

Ill-Posed Glacier Volume Estimation

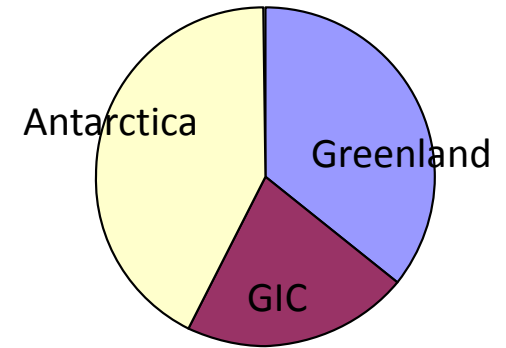
AKA: Please Don't Shoot the Messenger



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Who Gives A Rat's Derriere About GIC Volume?

- 200,000+ glaciers and ice caps (GIC), but we only have measured volumes for about 100. **Ouch.**
- GIC will contribute about 1/3 of SLR over next 100 years.
- GIC are a leading term in today's sea-level rise.
- Can't predict their contribution to SLR if we don't know their volume.



Approx. SLR contributions from ice over next 100 years

So How Do We Calculate Volume?

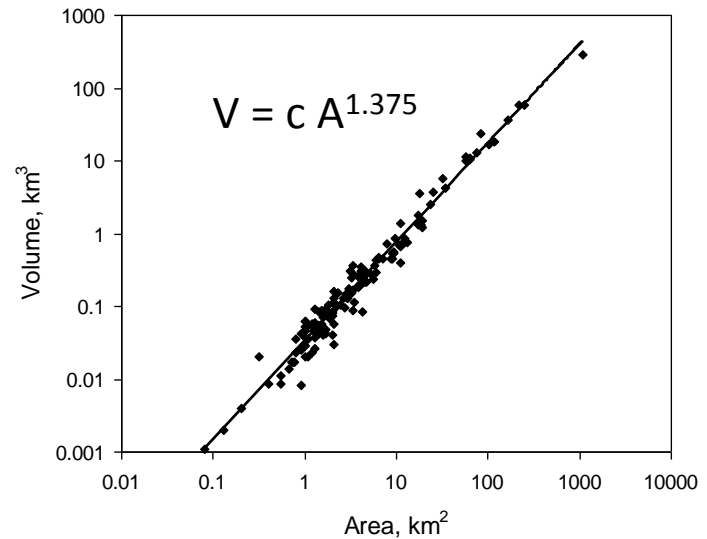
- Volume-Area Scaling

- Cute, simple, direct.
- **Imprecise?** (Stay tuned!)

- Numerical Inversions

- Complex but versatile.
- Hard to model 200,000+ glaciers, but it has been done!
 - Huss and Farinotti, 2012

- **Presumably more precise?**



$$h_i = \sqrt[n+2]{\frac{(1 - f_{sl}) \cdot q_i}{2A_f(T)} \cdot \frac{n + 2}{(F_{s,i} \rho g \sin \bar{\alpha}_i)^n}}$$

Which Approach is More Accurate?

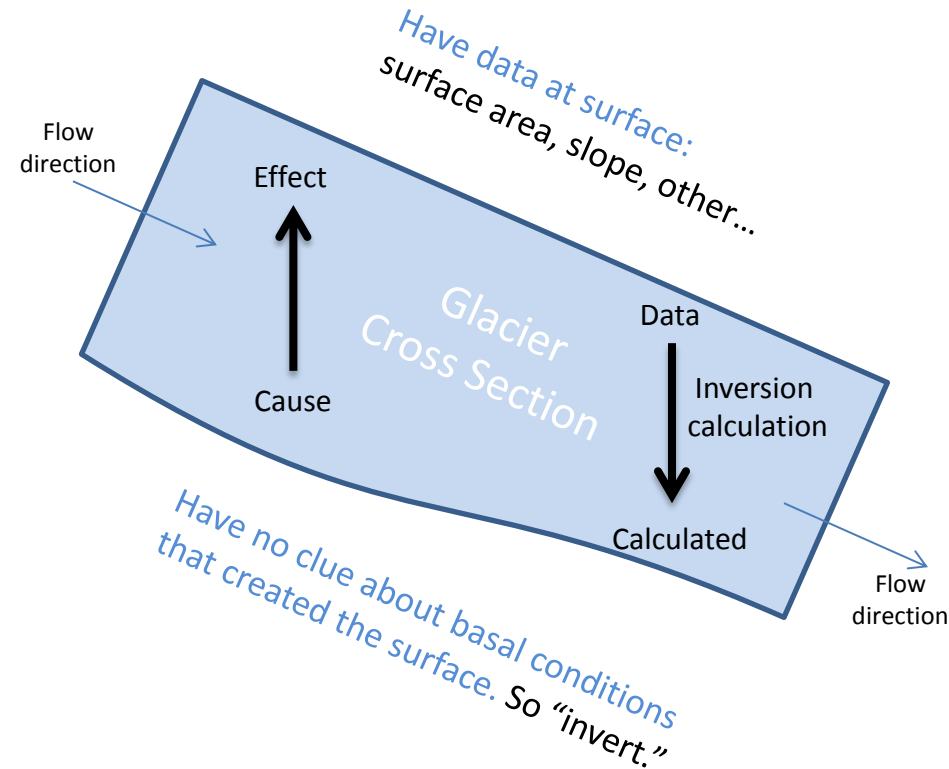
Modeling Versus Scaling

- Neither.
 - (Please don't shoot the messenger.)
- Both.
 - They give darn near the exact same accuracy and precision.
 - Can prove it.



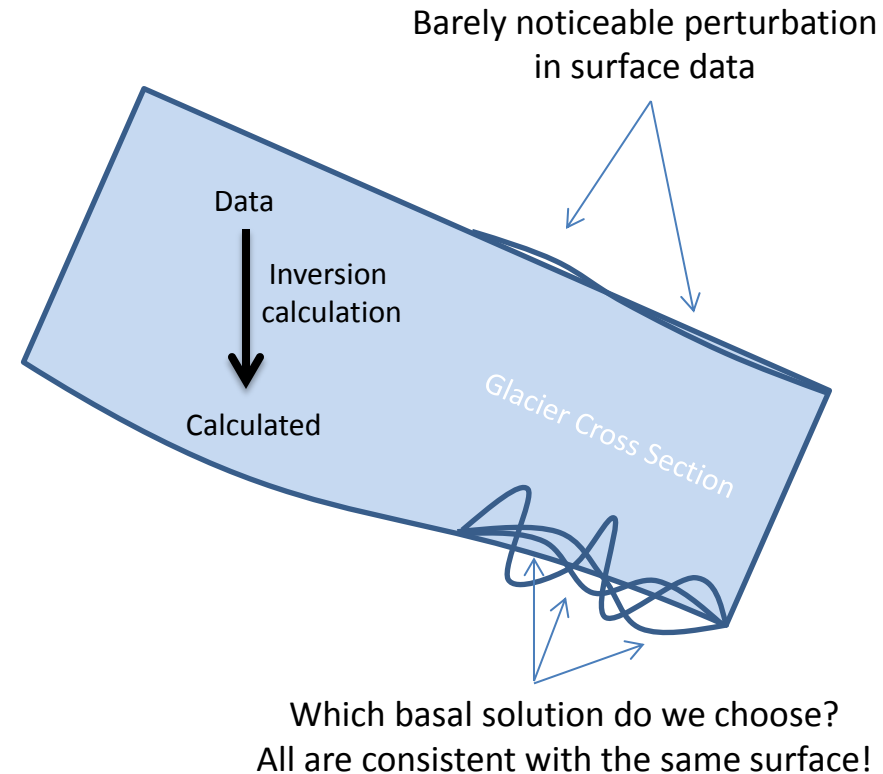
The Problem? Mathematical Inversion

- Basal conditions drive the surface conditions.
 - Slip at the bed determines velocity at surface.
 - Topography at bed causes bumps on the surface.
- So use observed “effect” to derive the “cause.”
 - Backwards!
- Called an inversion.



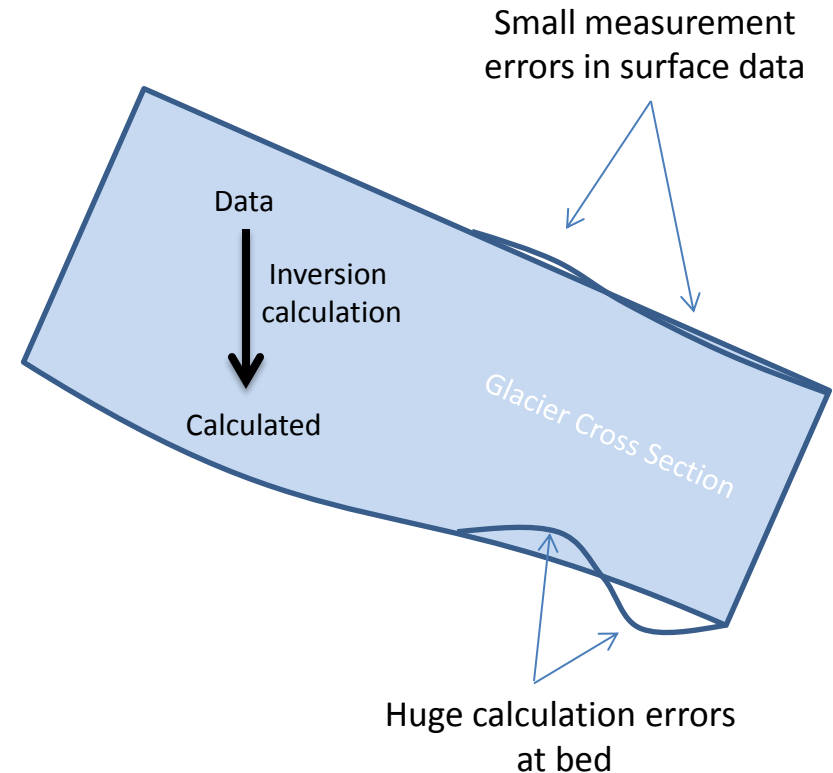
Cloud of Solutions

- Lots of bed conditions correspond to the same surface.
 - Different velocities, stresses, topographies...
 - Think of glacier as the viscous equivalent of a sponge. Wiggle the bottom of the sponge and will barely see it at the surface.
- So which one do we choose?
 - All fall within the measurement errors at surface.



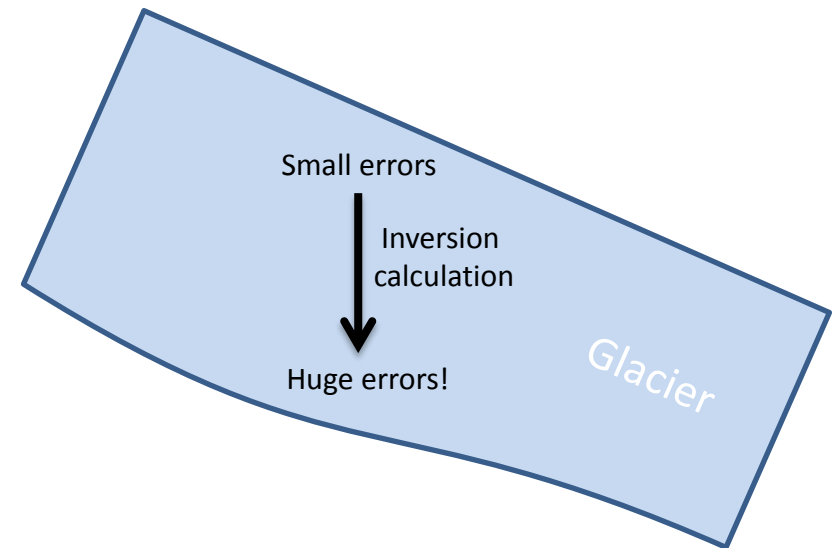
Unstable Solution!

- Small errors in surface data translate to huge calculation errors at the bed.
 - Unstable.
- Note: We assume that nothing whatsoever is known about the bed.
 - This is the case for most GIC.
 - For ice sheets we often know basal topography and that can be a game changer.



Glacier Inversion is Ill-Posed

- Numerical inversions are often “ill-posed.”
 - This does *not* mean we set up the problem incorrectly.
 - Most geophysical inversions are ill-posed.
- Ill-posed means the solution is unstable or not unique.
 - A cloud of possible solutions.
 - **Mathematically impossible** to pick the correct one.



Balise and Raymond (1985), **Lliboutry (1987)**, MacAyeal (1993), **Bahr et al (1994)**, Truffer (2004), Chandler et al (2006), and many more!

Worst at High Frequency And Deep Depths

$$\Delta \hat{\sigma}'_{ij}(k, z) = \Delta \hat{\sigma}'_{ij}(k, 0) e^{0.6kz}$$



Exponential is huge at large k and z .

Exponential is small at small k .

Want $0.6kz \ll 1$.

Want Small Exponential

$$0.6 k z \ll 1$$

Spatial frequency

$$\lambda \ll 0.6 (2\pi) z$$

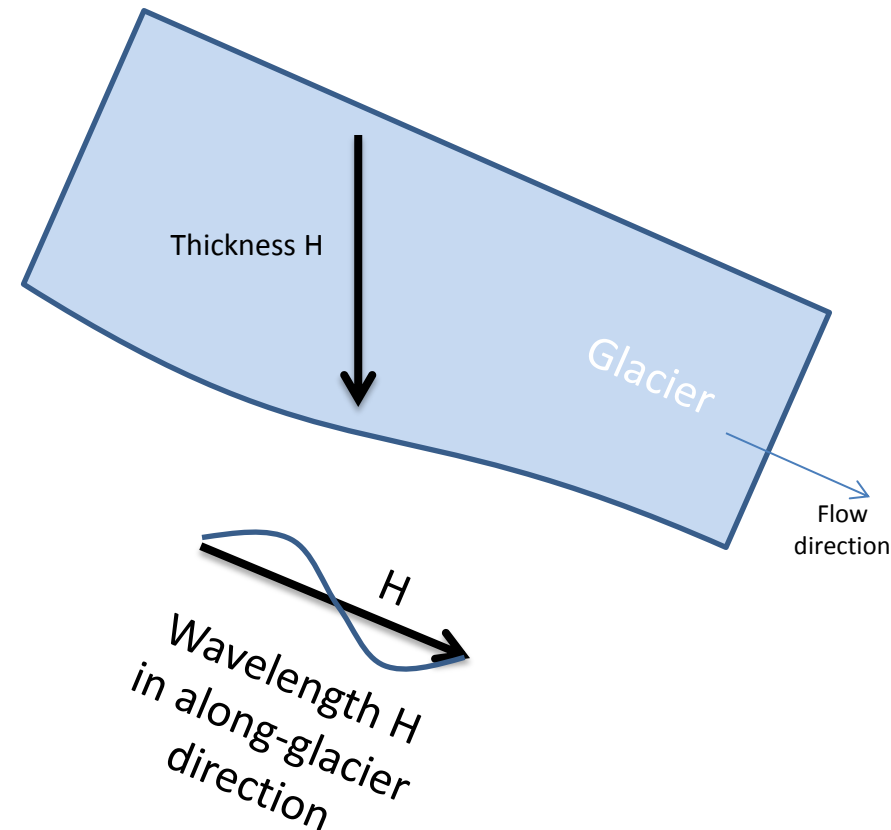
Equivalent spatial wavelength.

$$\lambda \ll 4 z$$

Wavelength should be order of magnitude (10 times) smaller than $4 z$.

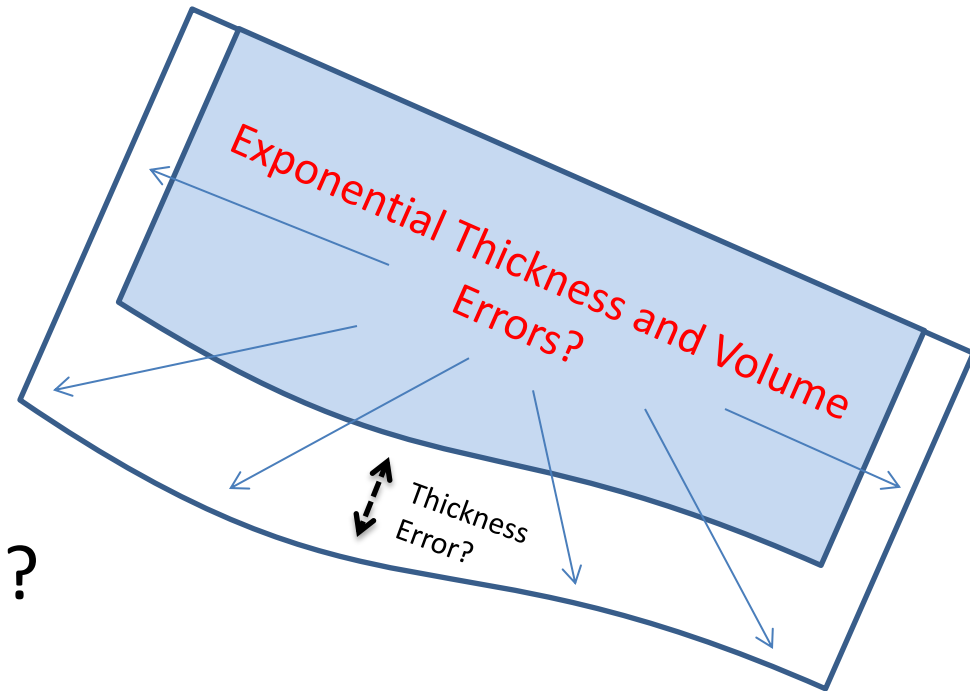
Practical Implication

- Want exponential to be negligible.
- High spatial frequencies cause big exponential.
- **So remove all spatial wavelengths less than $40 H$.**
 - In along-glacier direction. Not depth.



But Wait! That's Just Stress?

- Stress errors grow exponentially.
 - From perturbation analysis of continuum mechanics.
 - Bahr et al (1994).
- What about velocity, thickness, **volume**, etc.?
 - Remember? We want volume errors for SLR!



Scaling Solution!

- All glaciological continuum parameters scale with all others.
 - From dimensional analysis.
 - Also from stretching transformations of continuum equations.
 - Bahr (1997), Bahr et al (1997), Bahr and Rundle (1995)
- Any one implies all others.
 - Observed are in blue.

$$V = cA^\gamma$$

$$V = c_v \sigma'_{ij}{}^{\gamma_v}$$

$$H = c_H \sigma'_{ij}{}^{\gamma_H}$$

$$A = c_A \sigma'_{ij}{}^{\gamma_A}$$

$$L = bA^\beta$$

⋮

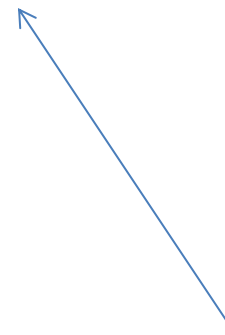
Lots of Little Devils in the Details

- But bottom line:
Can transform stress solution to any other parameter.
 - Detail: Errors in surface stress are *always* there, either measured or calculated by model.
 - So measured/calculated errors in stress translate to errors in volume, thickness, etc.

$$\Delta \hat{\sigma}'_{ij} = \Delta \hat{\sigma}'_{ij}(k, 0) e^{0.6kz}$$

$$\Delta \hat{V} \propto \Delta \hat{\sigma}'_{ij}(k, 0) e^{0.6kz}$$

$$\Delta \hat{H} \propto \Delta \hat{\sigma}'_{ij}(k, 0) e^{0.6kz}$$



Assorted scaling constants creep in here.

What's It Mean for Volume-Area Scaling?

- The wavelength that applies to scaling is $\lambda = 2L = 2A^{0.625}$.

- Substitute and get $\Delta\hat{V} = \Delta\hat{\sigma}'_{ij}(k, 0)e^{0.064 A^{-0.25}}$

- Effectively, no error growth!

Approaches zero for $A > 1$




What's It Mean For Numerical Inversions?

- Must filter volume solution at all wavelengths shorter than $40H$.
- Suppose dx = along-glacier grid spacing of model.
 - $\lambda = 2dx$
 - Grid spacing must be $20H$.
 - Can barely resolve small glaciers.
 - Will only have a few grid points in large glaciers.

Measured Area (km ²)	Expected Thickness (km)	Expected Length (km)	Grid Spacing dx (km)	Number of grid points dx in model
1	0.03	1	0.6	1
10	0.08	4	1.5	3
100	0.19	18	3.6	5
1000	0.45	75	8.5	9
10,000	1.08	316	20.3	16

Well, that's depressing.



What About Other Errors, Huh?

- Irrelevant compared to ill-posed errors.
 - Ill-posed errors are *exponential*.
 - Will swamp all other errors!
 - For example, $V = c A^{1.375}$, but c is poorly constrained. That means a *linear error* in V with linear error in c .
 - That's trivial compared to an exponential!
 - Unknown numerical model parameters. Ditto.

Upshot for Volume Calculations

- Can't do much better than scaling.
 - Simple 😊
- Numerical models work just as well *if filtered*.
 - But sooo complicated 😞
- Lots of other sources of error.
 - Irrelevant if ill-posed errors are not controlled!
- Many good reasons to use models.
 - E.g., estimate englacial velocity and stress.
 - But volume is not one of them.

