

# Use of CESM to quantify aerosol forcing from the Eyjafjallajökull volcanic eruptions

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Material from *Flanner et al* (2014), JGR, doi:1002/2014JD021977

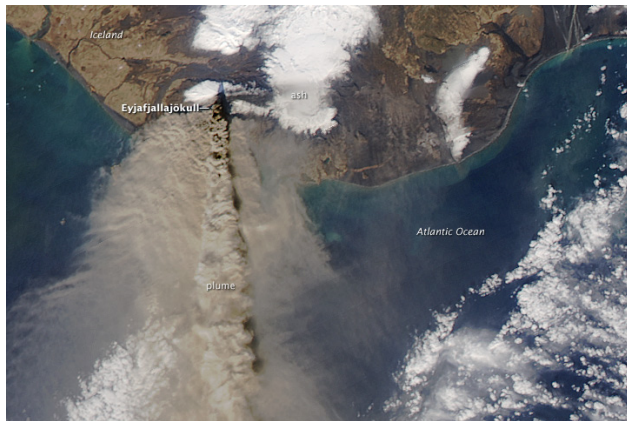


JPL



# Eyjafjallajökull eruption

MODIS, NASA Earth Observatory Image of the Day



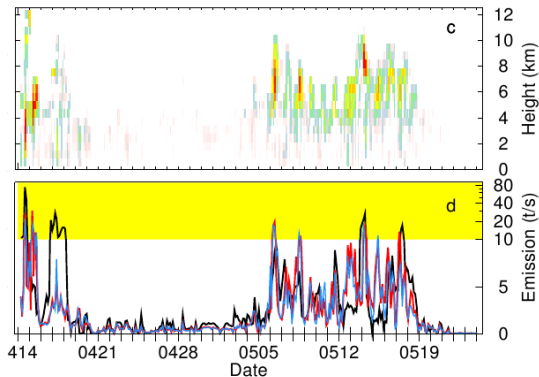
- Two eruption episodes in April and May 2010
- Disruption of air traffic. Small but unique impact on climate.

# Aerosol climate impacts

- Atmospheric ash forcing:
  - longwave: +
  - shortwave: ?
- Atmospheric sulfate forcing:
  - longwave: +
  - shortwave: -
- Deposition of ash to snow and sea-ice:
  - shortwave: +
- Insulation of snow: -
- Aerosol-cloud indirect effects:
  - shortwave: - (?)
  - longwave: + (?)



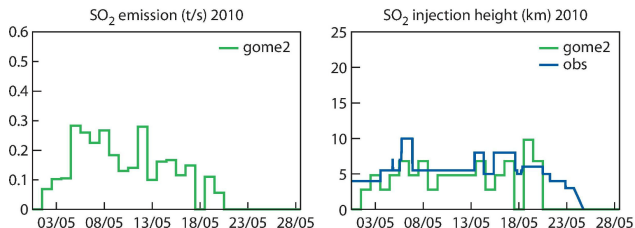
**Figure:** 10 cm thick ash layer overlying snow in September 2010. Courtesy of Steve Warren.

Ash emissions (*Stohl et al., 2011*)

- Total injected tephra mass:
  - $8.3 \pm 4.2$  Tg as “fine” ash ( $2.8 - 28 \mu\text{m}$  diameter)
  - $\sim 12$  Tg in  $2.5 - 250 \mu\text{m}$  size range
- Global annual black carbon emissions:  $\sim 8$  Tg

SO<sub>2</sub> emissions

- Estimates derived from OMI, SCIAMACHY, GOME-2, and ground radar (*Flemming and Inness, 2013*):



- Total SO<sub>2</sub> emissions: 0.25 (0.13 – 0.43) Tg, also from *Heard et al. (2012)*
- Only ~ 3% of SO<sub>2</sub> emissions occurred during April event (not shown)
- Very little injection of SO<sub>2</sub> into stratosphere

# Aerosol radiative forcing calculations

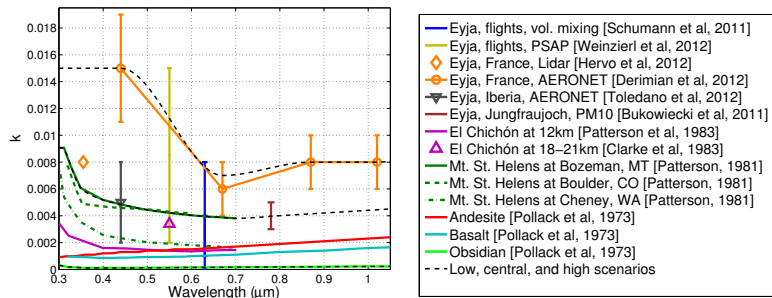
- CAM, CICE, and CLM employed in different capacities
- Modified CAM4 to accommodate 4 ash tracers and volcanic SO<sub>2</sub>/SO<sub>4</sub> tracers with new optical properties
- Vertically-resolved daily SO<sub>2</sub> emission fluxes from *Flemming and Inness*, (2013), oxidation to SO<sub>4</sub> simulated with CAM4-BAM with prescribed 2010 SSTs/sea-ice
- Daily 3-D ash fields from *Stohl et al.* (2011), 25 size bins, re-partitioned into 4 size bins, prescribed in CAM
- Atmospheric RF calculations using RRTMG with prescribed ash and SO<sub>4</sub> fields in CAM
- Added 4 ash particle species to CLM/SNICAR and CICE/Delta-Eddington (pre-existing: 2 BC species, 4 dust).
  - Particle size ranges partitioned to have roughly equal surface area
  - $r < 0.56 \mu\text{m}$ ,  $0.56 < r < 1.0 \mu\text{m}$ ,  $1.0 < r < 2.5 \mu\text{m}$ ,  $r > 2.5 \mu\text{m}$
- Daily ash deposition fluxes from *Stohl et al.*, (2011) prescribed in CLM4 and CICE4 simulations with 2010 forcing data

# Atmospheric ash (*Stohl et al., 2011*)

- Ash transport, wet+dry deposition simulated with the Lagrangian transport model FLEXPART, met fields from ECMWF and NCEP
- Forward dispersion modeling and satellite observations combined with inversion scheme to determine time-resolved ash emissions

# Ash optical properties

- Uncertainty in imaginary component of ash refractive index drives large uncertainty in forcing
- Low, central, and high absorptivity estimates derived from aircraft/PSAP measurements, sun photometer inversions, Lidar inversions, and measurements of previous events



- Mie optical properties weighted into RRTMG SW and LW spectral bands, and provided as supplementary data



# Optical properties

- Ash particles are often highly non-spherical

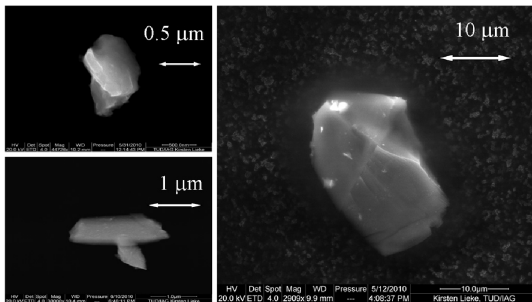
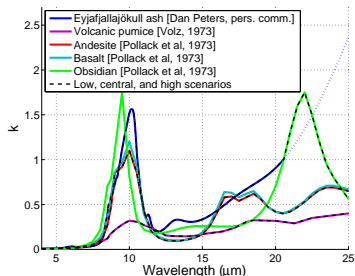
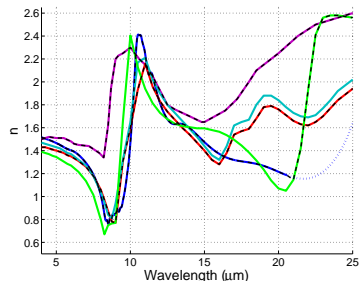


Figure: Schumann et al (2011), ACP

- Properties for equal-mass non-spherical particles simulated with T-Matrix code [*Mishchenko and Travis, 1998*]. MAC of Chebyshev particles, oblate/prolate spheroids, spheres differ by **16% at most**

# Optical properties

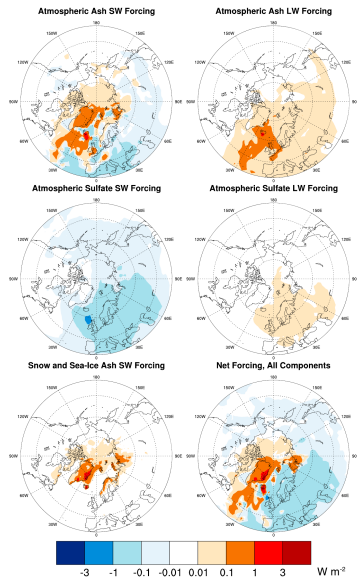
- Variability in ash refractive index in longwave spectrum:



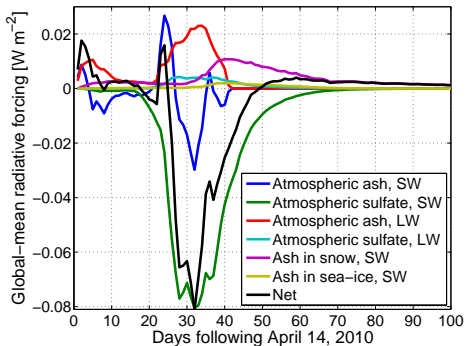
- Sulfate optical properties derived for three size distributions:  
 $r_e = 0.17, 0.27, 0.43 \mu\text{m}$  (*Rasch et al.*, 2008; *O'Dowd et al.*, 2011)

# Aerosol forcing components: Daily animations

# Aerosol forcing components: 2-month means



# Aerosol forcing components: Timeseries of global means



- Atmospheric ash SW forcing is noisy because of variable plume/cloud/cryosphere co-location and short residence time
- Ash LW forcing is substantial because particles are large
- Ash-in-snow forcing persists for months
- Negative SW forcing from sulfate dominates (in May)

## Aerosol forcing components: Means

Table: Global annual-mean radiative forcings [ $\text{mW m}^{-2}$ ]

Instantaneous top-of-atmosphere forcings								
	Ash SW	Ash LW	Ash in snow	Ash in sea-ice	Sulfate SW	Sulfate LW	Net	Net Effective
Variable optical properties, central emissions								
Low	-4.1	0.7	0.0	0.0	-4.1	0.2	-7.3	-7.2
Central	-0.3	1.1	0.8	0.1	-3.8	0.2	-1.9	-0.5
High	+2.7	1.2	1.3	0.2	-3.0	0.2	+2.8	+4.9
Variable emissions, central optical properties								
Low	-0.1	0.5	0.4	0.1	-6.1	0.4	-4.9	-4.3
Central	-0.3	1.1	0.8	0.1	-3.8	0.2	-1.9	-0.5
High	-0.6	2.1	1.5	0.3	-2.1	0.1	+1.2	+4.5

- Central estimates yield weakly negative forcing
- High ash absorption assumption produces *positive* net forcing
- Forcing sign of atmospheric ash component is uncertain
- Uncertainty in emissions of ash and  $\text{SO}_2$  are both  $\sim 2\times$

# Uncertainty due to clouds

- Cloud variability in different ensemble members drives large variation in atmospheric ash SW forcing, but has little impact on the other forcing terms

**Table:** Global annual-mean instantaneous top-of-atmosphere radiative forcings from different ensemble members [ $\text{mW m}^{-2}$ ]

Ensemble Member	Ash SW	Ash LW	Sulfate SW	Sulfate LW
E1	-0.29	1.06	-3.83	0.24
E2	-0.59	1.07	-3.73	0.24
E3	-0.45	1.03	-3.56	0.22
E4	-0.20	1.07	-3.62	0.22
E5	-0.28	1.05	-3.57	0.22
Mean	-0.36	1.06	-3.66	0.23

# Conclusions

- CESM is a useful tool for calculating RF of various volcanic aerosol components
- Net aerosol forcing from Eyjafjallajökull was nearly climate neutral, marking an unusual volcanic event in present climate
  - Negative sulfate forcing slightly exceeded positive ash forcing
  - Ash-in-snow forcing persisted longer than atmospheric forcing, but operated over a smaller spatial domain
  - Ash longwave forcing is non-negligible
- Ash absorptivity and emissions are largest sources of uncertainty. Ash/cloud covariance is large source of uncertainty for atmospheric ash SW forcing
- Beyond RF: Did large positive forcing over Arctic and Greenland enhance summer melt in 2010?
- Did latitudinal gradient in forcing alter atmospheric dynamics in a meaningful way (e.g., weakened westerlies)?