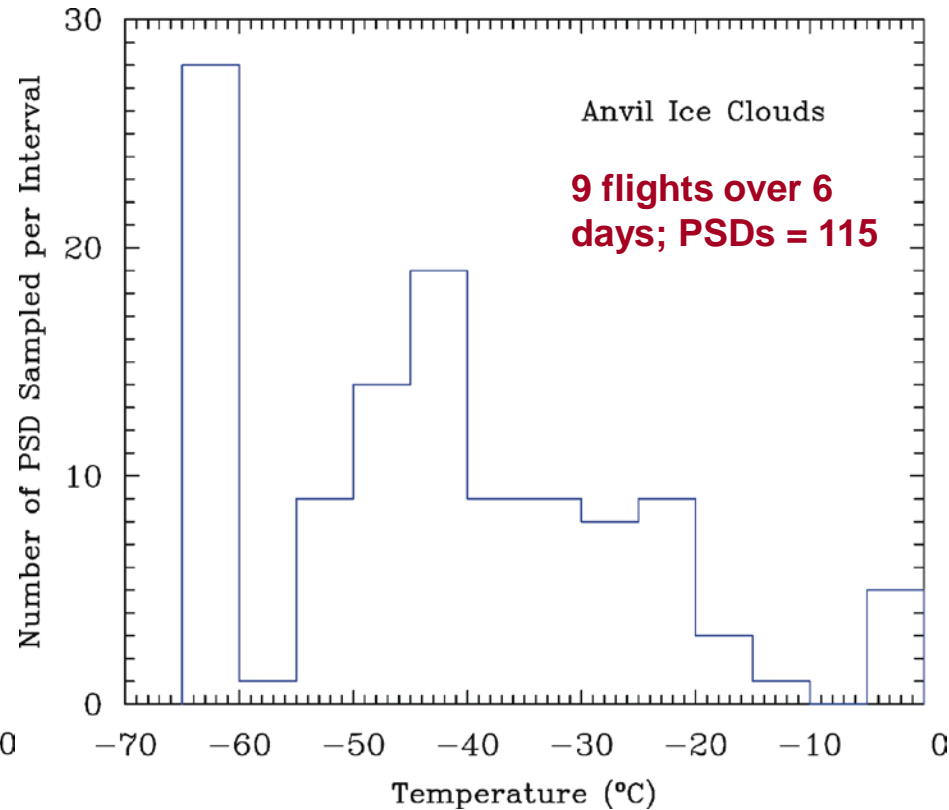
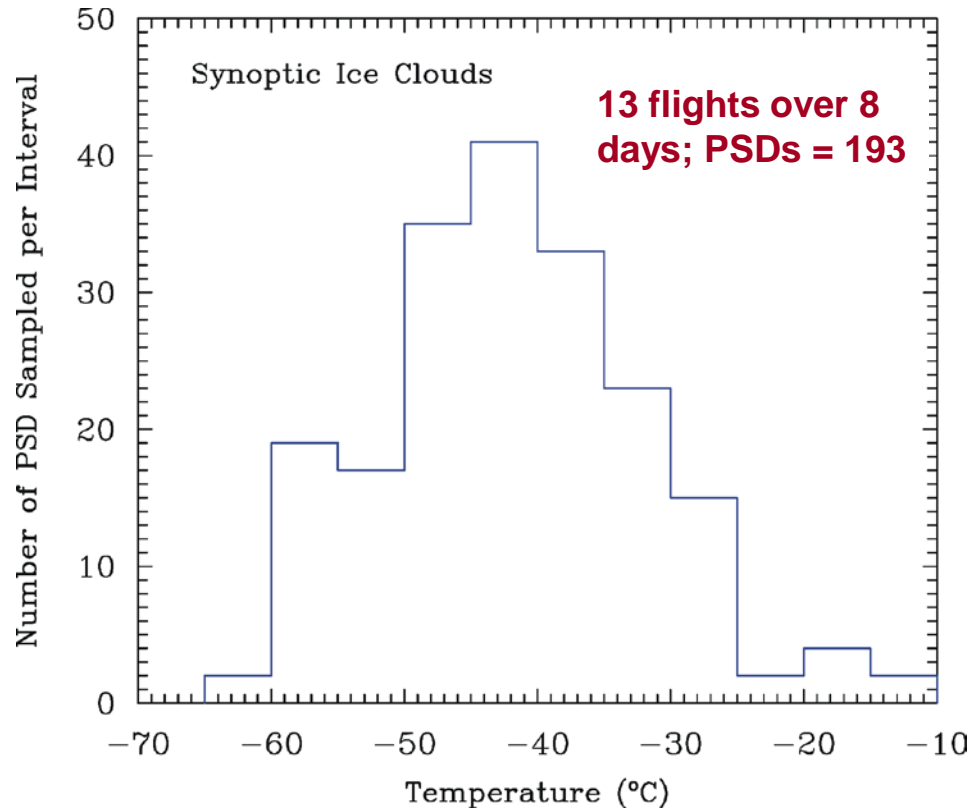


Parameterizing Ice Particle Mass and Area in Ice Clouds: Towards a Self-consistent Treatment of Ice Microphysics and Radiation

David L. Mitchell and Ehsan Erfani
Desert Research Institute, Reno, Nevada



SPARTICUS PSD SAMPLING FOR SELECTED SYNOPTIC AND ANVIL CIRRUS CLOUDS



Generating m-D and A-D Relationships Using the 2D-S Probe

The 2D-S probe measures the size resolved concentrations of ice particle number, projected area and estimated mass from 10 to 1280 μm , using the Baker-Lawson mass-area power law to estimate particle mass. The mean ice particle mass for each size-bin is given as

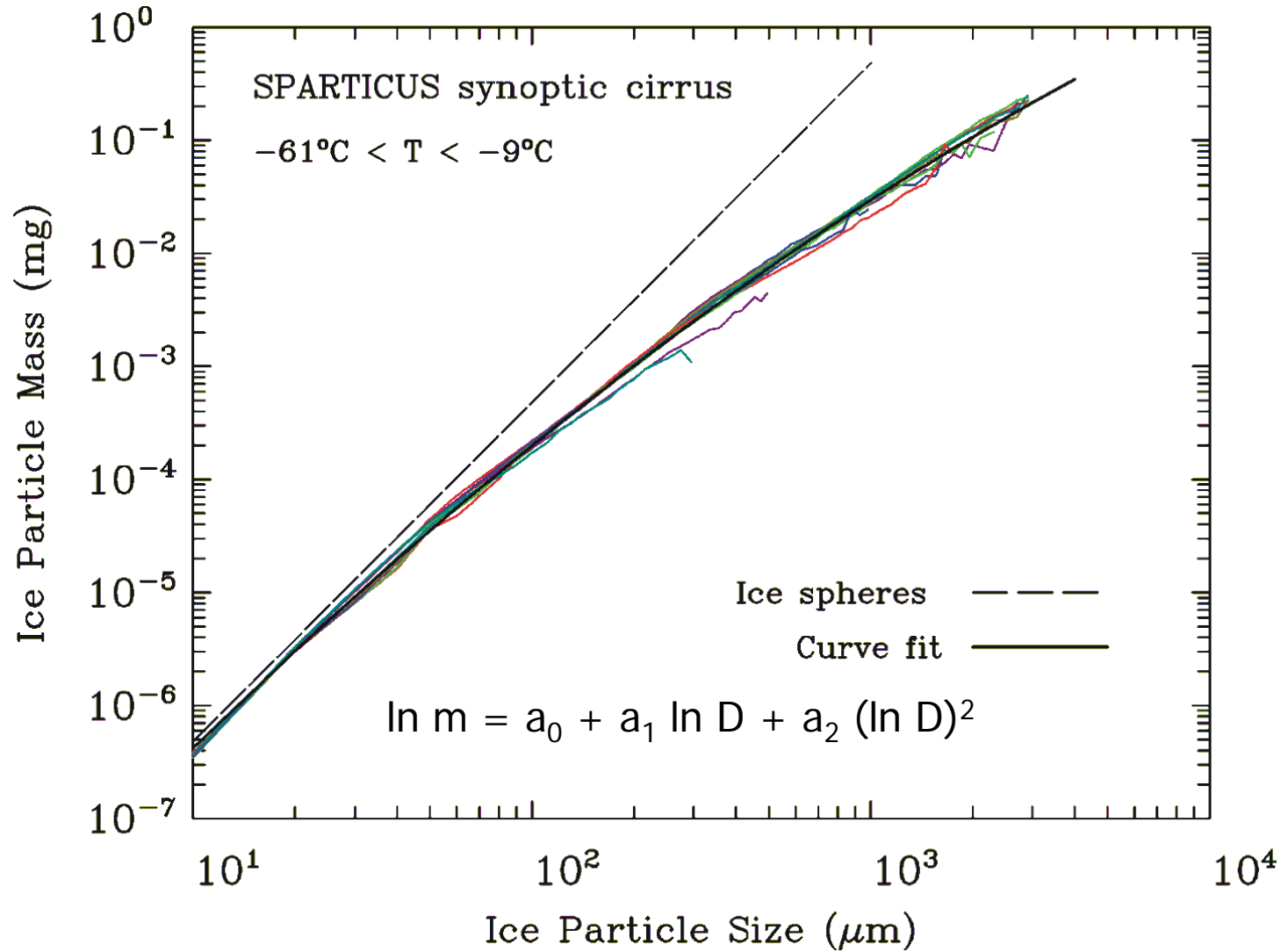
$$m = M/N$$

where M and N are mass and number concentration for a given size-bin. Relating m to the midpoint size of each bin, the dependence of m on D is revealed. The dependence of mean ice particle projected area on D is found in a similar manner.

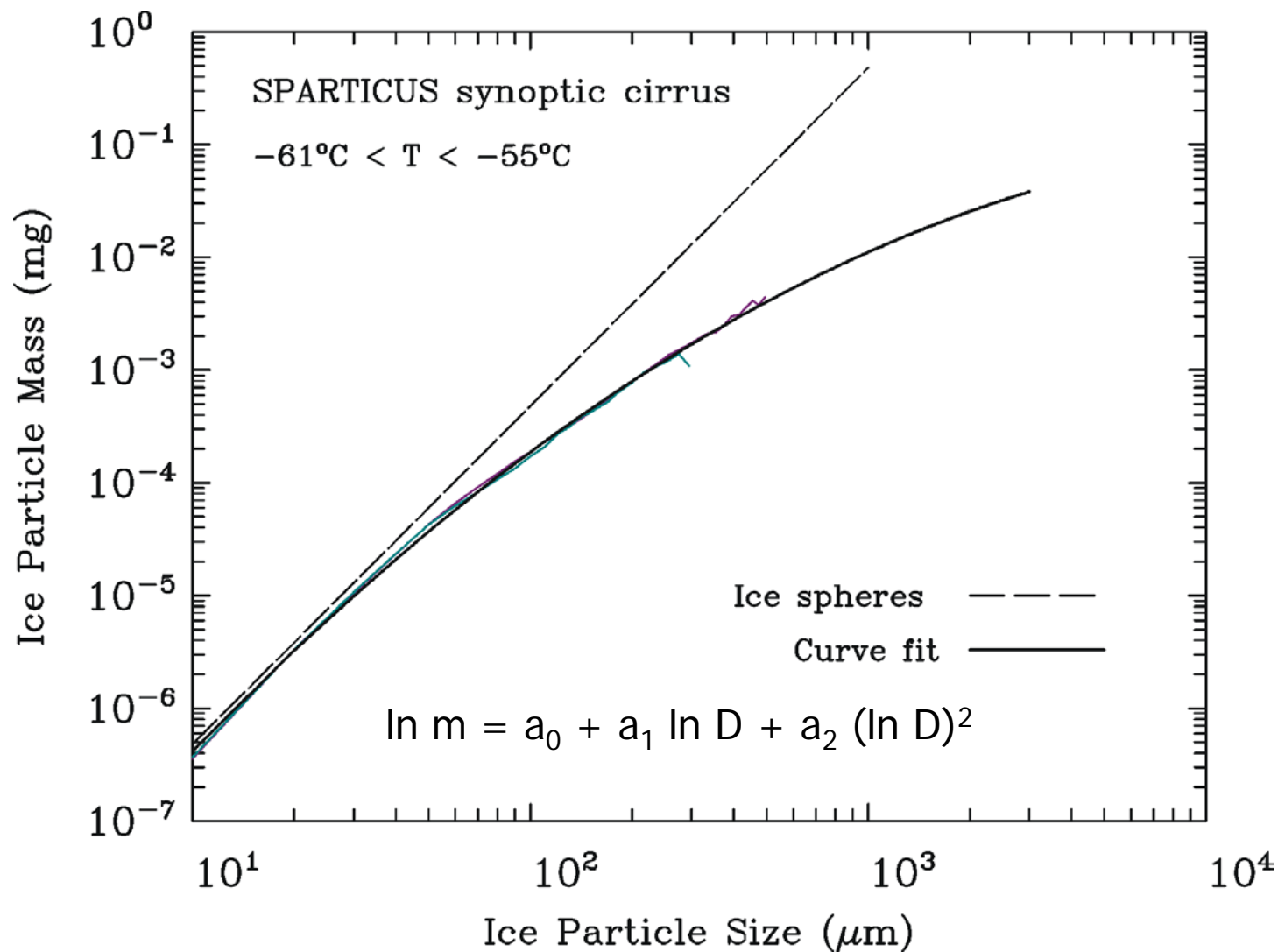
A mean PSD was produced for each 5°C T-interval that reduces scatter while focusing the curve fit. Temperature categories were identified where m-D & A-D curves were similar. A single m-D or A-D curve fit was produced per T-category.

A key objective here was to evaluate the dependence of m and A on D over a much greater range of ice particle size than done before (especially the size range between 10 and 100 μm).

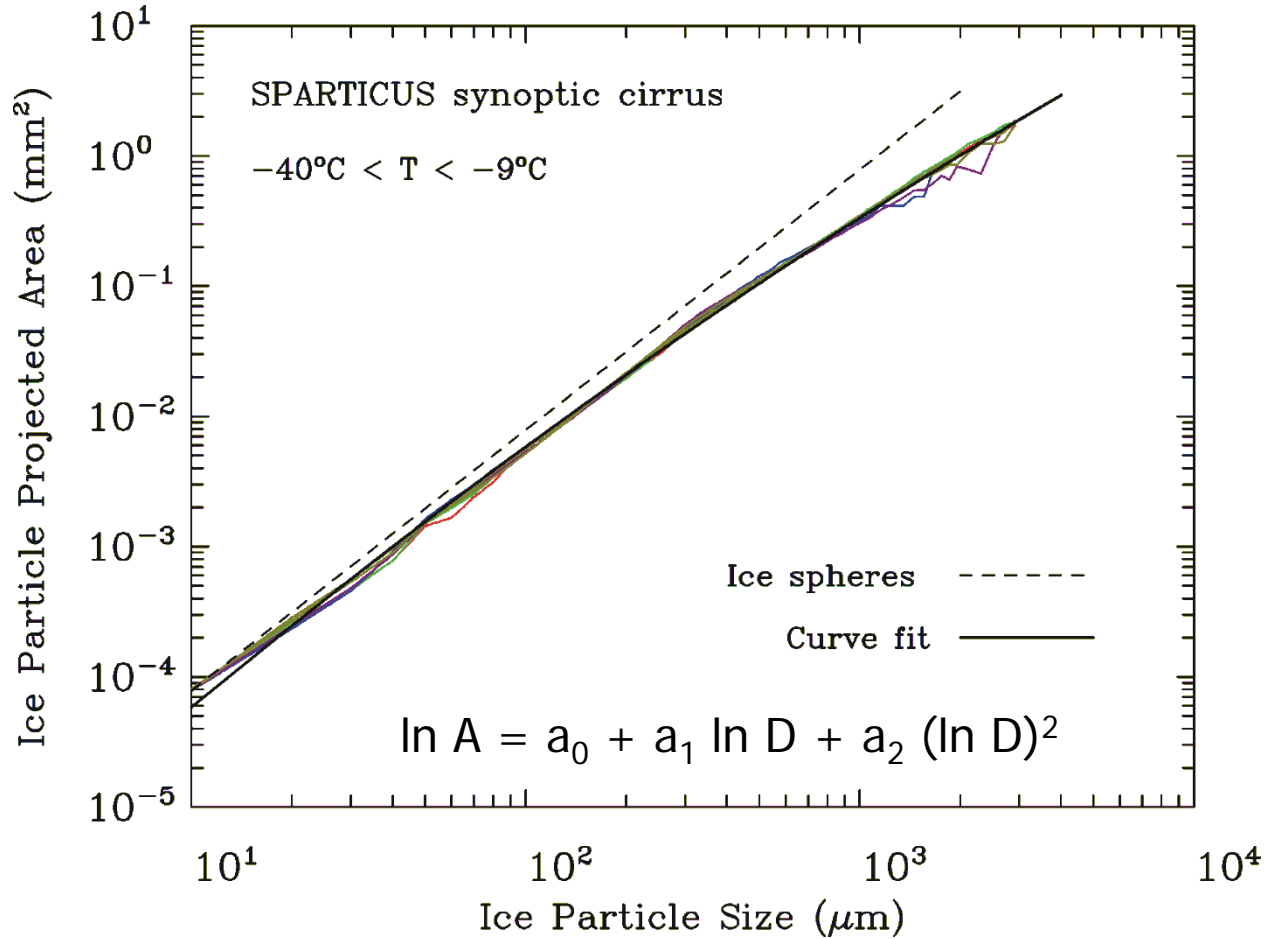
Dependence of $\log(m)$ on $\log(D)$ is not linear



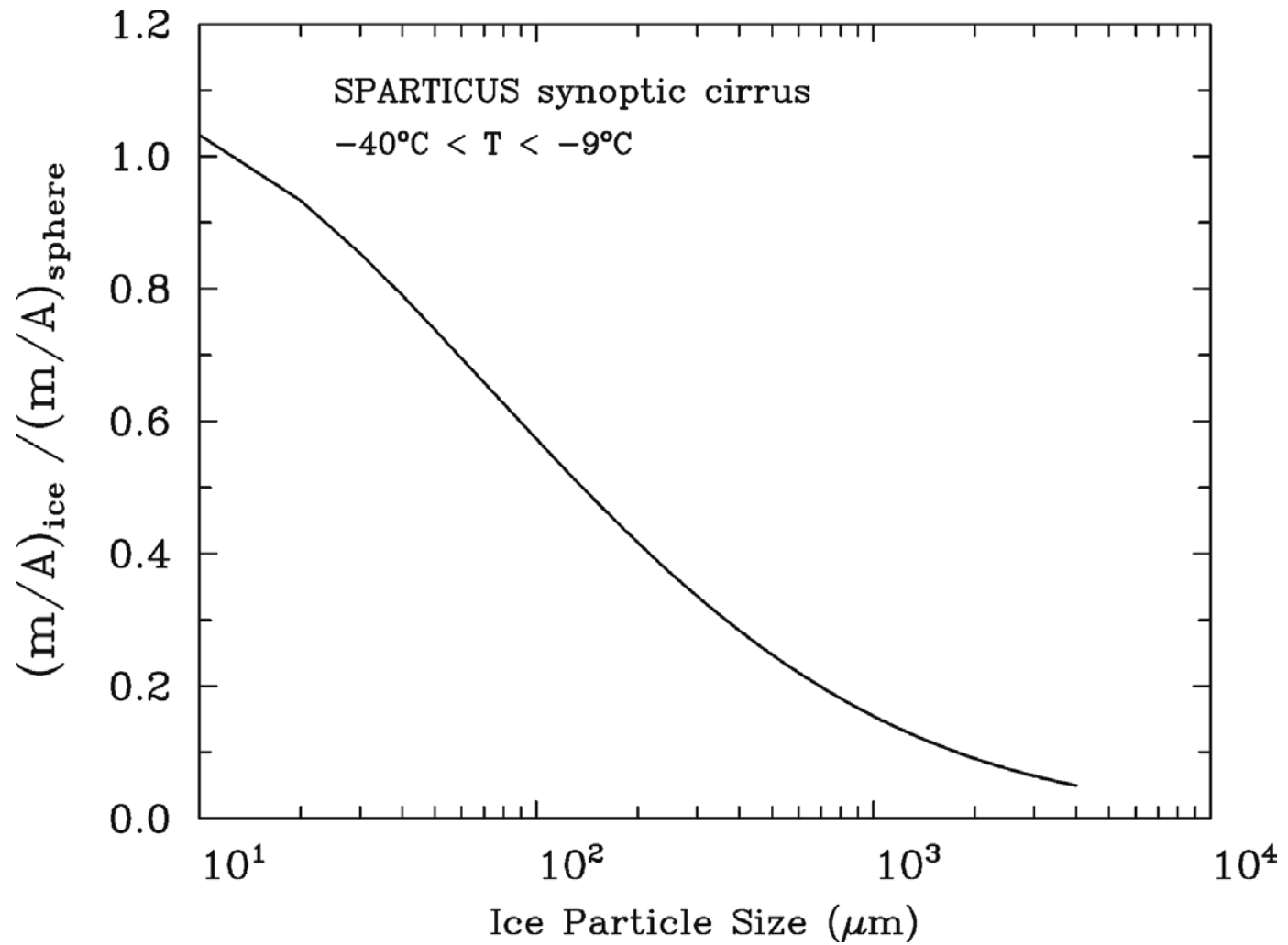
Data was grouped into 3 temperature-categories for synoptic and anvil cirrus, with a m-D and A-D curve fit for each category



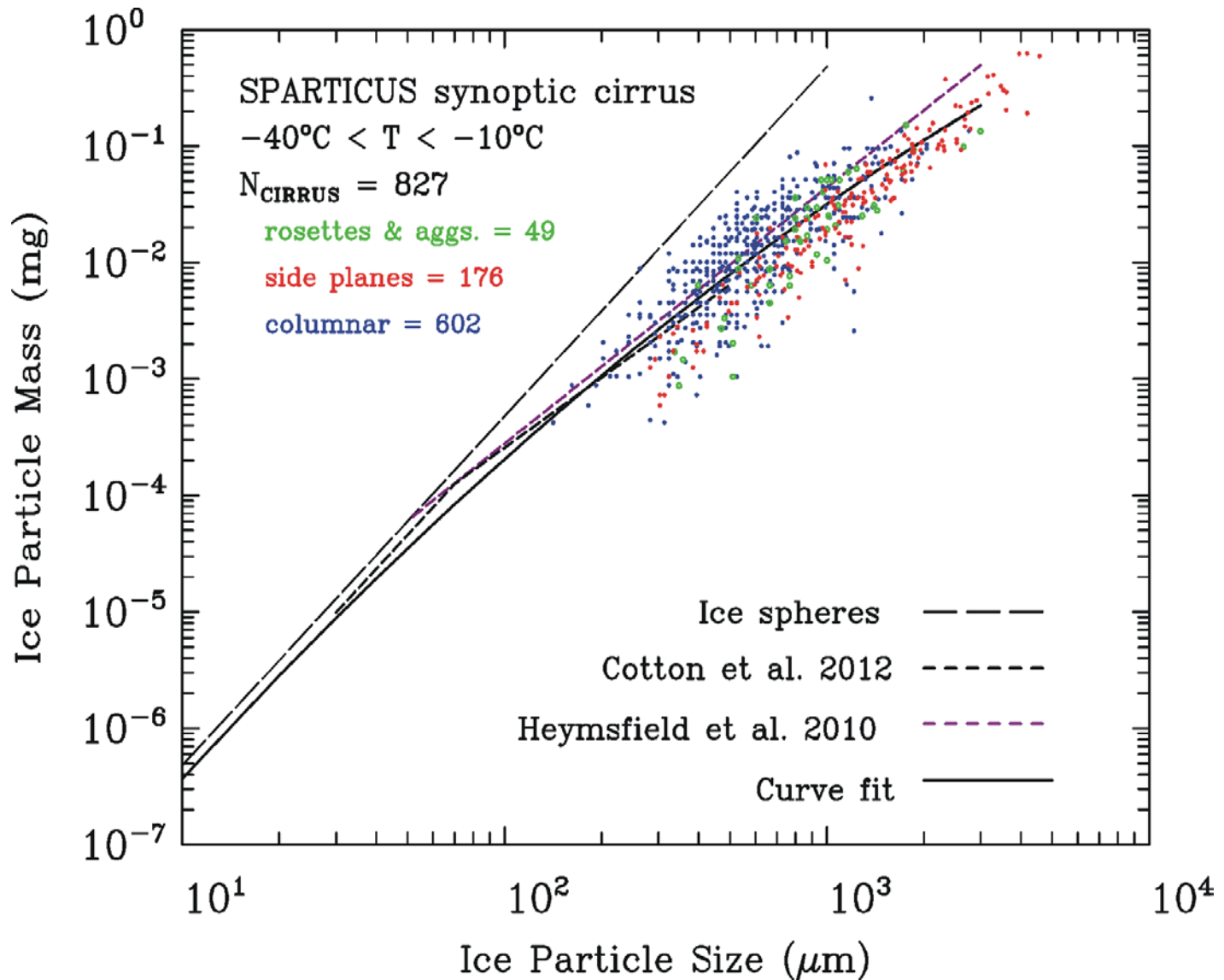
Example of A-D Curve Fit



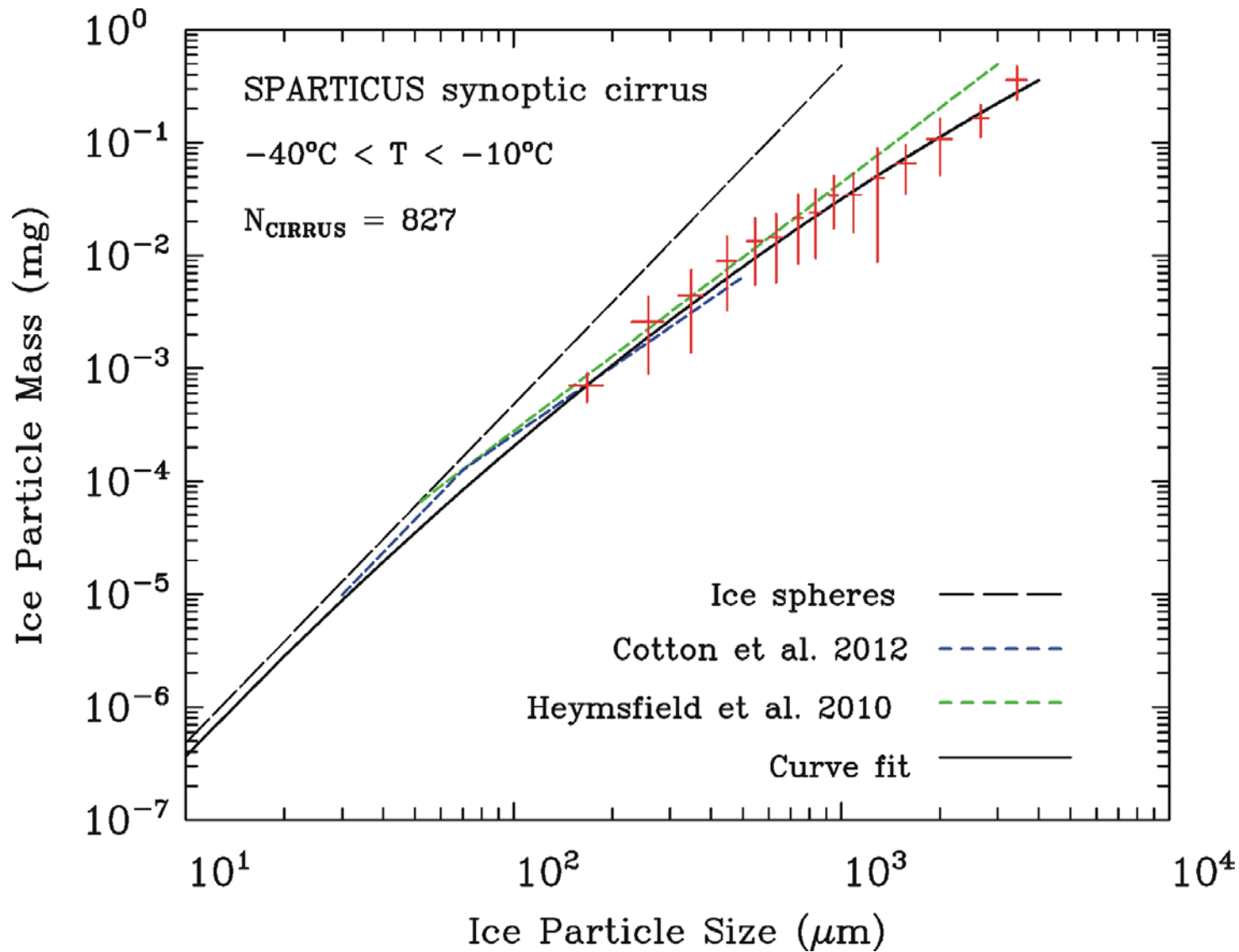
Testing for self-consistency and physical realism

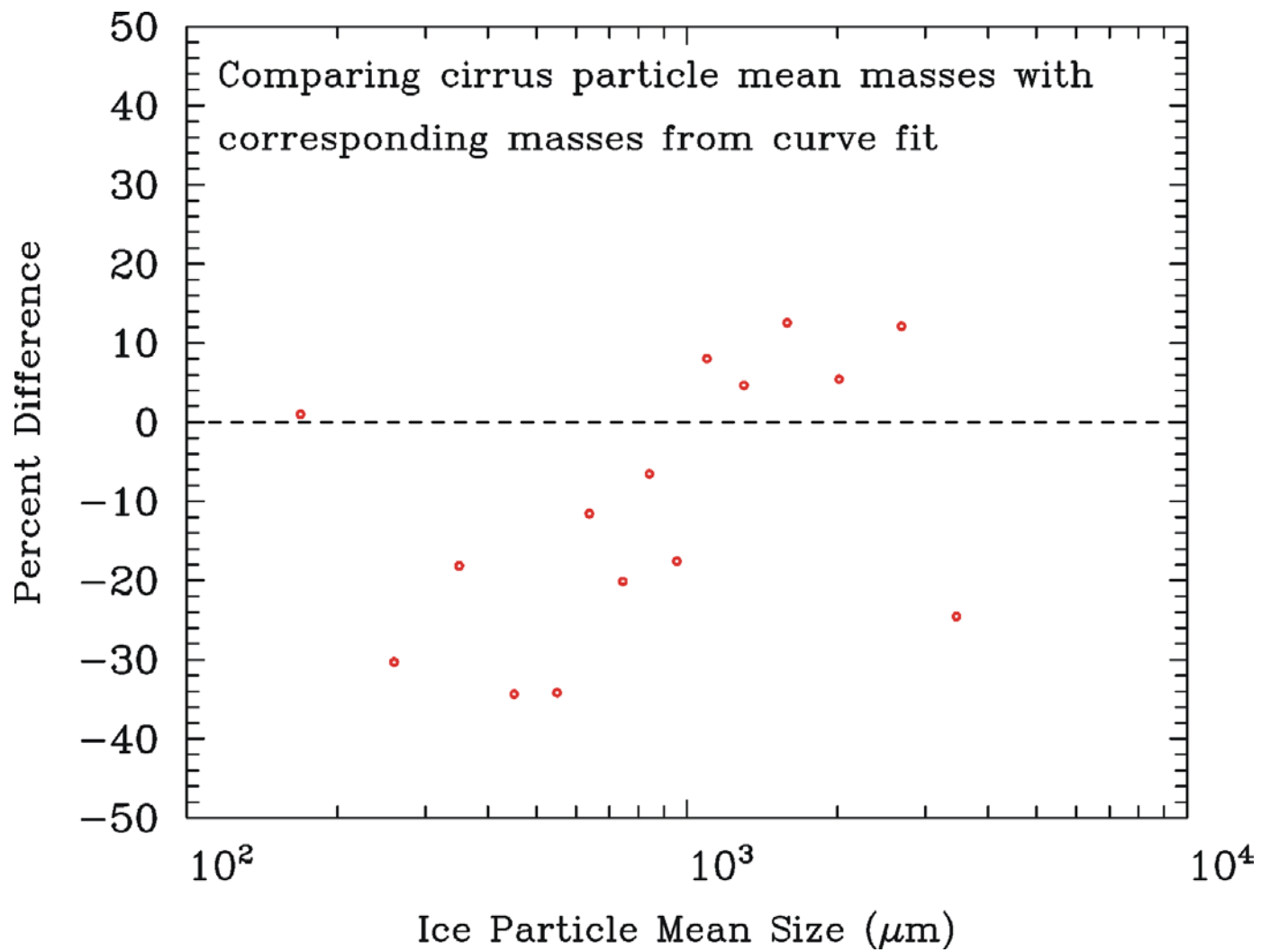


Comparisons with other field data and other studies



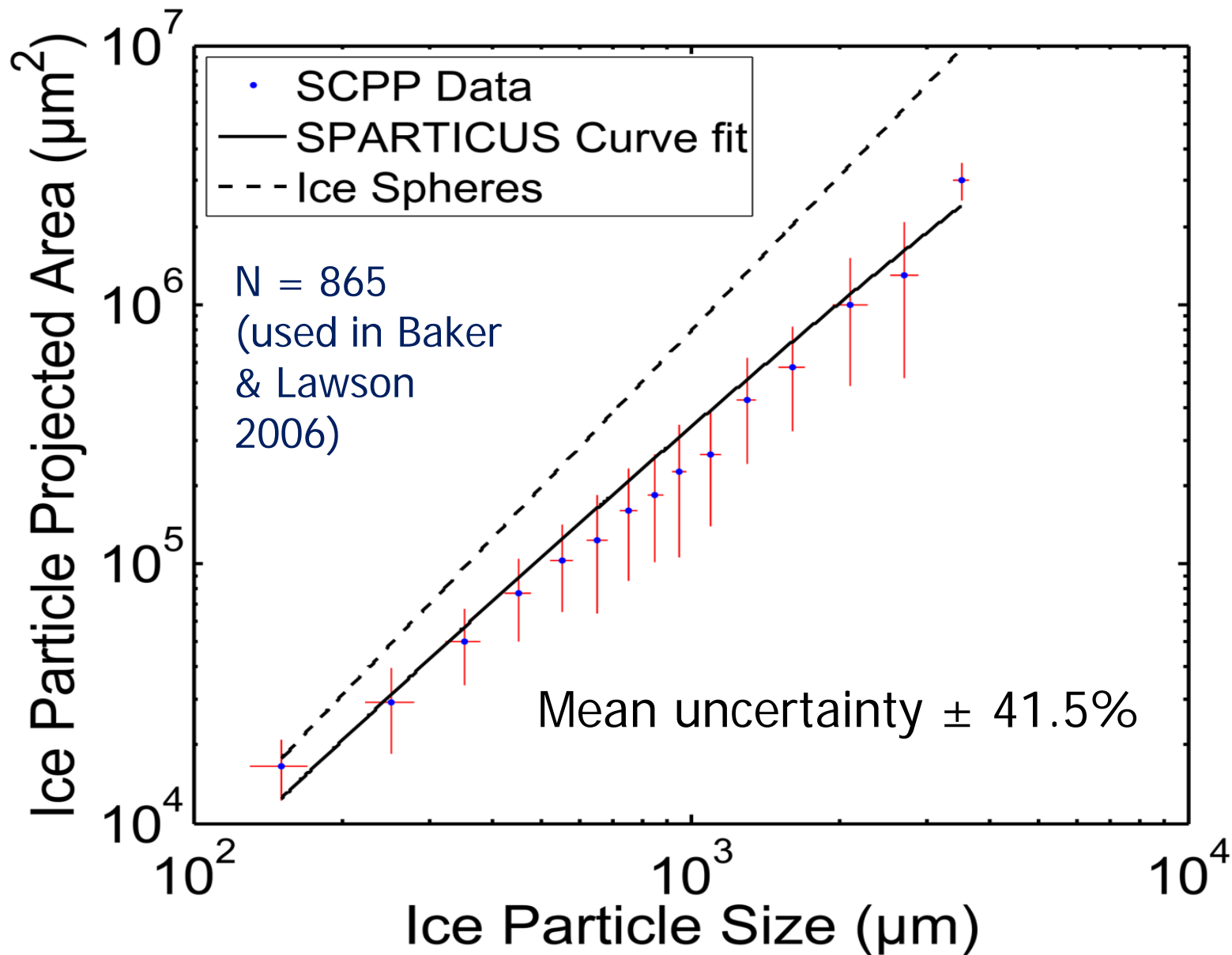
Mean uncertainty = $\pm 54.4\%$





- 10

-



$$m = \alpha D^\beta$$

$$A = \gamma D^\delta$$

For a given D, we can

obtain these power laws

$$\ln m = a_0 + a_1 \ln D + a_2 (\ln D)^2$$

$$d(\ln m)/d(\ln D) = \beta = a_1 + 2a_2 \ln D$$

$$\alpha = \exp[a_0 + a_1 \ln D + a_2 (\ln D)^2] / D^\beta$$

β uncertainty \ll α uncertainty

Synoptic cirrus, exponent of m-D power law

	Ice particle size (μm)				
	50	150	500	1500	4500
Temperature Range	Power β				
-40 < T \leq -10°C	2.632	2.368	2.079	1.814	1.550
-55 < T \leq -40°C	2.564	2.280	1.968	1.684	1.399
-65 < T \leq -55°C	2.477	2.057	1.597	----	----
Mean β	2.558	2.235	1.881	1.749	1.475
Standard deviation of β	0.064	0.131	0.206	----	----

Mean uncertainty for β as $100 \times \sigma / \text{mean value}$: 6.74%

Application to Cloud Modeling

$$N(D) = N_o D^v \exp(-\lambda D)$$

$$\lambda = \left(\frac{\alpha \Gamma(\beta+v+1) N}{\Gamma(v+1) IWC} \right)^{1/\beta}$$

Used in CAM5

To a good approximation, λ is obtained by evaluating α & β at $D = 500 \mu\text{m}$. Then estimate the D for the cloud property or process of interest by evaluating β and δ at $D = 500 \mu\text{m}$:

$$D_N = (v + 0.67)/\lambda$$

Median number conc. dimension

$$D_a = (\delta + v + 0.67)/\lambda$$

Median area dimension

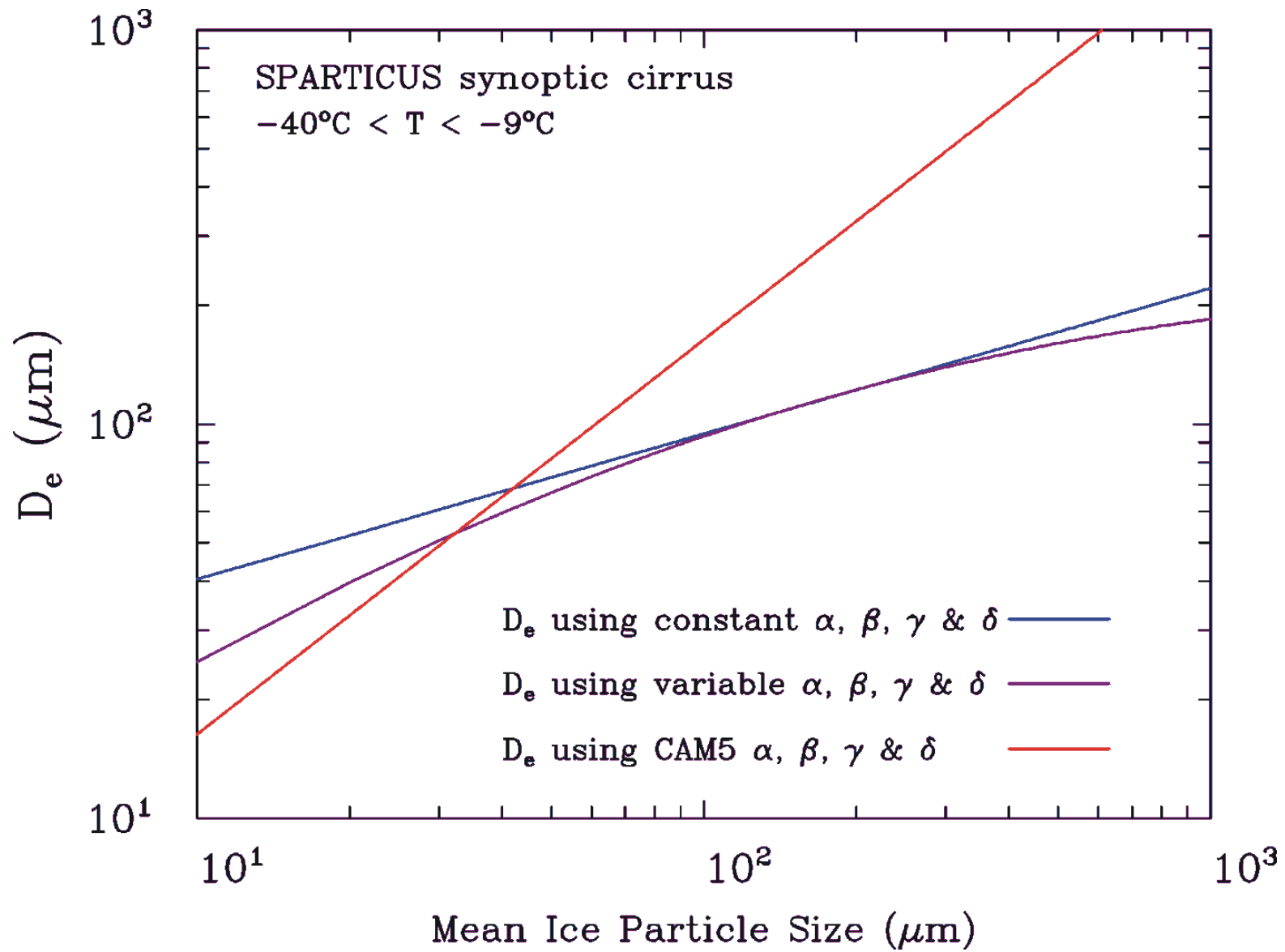
$$D_m = (\beta + v + 0.67)/\lambda$$

Median mass dimension

$$D_Z = (2\beta + v + 0.67)/\lambda$$

Median radar reflectivity dimension

Then calculate α , β , γ and δ for the selected D value.



Summary and Conclusions

1. When considering all sizes, ice particle mass and projected area are better estimated by 2nd order polynomial fits than by power laws.
2. Method's accuracy is evident through comparisons with two recent studies and ground field measurements.
3. To a 1st approximation, the uncertainties (σ) measured can be attributed to the prefactor of the m-D and A-D power laws.
 - This should be useful for estimating σ of cloud retrievals.
4. A means of obtaining accurate m-D and A-D power laws for modeling ice clouds was developed, thus avoiding major changes to model architecture.
5. This treatment of m-D/A-D relationships can be used in CAM5 to achieve consistency between the ice cloud microphysical and optical properties.

