

# Soil Carbon modeling in CLM4CN

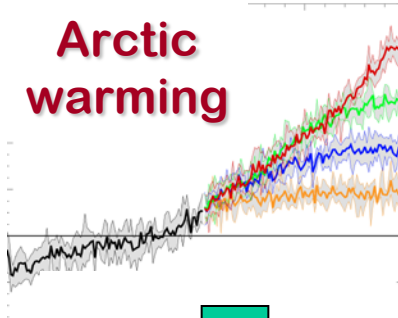
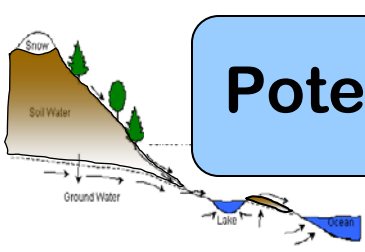
**David Lawrence,**

**Sean Swenson**

**NCAR, Boulder, CO**



# Potential Arctic terrestrial climate-change feedbacks



**Global warming**

**Arctic warming**

**Carbon sequester**

**Shrub growth**

**Enhanced [nitrogen]**

**Microbial activity increases**

**Permafrost warms and thaws**

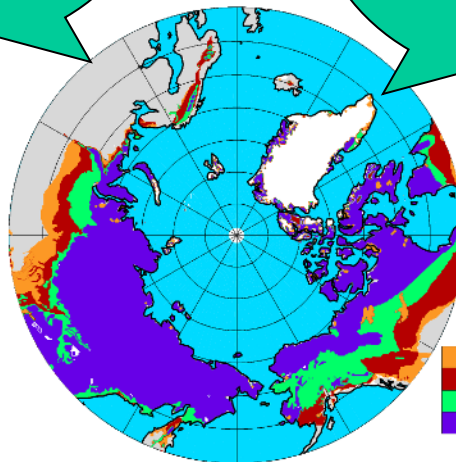
**Expanded wetlands**

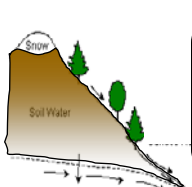
**Lakes drain, soil dries**

**CH<sub>4</sub> efflux**

**CO<sub>2</sub> efflux**

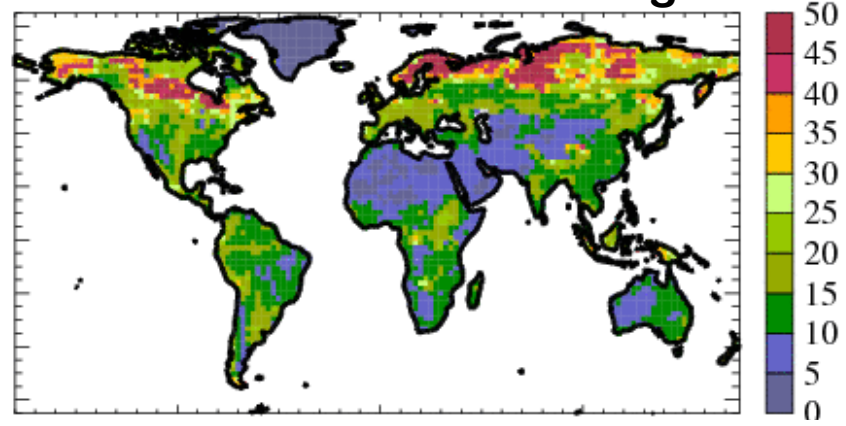
**Arctic runoff increases**



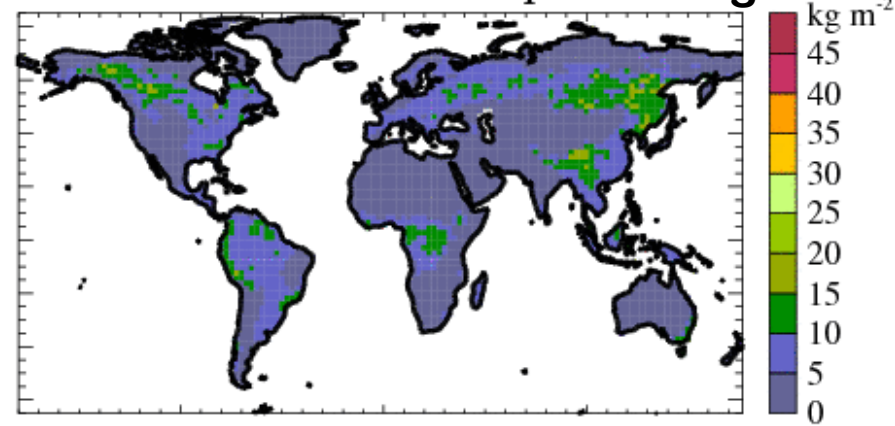


# CN Soil carbon compared to Global Soil Data Task obs

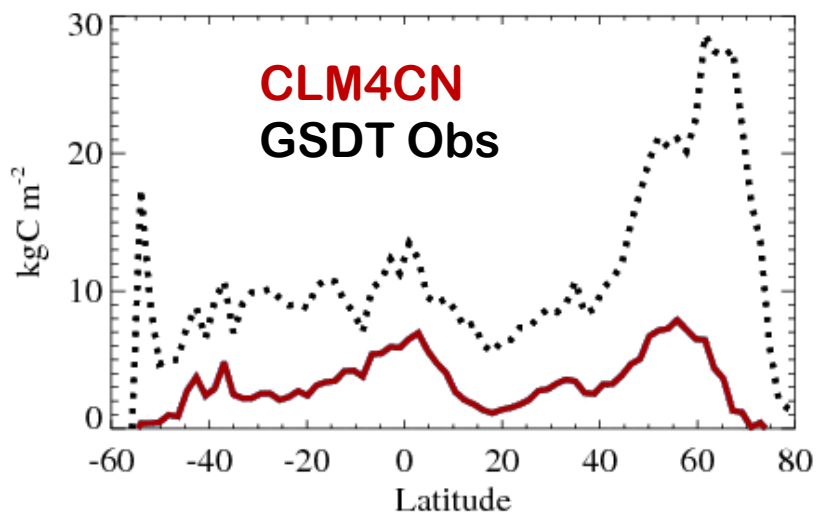
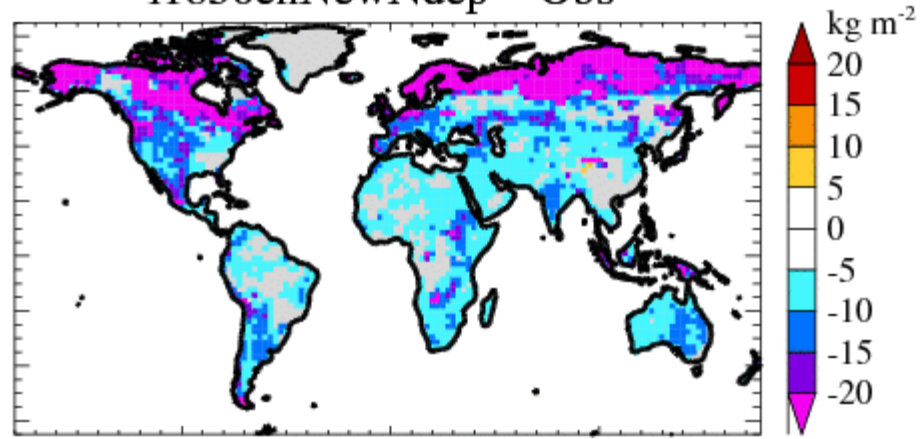
Obs ~2000 PgC

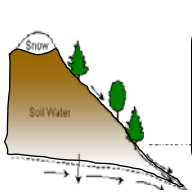


i1850cnNewNdep ~500 PgC

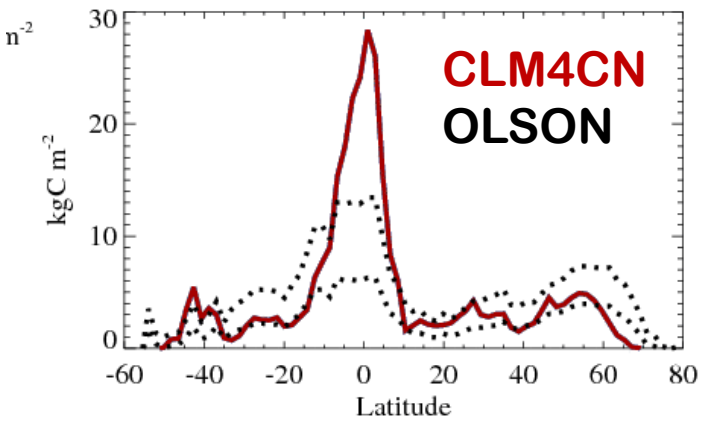
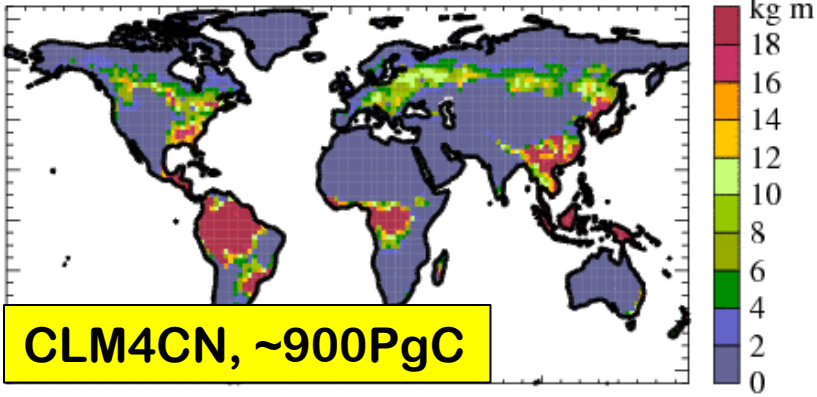
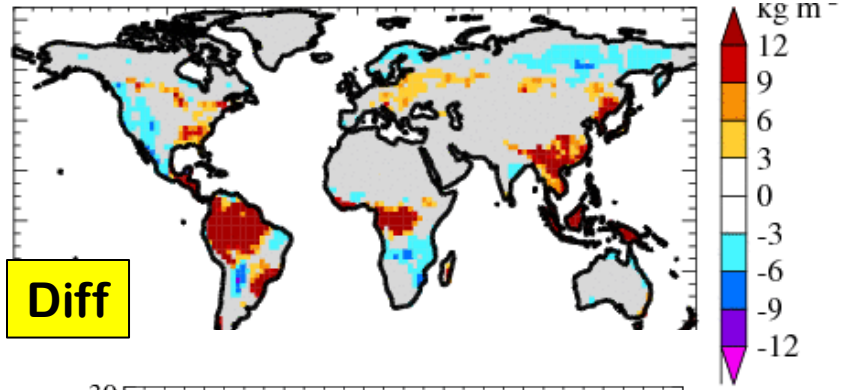
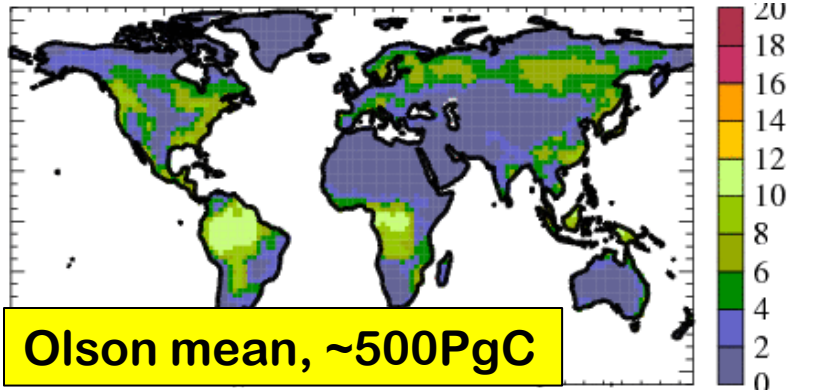
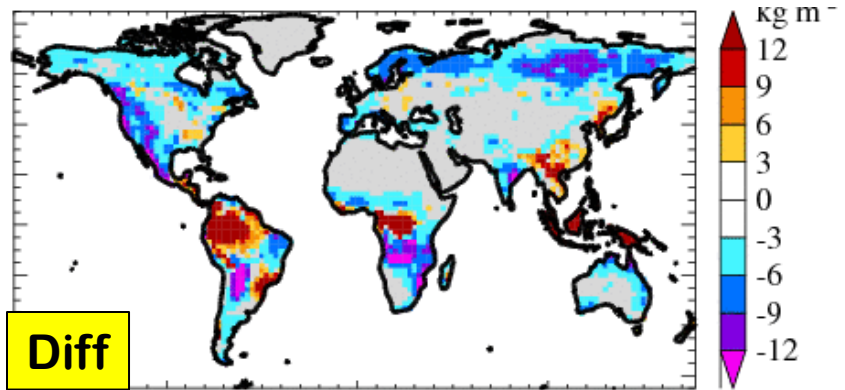
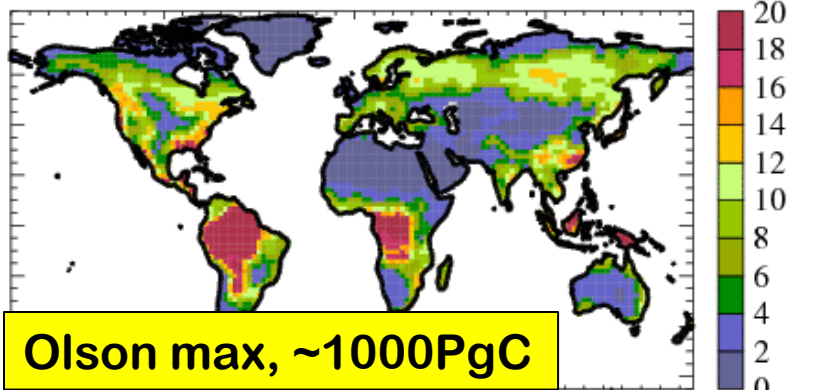


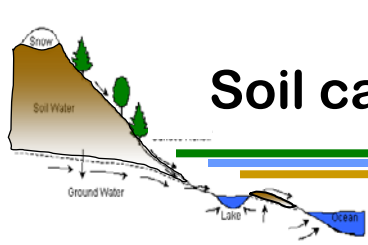
i1850cnNewNdep - Obs





# Vegetation carbon compared to Olson et al. (1985) (updated by Gibbs et al. 2007)

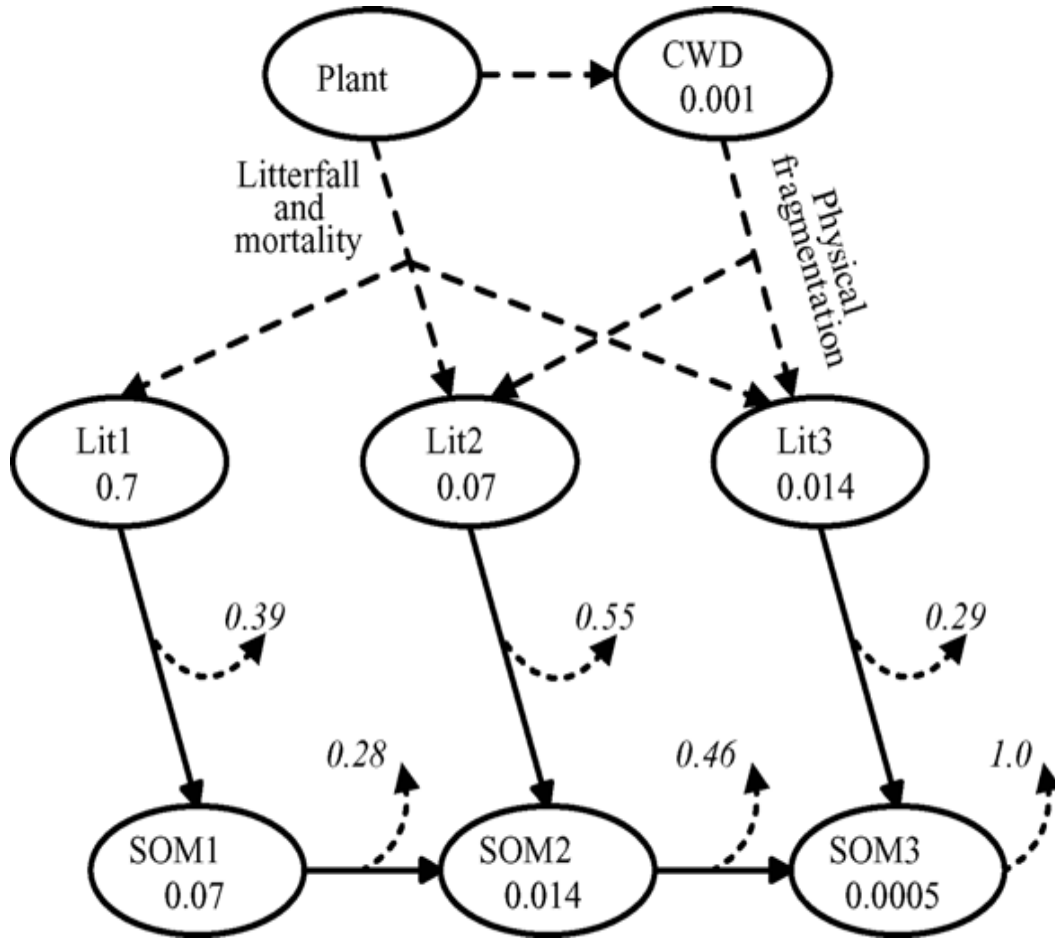
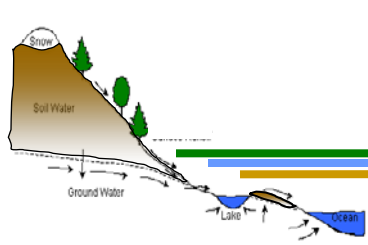




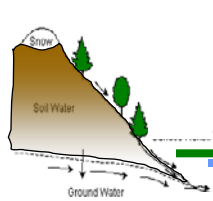
# Soil carbon: Issues from perspective of northern high latitudes

- In tundra zones, very low vegetation growth CLM4CN (at least partly due to hydrology problem)
- Soil decomposition rates
  - No limits due to anoxia at high water levels
  - Location of soil carbon is ‘virtual’ within top 5 model levels
- Large carbon stores result of thousands of years of accumulation (with differing initiation dates) in peatlands or similar systems
- Not representing unique biogeochemistry of peatlands

# Heterotrophic soil respiration in CLM-CN



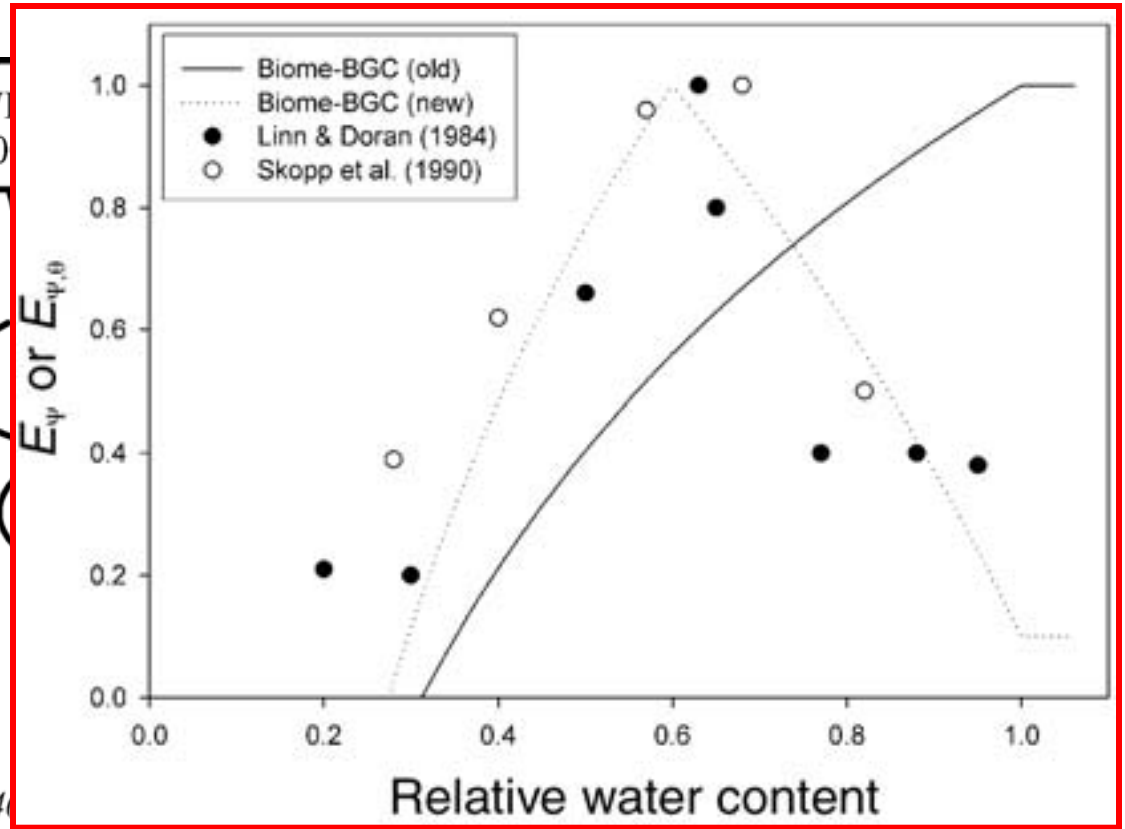
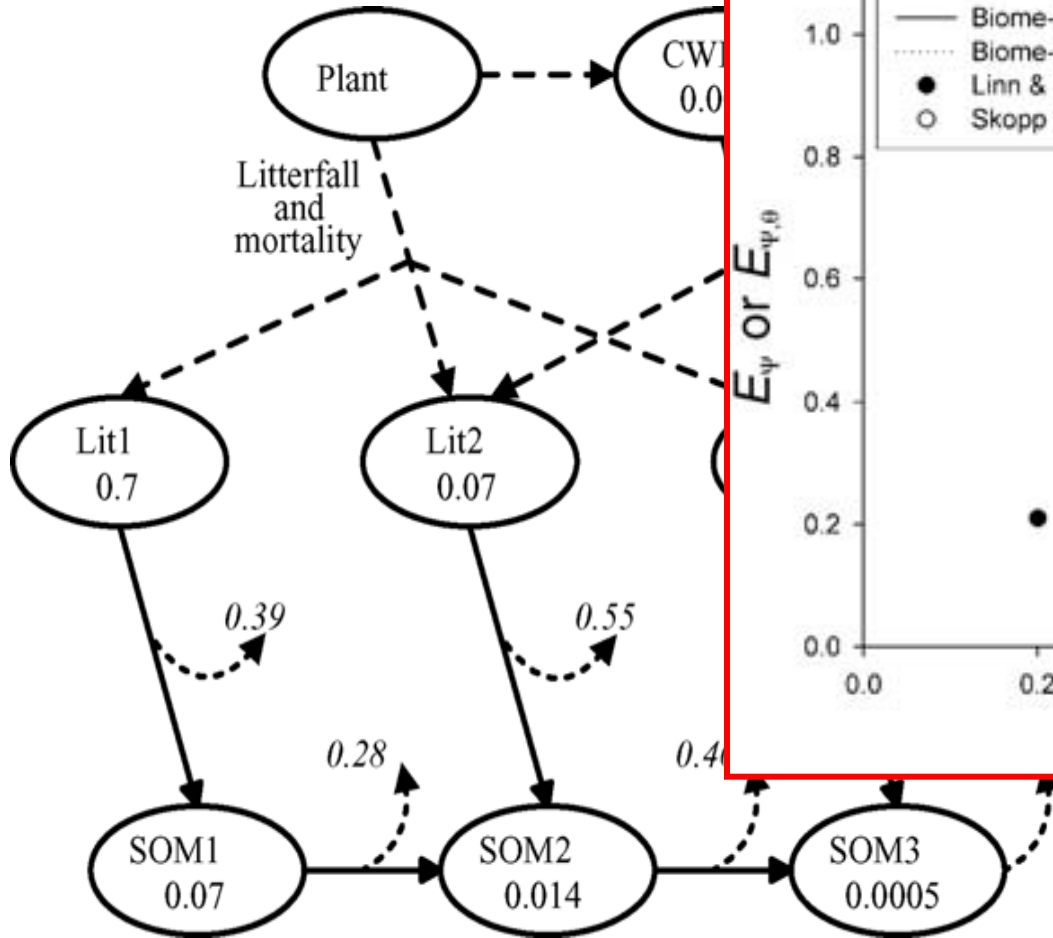
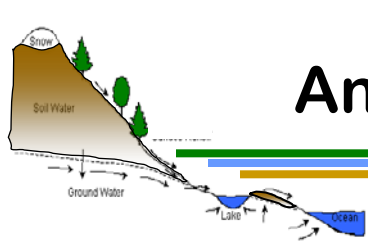
**Base decomposition rates for each SOM pool are modified by functions of water and temperature**



# Proposed modifications

- Cold region hydrology modifications from Sean
- Connect organic soil thermal and hydrologic properties (Lawrence and Slater, 2008) with prognostic CN soil carbon
  - Represent vertical decrease in hyd. conductivity from fabric to sapric peat - wetter soil in organic rich regions
- Incorporate anoxia limitation on decomposition rates
  - Sync up CLM soil suction with CN soil suction
- Account for impact of vertical distribution of soil carbon on decomposition rates
- Change Q10 from 1.5 to 2 or ???
- Assume that Arctic C3 grass more like moss – grows in nutrient-limited environs
- Initialize model with ‘observed’ soil carbon and slowly turn on carbon pool transfers

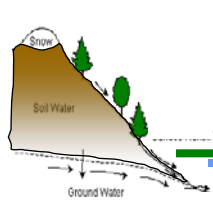
# Anoxia limitation on soil carbon decomposition



Bond-Lamberty et al., 2007

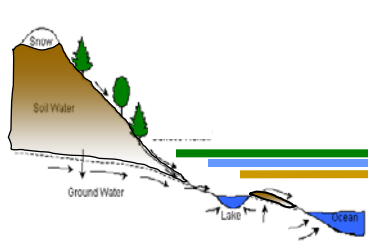


# Proposed modifications



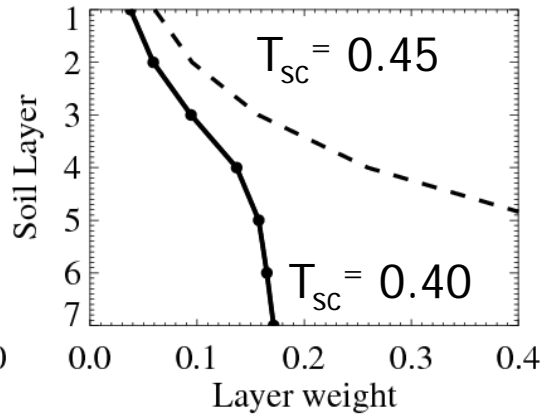
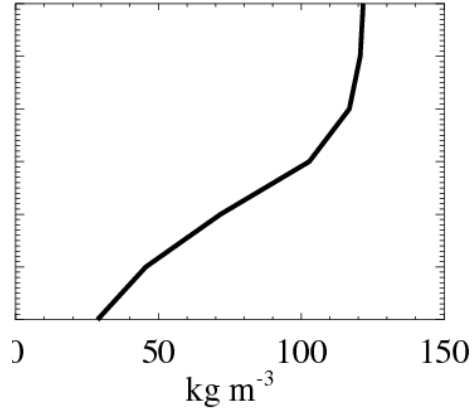
- Cold region hydrology modifications from Sean
- Shallower rooting profile for Arctic C3 grass and boreal shrubs
- Connect organic soil thermal and hydrologic properties (Lawrence and Slater, 2008) with prognostic CN soil carbon
  - Represent vertical decrease in hyd. conductivity from fabric to sapric peat - wetter soil in organic rich regions
- Incorporate anoxia limitation on decomposition rates
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- Assume that Arctic C3 grass more like moss – grows in nutrient-limited environs, leaf litter C:N ratio
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# Vertical distribution of carbon and impact on decomposition rates

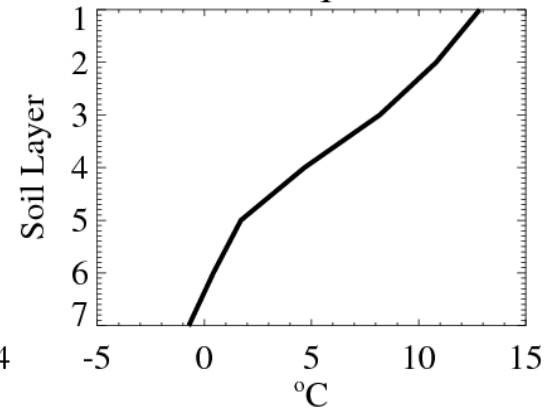


## Siberia peatland

Organic Matter Profile

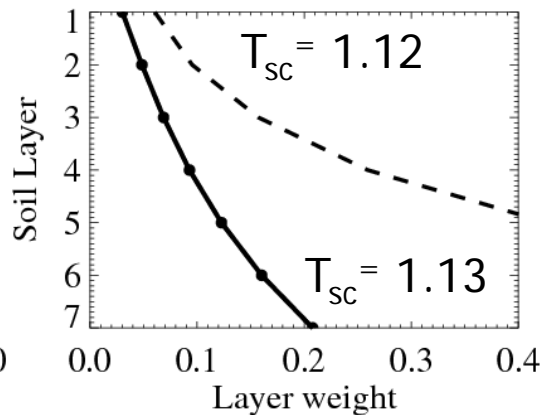
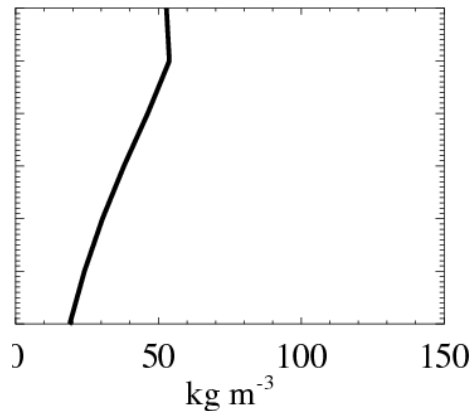


Soil Temperature

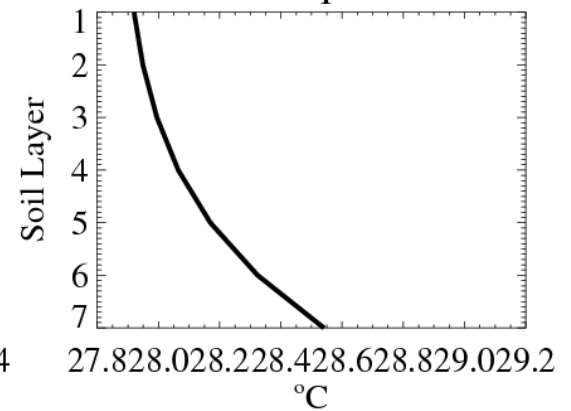


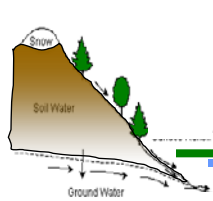
## Tropical Africa

Organic Matter Profile



Soil Temperature

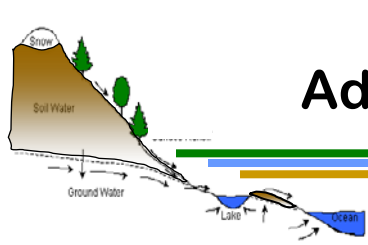




# Proposed modifications

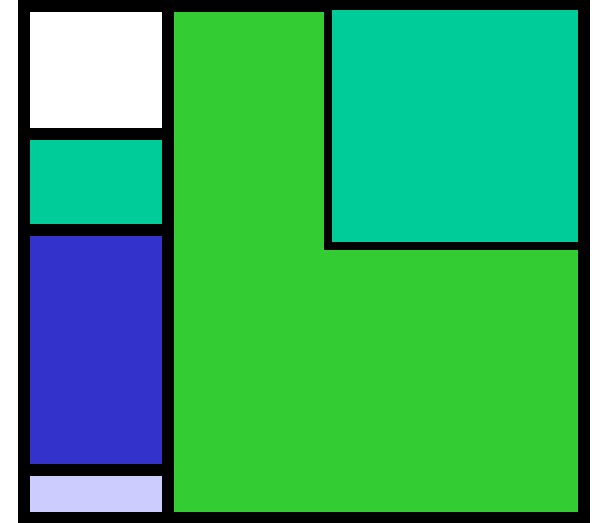
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# Additional proposed revisions – wetland carbon cycling



At each time step:

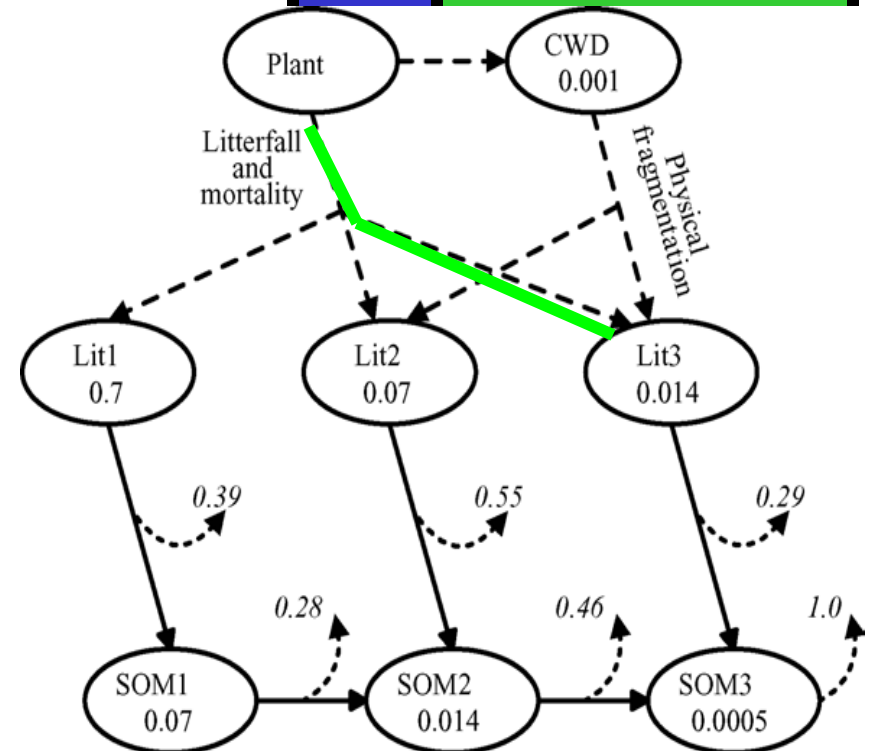
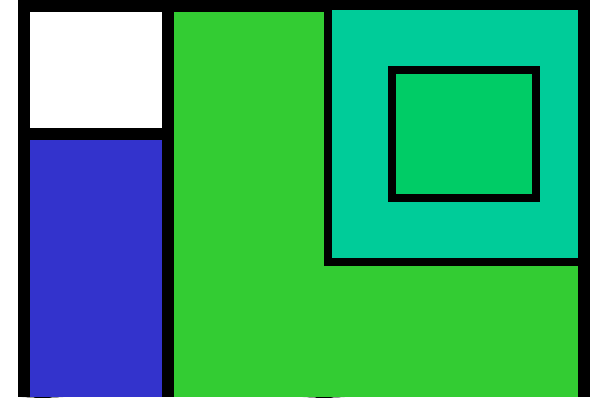
- Calculate saturated fraction of vegetated portion of grid cell (Sean's work)
- For unsaturated fraction of grid cell, soil respiration calculated as above
- For saturated fraction of grid cell, soil respiration at 10% of temperature regulated base rate

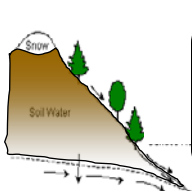


# What about wetland vegetation?

Ideally, need a new 'moss-like' PFT

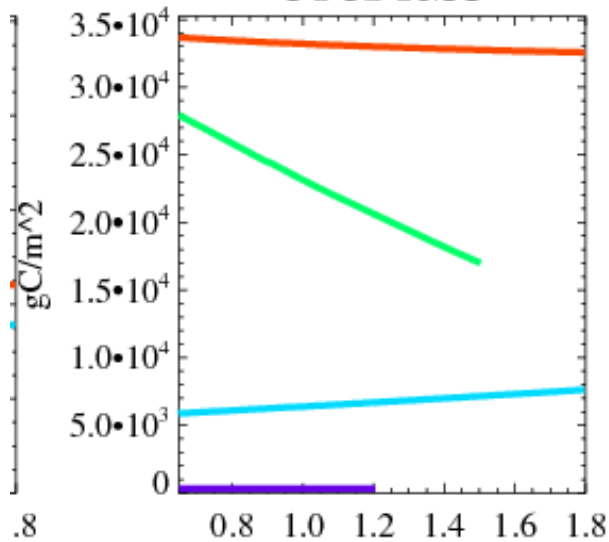
- Assume that moss preferentially inhabits the saturated fraction of grid cell
- Dead moss goes to recalcitrant litter pool
- Short cut: skip moss PFT and simply assume that litter from grass growing in saturated zone goes to recalcitrant litter pool



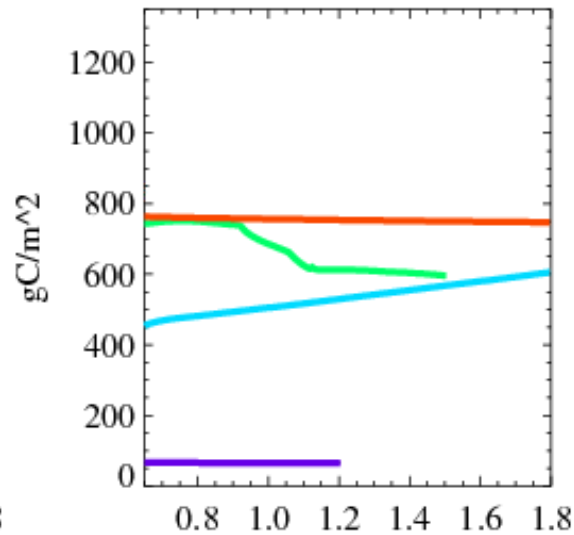


# Arbitrary point in Alaska Arctic

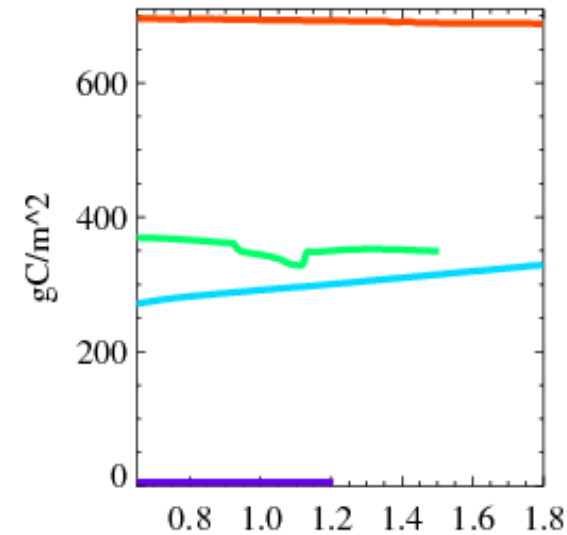
### TOTSOMC



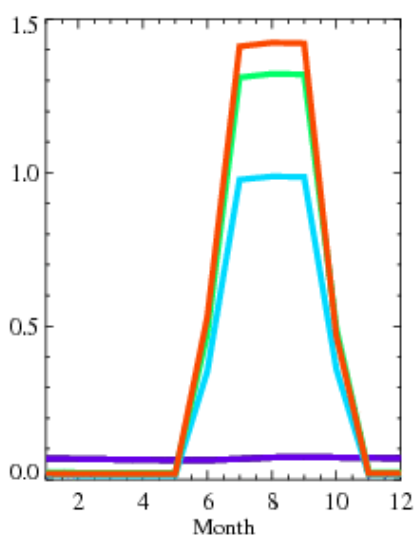
### TOTVEGC



### TOTLITC

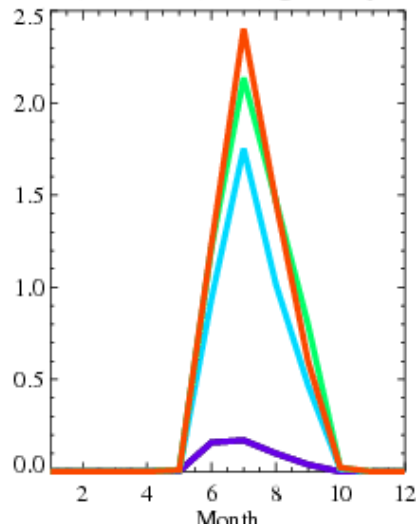


### TLAI



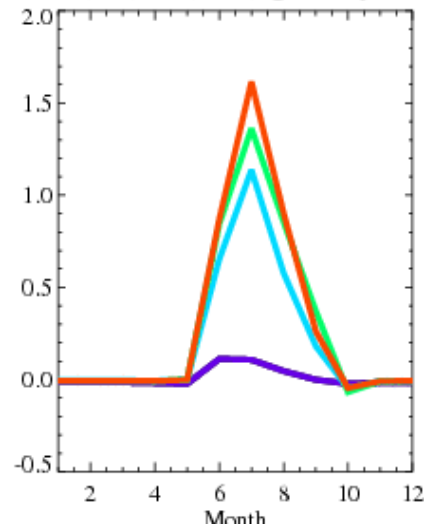
### GPP

14, 128, 172, 174,  $\text{gC m}^{-2} \text{ yr}^{-1}$



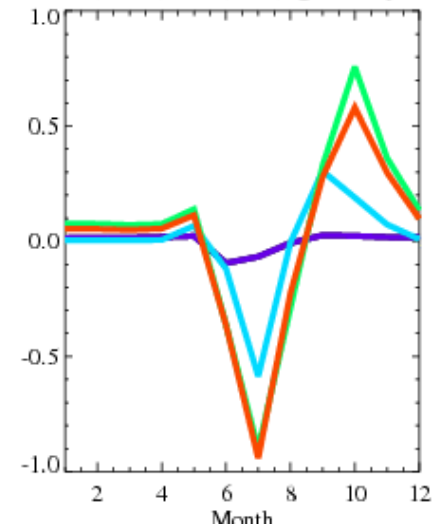
### NPP

4, 75, 101, 108,  $\text{gC m}^{-2} \text{ yr}^{-1}$

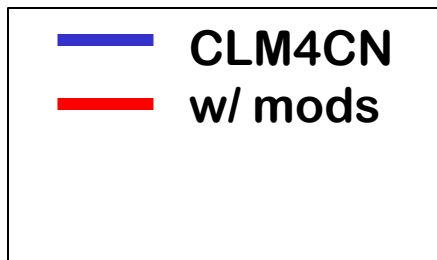
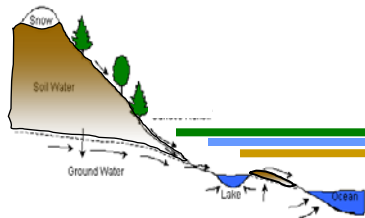


### NEE

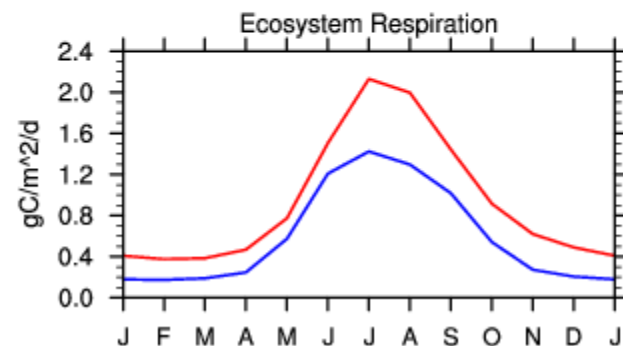
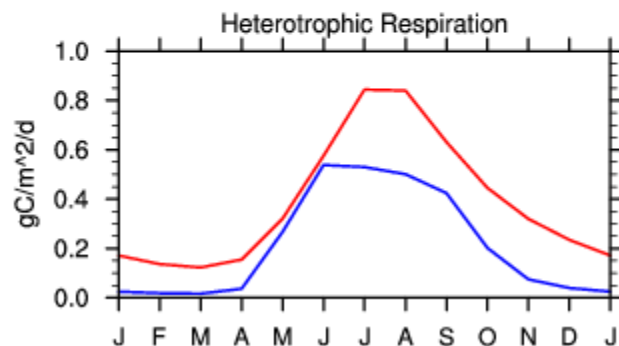
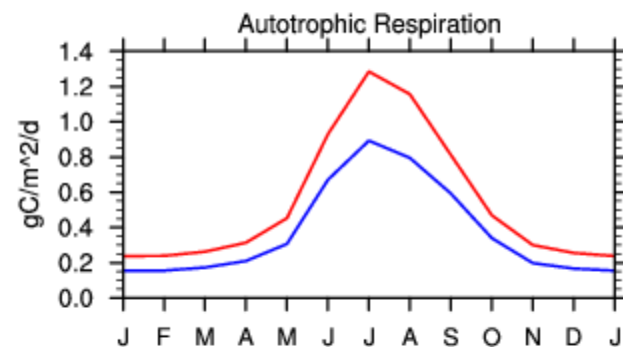
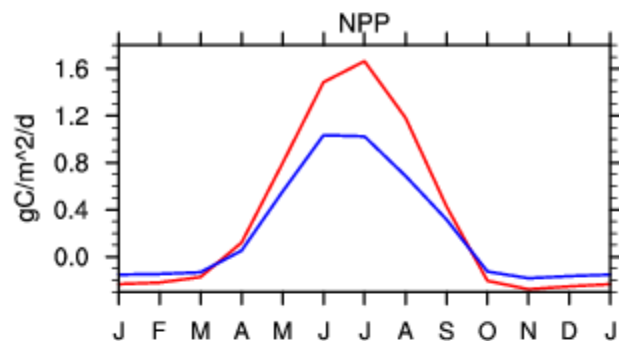
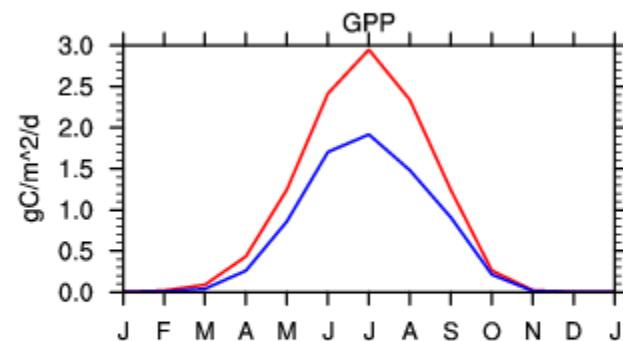
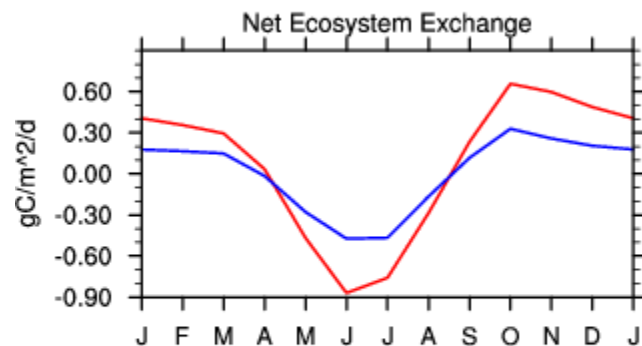
0.0, -1.5, 14.2, 1.5,  $\text{gC m}^{-2} \text{ yr}^{-1}$

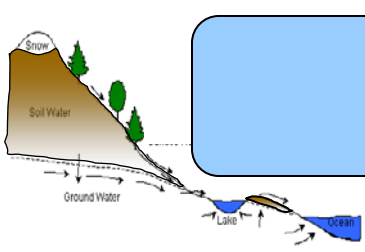


# Some results from global tests



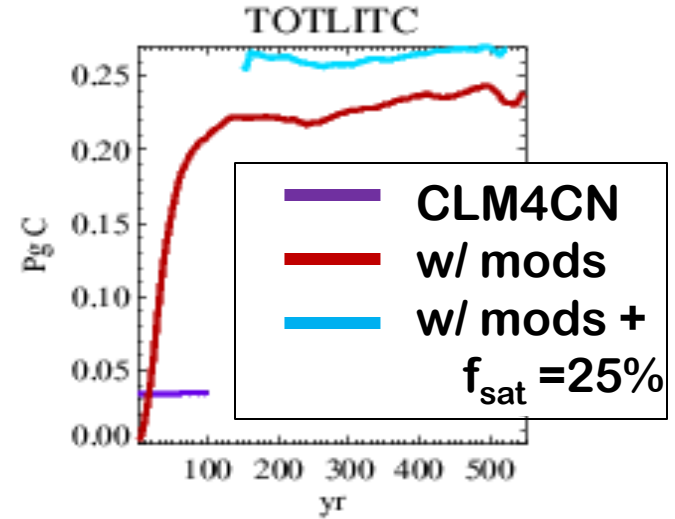
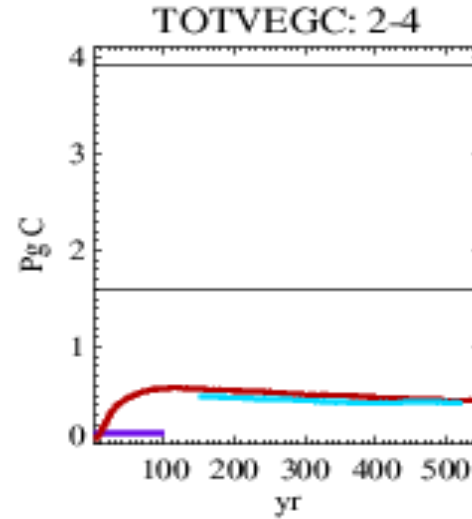
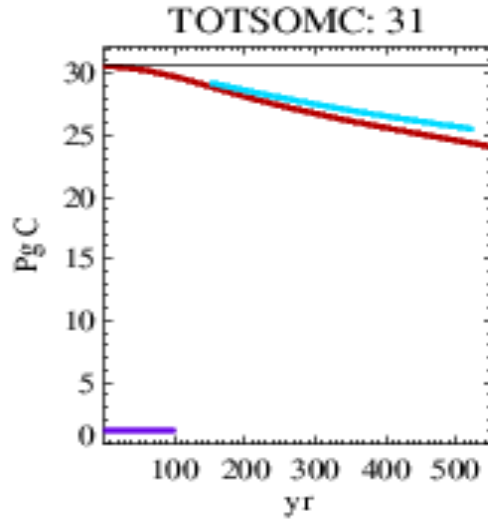
## Polar(60-90N,-180W-180E)



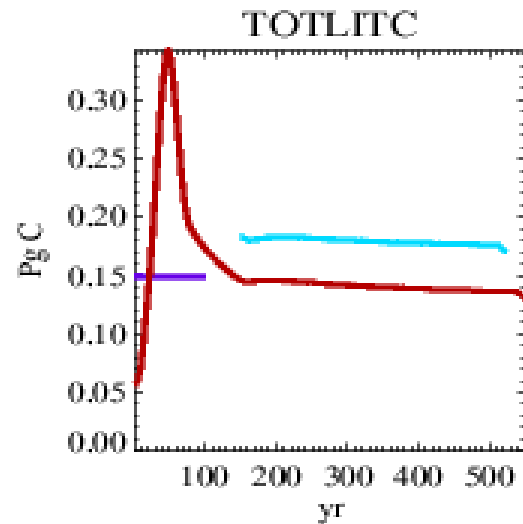
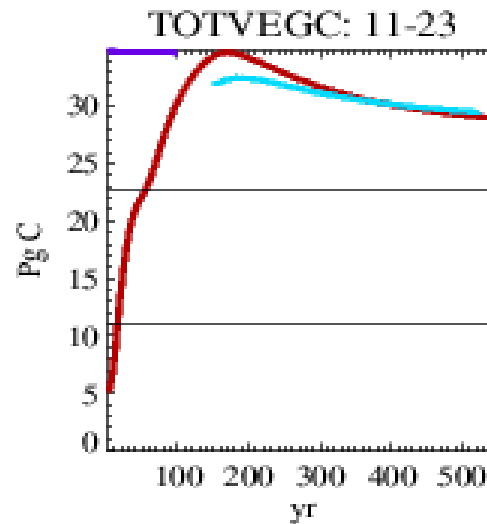
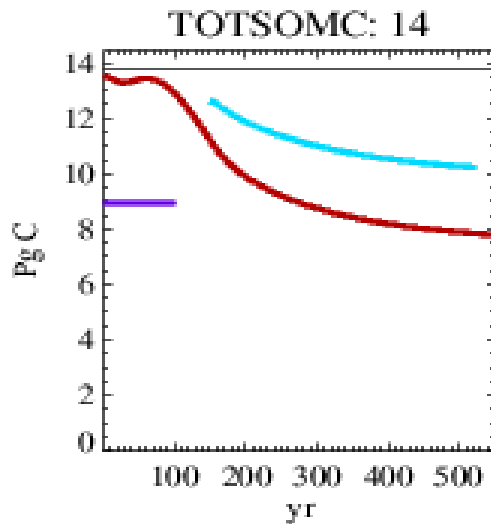


# Results from global runs

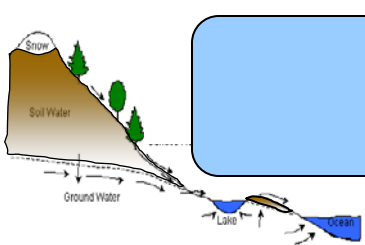
## Alaska Arctic



## Amazonia

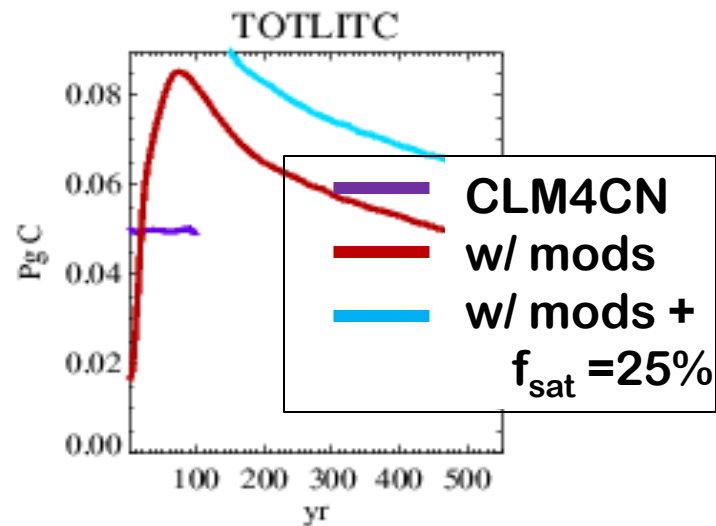
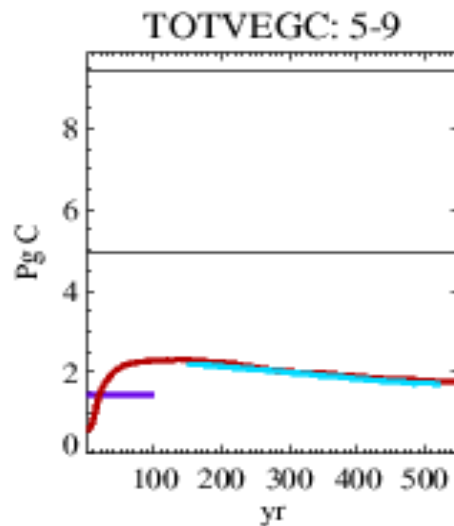
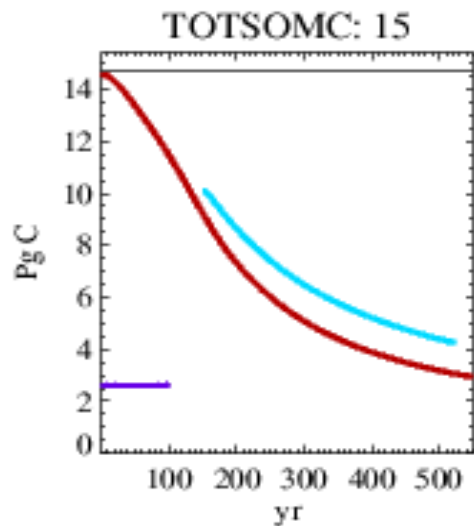




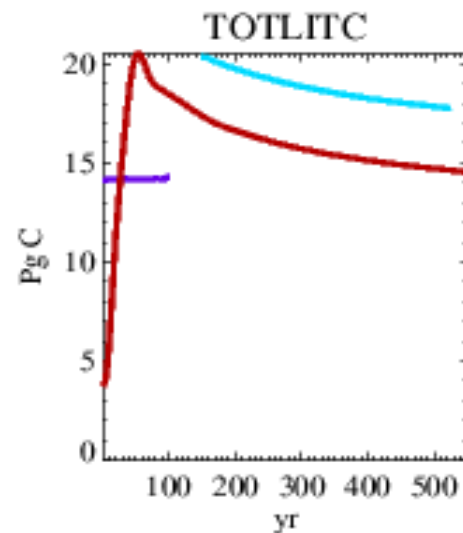
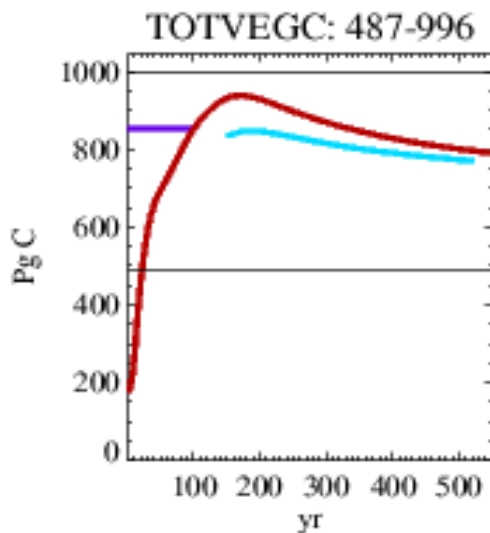
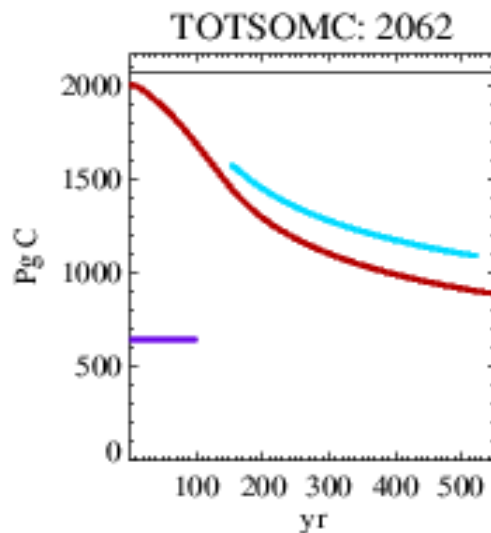


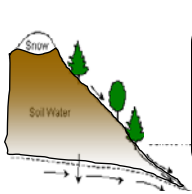
# Results from global runs

## Western US



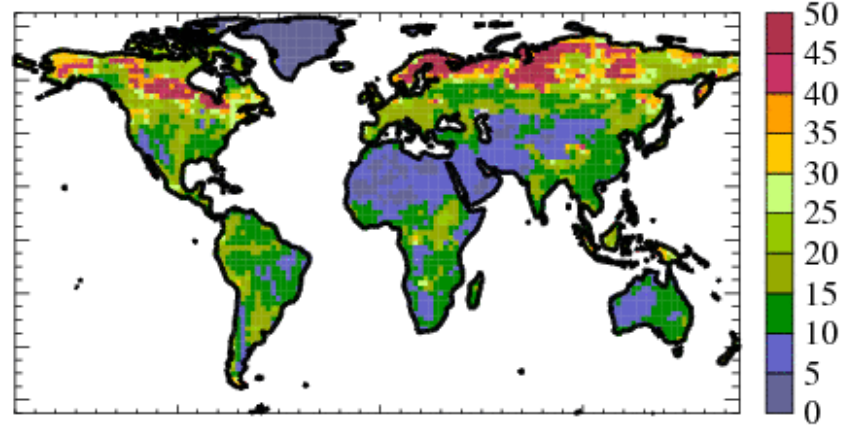
## Global



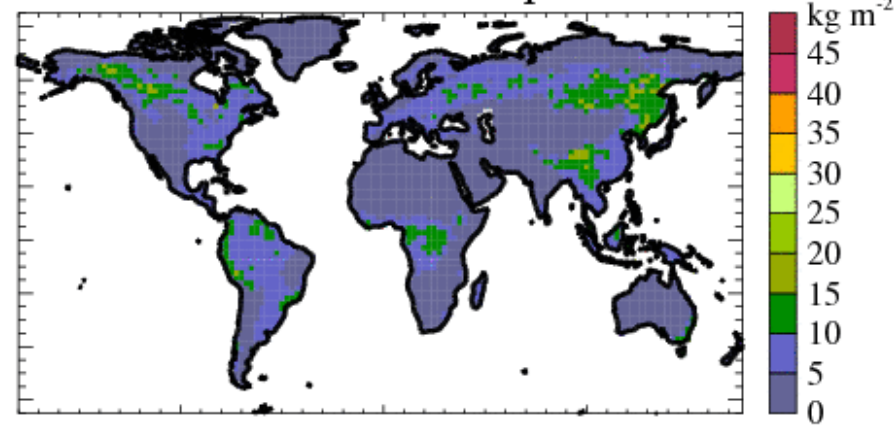


# CN Soil carbon compared to Global Soil Data Task obs

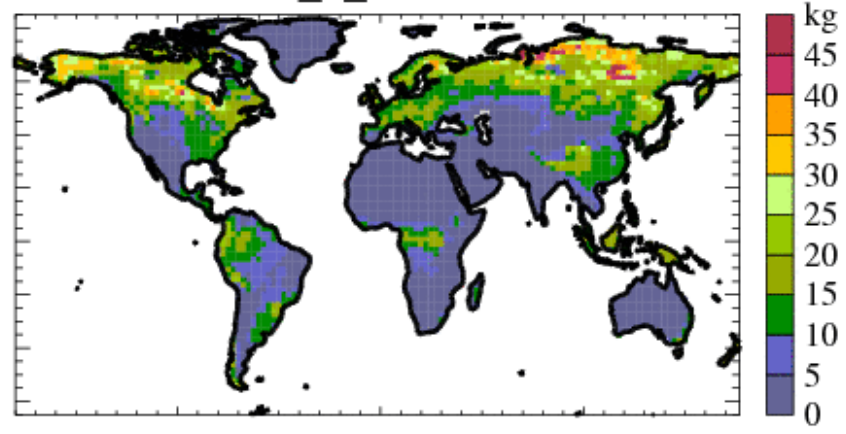
Obs



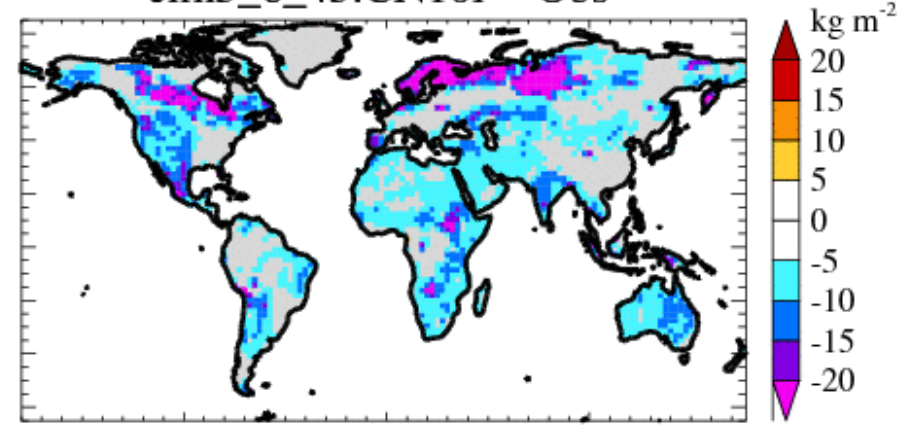
i1850cnNewNdep



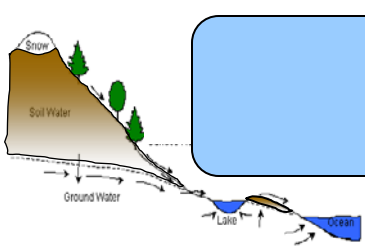
clm3\_6\_45.CN10r



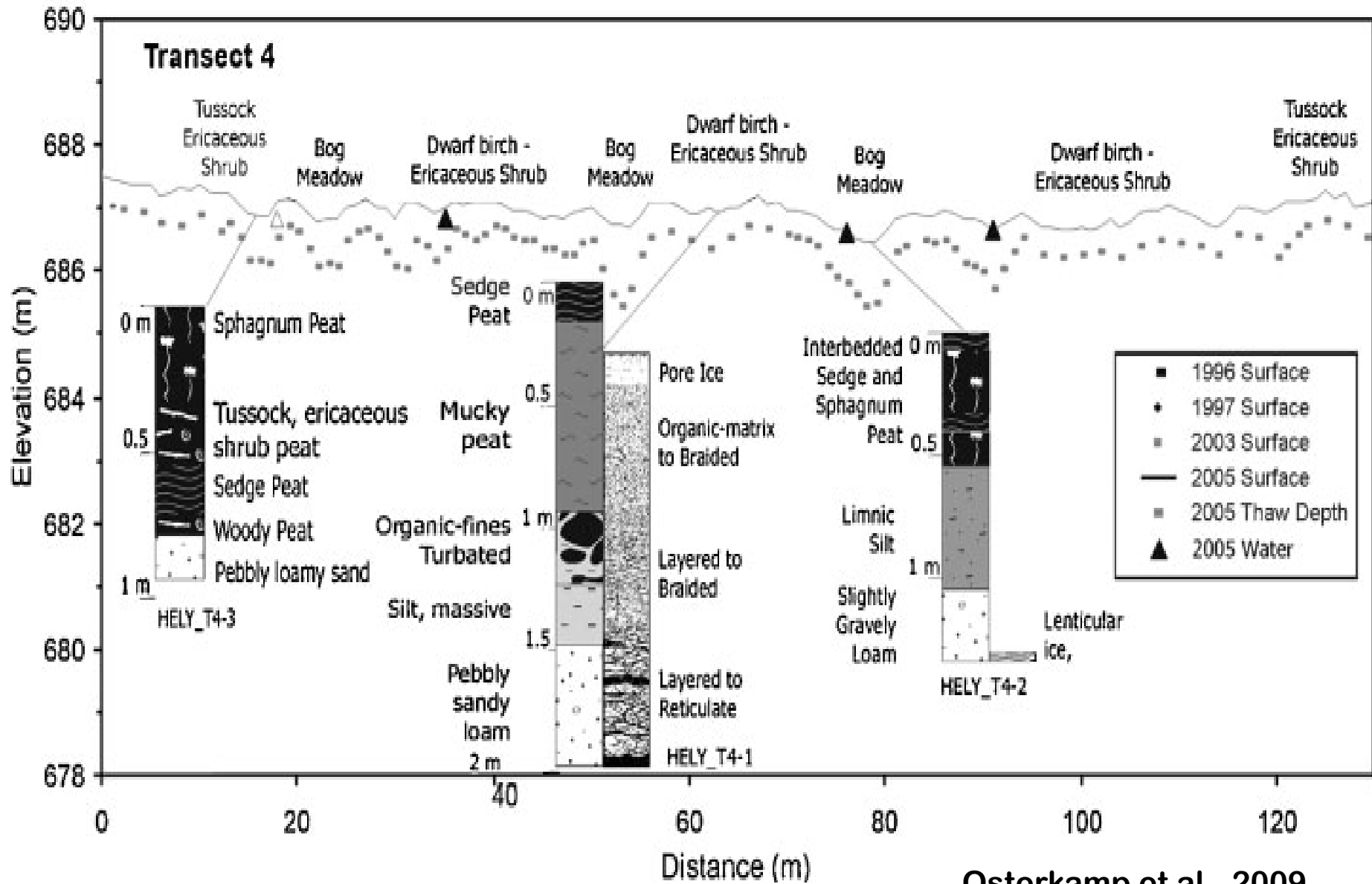
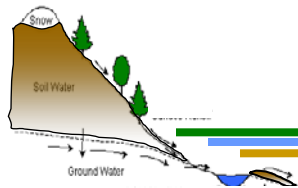
clm3\_6\_45.CN10r - Obs



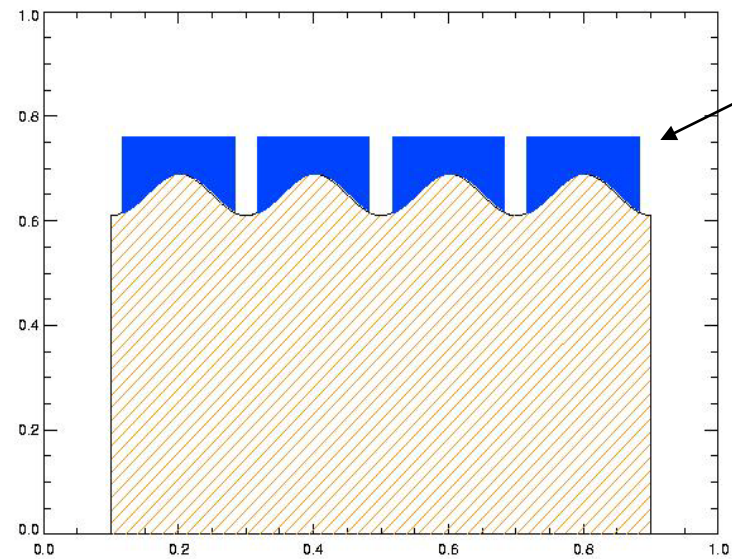
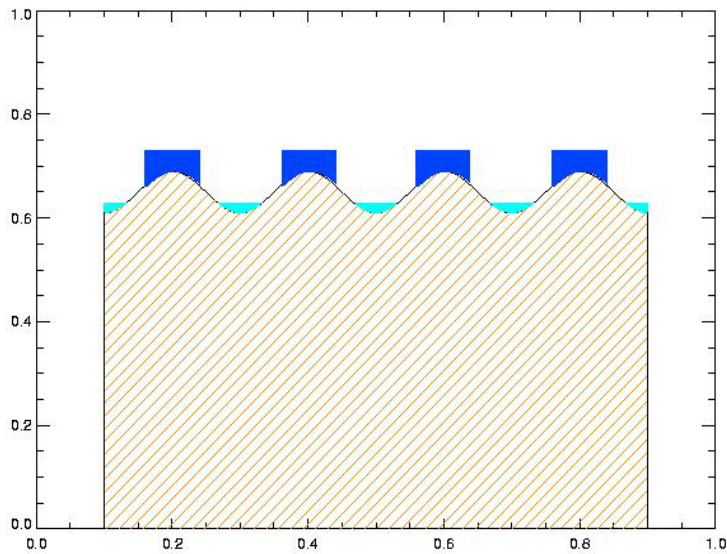
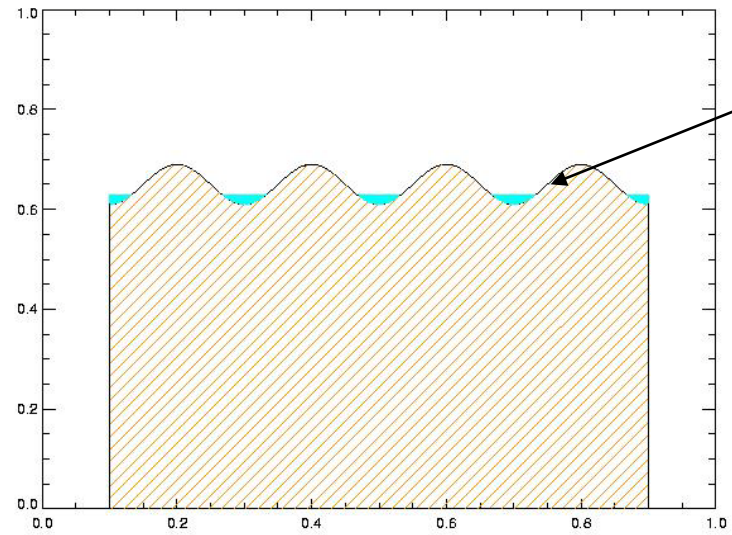
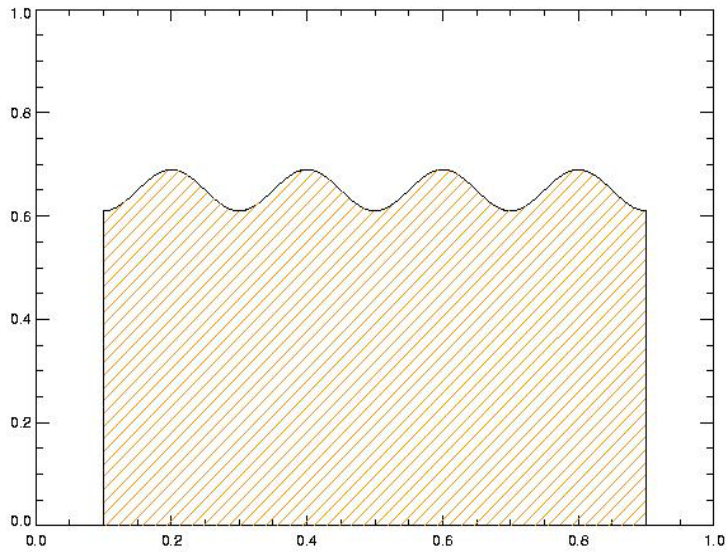
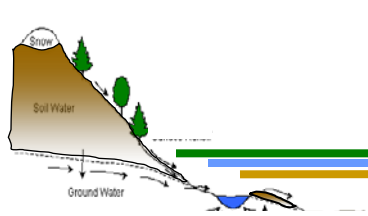
# Summary



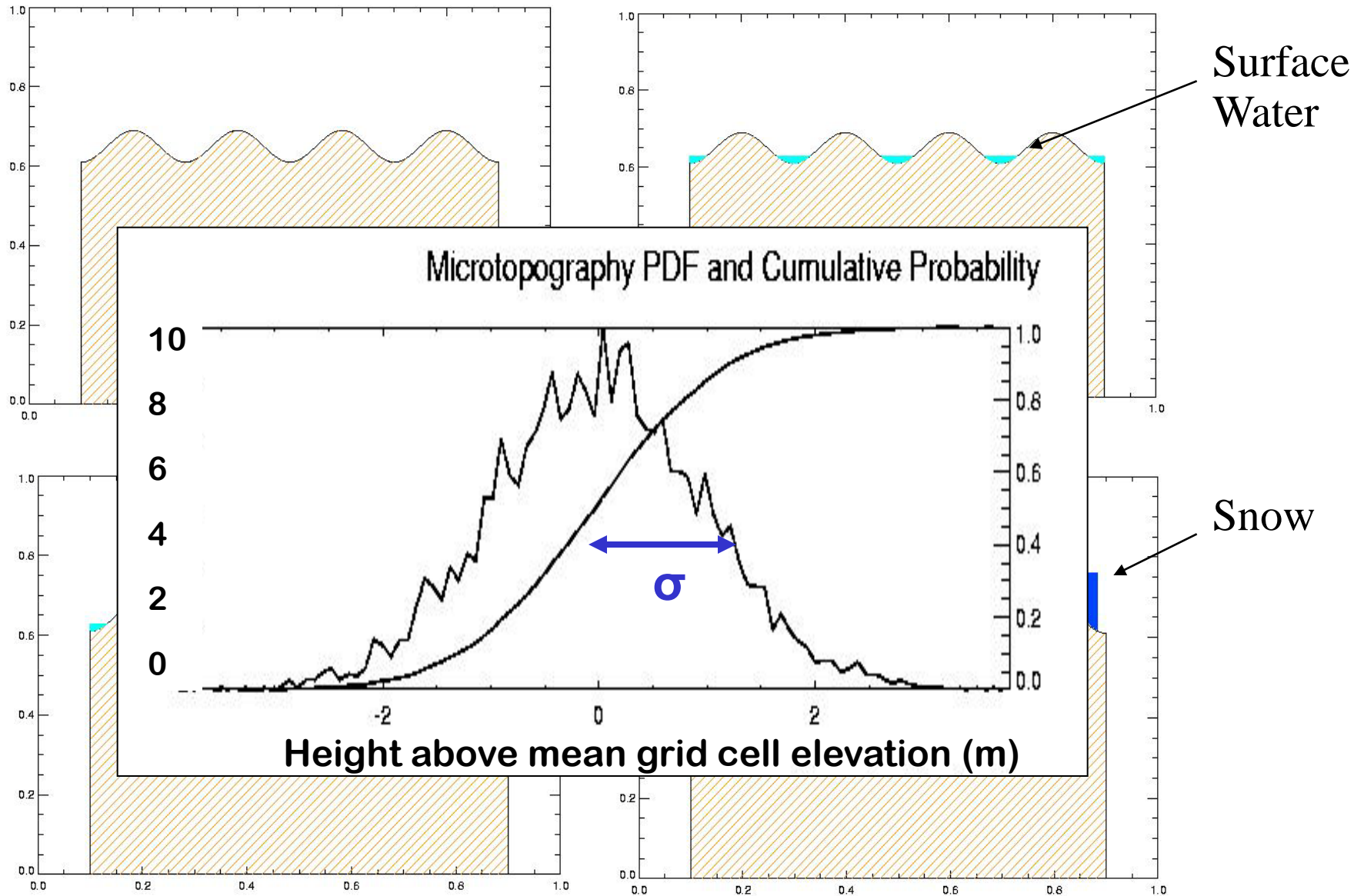
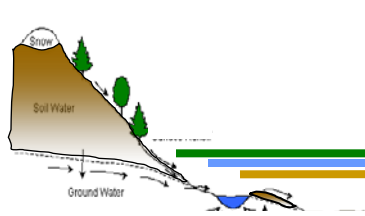
# Microtopography



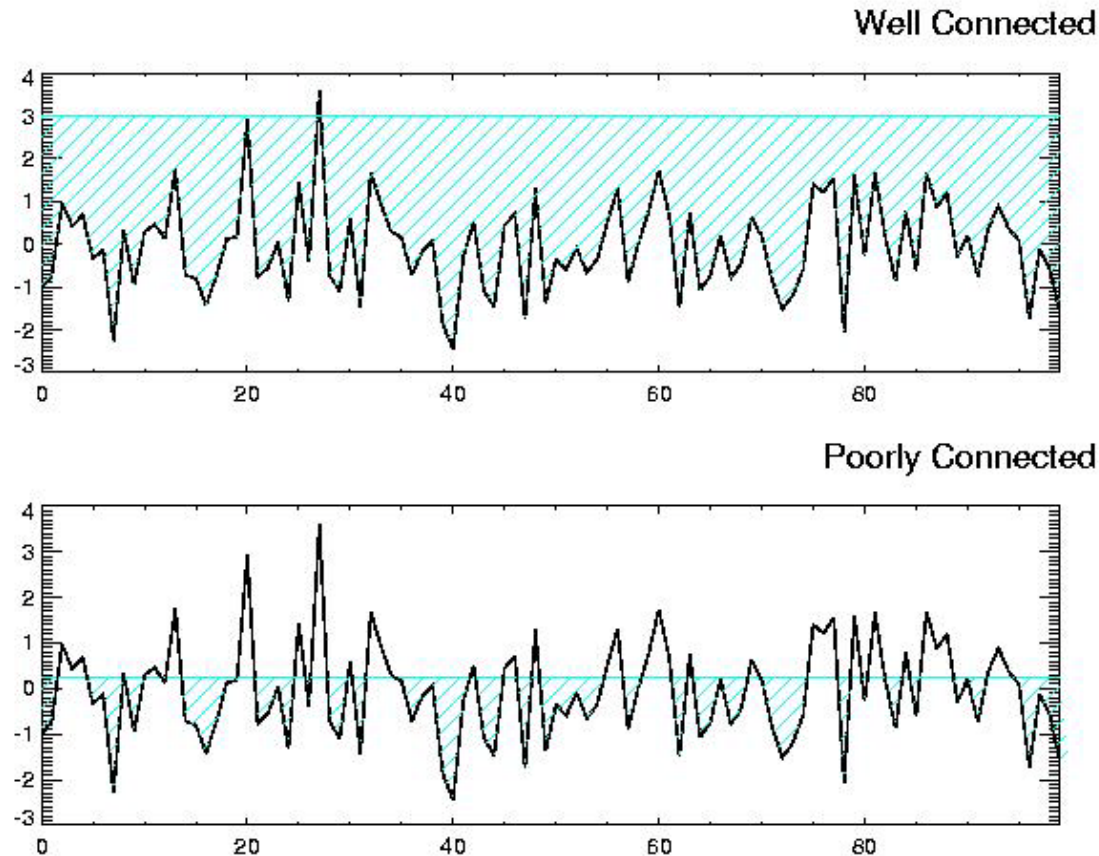
# Surface Water Component Concept



# Surface Water Component Concept



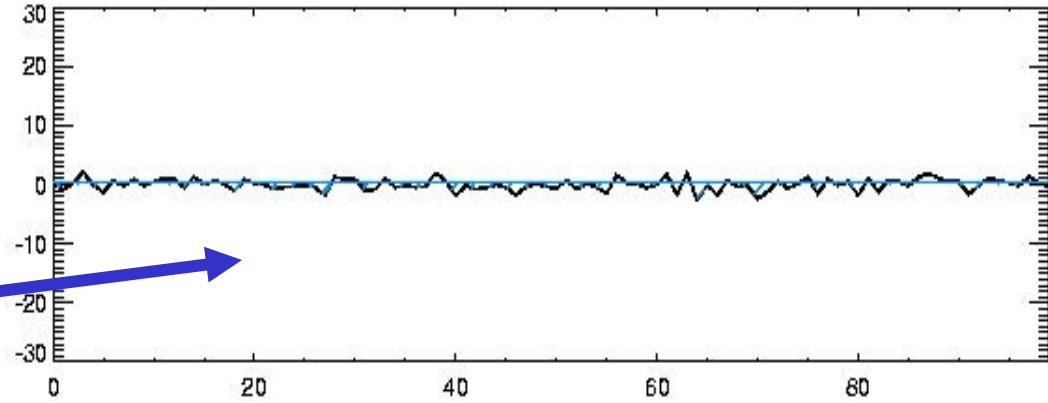
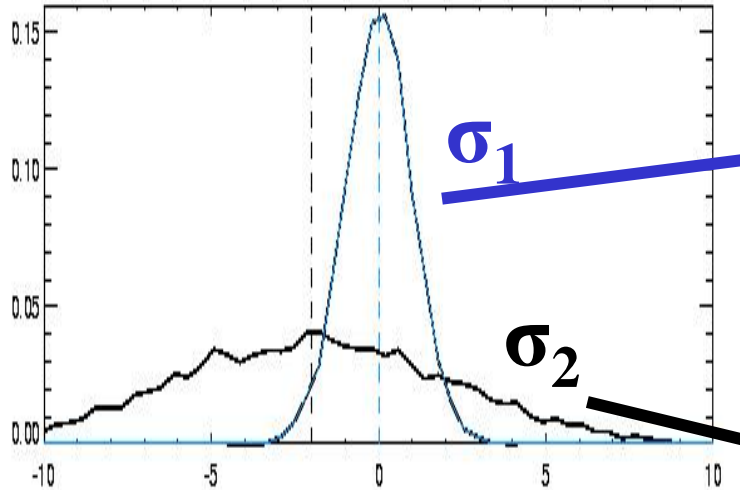
- When storage is large compared to microtopography, “wet” areas are well connected, and surface runoff is high.
- When storage is small compared to microtopography, “wet” areas are generally not connected, and surface runoff is low.



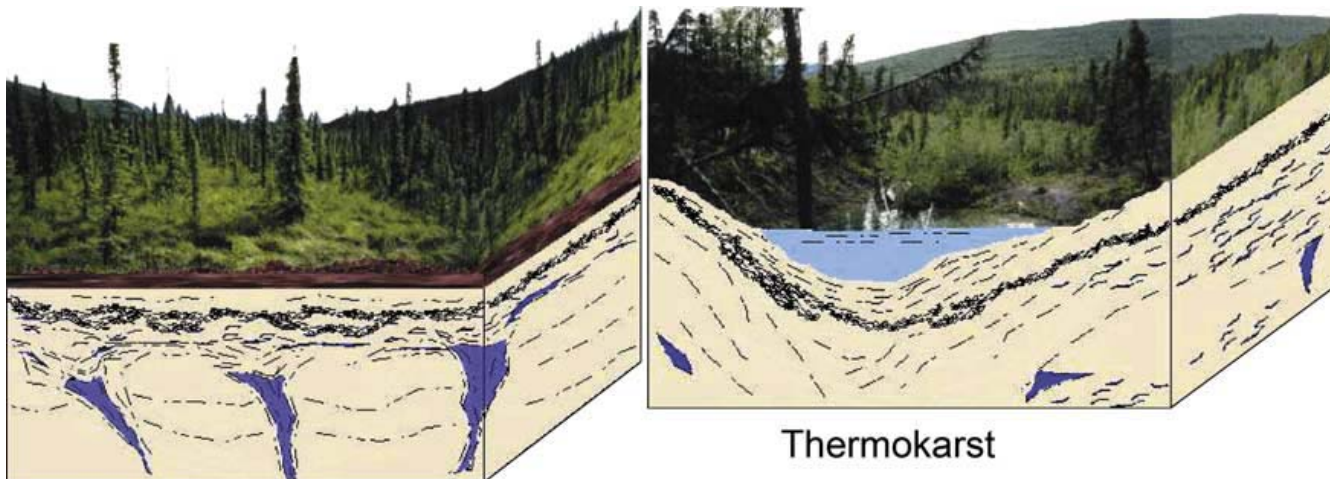
# Parameterizing Thermokarst

Low relief: e.g. Arctic coastal plains

## Microtopography PDF



Height above mean surface (m)



Thermokarst