ENSO Diversity in the NCAR-CCSM4

Antonietta Capotondi

NOAA/ESRL and University of Colorado/CIRES

What do we mean with ENSO diversity?

Warm pool (WP)

Cold Tongue (CT)

Others



Kug et al. 2009: "Two types of El Niño: Cold Tongue El Niño and Warm Pool El Niño" (NOAA-ERSST 1970-2005) Selection of events based on Niño3 and Niño4 indices

Definitions

"Dateline El Niño" (Larkin and Harrison 2005)

"El Niño Modoki" (Ashok et al. 2007)

"Central Pacific El Niño" (Kao and Yu 2009)

"Warm Pool El Niño" (Kug et al. 2009)

The types of El Niño may be the manifestation of a broader and perhaps continuum of ENSO flavors.

Why is ENSO diversity important?

Different locations of maximum warming are associated to different teleconnection patterns:

- It has been suggested as a forcing for the southernmost lobe of the NPO, which, in turn, appears to force the North Pacific Gyre Oscillation (NPGO, Di Lorenzo et al. 2008)
- It has been linked to changes in tropical cyclone activity (Kim et al. 2009), shifts in precipitation patterns (Weng et al. 2009), and warming in Antarctica (Lee et al. 2010, Ding et al. 2011)

It has been suggested that CP warming may become more pronounced with global warming (Yeh et al. 2009)

Objectives

- Use 500 years (800-1299) of Pre-industrial Control integration to assess how well CCSM4 simulates a broad spectrum of El Niño warming, adopting various approaches proposed in previous studies (Niño3 and Niño4 indices, EOF analysis)
- Use the model to have a more robust assessment of the processes involved in different types of events
- Examine teleconnection patterns associated with different patterns of warming

Compare model with SODA ocean reanalysis (1958-2007). Forced by ERA-40 (1958-2001) and QuickSCAT (2002-2007) Resolution: 0.4(lon) x 0.25(lat) x 40 vertical levels

CMIP3 analysis

SST obs: ERSST V.2 (1910-2009), Using JFM Niño3 and Niño4 indices



Zonal SST profiles



Precipitation - CMIP3 models



Models with a better separation of SST maxima also have shifted precipitation patterns. The lack of separation is attributed to the inability of the convection system to shift eastward due to the cold tongue bias.

Kug and Ham 2011

Can CCSM4 reproduce warming over a broad range of longitudes?



Zonal SST profiles



Can CCSM4 reproduce shifted precipitation patterns?



Diversity of individual events

CCSM4

WP El Nino





WP EI Nino



SST composites

CCSM4

SODA

SST (standard inds)

DA SST



WP

СТ

Thermocline depth (Z15) composites

CCSM4 _{Z15} SODA

SODA Z15 (Standard inds)



СТ





-12 -9 -6 -3 0 3 6 9 12 15 18 21 24 27 30

Evolution of zonally averaged thermocline depth



24

21

1B ·

15

12 -

205

CCSM4

-15

SODA

9

6

3

0

-3

Modified Niño3 and Niño4 indices (centers of Niño3 and Niño4 regions displaced 20° westward)



SST spatial patterns



Modified Niño3 and Niño4 indices Z15 evolution

Z15 odif)



Composite of Z15 depth anomalies







dified inds)



WP

СТ

EOF approach

SODA

CCSM4





EOF approach

SODA

SODA (1958-2007)



11.7%

7.1%







How are the two leading SST EOFs related to Z15 variability?

SST

EOF1 63.68%

20N

20S

20N

20S

20N

20S

Z15

CCSM4 (800-1299)





Z15 CCSM4 (800-1299)

CCSM4 PCs

PC1-PC2

PC1-PC3



PCs lag correlation

CCSM4

SODA









Interpretation of lag-correlation











Spectral characteristics



10⁰

Conclusions (1)

CCSM4 shows a broad range (continuum?) of longitudes where maximum warming occurs during El Niño events in relative good agreement with SODA.

Warming in the eastern Pacific is associated with large thermocline depth anomalies, and recharge-discharge processes. The thermocline involvement becomes weaker for events in which the warming occurs further west. For events with warming in the Warm Pool region thermocline anomalies are negligible, and the SST anomalies are longer-lived.

The WP events emerge as the third EOF of SST, with the first two EOFs representing the surface manifestation of the thermocline evolution during recharge-discharge processes. This mode appears to have relative more power at low frequencies.

Conclusions (2)

Other aspects to consider:

- 1. Are there atmospheric precursors, or oceanic preconditioning that lead to a specific ENSO flavor?
- 2. How do the teleconnections change as a function of the longitude of warming?

What aspects of ENSO diversity emerge from observations?

Sea surface height composites GODAS 1980-2005

Composite evolution of zonally averaged sea surface height





Kug et al. 2009