

# A Stochastic Case Study: Monte Carlo ice flow modeling projects a new stable configuration for Columbia Glacier, Alaska, by c. 2020

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**LIWG Workshop**  
**February 2012**

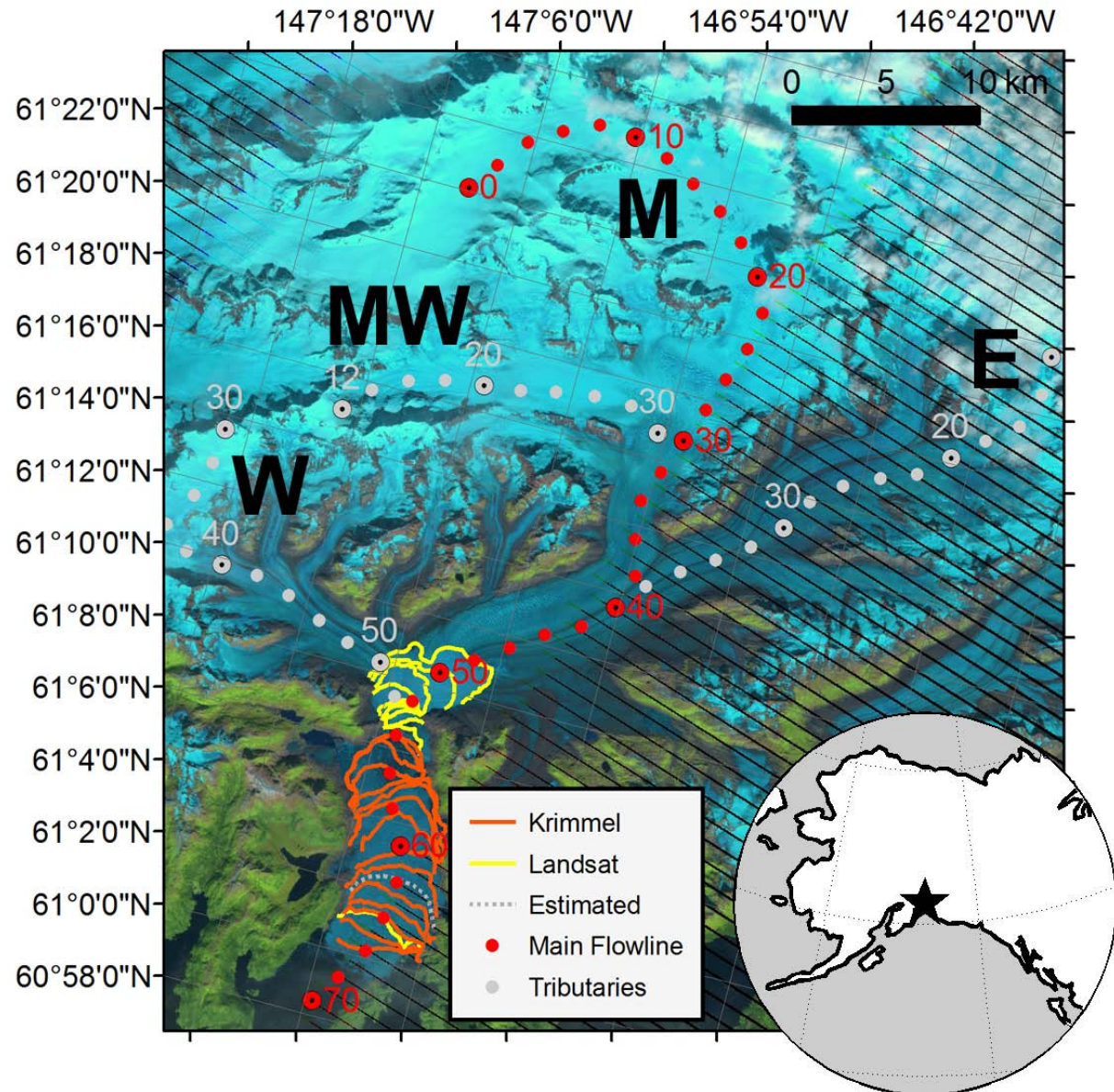
Online in *The Cryosphere Discussions* (tc-2012-15)



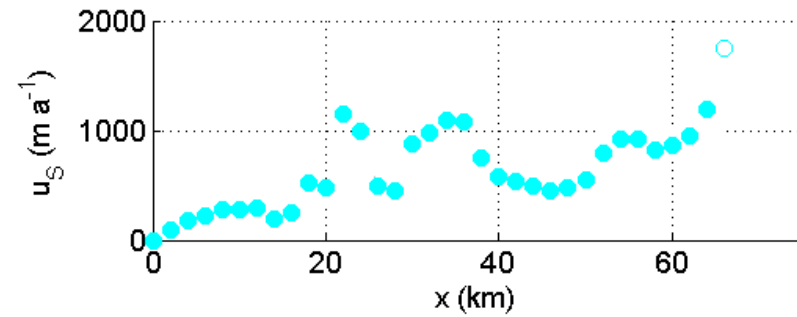
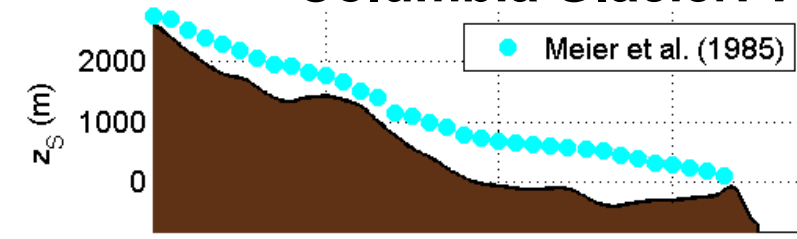
# Columbia Glacier: Highly transient modeling target

Projecting uncertainty benefits from models being executed many times ( $n > 1000$ ) under slightly different conditions.

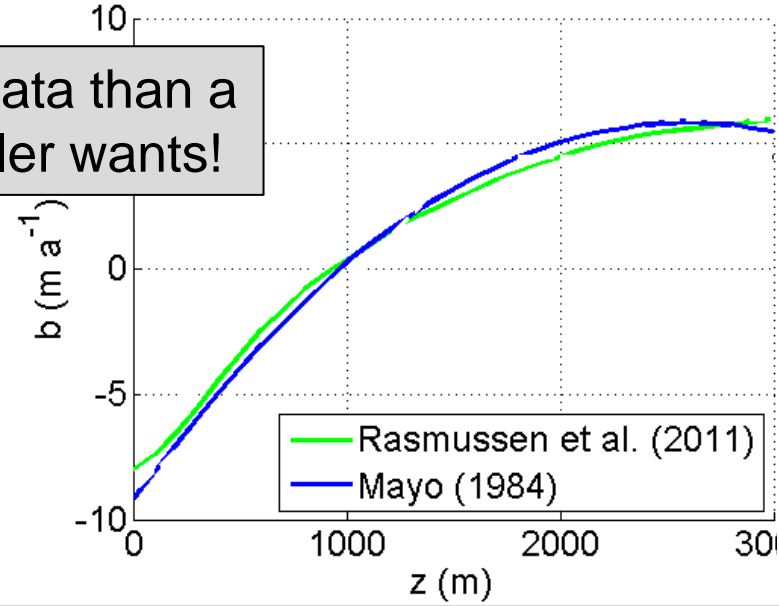
- ~ 18 km of retreat and ~ 100 km<sup>2</sup> loss of ice covered area since 1983.
- ~ 0.6 % of global sea level rise over the 2003-2007 period.



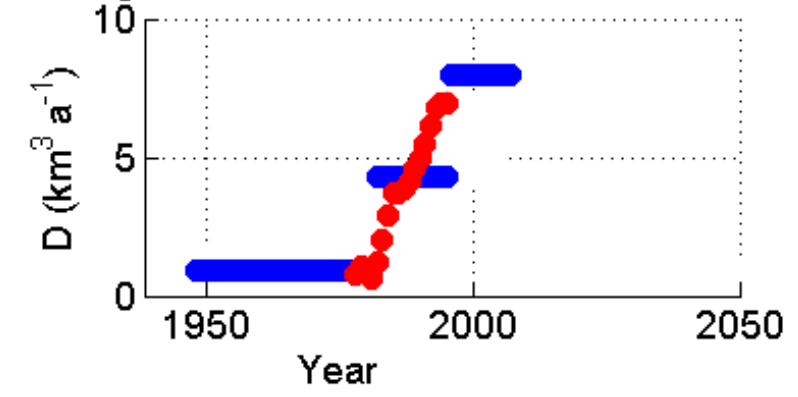
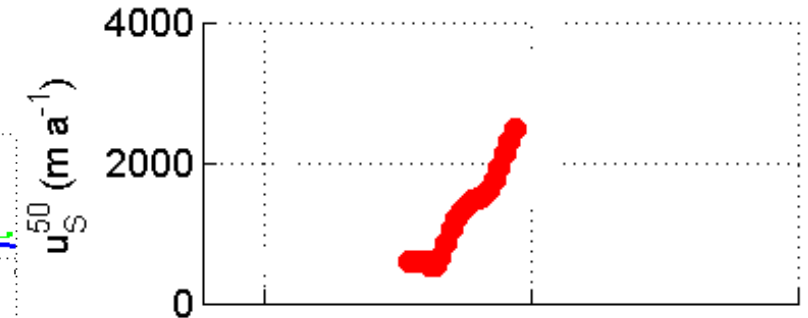
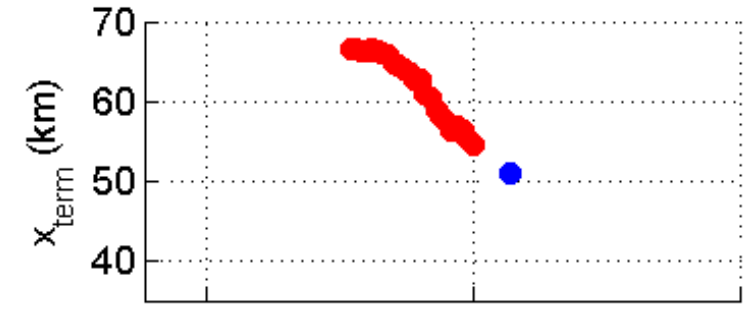
# Columbia Glacier: Well-documented modeling target



More data than a modeller wants!



- Krimmel (2001)
- Rasmussen et al. (2011)
- Mayo (1984)



## “Typical” 1D (depth-integrated) ice flow model

$$\frac{\partial H}{\partial t} = b - \frac{1}{w} \frac{\partial Q}{\partial x}$$

$$Q = Fw \left( u_b H + \frac{2A}{(n+2)} \left( \rho g \left| \frac{\partial z_s}{\partial x} \right| \right)^{(n-1)} \tau H^{(n+1)} \right)$$

$$\tau = -\rho g H \frac{\partial z_s}{\partial x} + 2 \frac{\partial}{\partial x} (H \bar{\tau}'_{xx})$$

$$0 = \bar{\tau}'_{xx}{}^3 \left\{ 2 \frac{\partial z_s}{\partial x} \left( \frac{\partial H}{\partial x} - \frac{\partial z_s}{\partial x} \right) + H \frac{\partial^2 z_s}{\partial x^2} - \frac{1}{2} \right\} + \bar{\tau}'_{xx}{}^2 \left\{ \tau \left( \frac{2}{3} \frac{\partial H}{\partial x} - \frac{3}{2} \frac{\partial z_s}{\partial x} \right) \right\} + \dots$$

$$\bar{\tau}'_{xx} \left\{ \tau^2 \left( 3 \frac{\partial z_s}{\partial x} \frac{\partial H}{\partial x} + \frac{3}{2} H \frac{\partial^2 z_s}{\partial x^2} - 2 \left( \frac{\partial z_s}{\partial x} \right)^2 - \frac{1}{6} \right) \right\} + \tau^3 \left( \frac{2}{5} \frac{\partial H}{\partial x} - \frac{1}{4} \frac{\partial z_s}{\partial x} \right) + \frac{1}{2A} \frac{\partial u_b}{\partial x}$$

Upstream BC: Type 2 (specified flux)  $Q = 0$  at flow divide

Downstream BC: Type 1 (specified head) of observed ice cliff height

Top BC: Statistical param. of surface mass balance (wide param. space)

Bottom BC: Statistical param. of basal sliding velocity (wide param. space)

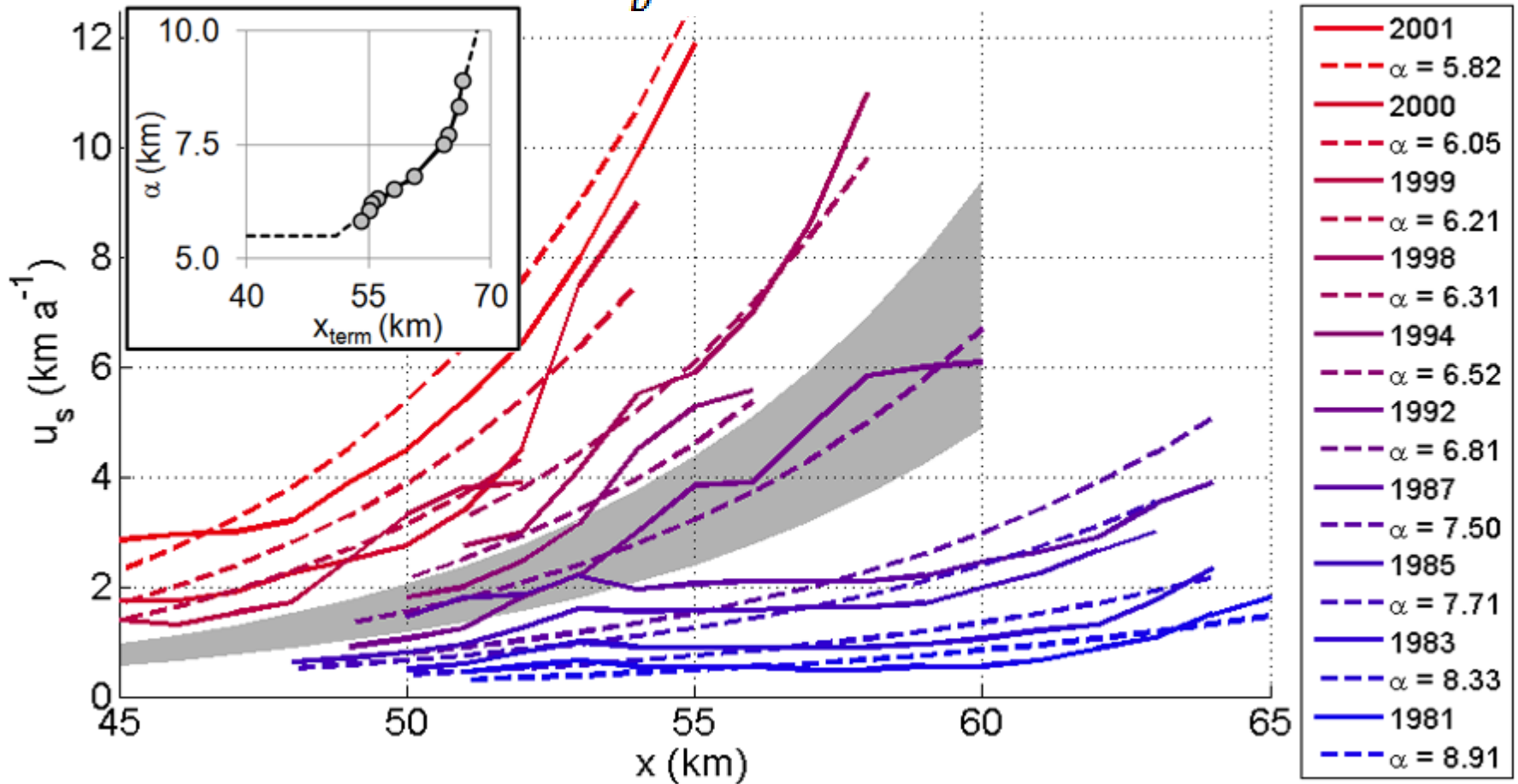




# Statistical Parameterization: Basal sliding velocity

$$u_b = k e^{(x/\alpha)}$$

No physics:  $f(x)$

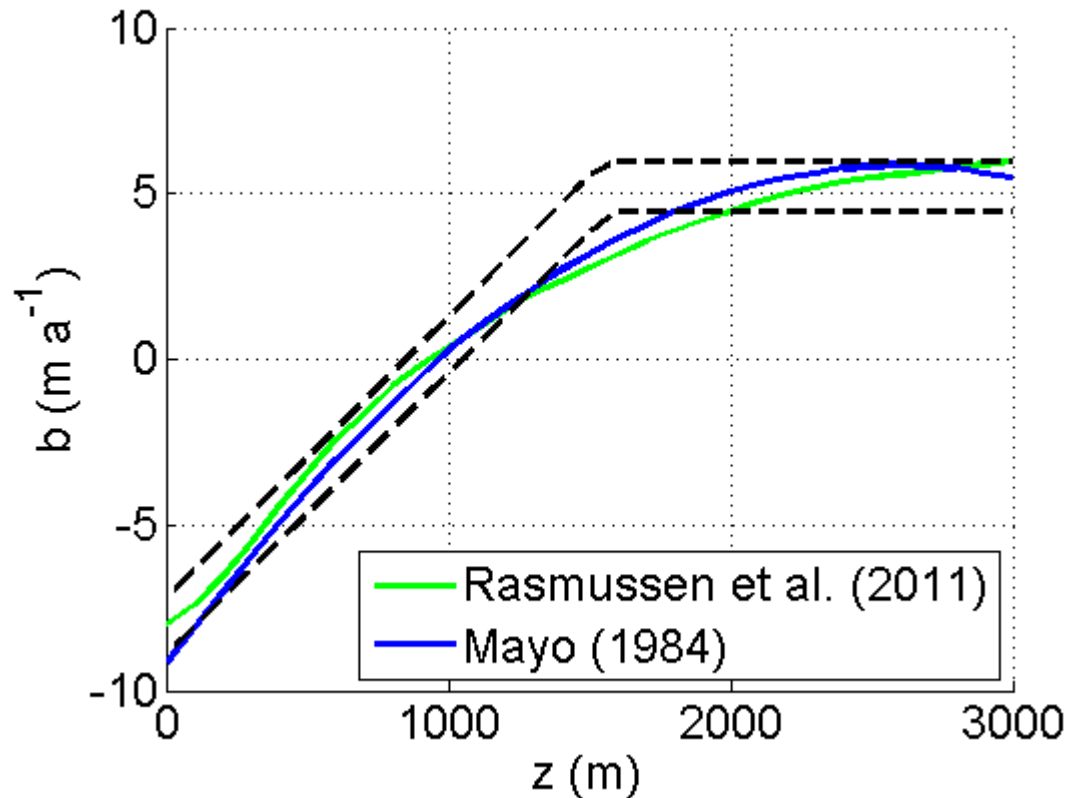


Length scale alpha varies within this range each simulation



# Statistical Parameterization: Surface mass balance

$$b = \begin{cases} \gamma(z_s - z_{ela}) & \text{if } b < b_{max} \\ b_{max} & \text{if } b \geq b_{max} \end{cases}$$



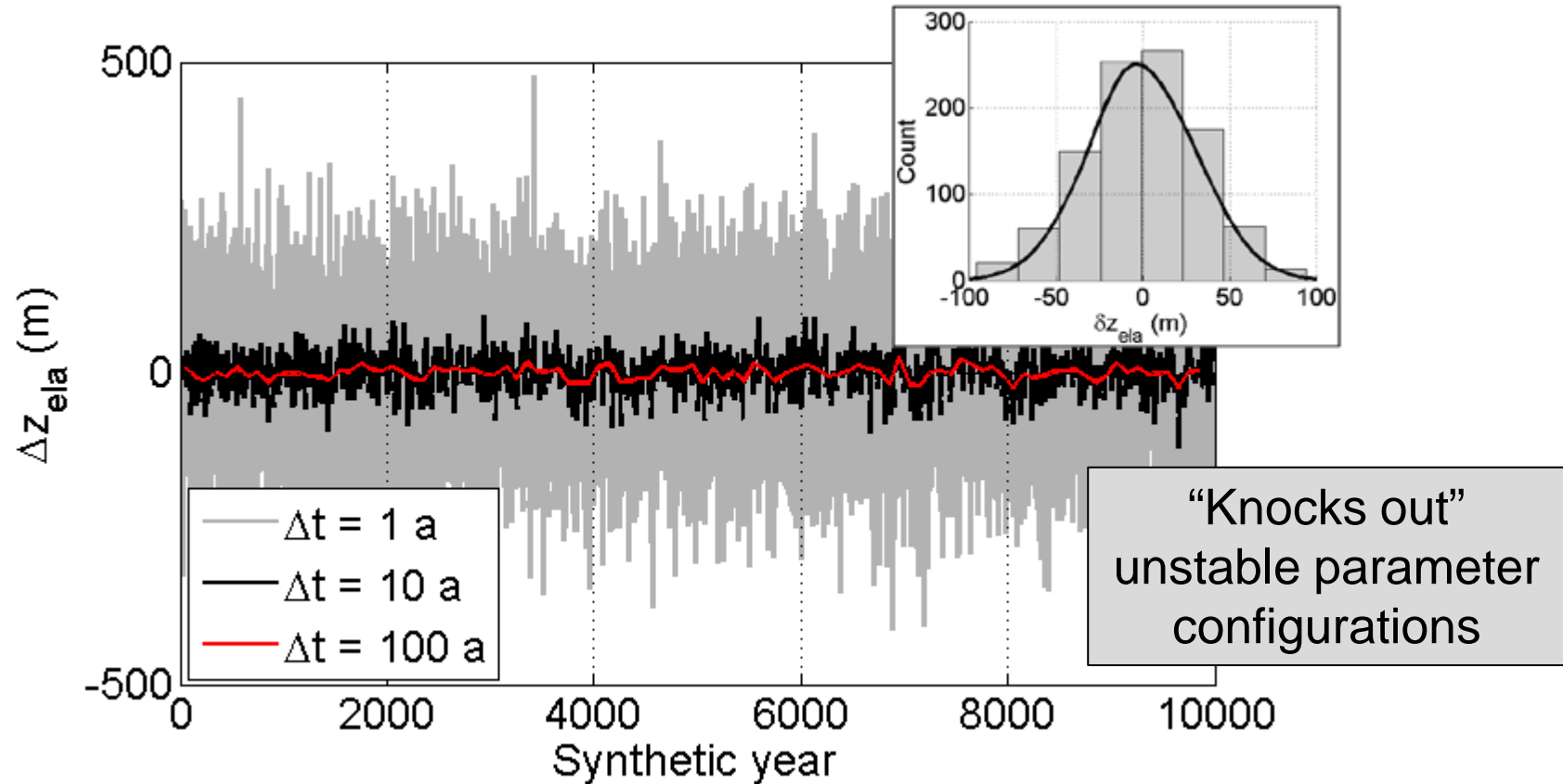
...also no physics:  $f(z)$

Surface mass balance varies within this range each simulation



# Statistical Parameterization: Climatic variability

$$\frac{\Delta z_{ela}}{\Delta t} = \left( \frac{\Delta T}{\Delta t} \right) \left( \frac{\Delta T}{\Delta z} \right)^{-1}$$

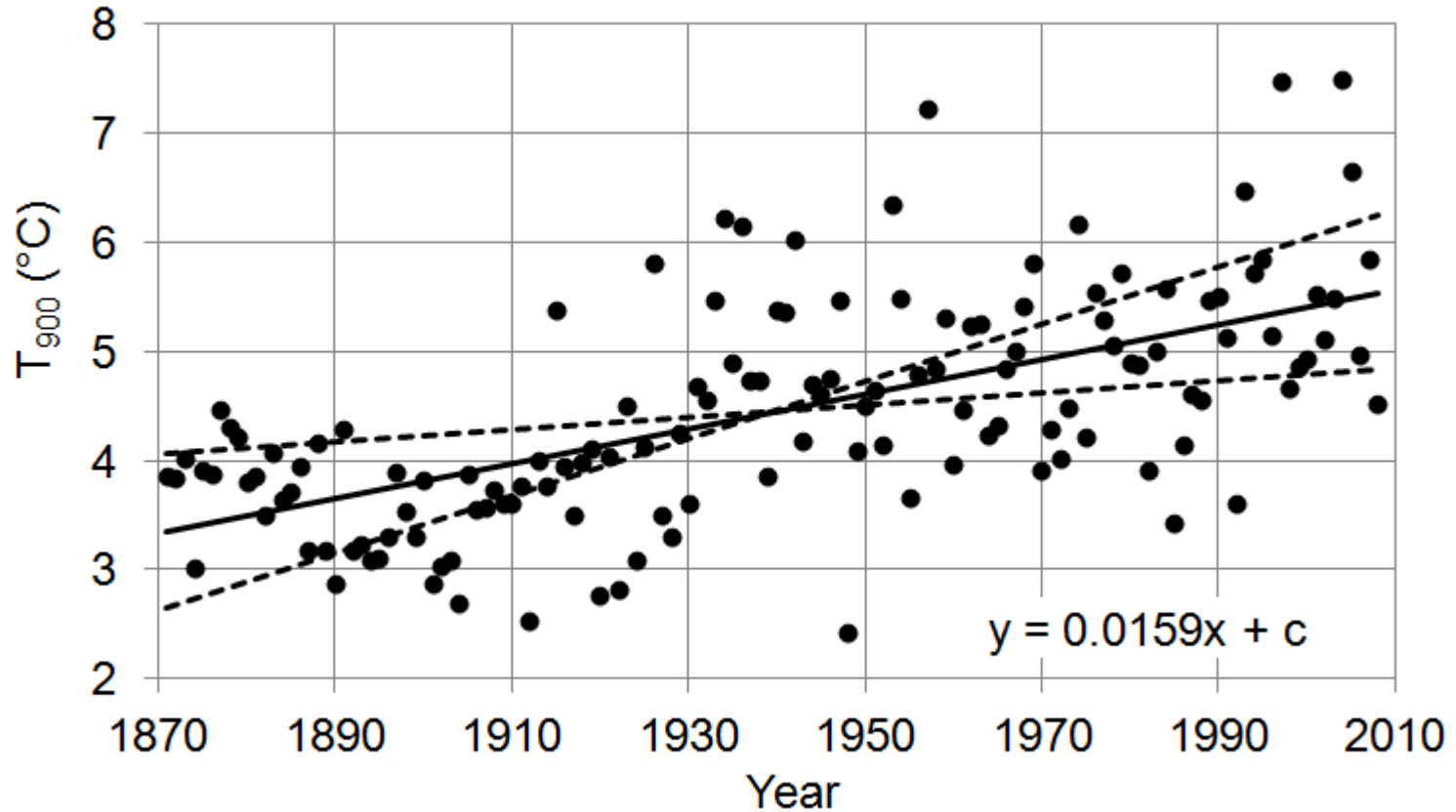


Randomly perturb ELA every decade



# Statistical Parameterization: Climate forcing

$$\frac{\Delta z_{ela}}{\Delta t} = \left( \frac{\Delta T}{\Delta t} \right) \left( \frac{\Delta T}{\Delta z} \right)^{-1}$$

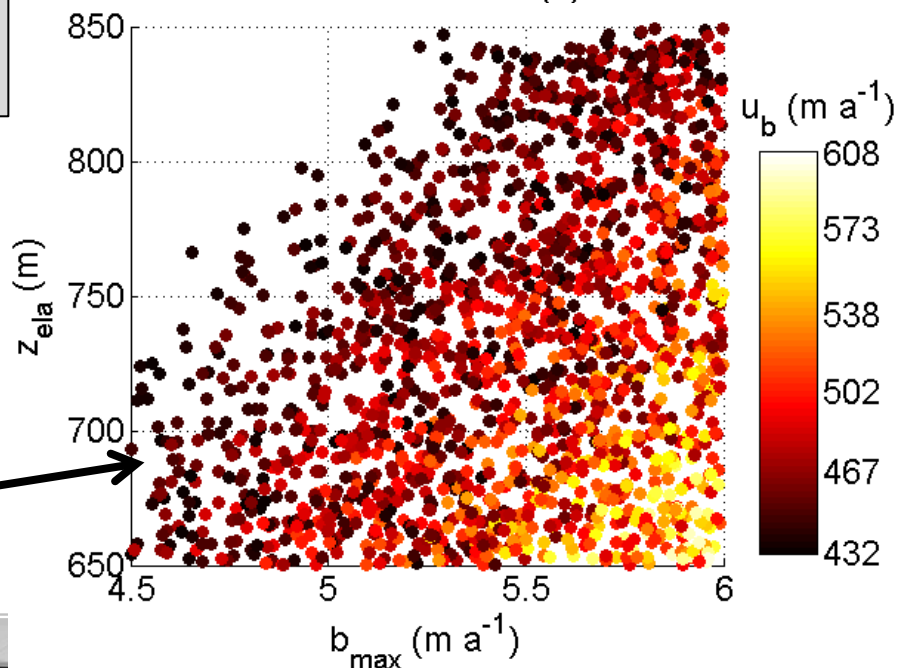
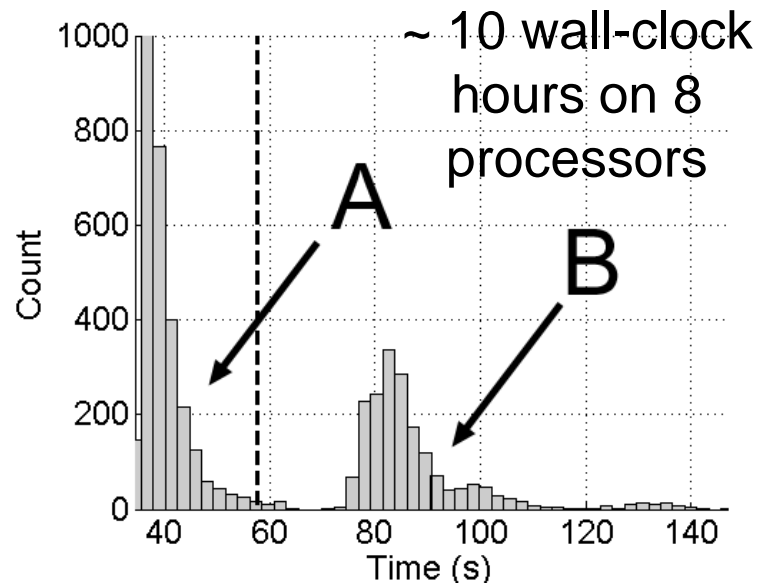
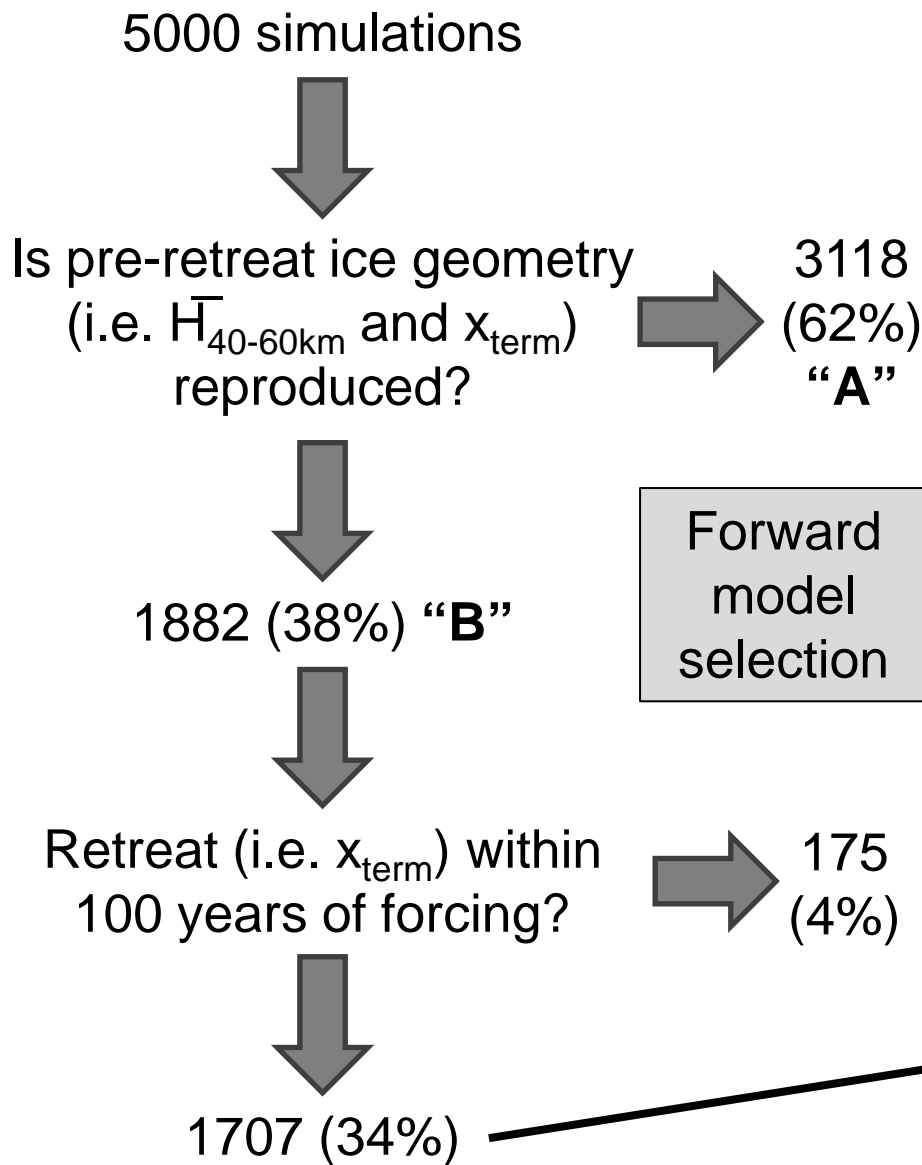


Randomly prescribe forcing from this range of 900mb air temperature increase

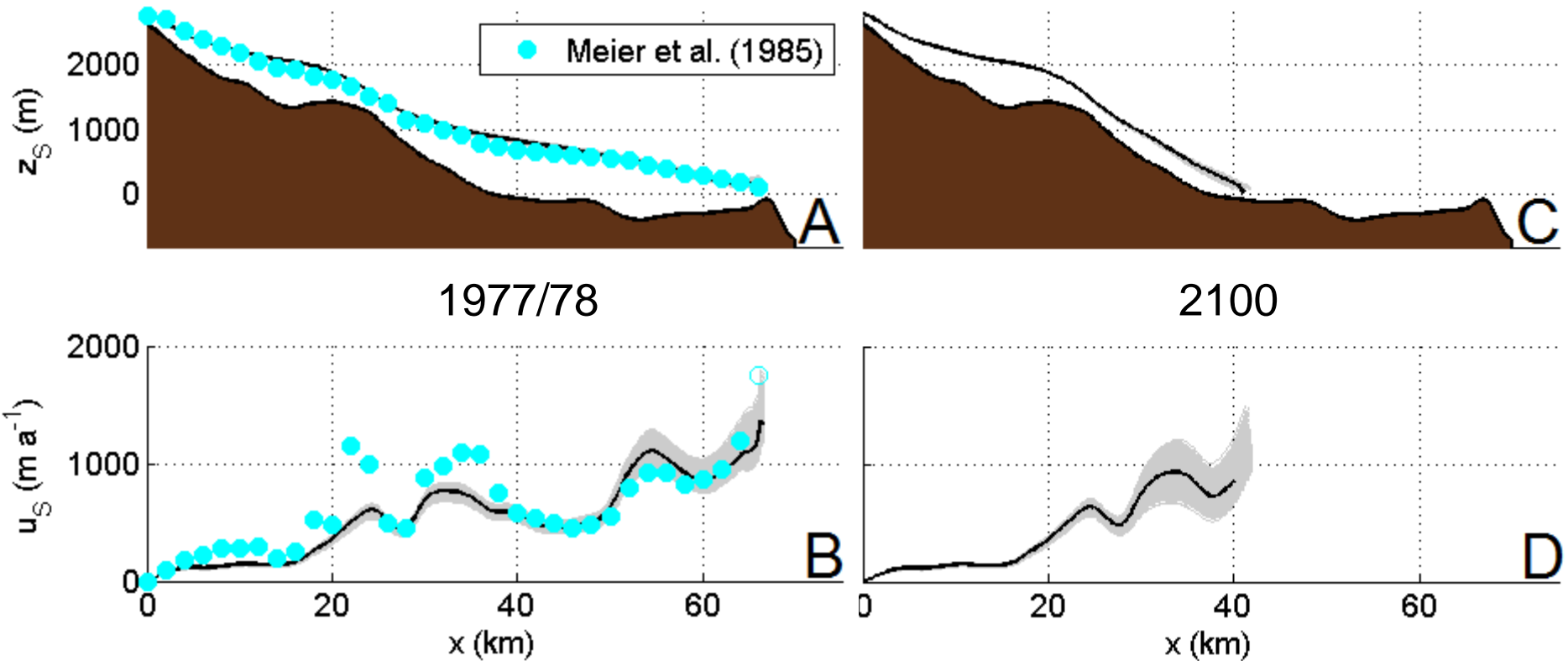




# Ensemble selection: A diverse population of Columbia Glaciers



# Results: Selected ensemble

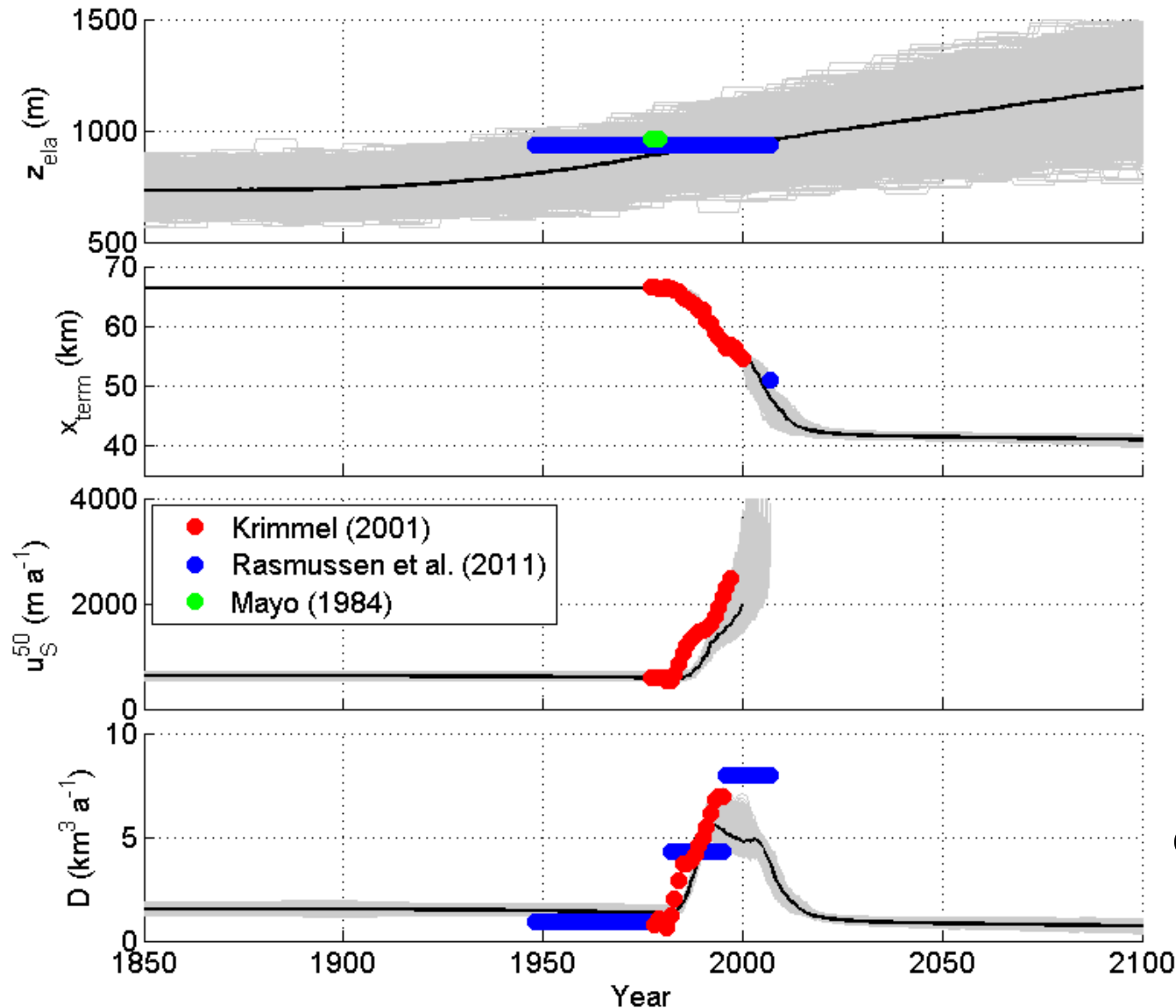


The momentum approx. fails in some spots, but overall good fit...

...confidence in future ice geometry: 83 % of ice loss occurred by 2007



# Results: Selected ensemble



Reasonable SMB  
and climate  
variability

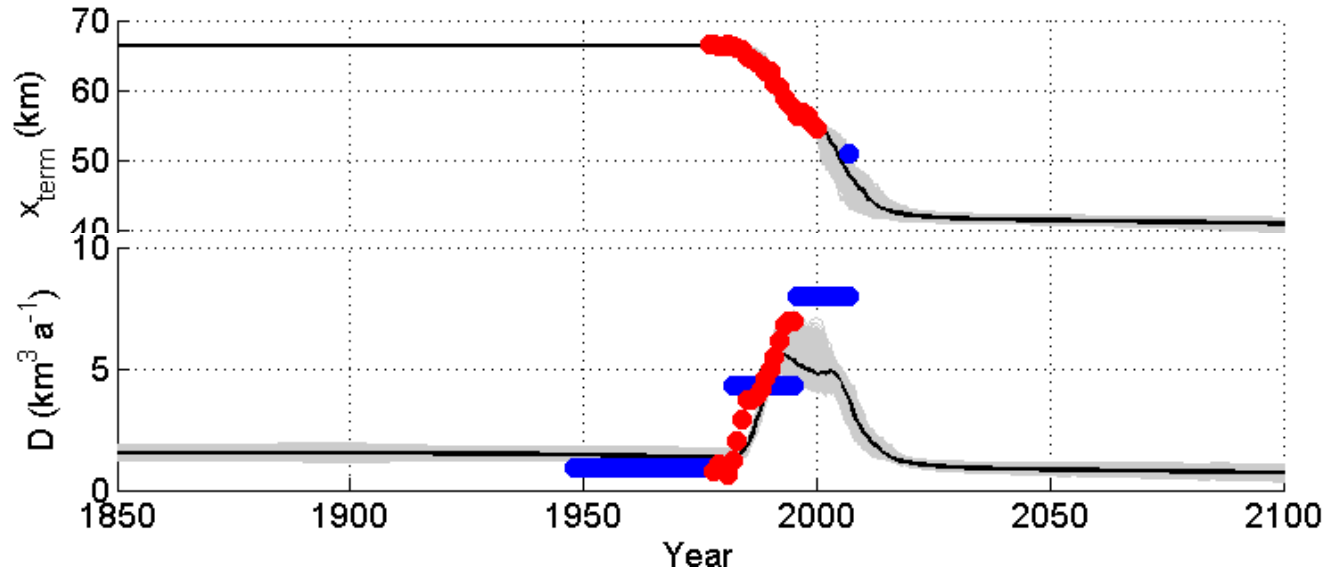
New stable  
geometry c. 2020

Not bad for blunt  
statistical  
parameterization

Return to dynamic  
equilibrium iceberg  
calving c. 2020



## Discussion points: Applicability beyond Columbia Glacier



High likelihood of spatially (*Venteris, 1997*) temporally (*Meier et al., 1994*) transient ice density at Columbia Glacier:

**Is “swelling” of remaining ice suppressing the apparent retreat rate?  
How can the momentum balance be modified to accommodate this?**

Highly transient iceberg calving can “turn off” just as quickly as it “turned on”:  
**Can we use rapid response time (i.e.  $\sim 40$  a) to model future ice volume in steady-state with future climate?**



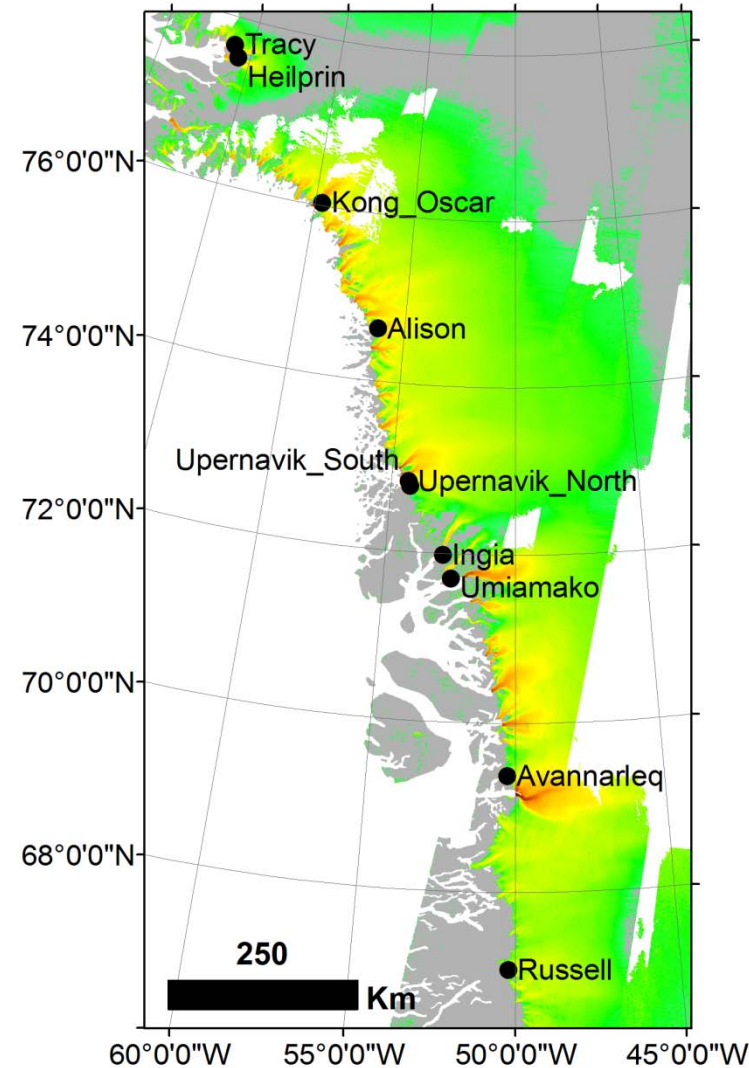
# Future work: West Greenland outlet glaciers

NASA ROSES

“Integrating IceBridge and ICESat data with a Monte-Carlo modeling framework to constrain the mechanisms of recent acceleration in West Greenland outlet glaciers.”

H. Rajaram, T. Phillips and W. Colgan

- 2D (cross-sectional) thermo-mechanical
- More selection filters (velocity,  $dH/dt...$ )
  - Five acceleration mechanisms:
    - (i) Meltwater-enhanced basal sliding
    - (ii) Loss of terminus back-stress
    - (iii) Adjustment to surface ablation
    - (iv) Decreased effective basal pressure
    - (v) Cryo-hydrologic warming





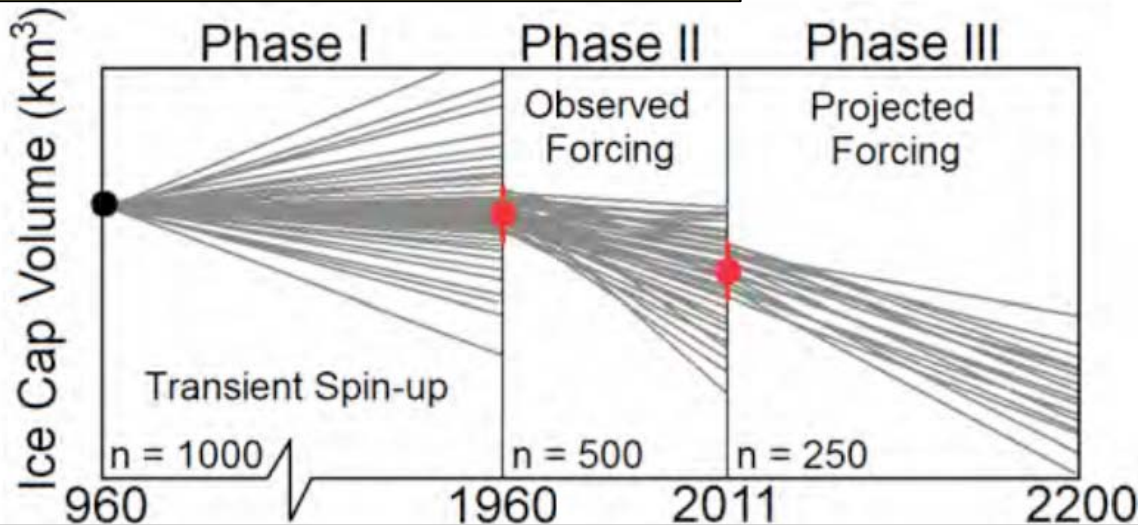
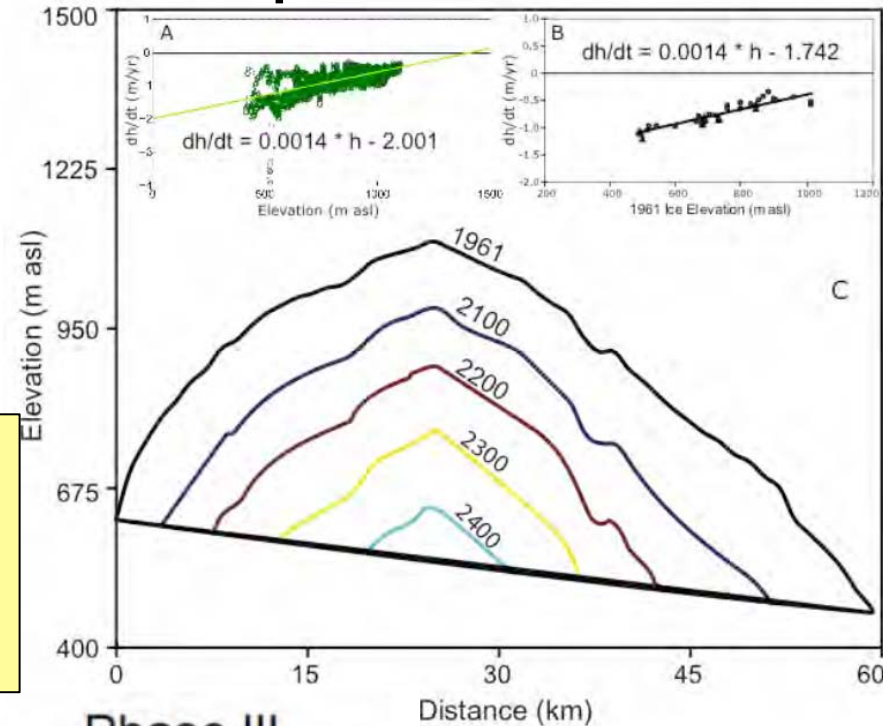
# Future work: Barnes Ice Cap

NSF OPP (Pending)

“The End of the Ice Age: Modeling the Disappearance of the Barnes Ice Cap”

G. Miller and W. Colgan

- 3D thermo-mechanical model (Elmer?)
  - Inverse approach to Wisconsin Ice
- Radiocarbon dates: disappearance will be a once in  $\sim 2.5$  Ma event?



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