### CMIP in the Stratosphere

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## **Basic Vocabulary**

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**CMIP** = Coupled [ocean-atmosphere-climate] Model Inter-comparison Project

**Stratosphere** = second-lowest layer of the atmosphere; layers separated by local maxima / minima of temperature as a function of altitude



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## Key 2013 Papers Assessing Stratospheric Performance of the CMIP5 Models

**Eyring V., et al.** (2013) "Long-term **Ozone** Changes and Associated Climate Impacts in CMIP5 Simulations," *J. Geophys. Res. 118*, 5029–5060, doi:10.1002/jgrd.50316

**Charlton-Perez A. J. et al.** (2013) "On the Lack of Stratospheric Dynamical Variability in Low-top Versions of the CMIP5 Models," *J. Geophys. Res., 118,* 10.1002/jgrd.50125

General assessment of T and dynamics for 10 hPa observed difference, correlation and r.m.s. error for q, T, u in 21 models\*

Kawatani Y. & Hamilton K. (2013) "Weakened Stratospheric Quasi-Biennial Oscillation Driven by Increased Tropical Mean Upwelling," *Nature 497*, 478-481

\*My 26 models\* = their 21 + 5 extras (CESM-WACCM, CanESM2/AM4, HadGEM2-A, MPI-ESM-MR).

### Vertical Levels in 26 CMIP5 Models [17 include 3-hr surface output\*]

- 17 / 26 models include all of the stratosphere
- 10 / 26 models include all of the stratosphere + all of the mesosphere



5 / 23 all-stratosphere

0 / 23 all-stratosphere+mesosphere (Cordero and Forster 2010) "The four models in CMIP5 that realistically simulate the QBO" (Kawatani and Hamilton 2013; they note that CESM1-WACCM also simulates a realistic QBO but uses nudging to do so).

## Surface-Pressure Amplitudes of Atmospheric Tides in Observations\* and Two CESM Models\*\*



Like the diurnal cycle of surface temperature, the <u>diurnal harmonic</u> of surface pressure is forced by tropospheric heating:  $H_2O(v)$  absorbing solar energy, latent heat release . . .

UVECO

\* A. Dai and J. Wang, 1999: "Diurnal and Semidiurnal Tides in Global Surface Pressure Observations, *Journal of the Atmospheric Sciences*, 56, 3874-3890
\*\* C. Covey, A. Dai, R. S. Lindzen, D. R. Marsh 2014: "Atmospheric Tides in the Latest Generation of Climate Models," in revision for *Journal of the Atmospheric Sciences*

The <u>semidiurnal harmonic</u> is forced by O<sub>3</sub> heating in the middle atmosphere . . .



... so the model with a complete middle atmosphere (WACCM4) gets a bigger amplitude. But apparently, compensating errors are occurring. Our papers show that *all* CMIP models overestimate semidiurnal amplitude. Compensating errors may be common!

#### Q: What sort of compensating errors could the models have in their tide simulations? $G(x) = A e^{\alpha x} + e^{-\alpha x} \Rightarrow$

A: One possibility involves reflection of waves from the top of the model:

From Lindzen et al. (1968): "when the atmosphere has an artificial top then there are an infinite number of h's for which [the **vertical structure equation**] has nontrivial solutions ... These constitute the free oscillations of the atmosphere; they include **Rossby-Haurwitz waves**. In an infinite atmosphere these are all associated with a single h ... [but] bounded models will have other Rossby-Haurwitz waves associated with the spurious h's."

(Horizontal lines show positions of the top-most "full levels" in 17 CMIP5 models.)



 $\alpha$  ( $\alpha$ )

REFERENCE: R. S. Lindzen, E. S. Batten and J.-W. Kim, 1968: "Oscillations in Atmospheres with Tops," *Monthly Weather Review*, **96**, 133-140

## The Quasi-Biennial Oscillation in Microwave Sounding Unit data

- QBO traditionally defined by the mean zonal wind →→→
- ... but also has a temperature component ≤ 3 – 5 K.



- Therefore, can complicate T-trend analysis in tropical upper troposphere.
- QBO signal appears in MSU temperatures after filtering / EOFs (Christy & Drouilhet 1994: *J. Climate 7*, 106; Fernandez et al. 2004: *J. Climate 17*, 3934).
- Now emerges directly from a VERY simplistic Fourier analysis of the complete 1979-2013 time series.

\*Graphic from <u>http://jra.kishou.go.jp/JRA-25/JRA25quality\_en.html</u> (2008)

### MSU Temperatures in the Tropical Lower Stratosphere (weighting function peak ≈ 70 hPa; 20°S – 20°N area average) Thanks to Ben Santer for data.

FOR OBSERVATIONS: Superimposed on a steady decline are times of elevated temperatures around 1983 (El Chichon), 1992-3 (Pinatubo), 1997-2000 (El Nino?) and 2011 (La Nina??). These occur roughly once a decade. Also apparent are temperature spikes occuring at roughly two-year intervals -- by definition "quasi-biennial."



Fourier analysis:

$$T_m = \sum_{s=-N/2+1}^{N/2} a_s e^{2\pi i s m/N}, m = 1, 2, 3, ..., N.$$





### **Observations of Structure and Dynamics\***



"Cold tropical tropopause and midlatitude *T*-max in the lower stratosphere in the winter hemisphere" "Rather uniform *T*-increase from winter pole to summer pole in the 30 – 60 km layer" [ $p \approx 1$  hPa] "Reversed *T*-gradient above 60 km [p < 0.2 hPa] with *T* increasing from summer to winter pole" "Intrusion of summer hemisphere easterlies into the winter hemisphere"

"Displacement of axis of the winter hemisphere westerly jet to about 60° (the polar night jet)"

"Semiannual oscillation of *u* has its maximum amplitude near the equatorial stratopause." "Very strong oscillation of *u* and *T* in the tropical lower stratosphere which has a somewhat irregular period, averaging about 26 months. Alternating pattern of downward propagating easterlies and westerlies."



\*Graphics from http://ds.data.jma.go.jp/gmd/jra/atlas/eng/indexe\_isobar13.htm and jra.kishou.go.jp/JRA-25/JRA25quality\_en.html (2008). Quotes: Holton, *Dynamic Meteorology of the Stratosphere and Mesosphere* (1975). The fundamental observations are robust from the rocketsonde era to the satellite / reanalysis era.

# Q: Do results from other CMIP models also imply compensating errors?

# A: Definitely! <u>All of them</u> over-predict semidiurnal tide amplitude <u>independent of their vertical coverage</u>:



Thick black line: observations Thin lines + symbols: CMIP5\*

\*For CMIP3 see Covey et al., 2011: "The Surface-Pressure Signature of Atmospheric Tides in Modern Climate Models, *Journal of the Atmospheric Sciences*, **68**, 3874.





Diurnal time of maximum (0→24 hrs LST)

Semi-diurnal time of maximum (0→12 hrs LST)

Diurnal time of maximum (0→24 hrs LST)

Semi-diurnal time of maximum (0→12 hrs LST)