An effective hyper-resolution pseudo-3D implementation of small scale hydrological features to improve regional and global climate studies

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### Introduction: Current hydrology in CLM





No lateral subsurface flow taken into account!

# Model resolution increases due to computational developments



Importance of:

- river network decreases,
- lateral subsurface flow increases,
- elevation differences increase.

Question: How to improve the CLM's hydrology while still making use of vertical column information and focus?



Separating the vertical and lateral hydrological response (pseudo 3-D):

- Vertical response:
  - Use CLM soil column,
  - Extend depth of column to bedrock,
- Lateral response:
  - Design a new model (no linear reservoir),
  - Defining a generic approach resolution independent,
  - Differentiating different hydrological response units,
  - Here: 1 km pixel resolution (hyperresolution),
- Using different resolutions for vertical and lateral hydrological response,
- Making use of high resolution datasets.





### Make use of high resolution datasets



- Vegetation data
- Digital elevation model (DEM)
- Soil databases

Broxton et al., 2014





#### Make use of high resolution datasets



Pelletier & Rasmussen, 2009

- Vegetation data
- Digital elevation model (DEM)
- Soil databases





#### Make use of high resolution datasets



30 m DEM



- Digital elevation model (DEM)
- Soil databases





### Differentiation between different hydrological units













### Make use of high resolution DEM to differentiate between hydrological units



Fraction of the different lateral hydrological components for each 1 km pixel.



QGW

QCH

QGW

Q<sub>CH</sub>

Q



#### A hybrid 3-D approach for the hillslope type



The 1-D Richards equation:

$$\frac{\partial \theta}{\partial t} = \frac{\partial}{\partial z} \left[ K_{\nu}(\Psi) \left( \frac{\partial \Psi}{\partial z} + 1 \right) \right] - G(h)$$

The 1-D hillslope storage Boussinesq equation (Troch et al., 2003)

$$\frac{\partial}{\partial t}(f(h) h) = \frac{1}{w} \frac{\partial}{\partial x} \left( w K_{l} h \left( \sin \alpha + \frac{\partial h}{\partial x} \cos \alpha \right) \right) + \cos \alpha R_{gw}$$

$$\sum_{\text{OF ARIZONA}} \text{THE UNIVERSITY} OF ARIZONA$$

### Comparison of hsB and CLM at the hillslope scale









## A hybrid 3-D approach for lateral interaction between pixels





The 1-D Richards equation:

$$\frac{\partial \theta}{\partial t} = \frac{\partial}{\partial z} \left[ K_{v}(\Psi) \left( \frac{\partial \Psi}{\partial z} + 1 \right) \right] - G(h)$$

The 2-D subsurface Boussinesq equations (Darcy)

$$A\frac{\partial(hf)}{\partial t} = -\partial \left\{ w_x h k_L \left( \frac{\partial h}{\partial x} + \frac{\partial E}{\partial x} \right) \right\}_x -\partial \left\{ w_y h k_L \left( \frac{\partial h}{\partial y} + \frac{\partial E}{\partial y} \right) \right\}_x + R_{gw} A$$
**NCAR**

## Computational scheme for hybrid 3-D hydrological model









- Landscape Evolution Observatory (LEO) in Biosphere 2
- Convergent hillslope (11x30 m)
- Recharge and drainage experiment
- Compare to 3D Richards model CATHY (Niu et al.,







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Good correspondence with observations!



a)





We are developing a new hybrid 3-D hydrological scheme for CLM:

- ► Keep the current vertical column framework of CLM,
- Differentiate between lateral and vertical response,
- Extending the depth of the vertical column to the bedrock,
- Identifying the response of different hydrological units (e.g. hillslope, flatland, river network and lake).

For next period:

- Finish implementation of shallow water equations for river network and lake,
- Couple to vertical column structure of CLM,
- Test the possibilities of this model at different scales (catchment, continental, global).
- ► For longer run: Add possibility of using multiple vertical columns.





Questions?

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