Update on CAM5 Microphysical Improvements Regarding Ice Particle Mass- and Area-dimensional Expressions

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#### Dependence of log(m) on log(D) is not linear



Data was grouped into 3 temperature-categories for synoptic and anvil cirrus, with a m-D and A-D curve fit for each category



Comparisons with other field data and other studies





#### Testing for self-consistency and physical realism



# 2D-S probe (and other optical probes) cannot accurately measure ice particle area ratios for D < 200 $\mu$ m.

Out-of-focus problem demonstrated for glass spheres having diameter of 100  $\mu$ m, but image size varies depending on proximity to depth-of-field.



Out-of-focus problem is greatly reduced for 200 µm glass spheres



**Proposed solution:** Use Cloud Particle Imager (CPI) data when ice particle length D < 200  $\mu$ m. The CPI and 2D-S probes were co-located during SPARTICUS cirrus cloud sampling. The CPI spatial resolution is 2.3  $\mu$ m and there is no out-of-focus problem. Ice particle area & area ratios are very accurate.

#### Calculation of ice particle mass for $D < 200 \mu m$ :

For each ice particle, define an ice sphere having the same projected area as the ice particle and calculate the mass of this ice sphere.

**Expected error:** For D < 200  $\mu$ m, ice particles are generally compact, quasi-spherical or blocky in shape, making the above mass estimate a good approximation. For aspect ratios < 3, errors should not exceed ~ 30%.

## 2<sup>nd</sup> Problem: The dimension the 2D-S reports is not always the maximum dimension. Solution: Change the data processing algorithm New Algorithm (M7)

#### M1 Method

 The M1 method uses the length scale L1, which is the maximum particle dimension along the direction of travel. All particles are included in area and number concentration calculations.

#### M7 Method

• The M7 method uses the length scale MaxLength, which is the maximum particle dimension of the projected image. Only particles that are 'all-in' are included (particles intercepting the edge of the 2DS field of view are excluded from area and number concentration calculations).

# Conclusions based on one synoptic cirrus and one anvil cirrus flight:

 New algorithm has a larger effect on particle size and total projected area for anvil cirrus than for synoptic cirrus, because particles are often 1mm or larger in anvil cirrus, and are therefore more likely to intercept the edge of the 2D-S field of view.

Using M1 instead of M7 results in:

- < 5% bias in sample area for synoptic cirrus
- < 13% bias in sample area for anvil cirrus

#### **Summary and Conclusions**

1. When considering all sizes, ice particle mass and projected area are better estimated by 2<sup>nd</sup> order polynomial fits than by power laws.

2. Method's accuracy is evident through comparisons with two recent studies and ground field measurements.

3. Uncertainties in ice particle area and mass are considerable for D < 200  $\mu$ m, but these uncertainties will be greatly reduced by using CPI imagary.

4. The standard 2D-S data processing estimates an ice particle's maximum dimension D, but the new SPEC algorithm will actually measure D. Also, only particles completely within the 2D-S field of view are counted and analyzed.

### Generating m-D and A-D Relationships Using the 2D-S Probe

The 2D-S probe measures the size resolved concentrations of ice particle number, projected area and estimated mass from 10 to 1280  $\mu$ m, using the Baker-Lawson mass-area power law to estimate particle mass. The mean ice particle mass for each size-bin is given as

m = M/N

where M and N are mass and number concentration for a given size-bin. Relating m to the midpoint size of each bin, the dependence of m on D is revealed. The dependence of mean ice particle projected area on D is found in a similar manner.

A mean PSD was produced for each 5°C T-interval that reduces scatter while focusing the curve fit. Temperature categories were identified where m-D & A-D curves were similar. A single m-D or A-D curve fit was produced per T-category.

The main objective here was to evaluate the dependence of m and A on D over a much greater range of ice particle size than done before (especially the size range between 10 and 100  $\mu$ m).





$$m = \alpha D^{\beta}$$
 For a given D, we can  

$$A = \gamma D^{\delta}$$
 obtain these power laws

$$\ln m = a_0 + a_1 \ln D + a_2 (\ln D)^2$$

$$d(\ln m)/d(\ln D) = \beta = a_1 + 2a_2 \ln D$$

 $\alpha = \exp[a_0 + a_1 \ln D + a_2 (\ln D)^2] / D^{\beta}$ 

#### $\beta$ uncertainty << $\alpha$ uncertainty

#### Synoptic cirrus, exponent of m-D power law

	Ice particle size (µm)				
	50	150	500	1500	4500
Temperature Range	Power β				
- <b>40 &lt; T ≤</b> -10°C	2.632	2.368	2.079	1.814	1.550
-55 < T ≤ -40°C	2.564	2.280	1.968	1.684	1.399
- <b>65 &lt; T ≤</b> -55°C	2.477	2.057	1.597		
Mean β	2.558	2.235	1.881	1.749	1.475
Standard deviation of $\beta$	0.064	0.131	0.206		

Mean uncertainty for  $\beta$  as 100  $\times$   $\sigma/$  mean value : 6.74%

#### **Application to Cloud Modeling**

$$N(D) = N_{o} D^{v} \exp(-\lambda D)$$
  
$$\lambda = \left(\frac{\alpha \Gamma(\beta + \nu + 1) N}{\Gamma(\nu + 1) IWC}\right)^{1/\beta}$$
Used in CAM5

To a good approximation,  $\lambda$  is obtained by evaluating  $\alpha \& \beta$  at D = 500 µm. Then estimate the D for the cloud property or process of interest by evaluating  $\beta$  and  $\delta$  at D = 500 µm:

$$\begin{split} D_{N} &= (\nu + 0.67)/\lambda \\ D_{a} &= (\delta + \nu + 0.67)/\lambda \\ D_{m} &= (\beta + \nu + 0.67)/\lambda \\ D_{Z} &= (2\beta + \nu + 0.67)/\lambda \end{split}$$

Median number conc. dimension Median area dimension Median mass dimension Median radar reflectivity dimension

Then calculate  $\alpha$ ,  $\beta$ ,  $\gamma$  and  $\delta$  for the selected D value.











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#### SPARTICUS PSD SAMPLING FOR SELECTED SYNOPTIC AND ANVIL CIRRUS CLOUDS



#### Example of A-D Curve Fit

