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NATIONAL LABORATORY

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# Evaluating the resolution dependence of aerosol-cloud interactions in model simulations using A-Train observations

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## Acknowledgement:

CESM aerosol liaison (Ghan), DOE early career grant (Gustafson), DOE SFA ACP (Rasch)  
NCAR CISL (Yellowstone, RDA), PNNL PIC, NERSC, NASA Langley ASDC  
Heng Xiao (PNNL), Francis Vitt (NCAR), Kai Zhang (PNNL)

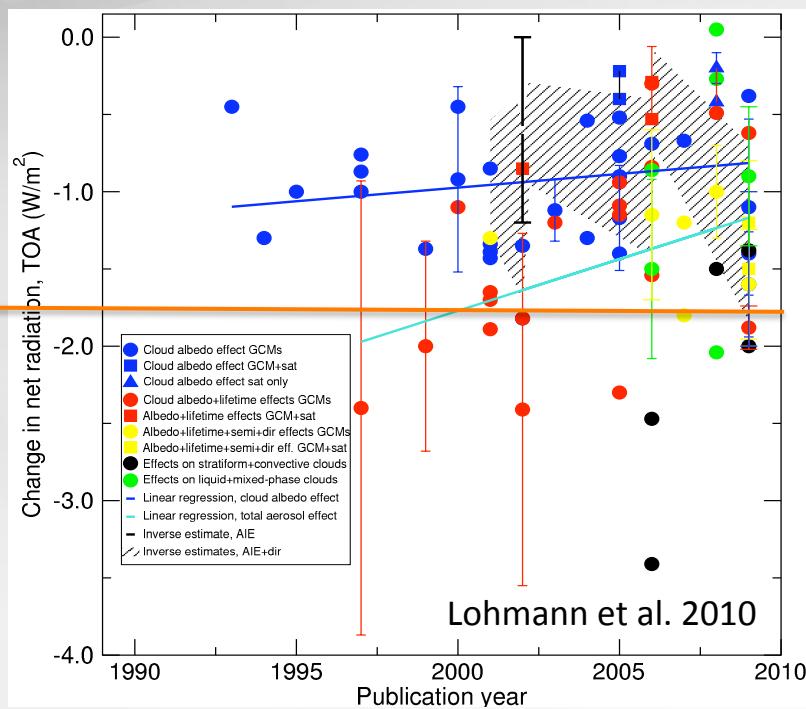


# Efforts to improve aerosol simulation

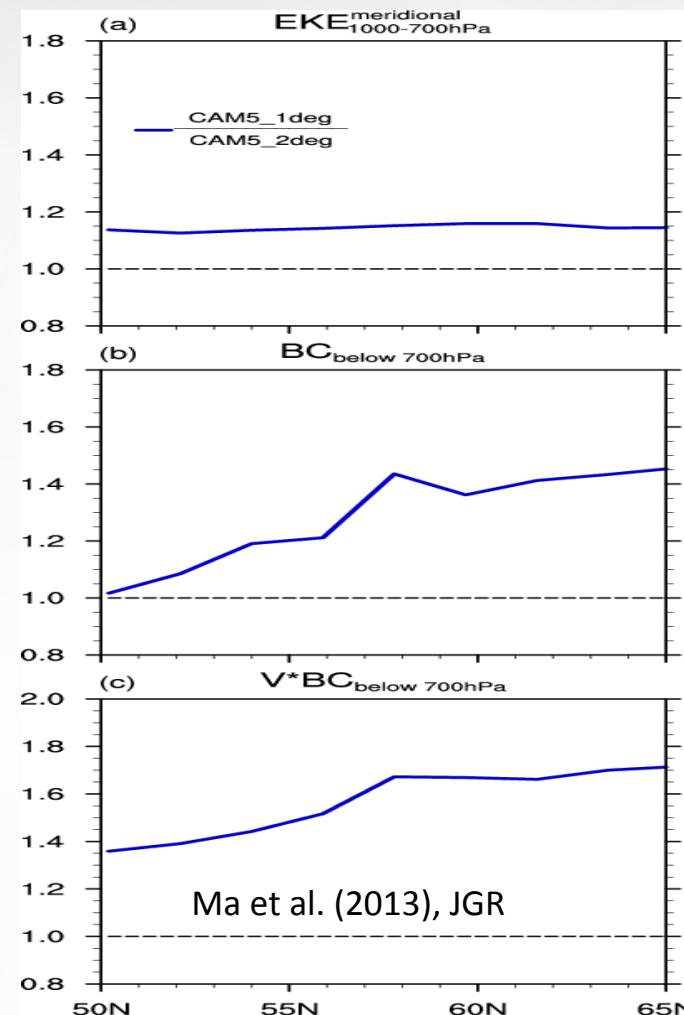
- Prescribed aerosol option (in cesm1.3)
- Diagnostic radiation for any MAM specie (in cesm1.3)
- AeroCom diagnostics (in cesm1.3 as history\_aero\_optics)
- MAM4: primary hydrophobic carbon mode added to MAM3 (PNNL, Wyoming)
- Less absorbing dust physprops file
- Unified treatment of convective transport and scavenging (PNNL)
- [Resolution dependence of aerosol simulation \(PNNL\)](#)
- Improved dust emission size distribution (Cornell, PNNL, Wyoming)
- Speciation of dust: optics (Cornell) & ice nucleation (Wyoming)
- More general aerosol thermodynamics (PNNL)
- Ammonium & nitrate (NCAR)
- Speciation of POM: hygroscopicity (PNNL)
- Ion-induced nucleation & subgrid homogeneous nuc (SUNY-Albany, PNNL)
- Marine organic sources (NC State, Harvard, LANL, Scripps, PNNL)
- Secondary organic aerosol intercomparison (MIT, NCAR, PNNL, LLNL, UM)
- Coupled fire smoke emissions (NCAR & PNNL)
- Coupled DMS emissions (LANL, ORNL, LLNL, PNNL)
- Coupling MAM to SNICAR (Flanner & PNNL)
- MAM volcanic aerosol (NCAR, PNNL)
- Geoengineering stratosphere, CCN (NCAR, PNNL)
- Frost flower sources (Scripps, LANL)

# Background

- Aerosol indirect forcing in CAM5 is **higher** than many other global models
- Resolution dependence** of aerosol transport may imply resolution dependence of aerosol distribution and aerosol-cloud interactions

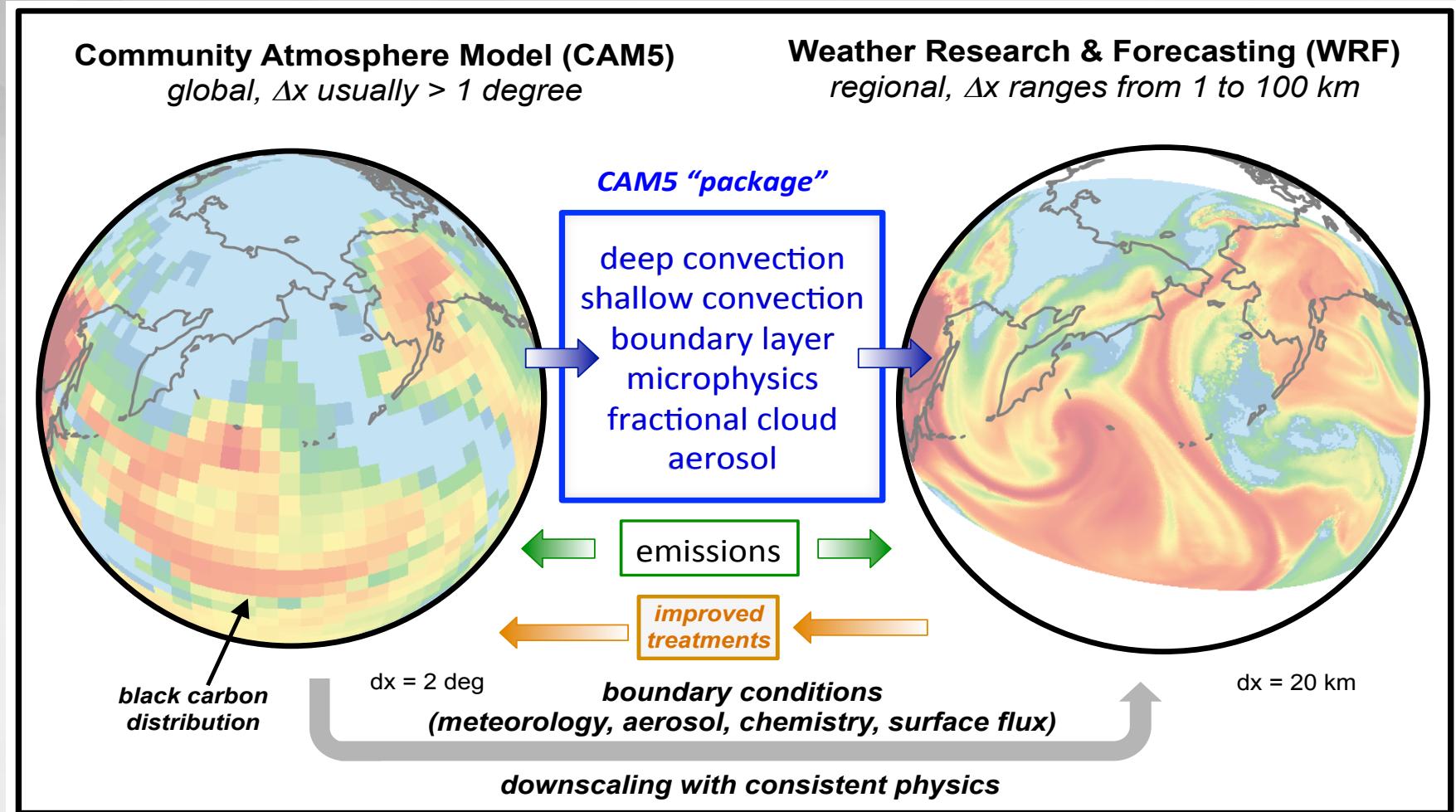


- Aerosol-cloud interactions can be amplified due to **data aggregation** (McComiskey and Feingold, 2012)



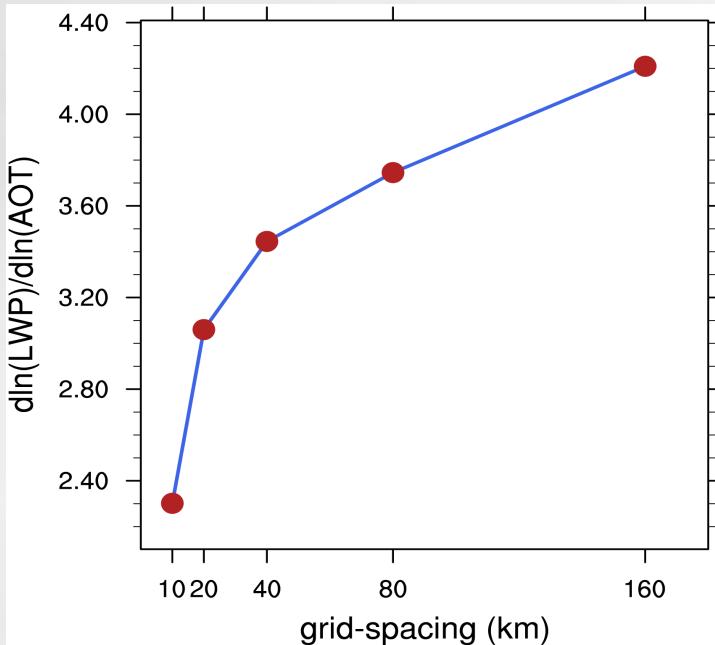


# Downscaling from CAM5 to WRF-Chem

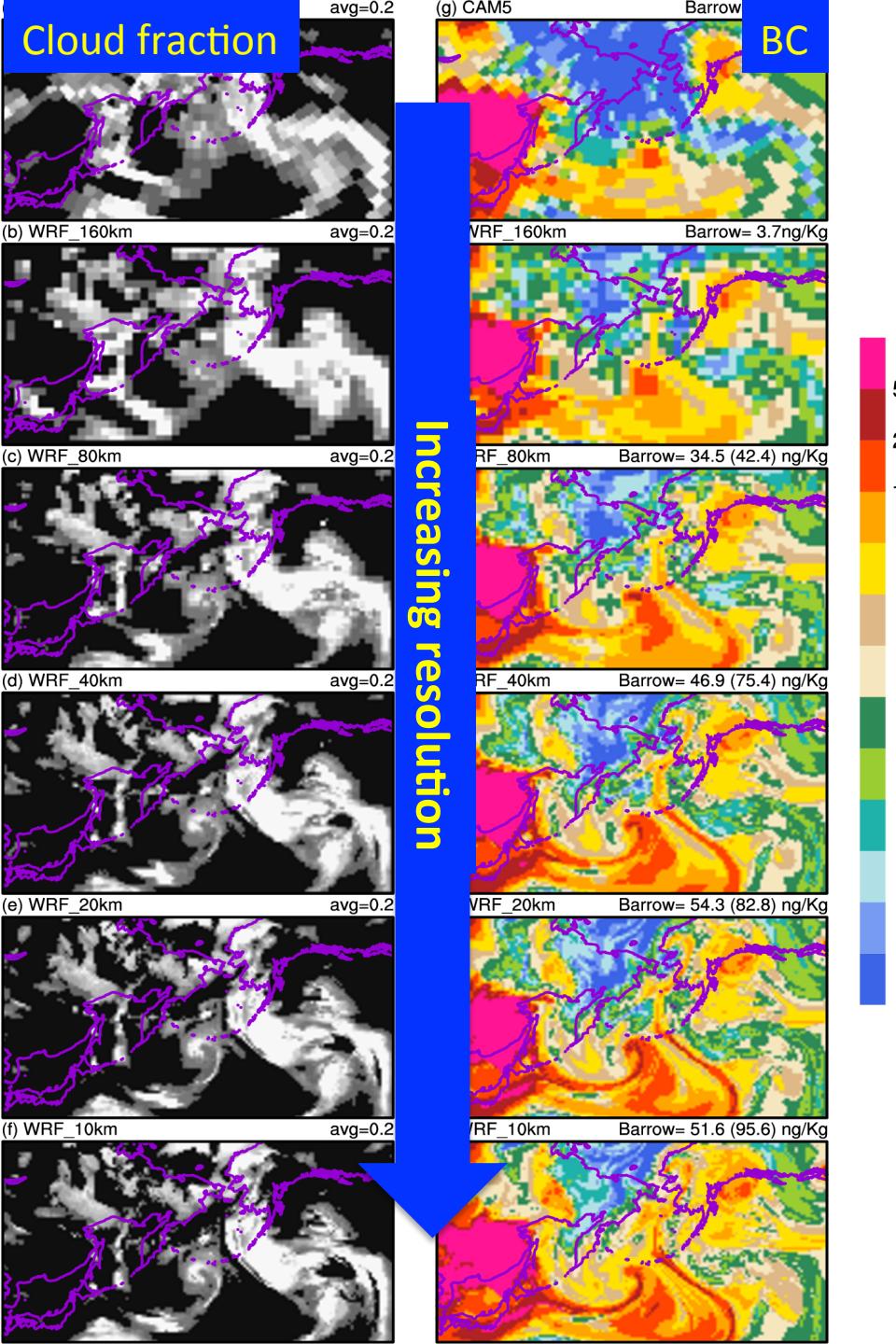


# Long-range transport of aerosols

- Filamentary structure in high-resolution runs
- Lower cloud susceptibility in high-resolution runs



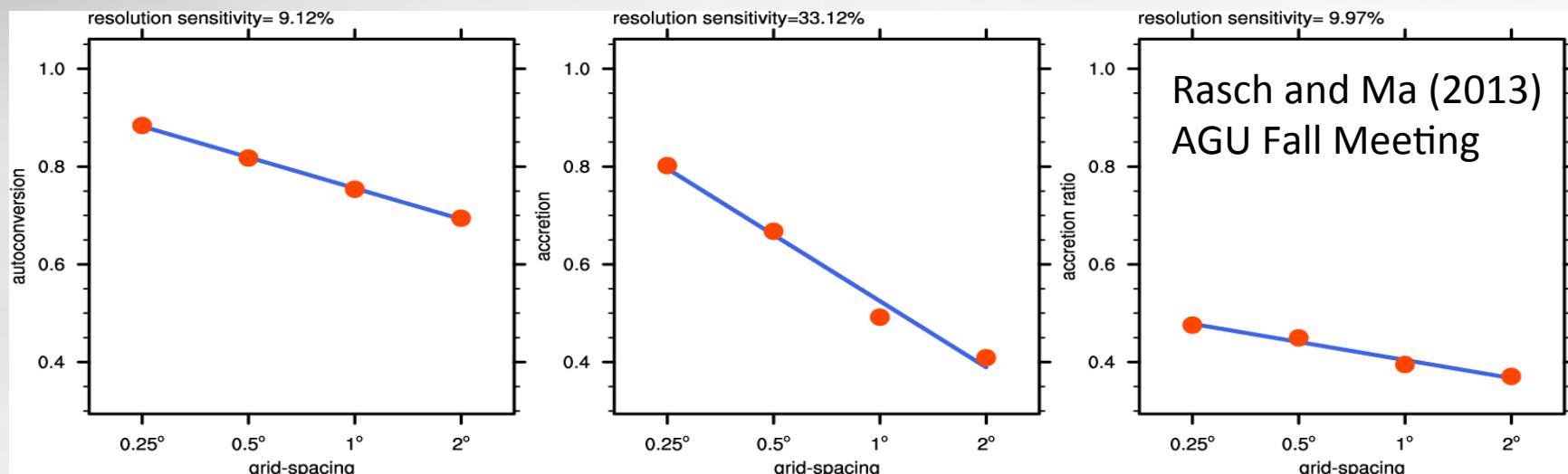
Ma et al. (2013), GMDD





# Higher resolution leads to

- precipitation intensity increases
- accretion rate increases (function of  $Q_r$ , not a function of  $N_c$ )
- drier atmosphere (less clouds), less scavenging of aerosols
- larger aerosol burden
- “convective” cloud fraction decreases + “stratiform” cloud fraction increases



## Science questions

- Is this behavior region- and season- dependent?
- What is the net effect on aerosol indirect forcing with increasing resolution?
- Do results converge with increasing resolution? Is it resolution-dependent?
- At what resolution the model produces the most realistic aerosol indirect forcing?



# Specified dynamics CAM5 simulations

- Standard offline meteorology methodology (3-dimensional wind and temperature, 2-dimensional surface wind stress, surface heat and moisture flux, surface temperature are prescribed. Wind-mass fixer is applied to ensure the time evolution of surface pressure and wind are consistent.)
- Year Of Tropical Convection ([YOTC](#)) analysis ( $0.15^\circ$ ), regridded to 2-deg, 1-deg, 0.5-deg, and 0.25-deg CAM5 grids using mass conservation interpolation
- Surface [moisture flux](#) comes from 0.25-deg offline CAM5 simulation, scaled to  $\sim 2.99$  global annual mean.
- Model [time step](#) and dynamical sub-stepping are kept the same for all resolutions
- [Model calibration](#) for aerosol, cloud, and convection parameterizations in the 2-deg configuration, and is kept the same for other resolutions.
- [AeroCom diagnostics](#) that compares properties at cloud-top.
- Direct comparison with satellite observations. Output [along satellite track](#).



# Challenges

- **Huge amount of data**
  - Pre-processing (YOTC analysis data processing)
  - I/O during runtime (writing instantaneous output is not fun)
  - Analysis (get a coffee, do emails, work on something else, go home!?)
- **Compensating errors through tuning in the free-running model due to unrealistic meteorology**
  - convectively unstable
  - cold bias in UTLS
  - weak surface wind
- **Physics not designed for very high resolution (model unstable and/or producing unrealistic results)**
- **A-Train data processing/interpretation (about 9.5 million 20-km footprints in 2009)**

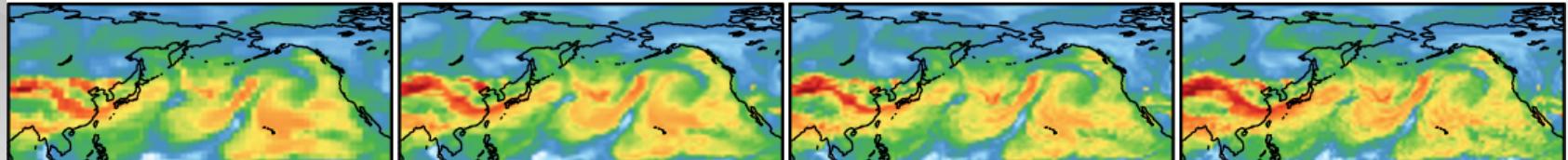


# SD CAM5 simulations

2009-03-24\_21Z

Aerosol Burden ( $\text{g/m}^2$ )

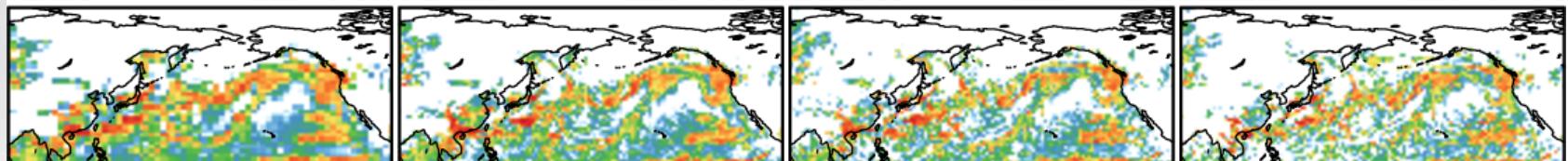
1.9x2.5 max = 3.5 0.9x1.25 max = 5.9 0.47x0.63 max = 10.2 0.23x0.31 max = 25.0



0.001 0.01 0.1 1

LWP ( $\text{g/m}^2$ )

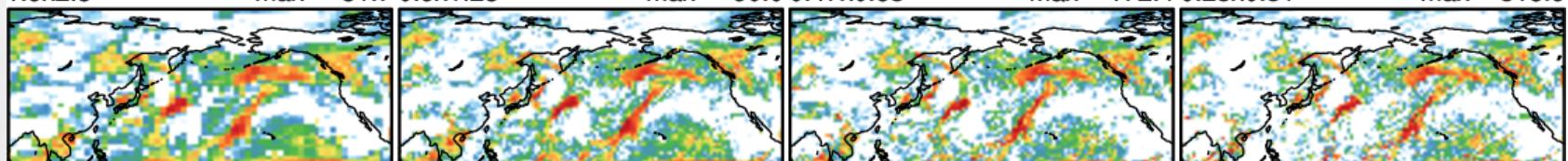
1.9x2.5 max = 949.4 0.9x1.25 max = 1067.8 0.47x0.63 max = 1157.6 0.23x0.31 max = 1771.6



1 10 100 1000

Precipitation ( $\text{mm/day}$ )

1.9x2.5 max = 51.7 0.9x1.25 max = 90.0 0.47x0.63 max = 172.4 0.23x0.31 max = 318.9

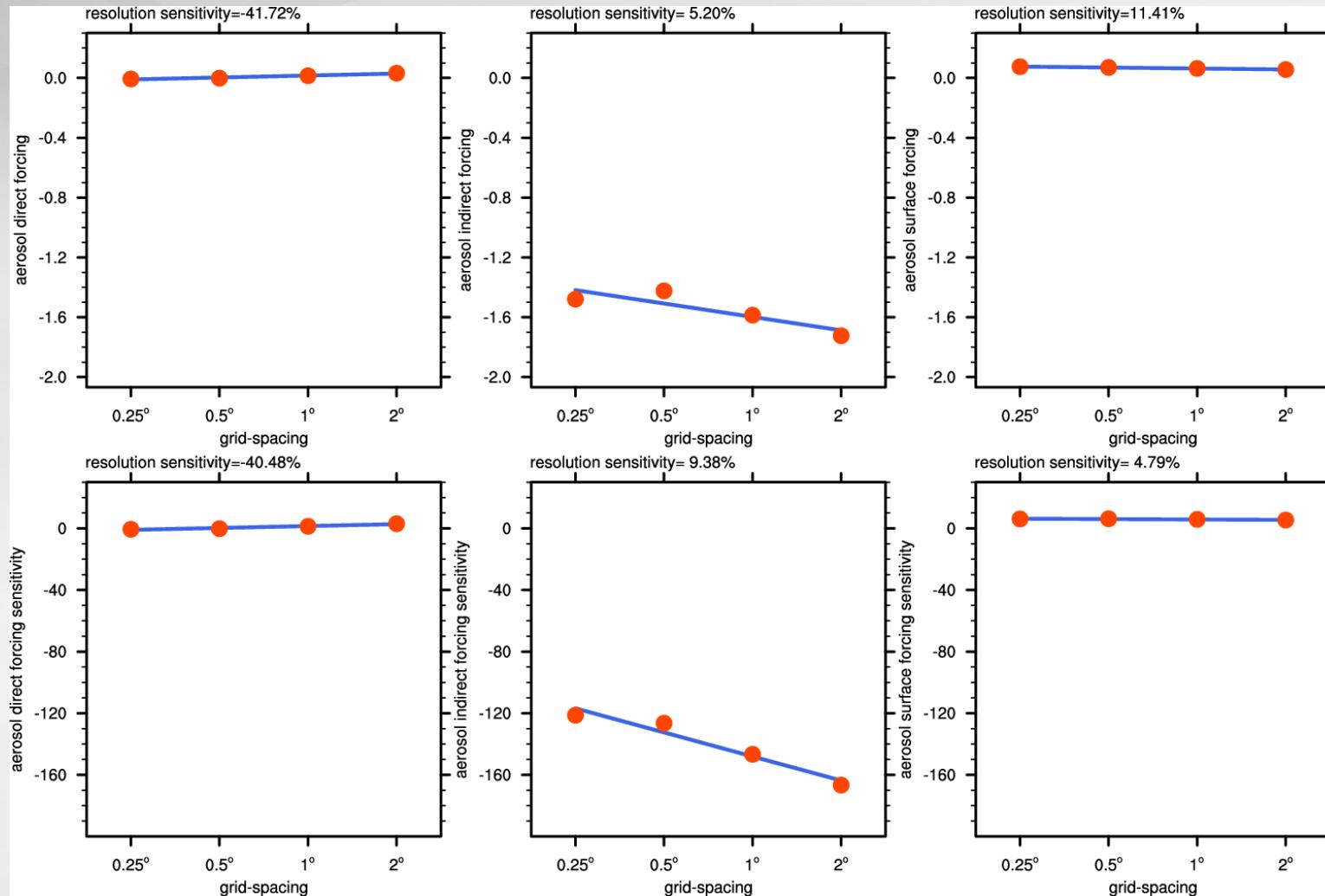


0.05 0.5 5 50



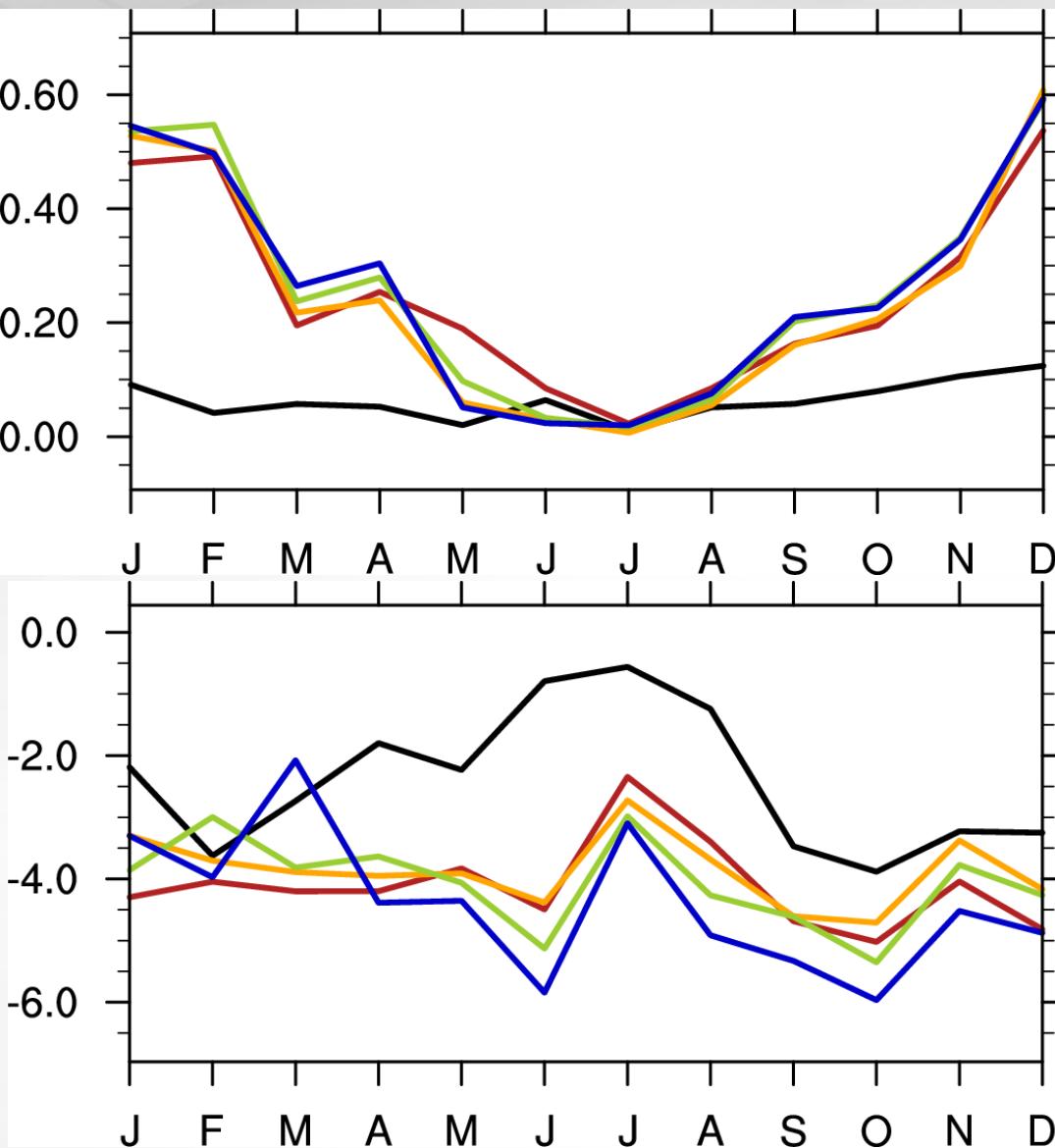
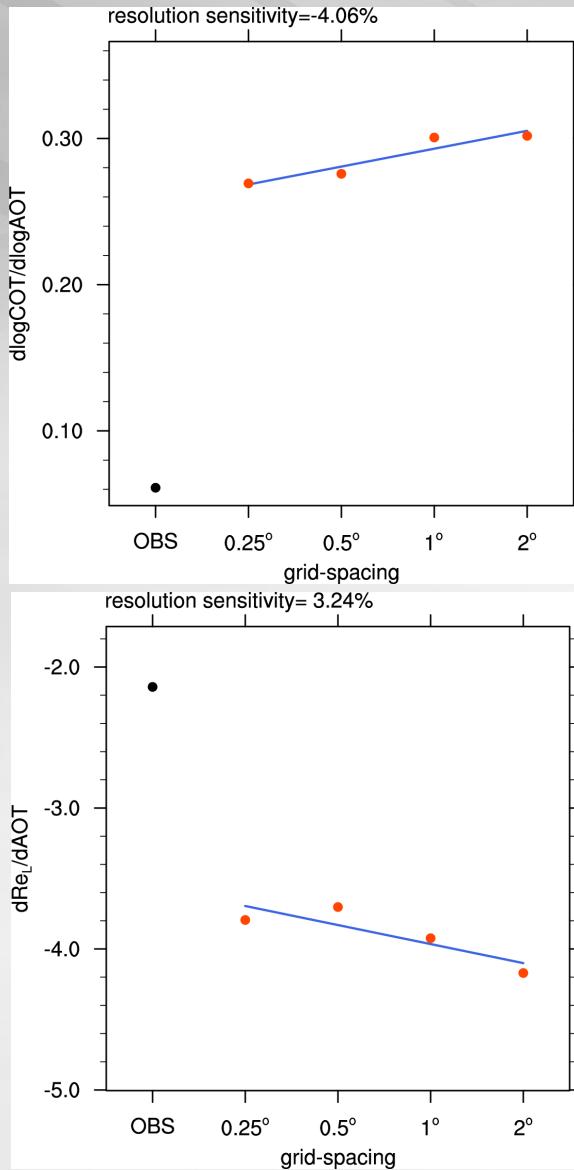
# Aerosol forcing decomposition

- Ghan (2013) methodology



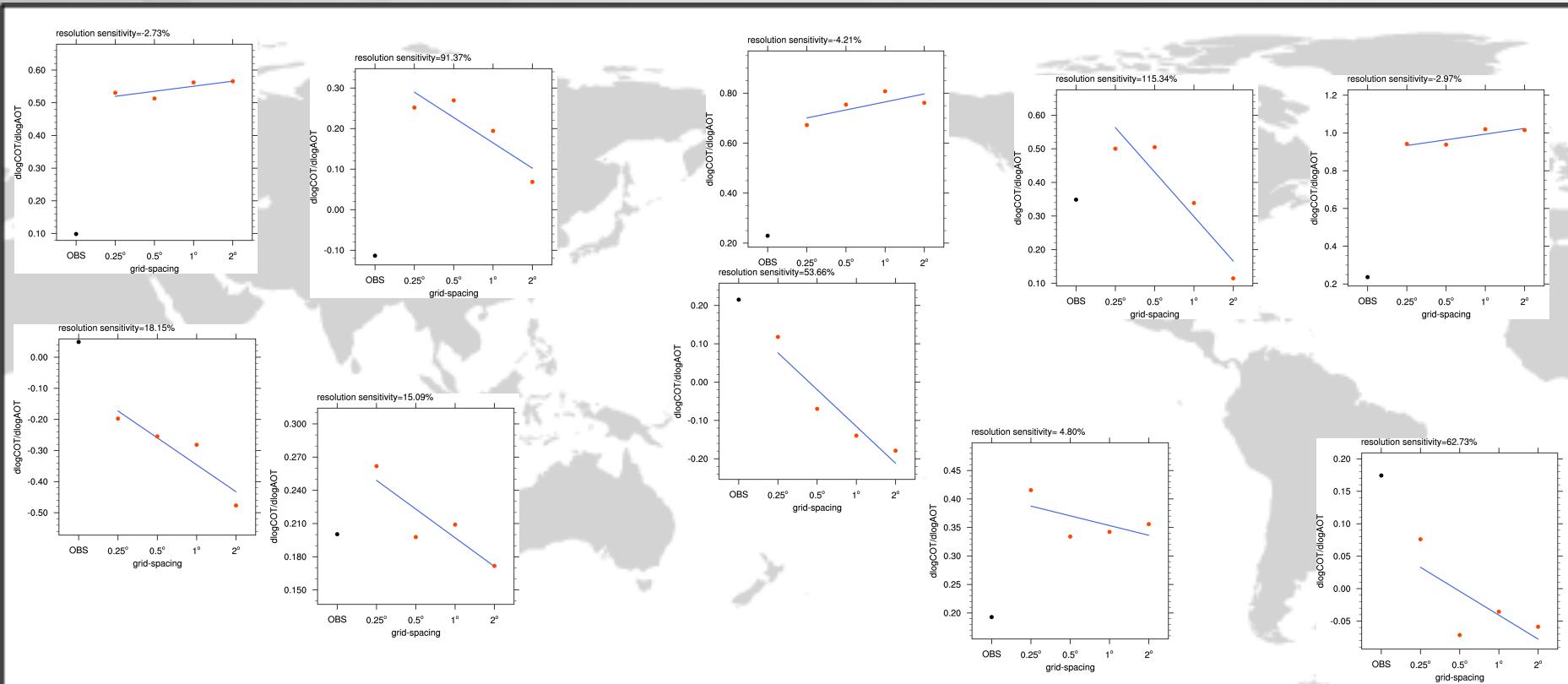


# Cloud susceptibility



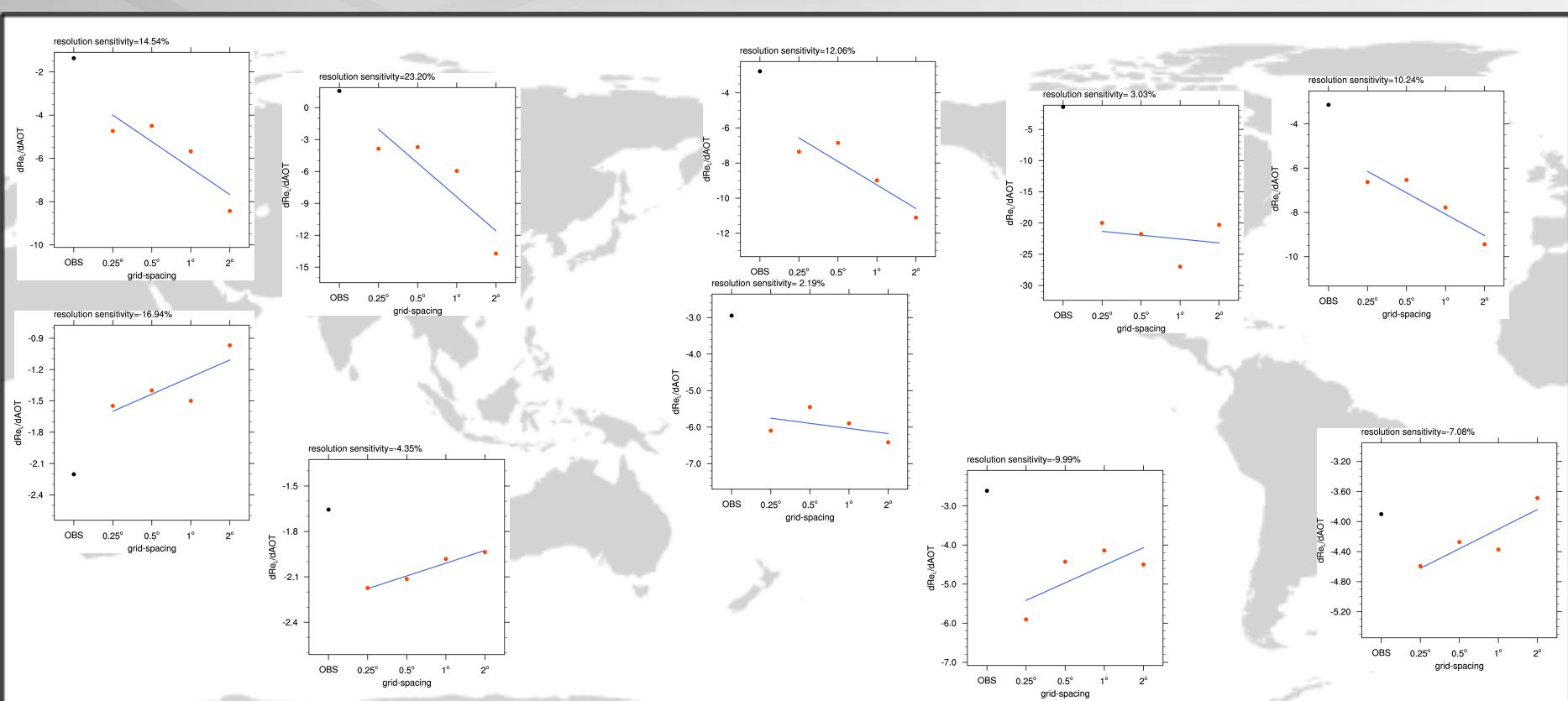


# Cloud susceptibility





# Cloud susceptibility first-indirect effect





# Next steps

- More careful analysis on region-dependence of cloud susceptibility
- $S_{\text{pop}}$  calculation to evaluate aerosol second indirect effect
- Ice cloud susceptibility evaluated against CloudSat and CALIPSO
- Gaussian process emulation to provide insights into physical mechanisms
- Understand the uncertainty range of the satellite observations
- Do I want to evaluate vertical profiles? Pros and cons to be considered.
- Weakly nudged runs to explore the behavior of the free-running model.
- Quantify bias due to sampling difference
- Use COSP to remove bias due to algorithm difference
- Aerosol transport into the Arctic