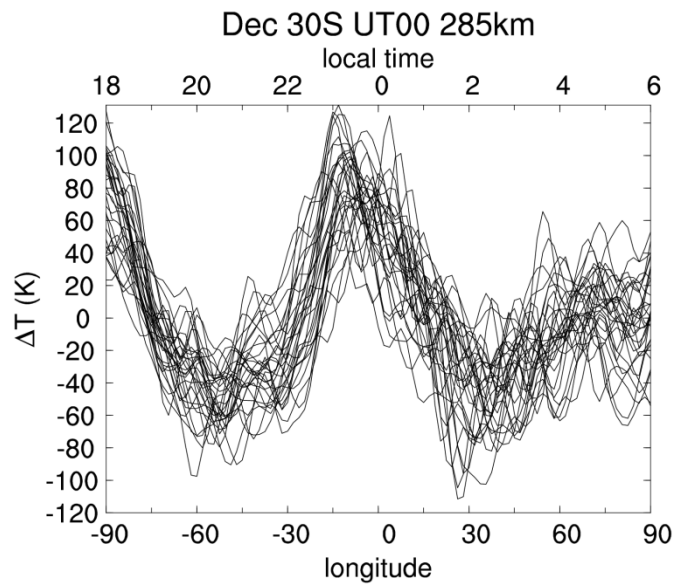
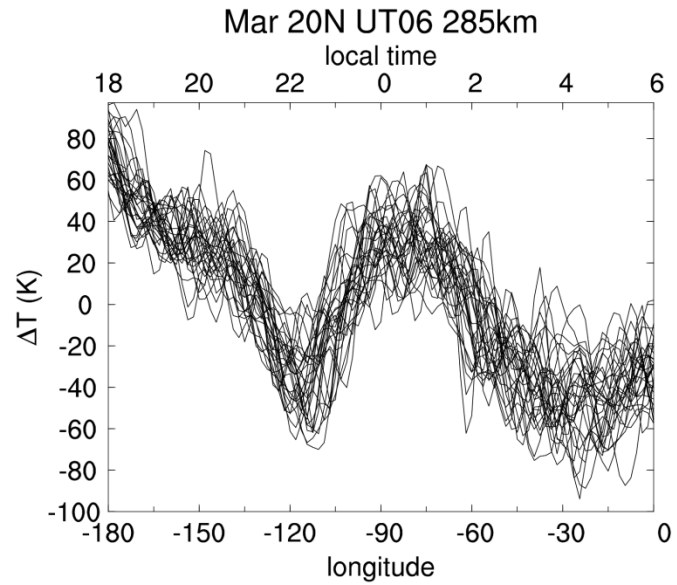
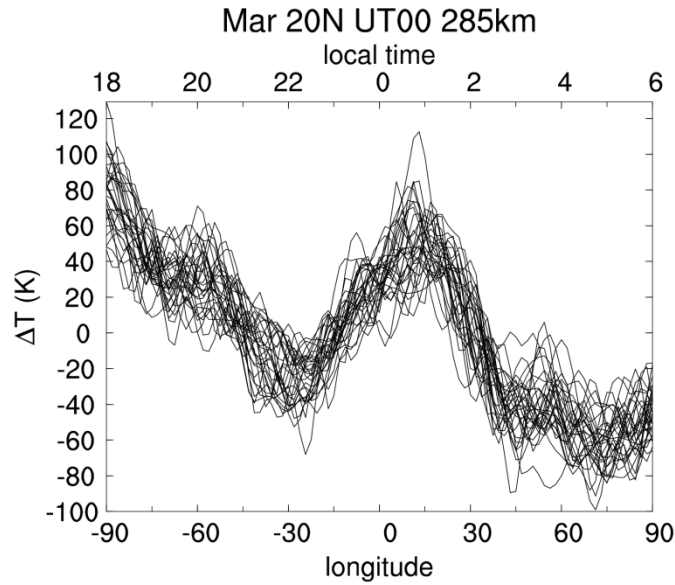


# Midnight temperature maximum (MTM)

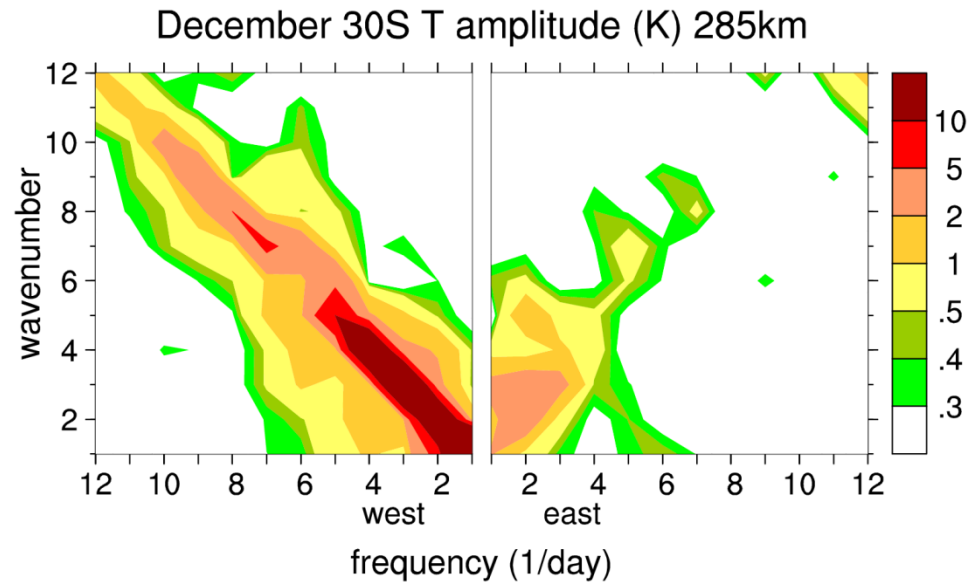
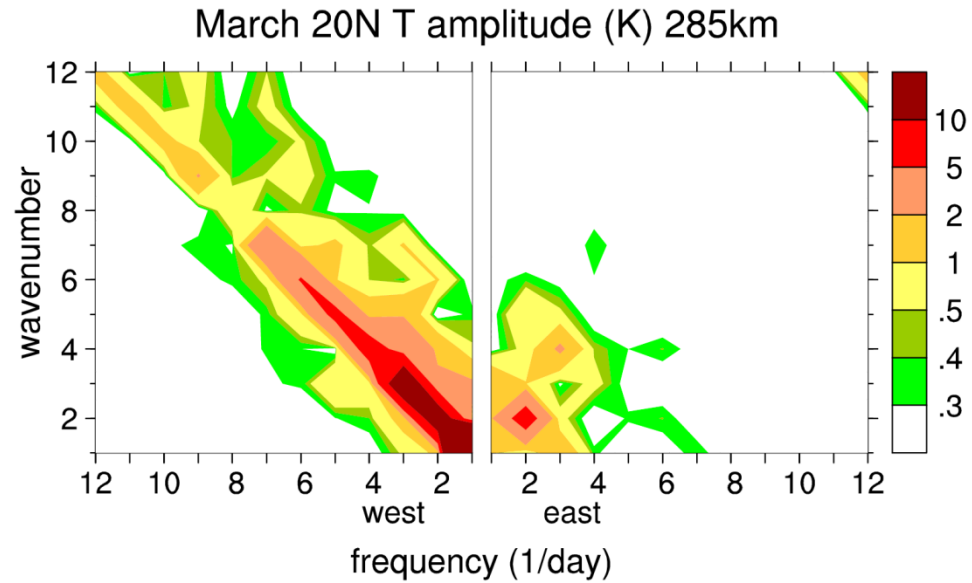


FPI observations at 37°S compared to TIEGCM & latitude simulations (Golev & Coppen, 2006), 1979)

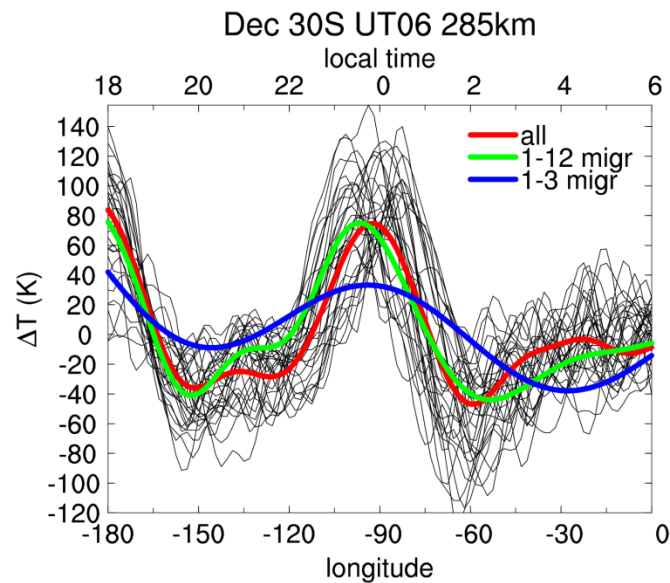
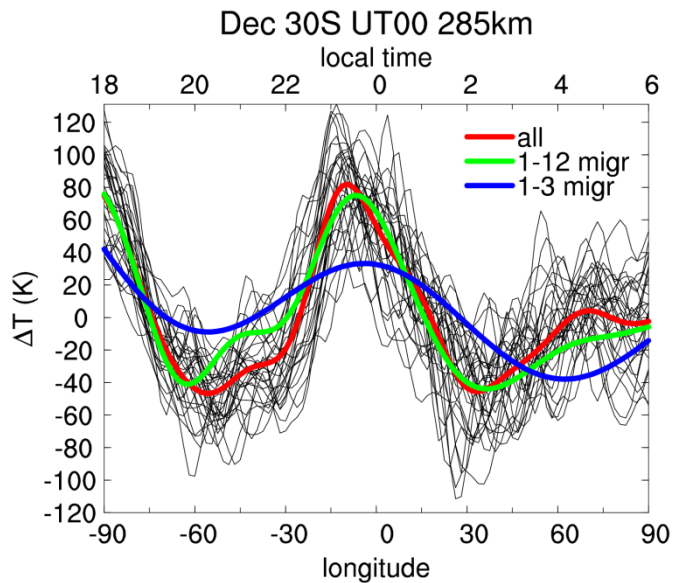
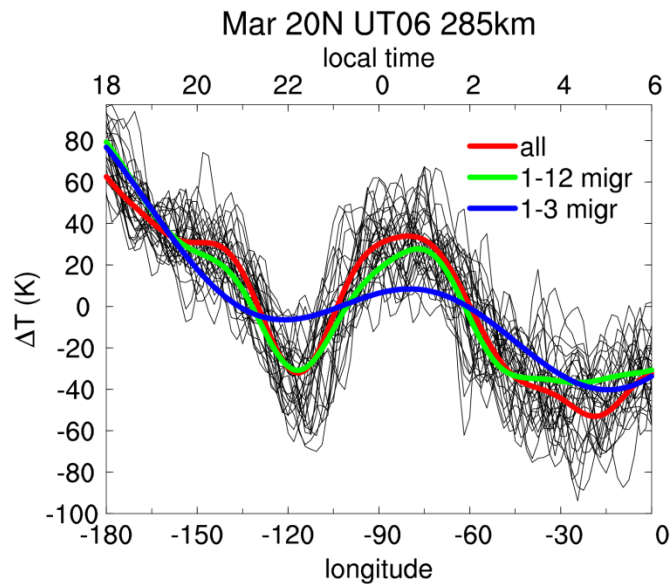
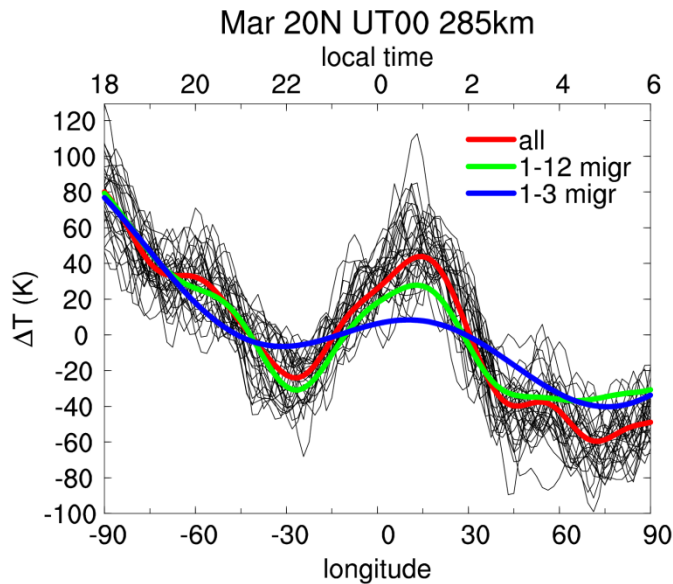
# MTM in WAM: Variability



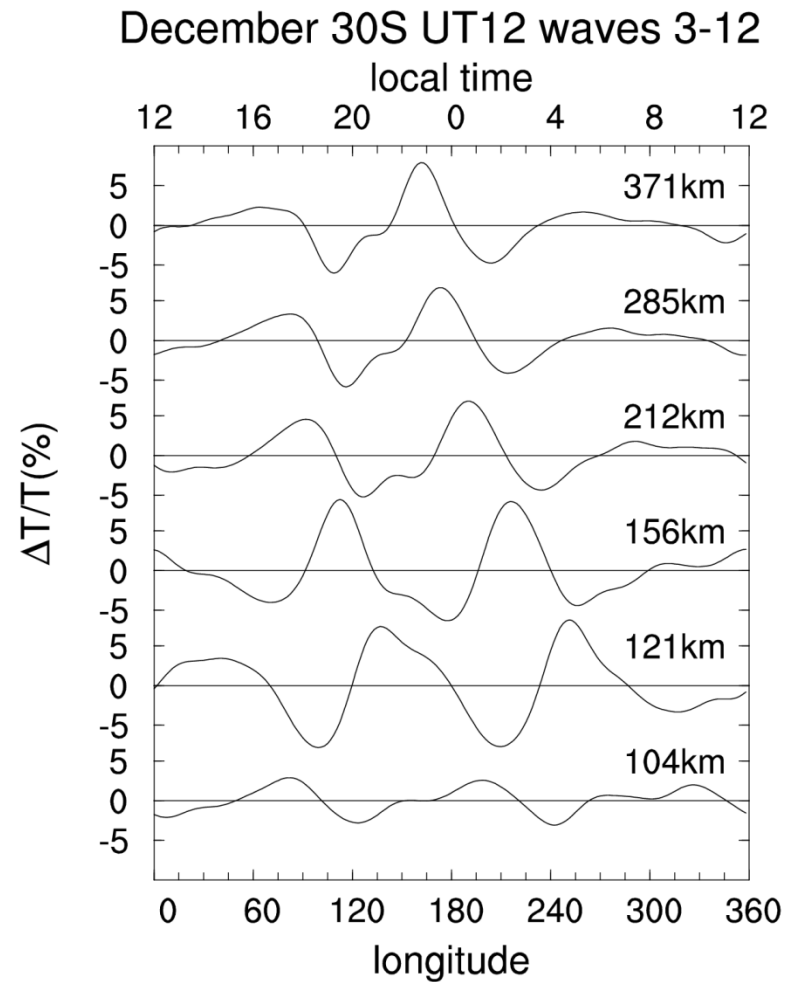
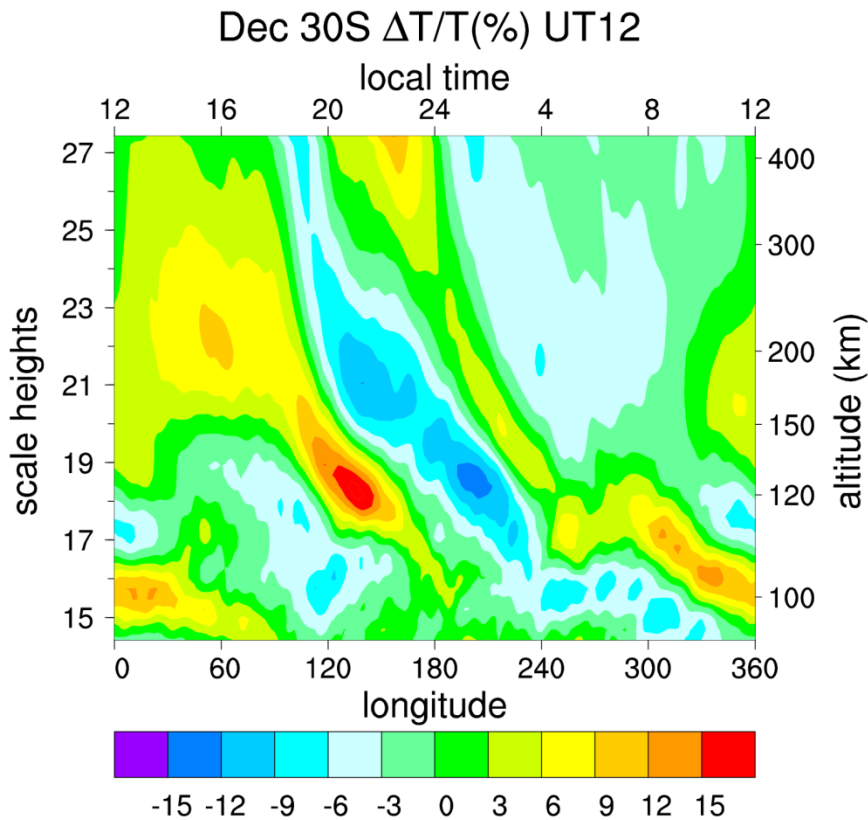
# MTM in WAM: Spectrum



# MTM in WAM: Spectral contributions



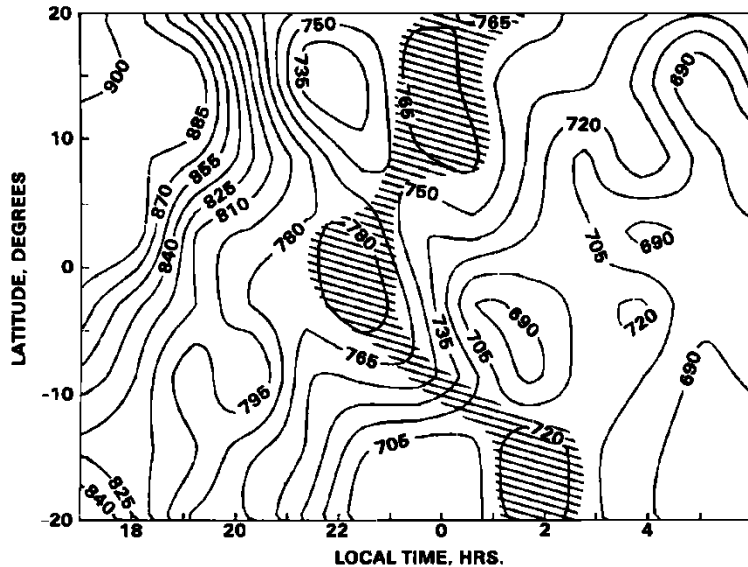
# MTM in WAM: Connection to lower thermosphere



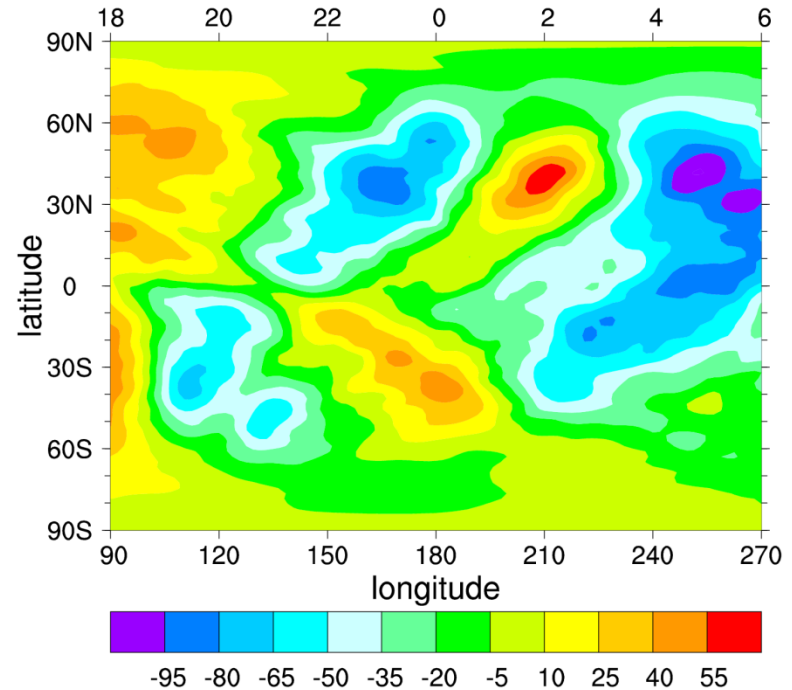
# MTM in WAM: Seasonal and latitudinal variability

THERMOSPHERIC TEMPERATURE MAP  
AE-DATA DAYS 77142 TO 77188

AVERAGE ALTITUDE 250 KM



Dec UT12  $\Delta T$  (K) 285km  
local time



AE-E data, May–July 1977  
(Herrero and Spencer, 1979)

# Summary

- **WAM has been validated on satellite observations of tides propagating from below to the thermosphere, including the non-migrating waves implicated in the observed global morphology of the ionosphere.**
- **WAM is the first comprehensive model to realistically reproduce the thermospheric MTM, including its variability. The MTM appears to extend into mid-latitudes in agreement with recent observations.**
- **Preliminary analysis indicates that the MTM is primarily produced by the terdiurnal and higher-order waves propagating upward from below and possibly interacting with the primary tidal harmonics and dissipative processes in the thermosphere.**