

Application of a higher-order flow model to outlet glacier dynamics in Greenland

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Outline

- (1) Introduce a tuning procedure that provides a useful initial condition for a higher-order, stand alone ice sheet model (applied to the GIS).

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- (3) From perturbation experiments, estimate SLR that is already “in” the ice sheet as a result of outlet glacier perturbations that have already occurred.
- (4) Scale the model results in order to estimate the potential future SLR from *all* recent outlet glacier perturbations.

Show tuned GIS model and compare with target velocities

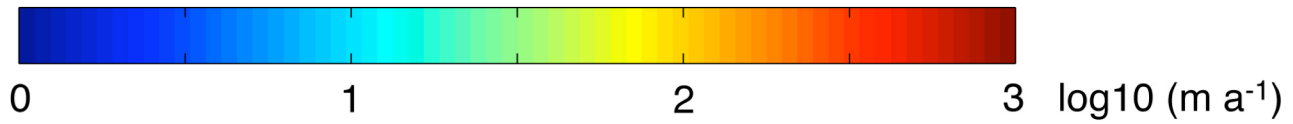
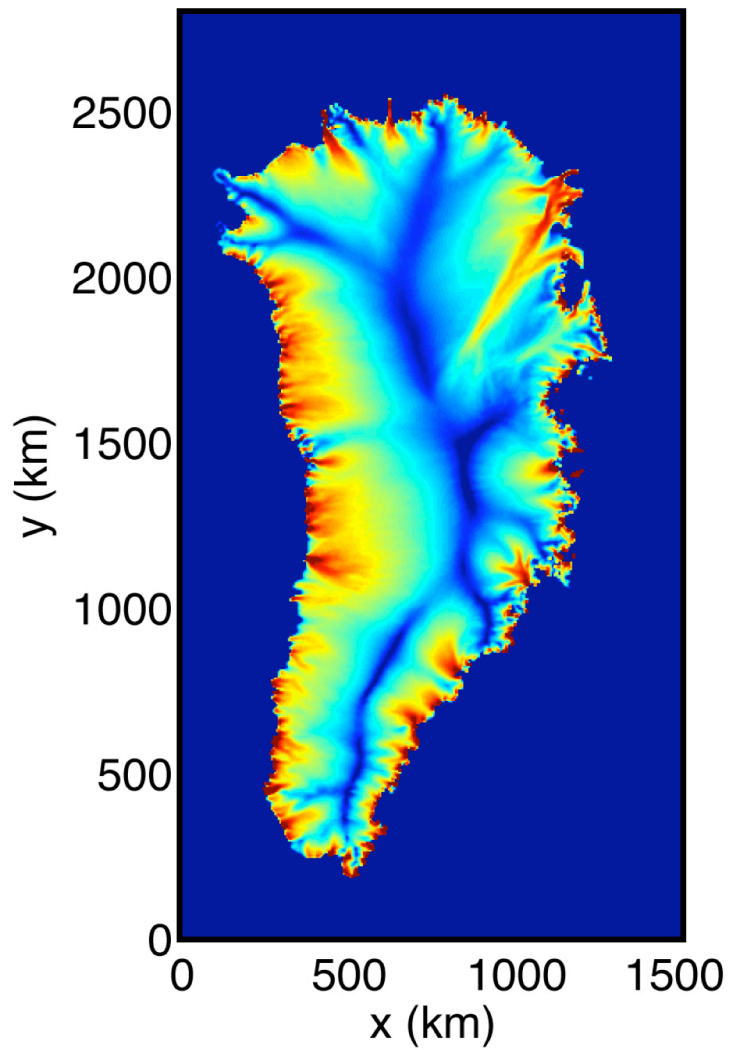
Apply model to outlet glacier perturbation experiments

Compare modelled and observed thinning rates

Estimate minimum future SLR from modelled glaciers

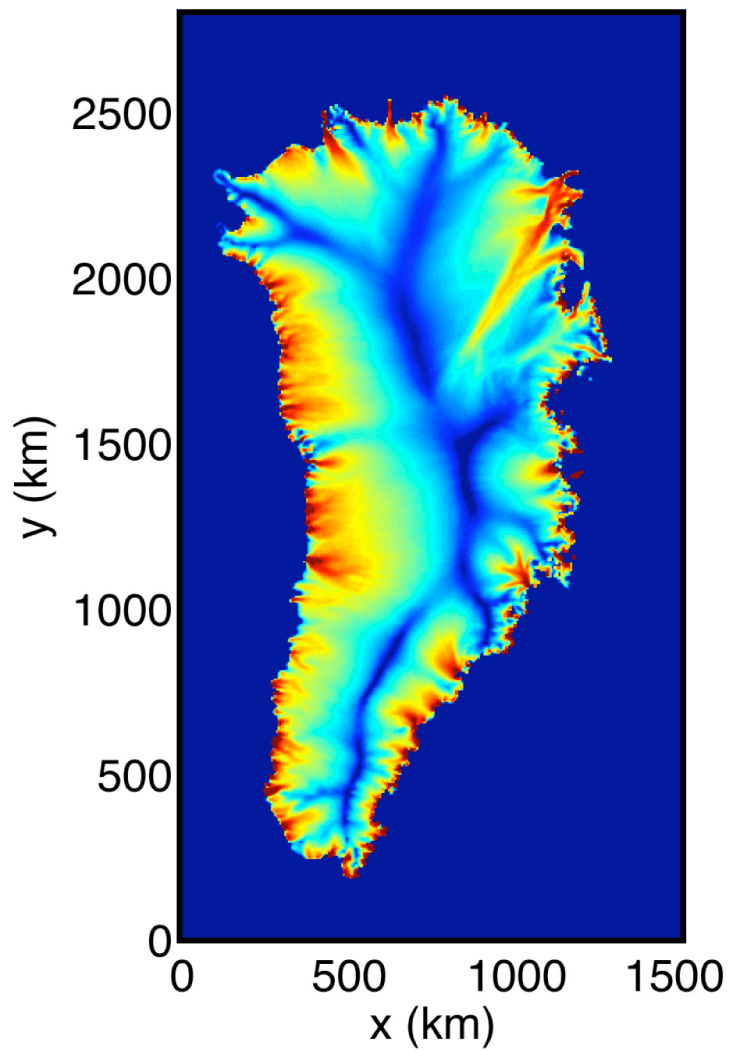
Scale estimate to entire ice sheet

balance velocities¹

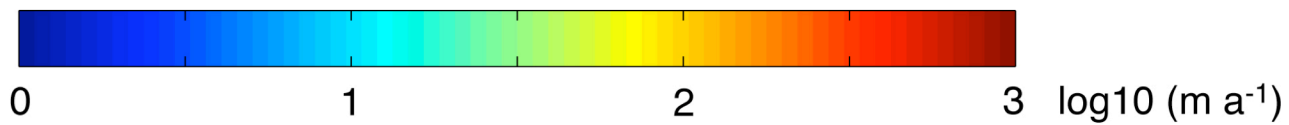
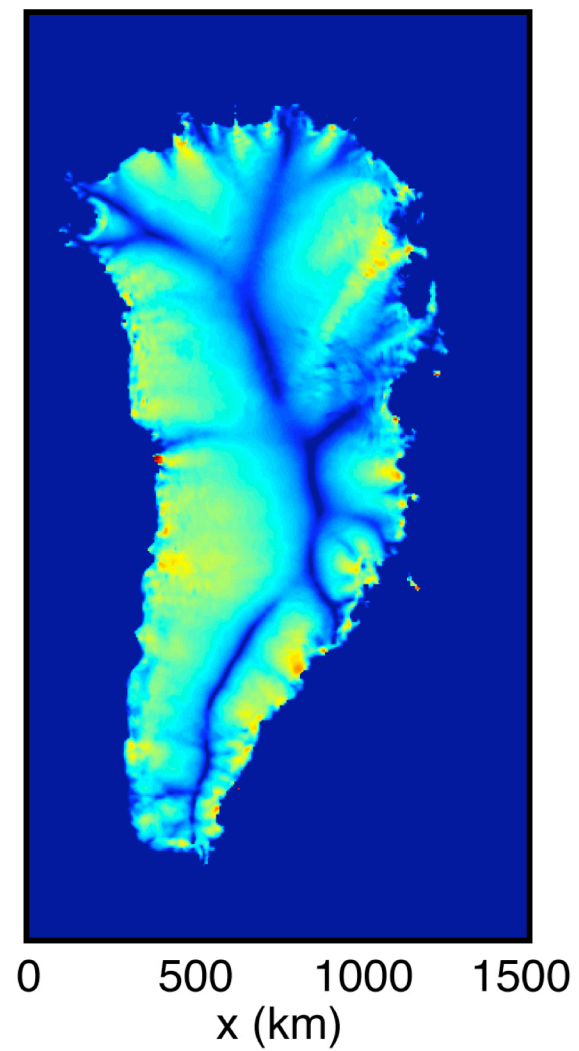


¹Bamber et al. (*J. Glac.*, **46**, 2000)

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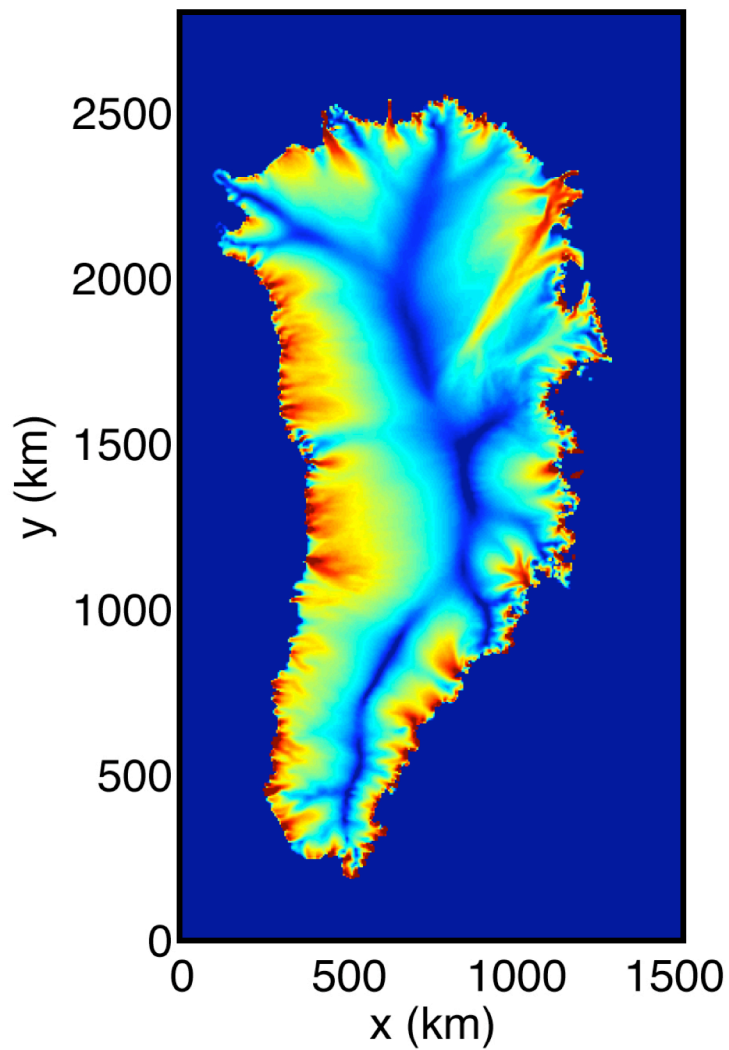


ss thermomech. velocities

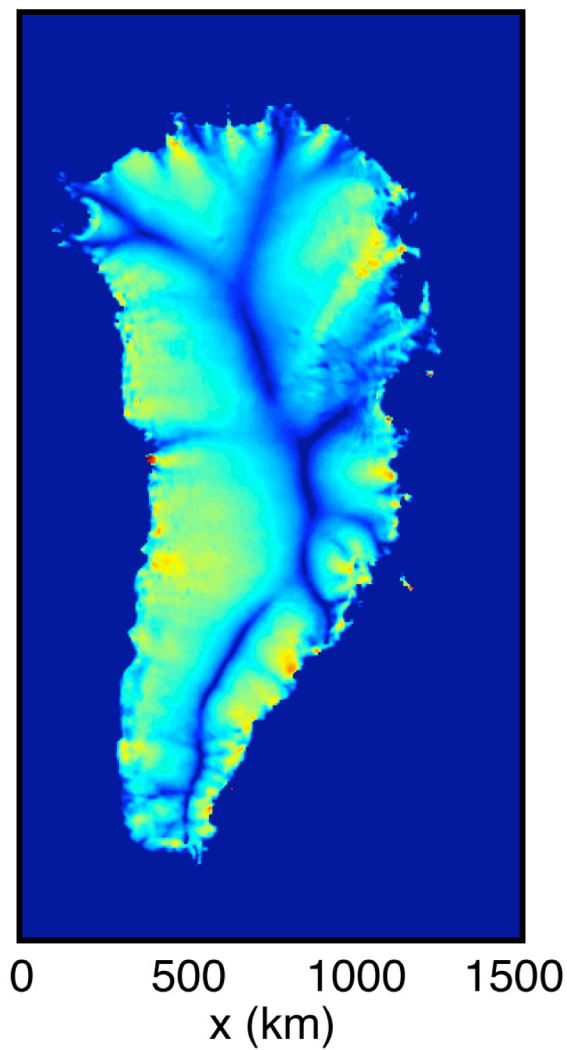


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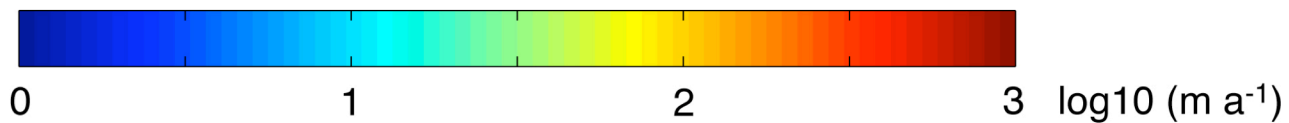
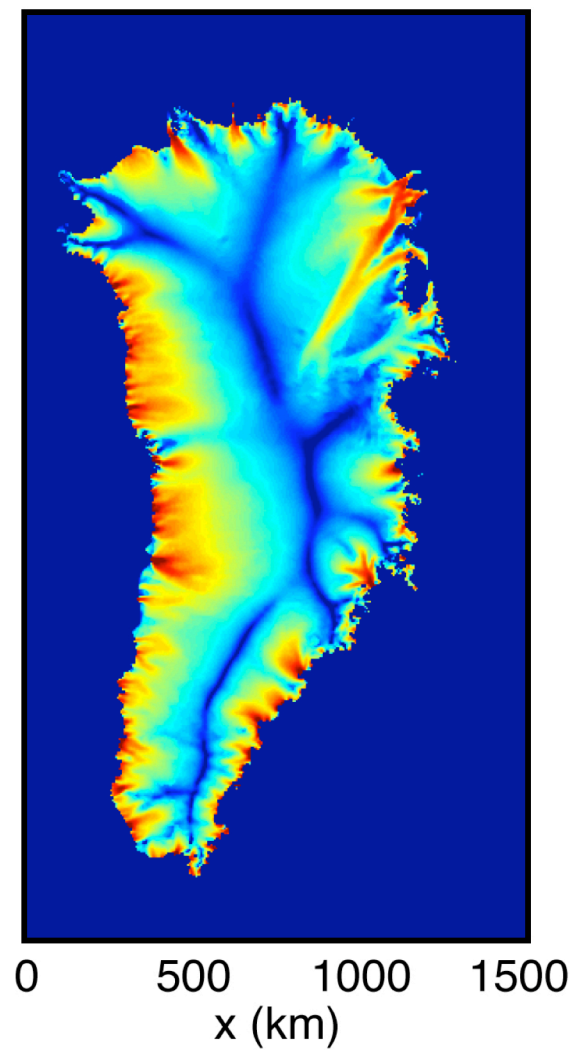
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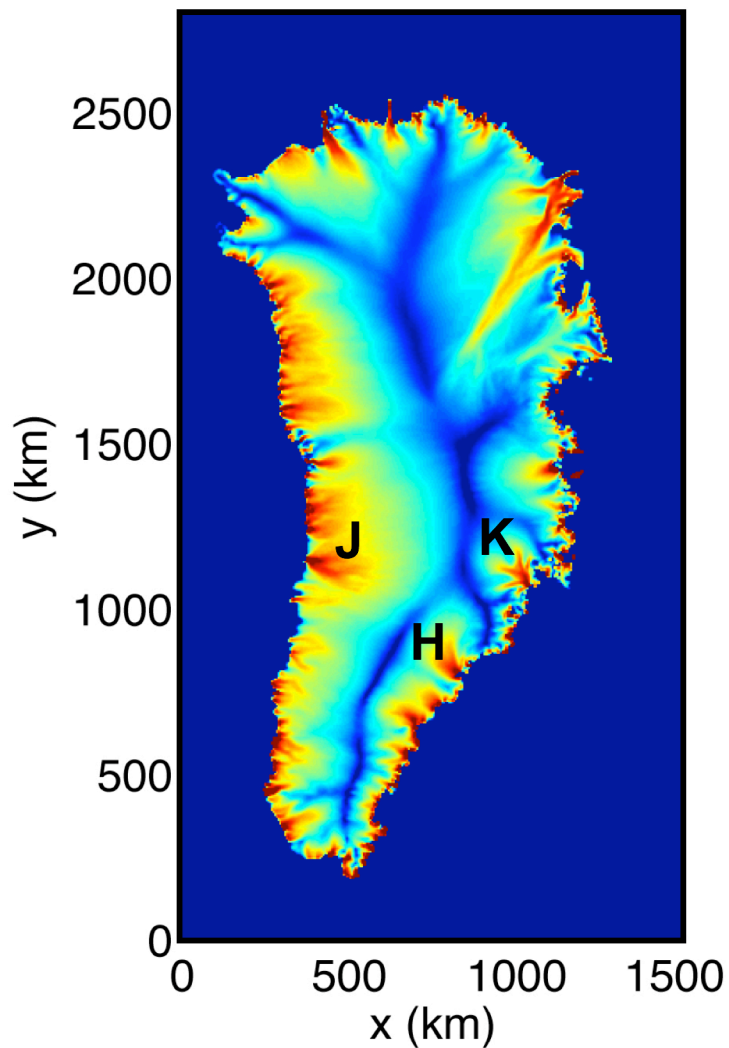


ss sliding velocities

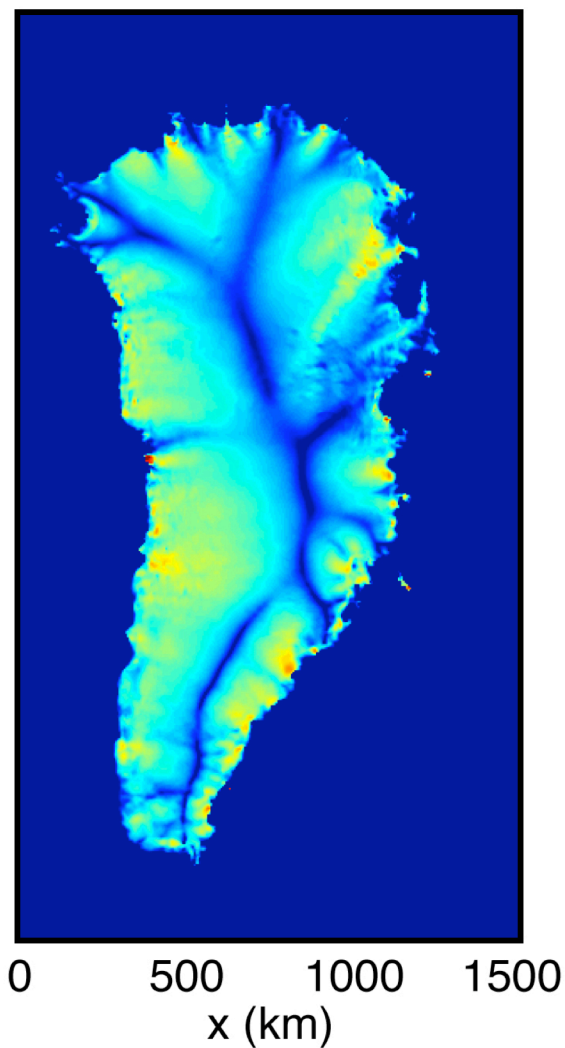


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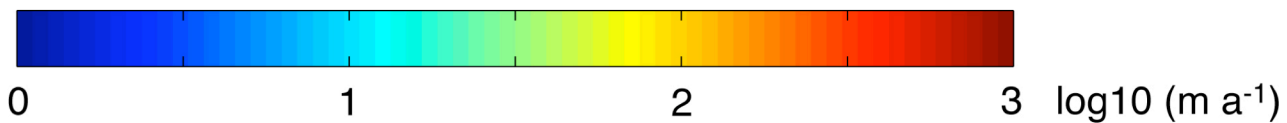
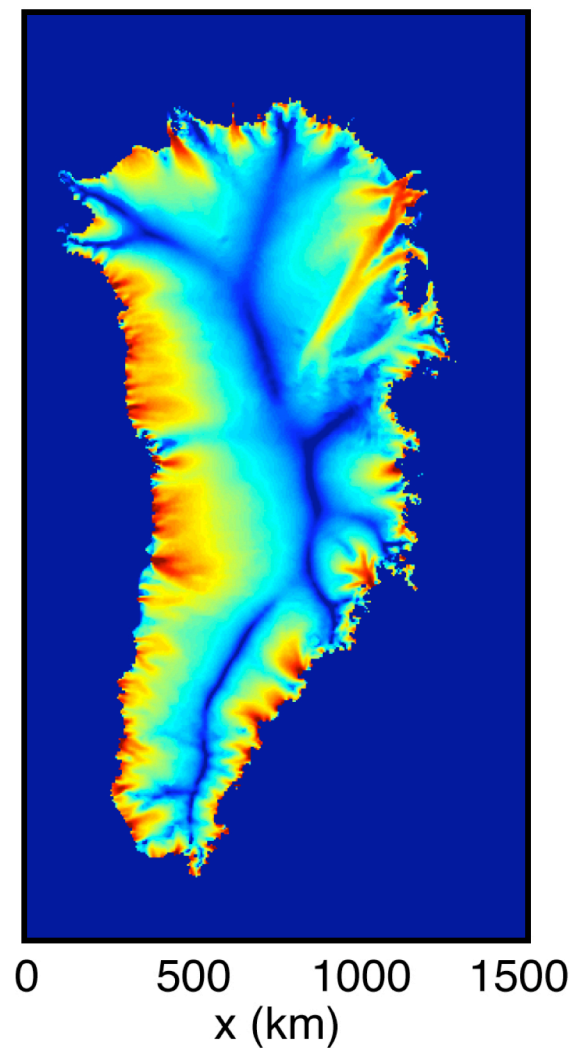
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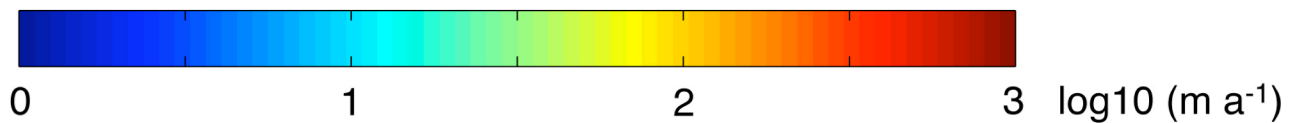
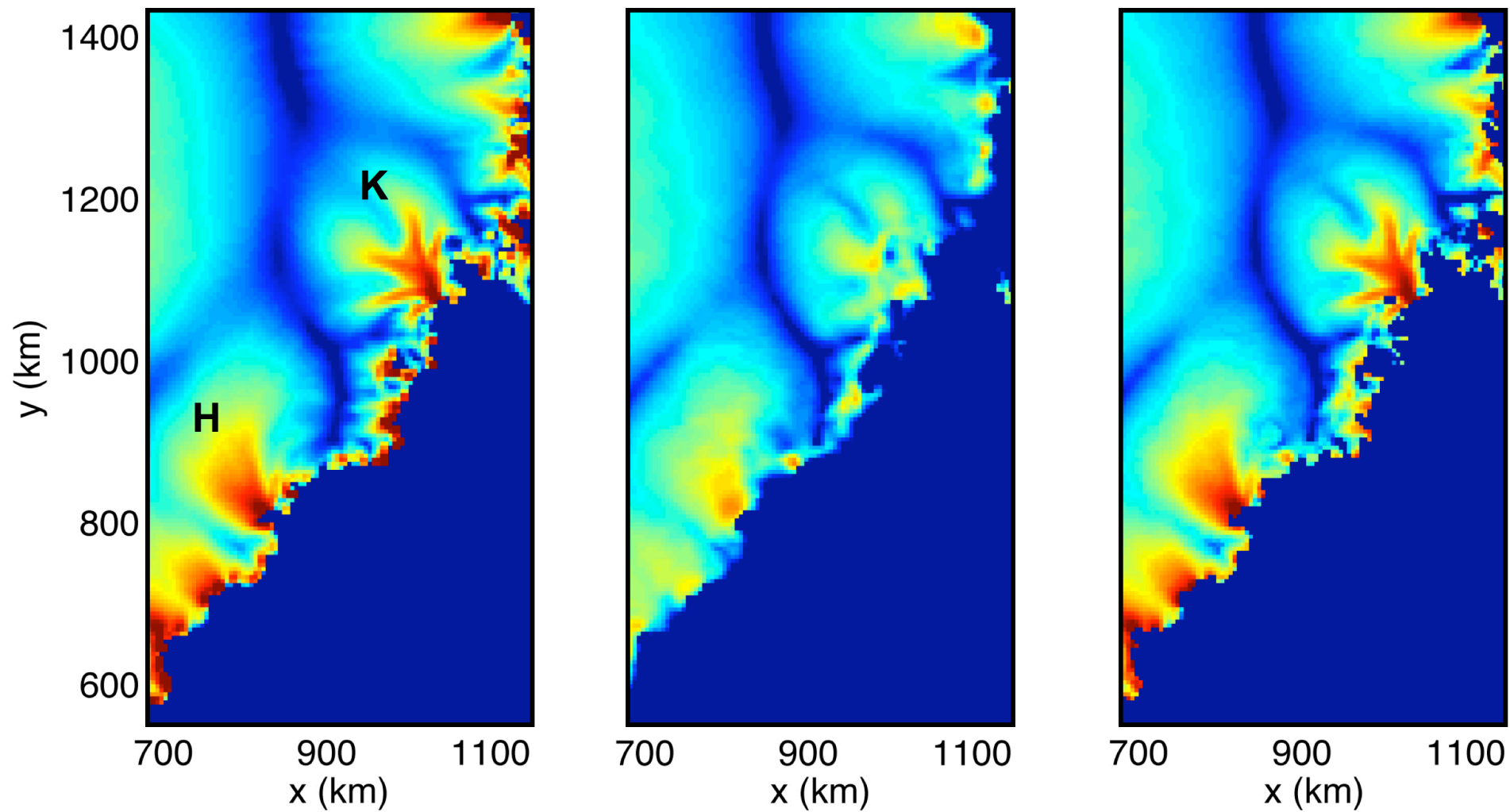


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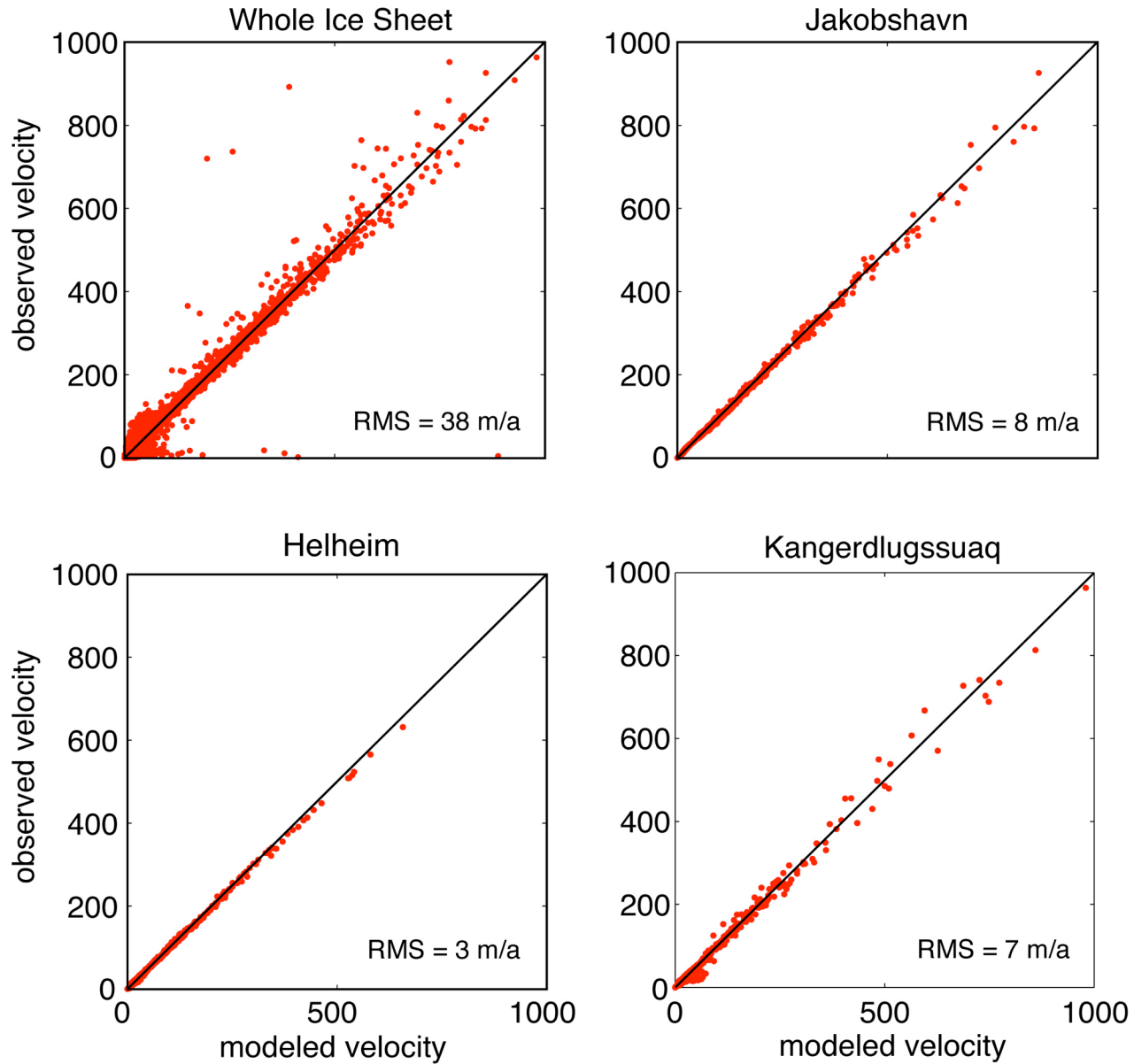
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Tuning Statistics





Show tuned GIS model and compare with target velocities

Apply model to outlet glacier perturbation experiments

Compare modelled and observed thinning rates

Estimate minimum future SLR from modelled glaciers

Scale estimate to entire ice sheet

Perturbation experiments

- (1) Using initial condition, perturb outlet BC for select outlet glaciers
(Jak, Hel, Kang)

Perturbation experiments: initial conditions

Assumed ~ss discharge ($\text{km}^3 \text{ a}^{-1}$) from select GIS basins:

<u>Basin</u>	<u>model (% error)</u>	<u>observations¹</u>
Jakobshavn:	23.5 (<1)	23.6
Helheim:	25.8 (-2)	26.3
Kangerdlugssuaq:	27.8 (0)	27.8

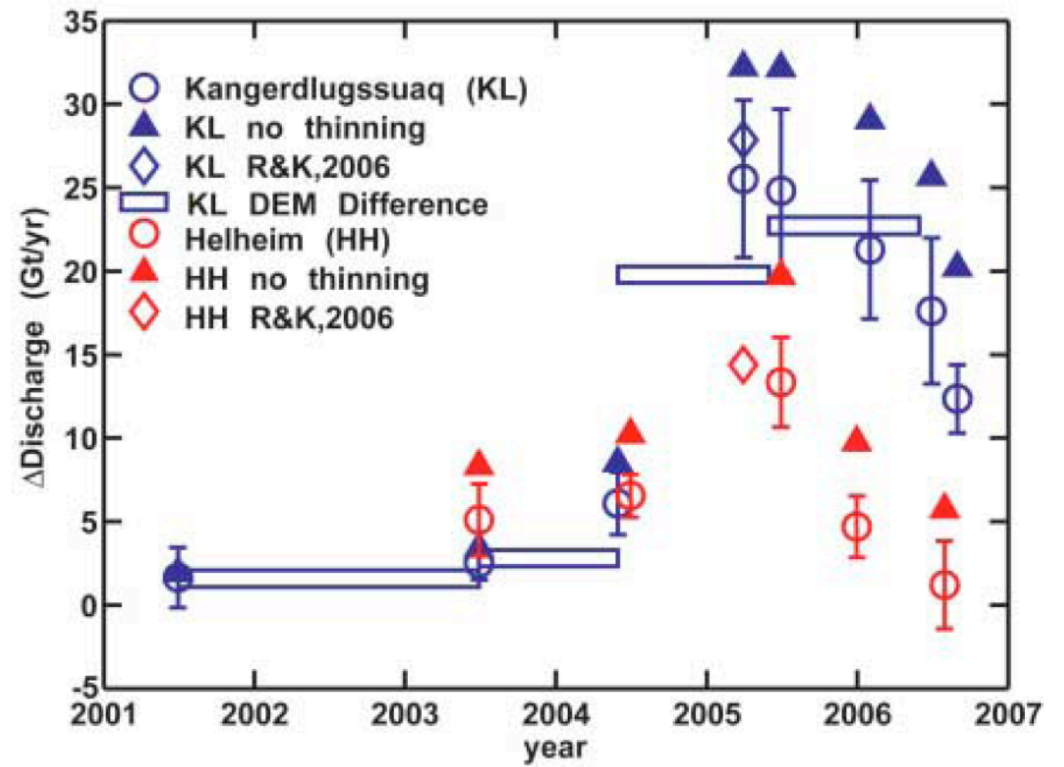
¹Rignot and Kanagaratnam (*Science*, **311**, 2006)

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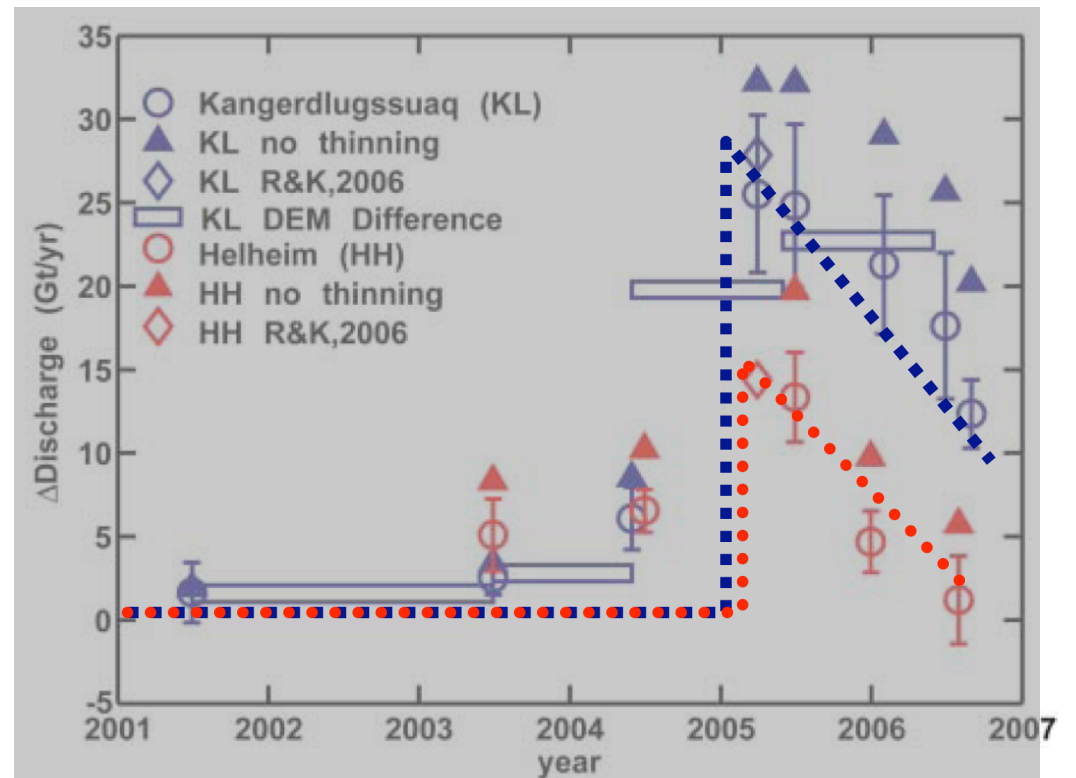
(2) perturbation: tune outlet BC so that flux change matches observations

Kang. and Hel.¹



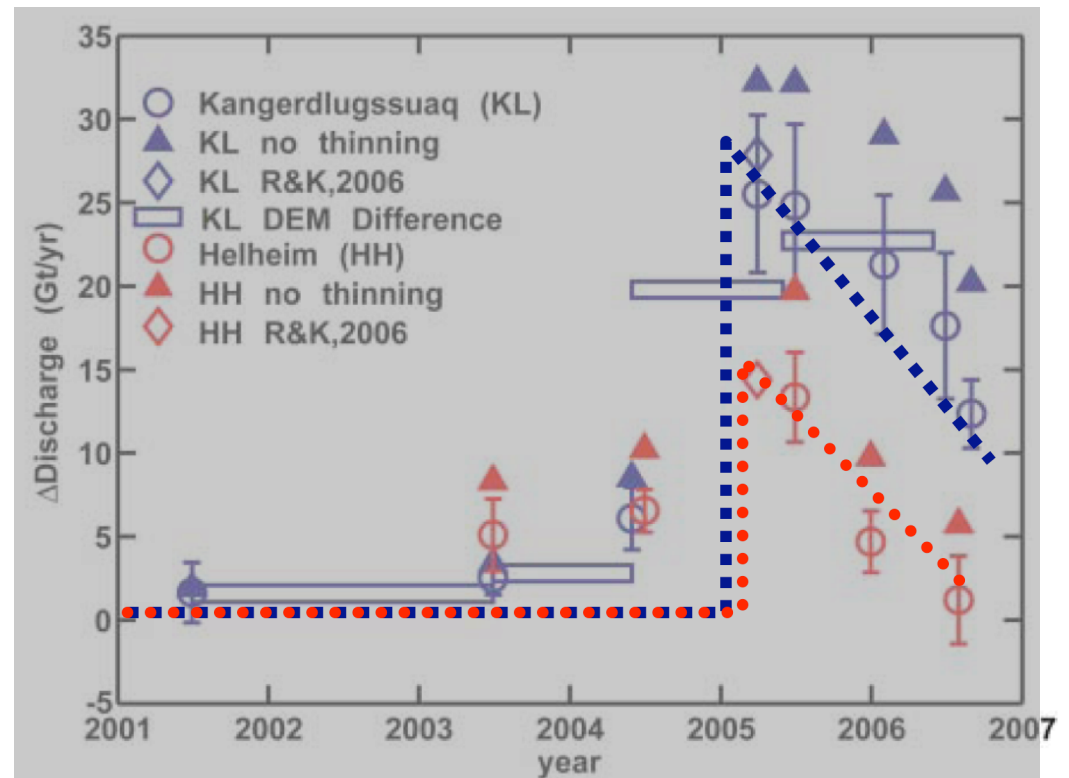
¹Howat et al. (*Science*, **315**, 2007)

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Jak:

Approx. flux doubling^{2,3},
from $\sim 25\text{-}50 \text{ km}^3 \text{ a}^{-1}$, in the
mid-late 90's

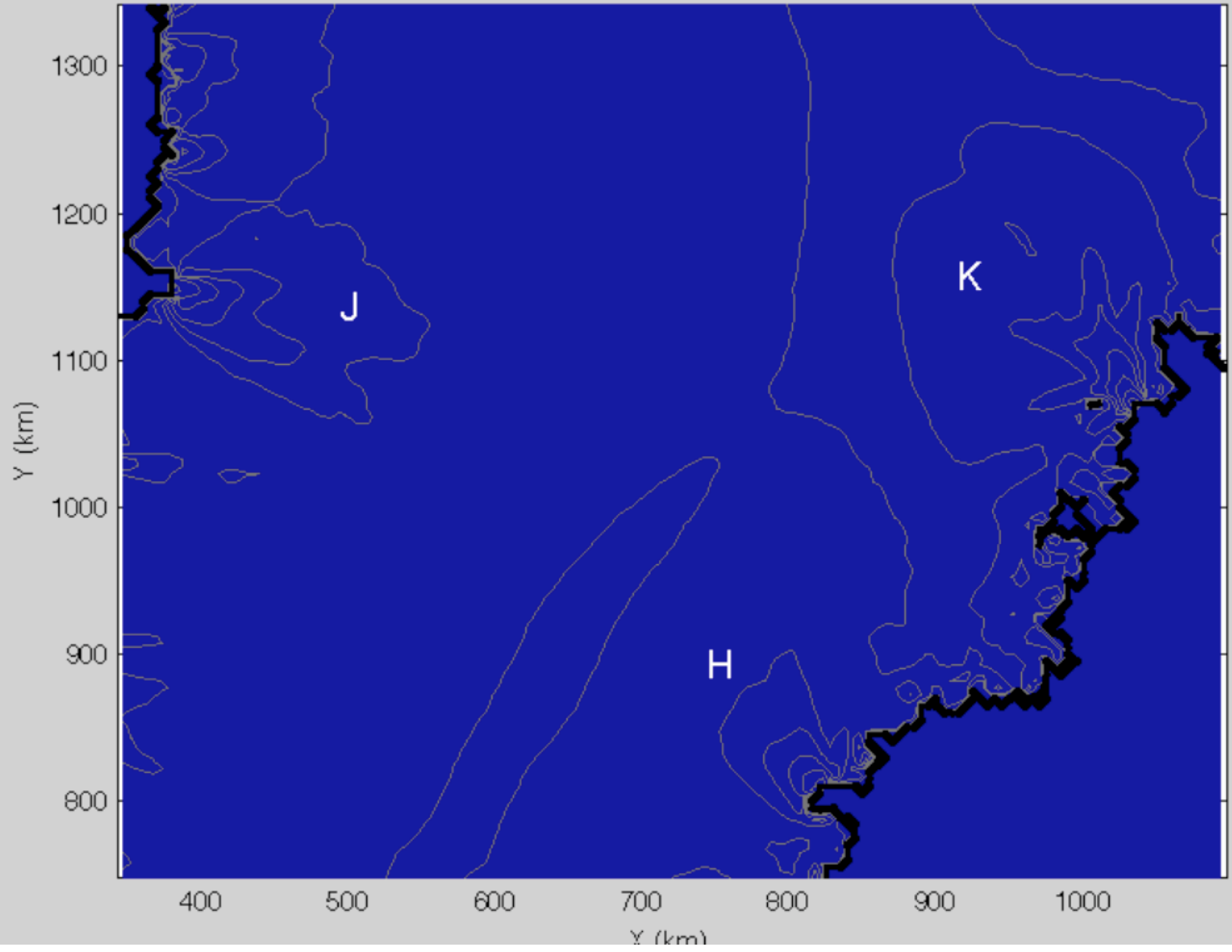
²Joughin et al. (*Nature*, **432**, 2004)

³Rignot and Kanagaratnam (*Science*, **311**, 2006)

Perturbation experiments (prognostic)

Next: movie of speed change over Jak., Kang. and Hel. glacier catchments from model years 0 to 25 (dt = 0.2 yrs) ...

0



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The result is an estimate for how past dynamic perturbations to these 3 basins will affect future GIS mass balance (i.e. long term, diffusive response).

This can be used to estimate a *minimum* future SLR contribution from GIS outlet glacier dynamics.

An aerial photograph of a glacier system, likely in Greenland, with a blue color overlay. The glacier is shown as a large, textured mass of ice, with various channels and features. The background shows a vast, flat landscape under a clear sky.

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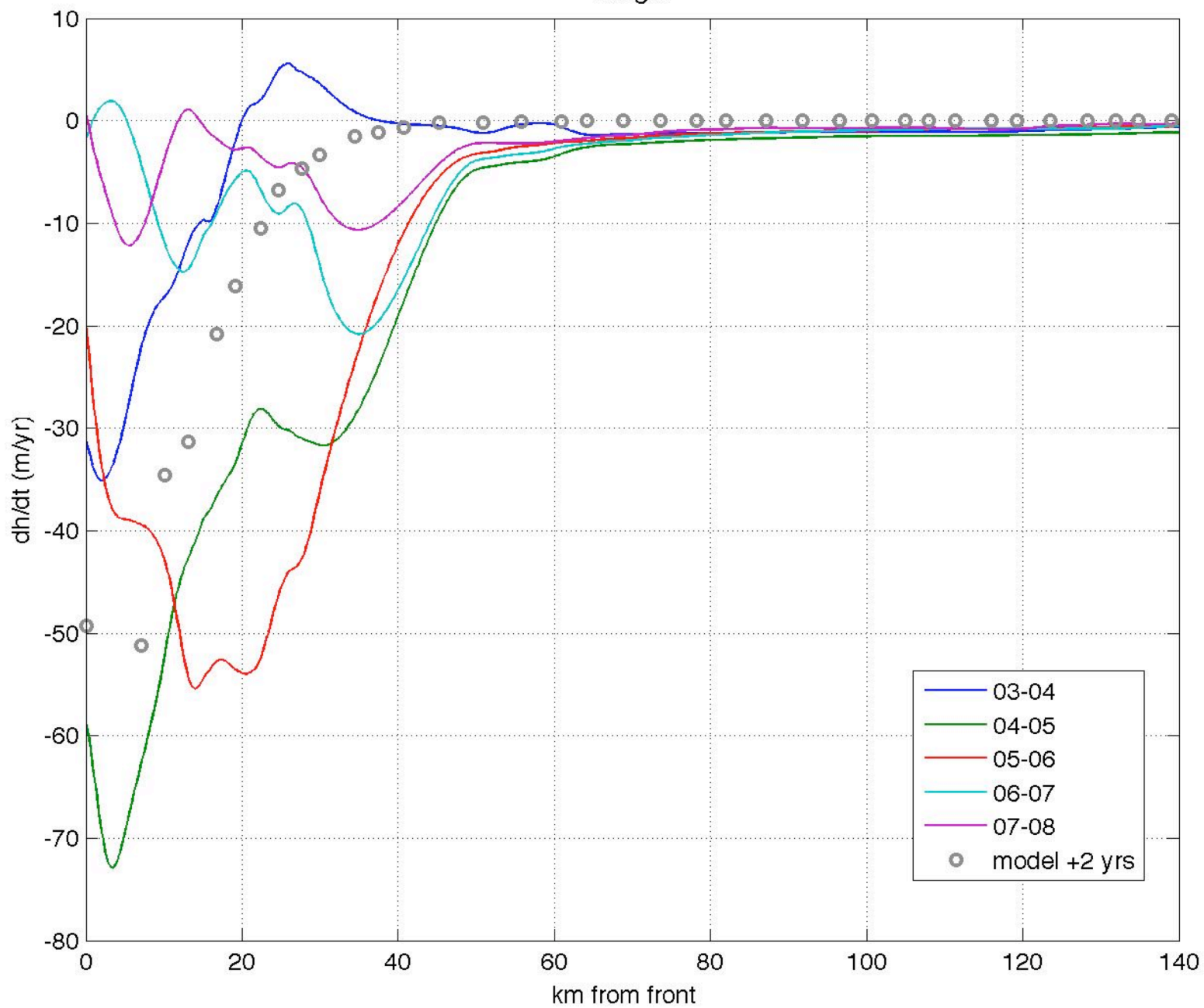
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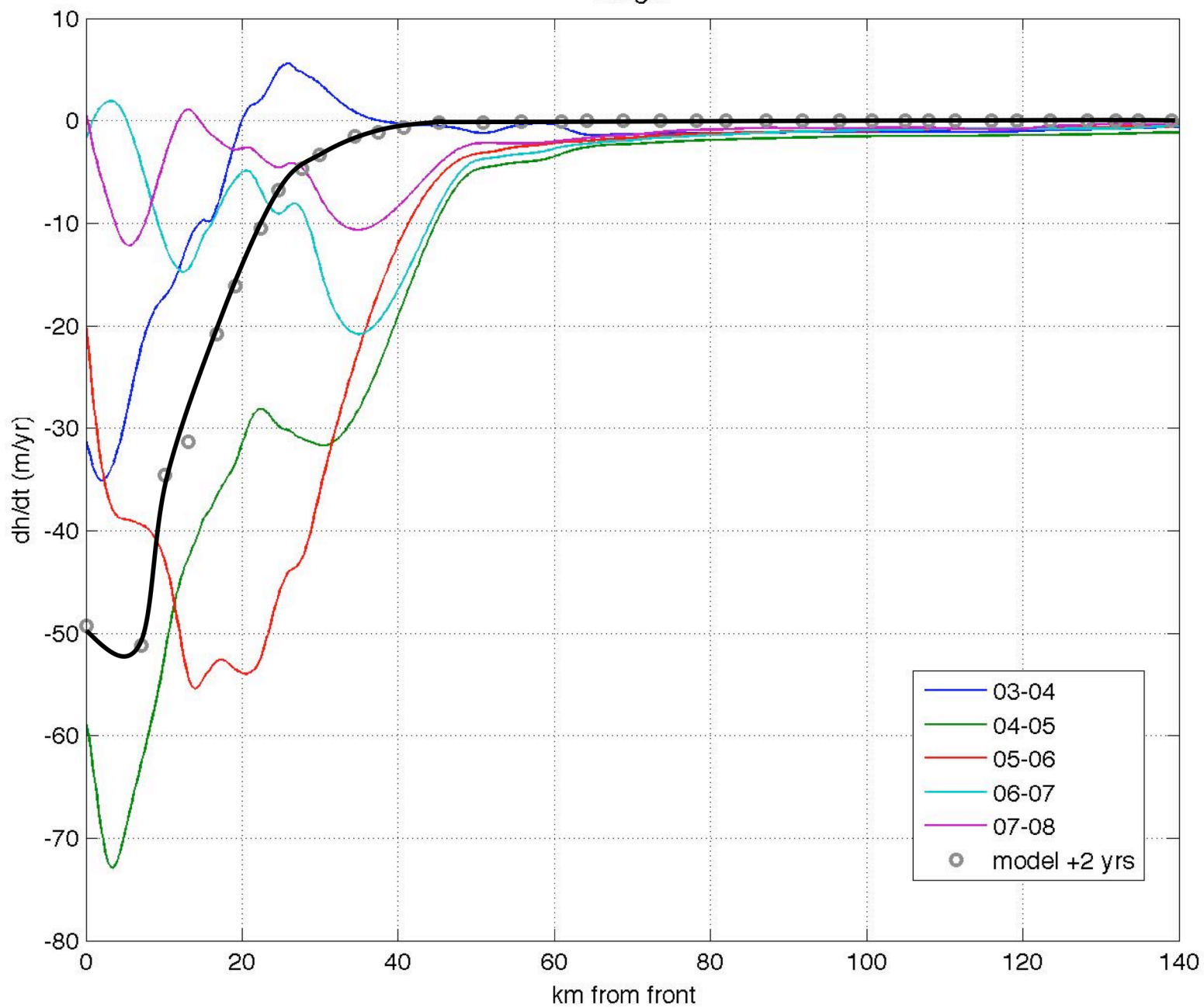
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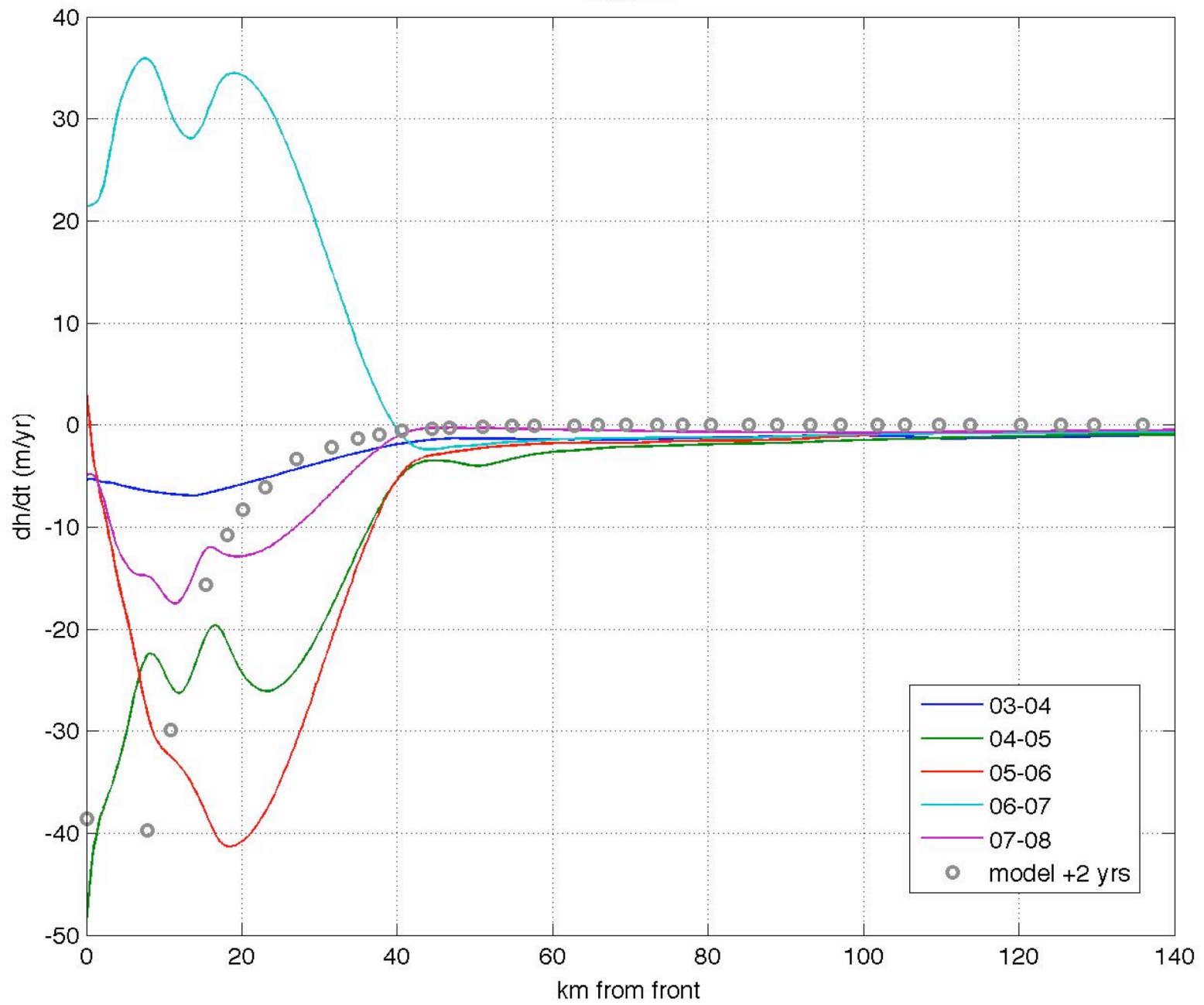
Kanger



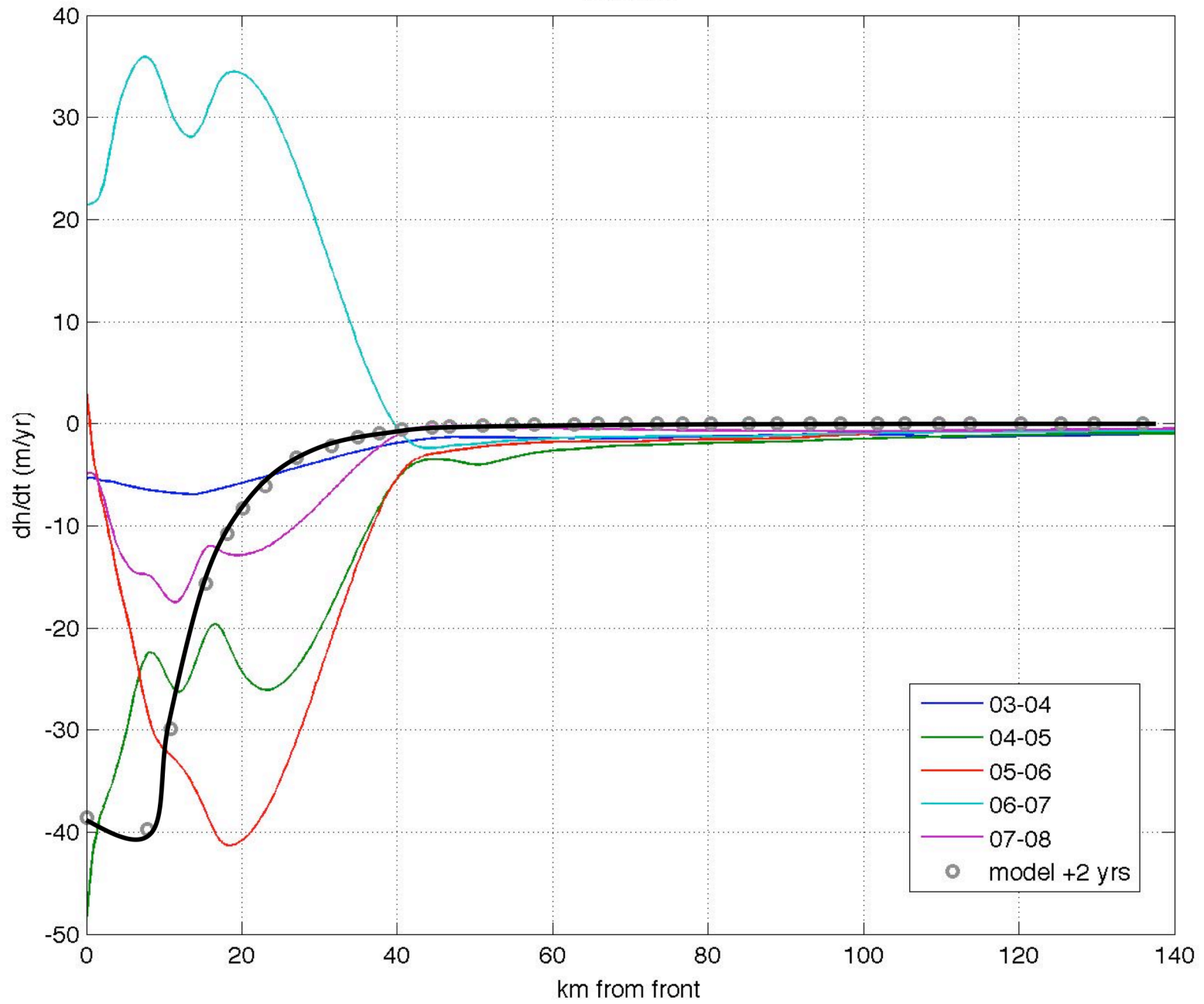
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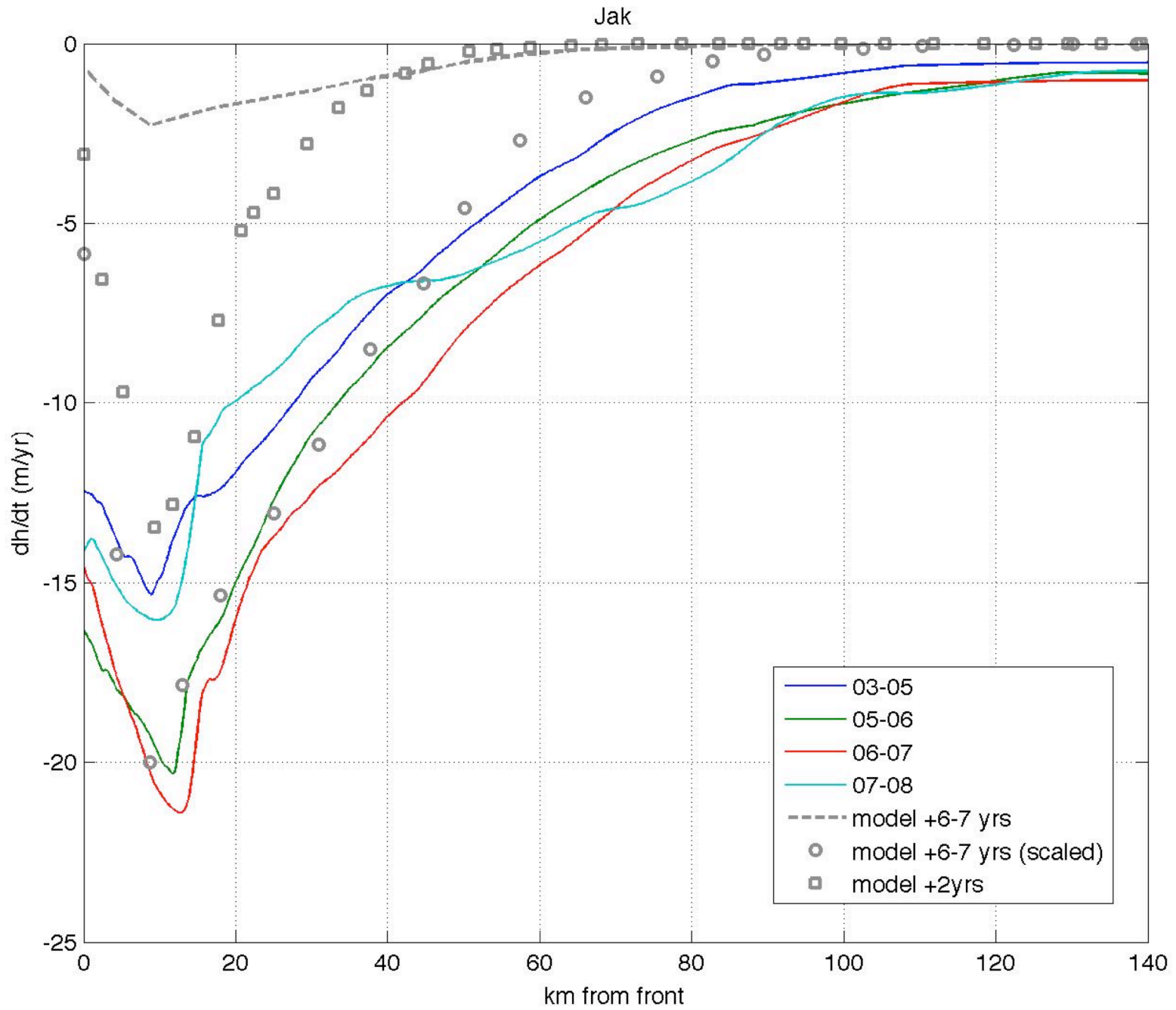


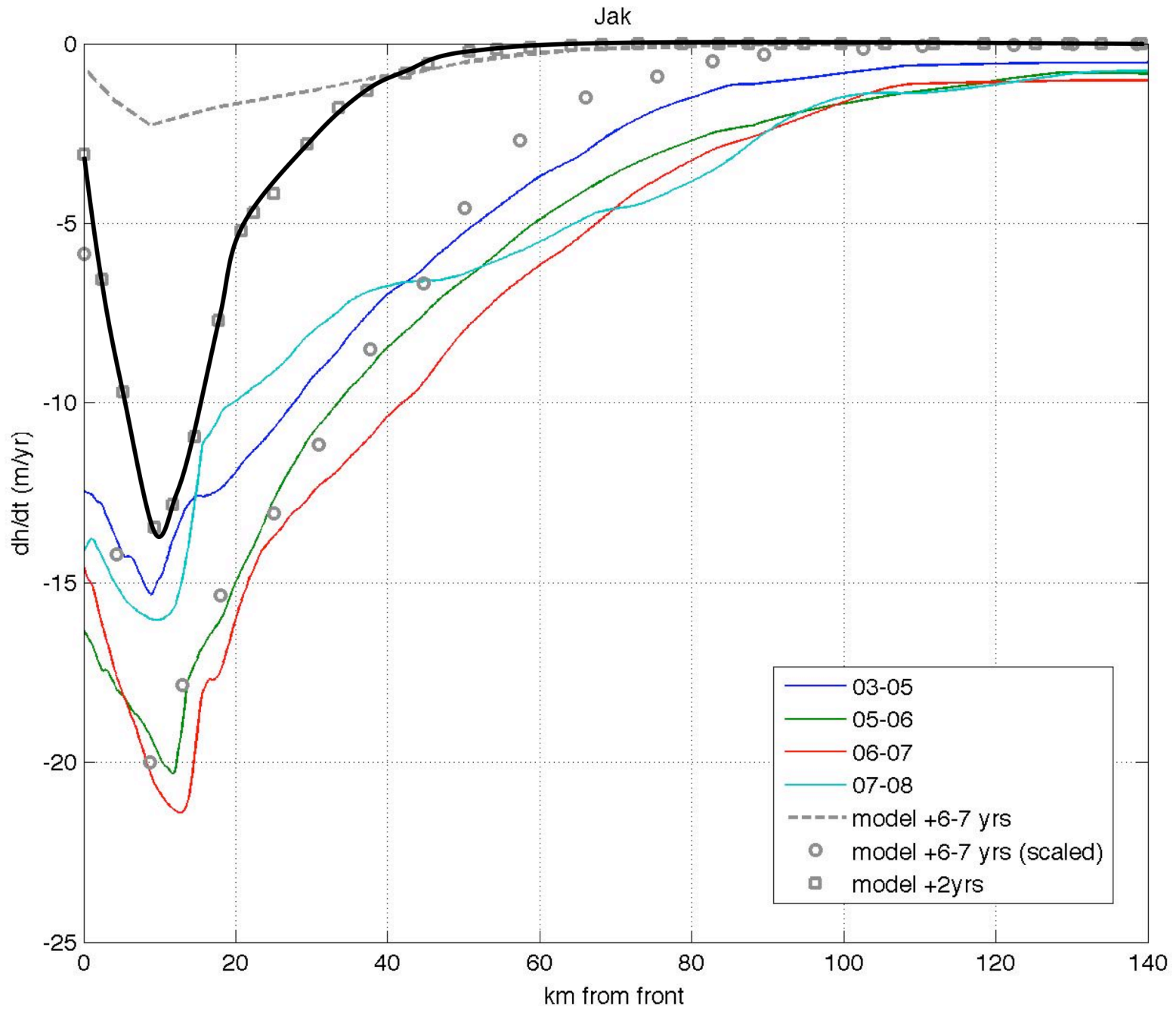
Helheim

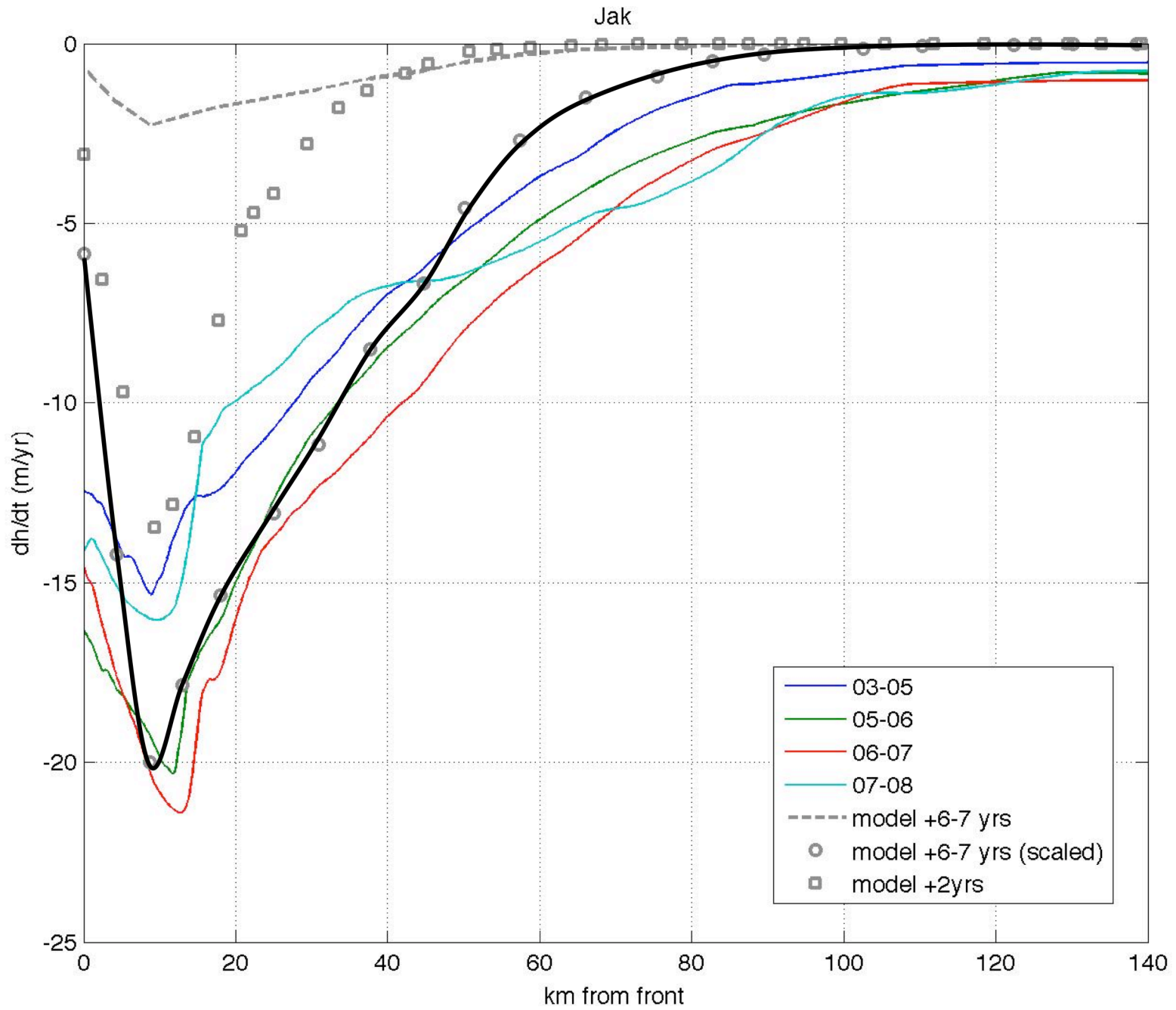


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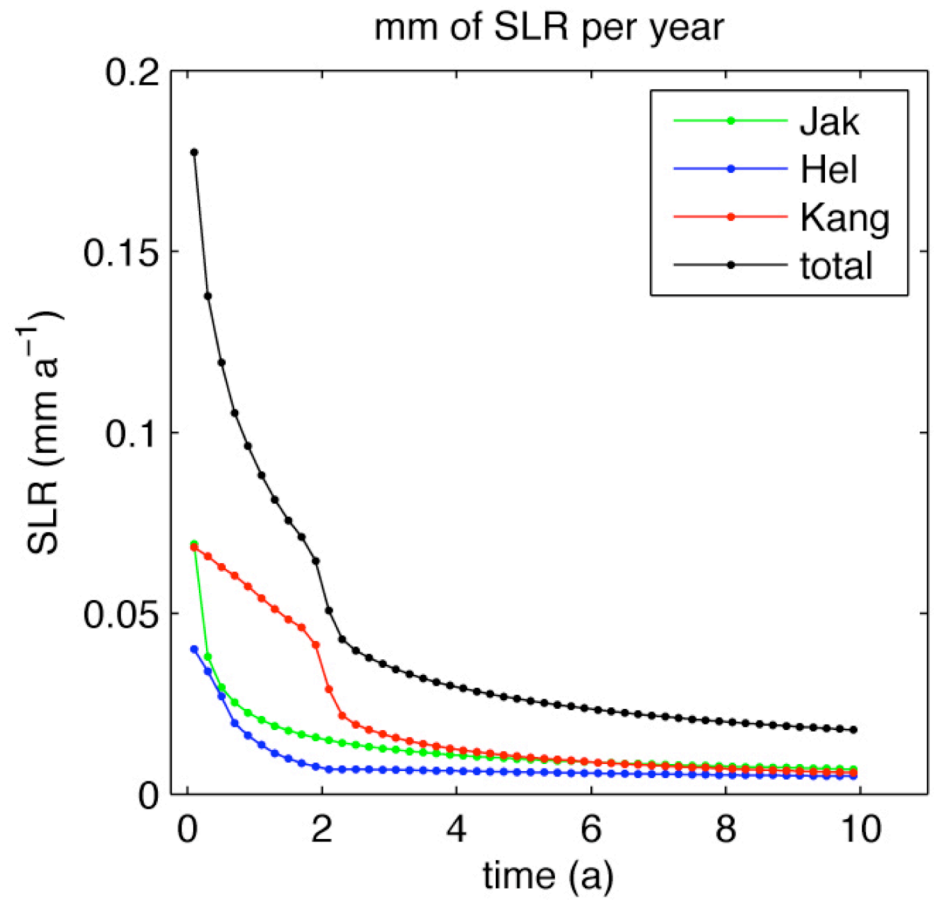
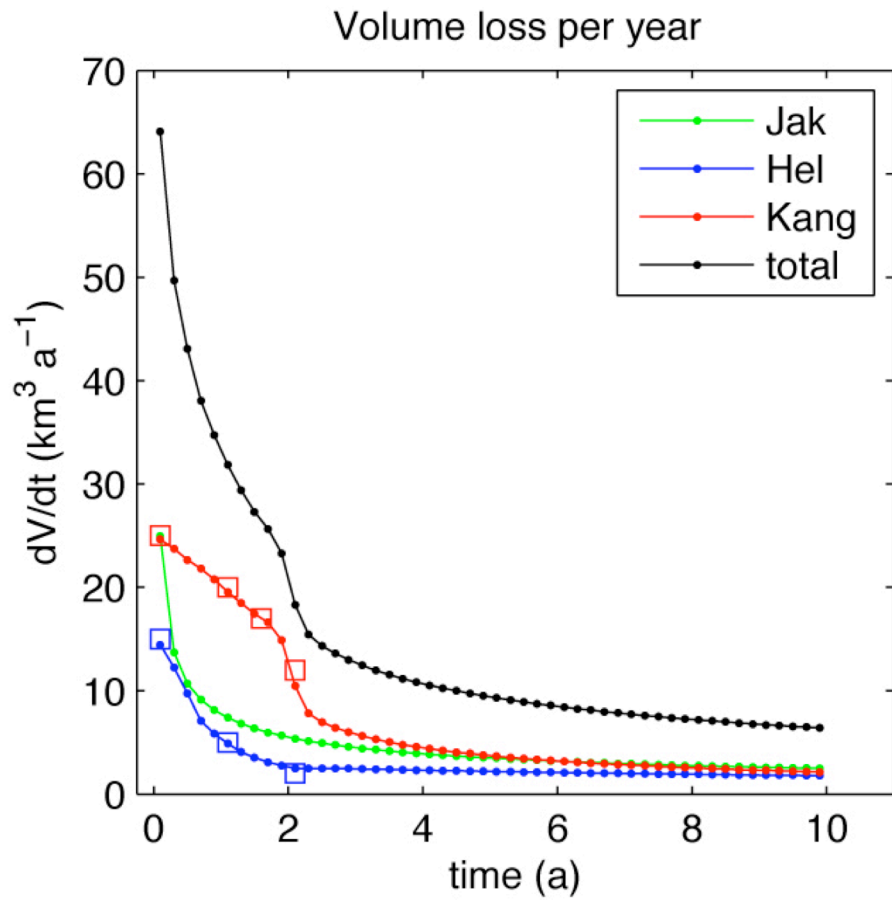
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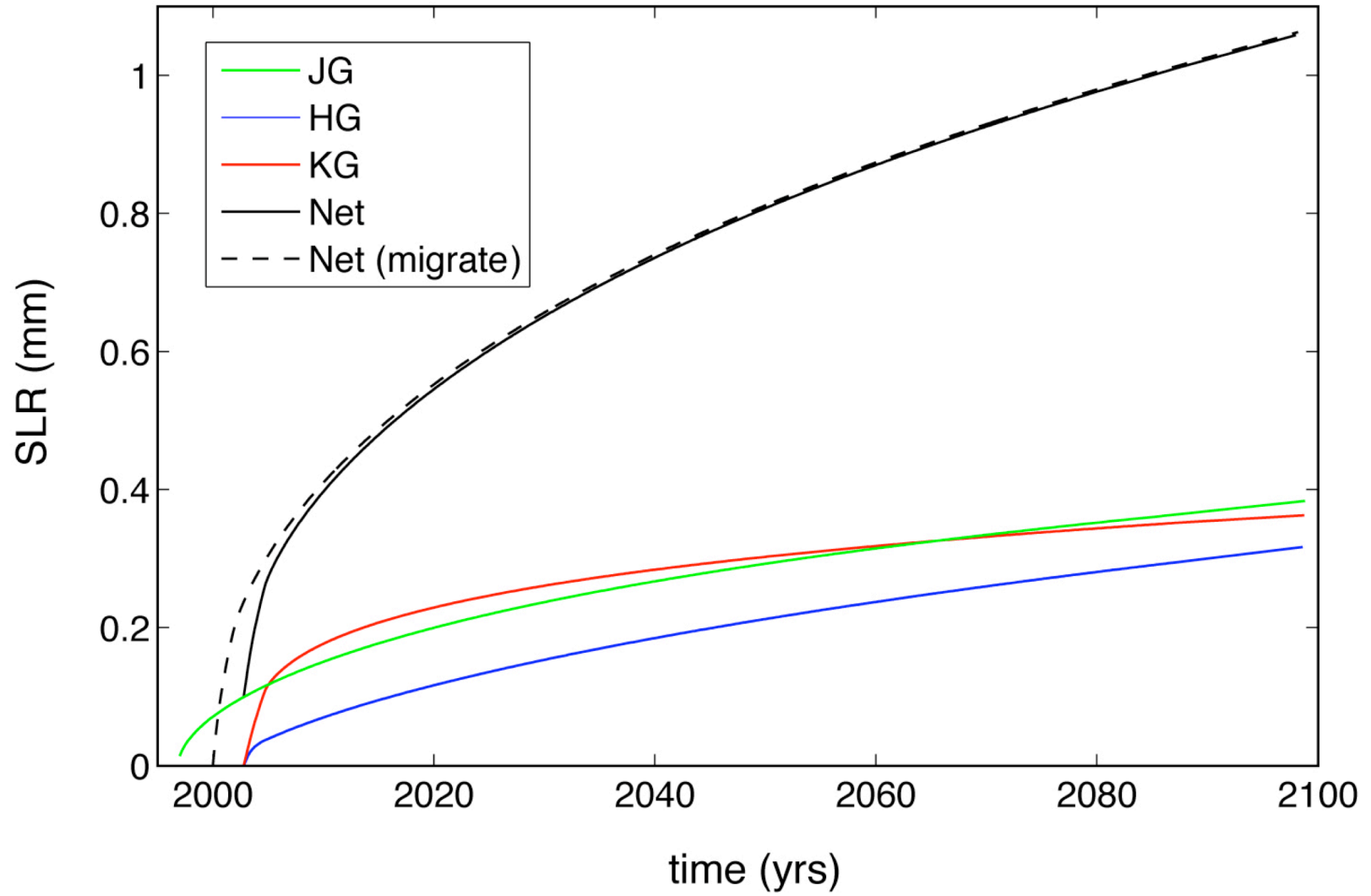
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¹Kang., Hel, obs. from *Howat et al. (Science, 315, 2007)*

cumulative SLR





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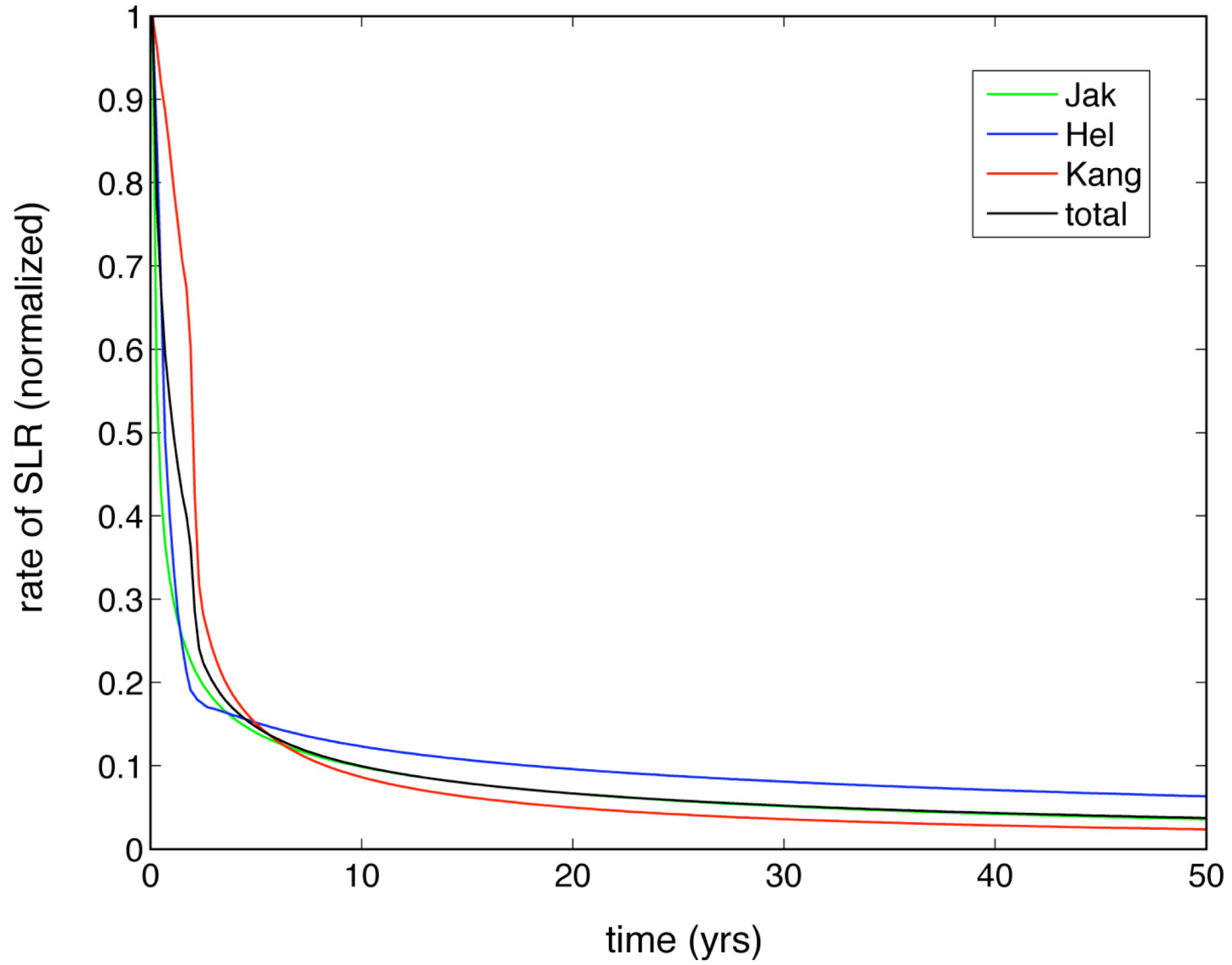
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normalized rate of SLR curves



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Normalized *rate* of SLR curves have the form ...

$$r(t) = r_0 (\alpha t + 1)^{-\frac{1}{2}}$$

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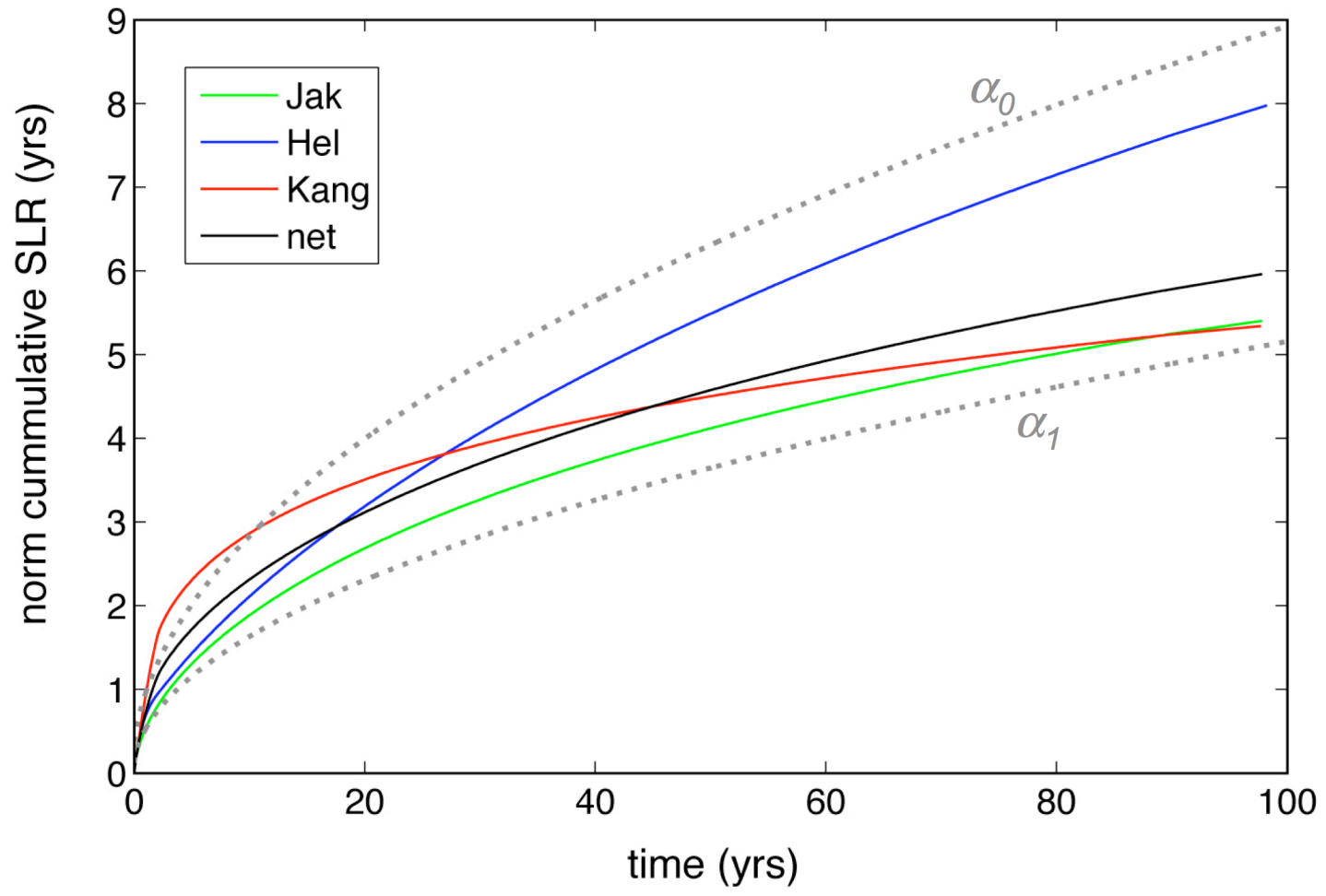
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$$SLR(t) = \int r_0 (\alpha t + 1)^{-\frac{1}{2}} dt \quad \dots = r_0 \left[\frac{2}{\alpha} (\alpha t + 1)^{\frac{1}{2}} + c \right] \quad \dots = r_0 f(t)$$

normalized cumulative SLR



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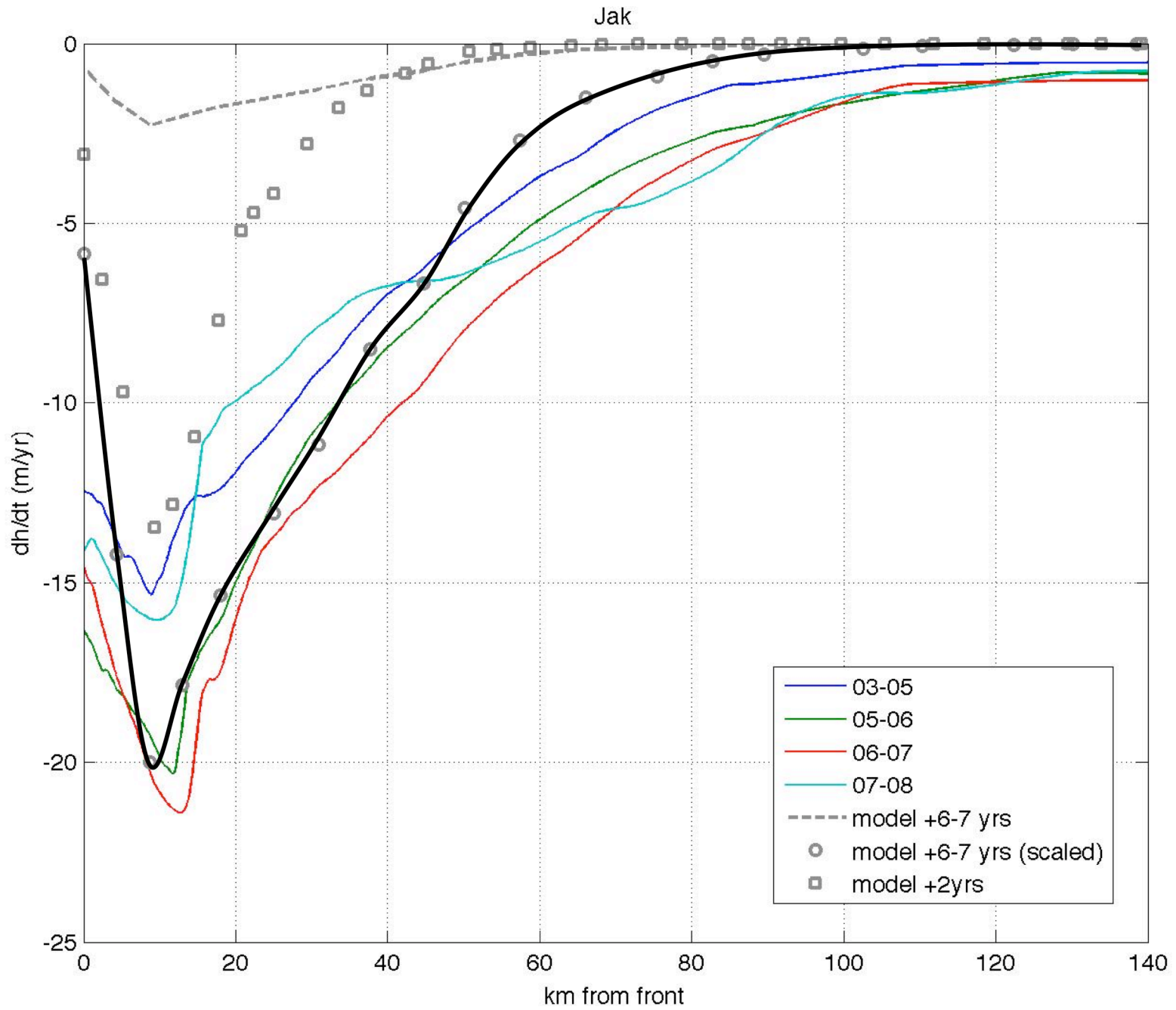
¹Rignot and Kanagar. (*Science*, **311**, 2006); ²Howat et al. (*J. Glaciol.*, **54**(187), 2009); ³Sole et al. (*Cryosphere*, **2**, 2008); ⁴Thomas et al. (*J. Glaciol.*, **55**(189), 2009); ⁵Nick/Vieli et al. (*Nat. Geosc.*, **2**, 2009);

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- Numerous marine-terminating outlet glaciers underwent acceleration and thinning between the mid-to-late '90's^{1,2,3,4} and the present.
- Where examined ^{2,3,4,5}, acceleration and thinning are best explained by perturbations of the type explored here.
- The current modelling captures the long-term affects of this type perturbation.

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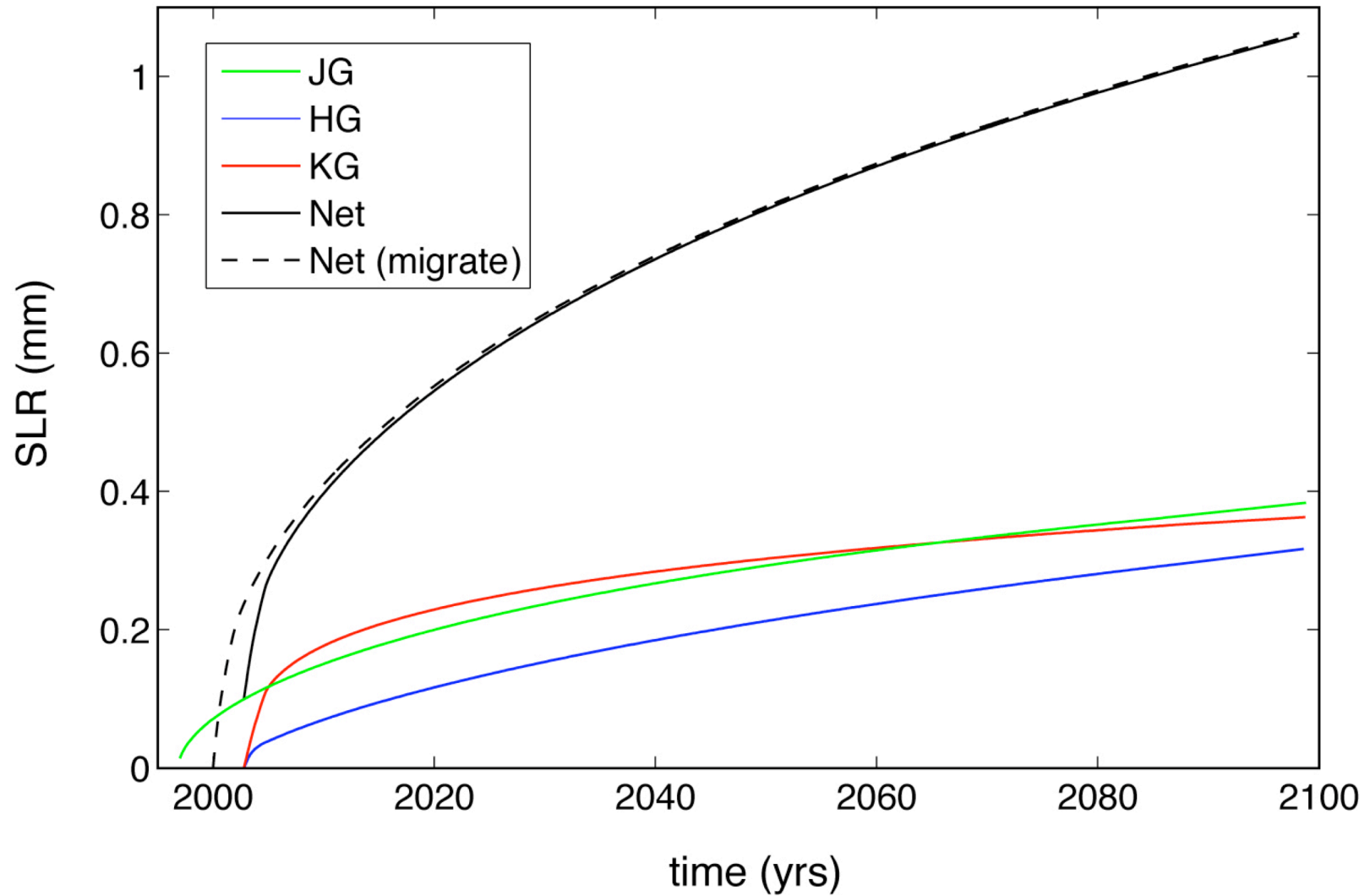


Can we scale these results?

The causative perturbations for this time period can be collapsed into a single, large perturbation.

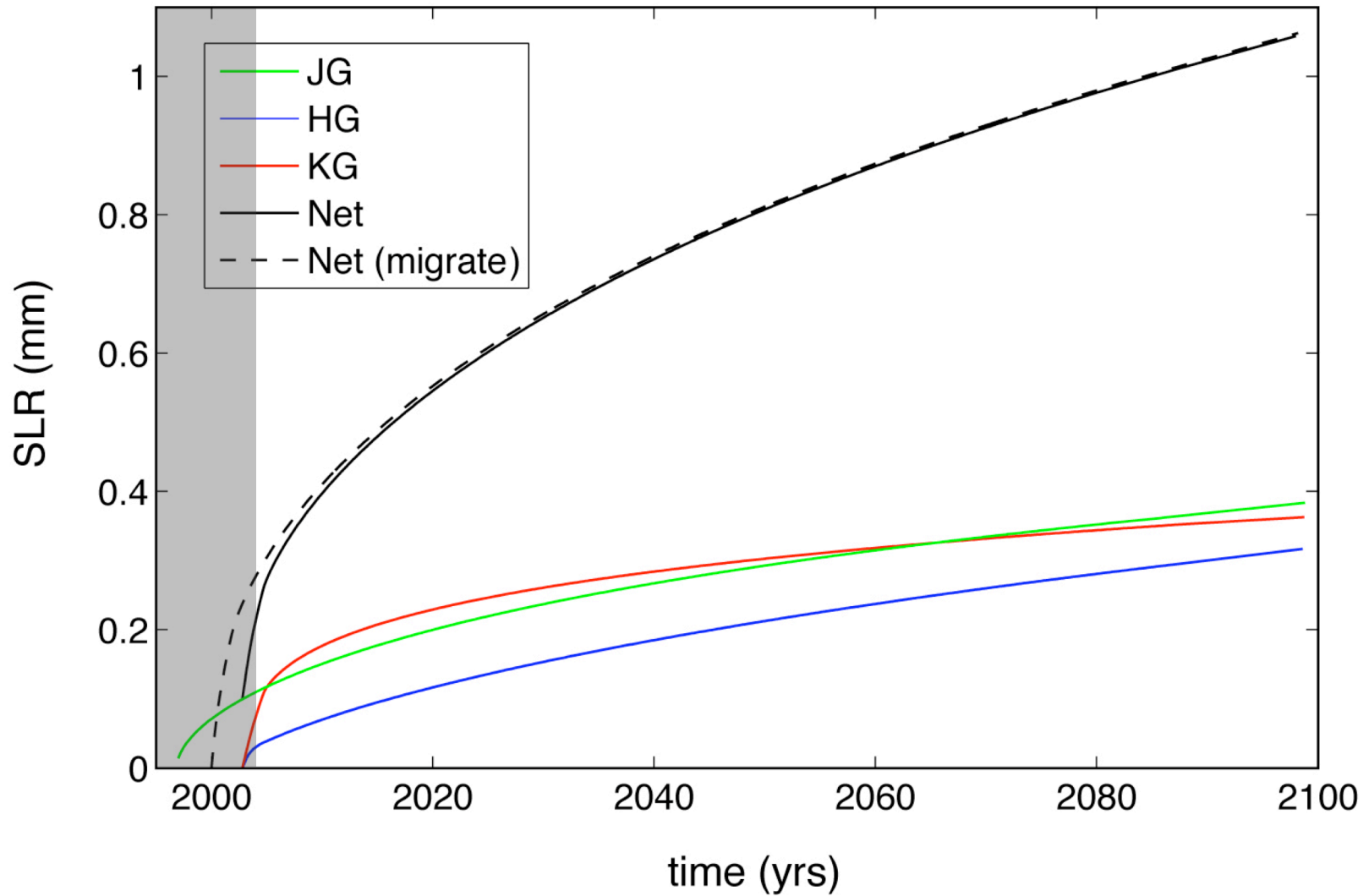
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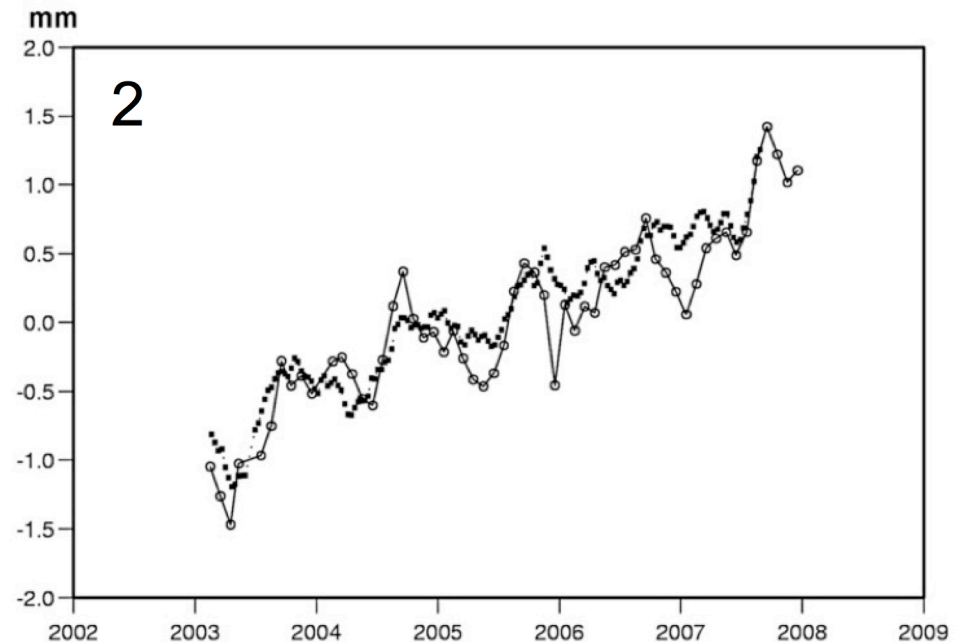
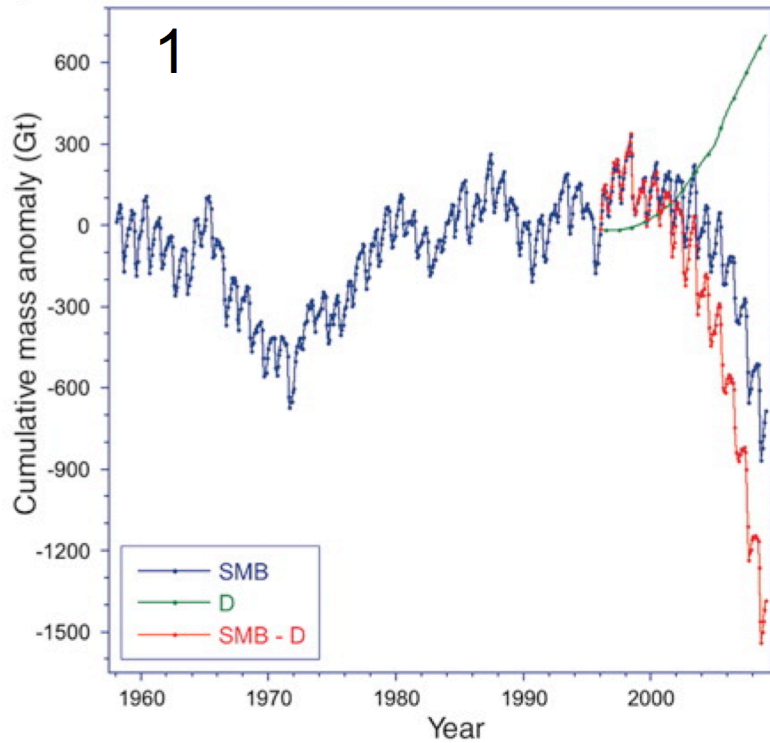
The model gives us an estimate for the **blue** term. How do we get an estimate for the **red** term?

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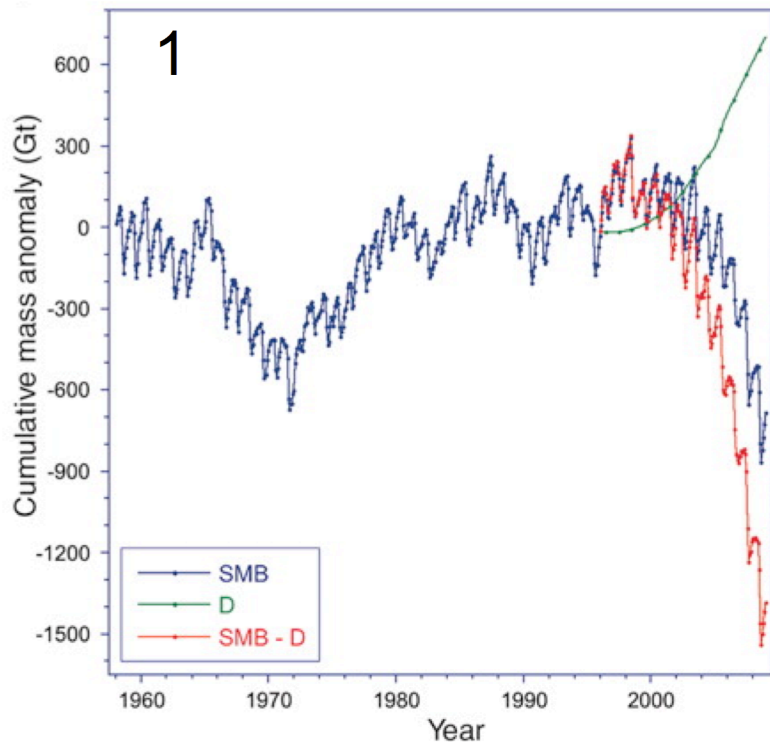


¹van den Broeke et al. (*Science*, **326**, 2009);

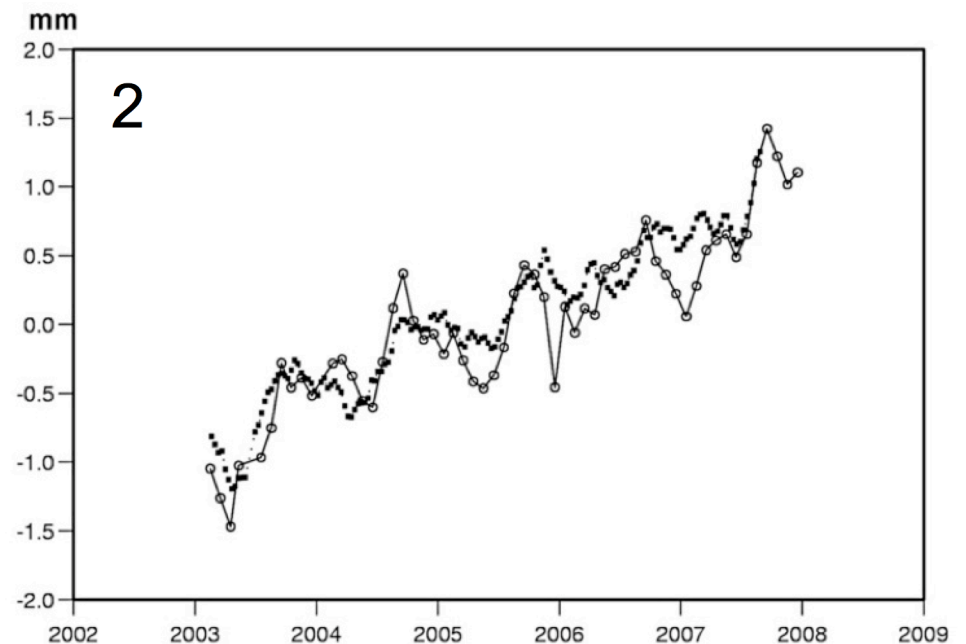
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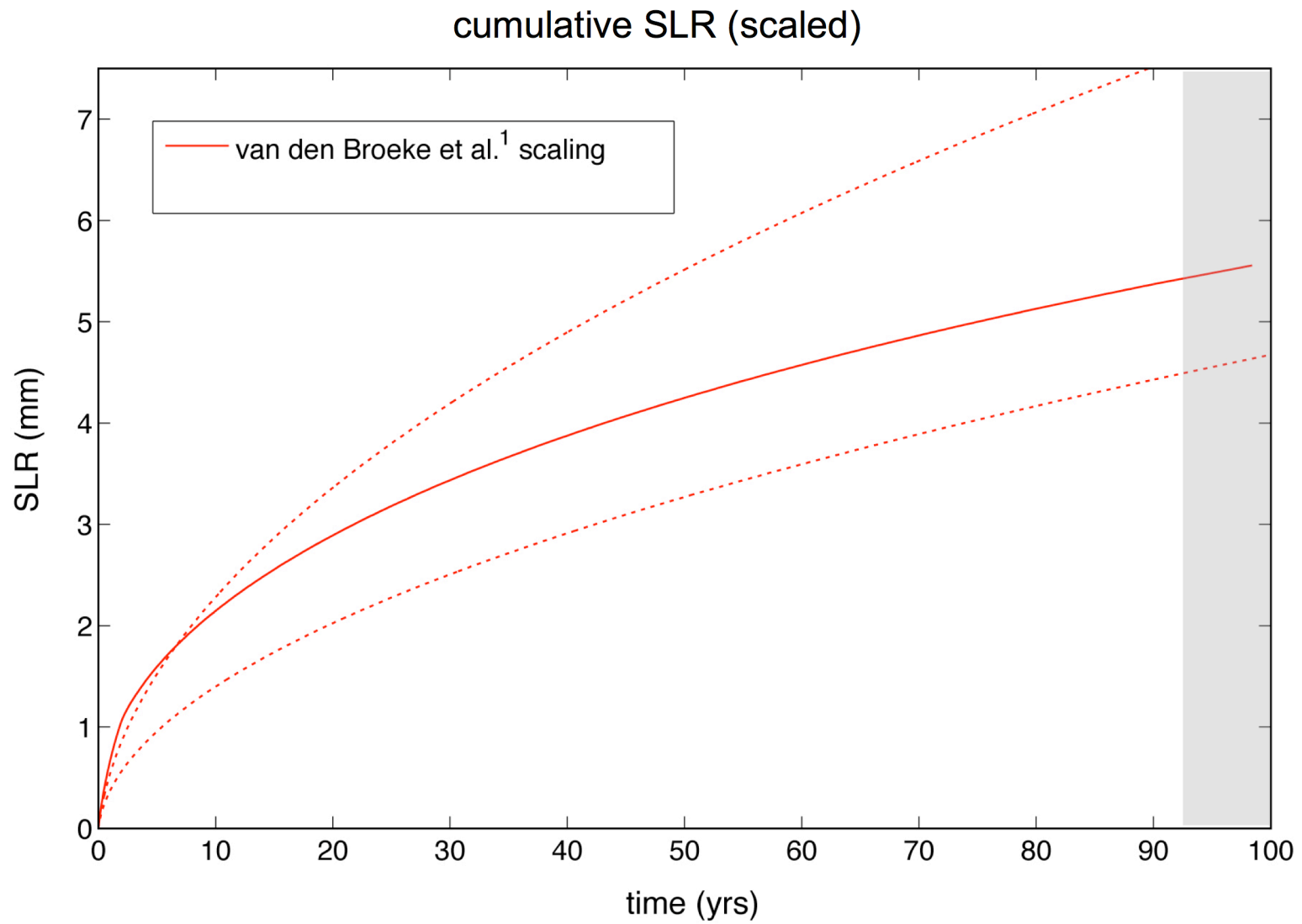
Mean rate of SLR from **2000-2008** is
 0.46 mm yr^{-1} , $\sim 1/2$ due to dynamics
 $\bar{r}_0 \sim 0.23 \text{ mm yr}^{-1}$



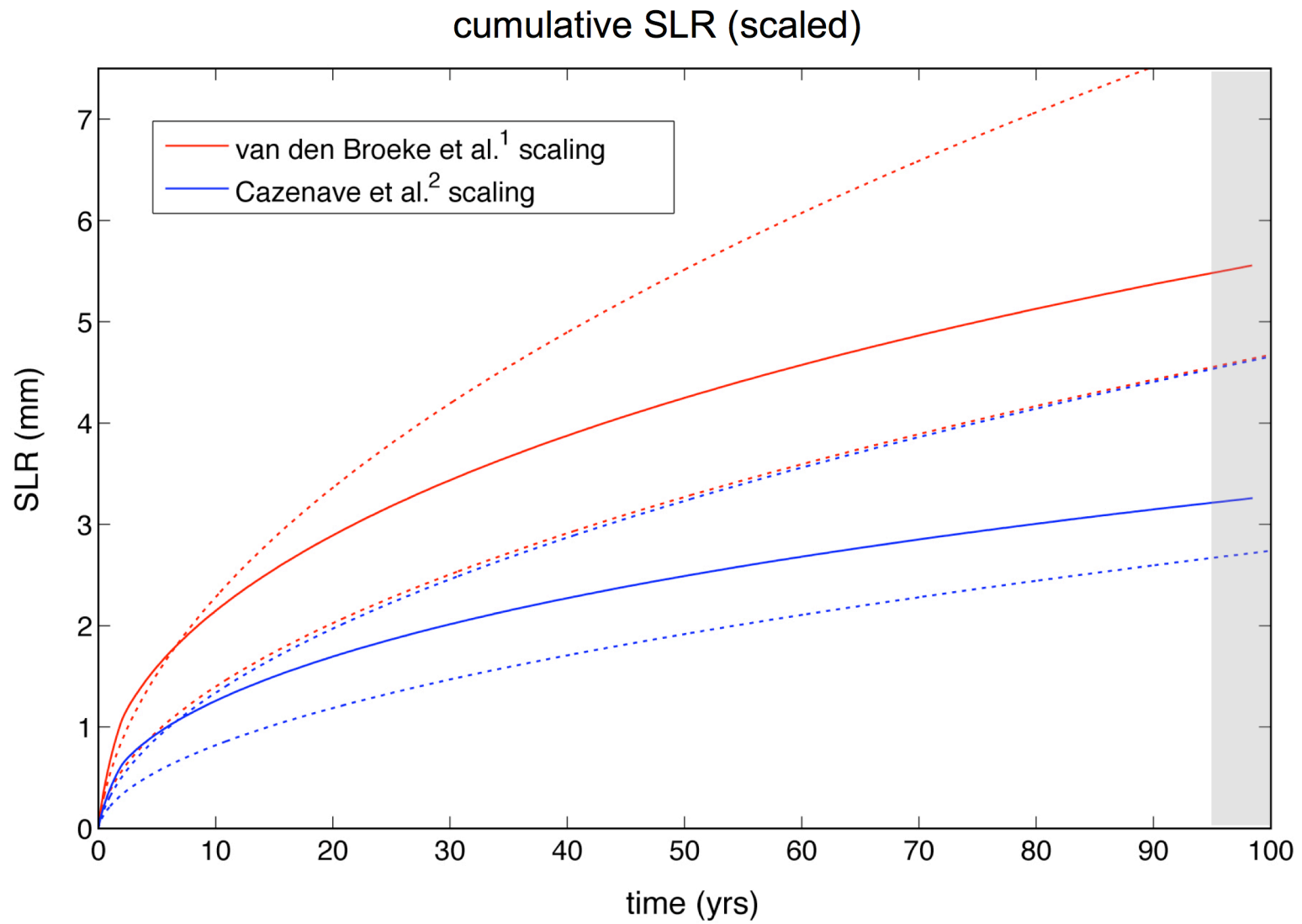
Mean rate of SLR from **2003-2008** is
 0.38 mm yr^{-1} (assume $1/2$ from dynamics)
 $\bar{r}_0 \sim 0.19 \text{ mm yr}^{-1}$

¹van den Broeke et al. (*Science*, **326**, 2009);

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If we scale the results for JG, HG, and KG to the entire ice sheet using estimates for the current rates of dynamic imbalance, we get SLR contributions over the next 100 yrs of 3-5 mm and 5-8 mm.

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Attributing all SLR from the 1st ~3 yrs (e-folding time) to the initial perturbation, 75% of the SLR in 100 yrs is from the diffusive response, yet to come.

END

