A finite element solver for icesheet dynamics to be integrated with MPAS



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Outline

- Introduction of finite element ice-sheets component
- Integration of finite element ice-sheet component with MPAS
- Results on ISMIP-HOM and Greenland/Antarctica geometries
- Non linear solvers: features and issues

Ice Sheet Modeling

Main components of an ice model:

Ice flow equations (momentum and mass balance)

$$\begin{aligned} -\nabla \cdot \sigma &= \rho \mathbf{g} \quad \text{and} \quad \nabla \cdot \mathbf{u} = 0, \\ \text{with } \sigma &= \tau - pI = 2\mu(\dot{\varepsilon}) \ \dot{\varepsilon} - pI, \\ \text{where } \mu \text{ viscosity, } \dot{\varepsilon} \text{ shear rate} \end{aligned}$$

Model for the evolution of the boundaries (thickness evolution equation) $\frac{\partial H}{\partial t} = H_{flux} - \nabla \cdot \int_{\Sigma} \mathbf{u} \, dz$

Temperature equation

$$\rho c \frac{\partial T}{\partial t} = \frac{\partial}{\partial z} \left(k \frac{\partial T}{\partial z} \right) - \rho c \mathbf{u} \cdot \nabla T + 2 \dot{\varepsilon} \sigma$$

Ice Sheet Modeling

"Reference" model: FULL STOKES¹

Approximations based on their accuracy with respect to the ice-sheet aspect ratio δ

- $O(\delta^2)$ **FO**, Blatter-Pattyn first order model² (3D PDE, in horizontal velocities)
- $O(\delta)$ Zeroth order, depth integrated models: **SIA**, Shallow Ice Approximation (slow sliding regimes), **SSA** Shallow Shelf Approximation (2D PDE) (fast sliding regimes)

 $\simeq O(\delta^2)$ High order, depth integrated (2D) models: *L1L2*³, (L1L1)...

¹Gagliardini and Zwinger, 2008. The Cryosphere. ²Dukowicz, Price and Lipscomb, 2010. J. Glaciol. ³Schoof and Hindmarsh, 2010. Q. J. Mech. Appl. Math.

Implementation Overview

ice-sheets FE dynamics component

LifeV: Parallel, object oriented, C++ Finite Element Library:

- linear and quadratic finite elements
- assembling of finite element matrices
- handling of boundary conditions

Implementation Overview

Trilinos:

- Parallel Data Structures (EPETRA)
- Parallel Linear Solvers (GMRES, CG...)
- Preconditioners (Multilevel, Multigrid, Incomplete LU)
- Nonlinear Solvers (NOX package: Newton, JFNK methods)

ice-sheets FE dynamics component

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Work in Progress

MPAS (land ice component):

- Voronoi unstructured grids
- Evolution equation solvers (temperature and thickness equation)



Interface MPAS-LIFEV

2D CVT mesh (Stereographic projection)

thickness/elevation/layers

temperature/ice flow factor

bedrock sliding coefficient

Solver options: model (FO, L1L2, SSA, SIA) nonlinear solver (Newton, Picard, JFNK) Boundary condition (free-slip, no-slip, robin, coulomb) LIFEV

velocity heat dissipation viscosity

Interface - Grids



Interface - Grids



Interface - Grids



Based on 2D grid and thickness and layers build vertically structured **3D grid**.

Build prisms with triangular base and split them in tetrahedra.



ISMIP-HOM: Test C

Surface velocity



Perego, Gunzburger, Burkardt, Journal of Glaciology, 2012.

Towards realistic simulations: Greenland, 5km resolution, sliding case

Comparisons between different models



Towards realistic simulations: Greenland, 2km resolution, no-slip



Towards realistic simulations: Antarctica, 5km resolution, no-slip



Convergence of the nonlinear method (test C ISMIP-HOM)

Solve F(u) = 0. Newton method: $J^k (\delta u)^{k+1} = -F(u^k)$. (NOX library)

The exact Jacobian matrix must be assembled at each nonlinear iteration.

In order to increase robustness of Newton method, the Newton step is halved to achieve monotonic convergence.



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Greenland, FO: convergence of the nonlinear method



First Order Equations

FO is a nonlinear system of elliptic equations in the horizontal velocities:

$$\begin{cases} -\nabla \cdot (2\mu \,\dot{\boldsymbol{\varepsilon}}_1) = -\rho g \frac{\partial s}{\partial x} & \mu = \frac{1}{2} A^{-\frac{1}{n}} \dot{\varepsilon}_e^{\left(\frac{1}{n} - 1\right)} \\ -\nabla \cdot (2\mu \,\dot{\boldsymbol{\varepsilon}}_2) = -\rho g \frac{\partial s}{\partial y}, & \dot{\varepsilon}_e = \sqrt{\dot{\varepsilon}_{xx}^2 + \dot{\varepsilon}_{yy}^2 + \dot{\varepsilon}_{xx}} \dot{\varepsilon}_{yy} + \dot{\varepsilon}_{xy}^2 + \dot{\varepsilon}_{xz}^2 + \dot{\varepsilon}_{yz}^2 \end{cases}$$

where s is the ice surface and,

$$\dot{\varepsilon}_{i,j} = \frac{1}{2} \left(\partial_j u_i + \partial_i u_j \right), \quad i,j \in \{x,y,z\}, \quad \dot{\varepsilon}_1 = \begin{bmatrix} 2\dot{\varepsilon}_{xx} + \dot{\varepsilon}_{yy} \\ \dot{\varepsilon}_{xy} \\ \dot{\varepsilon}_{xz} \end{bmatrix}, \quad \dot{\varepsilon}_2 = \begin{bmatrix} \dot{\varepsilon}_{yx} \\ \dot{\varepsilon}_{xx} + 2\dot{\varepsilon}_{yy} \\ \dot{\varepsilon}_{yz} \end{bmatrix}$$

Remark The nonlinear viscosity μ is singular when $\dot{\varepsilon}_e = 0$, however, $\mu \,\dot{\varepsilon}_1$ is not singular and the PDE is well defined.

Viscosity¹ regularization:

$$\dot{\varepsilon}_e^{-\left(1-\frac{1}{n}\right)} \approx \left(\sqrt{\dot{\varepsilon}_e^2 + \delta^2}\right)^{-\left(1-\frac{1}{n}\right)}$$

¹Goldsby D. L. Kohlsted, JGR, 2001.

Greenland, FO: convergence of the nonlinear method

Simplified Problem (similar kind of nonlinearity):

$$x |x|^{-\left(1-\frac{1}{n}\right)} = C.$$
 solution: $\alpha = C|C|^{n-1}$

Picard method: $x^{k+1} = C|x^k|^{\left(1-\frac{1}{n}\right)}$.

Newton method: $x^{k+1} = (1-n)x^k + nC|x^k|^{(1-\frac{1}{n})}$.

Picard convergence:
$$\lim_{k \to \infty} \frac{x^{k+1} - \alpha}{x^k - \alpha} = 1 - \frac{1}{n} \quad \left(= \frac{2}{3} \text{ when } n = 3 \right)$$

Newton convergence:
$$\lim_{k \to \infty} \frac{x^{k+1} - \alpha}{(x^k - \alpha)^2} = \frac{1}{2\alpha} \left(\frac{1}{n} - 1 \right)$$

regularization:
$$x |x|^{-(1-\frac{1}{n})} \approx x \left(\sqrt{x^2 + \delta^2}\right)^{-(1-\frac{1}{n})}$$

Greenland, FO: convergence of the nonlinear method (comparison using different regularizations)



Greenland, FO: convergence of the nonlinear method (comparison using different regularizations)



 δ variable : At each newton iteration δ is decreased from 1e-4 to 1e-9

Greenland, FO: convergence of the nonlinear method (LOCA continuation method)



The parameter δ is decreased by LOCA from 1e-4 to 1e-9

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