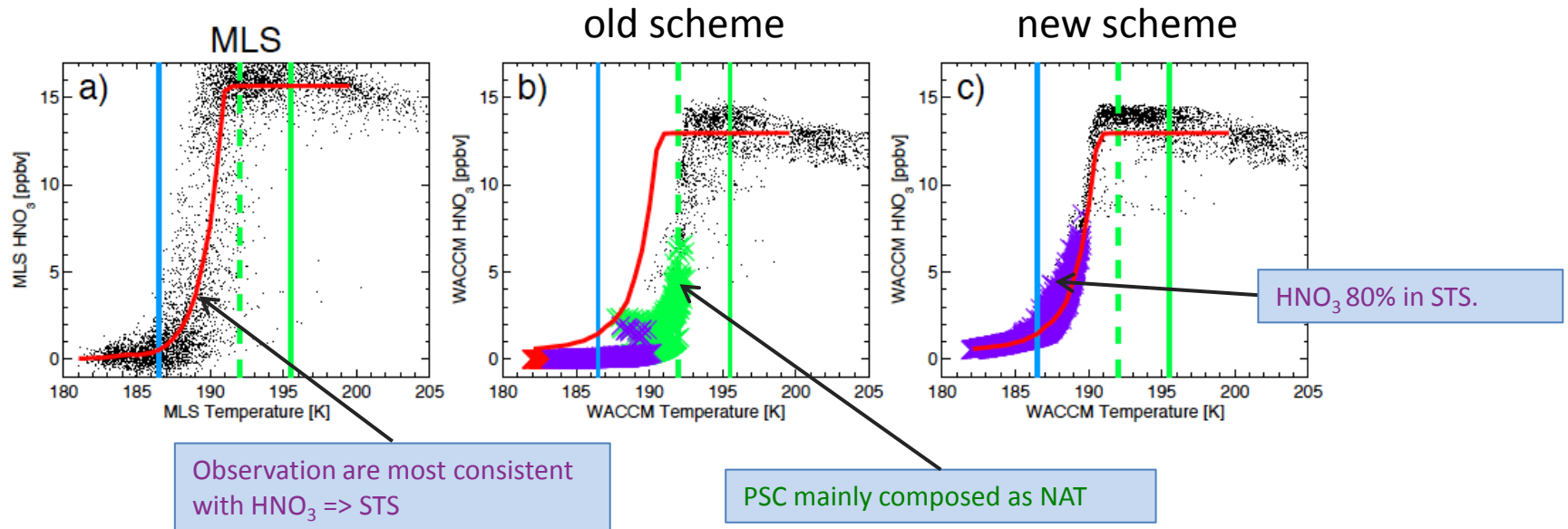


Extension of the WACCM gravity wave parameterization: motivation and results

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NCAR/ACD

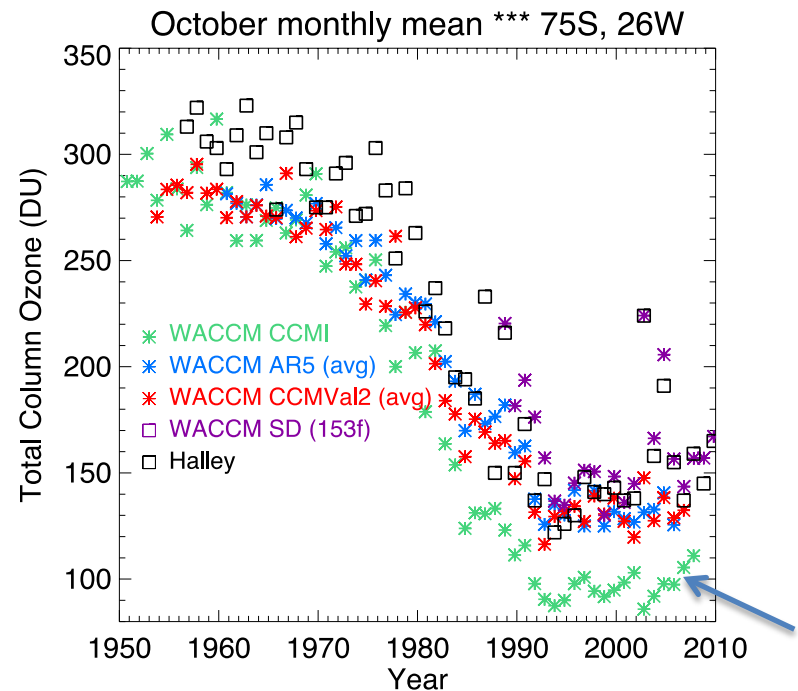
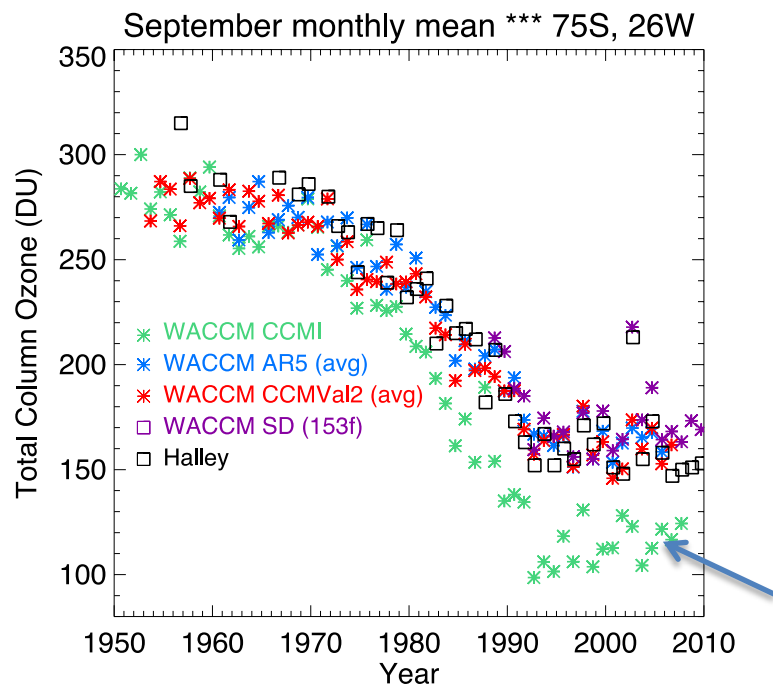
motivation: implementation of new heterogeneous chemistry module

Updated het chemistry changes partitioning of condensed-phase HNO_3 between **Nitric Acid Tri-hydrate (NAT)** and **Supercooled Ternary Solution (STS)** [see Wegner et al., JGR, 2013.]



- Updated het chemistry decreases the amount of irreversible denitrification by decreasing NAT and increasing STS
- Less denitrification allows reformation of ClONO_2 in Spring \rightarrow continued heterogeneous halogen activation
- Heterogeneous rate for halogen activation on STS is very T-dependent (the colder, the faster)
- Both these factors **require a more accurate representation of model winter/spring LS polar temperatures**

which leads to a problem

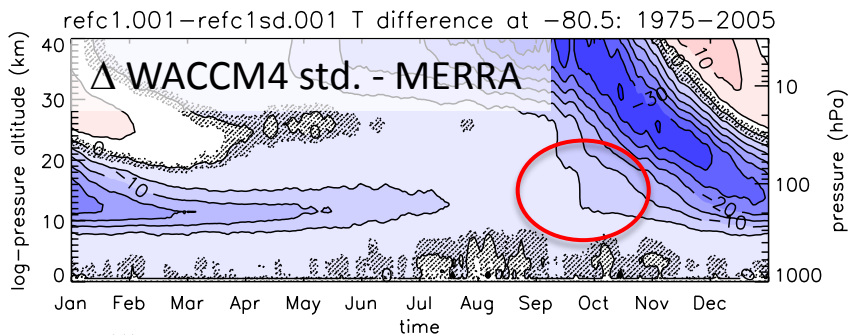
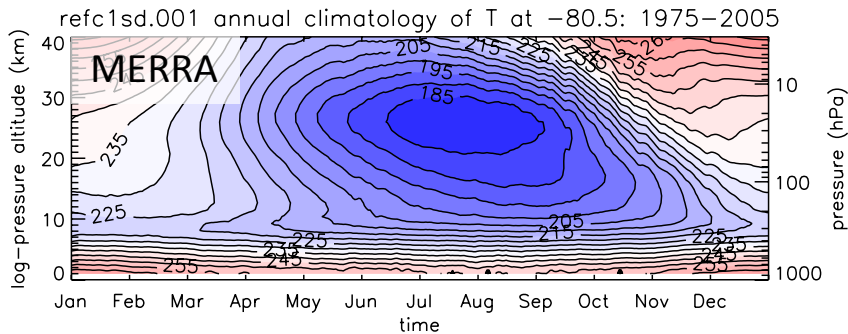
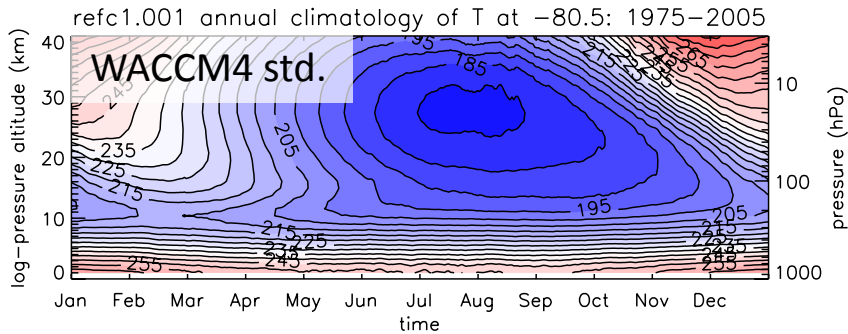


Observed and Calculated Ozone at Haley Bay

- model with **old chemistry** (red, blue) was reasonably consistent with observations
- model with **new chemistry** (green) produces unrealistically low ozone column because new het chemistry module is very sensitive the cold T

the ultimate cause of the problem

$\langle T \rangle$ (80°S) seasonal climatology 1975-2005

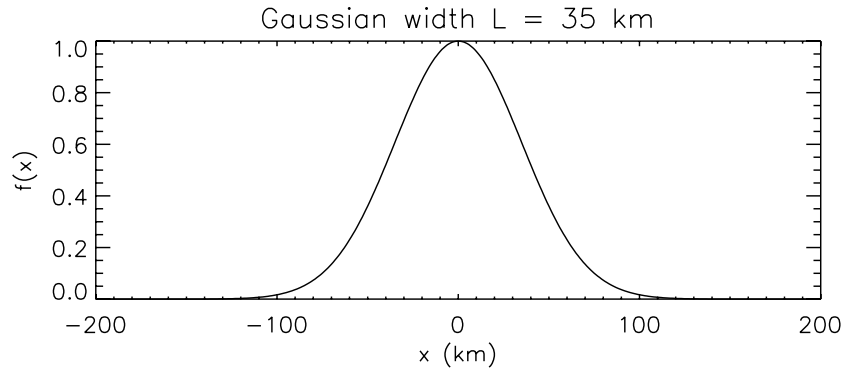


- standard version of WACCM4 has a **“cold pole” bias** in the SH
- T in ozone hole region/season is as much as 5-10 K colder than observed

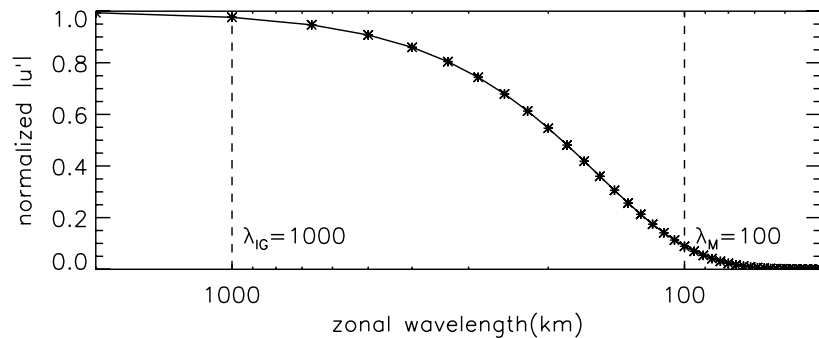
a possible solution

- polar temperatures are sensitive to wave-induced downwelling via adiabatic warming; this suggests that wave forcing is too weak in the SH
- *resolved wave* amplitudes and dissipation are not easily adjustable
- *parameterized gravity wave* forcing is adjustable, but “tuning” the parameterization to make GW break in the stratosphere degrades the simulation in the mesosphere
- parameterized GW in WACCM4 are “mesoscale” ($L_x = 100$ km); however, any physical source should excite a (“red”) spectrum in wavenumber
- → add a second spectrum of waves, with $L_x \sim 1000$ km (typical of the inertia-gravity range, IGW) to represent the effects of longer waves
- the longer IGW will have larger source amplitudes that can break in the stratosphere for reasonable values of the source stress

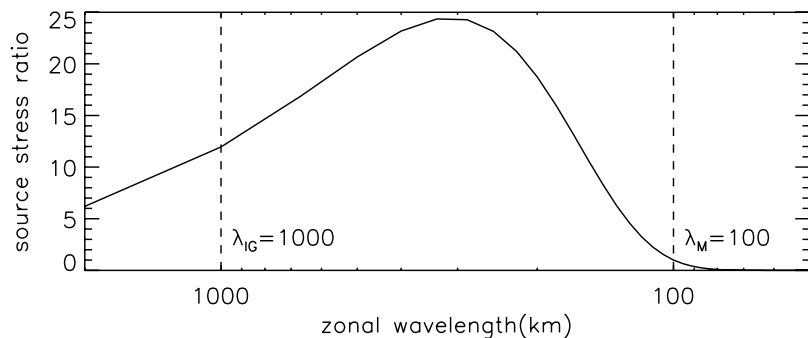
wavenumber spectrum



consider GW excitation by a “front”, idealized here as a Gaussian obstacle of width L : $w' = \langle U \rangle dh'/dx$ and, therefore, $|u'| = m \langle U \rangle |h'|$



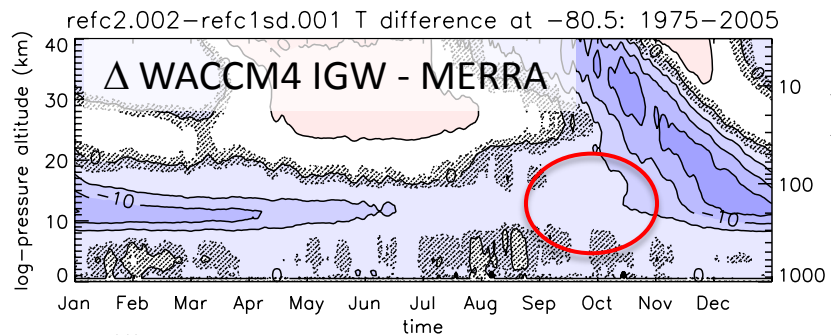
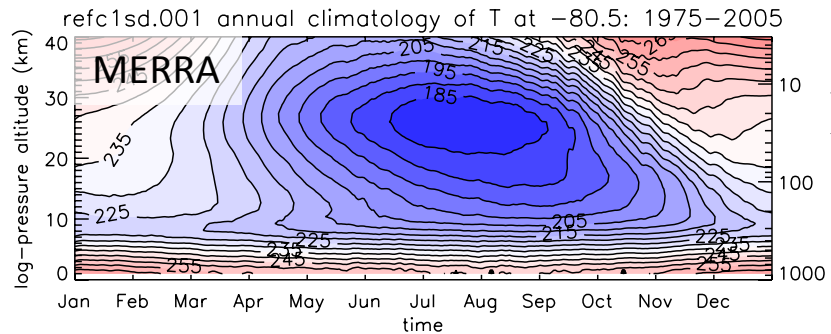
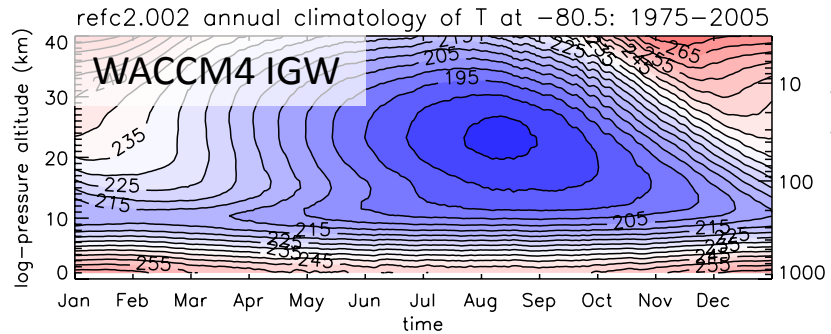
produces a Gaussian “red” spectrum in wavenumber (shown here as a function of horizontal wavelength); spectral amplitude $|u'|$ falls off rapidly at small wavelengths (large wavenumber, k)



the fall-off of spectral amplitude with decreasing wavelength means that longer waves tend to have larger source stress, $\tau = \rho (k/m) |u'|^2$, than mesoscale waves \rightarrow they break at lower altitude

<T> (80°S) with additional IGW spectrum

<T> (80°S) climatology 1975-2005



shaded: not significant at 95%

IGW spectrum uses source stress

$$\tau = 8 \times 10^{-3} \text{ Pa}$$

mesoscale GW spectrum uses

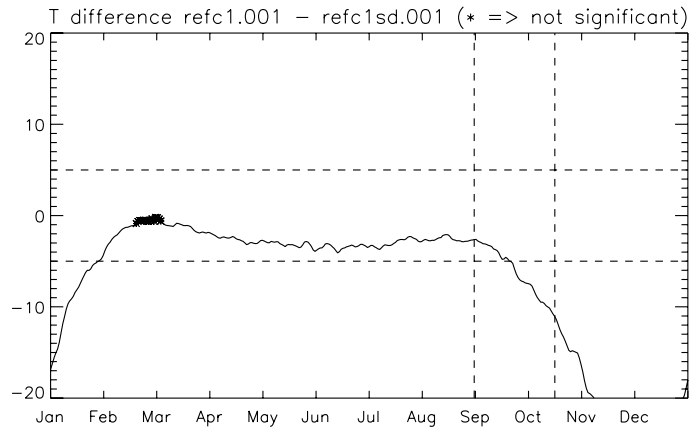
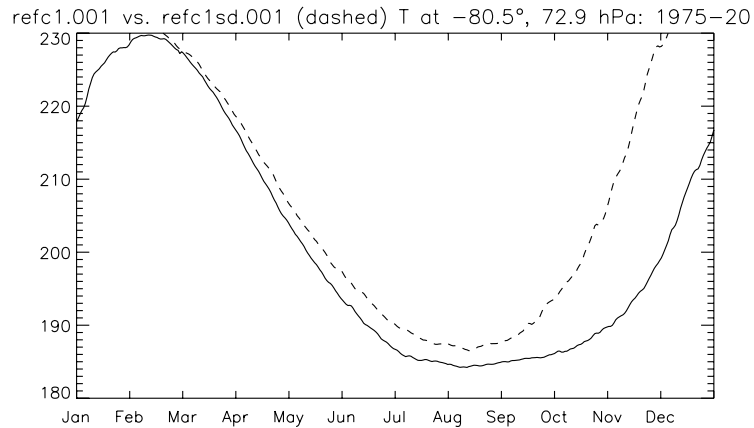
$$\tau = 1 \times 10^{-3} \text{ Pa}$$

these values are consistent with simple theoretical arguments outlined in previous slide

T in ozone hole region in SH spring is now much warmer

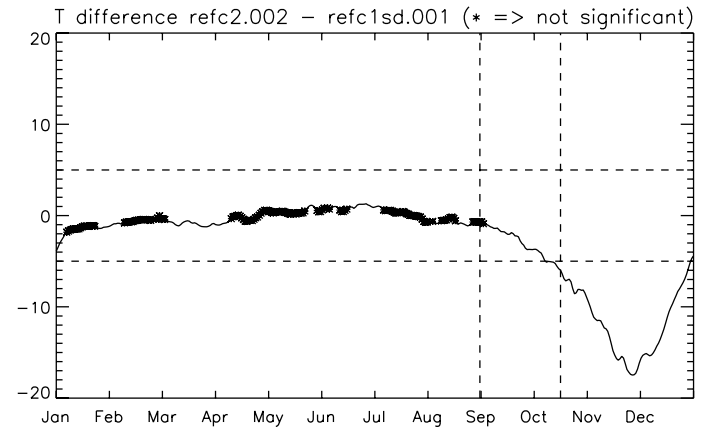
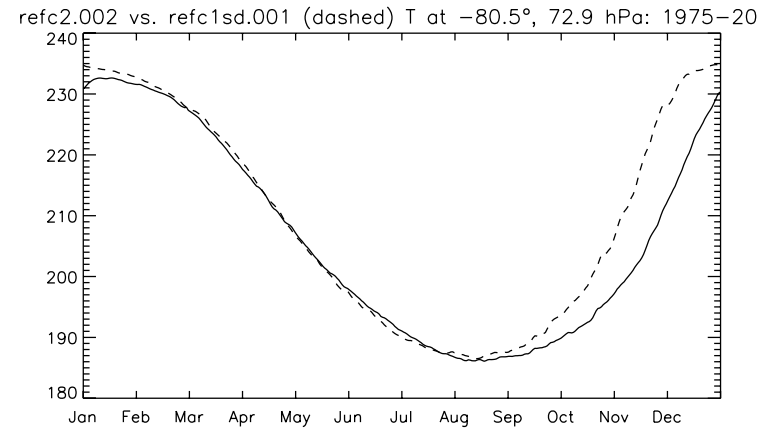
$\langle T \rangle$ @ 70 hPa, 80°S

WACCM4 std. vs. MERRA



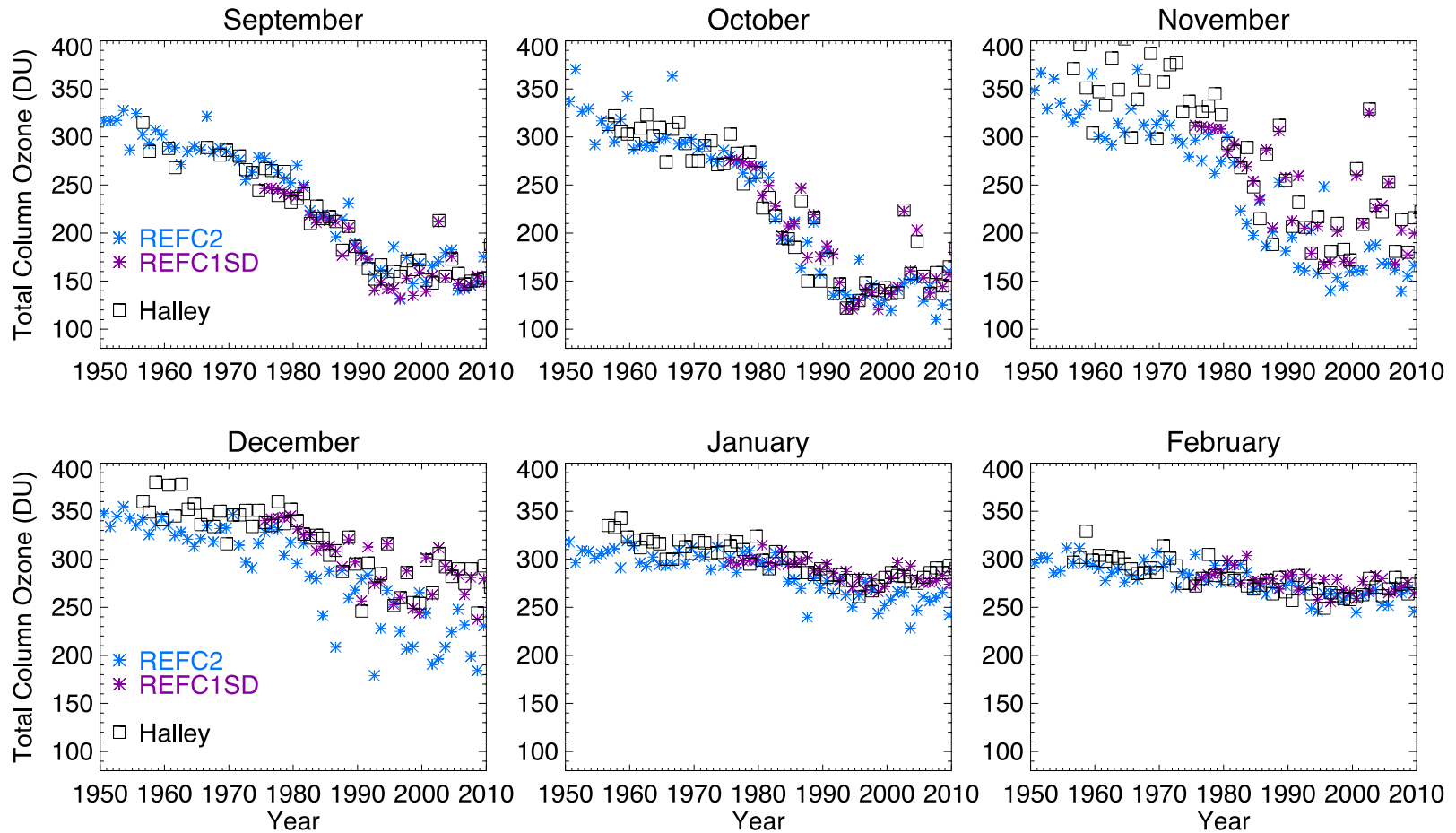
std. model is 5-10 K colder than MERRA from mid-September on

WACCM4 IGW vs. MERRA



model with IGW is within 5 K of MERRA through mid-October

ozone column in WACCM4 IGW

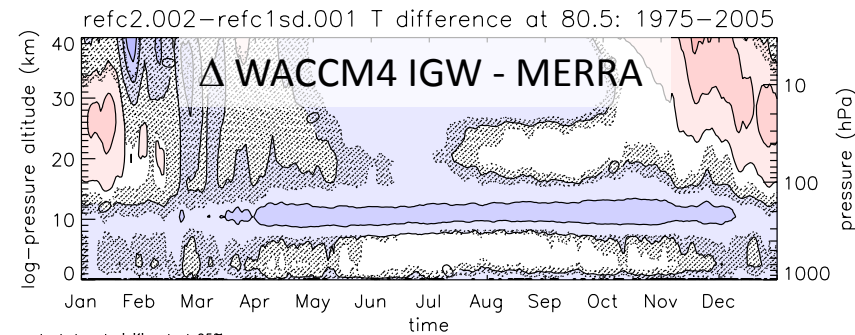
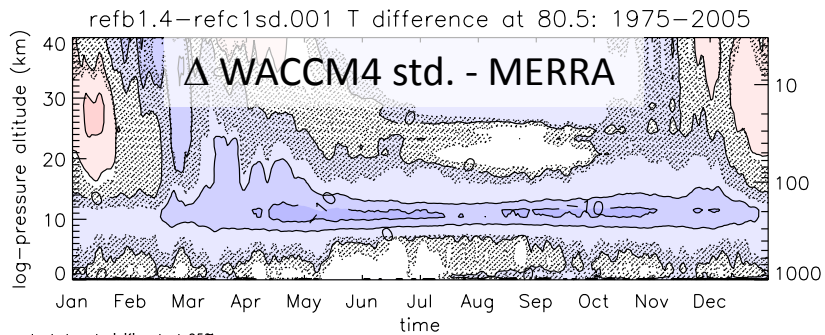
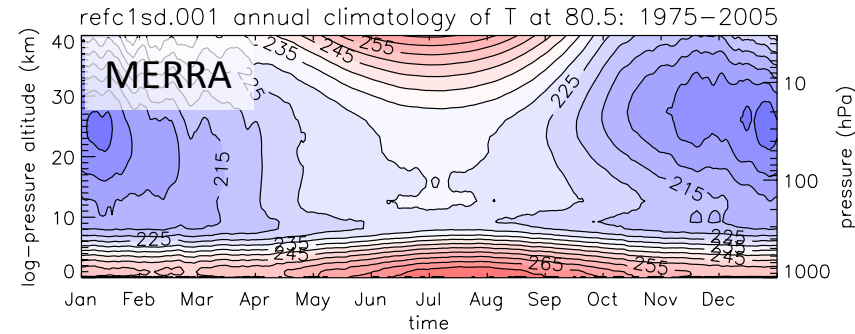
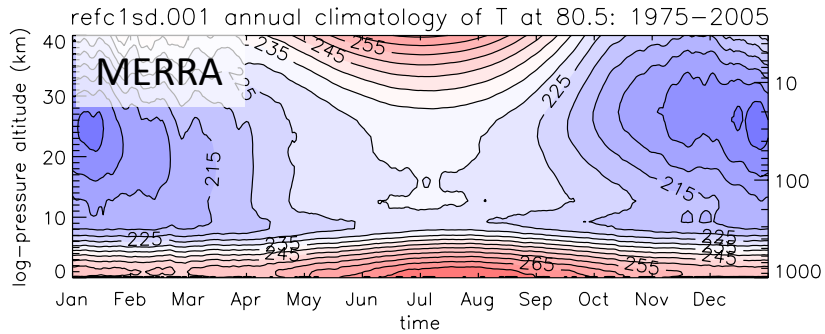
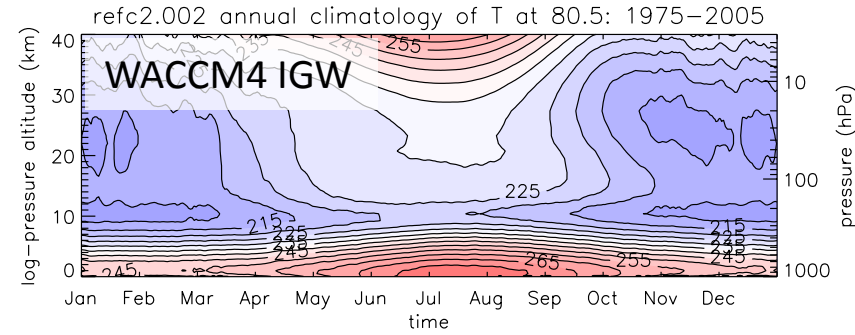
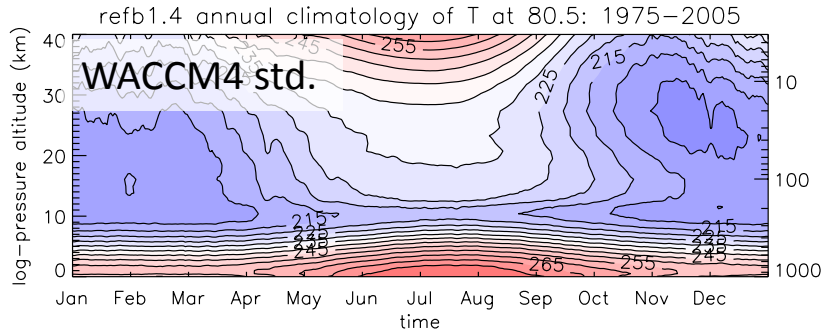


“REFC2”: WACCM4 IGW, free-running

“REFC1SD”: WACCM4 constrained with MERRA dynamics including IGW “solves” the low ozone problem

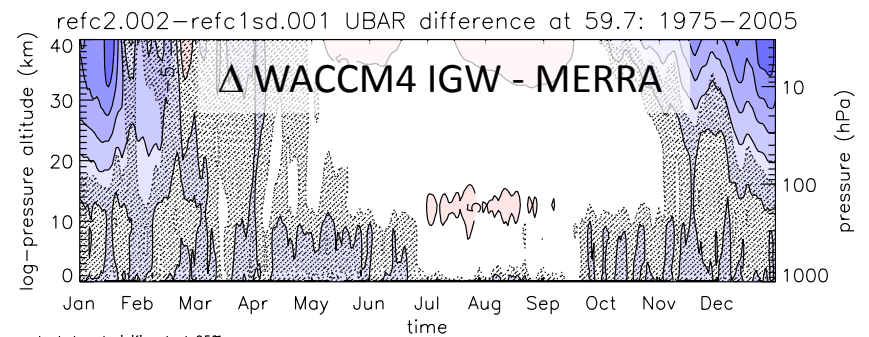
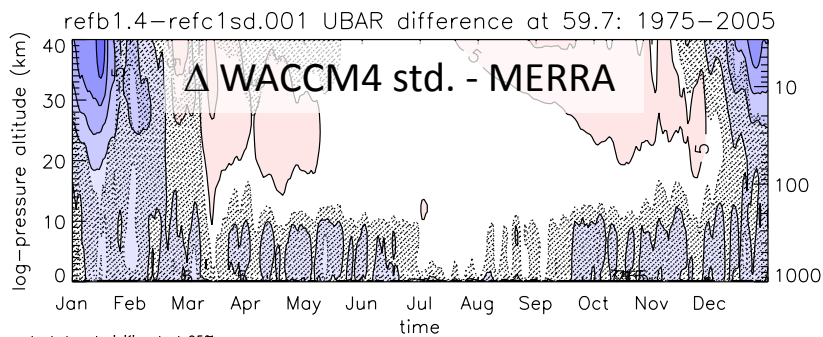
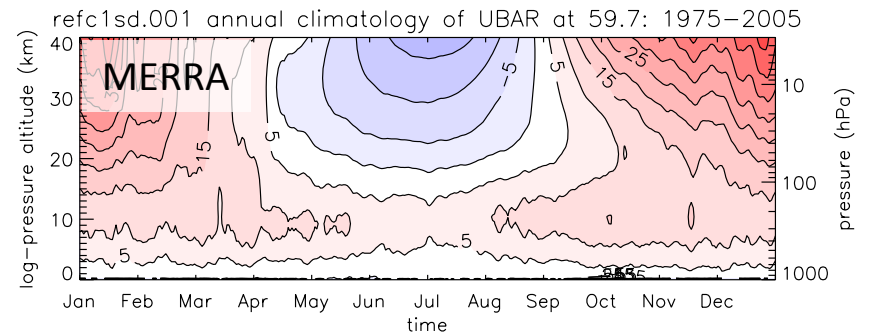
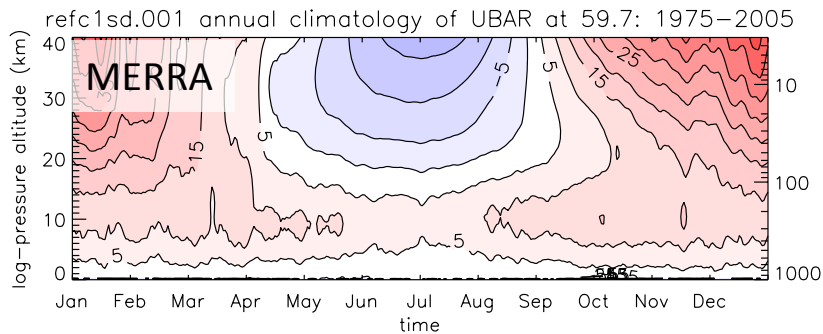
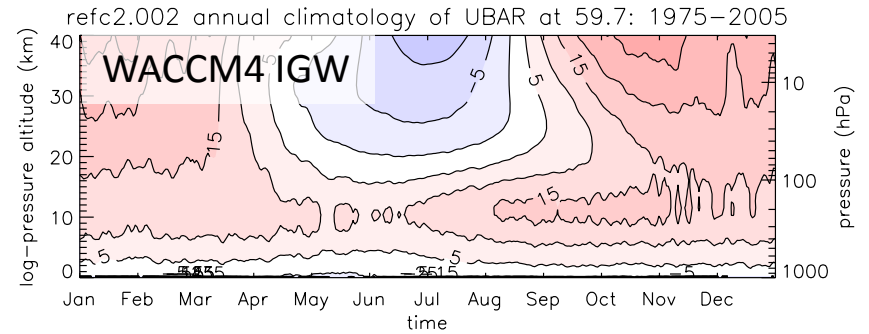
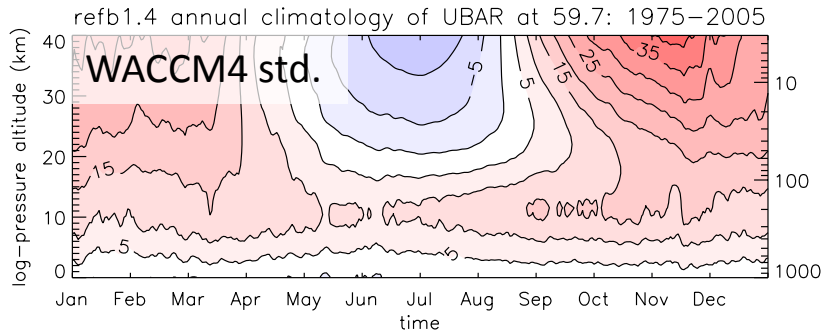
(except in December, because *final vortex breakdown is still too late*)

but what about T in NH winter?



differences in $\langle T \rangle$ (80°N) with respect to MERRA are small in both WACCM4 std. and WACCM4 IGW, although the latter is slightly warmer

similarly for the mean zonal wind at 60°N



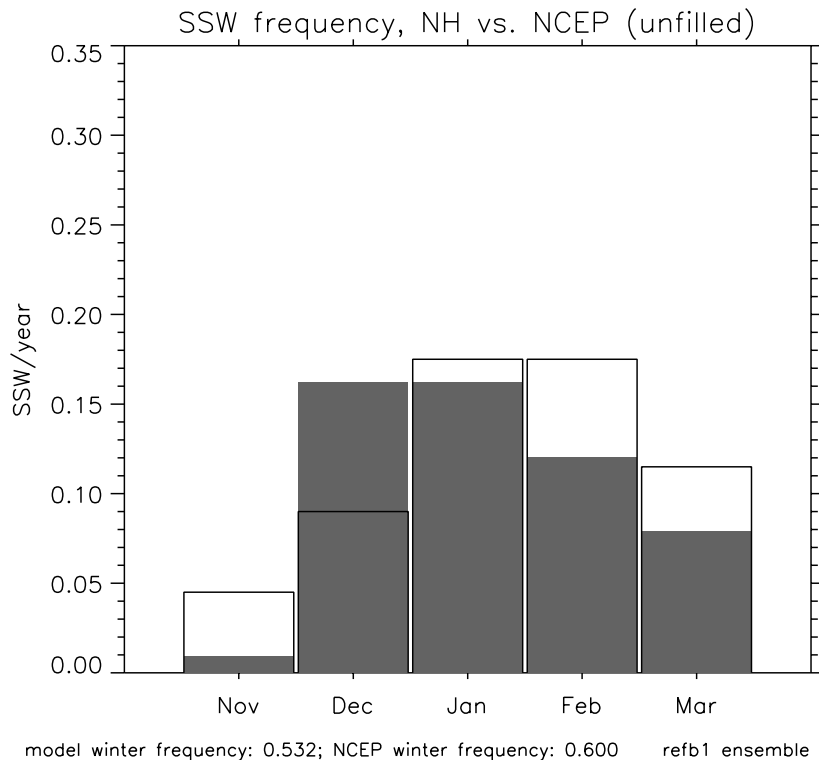
shaded: not significant at 95%

shaded: not significant at 95%

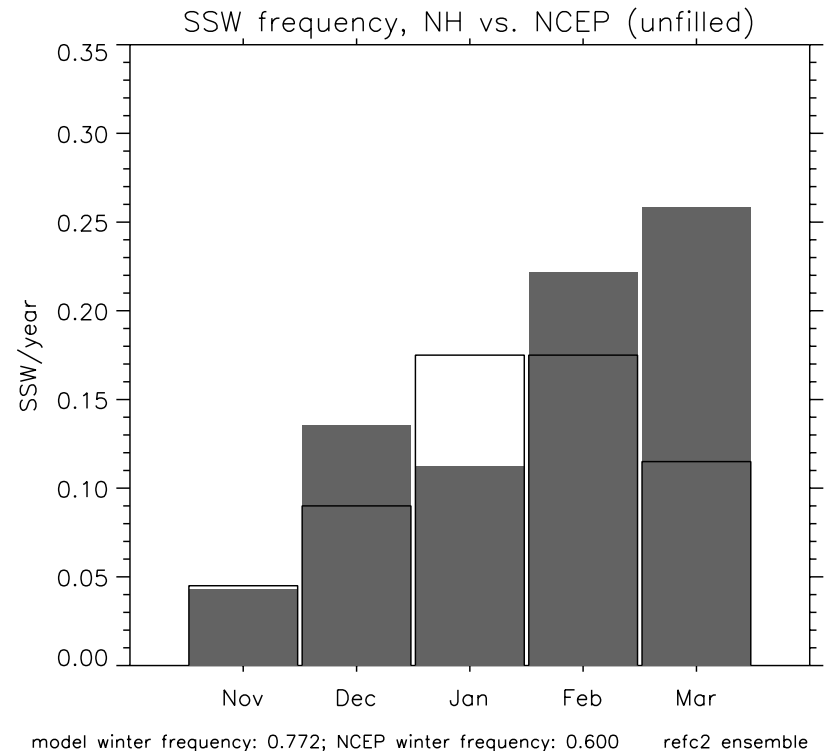
small differences overall, all of which suggests small impact of IGW in the NH. **However...**

SSW statistics: WACCM4 std. and IGW

WACCM4 std. 1975-2005
ensemble statistics

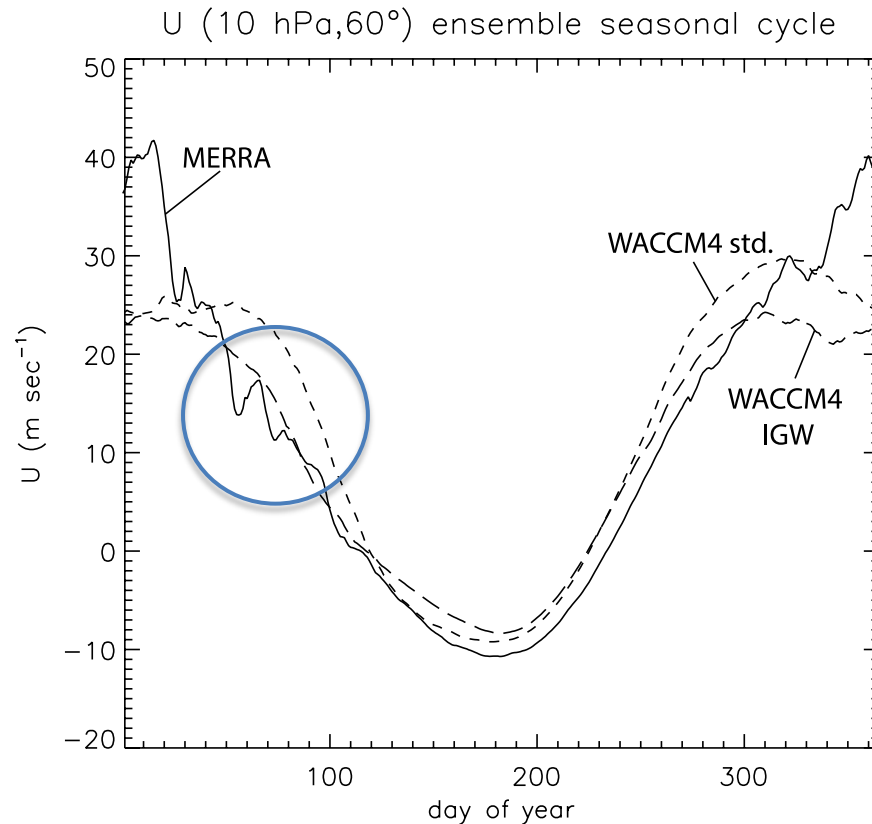


1975-2005 WACCM4 IGW
ensemble statistics



- WACCM4 IGW produces too many SSW late in the season (especially March)
- statistics are for a 3-member ensemble, so unlikely to arise by chance

a closer look at seasonal climatology: <U> (10 hPa, 60°N)



- $U(10 \text{ hPa}, 60^\circ\text{S})$ is stronger in MERRA than in either WACCM4 run in midwinter
- $U(10 \text{ hPa}, 60^\circ\text{S})$ is somewhat stronger overall in WACCM4 std. than in WACCM4 IGW and is noticeably stronger in NH spring (*but actually closer to MERRA*)

conclusions

- the addition of a second spectrum of parameterized gravity waves, in the IGW range, is physically reasonable
- it ameliorates the **SH** cold-pole problem sufficiently to allow realistic simulation of Antarctic ozone with the updated heterogeneous chemistry module
- it produces relatively minor changes in $\langle U \rangle$ and $\langle T \rangle$ in the **NH** compared to the standard version of WACCM4
- it agrees with MERRA data for the NH at least as well as the standard version
- **however**, it produces too many late-season SSW—this aspect of the simulation needs further study