Investigating the Tropical and South Atlantic variability during the last Millenium Volcanism and air-sea interaction processes in the equatorial Atlantic

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Outline

- Introduction
- Equatorial Atlantic variability
- Data and Methods
- Results
- Discussion -Conclusions



Pinatubo eruption, June 12, 1991



Volcanism impacts on climate

- Radiative cooling of the surface and warming of the stratosphere.
- Cooling of SSTs (ocean)
- SW radiation deficit on the surface
 - weakens the evaporation
 - Decrease of monsoonal precipitation (*Iles et al., 2013; Iles and Hegerl, 2014*)



The Atlantic Equatorial Mode (AEM)

- Interannual mode of variability: tongue-shaped spatial pattern in tropical Atlantic SST that peaks during JJA.
- ► The AEM mechanisms is similar to ENSO [Bjerknes feedback]



JJA climatology: SST ($^{\circ}$ C) in colors and winds (m/s) in arrows (both from Dee et al., 2011), PPT (mm/month)

in green lines (Adler et al., 2003).





Positive feedback loop reinforces initial anomaly



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Data and Methods

- ► CCSM4 LM run (Landrum et al., 2013)
- 1850 control
- ► AEM = (SST_{EAST} SST_{WEST})/σ which we regress with:
 ► SST
 - Zonal wind stress
 - Upwelling on the African coast (depth of the 20°C isotherm).



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Data and Methods (cont)

- Composite analysis
 - 1. Select the years of volcanic eruptions threshold of -2.5 W/m2 to select each eruption event
 - 2. Volcanic events of interest: the year before (-1), the eruption year (0), and four years after the eruption (+1 to +4).





Was the Bjerknes feedback active for the Atlantic?

Regression results



Regression of the AEM reconstructed index on (a and d) SST (°C); (b and e) t_X (N/m²); and (c and f) 20°C isotherm depth anomalies (Z₂₀).



Composite analysis



NCAR-CCSM4 volcanic composites. From (a) to (c): year (-1); from (d) to (f): year (0); from (g) to (i): year (+1); from (j) to (l): year (+4). First column: SST anomalies (°C); second column: t_{χ} anomalies (N/m²); third column: 20°C isotherm depth anomalies (m).



Sub-surface warming of the Atlantic



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Volcanism impacts on rainfall in SA Regression of the AEM-index onto PPT

 Stronger relationship between the AEM index and ppt-anomalies in the PI experiment



NCAR-CCSM4 AEM reconstructed index regressed onto precipitation anomalies (mm⁻¹/month). Panel (a) refers to **PI** results, and panel (b) refers to **LM** results. (p < 0.05).



Weak/Moderate/Explosive Eruptions



- Subperiods in colors refer to the intensity of volcanic activity. Blue: weak volcanic activity (850-1100 C.E.); pink: explosive volcanic activity (1100-1500); and yellow: moderate volcanic activity (1500-1850 C.E.).

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Impacts on Precipitation

AEM-index regressed onto precipitation anomalies (mm-1/month).



<u>Weak</u> volcanism (850-1100 C.E.): low impact on PPT over South America and the Atlantic Ocean.

Explosive volcanism (1100 to 1500 C.E.): intense radiative cooling in the tropical region weakens the relationship between the AEM and PPT over SA/Atl.

<u>Moderate</u> volcanism (1500-1850 C.E.): opposite effects whe compared to the explosive volcanic activity: intensifies the impact of the AEM and PPT over SA/Atl.



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Volcanic impact on PPT proxy

From stalagmites (SA cave)



(a) Central-eastern Brazil δ^{18} O record from Anjos1 stalagmite with volcanic sulfate aerosol accumulation¹⁷.



Discussion - Conclusions

Impact of the volcanic activity Radiative cooling

- Radiative cooling of the Eq. Atl. in the year of eruption is followed by warming in the Eq. Atl. four years later.
- Eq. Atlantic warming observed in the 4th year after the eruption.
- Positive AEM + volcanic activity weakens PPT in the Amazon basin and NE-Brazil.
- Pinatubo showed decrease in continental PPT in the Amazon and NE-Brazil (Trenberth and Dai, 2007).
- The radiative cooling caused by the explosive volcanic eruptions during the LM was so intense that it weakened (decoupled) the thermodynamics of air-sea interactions.



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THANK YOU

