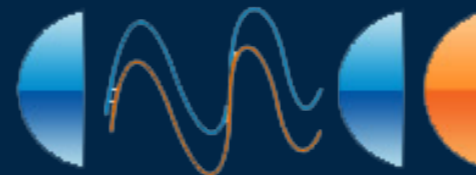


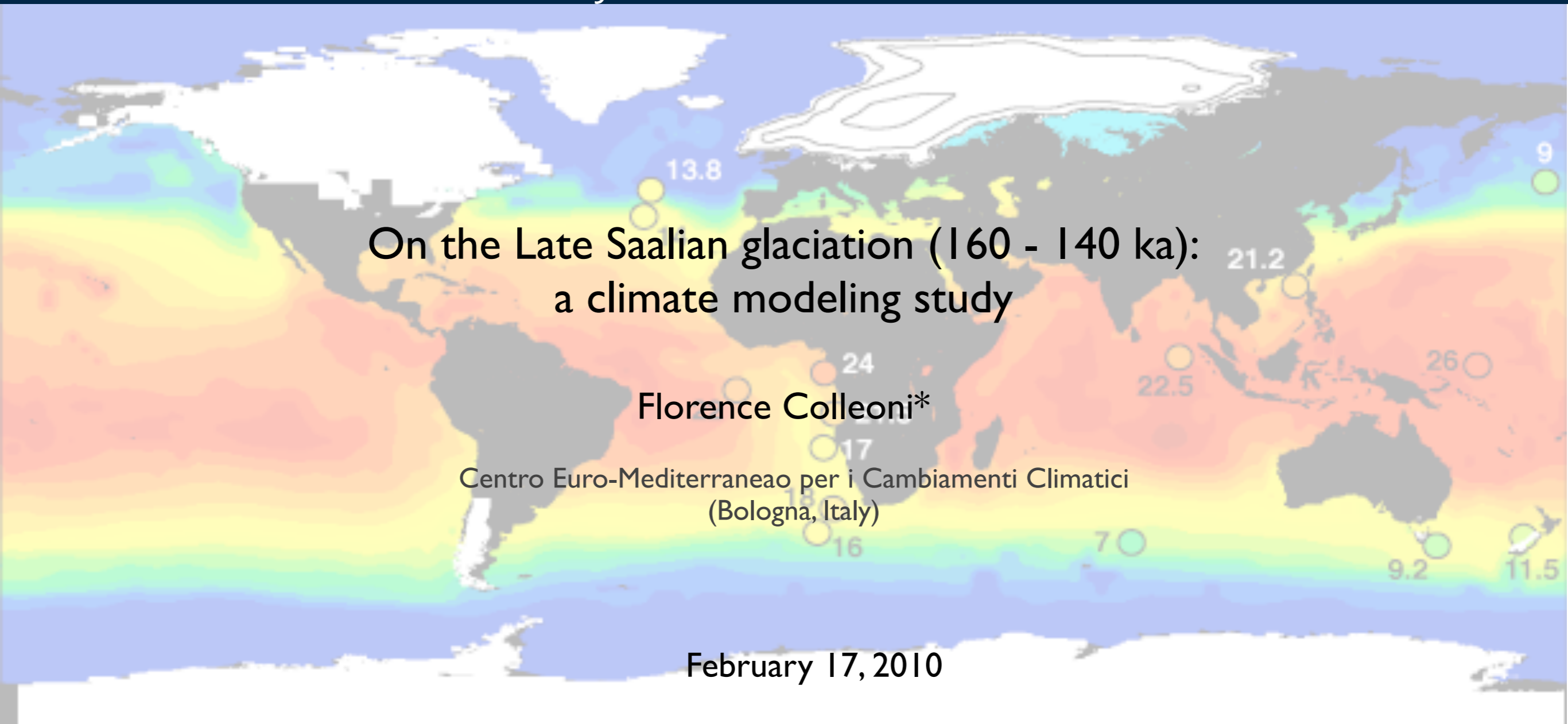
Laboratoire de Glaciologie et Géophysique de l'Environnement



Stockholm University



Centro Euro-Mediterraneo per i Cambiamenti Climatici



On the Late Saalian glaciation (160 - 140 ka): a climate modeling study

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Centro Euro-Mediterraneo per i Cambiamenti Climatici
(Bologna, Italy)

February 17, 2010

Work done at:

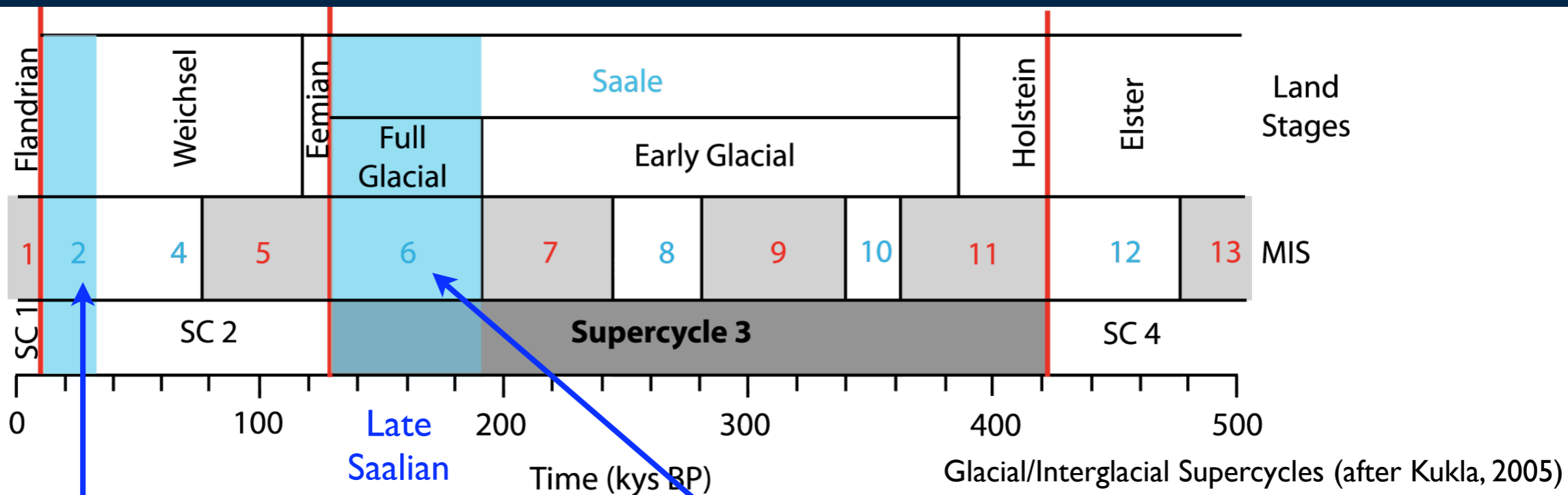
Laboratoire de Glaciologie et Géophysique de l'Environnement, CNRS, UJF, Grenoble (France)

Department of Geology and Geochemistry, Stockholm University (Sweden)

floccolleoni@gmail.com

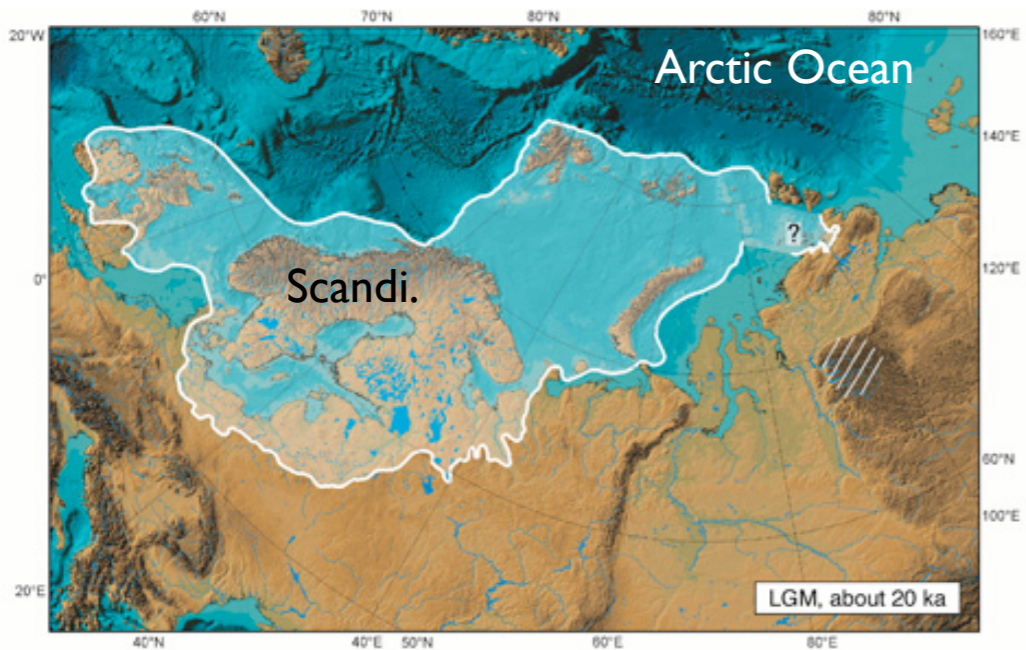


The Quaternary Environment of the Eurasian North Project (1999 - 2003)



Last Glacial Maximum (~ 21 ka)

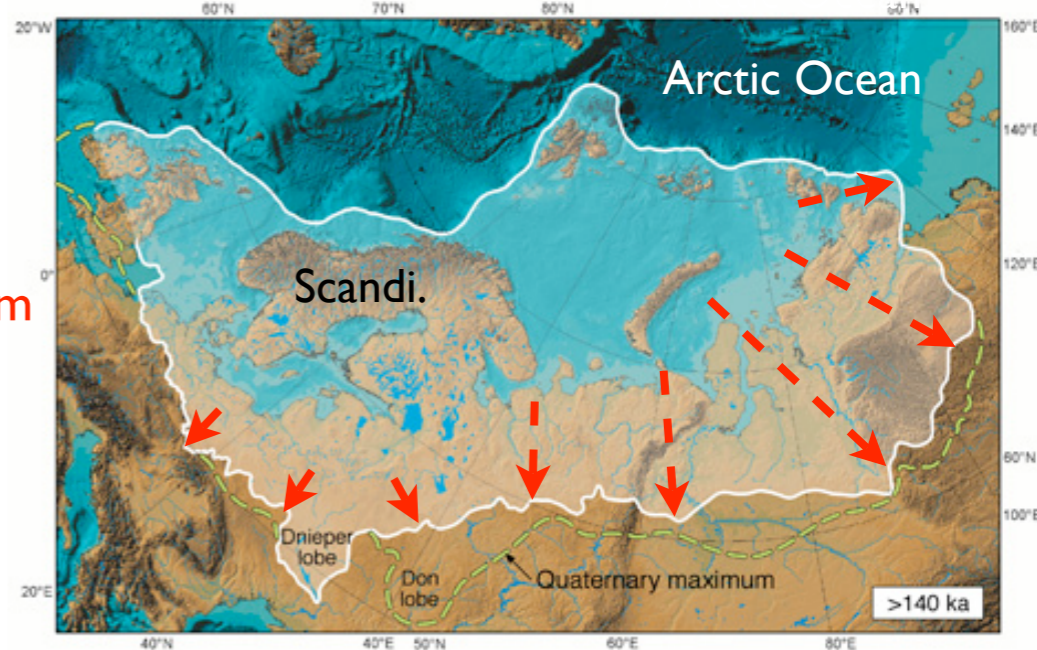
Late Saalian (~ 160 - 140 ka)

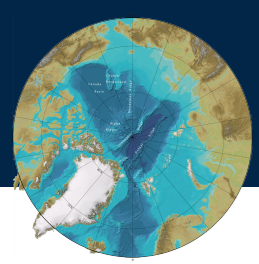


(Svendsen et al., 2004)

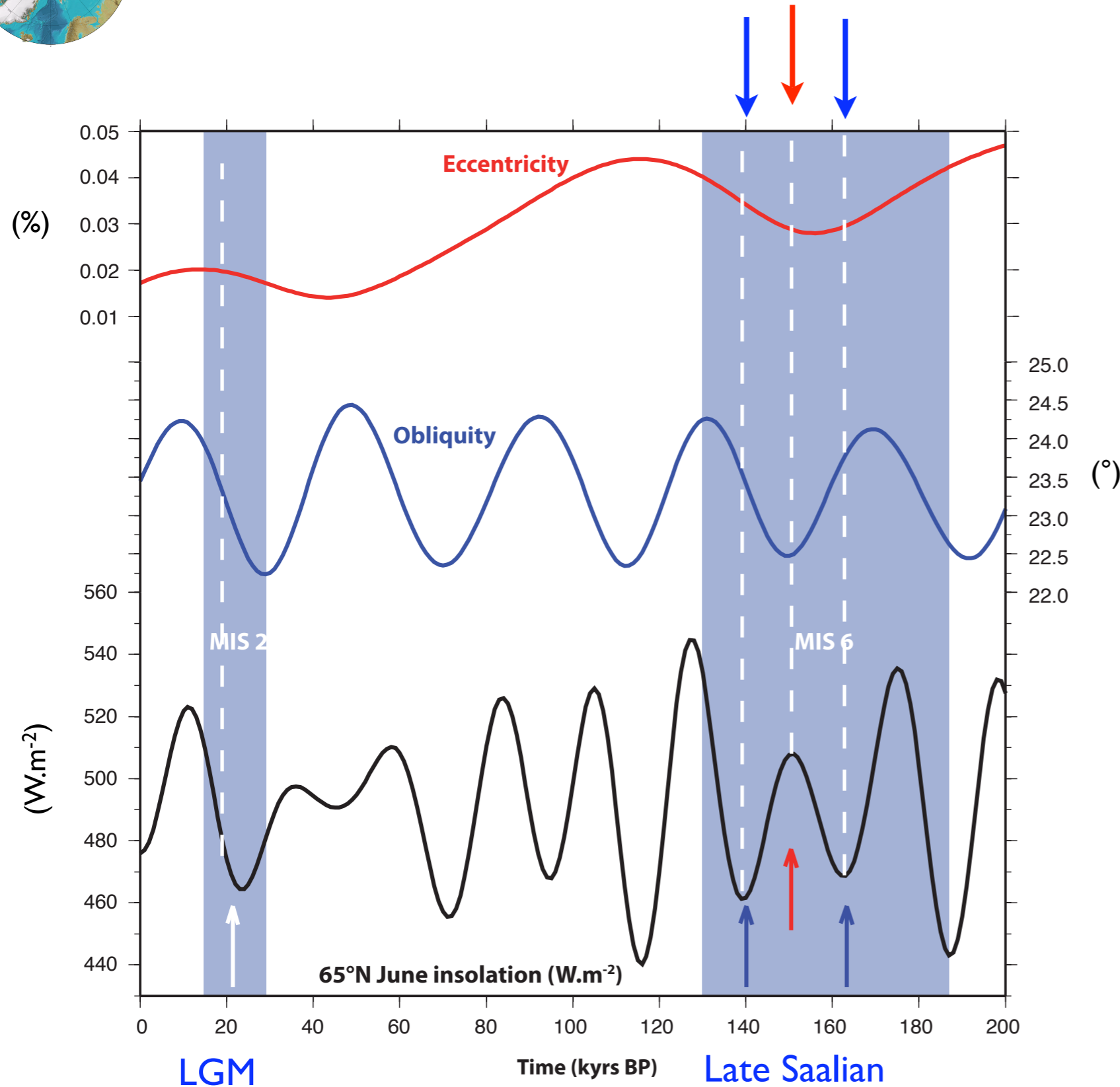
> 2 x larger
 ~ 22 m ESL* ← ~ 60 m ESL

*Equivalent Sea Level





The Late Saalian astronomical forcing



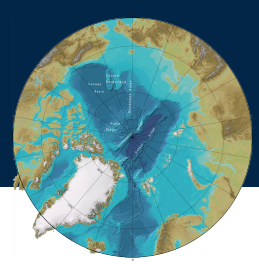
- **140 ka:** highest eccentricity
Perihelion: December 6*
- **150 ka:** lowest eccentricity
Perihelion: June 28
- **160 ka:** largest obliquity
Perihelion January 7

Perihelion = shortest distance to the sun

Late Saalian: larger eccentricity than during the LGM

*LGM: Perihelion January 17

After Berger and Loutre (1991)



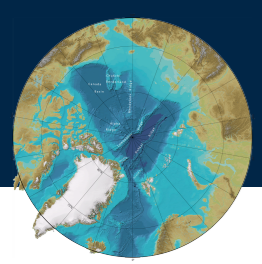
How could the ice sheet survive the insolation peak at ≈ 150 ka?

↳ **Reconstruction of the boundary conditions:**

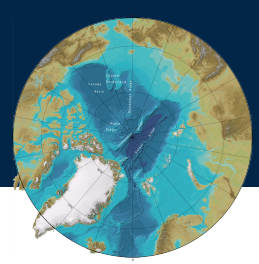
- Global Late Saalian topography
- Dust deposition distribution
- Vegetation cover in equilibrium
- Sea surface temperatures



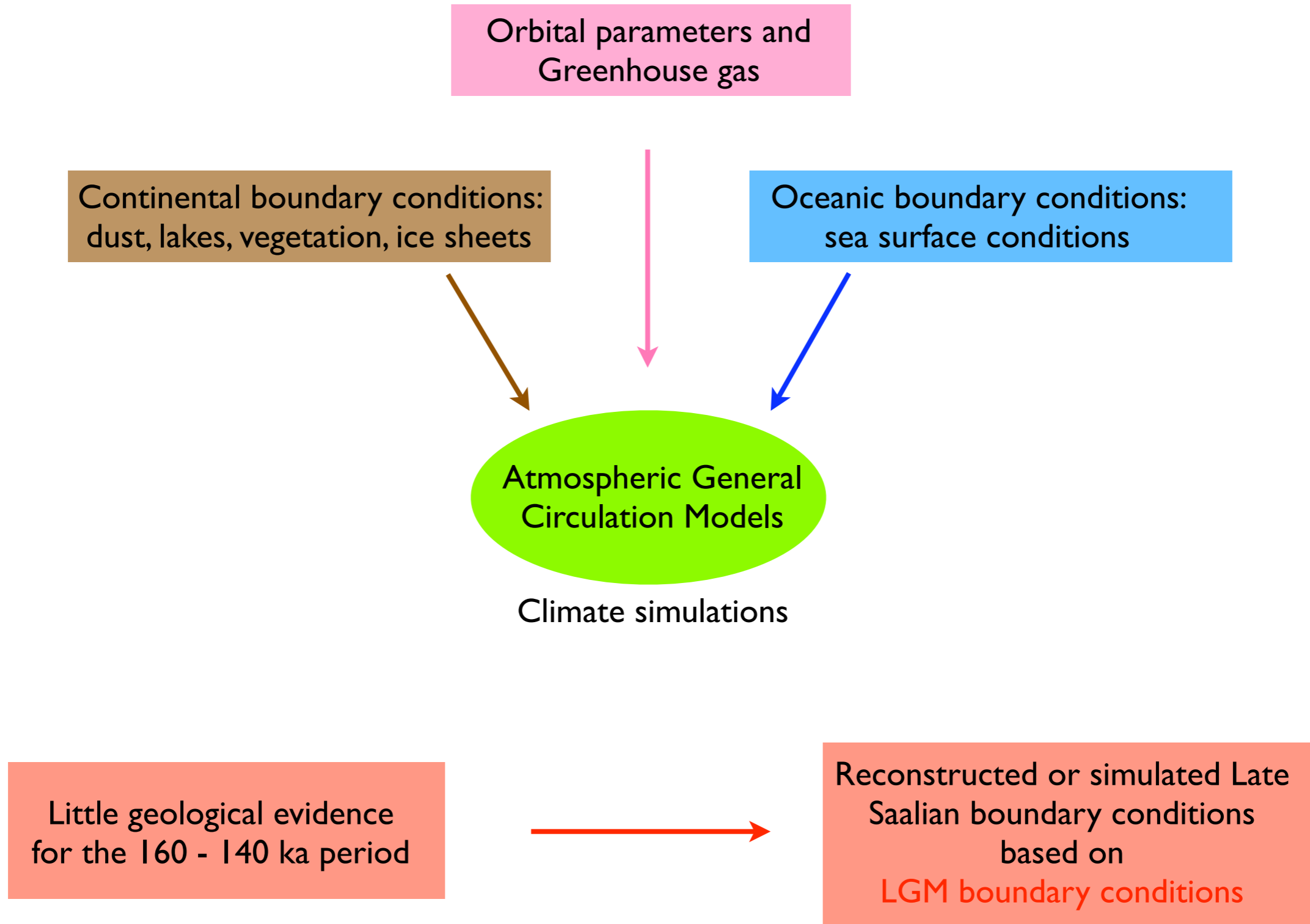
simulating 160 ka, 150 ka and 140 ka climate

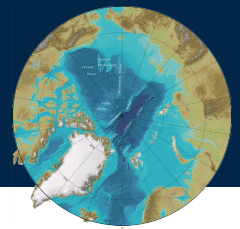


Numerical models and boundary conditions



The Late Saalian boundary conditions





Simulated Late Saalian boundary conditions:

BIOME4 (Kaplan et al. 2004)
 Equilibrium vegetation model
 simulated vegetation cover

Planet simulator (Fraedrich et al., 2005)
 AGCM + oceanic mixed layer
 simulated sea surface conditions

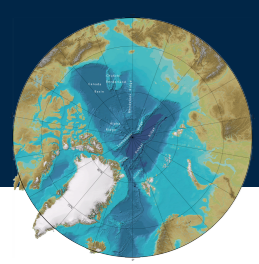
Initialised using:
 LMDZ4 climatology (~21 model yrs)
 forced with a LGM vegetation map
 (Crowley, 1995)
 Iterative scheme until equilibrium

Horizontal resolution: T42 (128*64)
 Initialized using:
 LGM monthly SST from Paul & Schaefer-Net (2003)
 Return SST through heat fluxes calculation
 Length of simulation: 50 model years



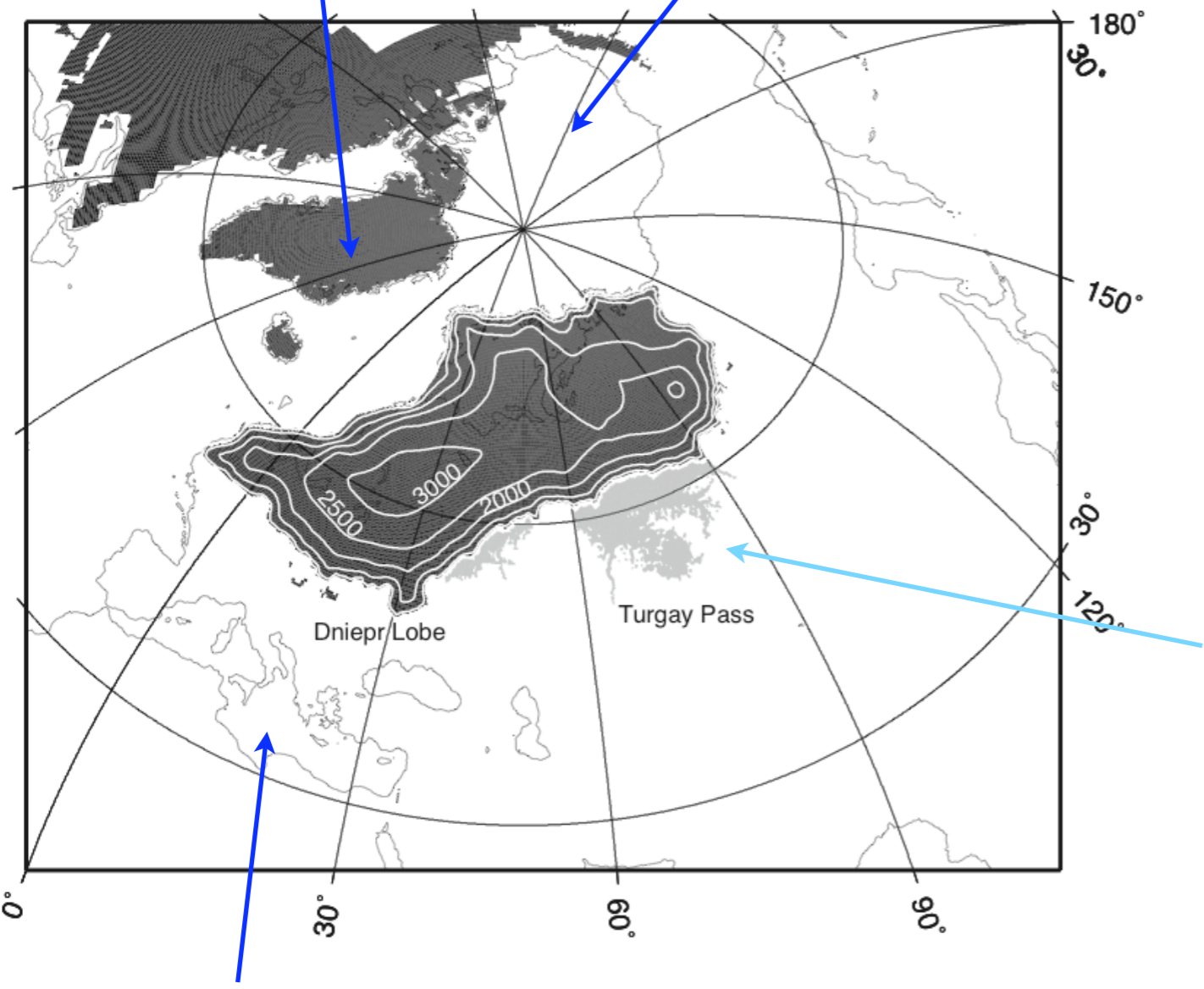
Horizontal resolution: 96 x 72
 Vertical resolution: 19 layers

Zoom over Eurasia: (grid space ~ 100 km)
 Length of simulations: 21-year snap-shots



The Late Saalian Northern Hemisphere topography

Greenland Arctic Ocean



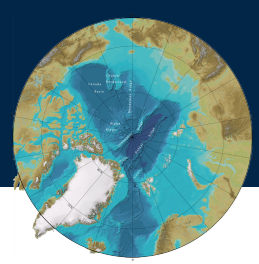
Eurasian ice sheet: 3000 m high reconstruction using GRISLI ice sheet model (Peyaud, 2006; Ritz et al. 1991)

Laurentide and Antarctica ice sheets: LGM ICE-5G (Peltier, 2004)

→ problem for the global ESL: excess of mass

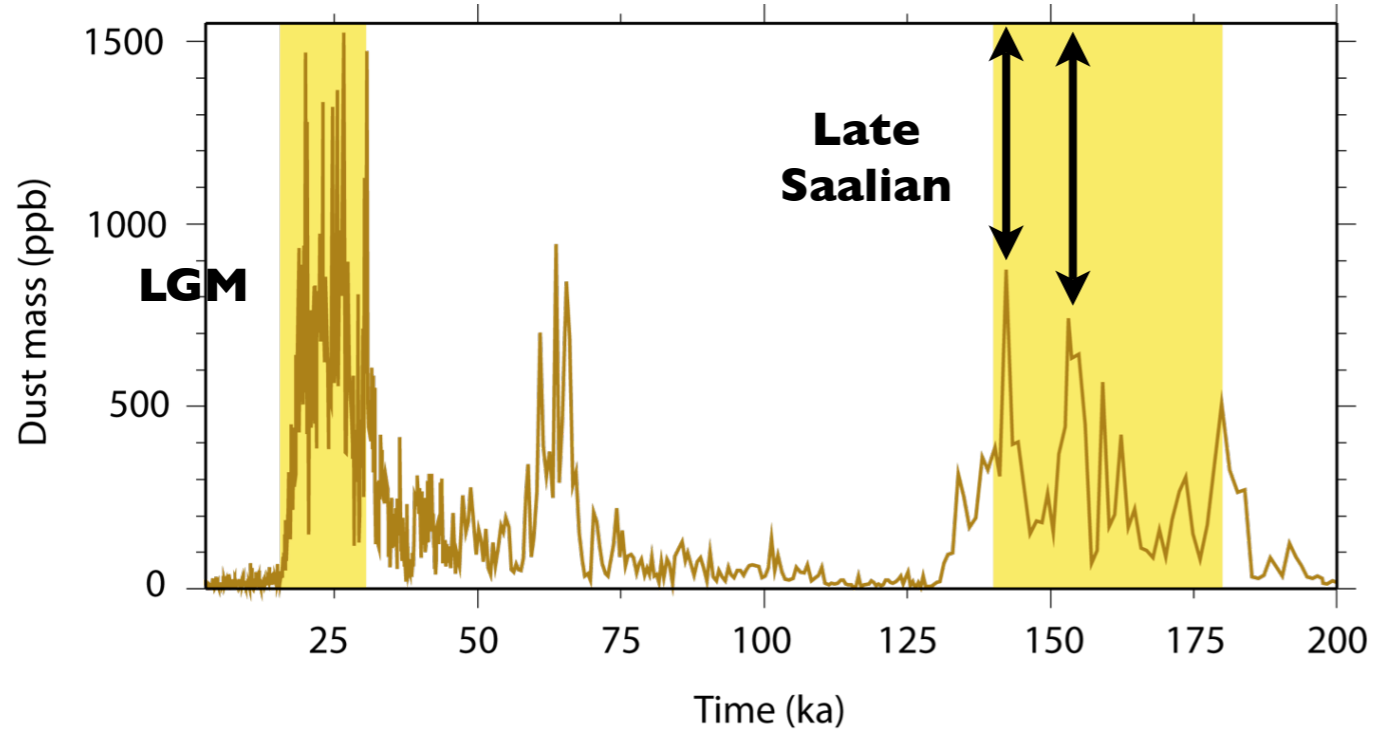
Proglacial lakes: after Mangerud et al. (2001)
- developed in the topographic basin in front of the ice sheet
- damming of the rivers network

Mediterranean



The Late Saalian dust distribution

Epica Dome C, East Antarctica (Delmonte et al., 2004)



Dust concentration: comparison with LGM:

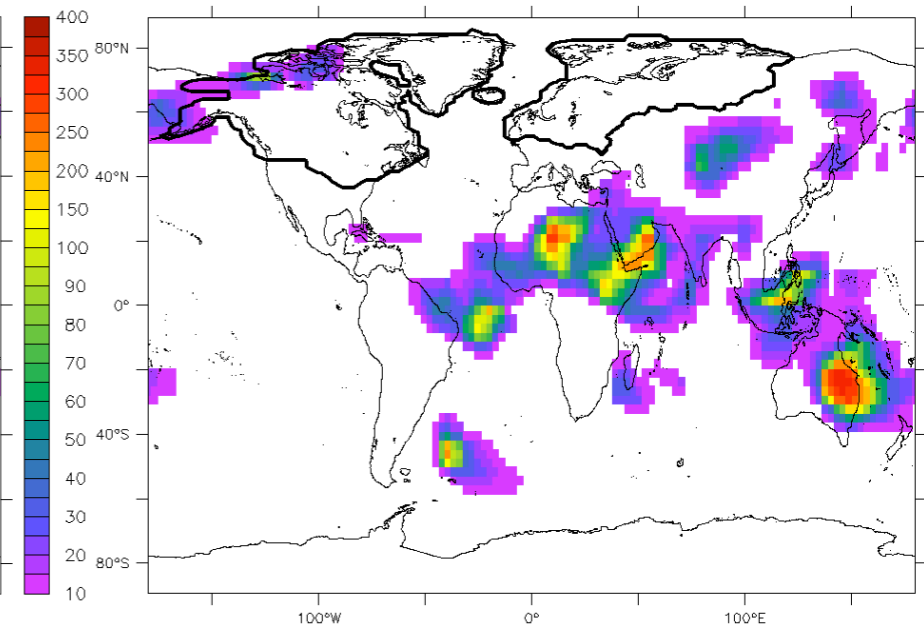
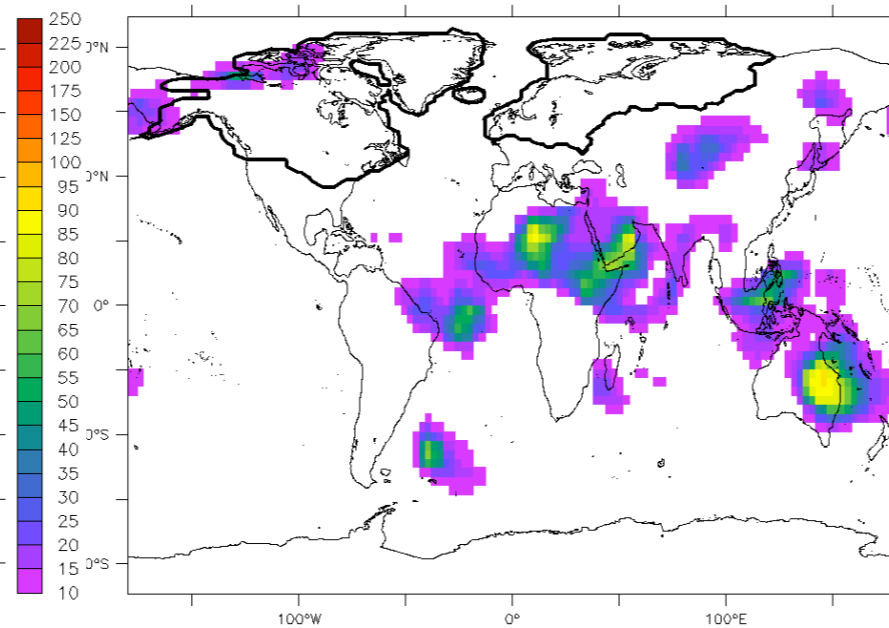
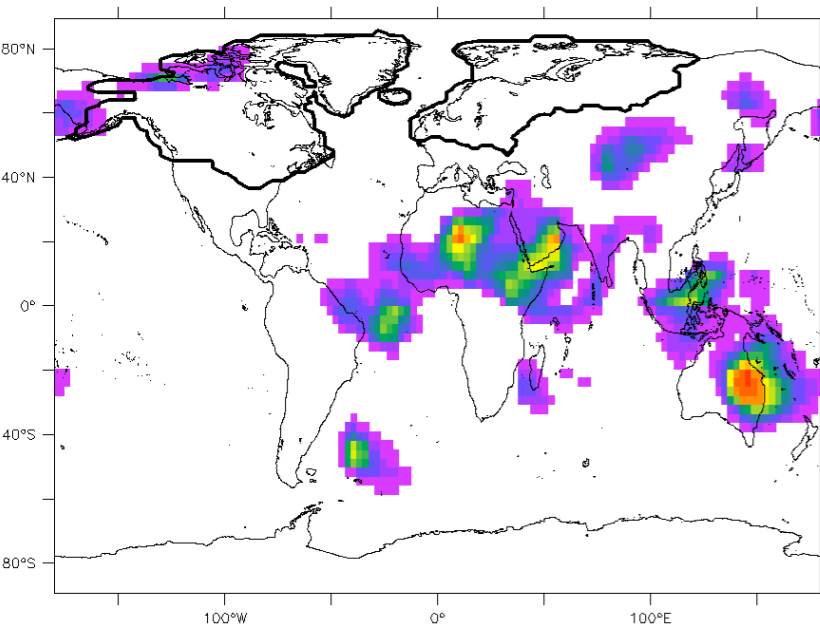
- 140 ka: 40% lower than LGM
- 150 ka: 80% lower than LGM
- 160 ka: 50% lower than LGM

In agreement with the Siberian loess stratigraphy (Chalchula, 2003)

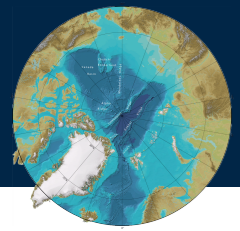
160 ka

150 ka

140 ka

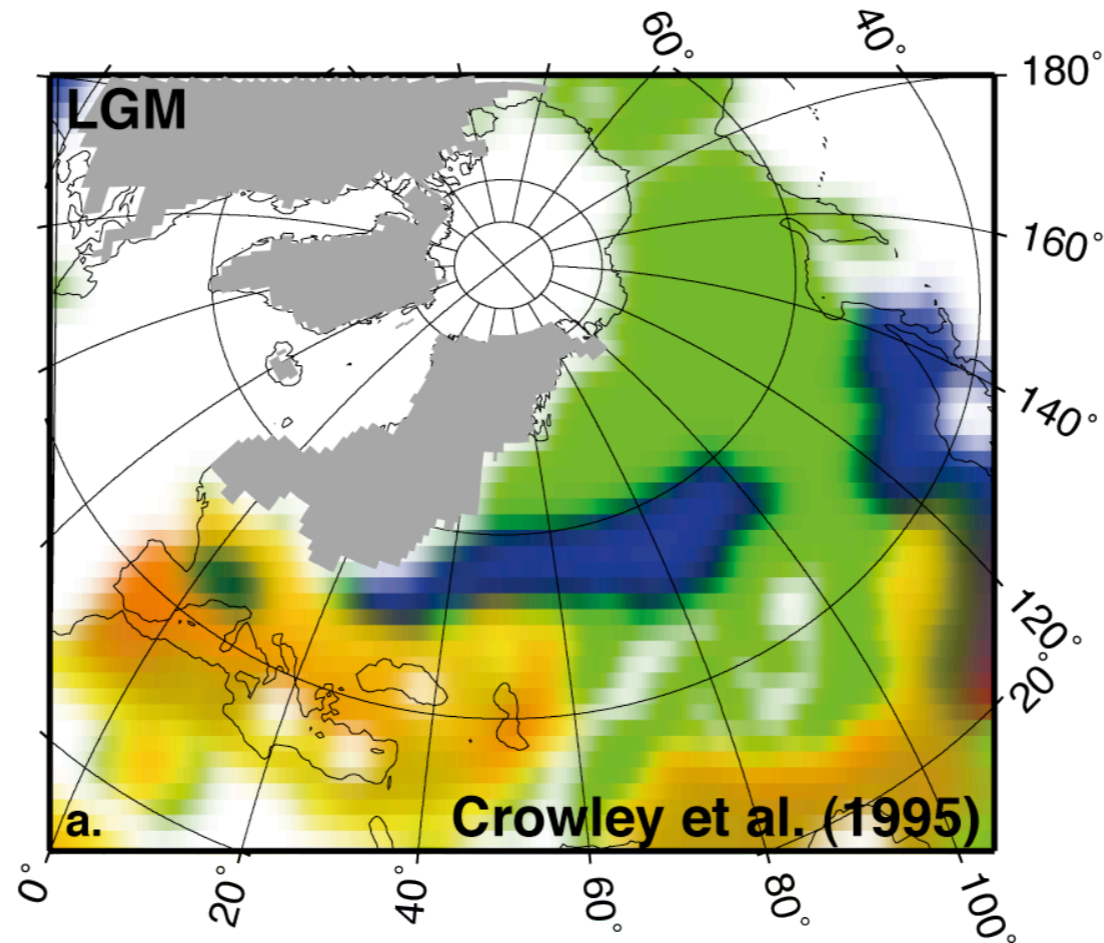


Dust concentration g/m²/yr

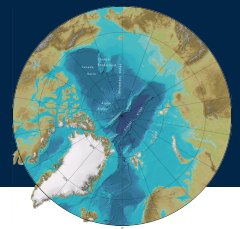


Vegetation cover: Crowley LGM vegetation map (1995)

Reference vegetation map used in all the Late Saalian reference simulations

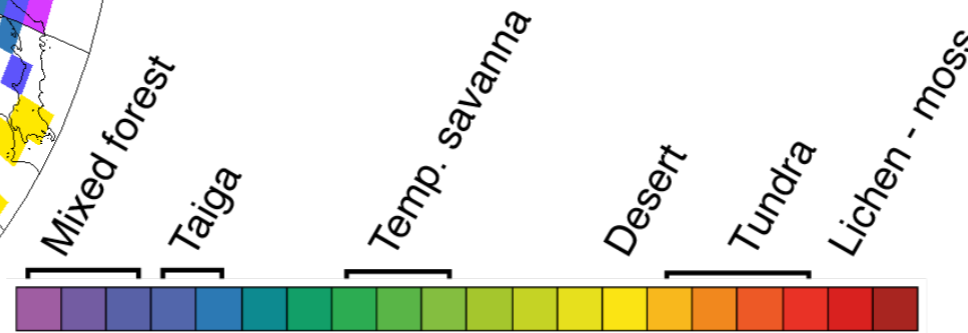
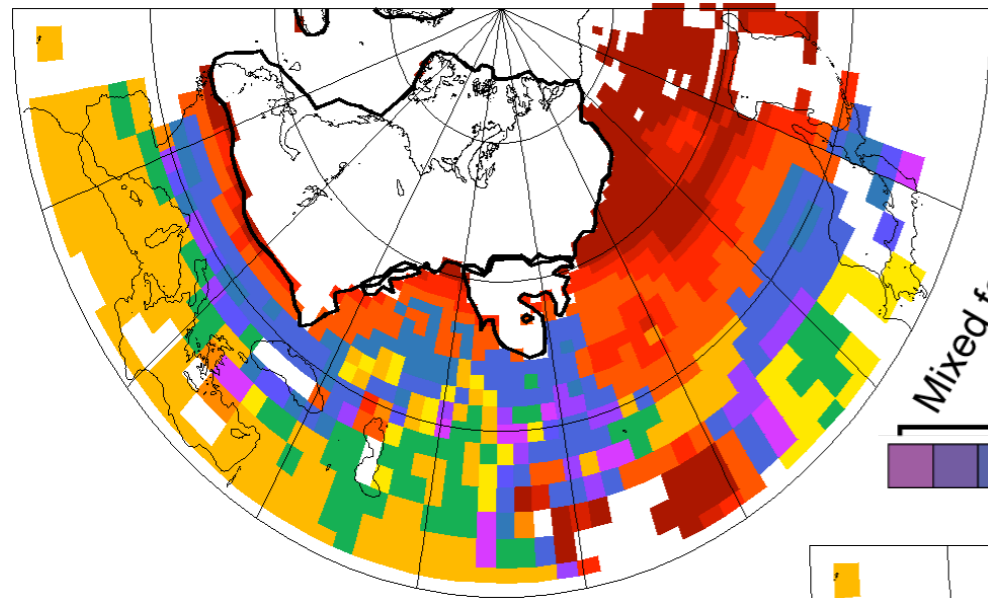


-  Needle-leaf trees
-  Broad-leaved trees
-  Grass land
-  Mixed forest



Vegetation cover: using BIOME 4

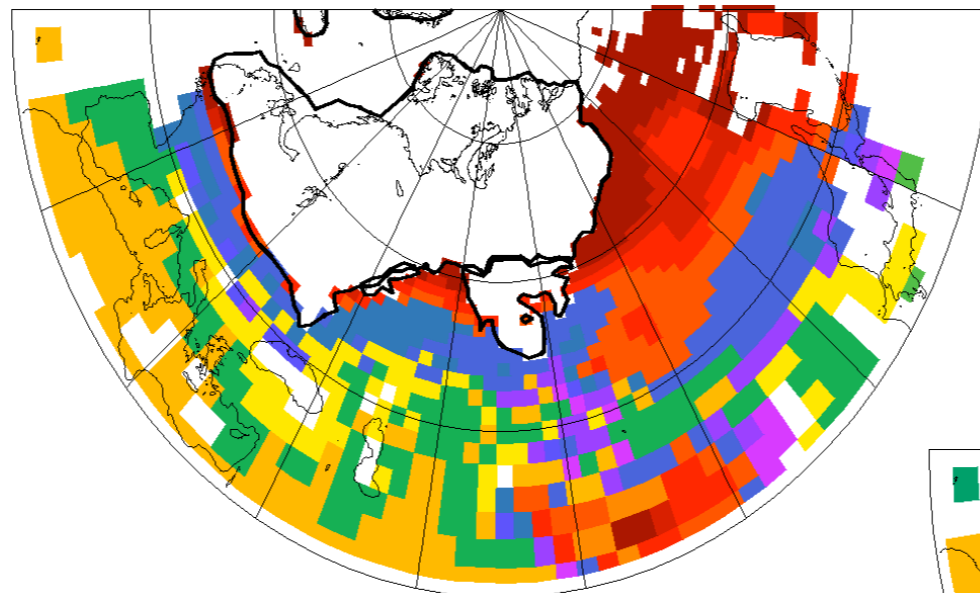
160 ka



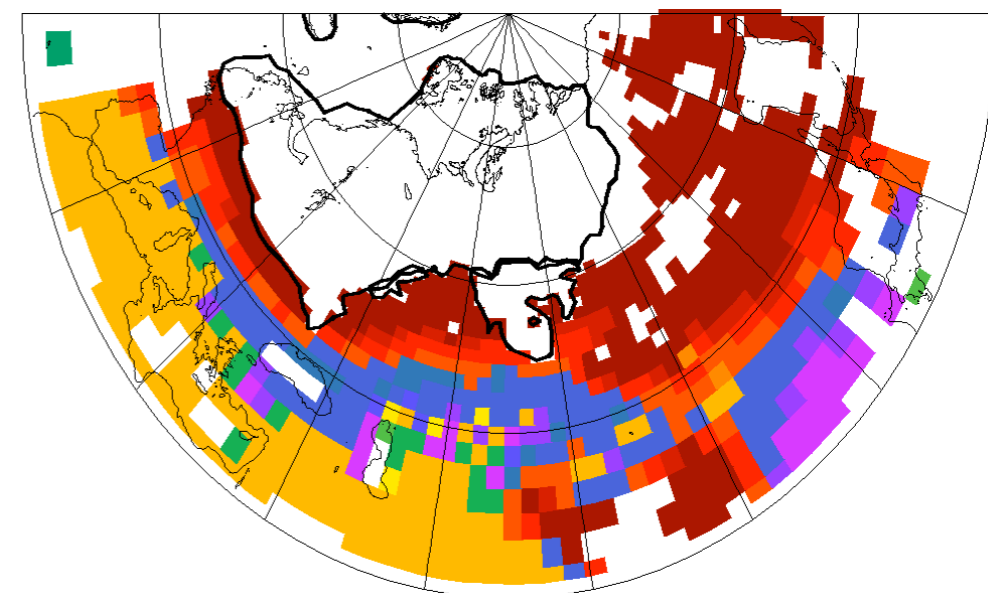
BIOME 4 /
LMDZ4
iterations
x 3

21-year
snap-shots

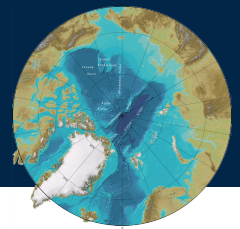
150 ka



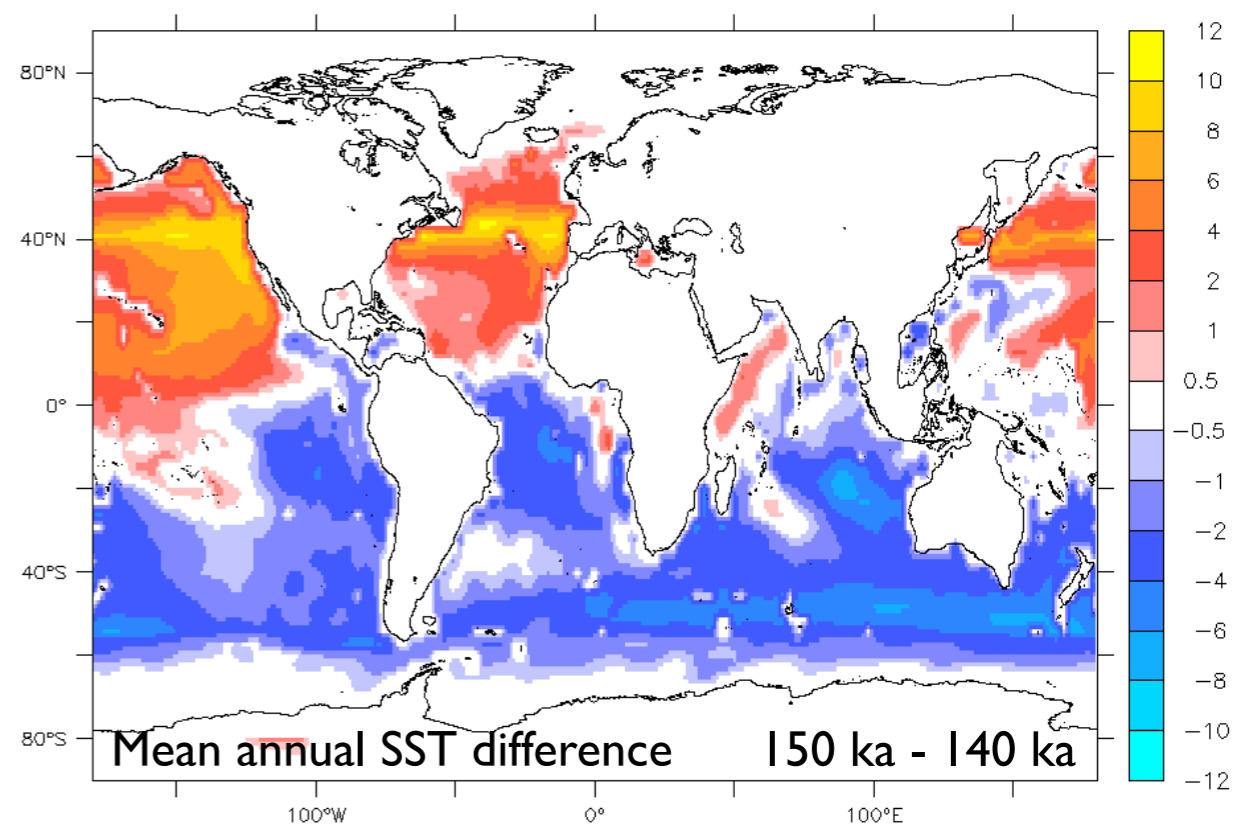
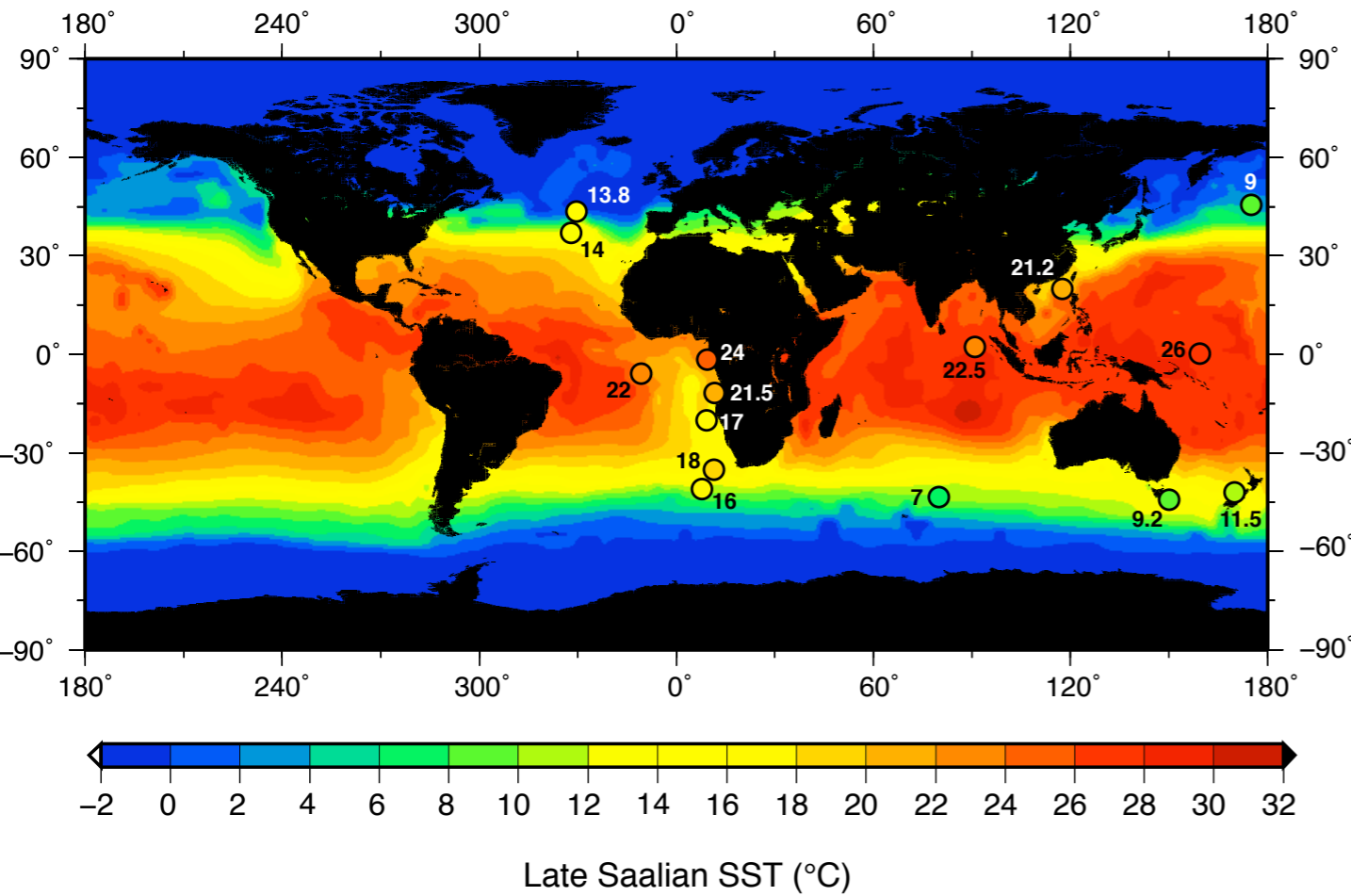
140 ka



- 160 ka: large tundra extent
- 150 ka: tundra replaced by mixed forests
- 140 ka: progression of barren soils: glacial maximum



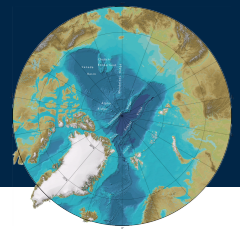
Late Saalian sea surface conditions



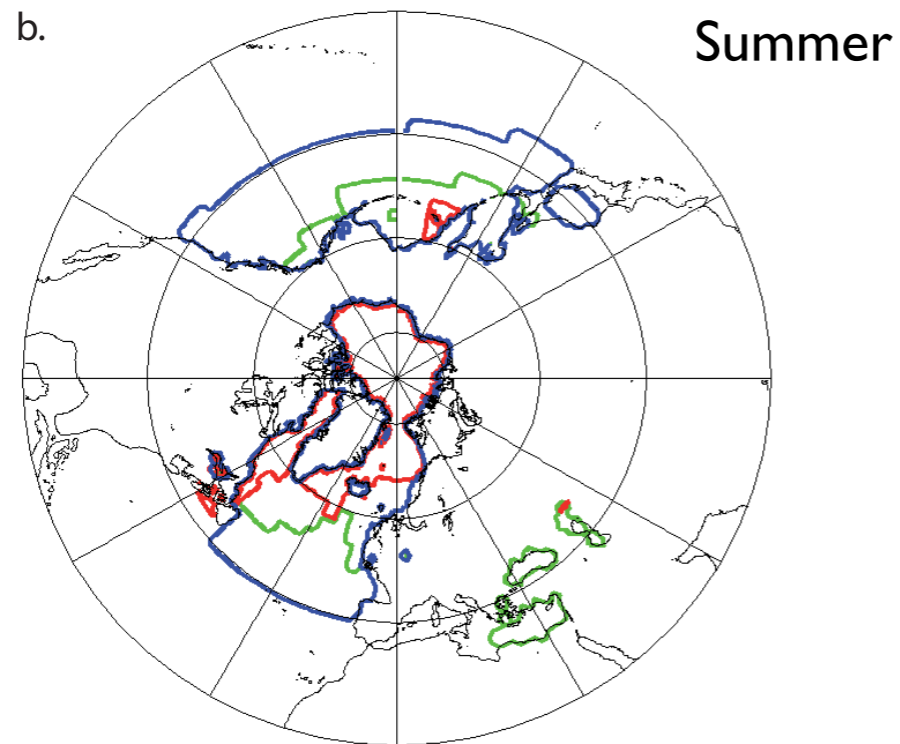
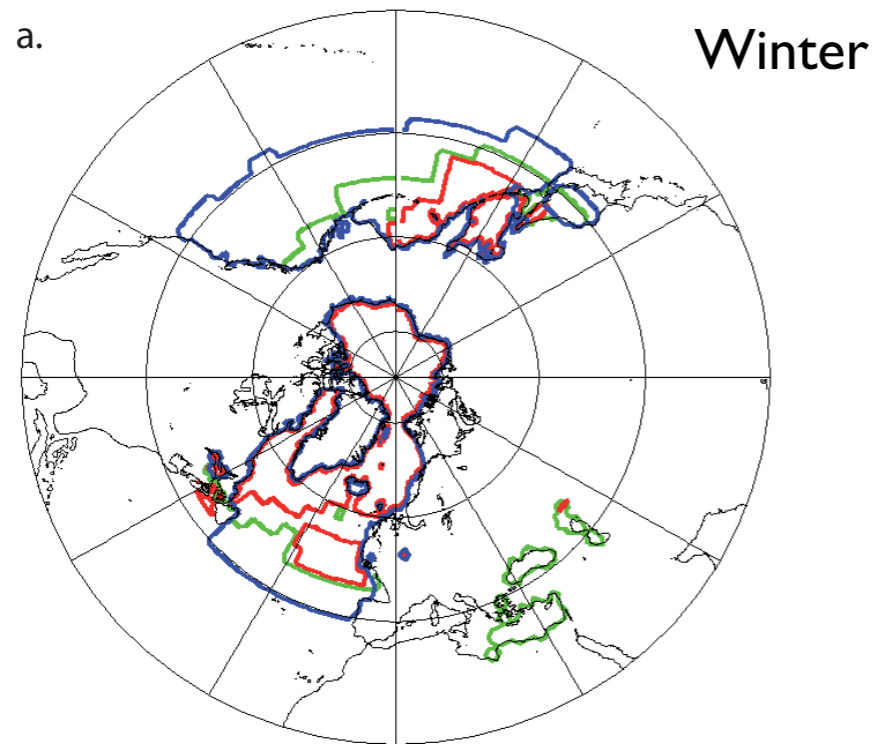
- Equatorial / Tropical SST: **good agreement with data**
- High latitudes Atlantic SST: **underestimate data**
- Large SIC extent in NH high latitudes: **~ 40°N**

- Asymmetry due to 140 ka SST large sea-ice cover: albedo feedback
- At 160 ka: similar to 150 ka (slightly warmer falls at 150 ka)

(Colleoni et al., sub.)



Sea ice cover of the Late Saalian period (160 -140 ka)



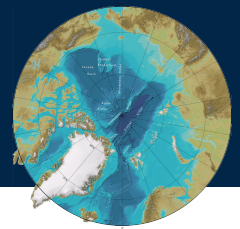
140 ka

150 ka

160 ka

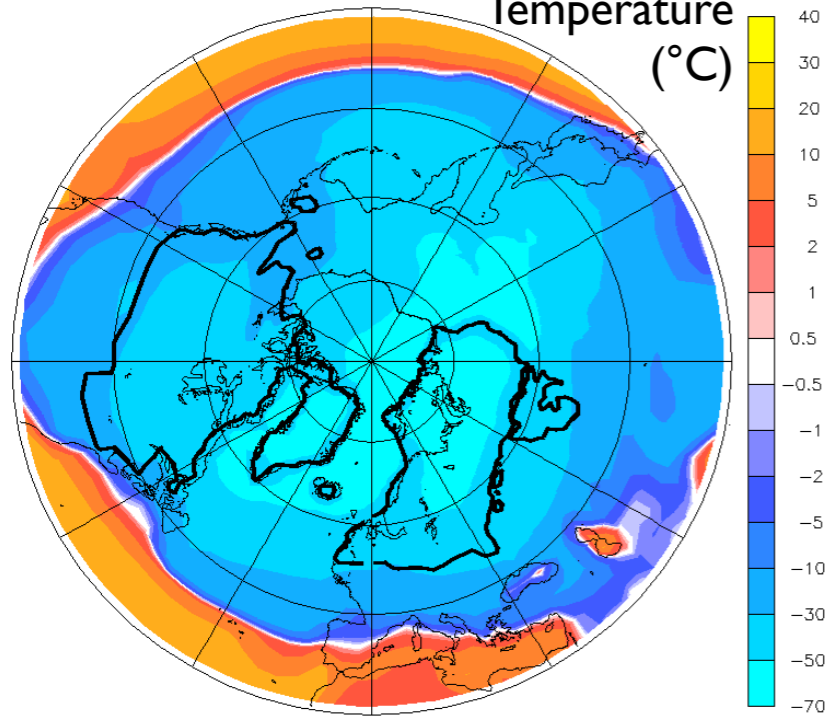
Northern Hemisphere Sea-ice Cover:

- **Largest extent:** 140 ka
- **Smallest extent:** 150 ka more or less similar to present-day extent
- **Intermediate extent:** 160 ka almost similar to CLIMAP LGM reconstructions

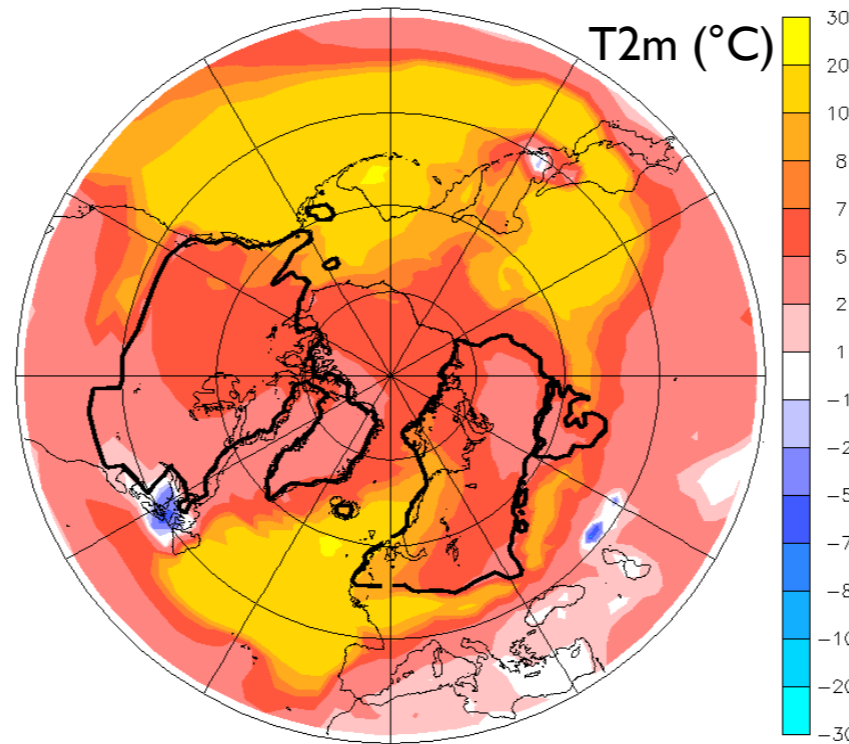


The Late Saalian mean annual climatology (160 ka to 140 ka)

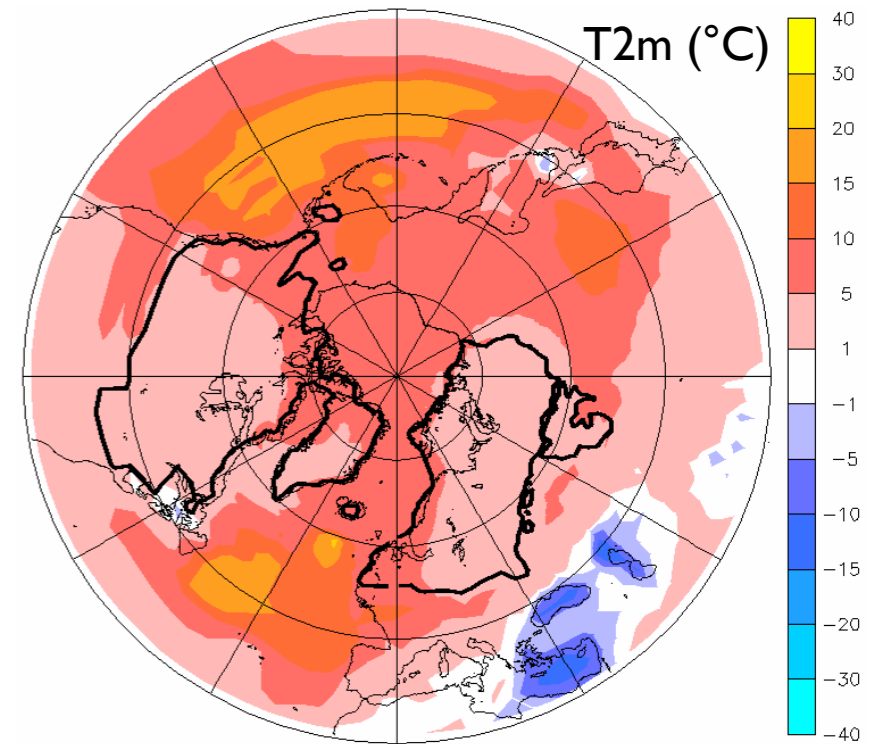
140 ka



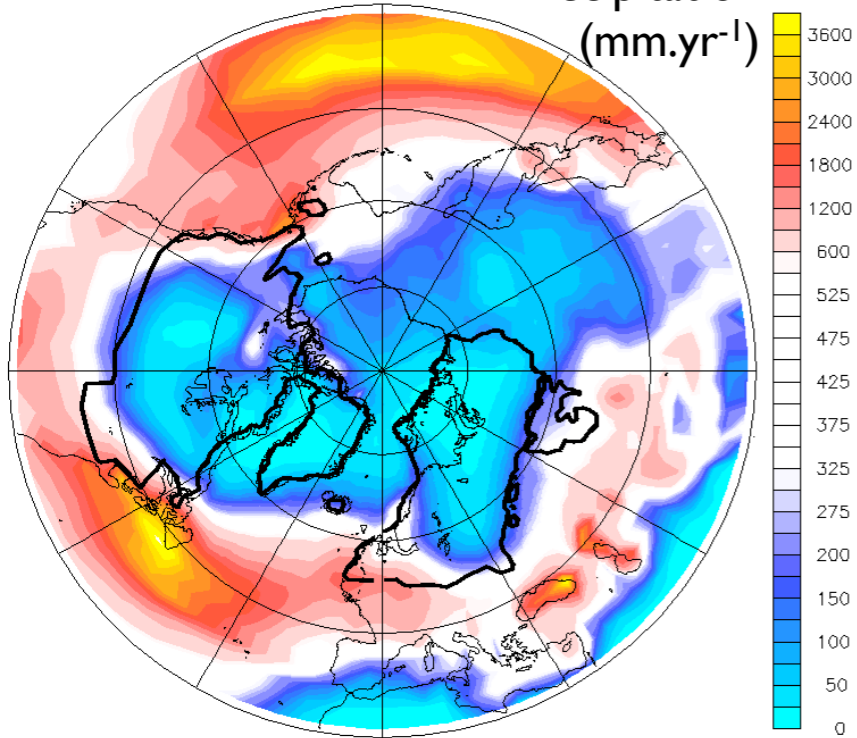
150 - 140 ka



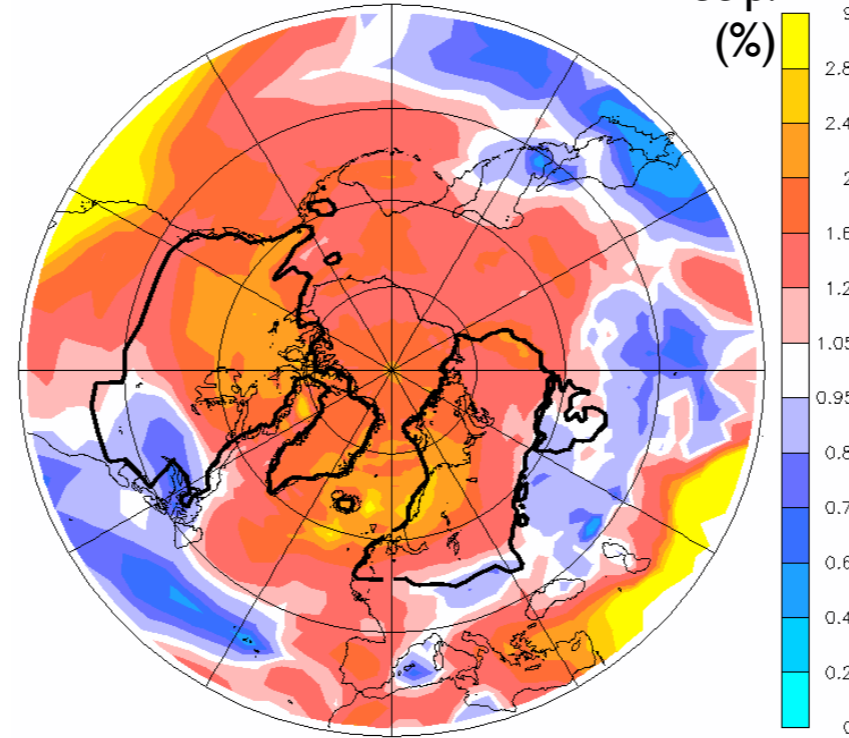
160 - 140 ka



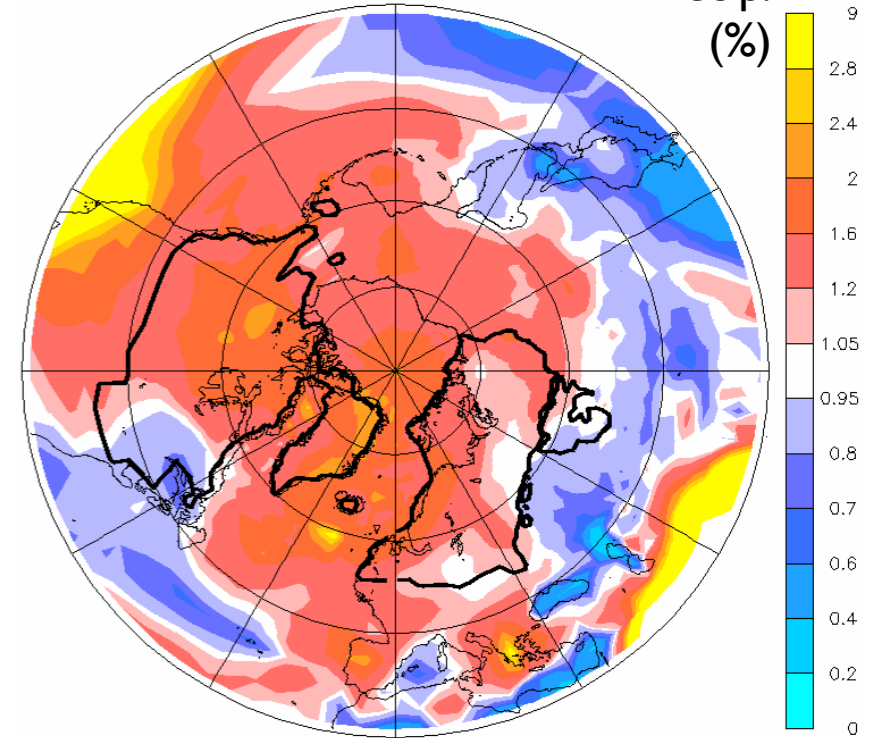
Precipitation (mm.yr⁻¹)

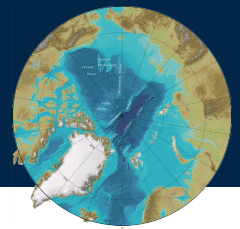


Precip. (%)

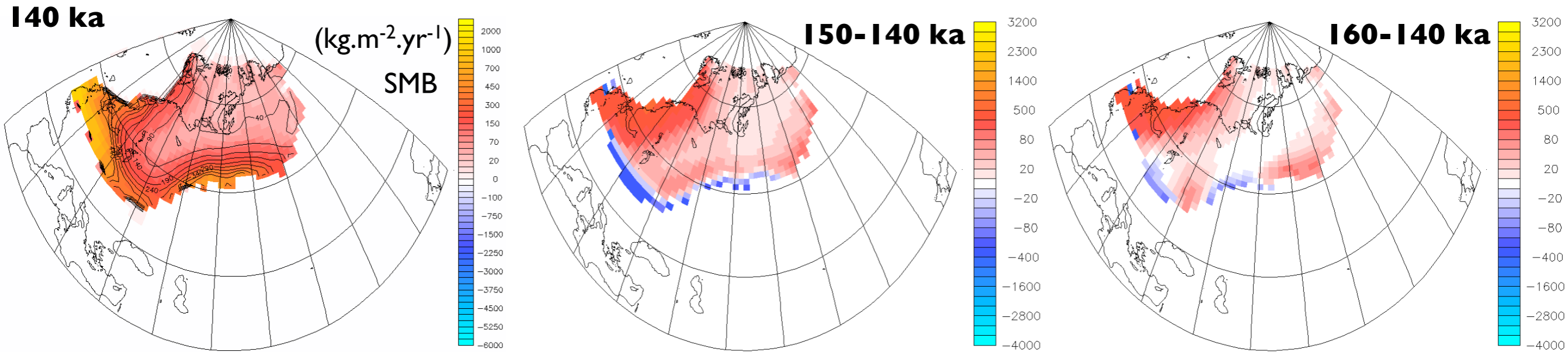


Precip. (%)

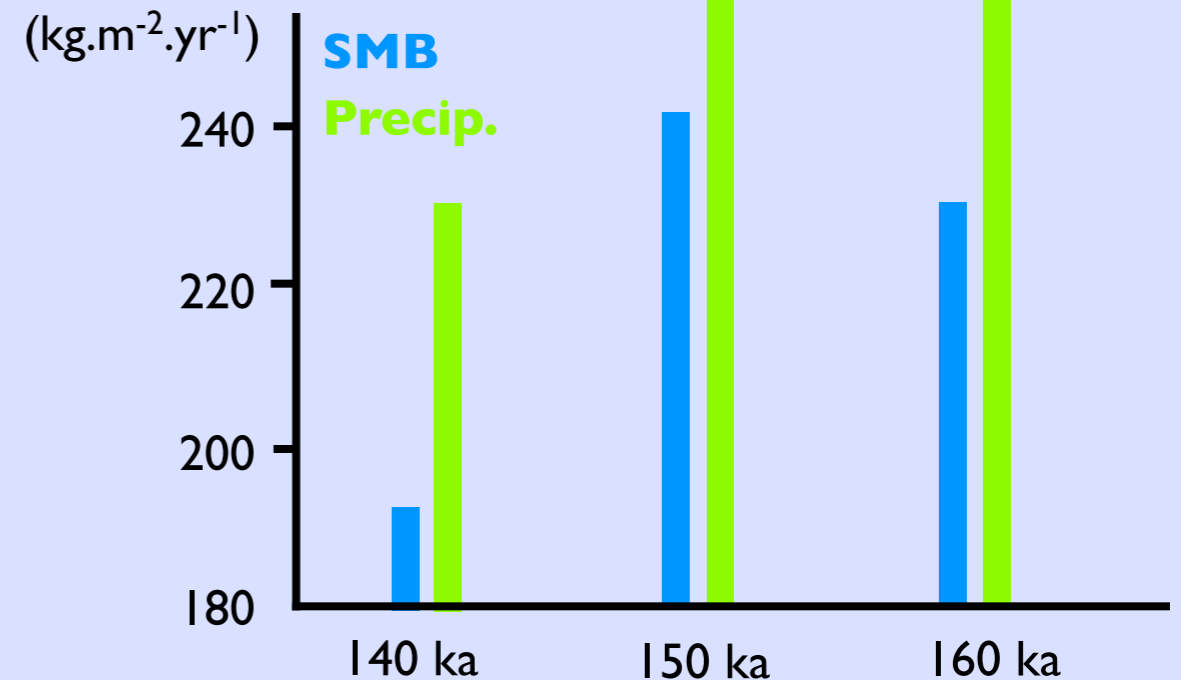


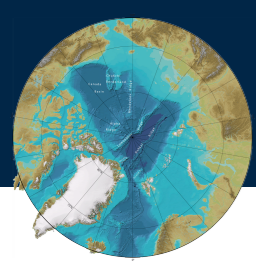


The Late Saalian Eurasian ice sheet mean surface mass balance



($\text{kg.m}^{-2}\text{.yr}^{-1}$)	140 ka	150 ka	160 ka
Tot. precip.	226	329	289
Snow	226	310	282
Evaporation	33	53	44
Ablation	0	15	5
SMB	193	243	233





Conclusions:

How could the ice sheet survive the insolation peak at ≈ 150 ka?

warmer mean annual climate at 150 ka than before:

- More open sea surface conditions: larger moisture advection
- Larger forest extent: warmer regional climate



Larger ablation compensated for by a larger accumulation along the oceanic margins