

Westward Large-Scale Dynamical Behavior of
the Lower Thermosphere Simulated by
WACCMX-SD

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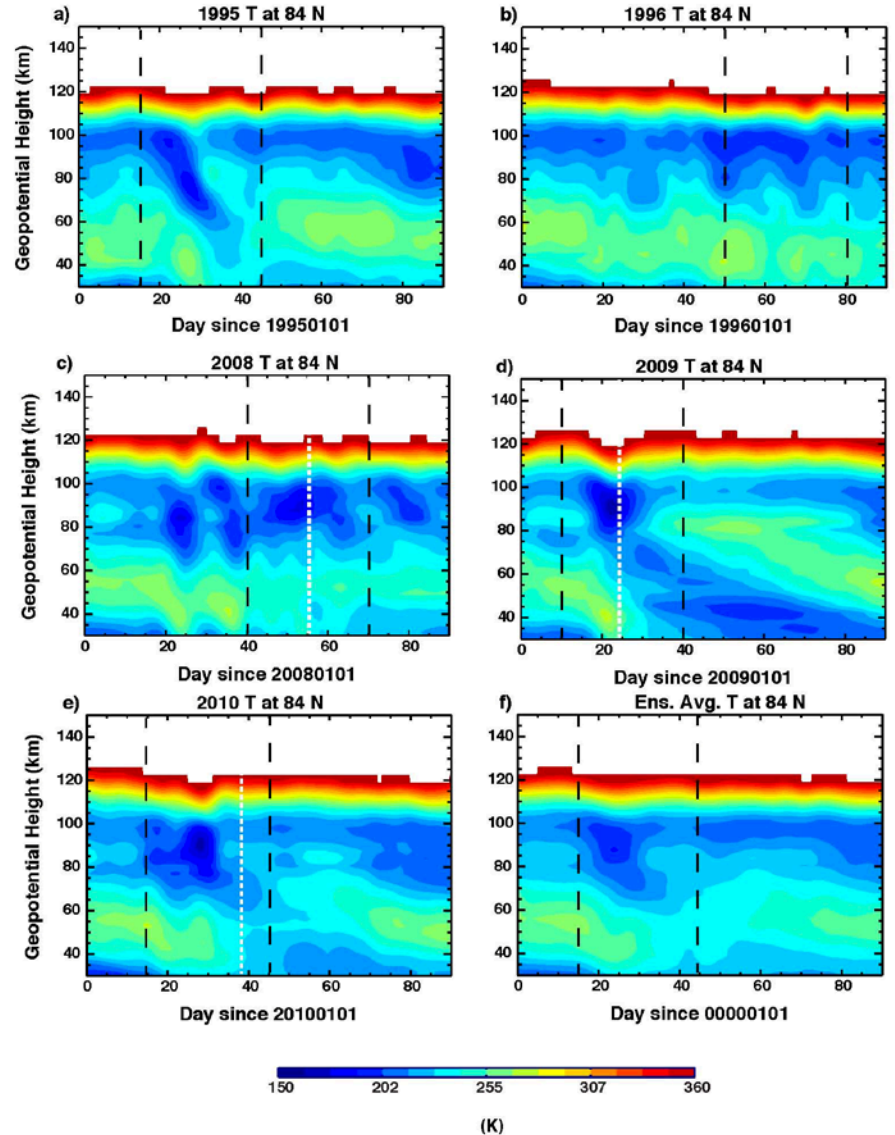
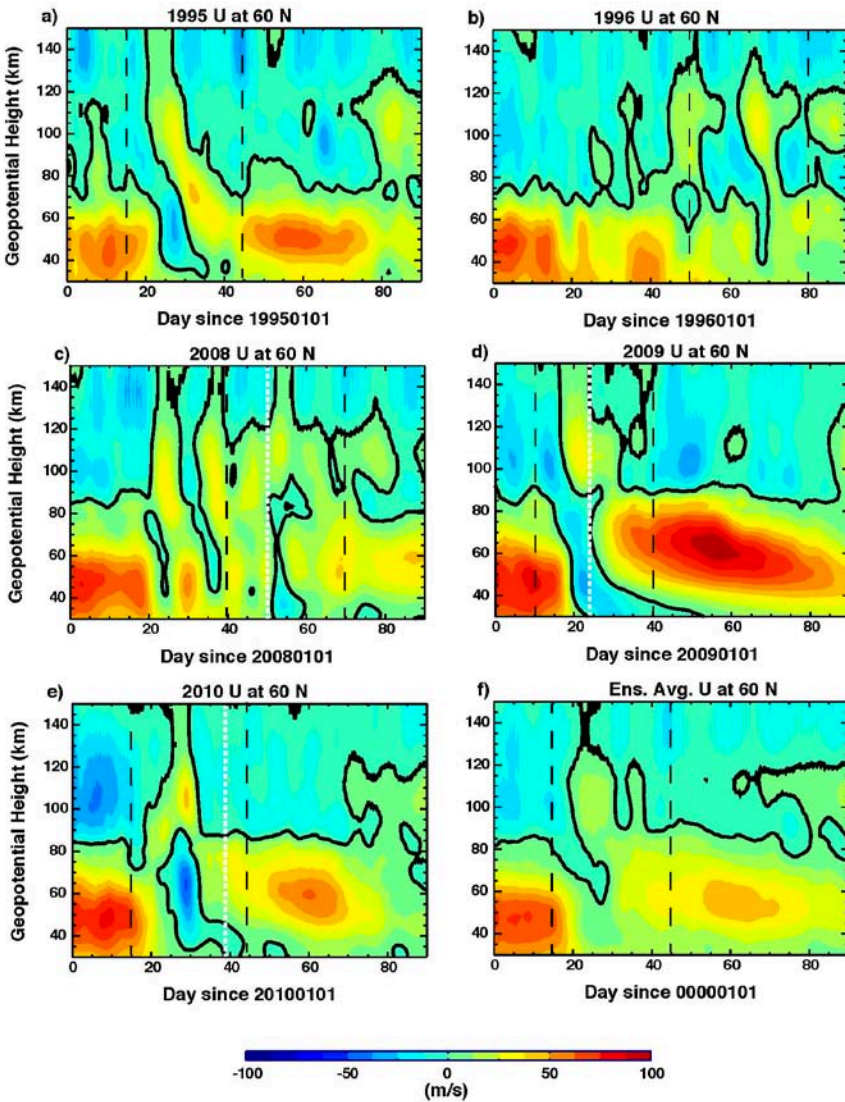
Motivation

- Understand and quantify the role of travelling planetary-scale disturbances in the lower thermosphere.
- Use SD configuration to generate realistic lower atmospheric conditions.
- Execute simulations of different boreal winters to build an ensemble of dynamical conditions
- Illustrate how traveling, planetary-scale waves - ubiquitous during boreal winter - can affect tidal variability.

Simulations

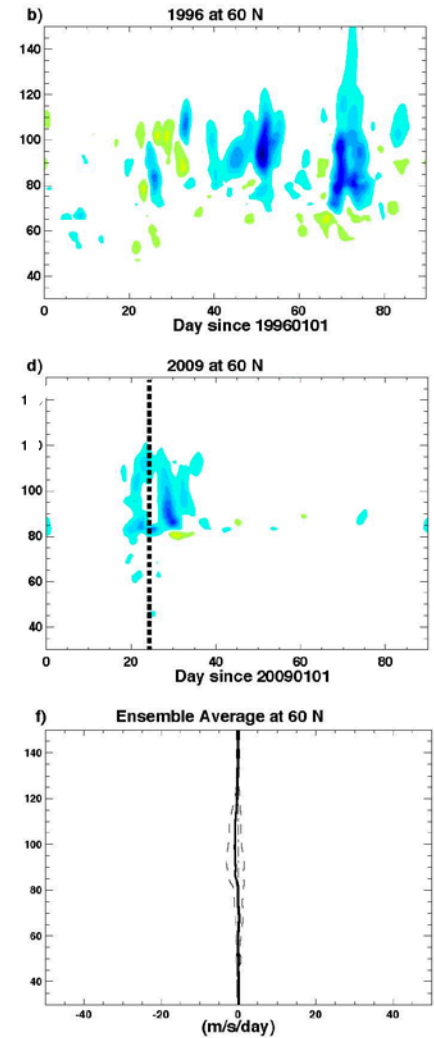
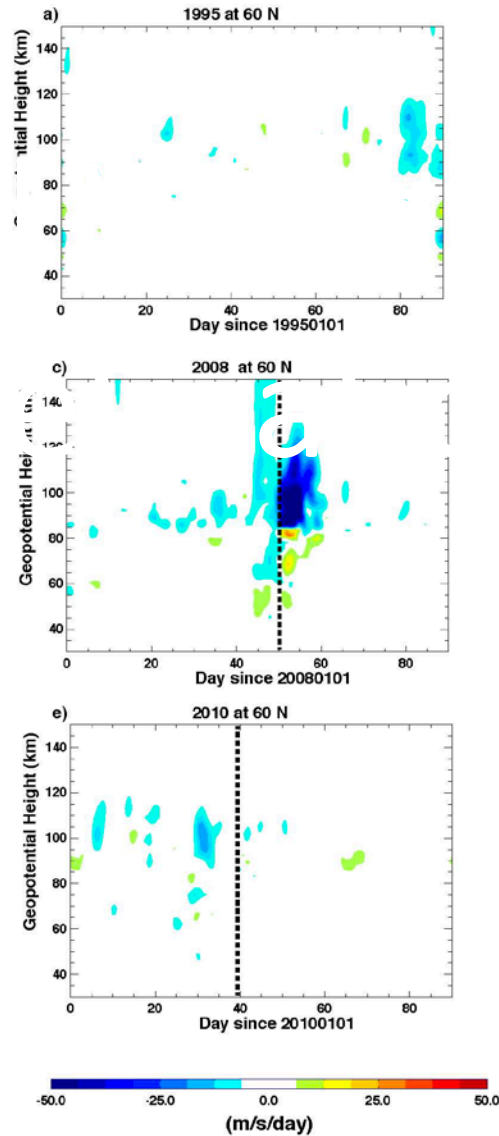
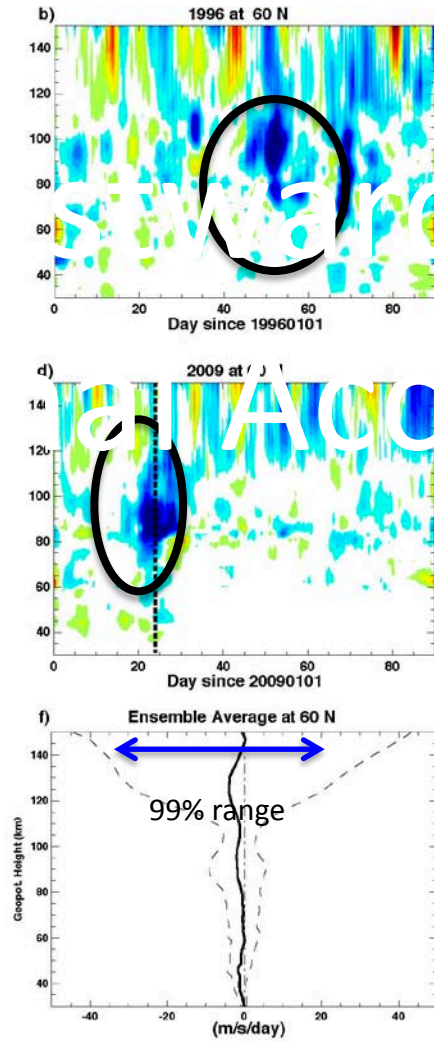
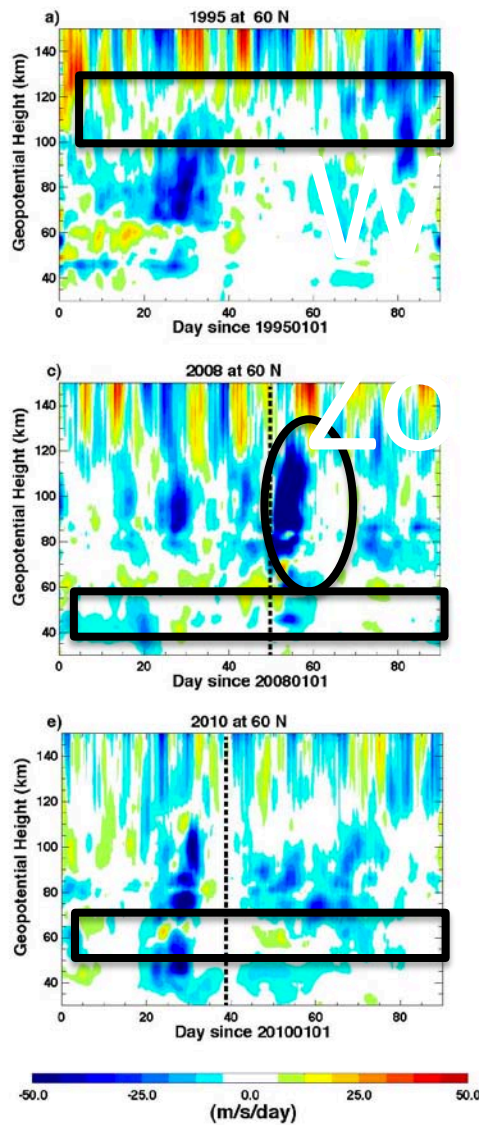
- Five boreal winters:
 - 1995: minor warming
 - 1996: no warming
 - 2008: wave-1 SSW
 - 2009: wave-2 SSW
 - 2010: wave-1 SSW
- These are all solar minimum winters
- SD configuration exploits Navy's and NASA's analysis (2008, 2009, 2010) or solely NASA's (1995 and 1996).

U at 60°N Climatology at 84°N



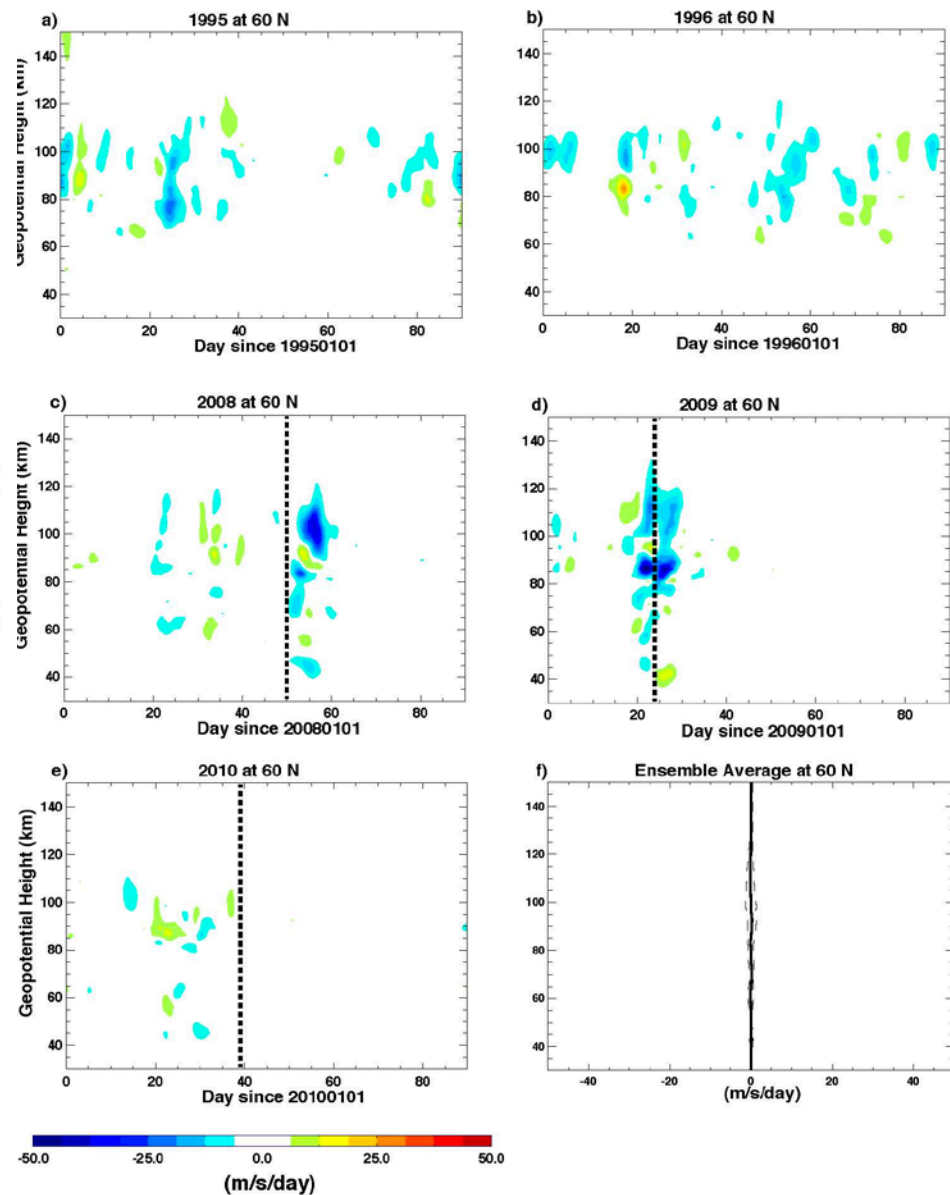
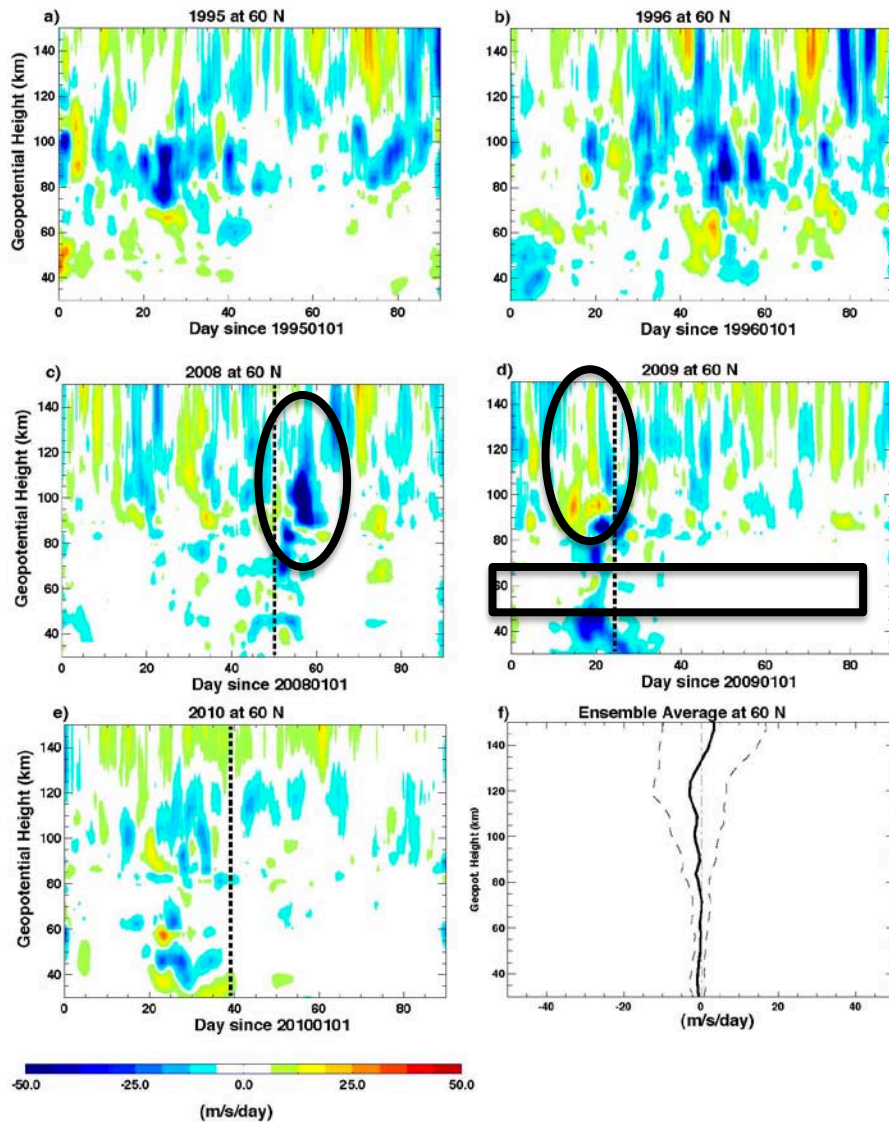
Westward wave-1 @ 60 N

Westward wave-1 ; $0.1 \leq \sigma \leq 0.5$ cpd @ 60 N



Westward wave-2 @ 60 N

Westward wave-2 ; $0.1 \leq \sigma \leq 0.5$ cpd
@ 60 N



Arrows: $(F^{(y)}, F^{(z)})$

Red shade: $EPD(0.1 \text{ cpd} \leq \sigma \leq 0.5 \text{ cpd}) \geq 0$

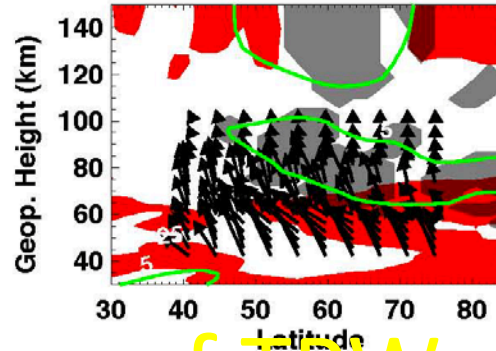
Gray shade: $(PV)_y \leq 0$

Green line: $U = 0$

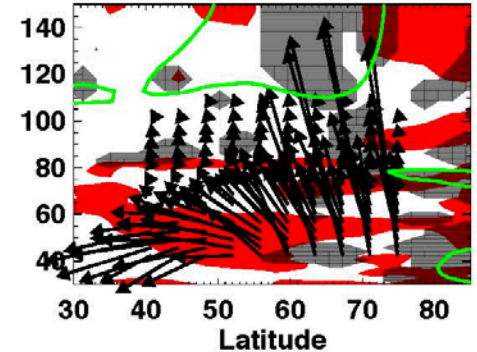


Source of TPWs: Dynamical Instability of the Background Zonal Flow?

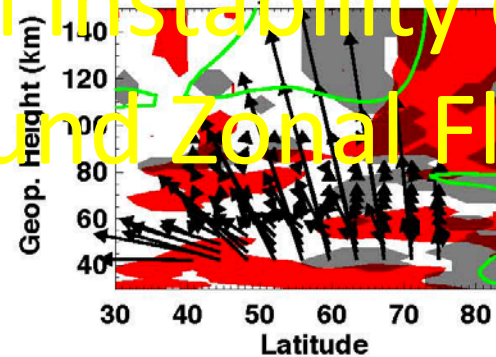
a) 1996: PVY and t1ZA
Days 50 - 80



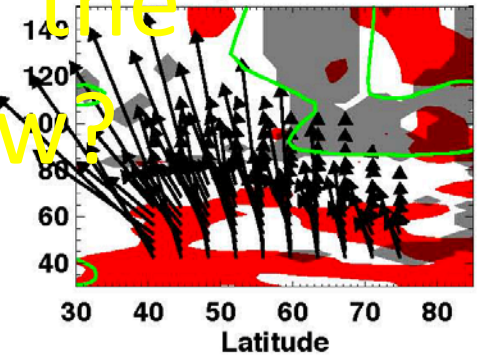
b) 2008: PVY and t1ZA
Days 40 - 70



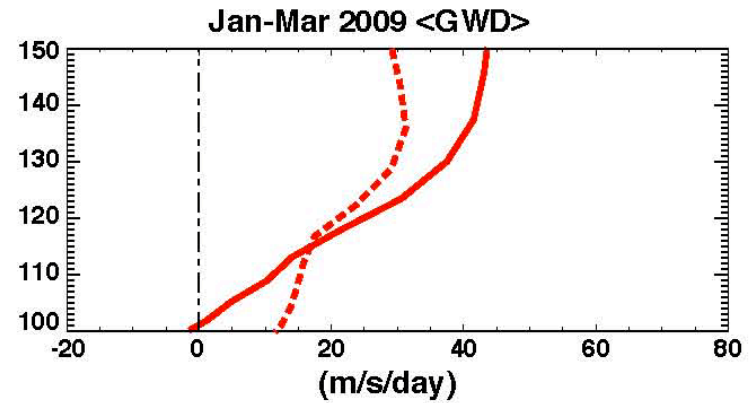
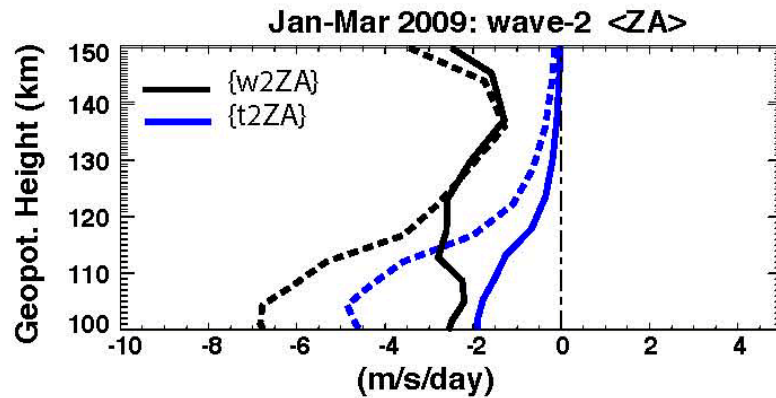
c) 2008: PVY and t2ZA
Days 40 - 70



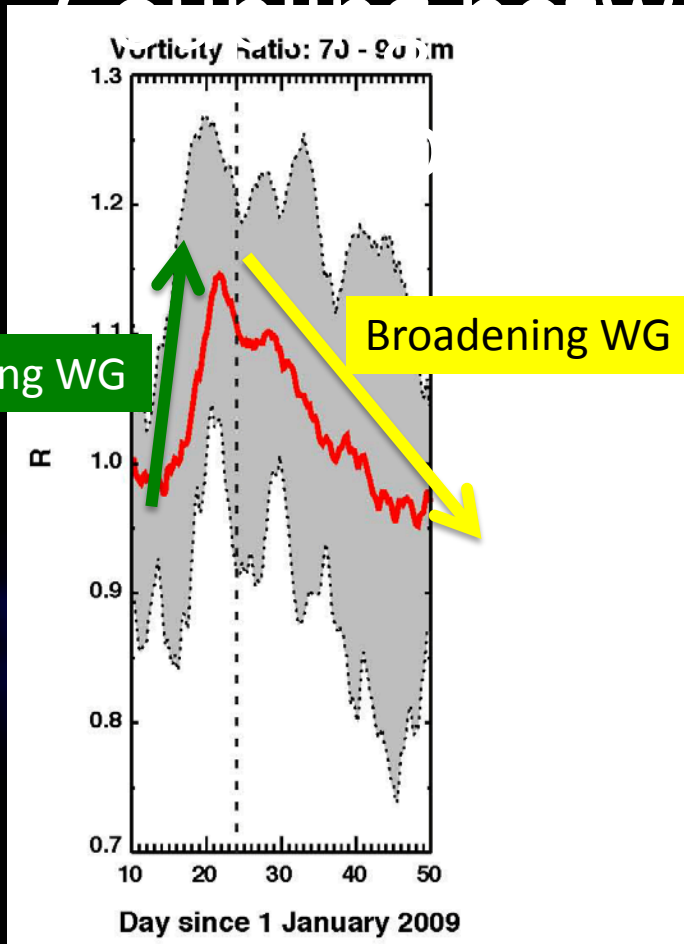
d) 2009: PVY and t2ZA
Days 10 - 40



Thermospheric Impacts



Coupling between high and low latitudes



McSignares (2002) suggests that the background zonal wind can change the vorticity in the tropics narrowing or broadening the wave guide for tropically trapped tides, and this mechanism can explain the intra-seasonal change of tidal amplitude.

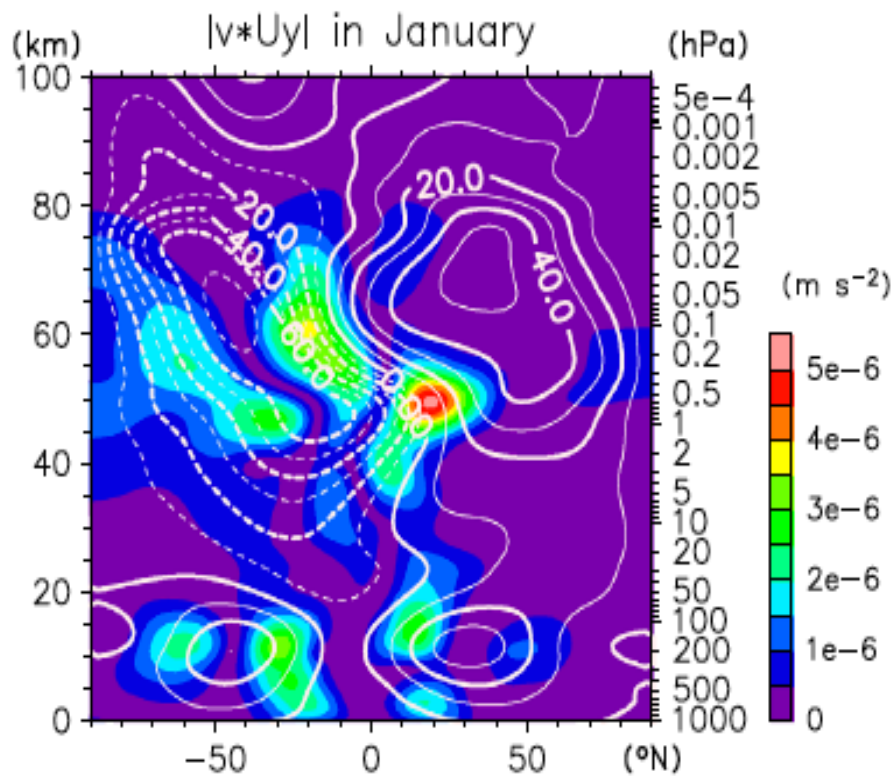
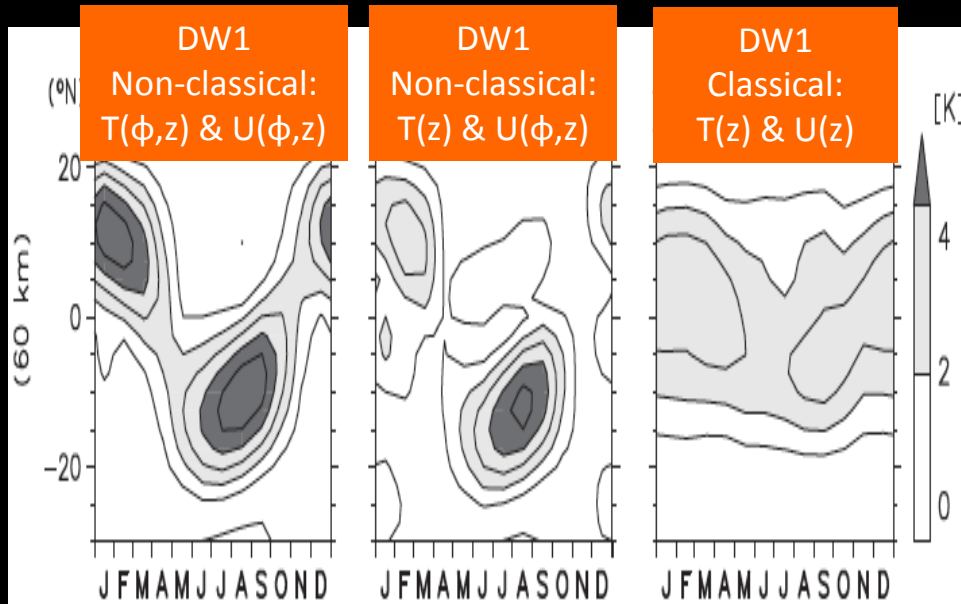
- Sassi et al. (2013) show that a SD-WACCM-X simulation reproduces the same mechanism on shorter time scales and is reflected in consistent changes of the tidal amplitudes.

$$R = (f - Uy) / f ; Uy = (Uy)_0 + \partial Uy$$

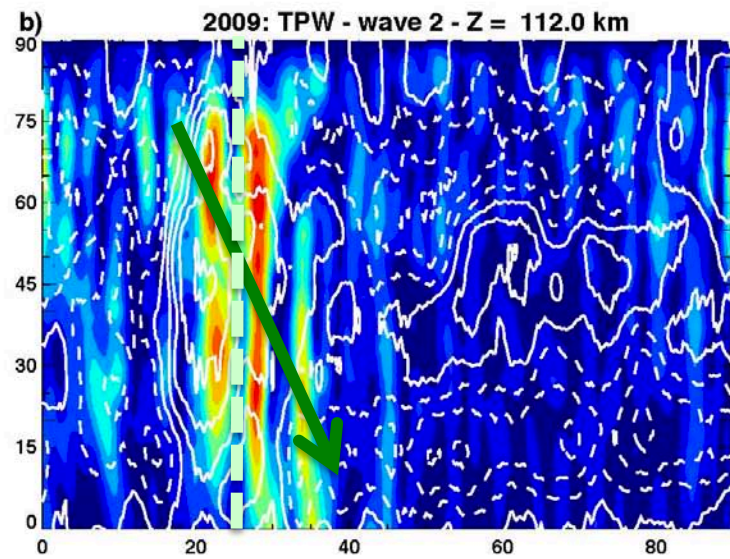
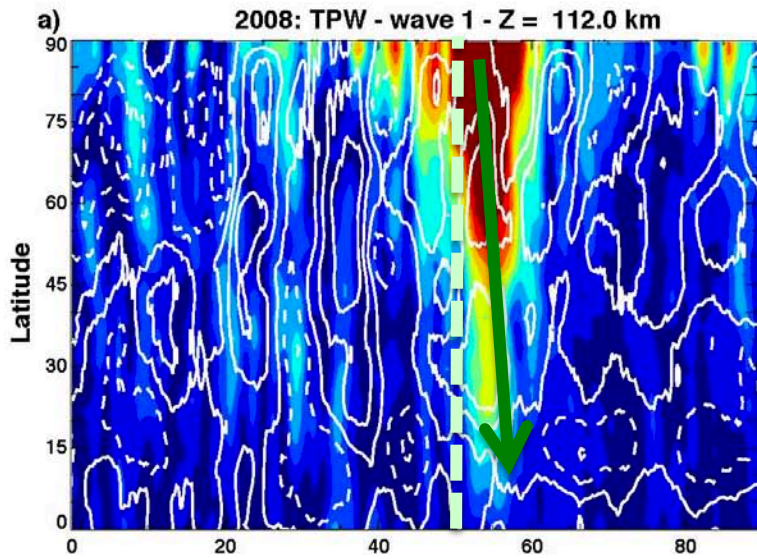
$$\partial Uy = 0 \rightarrow R_0$$

$$\partial Uy > 0 \rightarrow R < R_0 ; R \text{ decreases} \rightarrow \text{slower rotation / broad waveguide}$$

$$\partial Uy < 0 \rightarrow R > R_0 ; R \text{ increases} \rightarrow \text{faster rotation / narrow waveguide}$$



- *Sakazaki et al.* (2013) use a perturbation technique to find the solution to the horizontal structure tidal equation in the presence of a non-uniform (non-classical) background state: they show that non-uniform winds account for the difference from the classical solution.
- A term proportional to the wind shear behaves as a mechanical source of momentum in the upper atmosphere and modifies the seasonal evolution of the classical (1,2) Hough mode.



0.0 4.2 8.3 12.5 16.7 20.8 25.0
(m/s)

Color shade: RMS amplitude of v-wind
between $0.1 \leq \sigma \leq 0.5$ cpd westward

White contours: U

Conclusions

- TPWs emerge in the thermosphere from the upper mesosphere; their zonal acceleration can be prominent both at times of SSWs or during dynamically quiet times.
- The direct impact of TPWs to the momentum budget of the thermosphere is secondary compared to the parameterized GWD.
- Instability of the background flow as a source of TPWs remains uncertain, at least in a time-mean sense during the most disturbed periods.
- TPWs are associated with propagation of wave activity from mid-latitudes toward the tropics. Following *Sakazaki et al. (2013)* and *Sassi et al. (2013)*, the associated perturbation of the zonal flow are qualitatively consistent with tidal variability.
- It remains to be explained exactly how TPWs interact with the tropical circulation, especially which mechanisms is more important to modify the tidal amplitudes: mechanical forcing (*Sakazaki*) or width of the waveguide (*McLandress*).

Acknowledgements:

Support for this work was provided by NASA/LWS grant NNH12AT21I and ONR 6.1 Base funding

THANK YOU!