Improving the numerical convergence of viscous-plastic sea ice models with the Jacobian-free Newton-Krylov method

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The model

Dynamic/thermodynamic VP rheology, ellipse (Hibler, 1979) Domain: Arctic, North Atlantic and CAA Resolutions: 10, 20, 40 or 80 km Coupled to a slab ocean model • Forcing: - geostrophic winds NCEP/NCAR (6h) - climatological currents

Viscous-plastic formulation



$$\sigma_{ij} = 2\eta \varepsilon_{ij} + [\zeta - \eta] \varepsilon_{kk} \delta_{ij} - P \delta_{ij} / 2$$

$$\zeta = P/2\Delta, \ \eta = \zeta e^{-2}$$
 $\Delta = \sqrt{f(\varepsilon_{ij})}$

$$\Rightarrow \zeta = \min(P/2\Delta, \zeta_{\max})$$

The standard solver At time t, we want to solve the nonlinear system of equations:

F(u)=A(u)u-b=0

u⁰: initial iterate

do k=1, k_{max} Solve A(u^{k-1})u^k=b with a linear solver if ||F(u^k)|| < γ_{nl} ||F(u⁰)|| stop enddo

Numerical convergence of the nonlinear solution (6 January1997 00Z)

 10^{2}

norm

residual



 $\Delta x = 10 \text{km}$

Why is the convergence so slow?

- The momentum equation with a VP formulation is highly nonlinear
- The equation is not continuously differentiable because of a capping of the viscous coefficients
- The linearization is not optimal

Continuously differentiable formulation

Discontinuous :

$$\zeta = \min(P/2\Delta, \zeta_{\max}), \ \eta = \zeta e^{-2}$$

Continuous : $\zeta = \zeta_{\text{max}} \tanh(P/2\Delta\zeta_{\text{max}}), \ \eta = \zeta e^{-2}$



Standard linearization versus Newton F(u) $F(u) = \zeta(u)u - b = 0$ Taylor series: $F(u^{k-1}+\delta u^k) = F(u^{k-1})+F'\delta u^k$ uk **,** k-1 **Requiring** $F(u^{k-1}+\delta u^k) = 0$, we get $\delta u^{k} = u^{k} - u^{k-1} = -F(u^{k-1})/F'$ where $F' = \zeta' (u^{k-1})u^{k-1} + \zeta(u^{k-1})$ \implies SNewtool: $u^{k} = u^{k-1} - F(u^{k-1}) / (\zeta'(u^{k-1}) + \zeta(u^{k-1}))$

The JFNK solver

At time t, we want to solve the nonlinear system of equations:

F(u)=A(u)u-b=0

u⁰: initial iterate

do k=1, k_{max} "Solve" J(u^{k-1}) δ u^k=-F(u^{k-1}) with preconditioned GMRES u^k = u^{k-1} + δ u^k if ||F(u^k)|| < γ_{nl} ||F(u⁰)|| stop enddo

 $J(u^{k-1})v \sim (F(u^{k-1}+\epsilon v) - F(u^{k-1})) / \epsilon$

Comparison of the JFNK and standard solver



Computational gain of JFNK over the standard solver



Conclusions

 The standard solver of VP models implies a very slow convergence

The convergence can be improved by:

 having a continuously differentiable equation
 using a Newton method

Both standard and JFNK solvers show a certain lack of robustness

Typical shear deformation field (10 km)



Thank you!

Failures of the two solvers



The fully converged solution 6 January1997 00Z



In matrix form... F(u)=A(u)u-b=0

Standard method (iteration k) $A(u^{k-1})u^{k} = b$ $\bigcup_{k=1}^{k=1} A^{-1}(\bigcup_{k=1}^{k-1})F(\bigcup_{k=1}^{k-1})$ $u^{k} = u^{k-1} - A^{-1}(\bigcup_{k=1}^{k-1})F(\bigcup_{k=1}^{k-1}).$ Newton method (iteration k) $J(u^{k-1})\delta u^{k} = -F(u^{k-1})$ $u^{k} = u^{k-1} - J^{-1}(u^{k-1})F(u^{k-1})$

 $u^{k} = u^{k-1} - (A(u^{k-1}) + G(u^{k-1}))^{-1}F(u^{k-1})$

Failure of the line search method



18

Globalization method for the standard solver



Coming soon...

 méthode de convergence globale (Paul Tupper)
préconditionneur multi-grille algébrique (Paul Tupper)
parallélisation (Andy Pintar)

Dynamique de la glace de mer

 $\rho_i h \frac{Du}{Dt} = -\rho_i h f \hat{k} \times u + \tau_a - \tau_w - \rho_i h g \nabla H_d + \nabla \cdot \sigma$

 $\tau_a = \rho_a C_{J_a} |\mathcal{U}_a^g| (\mathcal{U}_a^g \cos \theta_a + k \times \mathcal{U}_a^g \sin \theta_a)$

 $\tau_{w} = \rho_{w} C_{dw} | u - u_{w}^{g} | | (u - u_{w}^{g}) \cos \theta_{w} + k \times (u - u_{w}^{g}) \sin \theta_{w} |$

terme de

rhéologie

Schéma numérique explicite?

 $\rho_i h \frac{Du}{Dt} = -\rho_i h f \hat{k} \times u + \tau_a - \tau_w - \rho_i h g \nabla H_d + \nabla \cdot \sigma$

Analyse de stabilité

 $\rho_{i}h\frac{\partial u}{\partial t} = \zeta_{\max}\frac{\partial^{2}u}{\partial x^{2}} + R \implies \Delta t \leq \frac{\rho_{i}h}{2\zeta_{m}}\Delta x^{2}$

Schéma numérique implicite Nous allons résoudre de façon implicite au temps: ...t- Δ t, t, t+ Δ t, t+2 Δ t, ...

Au temps t nous avons:

 $-\frac{\rho_{i}hu^{t}}{\Delta t} + \rho_{i}hfv_{avg}^{t} + \frac{\partial}{\partial x}\left[\left(\eta + \zeta\right)\frac{\partial u^{t}}{\partial x}\right] + \frac{\partial}{\partial x}\left[\left(\zeta - \eta\right)\frac{\partial v^{t}}{\partial y}\right] + \frac{\partial}{\partial y}\left[\left(\zeta - \eta\right)\frac{\partial v^{t}}{\partial y}\right] +$ $\frac{\partial}{\partial v} \left| \eta \frac{\partial u^{t}}{\partial v} \right| + \frac{\partial}{\partial v} \left[\eta \frac{\partial v^{t}}{\partial x} \right] - C_{w} \left(u^{t} \cos \theta_{w} - v_{avg}^{t} \sin \theta_{w} \right) = b_{u}$

Schéma numérique implicite

Nous laissons tomber l'indice t. À l'itération k nous avons:

$$-\frac{\rho_{i}hu^{k}}{\Delta t} + \rho_{i}hfv_{avg}^{k} + \frac{\partial}{\partial x}\left[\left(\eta(u_{l}) + \zeta(u_{l})\right)\frac{\partial u^{k}}{\partial x}\right] + \frac{\partial}{\partial x}\left[\left(\zeta(u_{l}) - \eta(u_{l})\right)\frac{\partial v^{k}}{\partial y}\right] + \frac{\partial}{\partial y}\left[\eta(u_{l})\frac{\partial v^{k}}{\partial x}\right] - C_{w}(u_{l})\left(u^{k}\cos\theta_{w} - v_{avg}^{k}\sin\theta_{w}\right) = b_{u}$$

Rhéologie de la glace de mer en 1D



Rhéologie de la glace de mer en 1D



Méthode du Résidu Minimal Généralisé (GMRES) avec préconditionneur

- méthode de sous-espaces de Krylov
- nécessite peu de mémoire
- la symétrie n'est pas nécessaire
- bonne convergence avec préconditionneur
- parallélisable

Optimisation et comparaison des solveurs linéaires



28

Taux de convergence des solveurs linéaires



29

Méthodes stationnaires: Jacobi, Gauss-Seidel, SOR et LSOR



W. L. Briggs, a multigrid tutorial³⁰

Méthodes stationnaires: un exemple avec la méthode Jacobi



W. L. Briggs, a multigrid tutorial ³¹

Traitement implicite du terme de Coriolis



32



С







Fig 3-7

- 0.5 - 0.4 - 0.3 - 0.2 - 0.1

> -0.1 -D.2 0.3

0.5

3.5

0.4 1.3

0.2 - 0.1 0

-0.1 -0.2 0.3 -0.4

-0.5

33

Fig 3-11





Méthodes stationnaires: un exemple avec la méthode Jacobi



W. L. Briggs, a multigrid tutorial ³⁵

Traitement implicite du terme de Coriolis



36

Erreurs après 2 itérations 6 janvier 1997 00Z





 $\Delta t = 6h$ $\Delta x = 10km$

Erreurs après 10 itérations 6 janvier 1997 00Z



Erreurs après 40 itérations 6 janvier 1997 00Z



Erreurs après 2 et 10 itérations avec un pas de temps de 30 minutes (6 janvier 1997 00Z)

2 itérations

10 itérations



 $\Delta t = 30 \min$ $\Delta x = 10 km$

Déformation (cisaillement) Jan 6 1997 00Z $\Delta t = 6h$

 $\Delta x = 10$ km





Stess State vs Pseudo Time Step 0 5 15



LSR

ADI

A very short course on sea ice rheology: the dilatancy effect

N

T.





$$p = -\left(\frac{\sigma_{11} + \sigma_{22}}{2}\right)$$

$$q = \sqrt{\left(\frac{\sigma_{11} - \sigma_{22}}{2}\right)^2 + \sigma_{12}^2}$$

p and q are called the stress invariants

Traitement implicite (iteration k, cut v)

$$-\frac{\rho_{i}hu^{k}}{\Delta t} + \rho_{i}hfv_{avg}^{k} + \frac{\partial}{\partial x}\left[\left(\eta(u_{l}) + \zeta(u_{l})\right)\frac{\partial u^{k}}{\partial x}\right] + \frac{\partial}{\partial x}\left[\left(\zeta(u_{l}) - \eta(u_{l})\right)\frac{\partial v^{k}}{\partial y}\right] + \frac{\partial}{\partial y}\left[\eta(u_{l})\frac{\partial v^{k}}{\partial x}\right] - C_{w}(u_{l})\left(u^{k}\cos\theta_{w} - v_{avg}^{k}\sin\theta_{w}\right) = b_{u}$$

$$-\frac{\rho_{i}hv^{k}}{\Delta t} - \rho_{i}hfu^{k}_{avg} + \frac{\partial}{\partial y}\left[\left(\eta(u_{l}) + \zeta(u_{l})\right)\frac{\partial v^{k}}{\partial y}\right] + \frac{\partial}{\partial y}\left[\left(\zeta(u_{l}) - \eta(u_{l})\right)\frac{\partial u^{k}}{\partial x}\right] + \frac{\partial}{\partial x}\left[\eta(u_{l})\frac{\partial v^{k}}{\partial y}\right] - C_{w}(u_{l})\left(v^{k}\cos\theta_{w} - u^{k}_{avg}\sin\theta_{w}\right) = b_{v}$$