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Energetically based mixing schemes

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Institut für Meereskunde, Universität Hamburg



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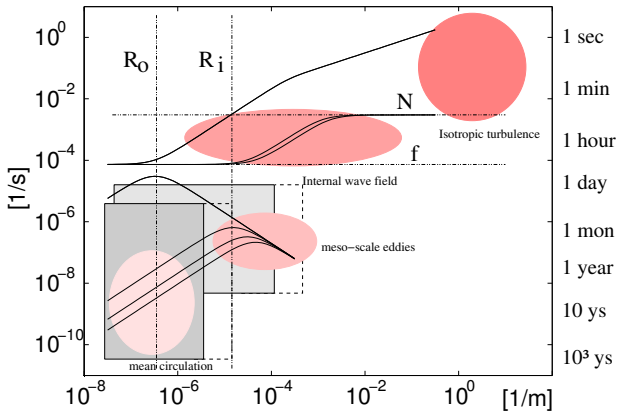
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Energetically based mixing schemes or how to construct a consistent ocean model

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- solid lines: dispersion relations of linear wave solutions
- red ellipses: dynamical regimes, grey boxes: ocean models
- R_o , R_i : Rossby radii of deformation,
 N buoyancy frequency, f Coriolis frequency

- assume field of weakly interacting internal (WKB) waves

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- note that waves with negative m move upward (with $\dot{z} = c$), those with positive m move downward (with $\dot{z} = -c$)
- integrate over all vertical wavenumbers $m > 0$ and $m < 0$

$$E^+ = \int_{-\infty}^0 dm \int d^2 k_h \mathcal{E}^+(\mathbf{k}_h, m, z, t)$$
$$E^- = \int_0^{\infty} dm \int d^2 k_h \mathcal{E}^-(\mathbf{k}_h, m, z, t)$$

- integration of spectral energy $\mathcal{E}(\mathbf{k}, \mathbf{x}, t)$

$$\frac{\partial \mathcal{E}}{\partial t} + \nabla_h \cdot (\dot{\mathbf{x}}_h \mathcal{E}) + \frac{\partial}{\partial \mathbf{k}_h} \cdot (\dot{\mathbf{k}}_h \mathcal{E}) + \frac{\partial}{\partial z} (\dot{z} \mathcal{E}) + \frac{\partial}{\partial m} (\dot{m} \mathcal{E}) = \omega S$$

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$$\frac{\partial E^+}{\partial t} + \frac{\partial J^+}{\partial z} = \int_{-\infty}^0 dm \int d^2 k_h \omega (S_{diss} + S_{ww})$$

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- work with total energy $E = E^+ + E^-$ and $\Delta E = E^+ - E^-$
- S_{ww} conserves total energy E
- assume that S_{ww} damps ΔE and that S_{diss} is symmetric in m

- total energy E and of difference of up/downward waves ΔE

$$\frac{\partial E}{\partial t} + \frac{\partial}{\partial z} c_0 \Delta E = -\epsilon_{iw}, \quad \frac{\partial \Delta E}{\partial t} + \frac{\partial}{\partial z} c_0 E = -\Delta E / \tau_v$$

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$$\epsilon_{iw} = \mu f E^2 / c_*^2$$

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- similar form of ϵ_{iw} is used for analysis of fine-structure measurements (Gregg, 1989)

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- assumptions:
 - S_{diss} is nearly symmetric in m
 - effect of S_{ww} is to eliminate asymmetries, i.e. to damp ΔE
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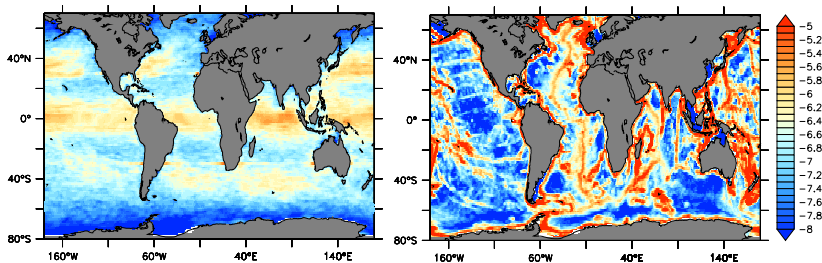
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- forcing by energy flux $c_0 \tau_v \partial(c_0 E) / \partial z$ at surface/bottom

Forcing IDEMIX

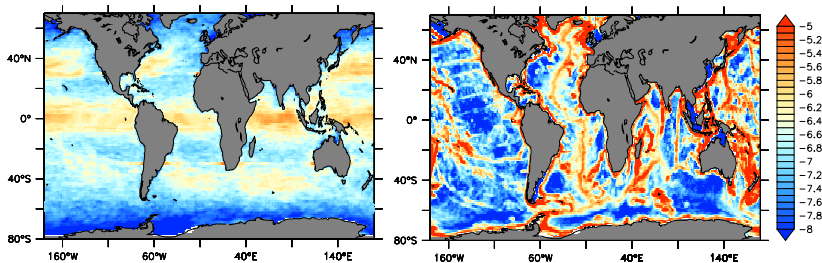
- external forcing functions in $\log_{10}(F/m^3s^{-3})$



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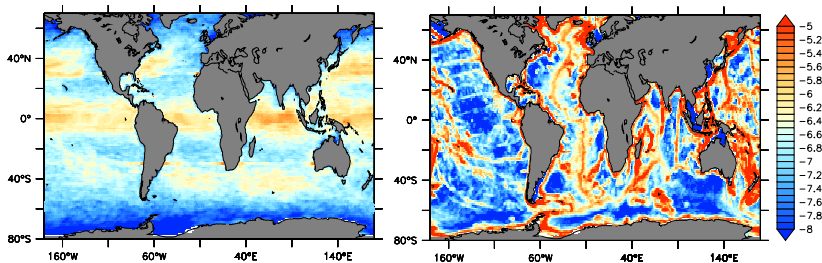
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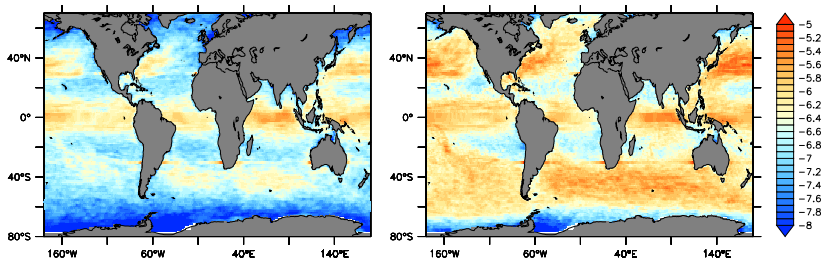
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- tidal flow over topography
1.7 TW from Jayne (2009)
- and other sources

Forcing IDEMIX

- external forcing functions in $\log_{10}(F/m^3s^{-3})$



0.14 TW from NCEP (6 hourly, 1.875°)

but 0.38 TW from CVFS (1 hourly, 0.35°) (Rimac et al, 2012)

- small-scale turbulent kinetic energy (TKE) equation

$$\frac{D\bar{E}}{Dt} = -\frac{\partial}{\partial z} (\text{fluxes}) - \overline{\mathbf{u}'w'} \frac{\partial \bar{\mathbf{u}}}{\partial z} + \overline{w'b'} - \nu \overline{(\nabla \mathbf{u})^2}$$

- with $\bar{E} = \overline{(u'^2 + v'^2 + w'^2)}/2$
- vertical shear production $\overline{\mathbf{u}'w'} \partial \bar{\mathbf{u}} / \partial z$
- exchange with potential energy $\overline{w'b'}$
- dissipation, conversion to internal energy

TKE model

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- assume that

$$\overline{b'w'} = -c_b K \frac{\partial \bar{b}}{\partial z}, \quad \overline{\mathbf{u}'w'} = -c_u K \frac{\partial \bar{\mathbf{u}}}{\partial z} \quad \text{and} \quad \nu \overline{(\nabla \mathbf{u})^2} \approx c_\epsilon \bar{E}^{3/2} L^{-1}$$

with dimensionless parameters $c_u, c_b, c_\epsilon = O(1)$
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$$\frac{D\bar{E}}{Dt} = \frac{\partial}{\partial z} c_E K \frac{\partial \bar{E}}{\partial z} + c_u K \left(\frac{\partial \bar{\mathbf{u}}}{\partial z} \right)^2 - c_b K N^2 - c_\epsilon \bar{E}^{3/2} L^{-1}$$

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- but shear production is related to breaking internal waves
→ add ϵ_{iw} from IDEMIX and set $E_{min} = 0$

Global model

- global model, MITgcm, with $1^\circ \times 1^\circ$ horizontal resolution, 115 vertical layers and thus small implicit numerical diffusion

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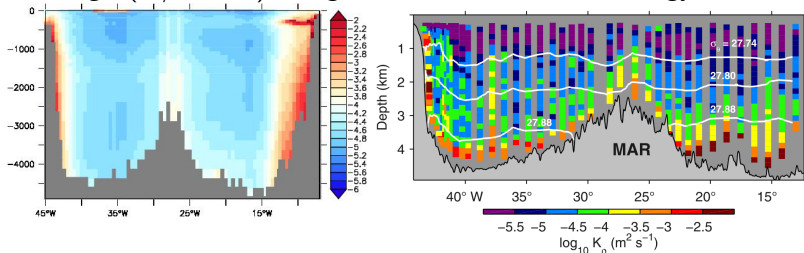
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- small-scale turbulence closure with TKE equation forced by
 - direct wind input of TKE
 - internal wave breaking

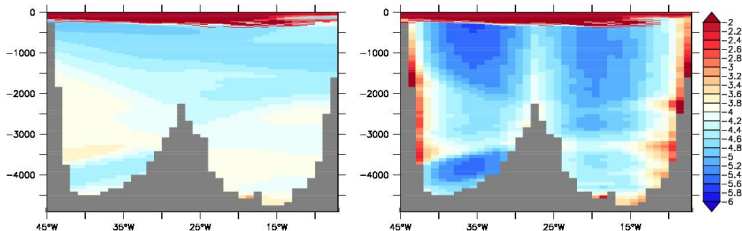
$$\frac{DE_{tke}}{Dt} = \nabla \cdot (\text{fluxes}) + c_u K \left(\frac{\partial \bar{\mathbf{u}}}{\partial z} \right)^2 + \epsilon_{iw} - c_b KN^2 - \epsilon_{tke}$$

Performance of IDEMIX

- $\log_{10}(K/m^2s^{-1})$ using N^2 from WOCE climatology at $48^\circ N$



- K using TKE eq. (left) and TKE + IDEMIX (right) in model



More energy for mixing

- meso-scale eddy kinetic energy

$$\frac{D\bar{E}}{Dt} = -\nabla \cdot (\text{fluxes}) + \bar{S} + \overline{b'w'} - \epsilon_{eke}$$

- with meso-scale eddy kinetic energy (EKE) $\bar{E} = \overline{(u'^2 + v'^2)}/2$
- and barotropic instability production $\bar{S} = -\overline{\mathbf{u}'\mathbf{u}'} \cdot \nabla_h \bar{\mathbf{u}}$
- and baroclinic instability production $\overline{b'w'}$
- and dissipation of EKE ϵ_{eke}

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- meso-scale eddy potential energy $\bar{P} = \overline{b'^2}/(2N^2)$

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- adding and assuming $\overline{\mathbf{u}'_h b'} = -K_{gm} \nabla_h \bar{b}$

$$\frac{D}{Dt} (\bar{E} + \bar{P}) = -\nabla \cdot (\text{fluxes}) + \bar{S} + K_{gm} (\nabla_h \bar{b})^2 N^{-2} - \epsilon_{eke}$$

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 - and dissipation of EKE ϵ_{eke}
- meso-scale eddy potential energy $\bar{P} = \overline{b'^2}/(2N^2)$

$$\frac{D\bar{P}}{Dt} = -\nabla \cdot \text{flux} - \overline{\mathbf{u}'_h b'} \cdot \nabla_h \bar{b} N^{-2} - \overline{b'w'}$$

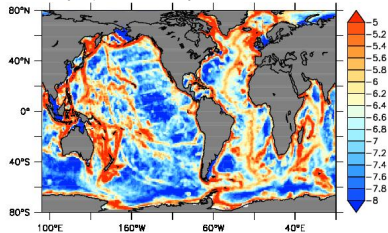
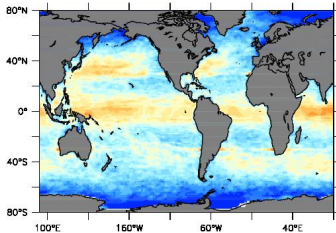
- adding and assuming $\overline{\mathbf{u}'_h b'} = -K_{gm} \nabla_h \bar{b}$

$$\frac{D}{Dt} (\bar{E} + \bar{P}) = -\nabla \cdot (\text{fluxes}) + \bar{S} + K_{gm} (\nabla_h \bar{b})^2 N^{-2} - \epsilon_{eke}$$

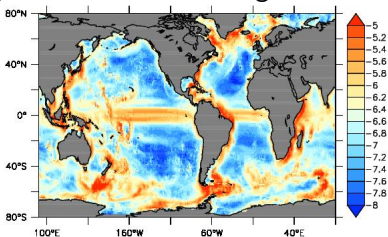
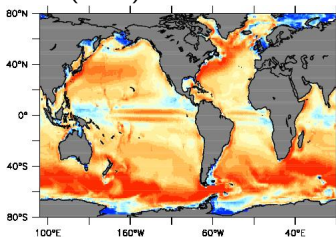
- close with $K_{gm} = c_g \sqrt{\bar{E}} L$, and $\epsilon_{eke} = c_\epsilon \bar{E}^{3/2}/L$ as in (Eden + Greatbatch, 2008)

consistent ocean model

- external forcing functions in $\log_{10}(F/m^3s^{-3})$



0.14 (0.38) TW wind forcing, 1.7 TW tidal forcing



0.67 TW of $K_{gm}(\nabla_h \bar{b})^2 N^{-2}$ and 0.49 TW of \bar{S}

IDEMIX forced with eddies

- internal wave closure (IDEMIX)

$$\frac{DE_{iw}}{Dt} = \nabla \cdot (\text{fluxes}) + \epsilon_{eke} - \epsilon_{iw}$$

- meso-scale eddy closure

$$\frac{DE_{eke}}{Dt} = \nabla \cdot (\text{fluxes}) + \bar{S} + K_{gm}(\nabla_h \bar{b})^2 N^{-2} - \epsilon_{eke}$$

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- assume local dissipation (interior loss of balance?)

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$$\frac{DE_{eke}}{Dt} = \nabla \cdot (\text{fluxes}) + \bar{S} + K_{gm}(\nabla_h \bar{b})^2 N^{-2} - \epsilon_{eke}$$

- assume local dissipation (interior loss of balance?)
- or assume $\int_{-h}^0 \epsilon_{eke} dz$ is injected at bottom only (generation of lee waves by meso-scale eddies?)

IDEMIX forced with eddies

- internal wave closure (IDEMIX)

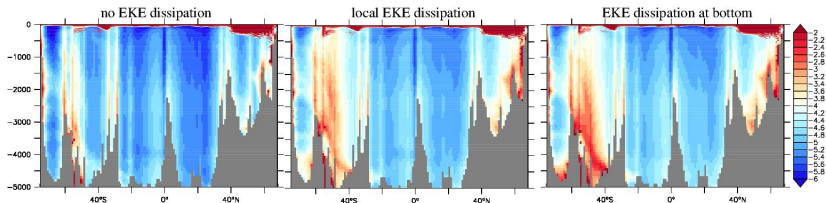
$$\frac{DE_{iw}}{Dt} = \nabla \cdot (\text{fluxes}) + \epsilon_{eke} - \epsilon_{iw}$$

- meso-scale eddy closure

$$\frac{DE_{eke}}{Dt} = \nabla \cdot (\text{fluxes}) + \bar{S} + K_{gm}(\nabla_h \bar{b})^2 N^{-2} - \epsilon_{eke}$$

- assume local dissipation (interior loss of balance?)
- or assume $\int_{-h}^0 \epsilon_{eke} dz$ is injected at bottom only (generation of lee waves by meso-scale eddies?)

$\log_{10}(K/m^2s^{-1})$ at 30°W



Summary

- construct a consistent ocean model without spurious energy sources and sinks

$$\begin{aligned}\frac{DE_m}{Dt} &= \nabla \cdot (\text{fluxes}) + \bar{b}\bar{w} - \bar{S} - c_u K \left(\frac{\partial \bar{\mathbf{u}}}{\partial z} \right)^2 \\ \frac{DP_m}{Dt} &= \nabla \cdot (\text{fluxes}) - \bar{b}\bar{w} - K_{gm} (\nabla_h \bar{b})^2 N^{-2} + c_b K N^2 \\ \frac{DE_{eke}}{Dt} &= \nabla \cdot (\text{fluxes}) + \bar{S} + K_{gm} (\nabla_h \bar{b})^2 N^{-2} - \epsilon_{eke} \\ \frac{DE_{iw}}{Dt} &= \nabla \cdot (\text{fluxes}) + \epsilon_{eke} - \epsilon_{iw} \\ \frac{DE_{tke}}{Dt} &= \nabla \cdot (\text{fluxes}) + c_u K \left(\frac{\partial \bar{\mathbf{u}}}{\partial z} \right)^2 + \epsilon_{iw} - c_b K N^2 - \epsilon_{tke}\end{aligned}$$

Summary

- construct a consistent ocean model without spurious energy sources and sinks
- IDEMIX connects different forcing functions, small-scale, and meso-scale turbulence closures

$$\begin{aligned}\frac{DE_m}{Dt} &= \nabla \cdot (\text{fluxes}) + \bar{b}\bar{w} - \bar{S} - c_u K \left(\frac{\partial \bar{\mathbf{u}}}{\partial z} \right)^2 \\ \frac{DP_m}{Dt} &= \nabla \cdot (\text{fluxes}) - \bar{b}\bar{w} - K_{gm} (\nabla_h \bar{b})^2 N^{-2} + c_b K N^2 \\ \frac{DE_{eke}}{Dt} &= \nabla \cdot (\text{fluxes}) + \bar{S} + K_{gm} (\nabla_h \bar{b})^2 N^{-2} - \epsilon_{eke} \\ \frac{DE_{iw}}{Dt} &= \nabla \cdot (\text{fluxes}) + \epsilon_{eke} - \epsilon_{iw} \\ \frac{DE_{tke}}{Dt} &= \nabla \cdot (\text{fluxes}) + c_u K \left(\frac{\partial \bar{\mathbf{u}}}{\partial z} \right)^2 + \epsilon_{iw} - c_b K N^2 - \epsilon_{tke}\end{aligned}$$

Summary

- construct a consistent ocean model without spurious energy sources and sinks
- IDEMIX connects different forcing functions, small-scale, and meso-scale turbulence closures
- unknown dissipation of meso-scale eddy energy

$$\frac{DE_m}{Dt} = \nabla \cdot (\text{fluxes}) + \bar{b}\bar{w} - \bar{S} - c_u K \left(\frac{\partial \bar{\mathbf{u}}}{\partial z} \right)^2$$

$$\frac{DP_m}{Dt} = \nabla \cdot (\text{fluxes}) - \bar{b}\bar{w} - K_{gm} (\nabla_h \bar{b})^2 N^{-2} + c_b KN^2$$

$$\frac{DE_{eke}}{Dt} = \nabla \cdot (\text{fluxes}) + \bar{S} + K_{gm} (\nabla_h \bar{b})^2 N^{-2} - \epsilon_{eke}$$

$$\frac{DE_{iw}}{Dt} = \nabla \cdot (\text{fluxes}) + \epsilon_{eke} - \epsilon_{iw}$$

$$\frac{DE_{tke}}{Dt} = \nabla \cdot (\text{fluxes}) + c_u K \left(\frac{\partial \bar{\mathbf{u}}}{\partial z} \right)^2 + \epsilon_{iw} - c_b KN^2 - \epsilon_{tke}$$