



Southern Ocean biases

Matthew C. Long Gokhan Danabasoglu, Peter Gent, and William Large

Climate and Global Dynamics Division National Center for Atmospheric Research

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Outline

Southern Ocean biases

Weak tracer uptake (CFCs,C_{ant})

Shallow mixed layer depths

Hypotheses:

Problems with:

- 1. missing forcing;
- 2. vertical physics;
- 3. mesoscale mixing.



Regional-mean mixed layer depths



Mechanisms controlling uptake in the Southern Ocean Vertical CFC11 transport

within Antarctic Circumpolar Current



- GM-bolus term opposes Eulerian-mean advection;
- Diapycnal mixing and isopycnal diffusion (Redi) dominate transport into the interior.

Seasonal handoff: boundary layer to isopycnal mixing

Vertical fluxes in ACC



Missing physics? September mixed layer depth: Obs and biases in CORE-forced runs



Downes et al. in prep

Hi-res ML biases

August-October MLD biases



 0.25° Community Atmosphere Model (CAM5, spectral element) 0.1° Parallel Ocean Program (POP2, 62 levels)

480

440

400

360

320

280

240

200

160

120

80

40

0

-40 -80

-120 -160

-200

-240

-280

-320

-360

-400

-440 -480

Vertical physics: K profile parameterization

Monin-Obukhov similarity theory

The vertical variation of turbulence characteristics in the **'surface layer'** depends only on the surface momentum flux (u^*) , buoyancy flux (B_f) , and distance from the boundary (d).

Key parameters

$$u^* = (|\tau_0|/\rho_0)^{1/2}$$
 $S^* = -\overline{ws_0}/u^*$ $L = u^{*3}/(\kappa B_f)$
friction velocity scalar fluctuation scale Monin-Obukhov length

Stability functions

$$\phi_m(\zeta) = \frac{\kappa d}{u^*} \frac{\partial U}{\partial z} \text{ (momentum)}$$

$$\phi_s(\zeta) = \frac{\kappa d}{S^*} \frac{\partial S}{\partial z} \text{ (scalar)}$$

are nondimensional and relate momentum and scalar fluxes to mean gradients, expressed as a function of the stability parameter: $\zeta = d/L$.

Vertical physics: K profile parameterization $w_x/(\kappa u^*)$ $G(\sigma)$ 0 Turbulent fluxes ε $\overline{wx}(d) = -K_x(\frac{\partial X}{\partial z} - \gamma_x)$.2 Diffusivity ($K \sim U \cdot L$) $K_{x}(\sigma) = w_{x}(\sigma) h G(\sigma)$ σ $(\sigma = d/h)$.6 Velocity scale $w_{x}(\sigma) = \frac{\kappa u^{*}}{\phi_{x}(\sigma h/L)}$ -5 = h/L.8 $(\zeta = d/L = \sigma h/L)$ 1.0

"In unstable conditions the turbulent velocity scales beyond the surface layer are assumed to remain constant at their $\sigma = \epsilon$ values. Without this constraint, unstable w_x values would become very large..."

Sensitivity to doubling w_x

 $2 \times w_{v}$: $\Delta HMXL$







Sensitivity to $\epsilon = 1$ (removing limits)





Zonal-mean σ_0



Unlimited with GM-Redi change



Representing w_x : Nondimensional flux profiles



Large et al. 1994

¹Large et al. 1994







Moeng & Sullivan 1994

Unstable to near neutral regions Stability parameter (1/L < 0)



Proposal: surface layer depth as a function of stability



Surface layer depth parameterization

MLD change



Prognostic thickness and isopycnal mixing coefficient

Horizontal uniformity in surface layer, attenuation with depth



Large-scale dynamics control mixed layer biases



Mixing tensor diagnosed in hi-res POP



Fox-Kemper et al. 2013

Prescribed diffusivity results



Summary

- Model solutions are sensitive to turbulent velocity scale in KPP, which has weak observational constraints;
- Surface layer thickness is likely a function of the stability regime;
- Southern Ocean mixed layers are sensitive to deep isopycnal and thickness mixing; need improved scheme governing horizontal and vertical variation;
- Waves may be an important missing forcing.