



funded in by NSF Climate and
Large-Scale Dynamics



Quantifying the seasonal sensitivity of the Northern Hemisphere jet-streams to Arctic warming

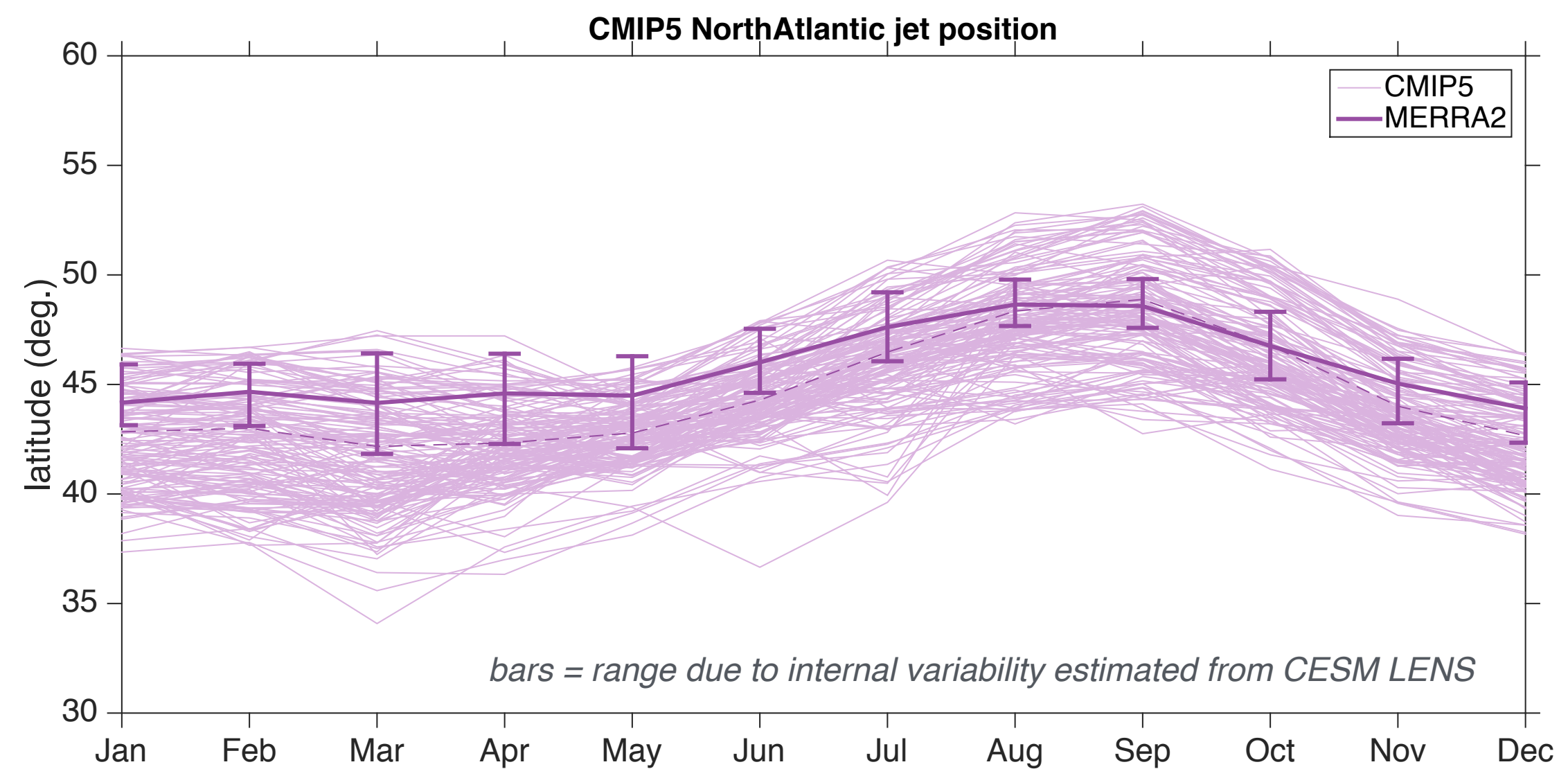
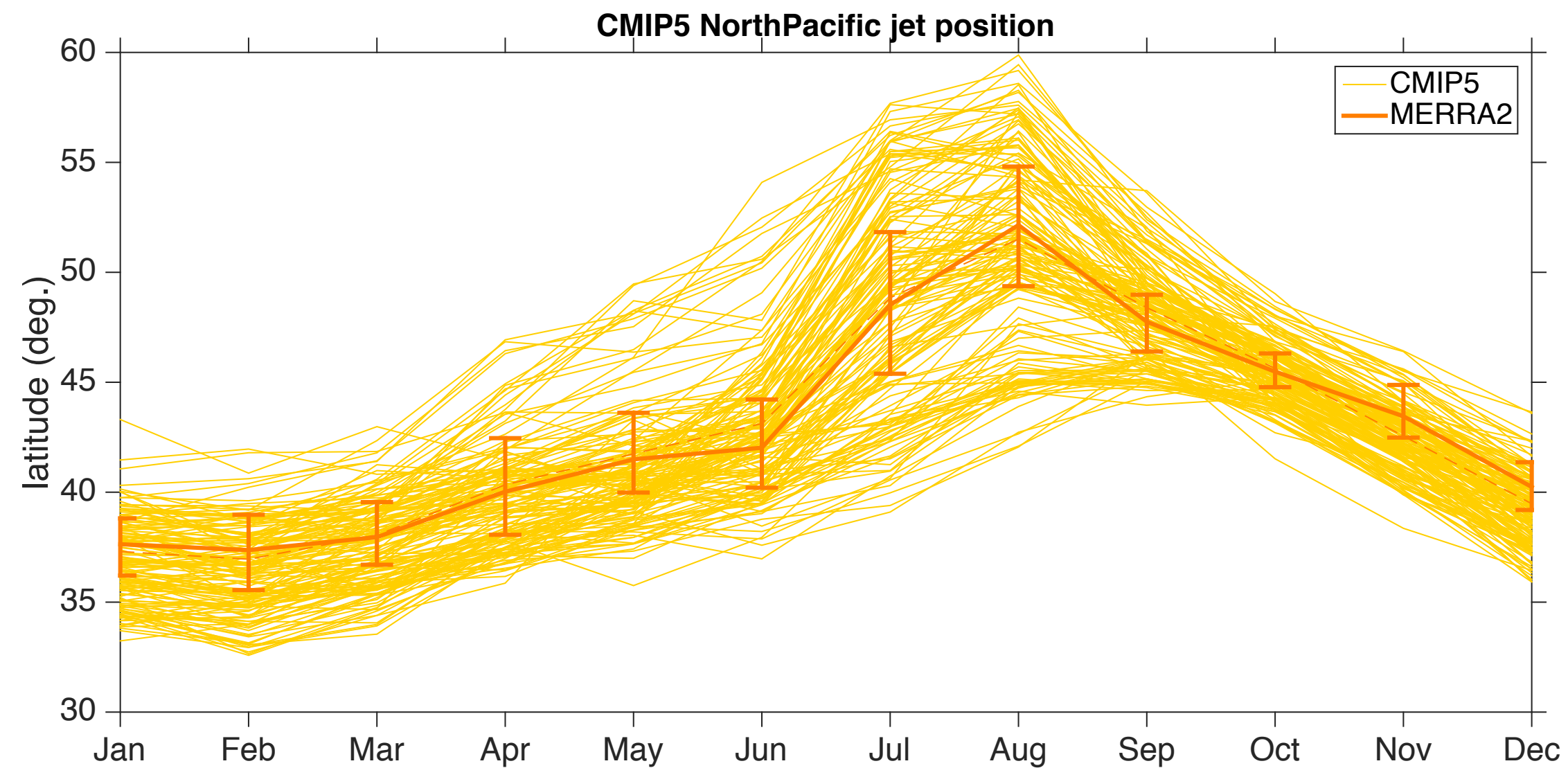
Elizabeth Barnes
Colorado State University

Isla Simpson
NCAR

along with graduate student collaborators

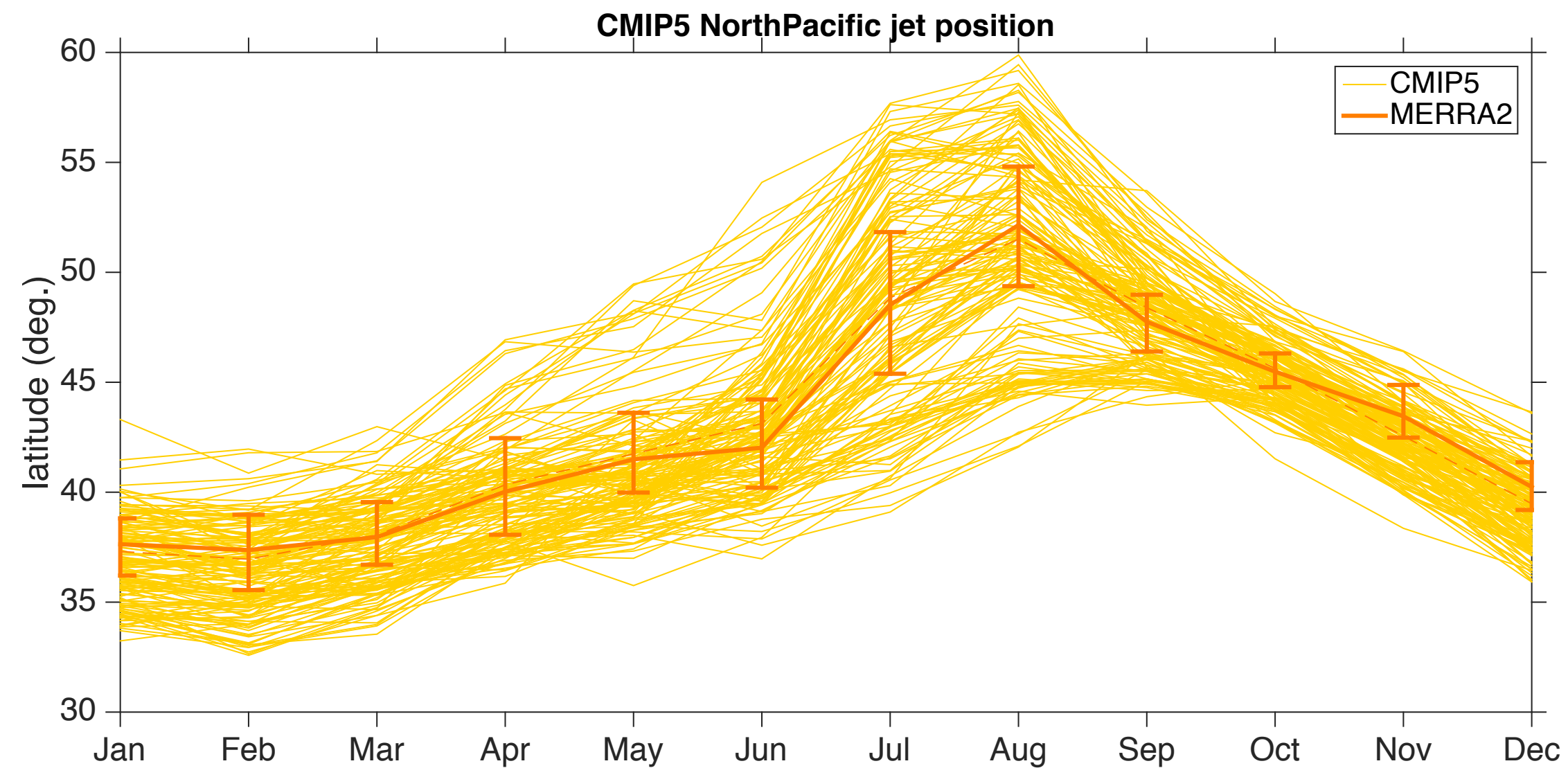
Marie McGraw, CSU
Bryn Ronalds, CSU

Seasonality of **the mean circulation**

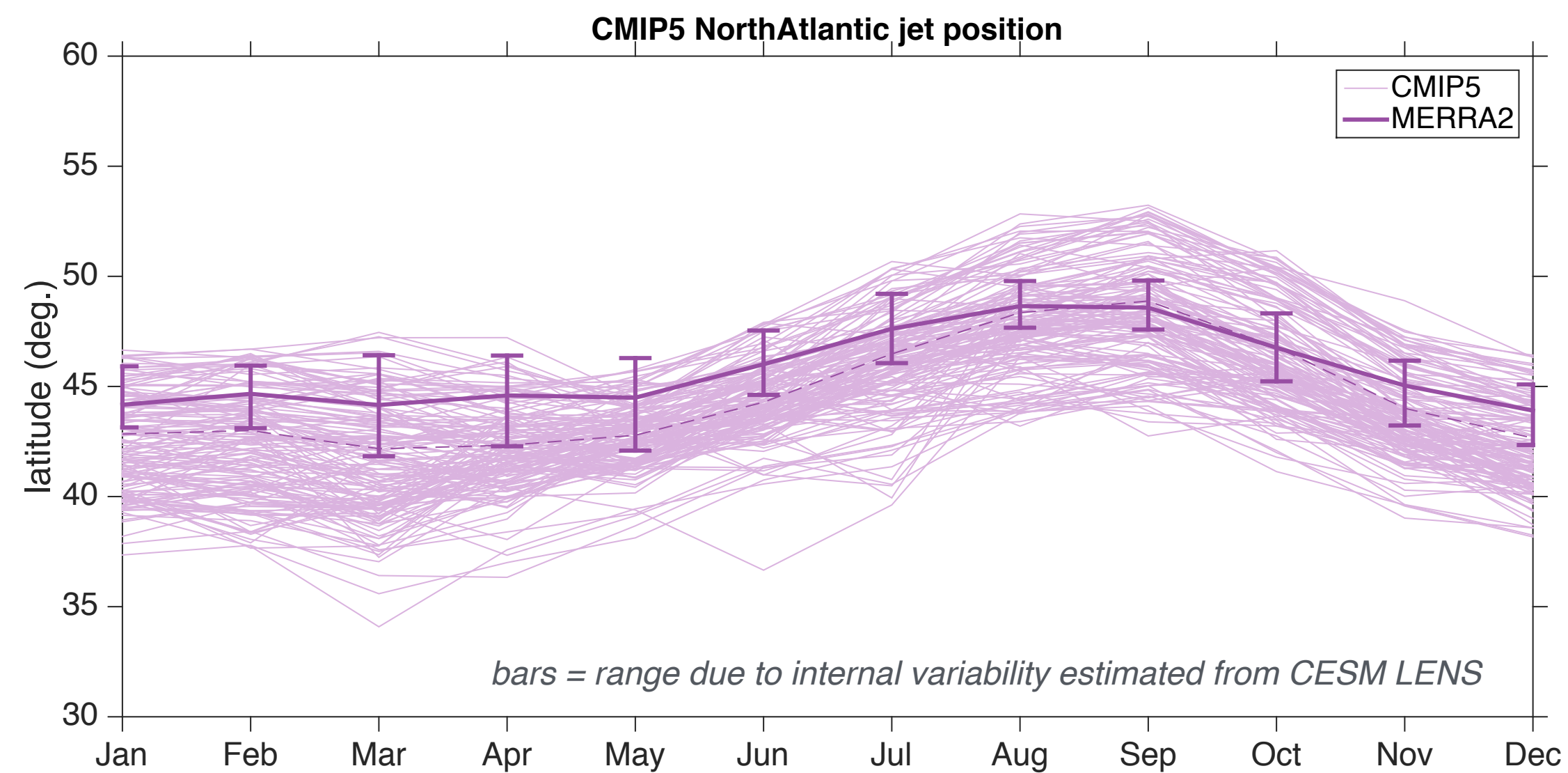


jet-streams exhibit seasonality
in their latitude and variability

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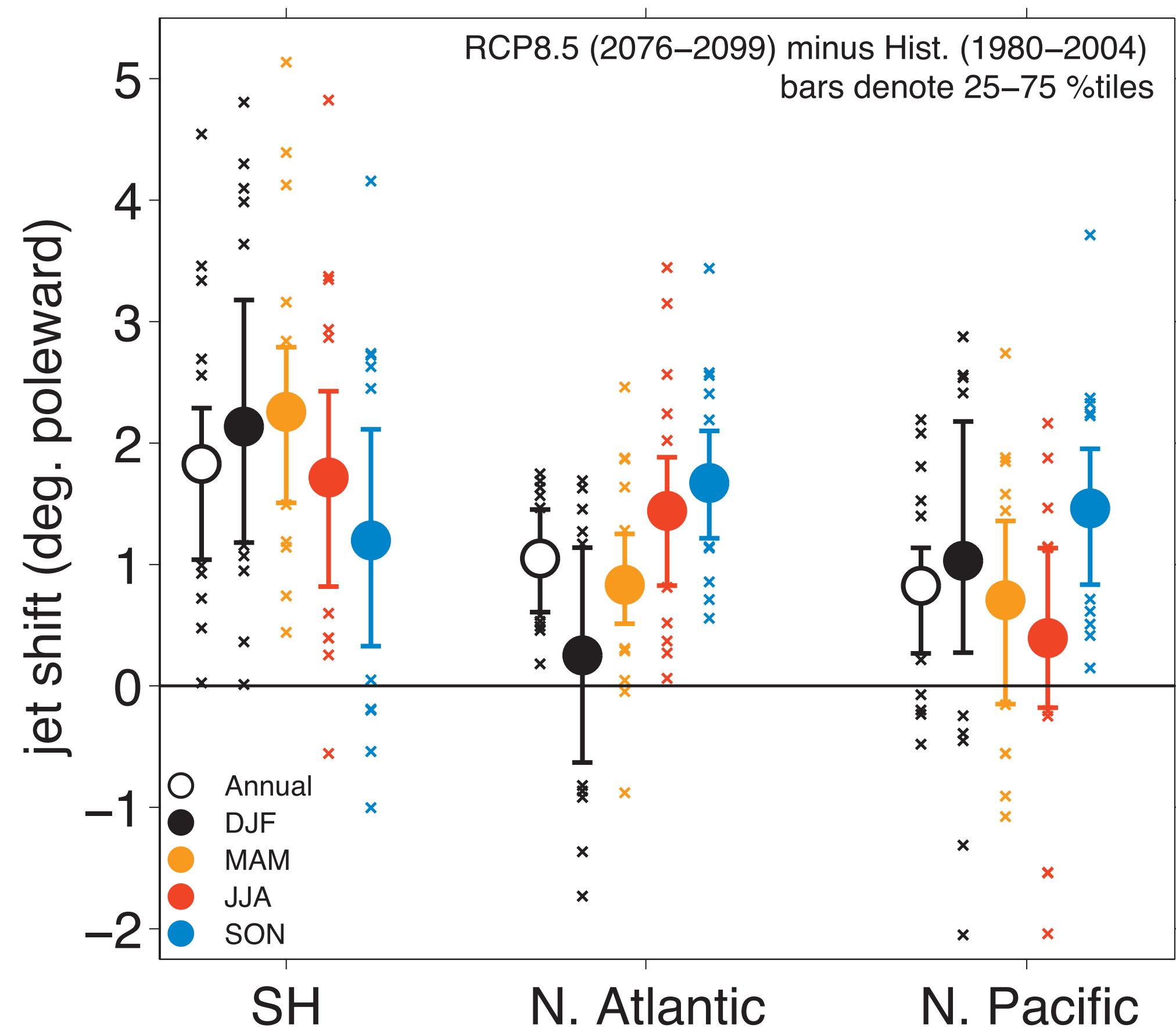


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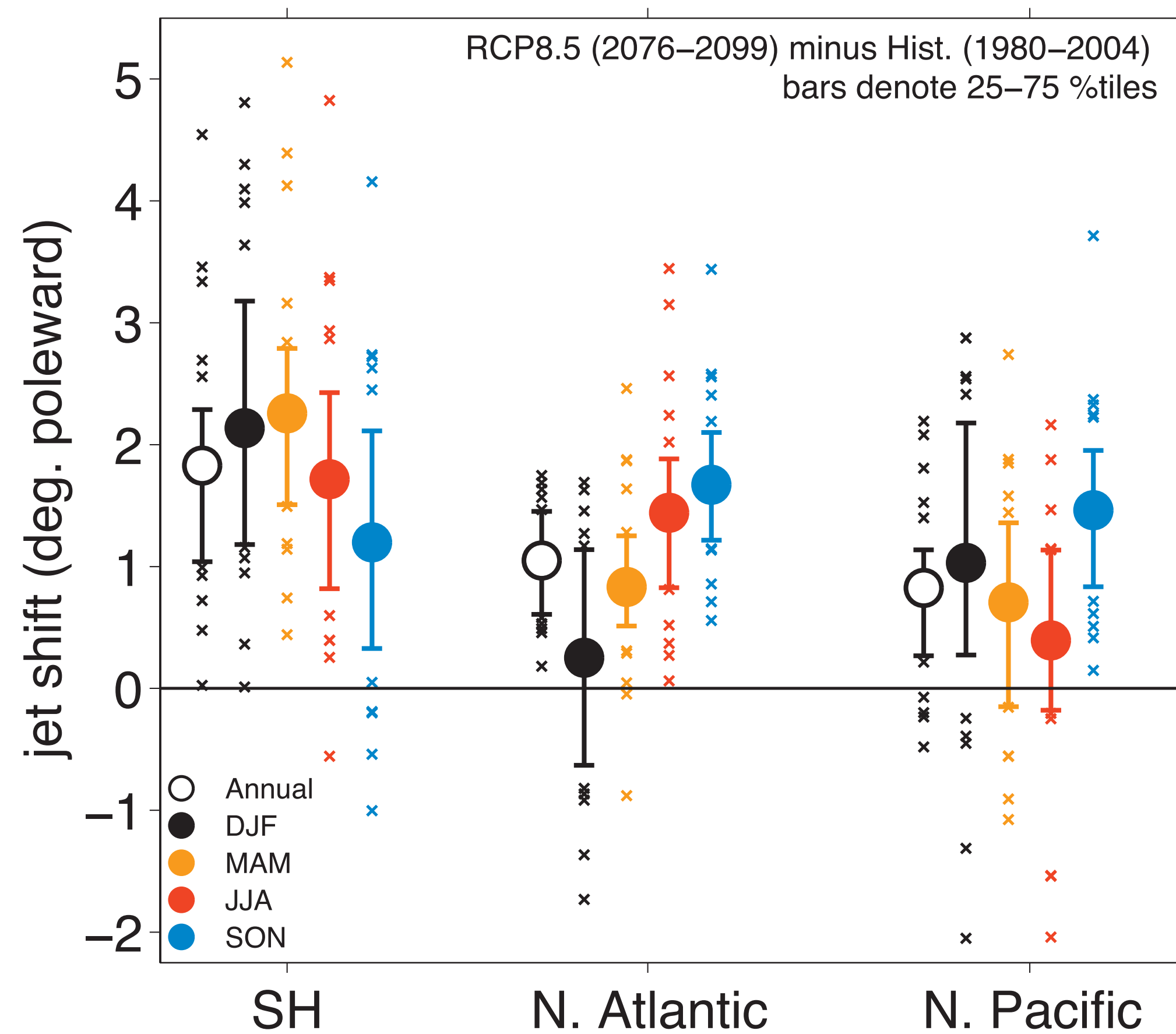


models exhibit seasonal biases
in the mean jet position

Seasonality of **future jet shifts**



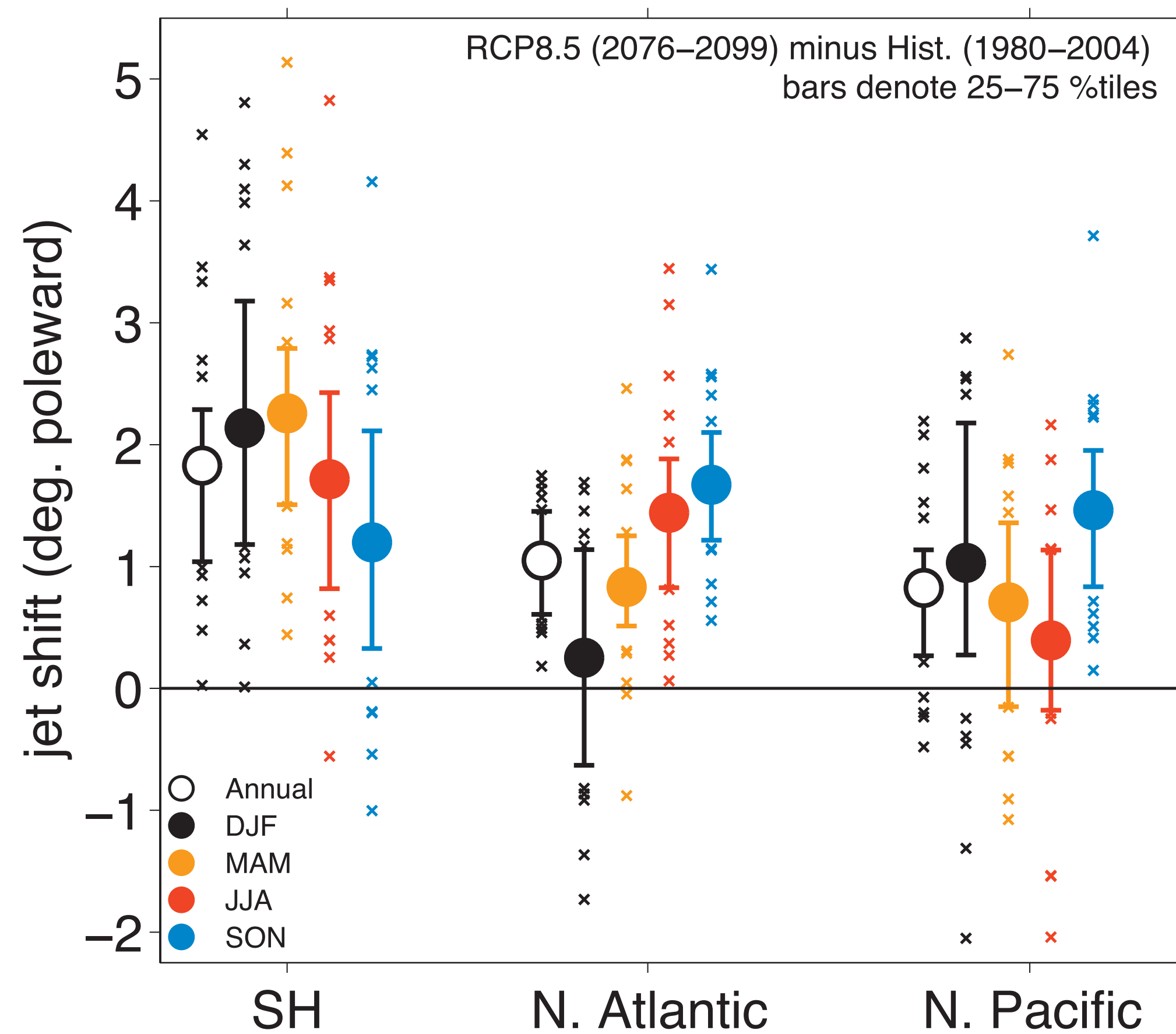
Seasonality of **future jet shifts**



jet shift has a rich seasonality that could be due to a few factors

- (1) **seasonality of forcing**
(e.g. sea ice loss)

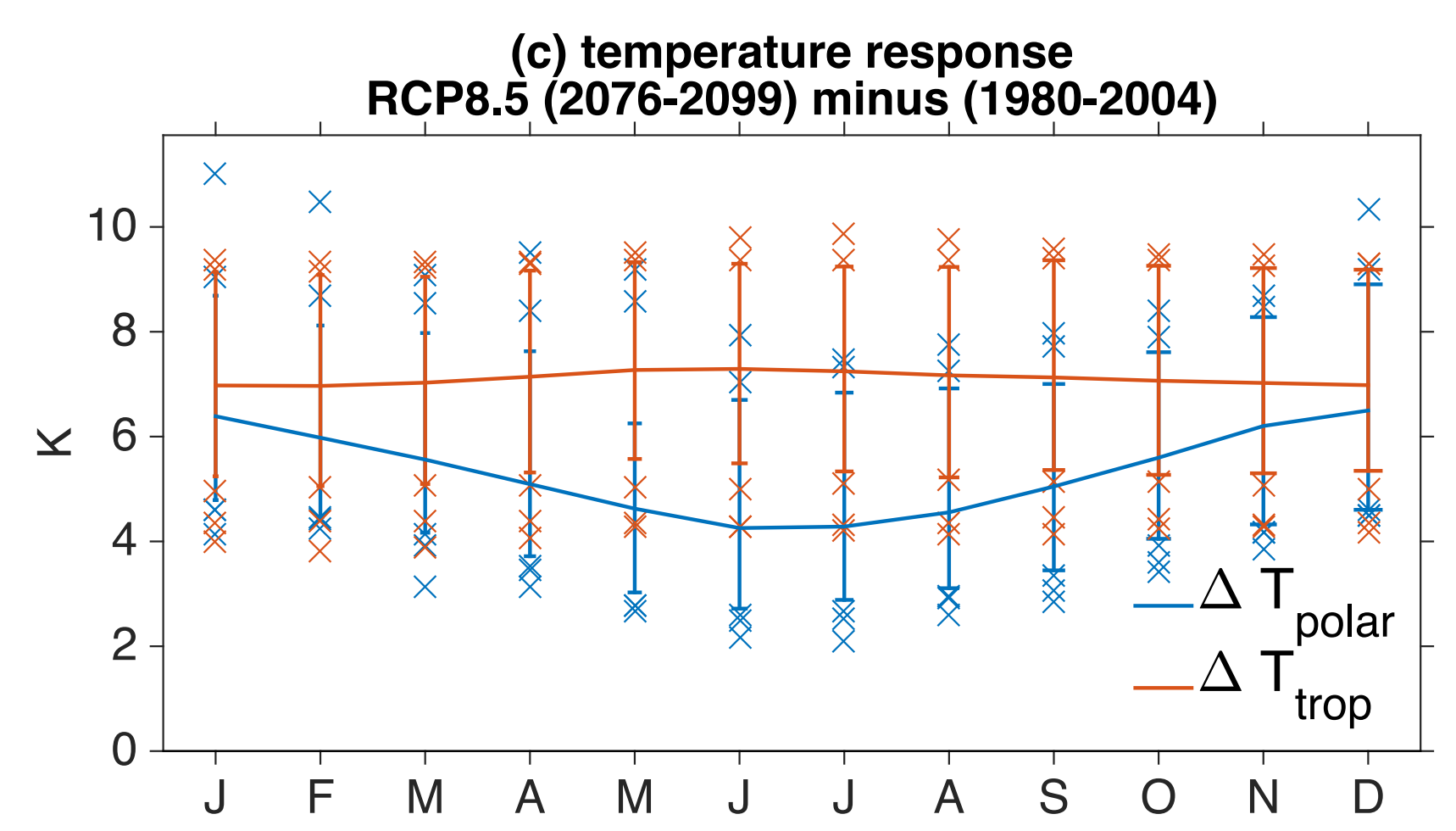
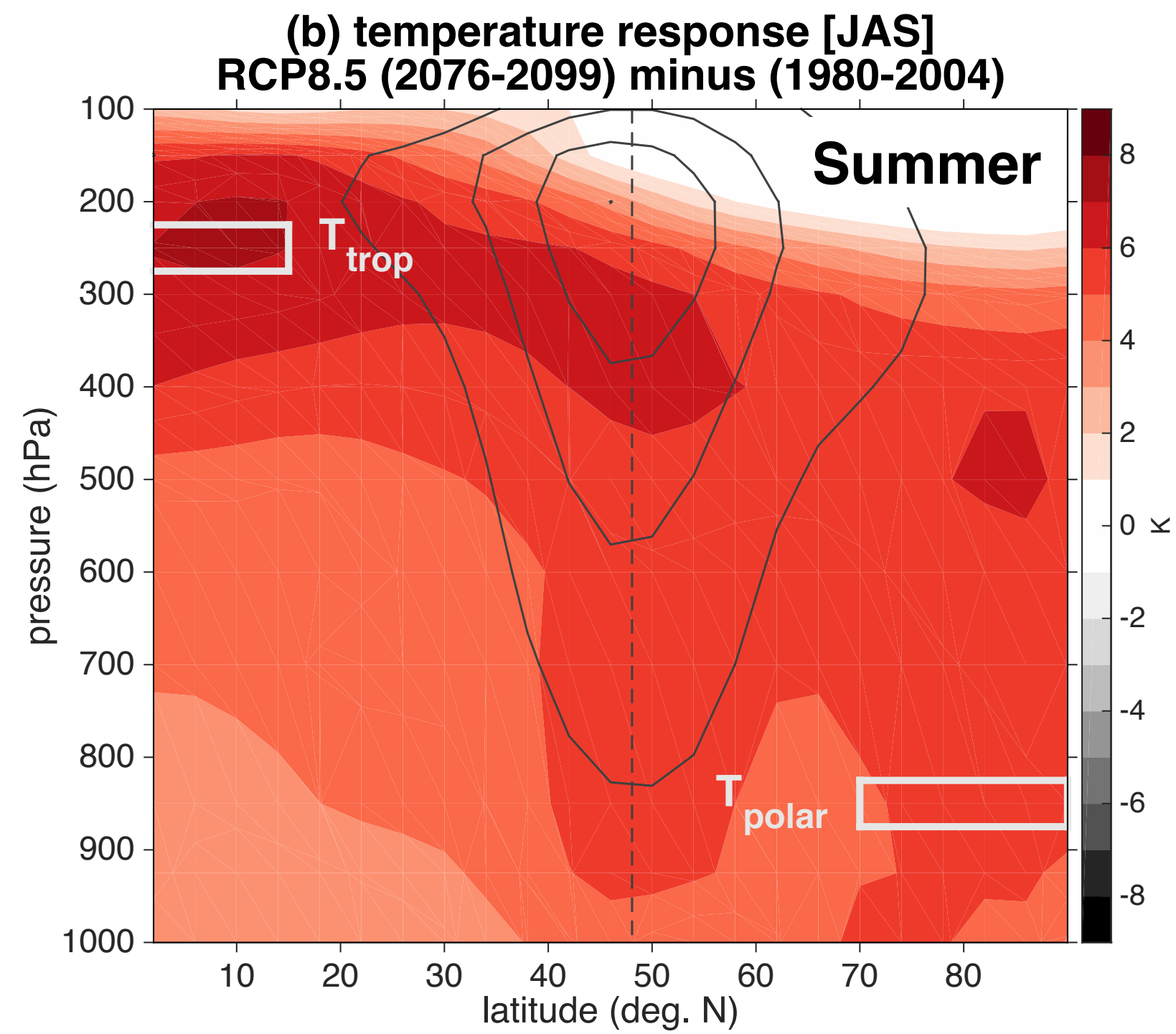
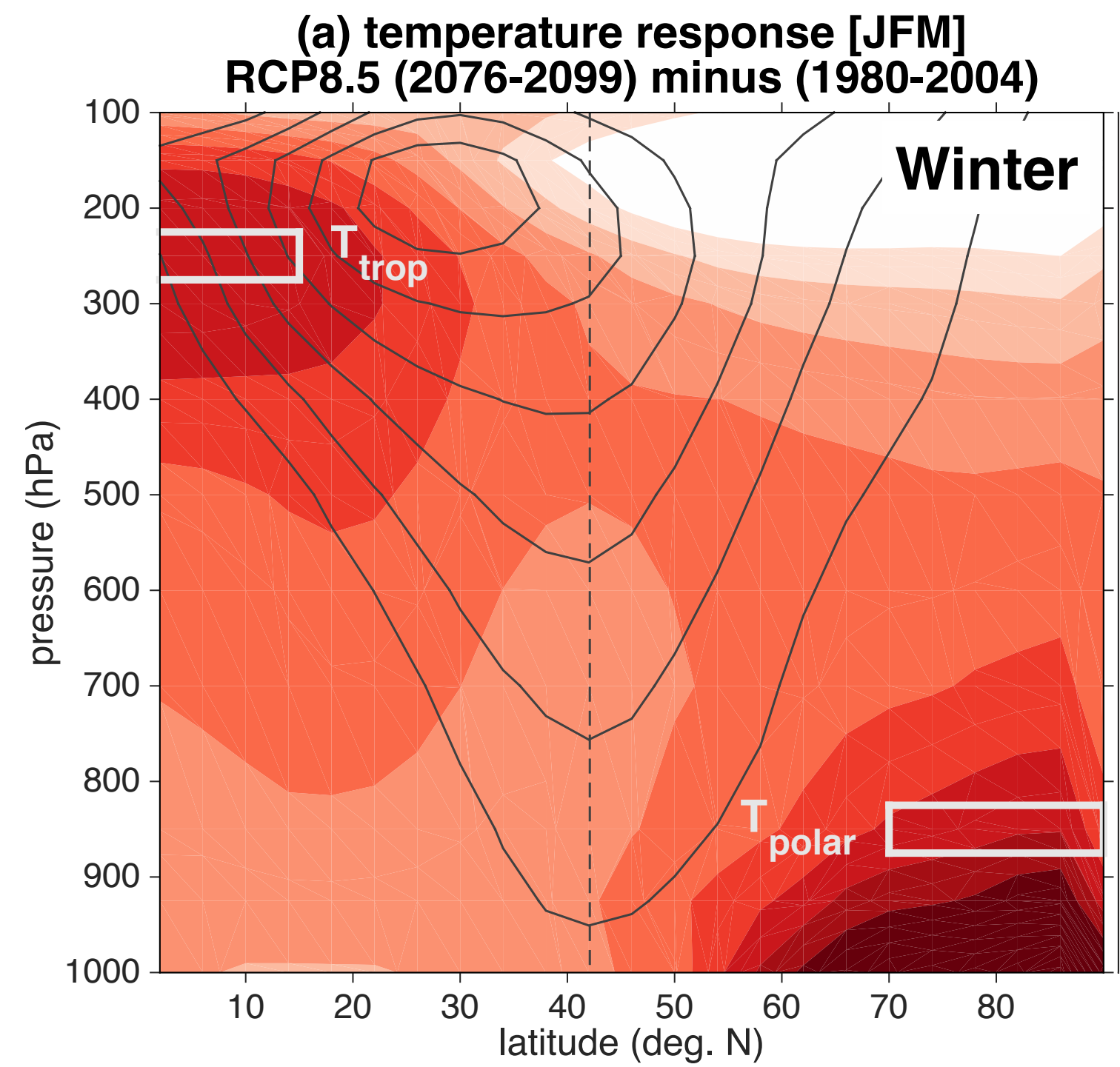
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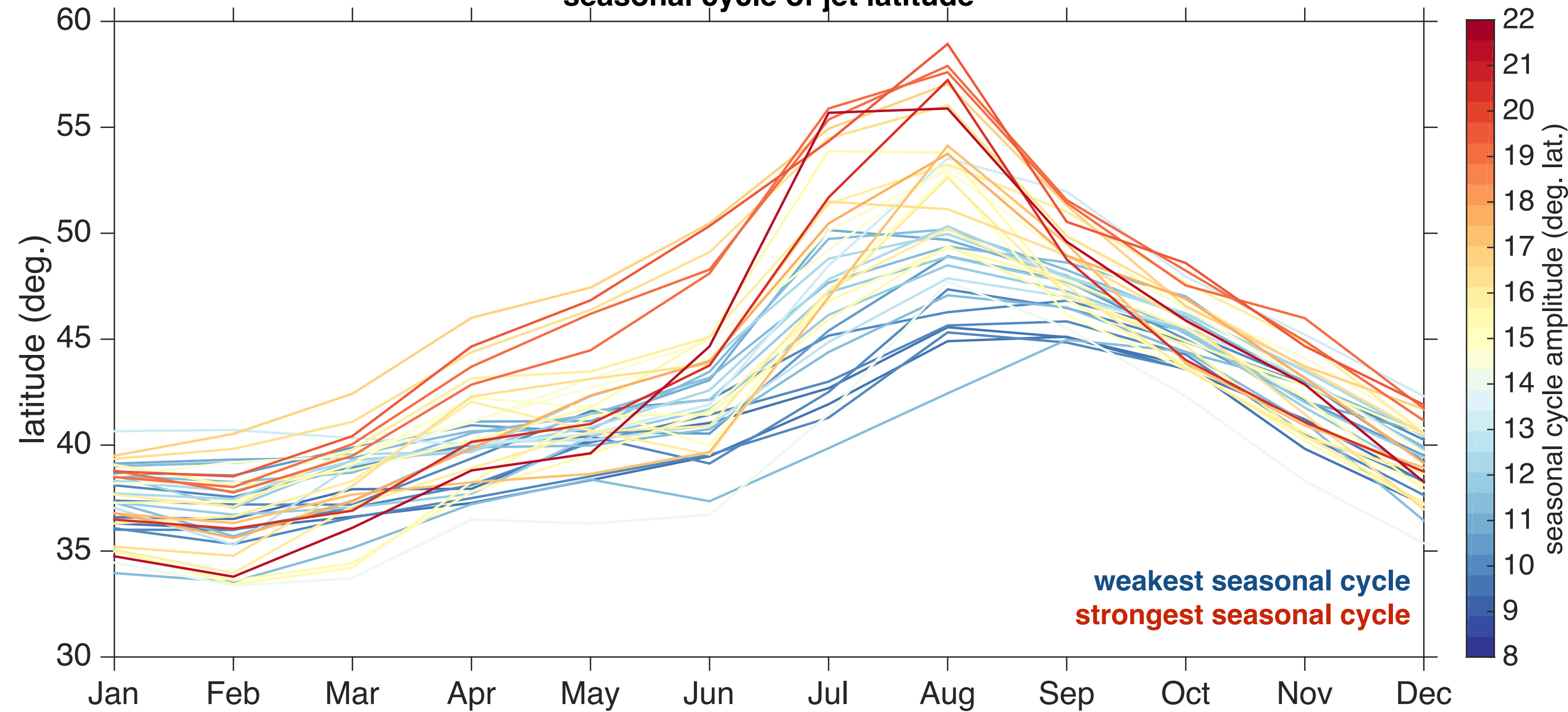
- (1) **seasonality of forcing**
(e.g. sea ice loss)
- (2) **seasonality of the circulation**
(e.g. even for constant forcing)

Seasonality of future warming



Seasonality of **the mean circulation**

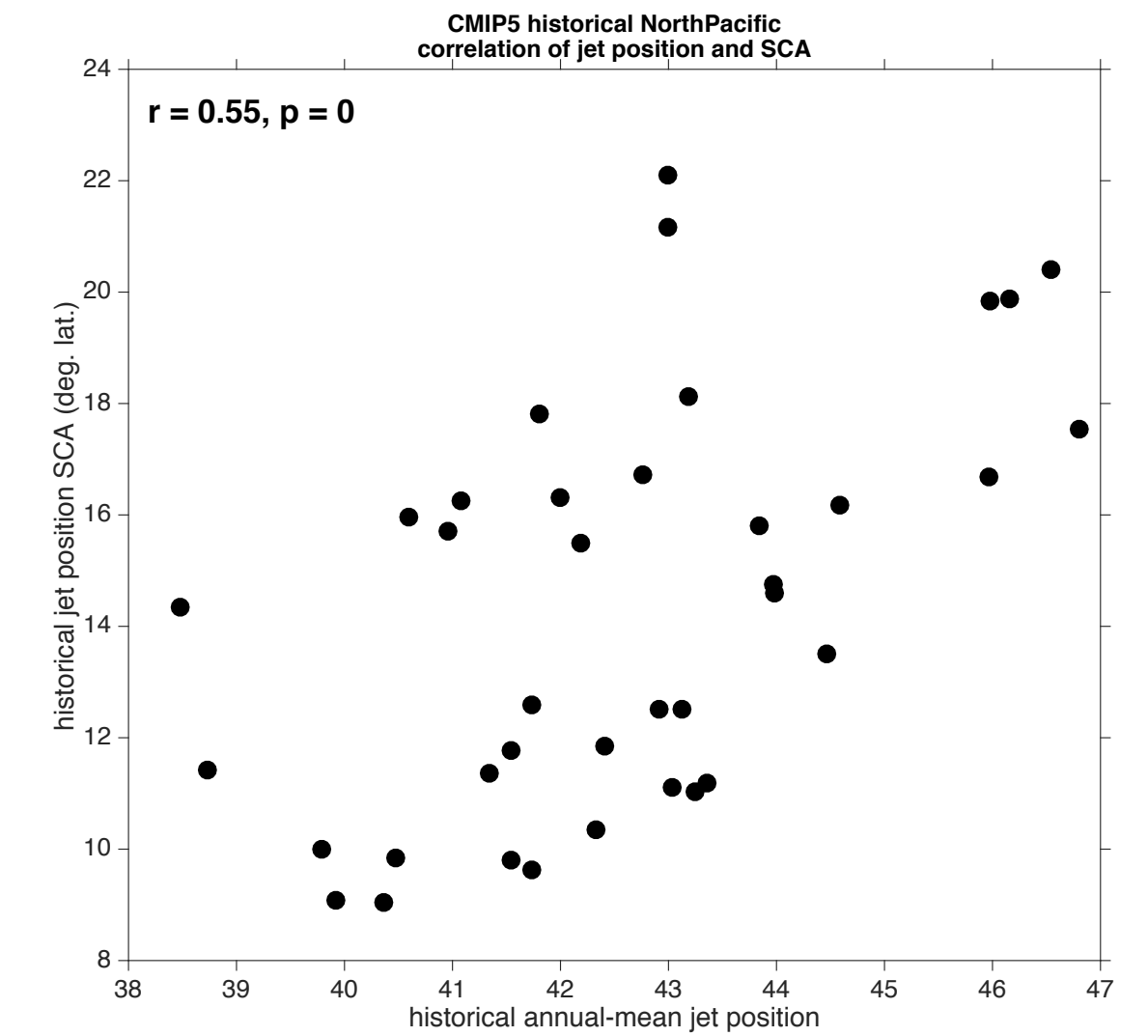
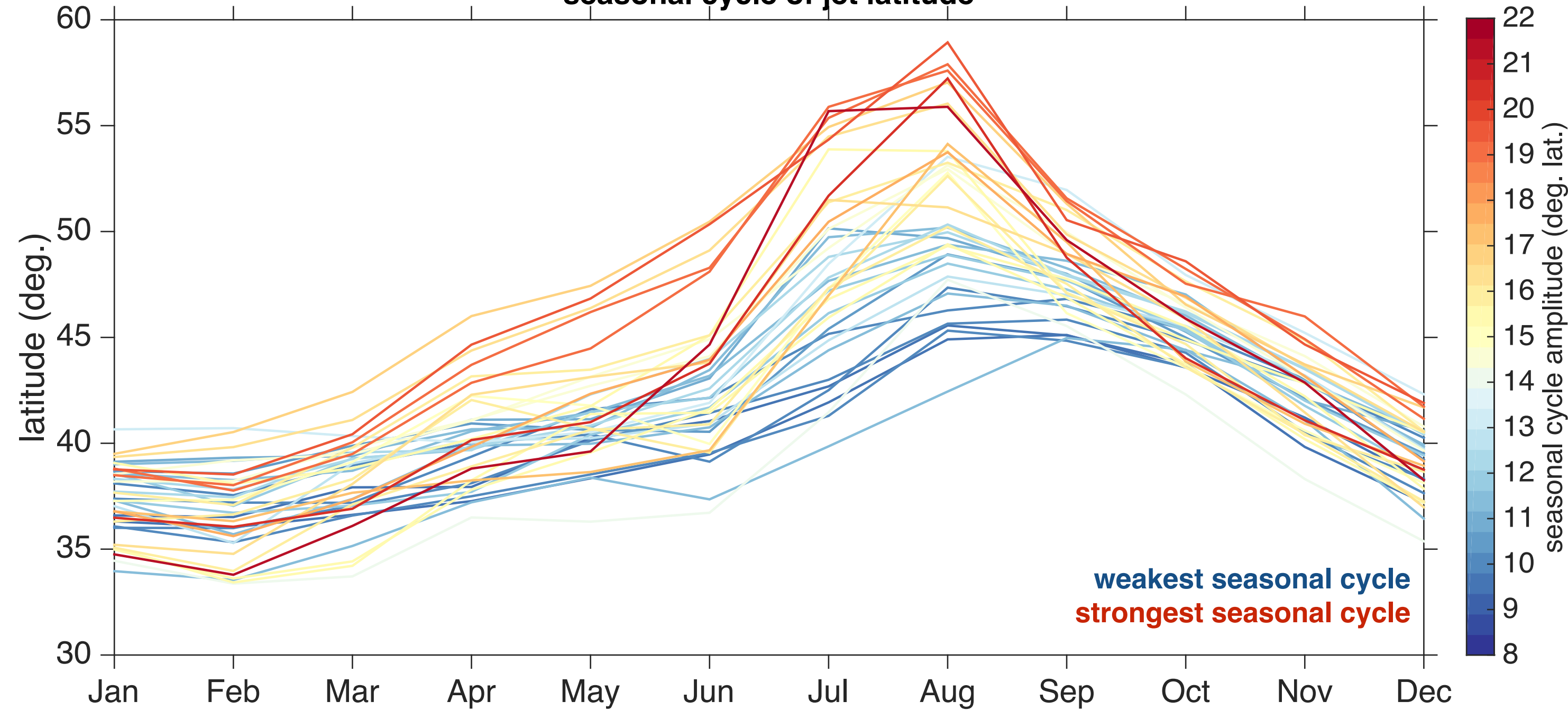
CMIP5 historical NorthPacific
seasonal cycle of jet latitude



differences in model ability to simulate the seasonality of the jet will likely impact the projected response

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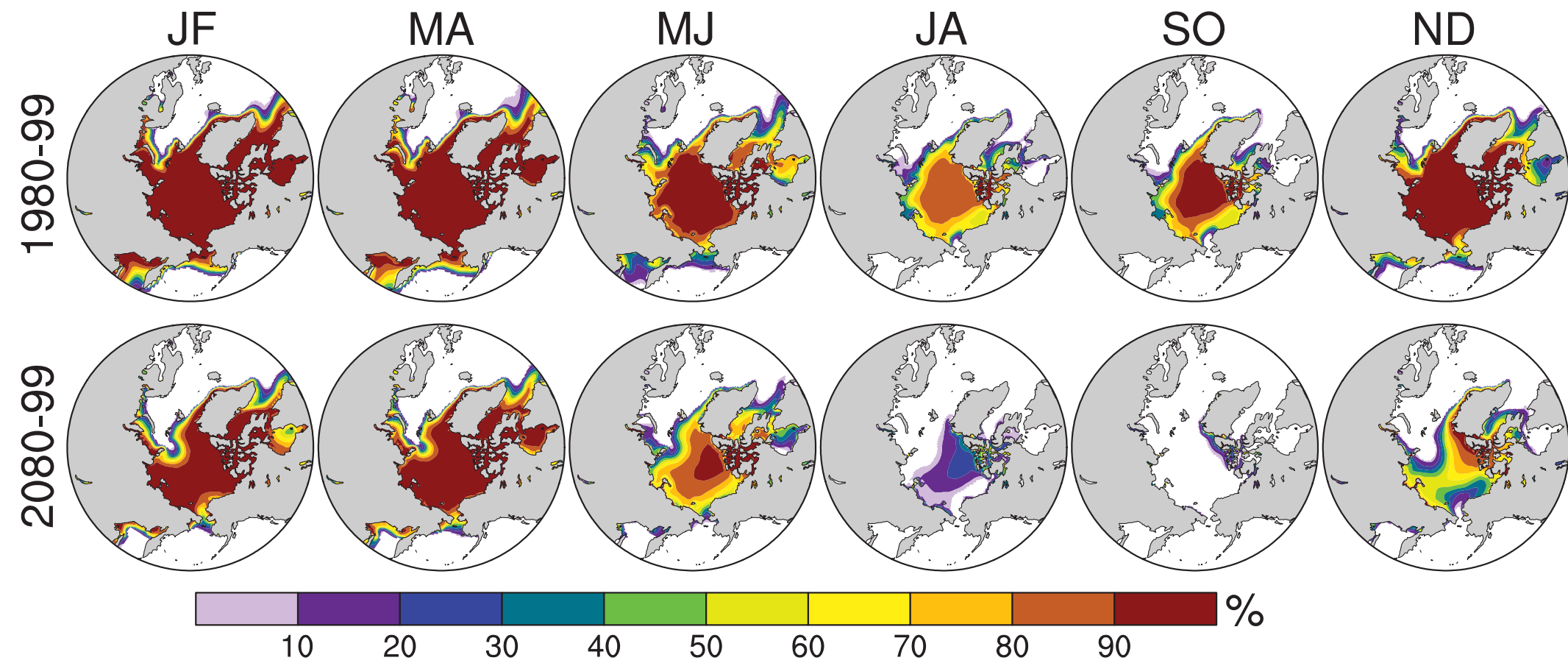
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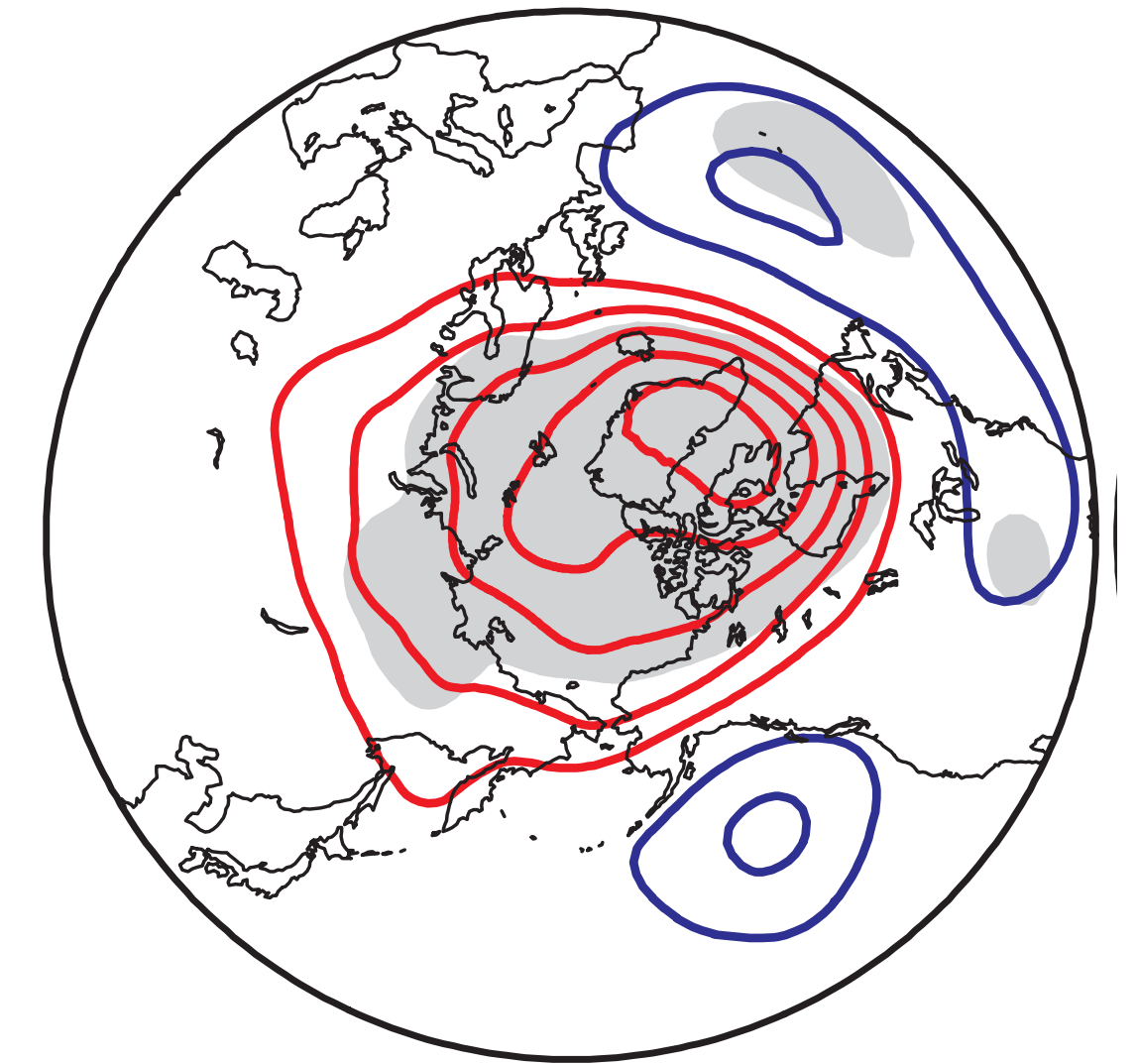
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Jet response to sea ice loss

a) Sea Ice Concentration



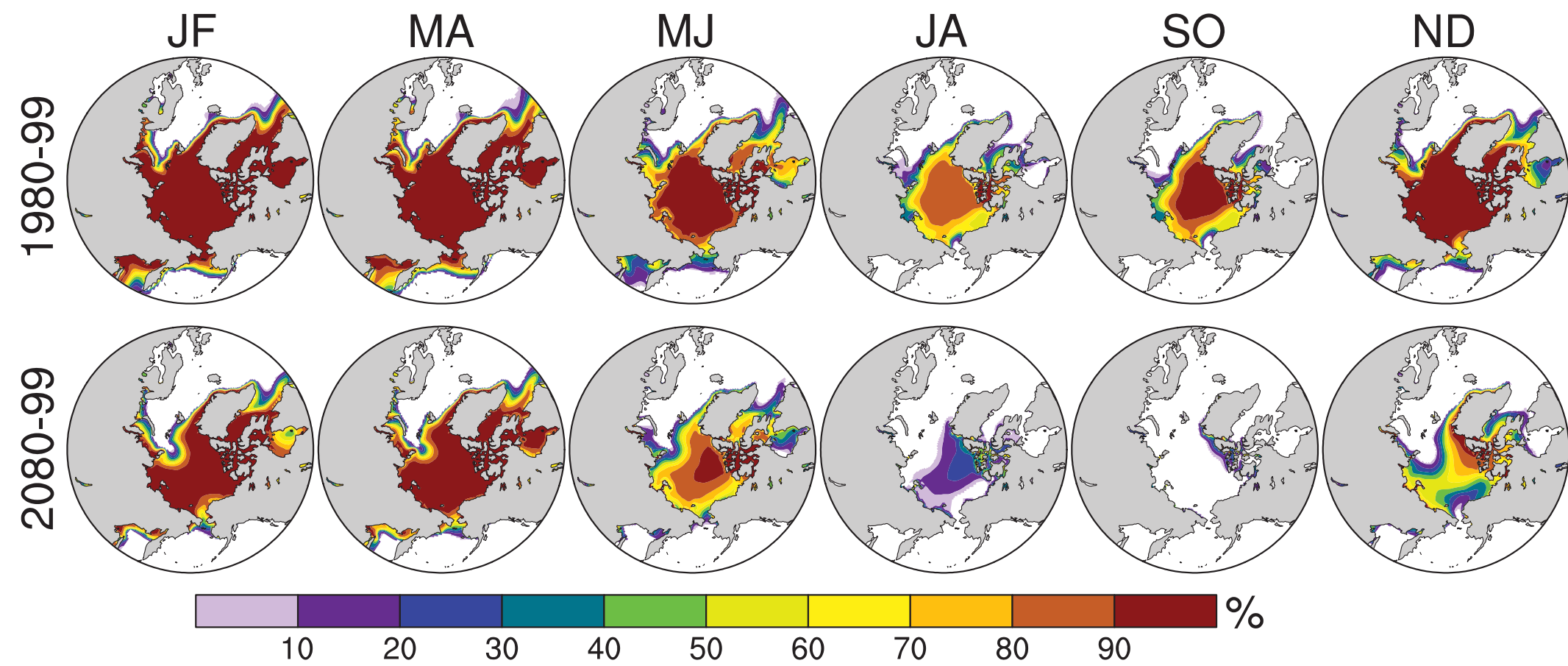
500 hPa geopotential height response in Jan.-Feb.



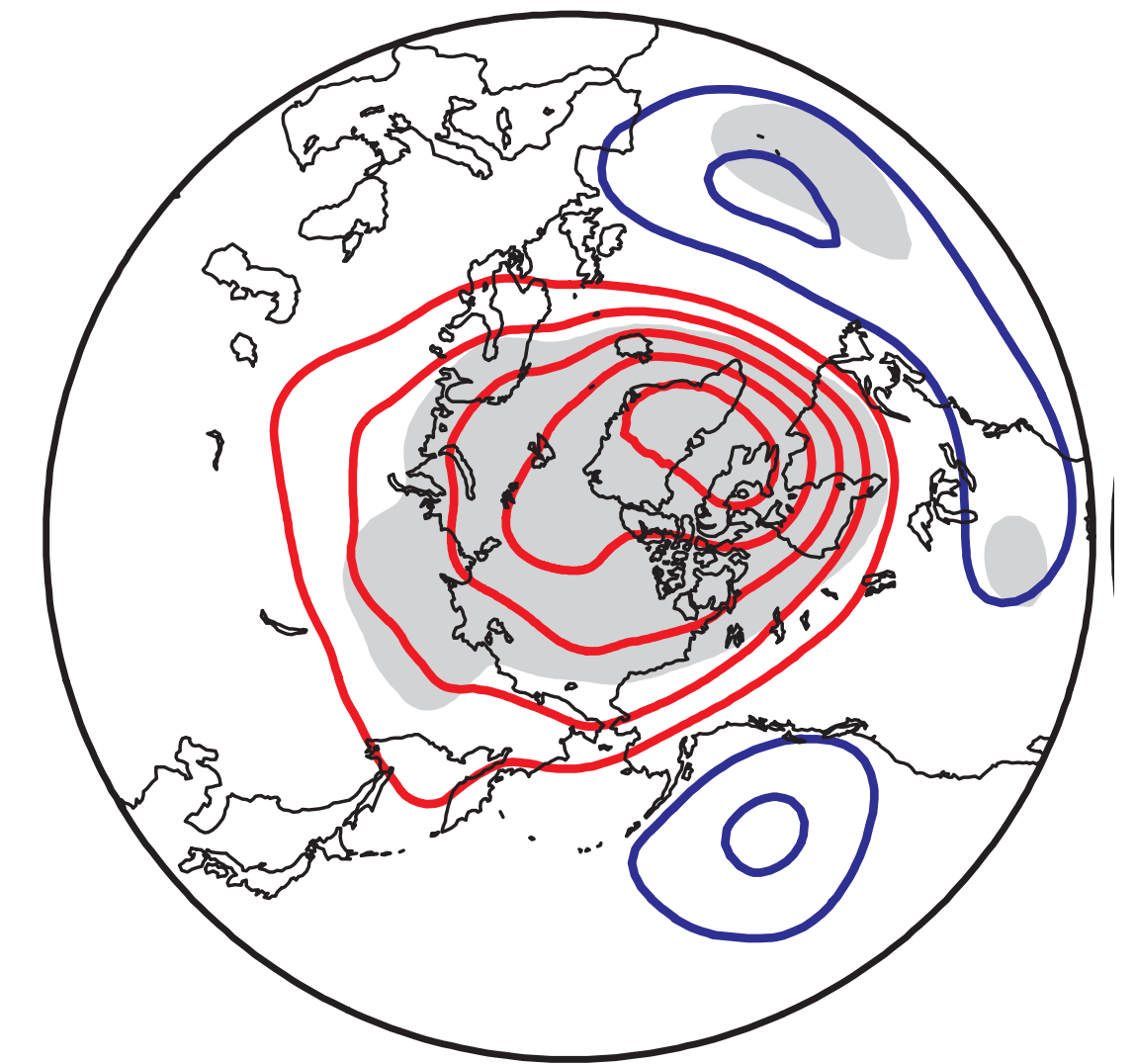
atmosphere-only CAM3 simulations
Deser, Tomas, et al. (2010; JCLI)

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abstract of Deser et al. (2010):

“The loss of Arctic sea ice is greatest in summer and fall, yet the response of the net surface energy budget over the Arctic Ocean is largest in winter.”

...

“[The circulation] response resembles the negative phase of the North Atlantic Oscillation in February only.”

atmosphere-only CAM3 simulations
Deser, Tomas, et al. (2010; JCLI)

Ultimate Goal

- a lot of work has been done understanding the net response of the jet to GHG warming...**this is not our goal here.**
- instead, we wish to quantify the **seasonal sensitivity to a 1K warming in a particular region**

$$\Delta \mathcal{J} = r \Delta T$$

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Ultimate Goal

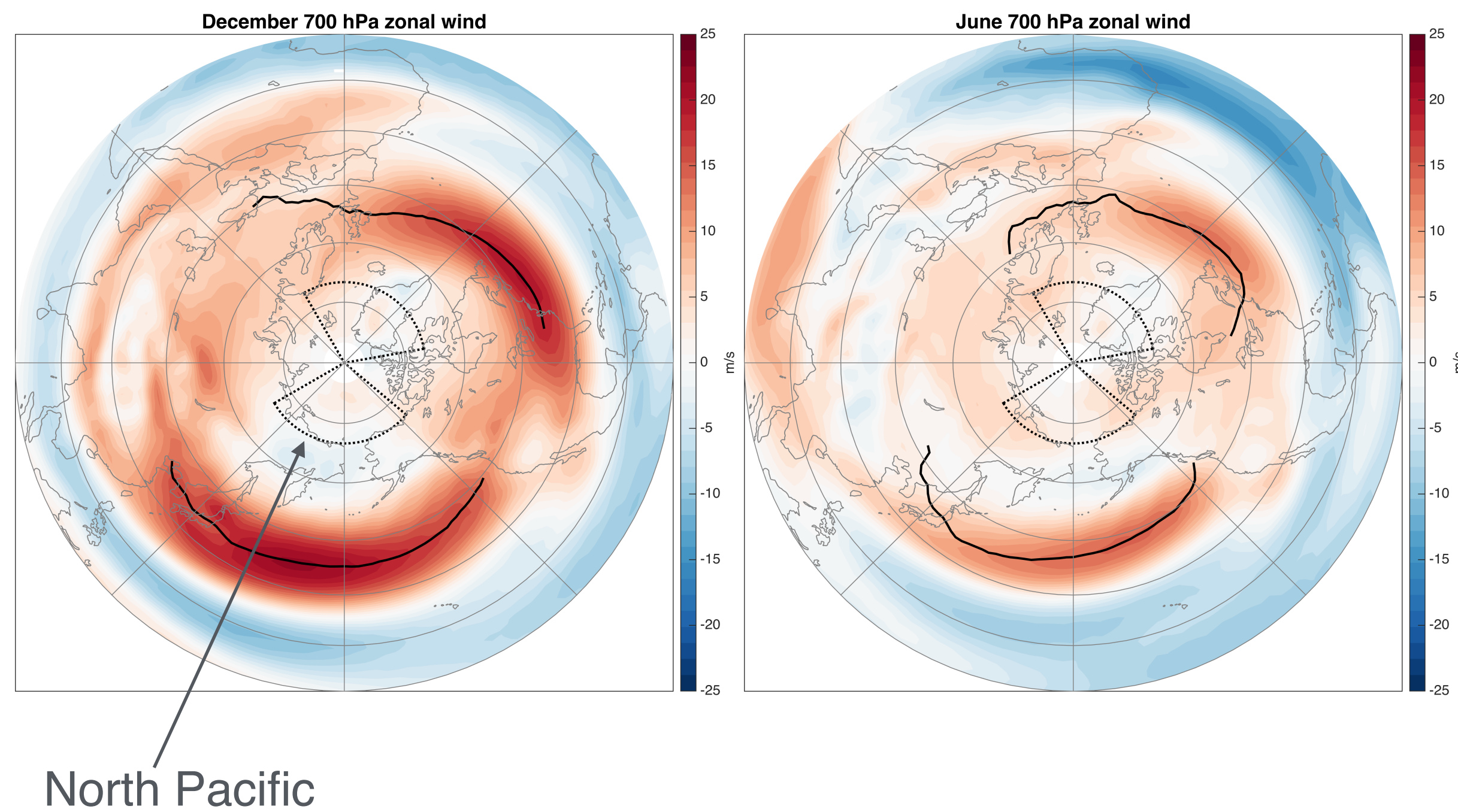
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Data



- CMIP5 (historical + RCP8.5: *detrended*)
- daily data (averaged into 10-day chunks*)
- 700 hPa zonal wind
- 850 hPa air temperature

Approach

$$\mathcal{J}_t = rT_t + \epsilon$$

time
(within a particular month)

...tried this, but got worried about the direction of causality given the lag 0

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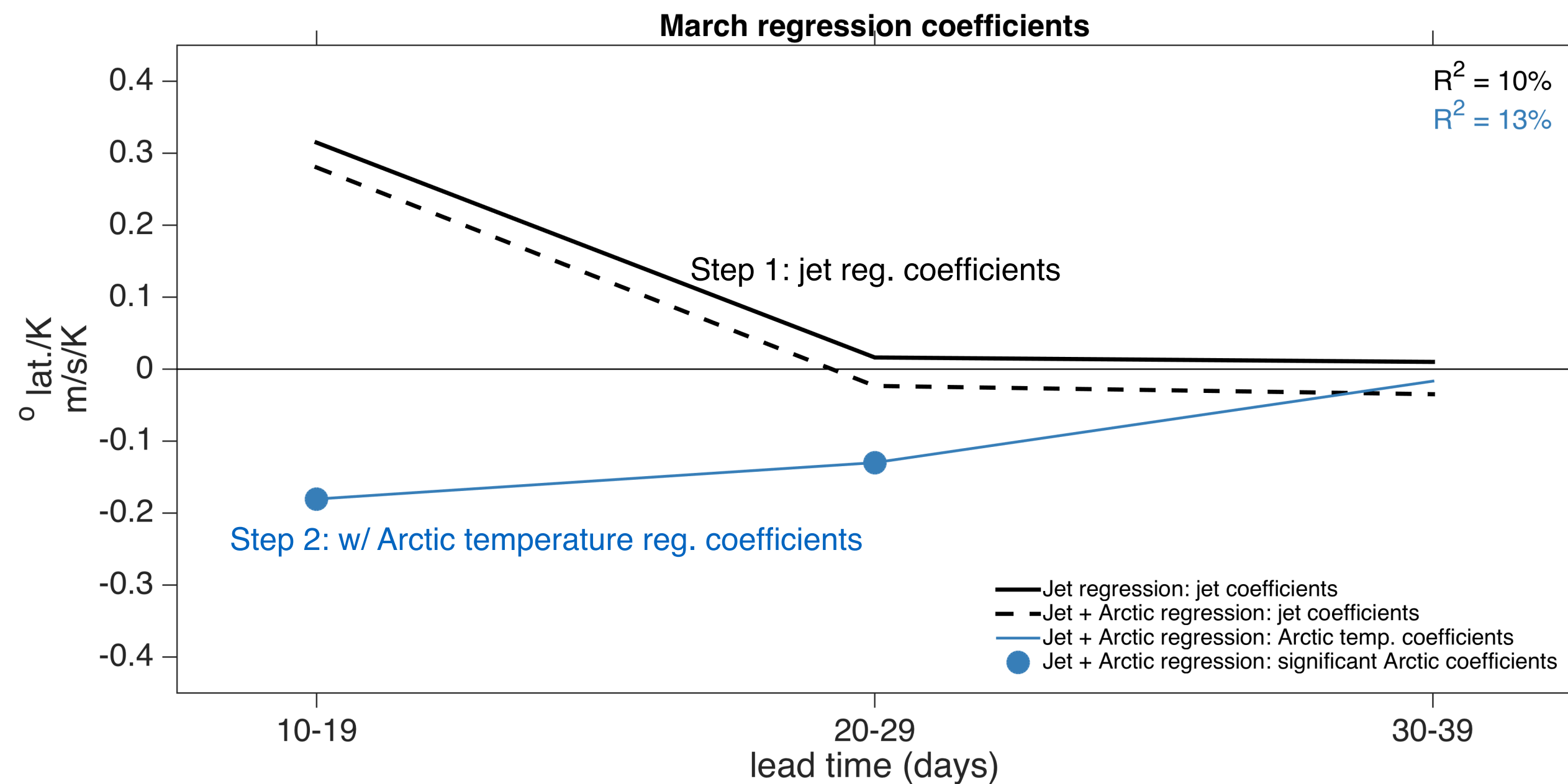
Granger-causality

Step 1
$$\mathcal{J}_t = a_0 + a_1\mathcal{J}_{t-1} + a_2\mathcal{J}_{t-2} + \dots + a_k\mathcal{J}_{t-k} + \epsilon_t$$

Step 2
$$\mathcal{J}_t = c_0 + c_1\mathcal{J}_{t-1} + c_2\mathcal{J}_{t-2} + \dots + c_k\mathcal{J}_{t-k} + b_1T_{t-1} + b_2T_{t-2} + \dots + b_kT_{t-k} + \epsilon_t$$

- there exists at least one significant “b” according to a t-test
- all of the “b” terms collectively add power to the regression

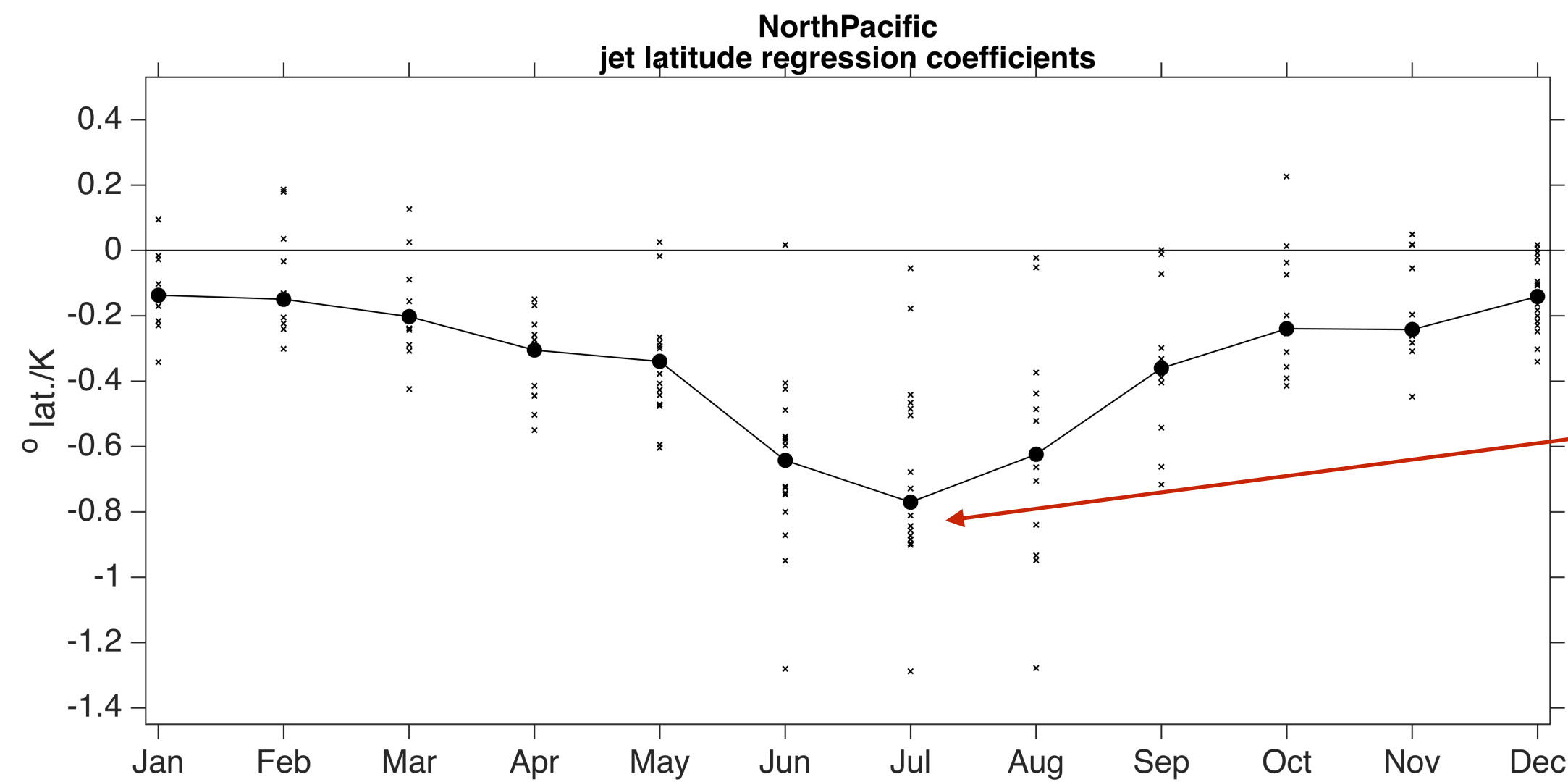
Approach



- jet shifts equatorward when Arctic was warm 10-30 days earlier
- first two coefficients are significant

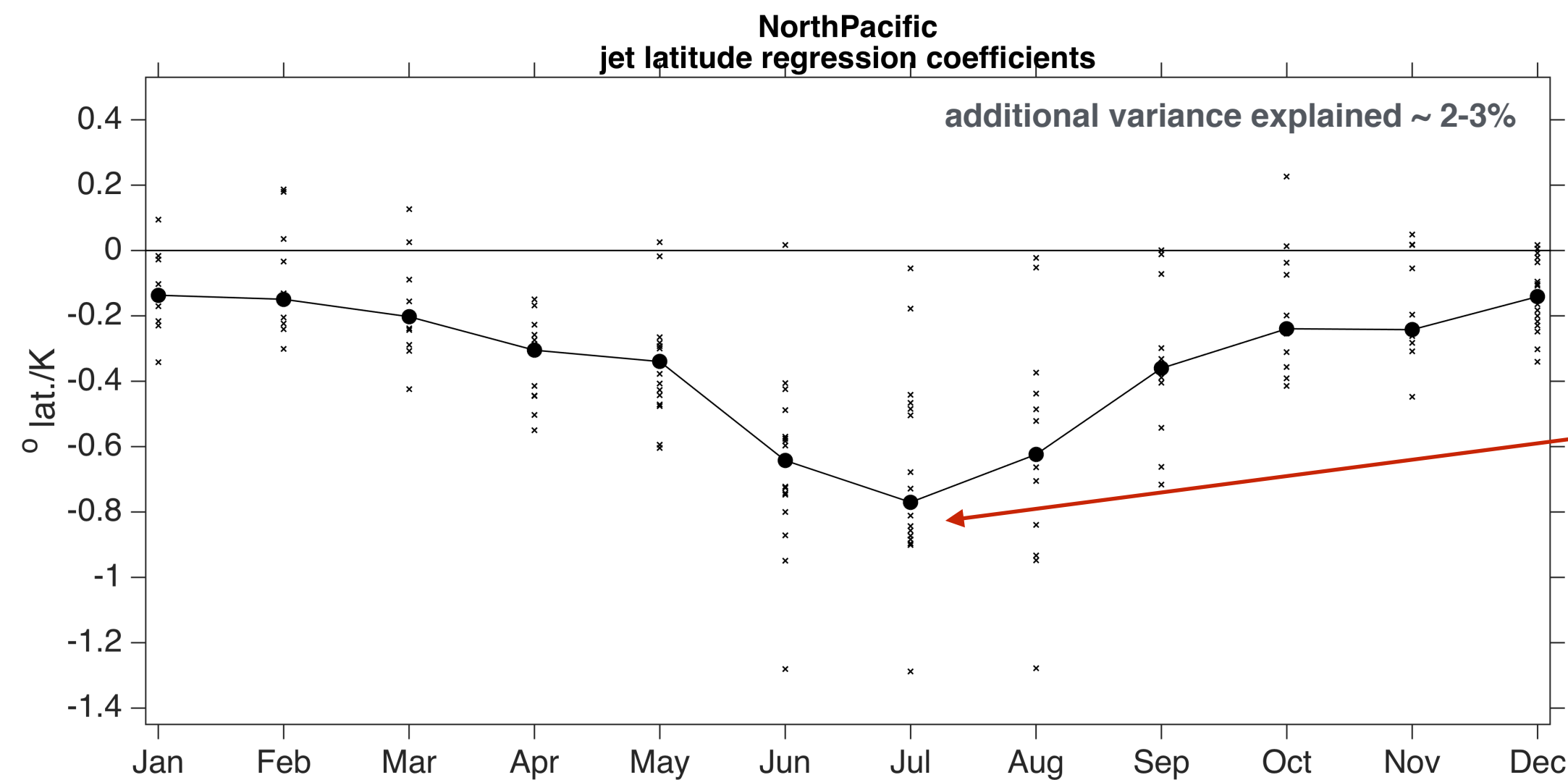
An example Granger causality calculation for North Pacific jet latitude in March from the CanESM2 model. For this example, we use the combined Historical + RCP8.5 time series.

Seasonal sensitivity: **jet position**

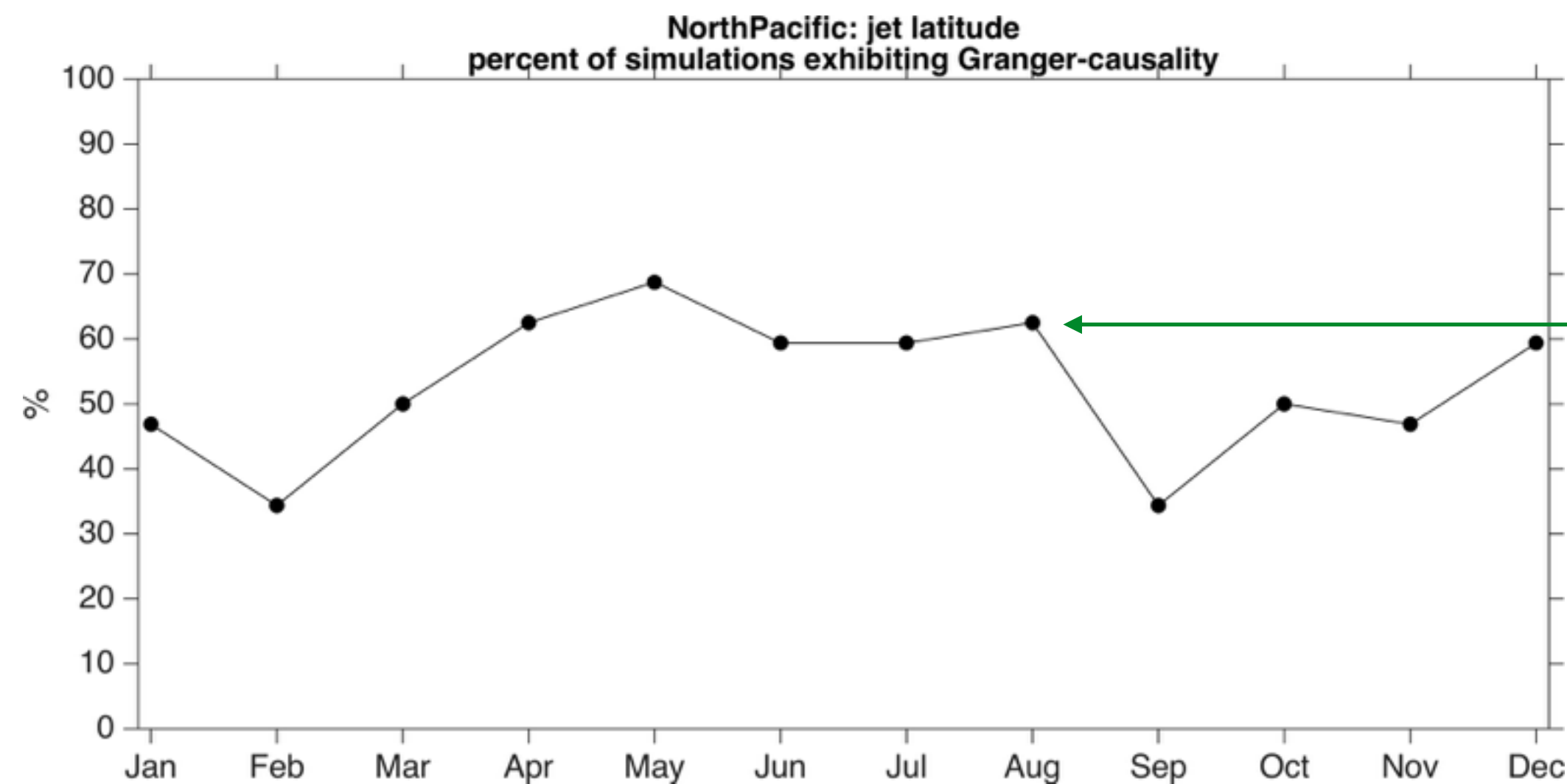


• jet shifts equatorward more in warm months

Seasonal sensitivity: jet position

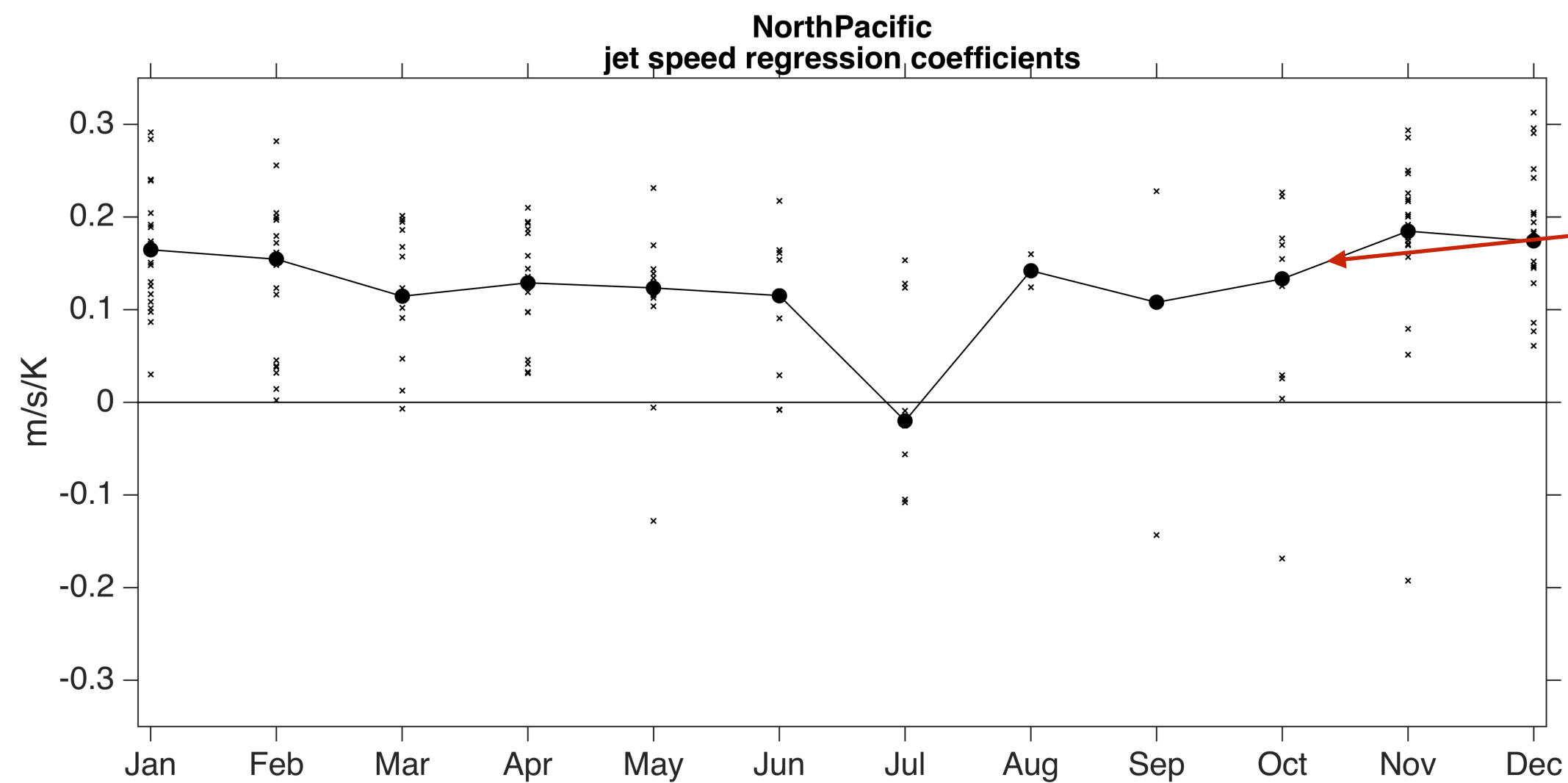


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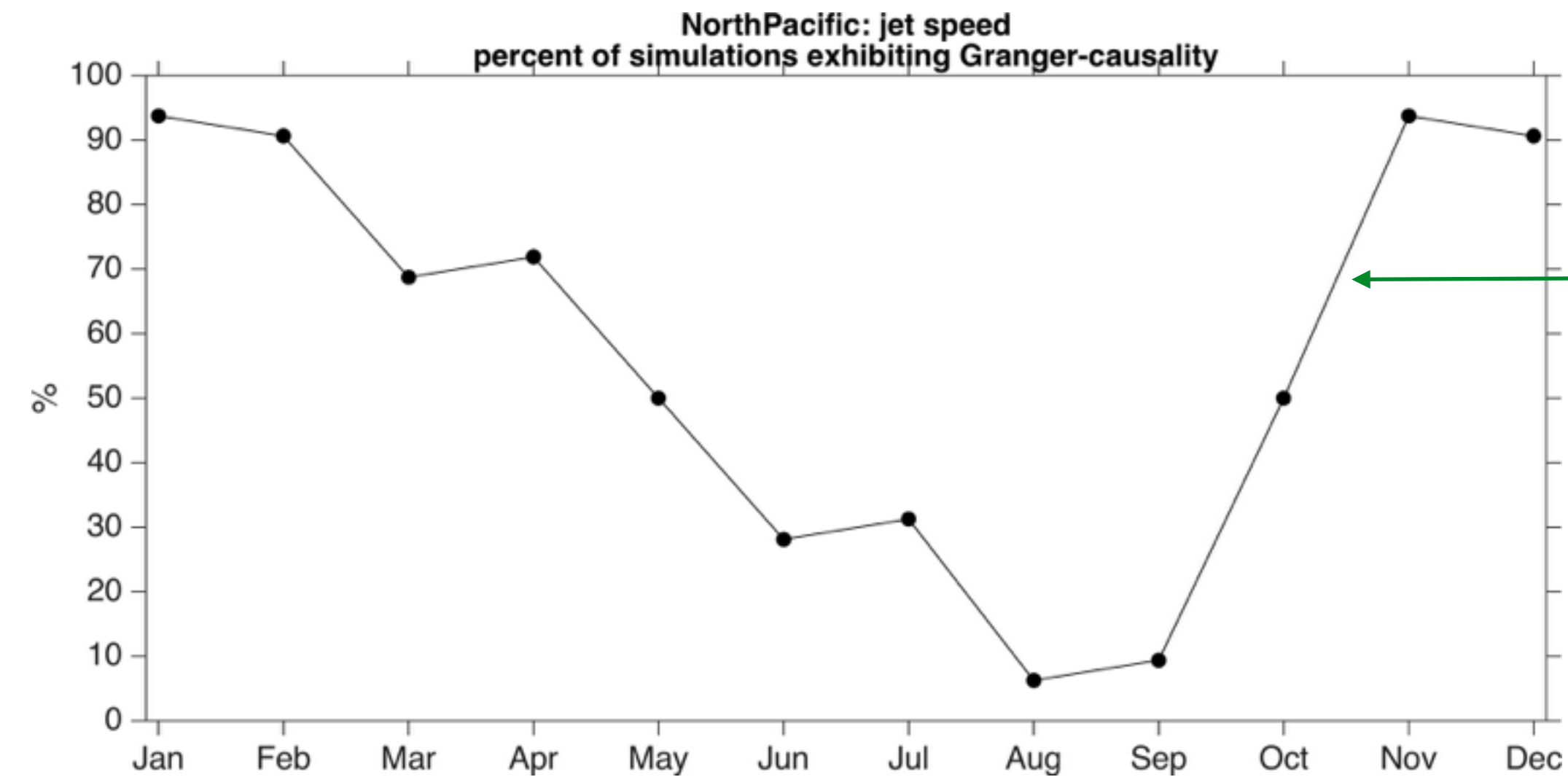


40-60% of simulations exhibit Granger-causality

Seasonal sensitivity: **jet speed**

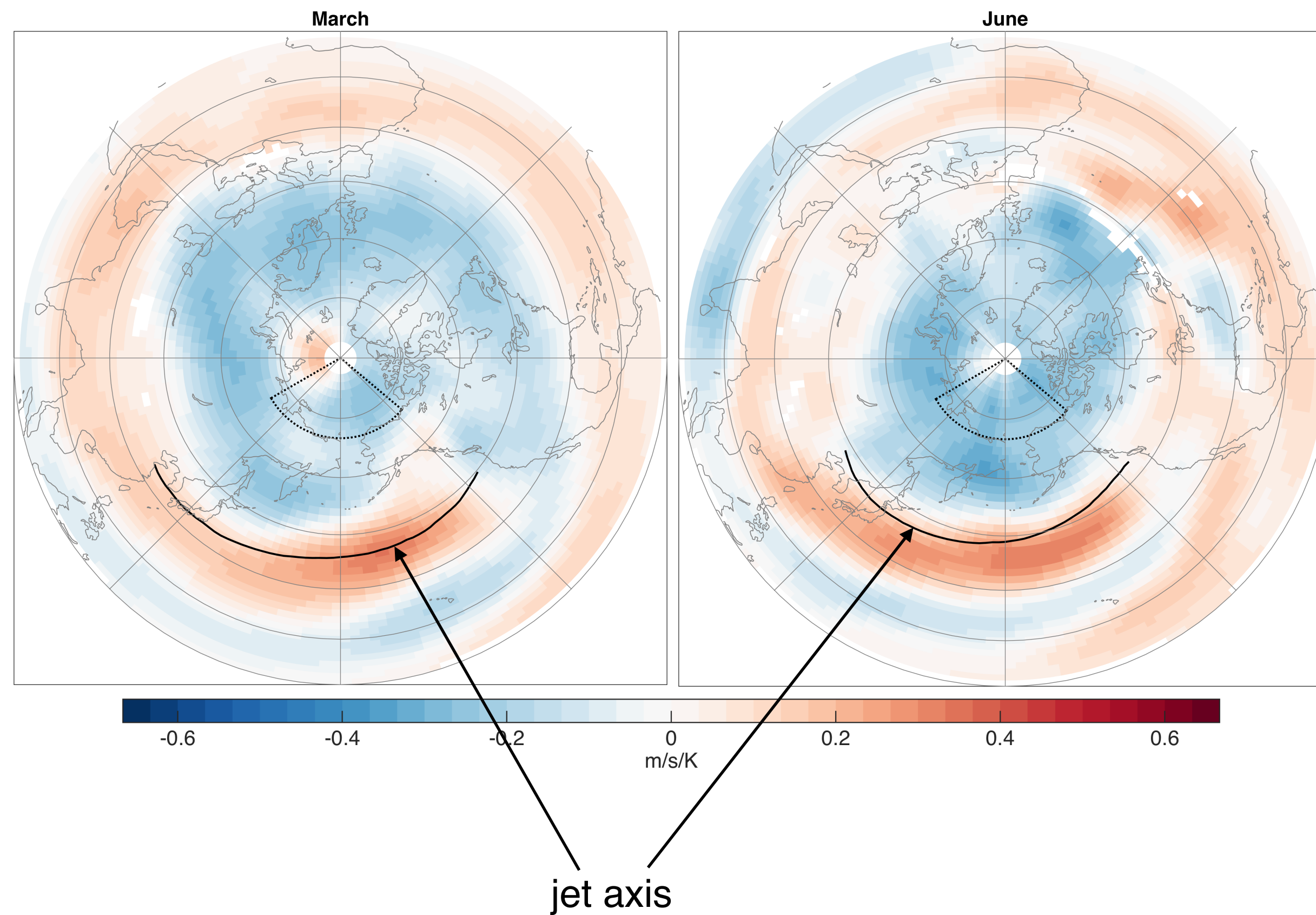


jet strengthens in most months



large seasonality in number of models exhibiting Granger-causality

Explaining the seasonality

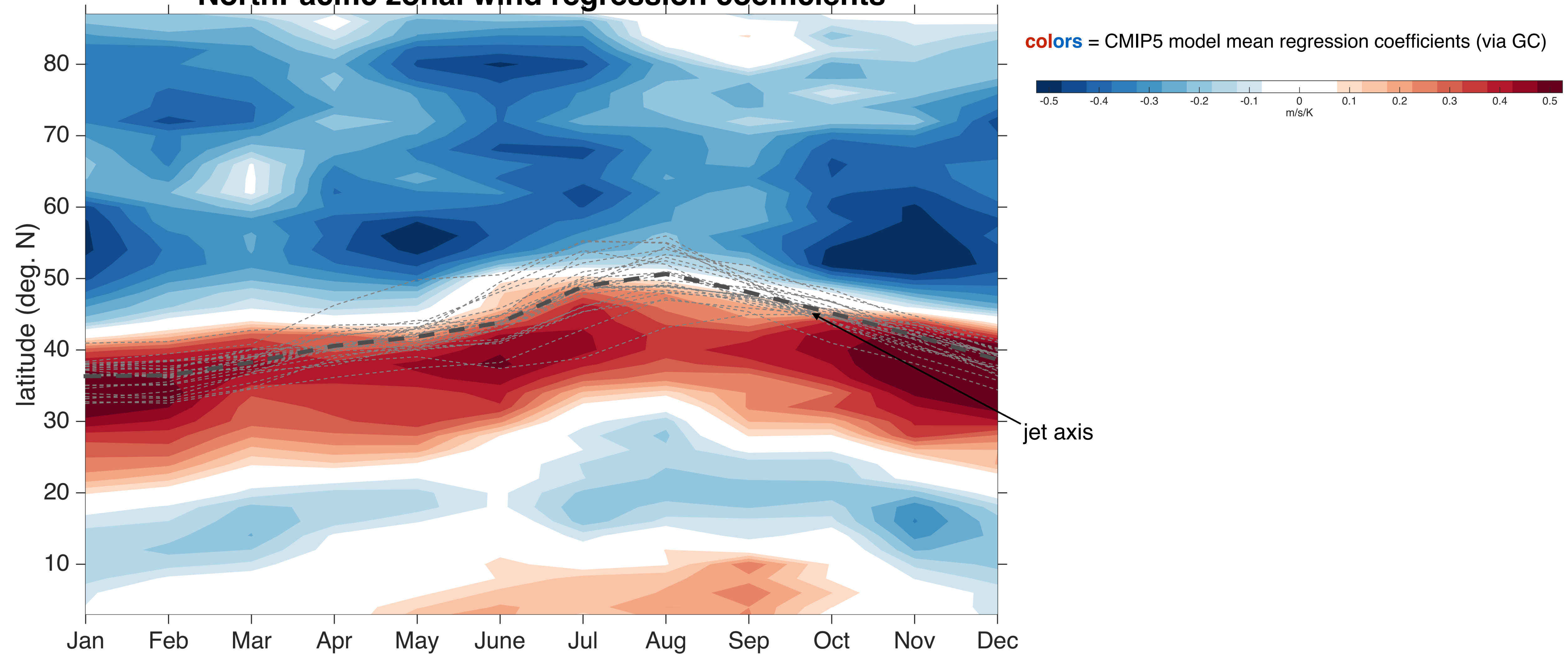


- jet shifts with the seasonal cycle
- wind anomalies remain relatively fixed in latitude

colors = CMIP5 model mean regression coefficients (via Granger-causality)

Explaining the seasonality

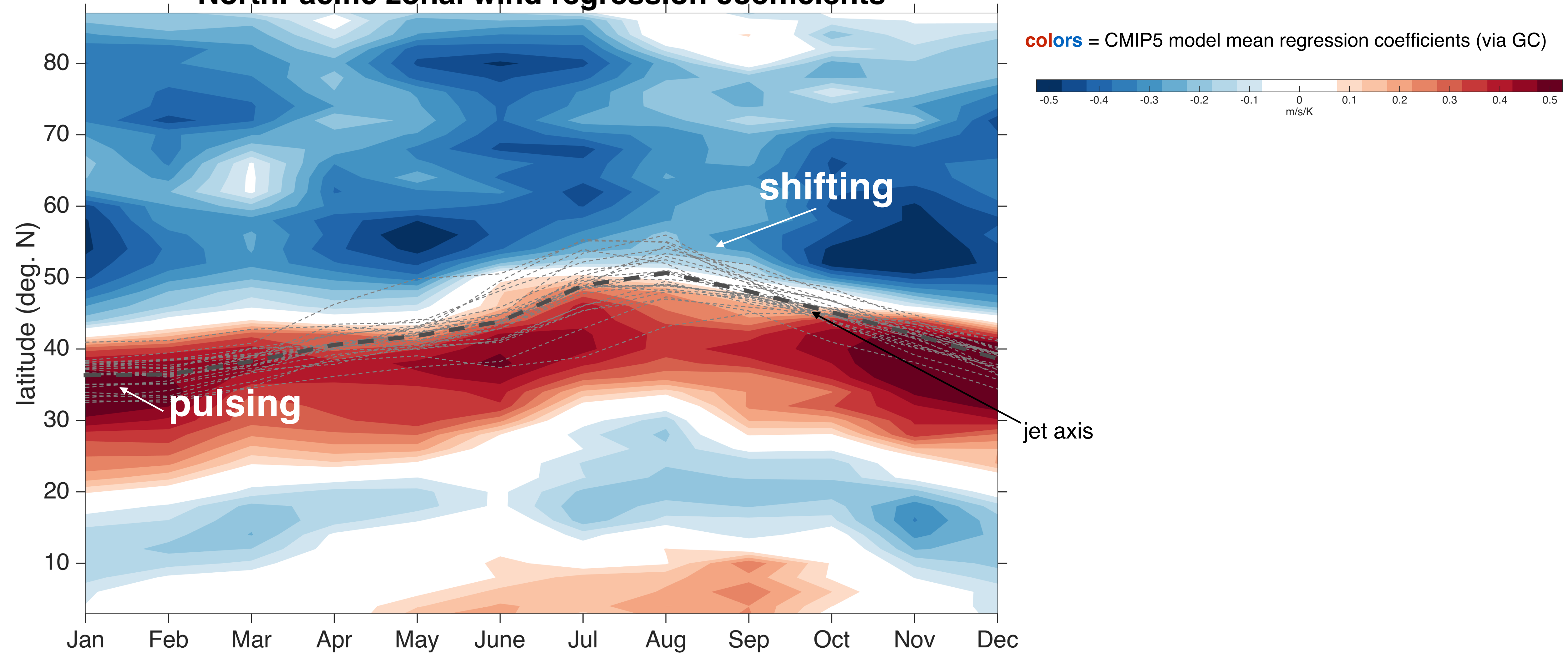
NorthPacific zonal wind regression coefficients



Barnes and Simpson (in prep)

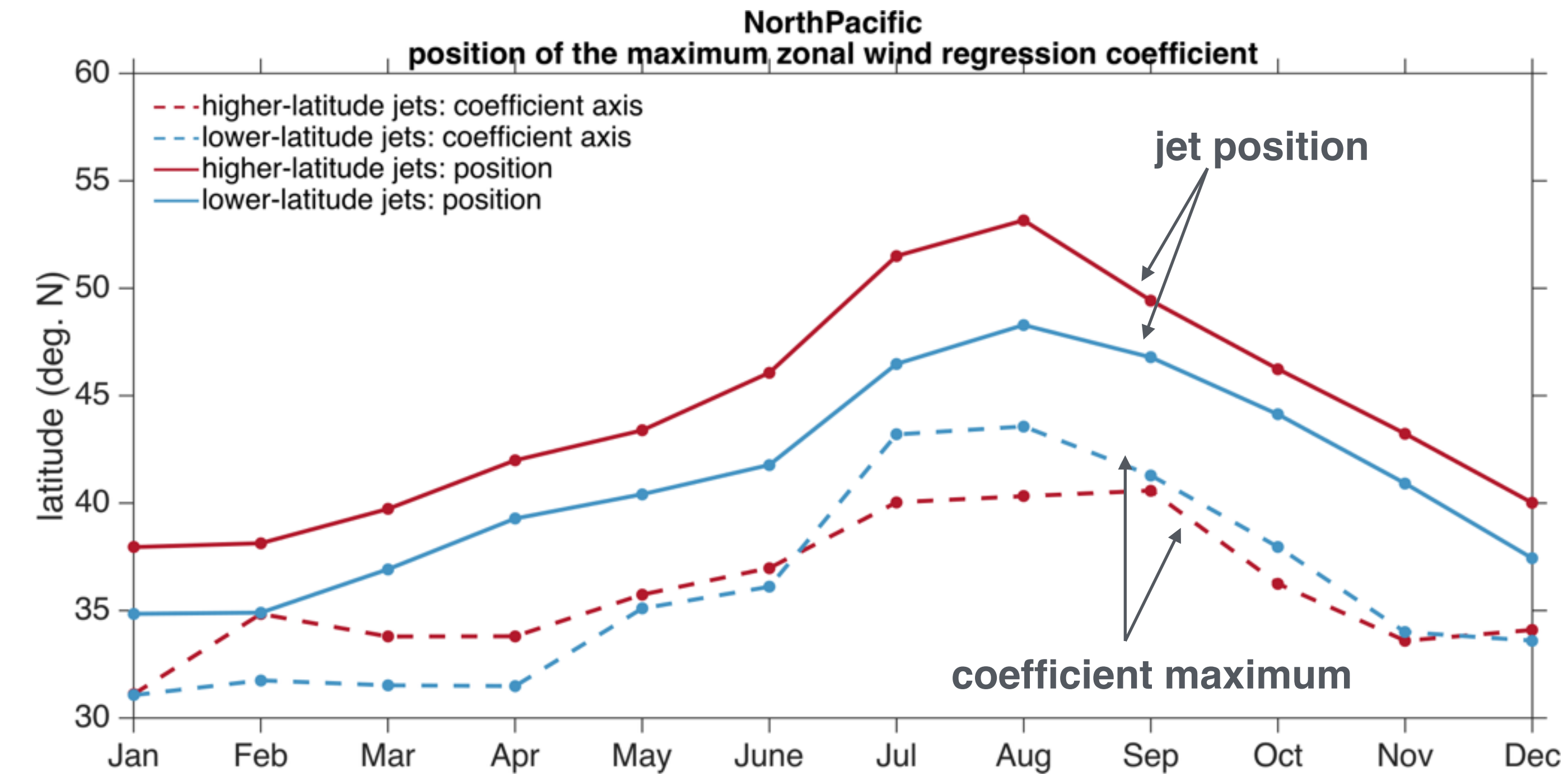
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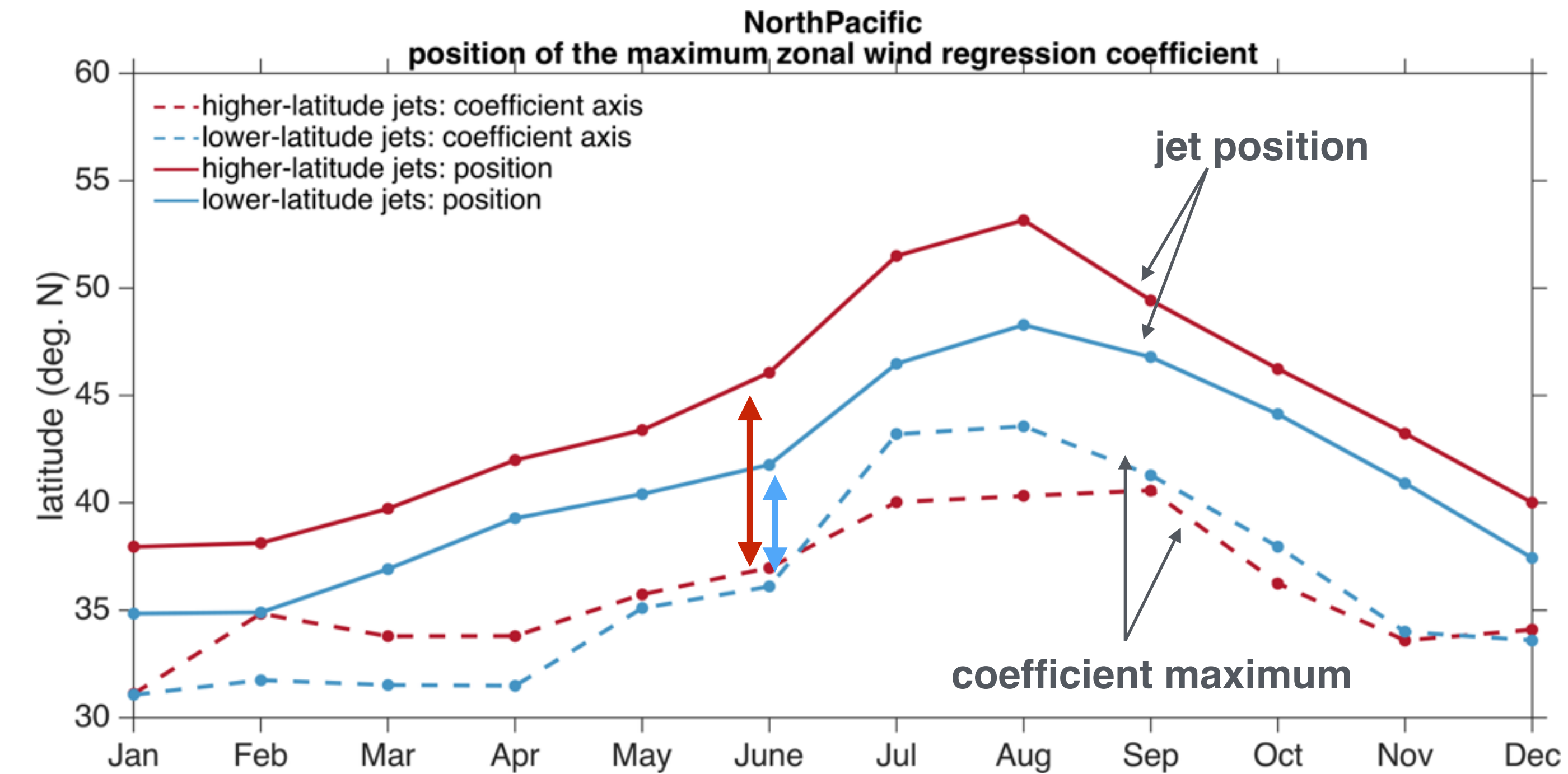


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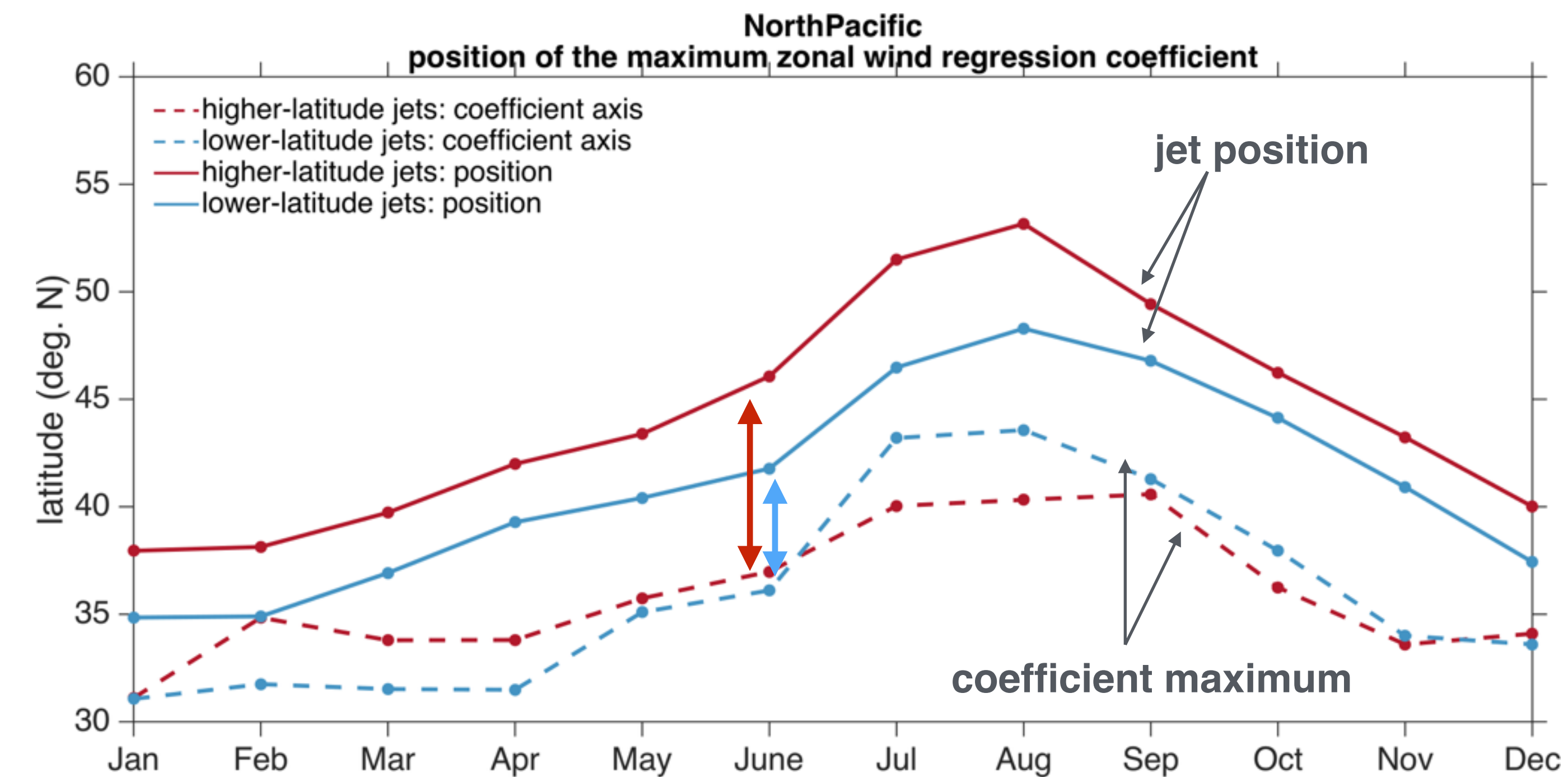
Implications for models biases



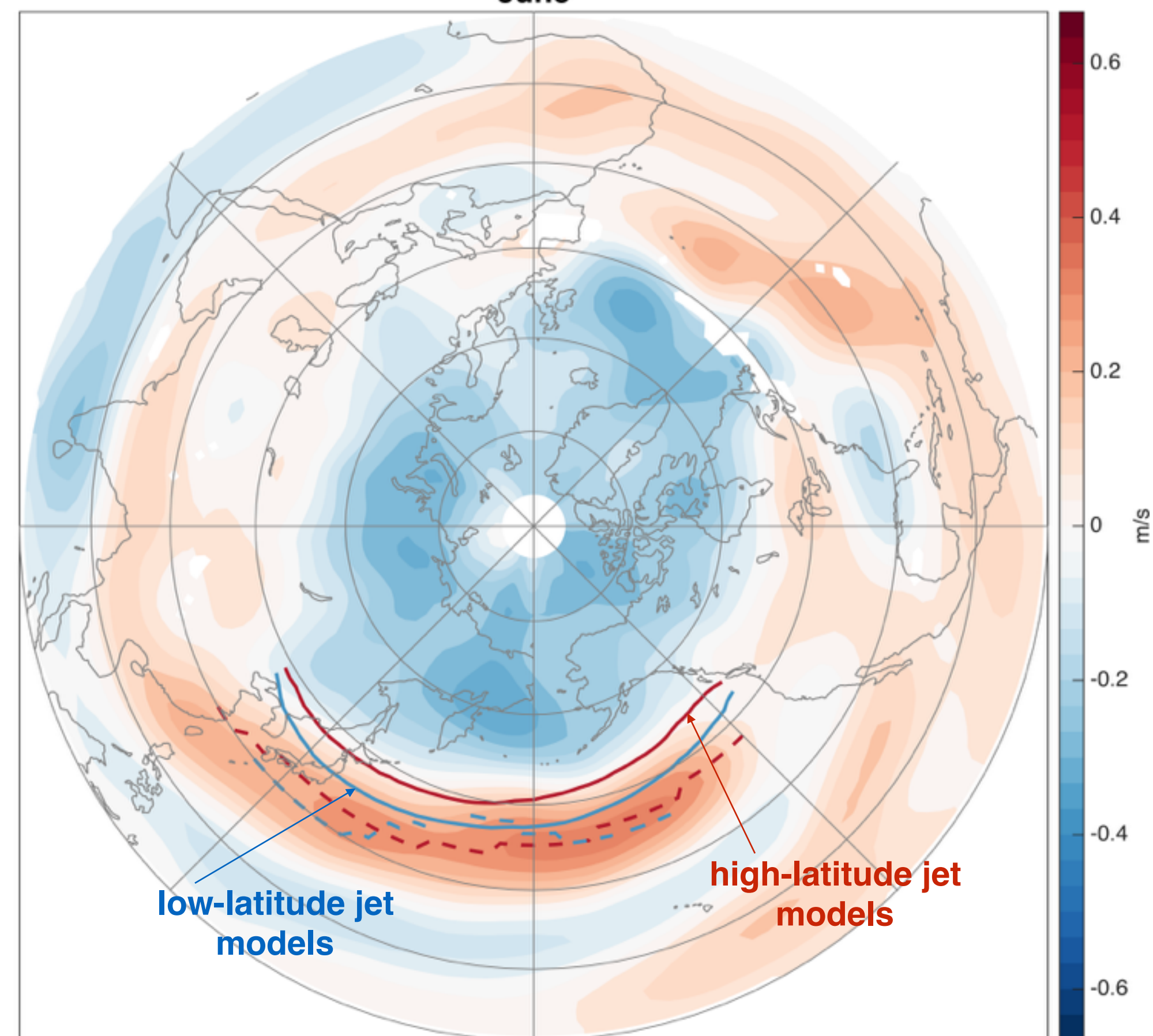
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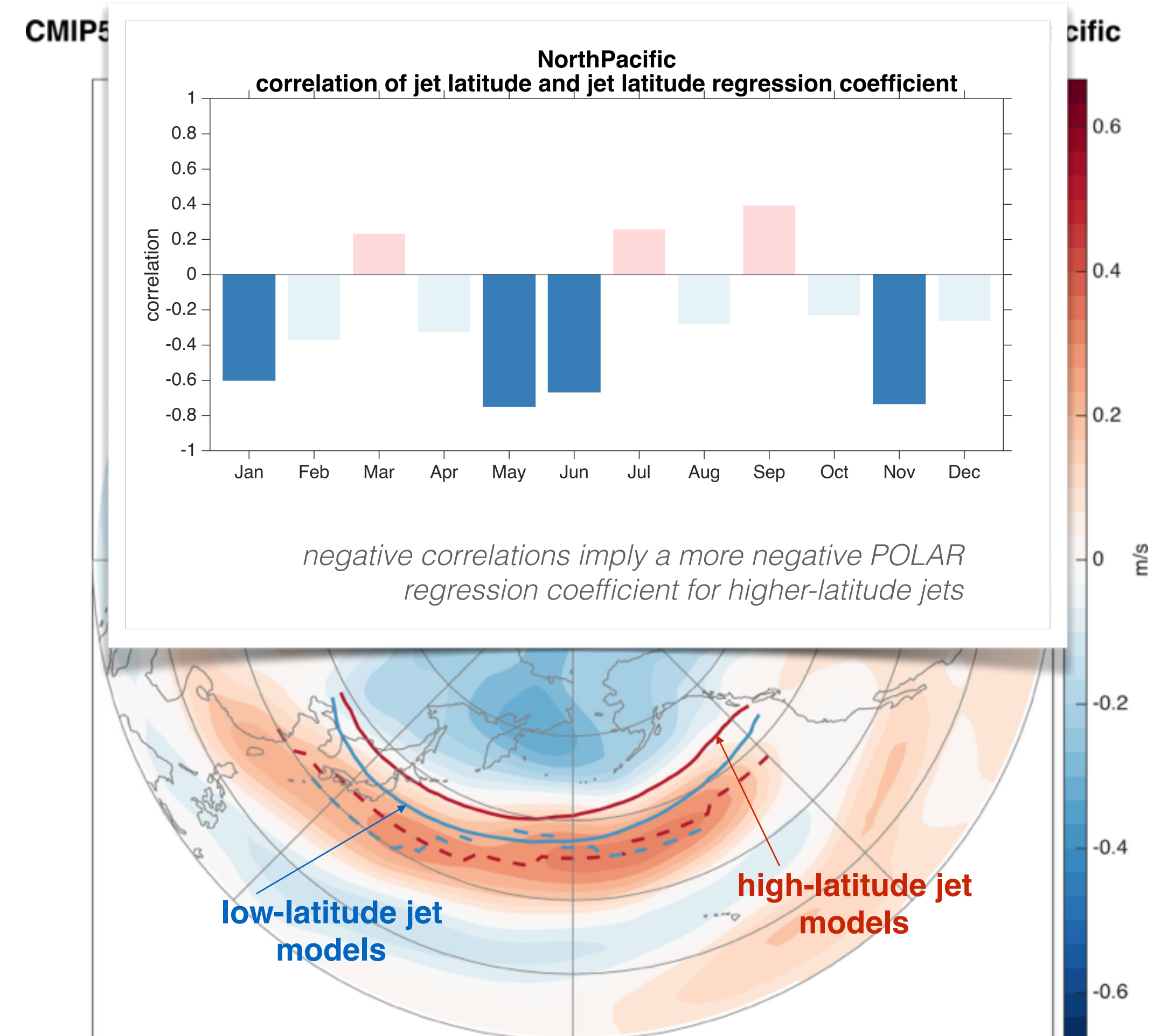
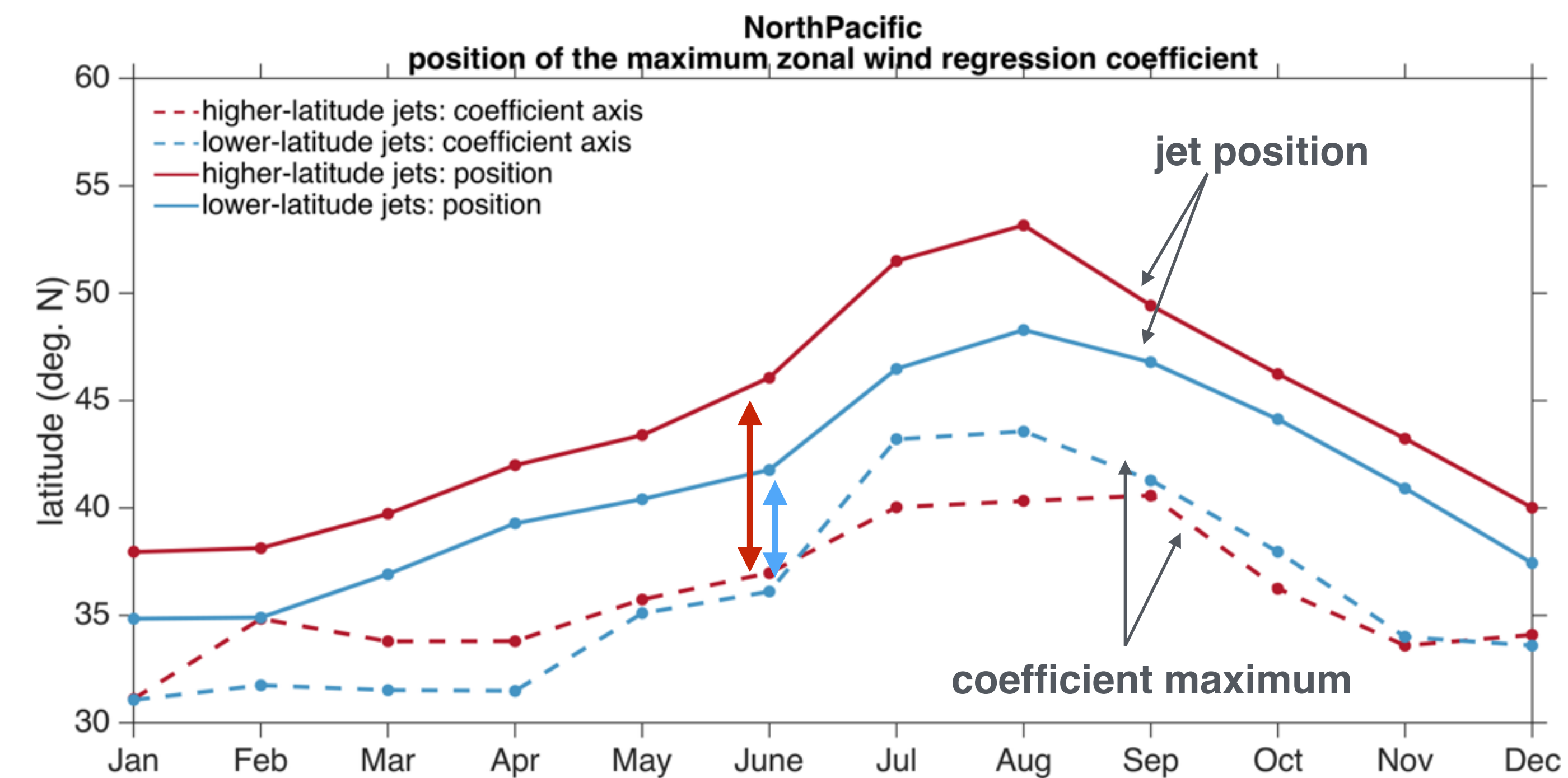


CMIP5 RCP8.5 u700 response to 1K POLE warming for warming in NorthPacific June



Barnes and Simpson (in prep)

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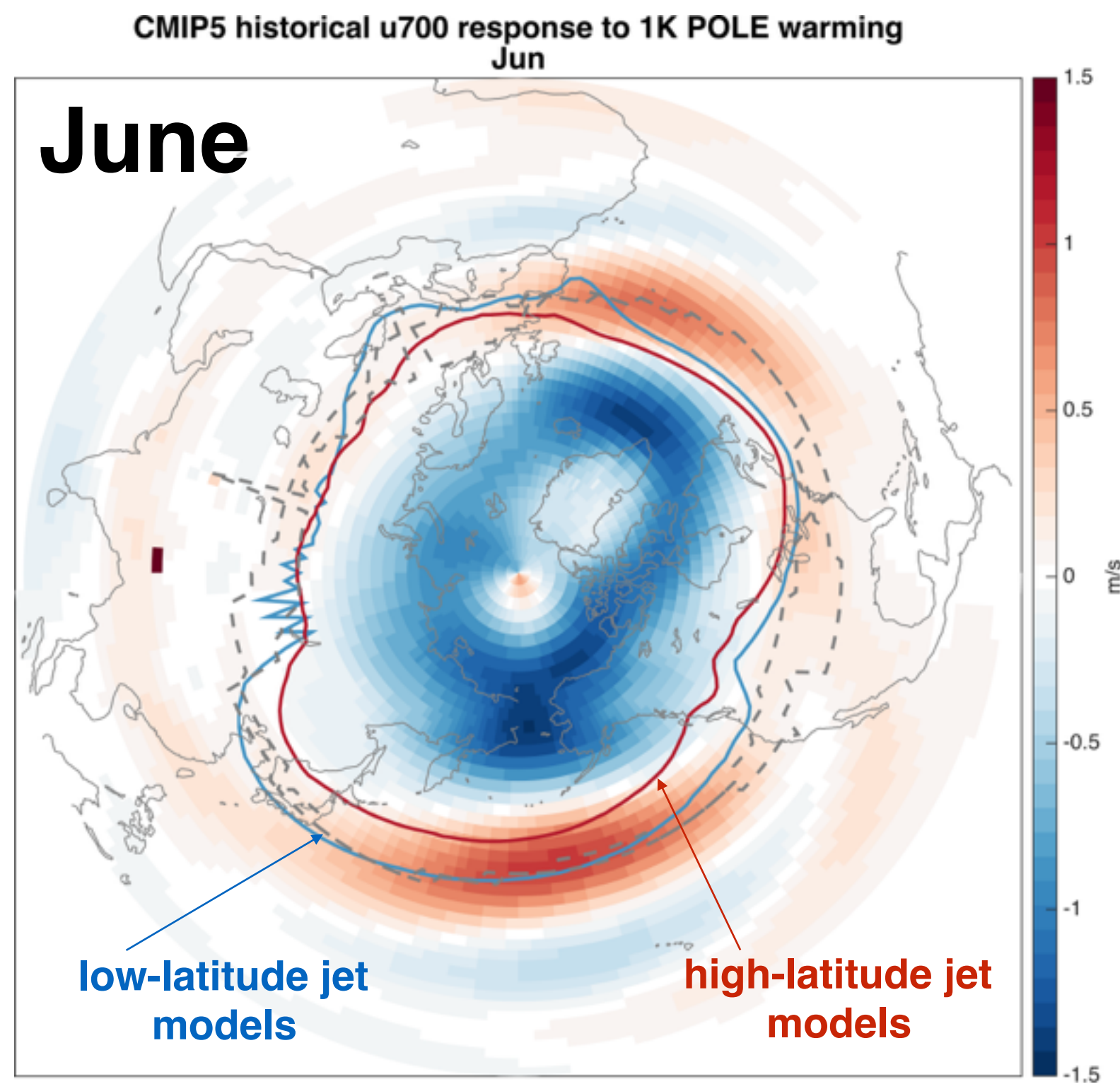


Barnes and Simpson (in prep)



- Granger-causality approach shows that Arctic warming can influence the North Pacific jet latitude and speed
- The sensitivity of the jet-stream to ~weekly variations in Arctic temperatures varies as a function of season
 - ▶ e.g. the North Pacific jet latitude is most sensitive to Arctic warming in the summer months
- The seasonality of the jet position sensitivity to Arctic warming can be understood by the jet shifting in and out of the anomalies. This is shown to have implications for jet-stream biases in the CMIP5 models.

Implications for models biases (monthly, lag = 0 analysis)

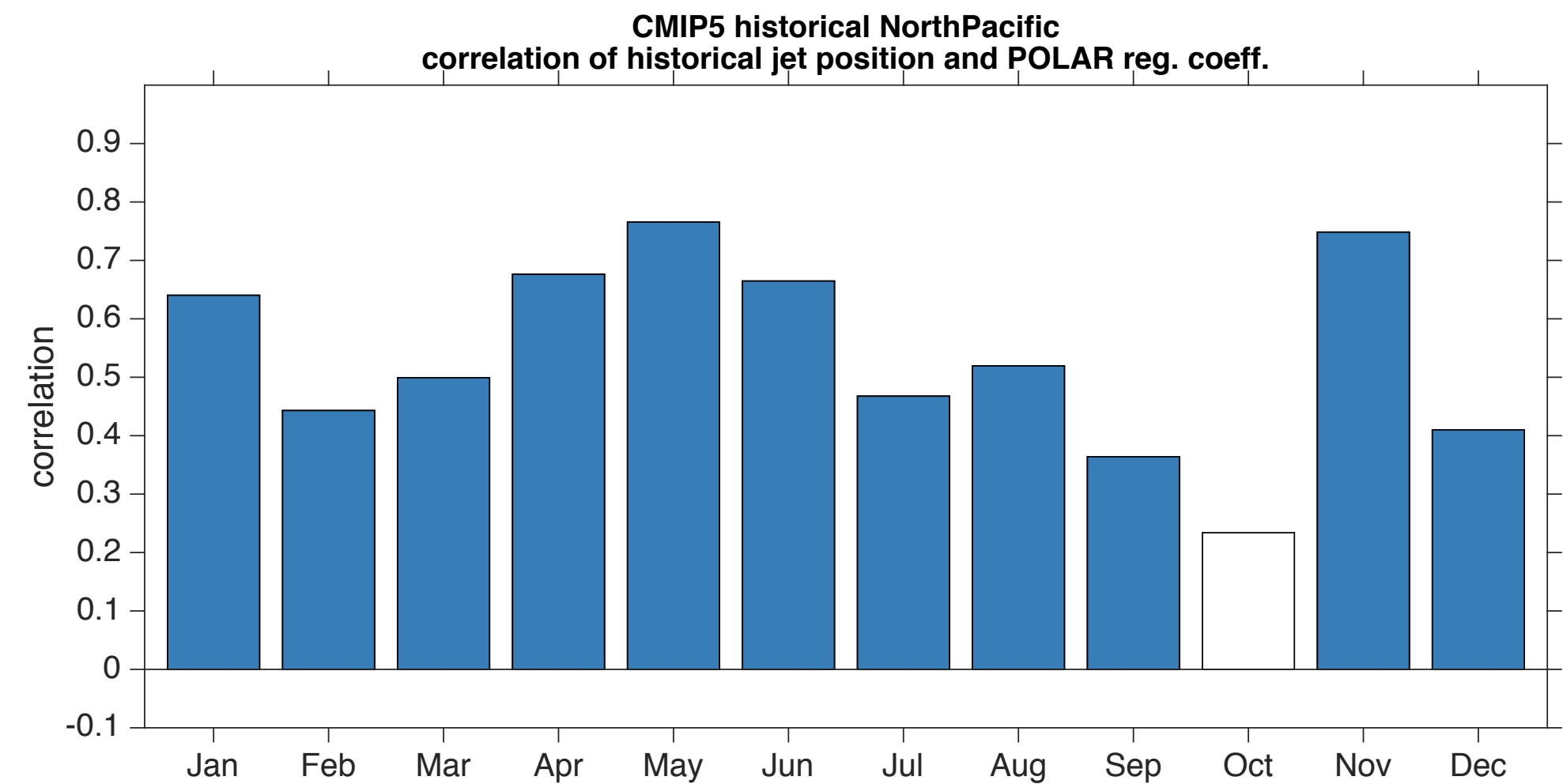
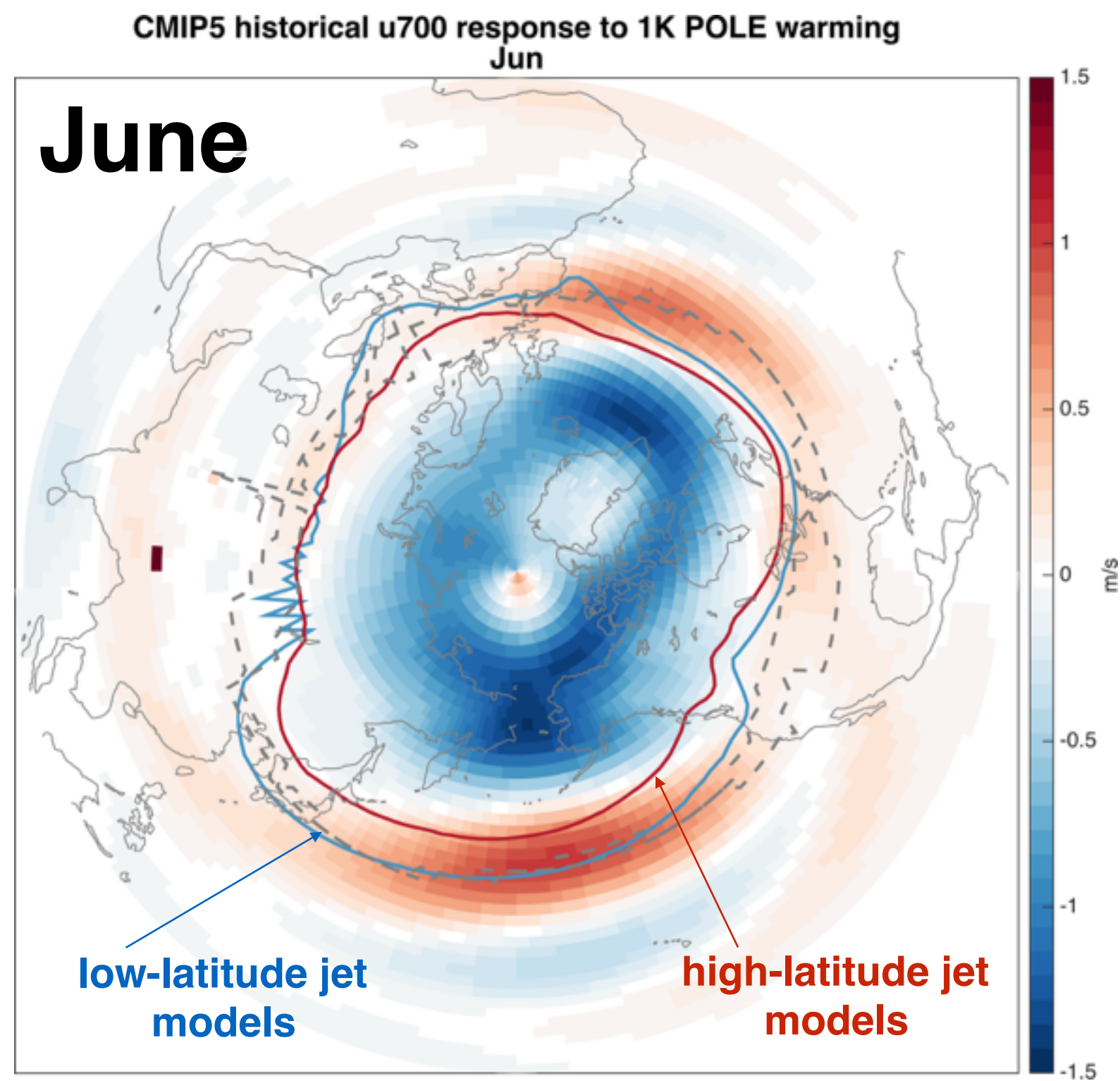


25% most equatorward CMIP5 jet-streams

25% most poleward CMIP5 jet-streams

latitude of maximum regression coefficient

Implications for models biases (monthly, lag = 0 analysis)

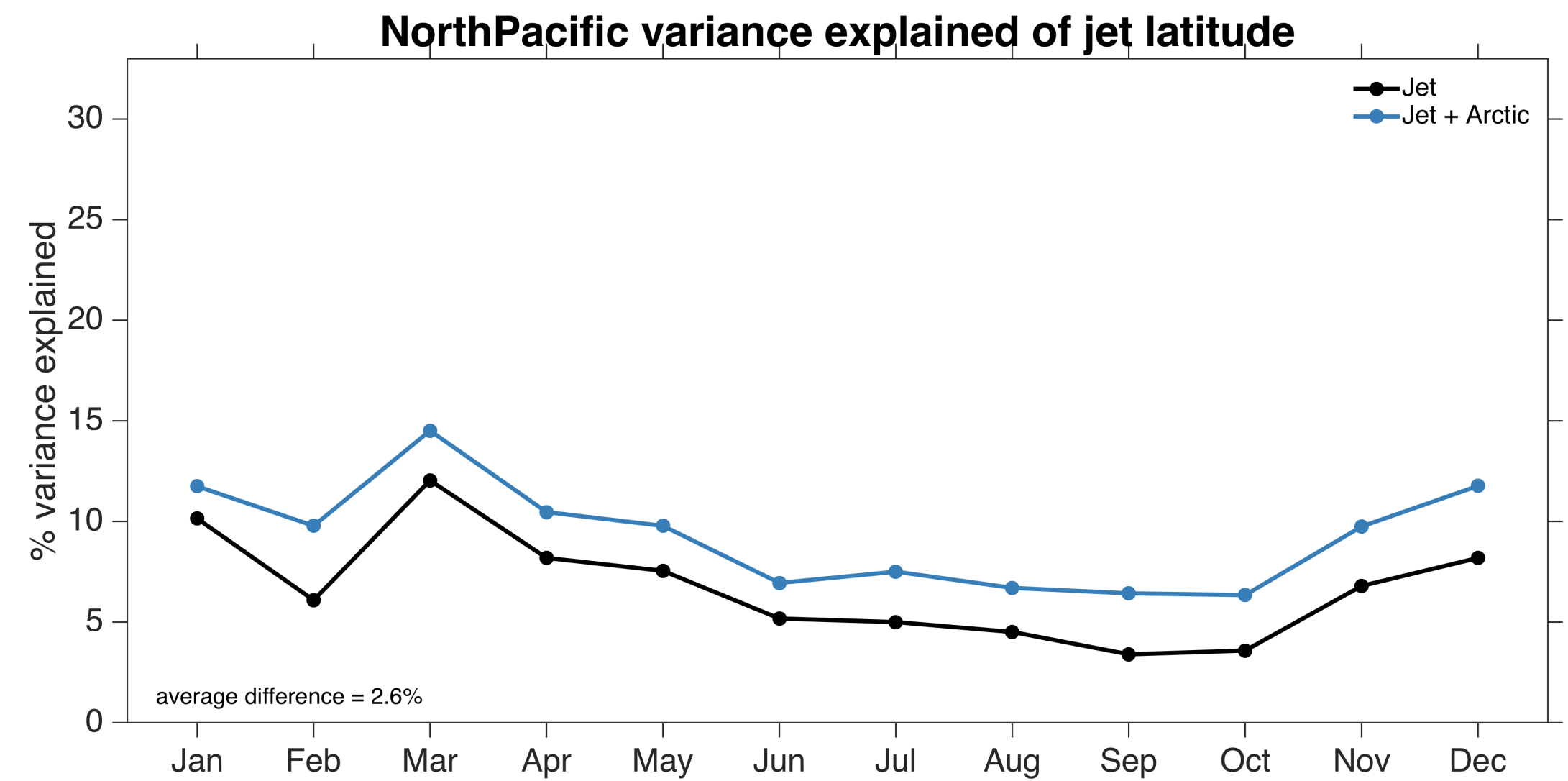
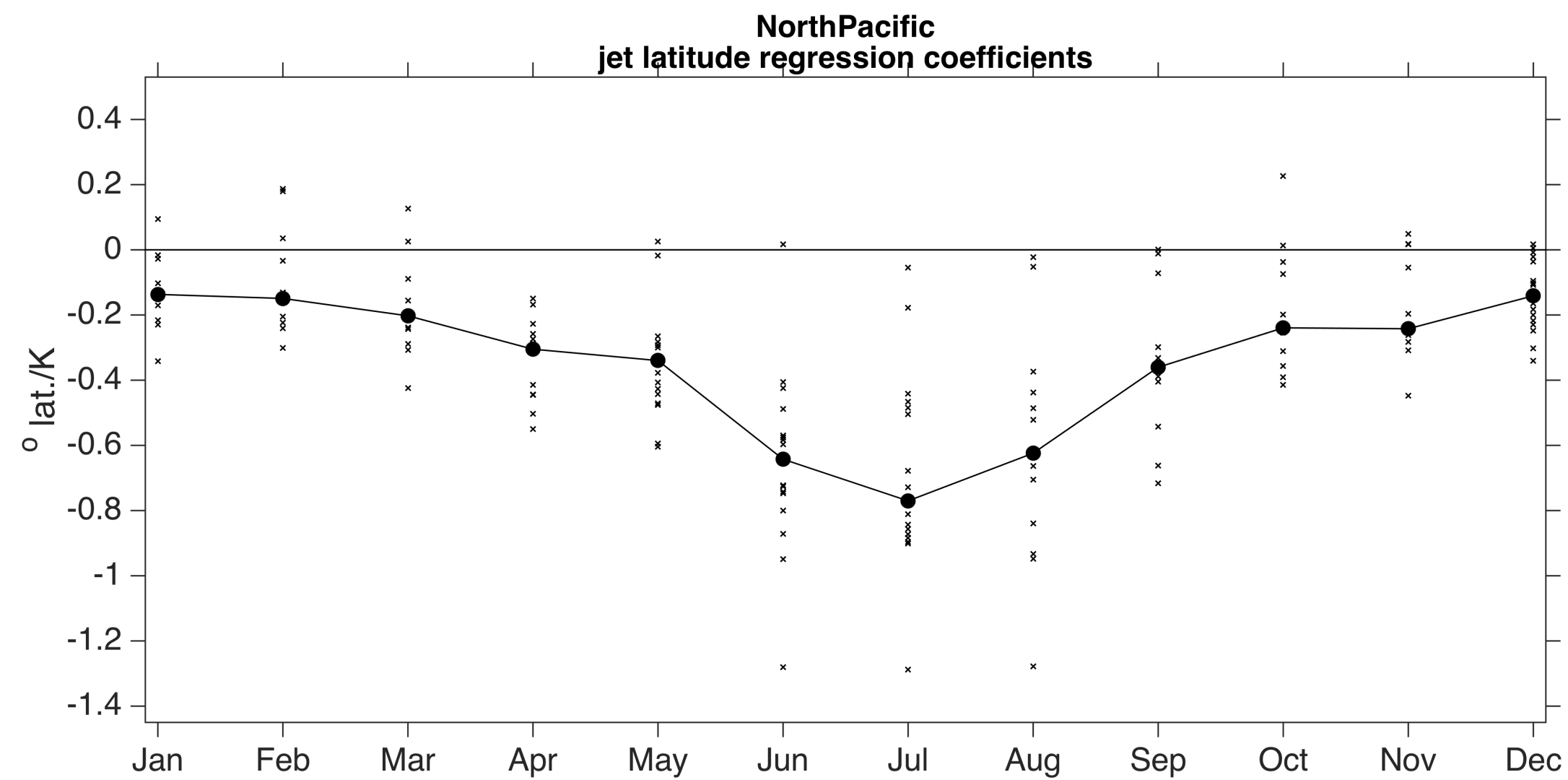


positive correlations imply a more negative POLAR regression coefficient for higher-latitude jets

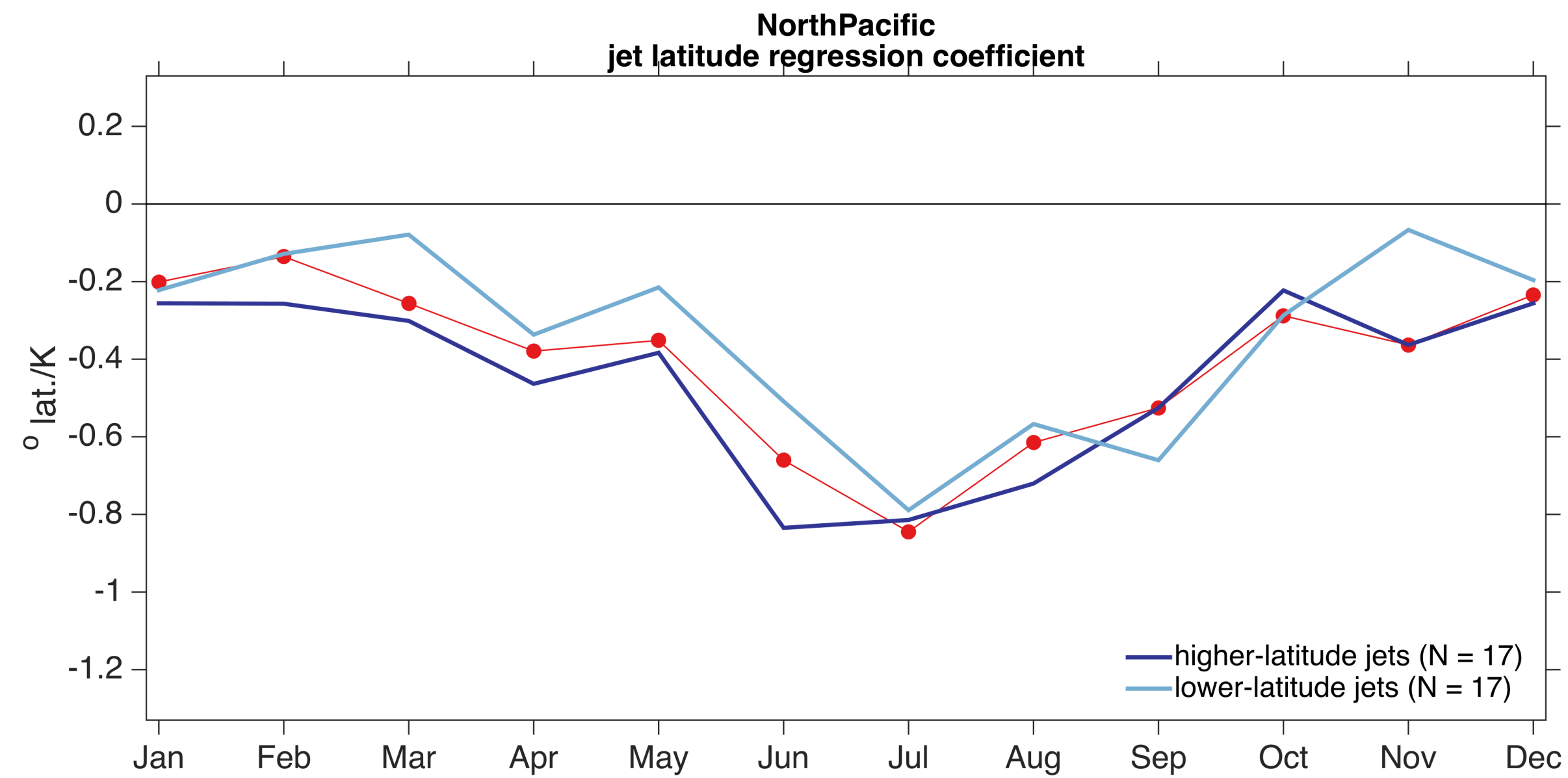
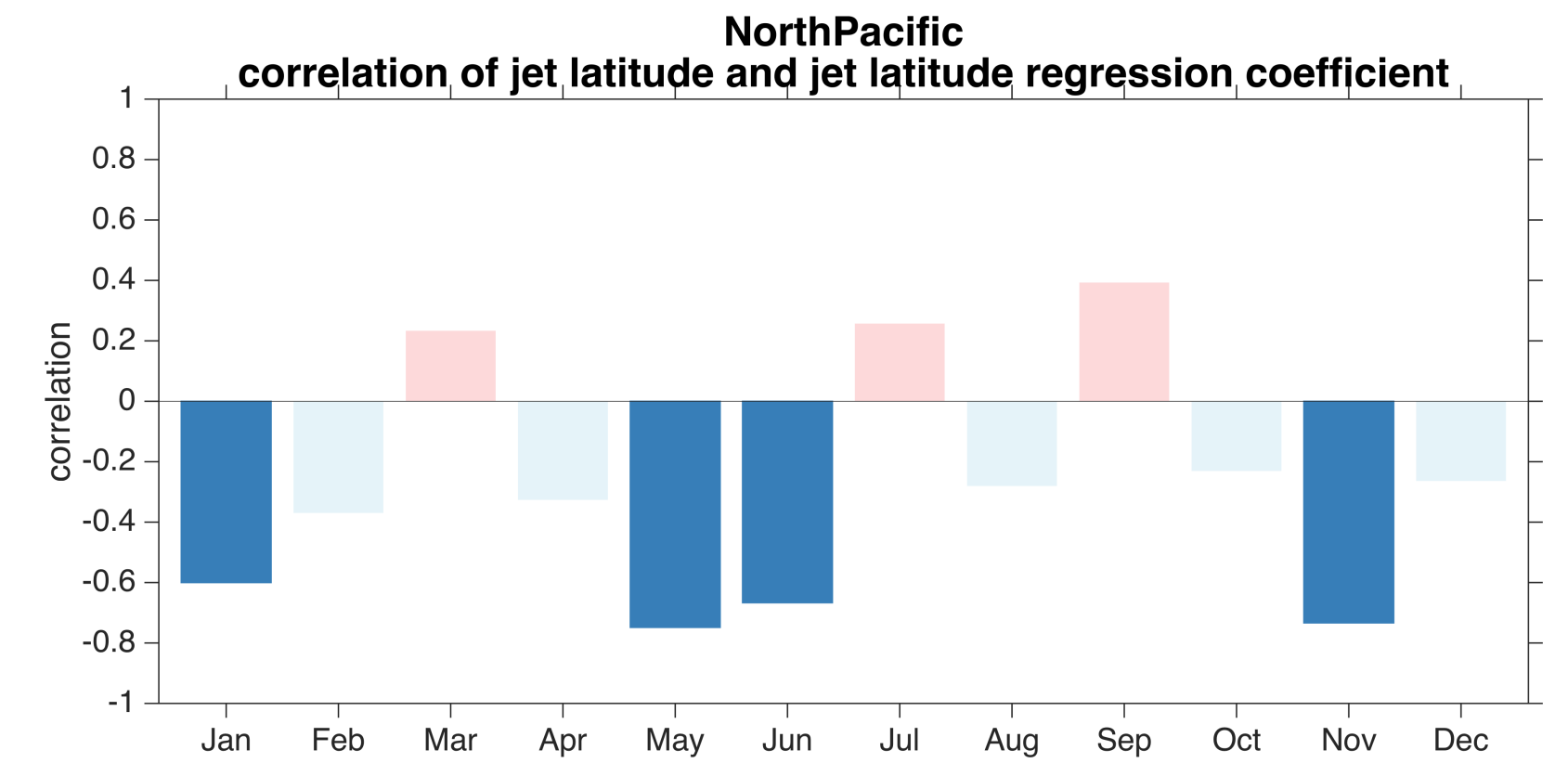
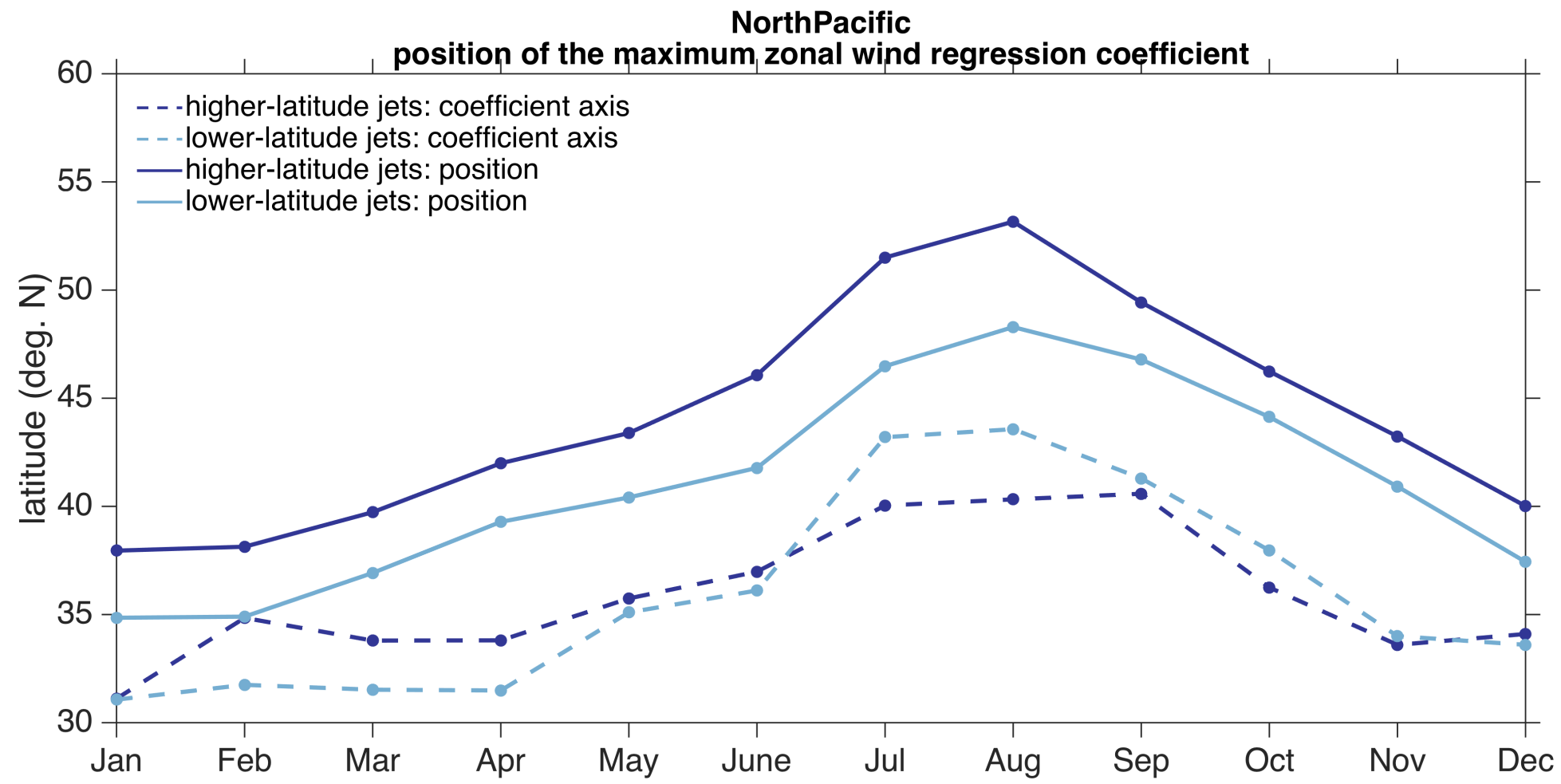
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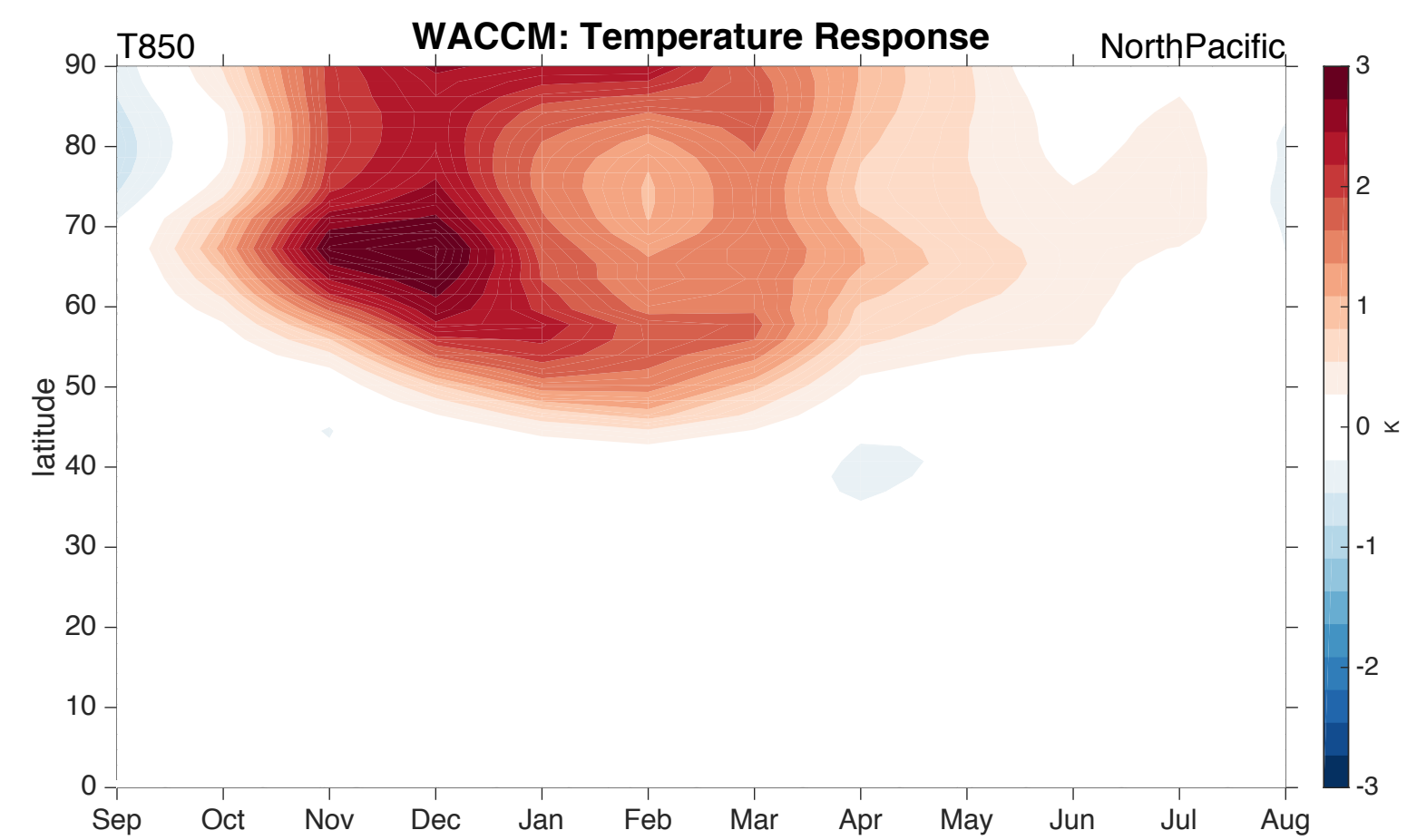
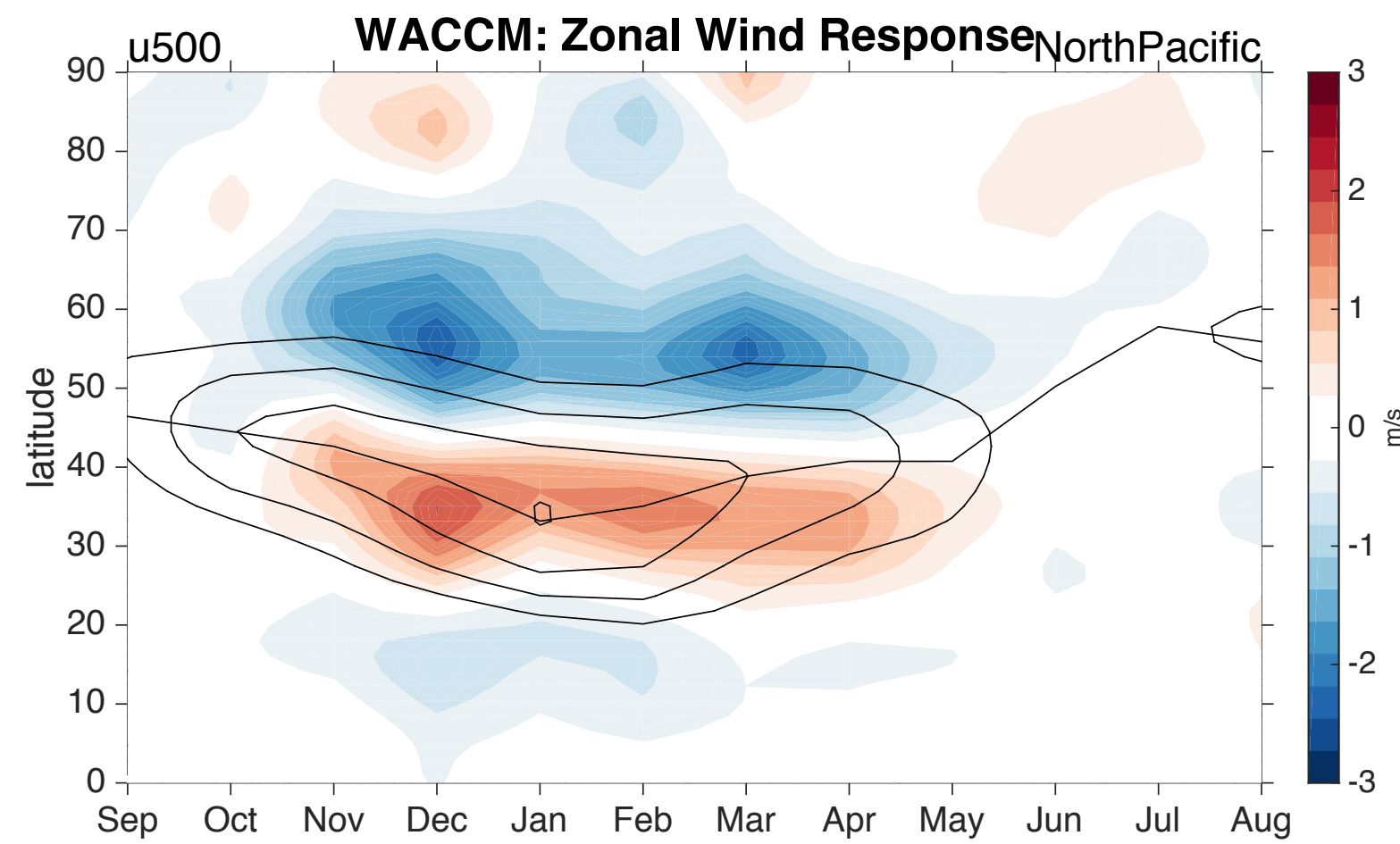
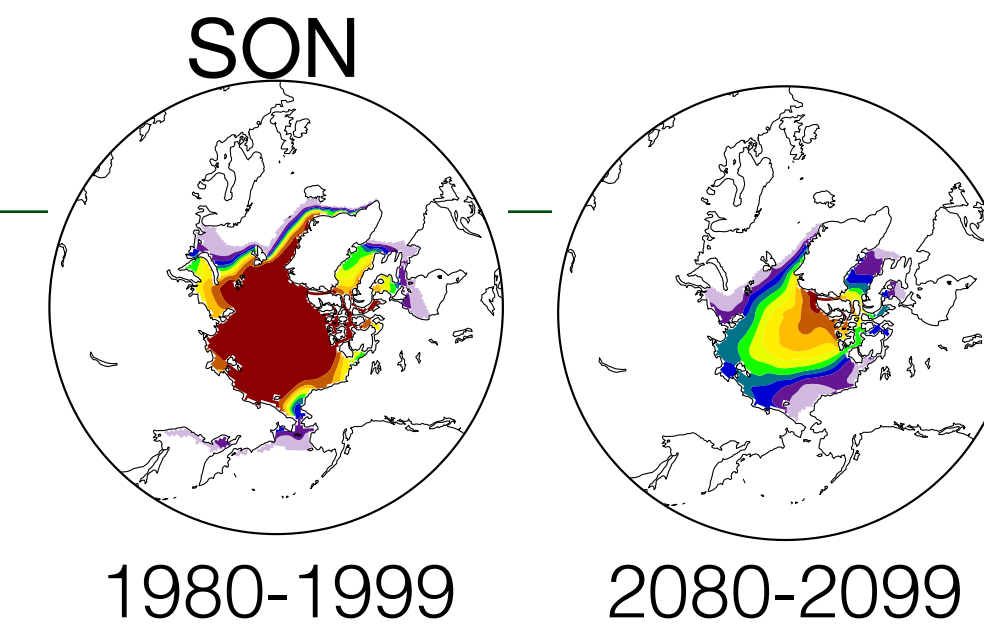
Seasonal sensitivity: jet position



Implications for models biases

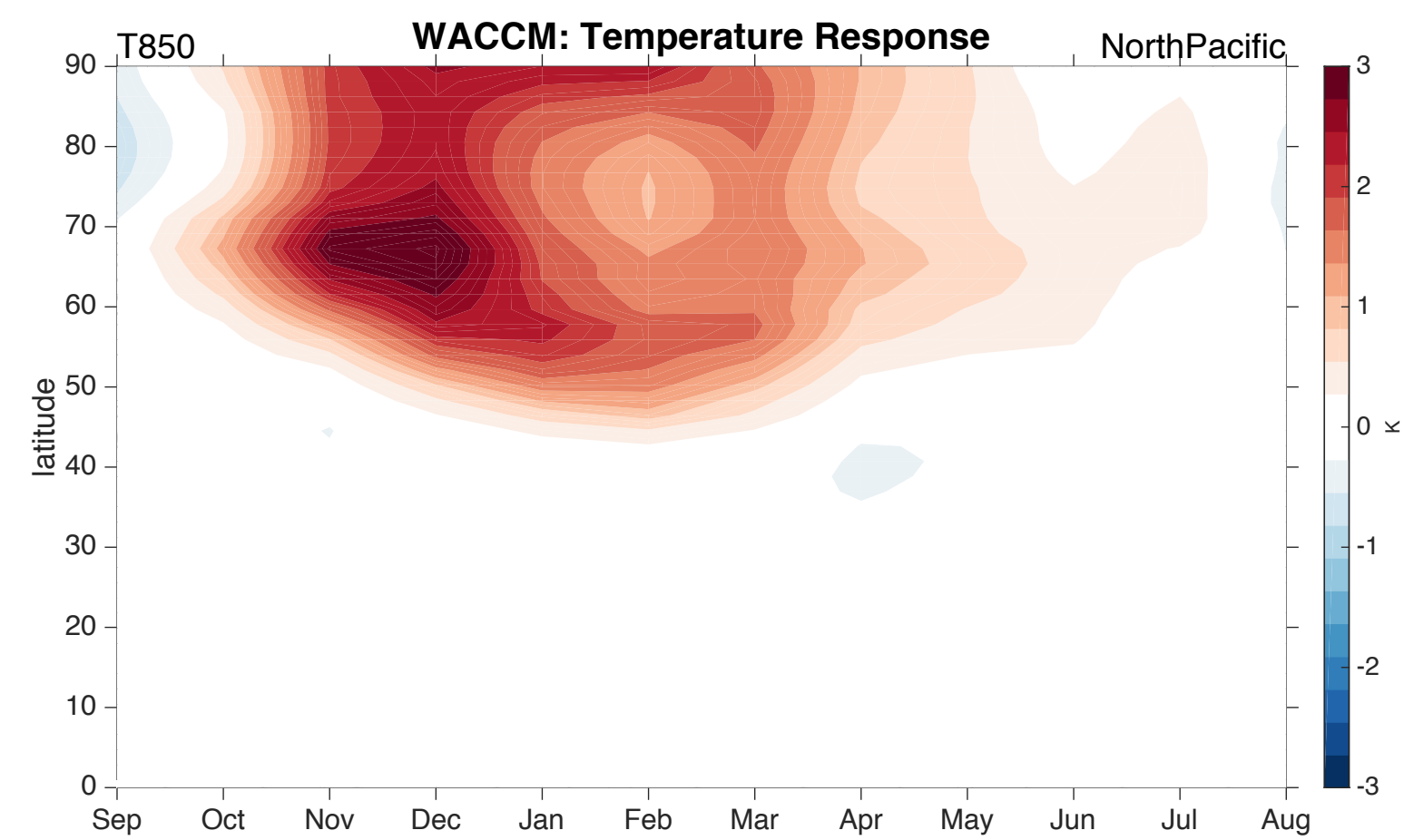
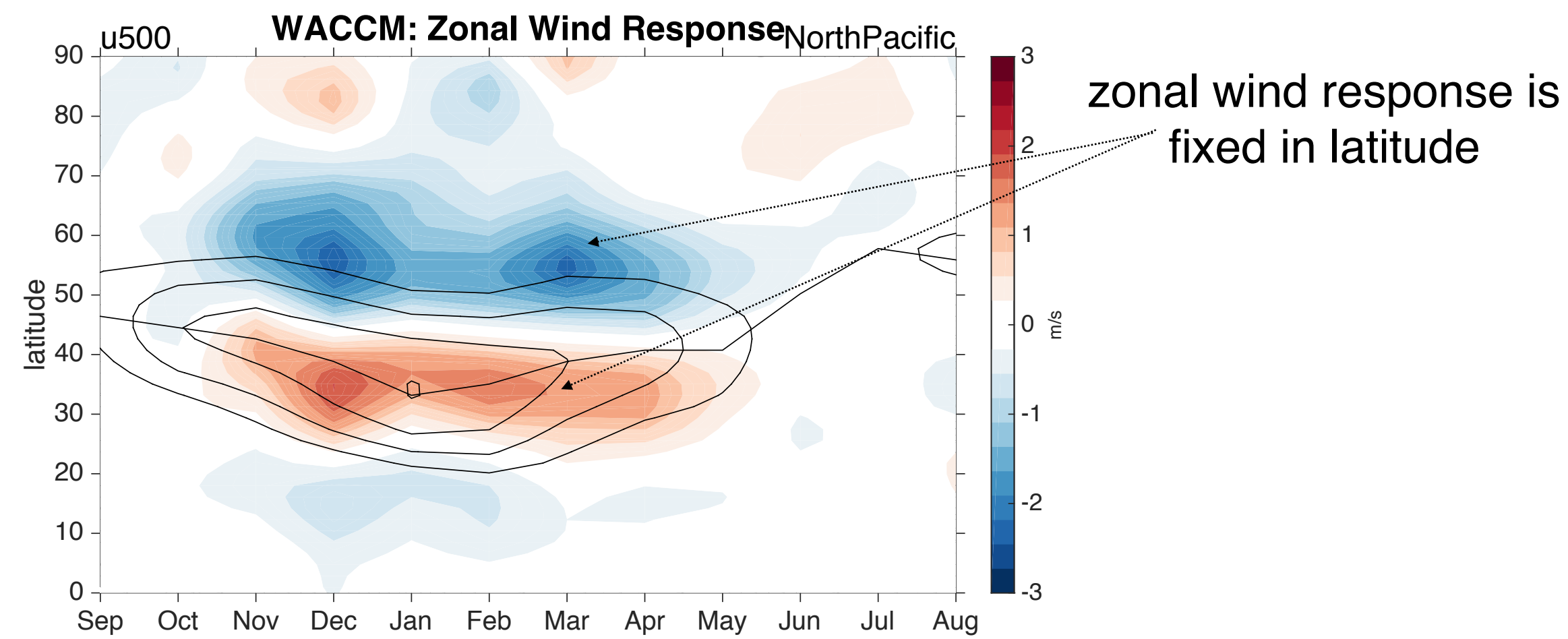
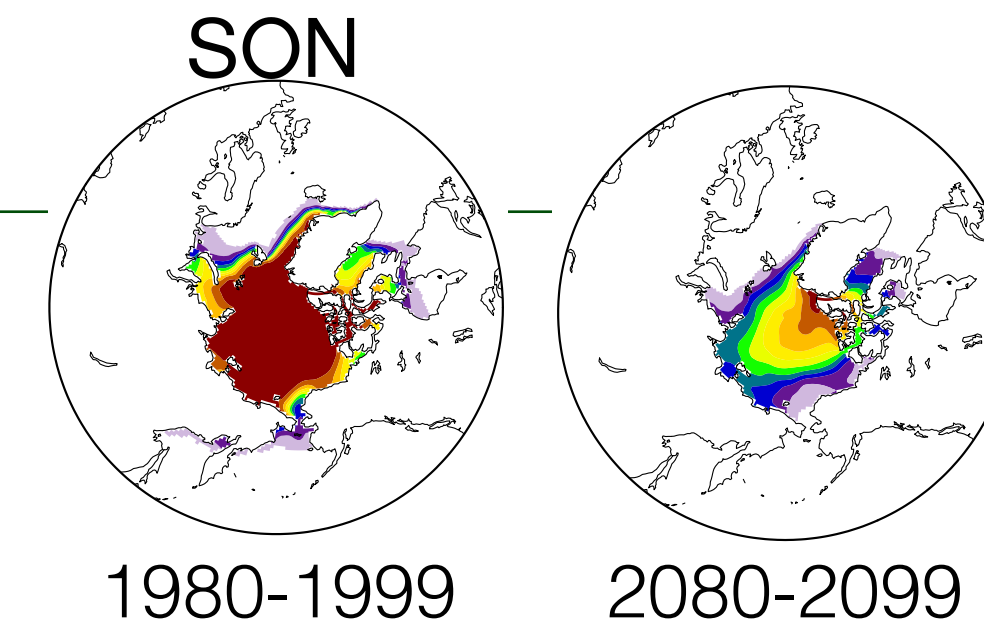


Jet response to sea ice loss



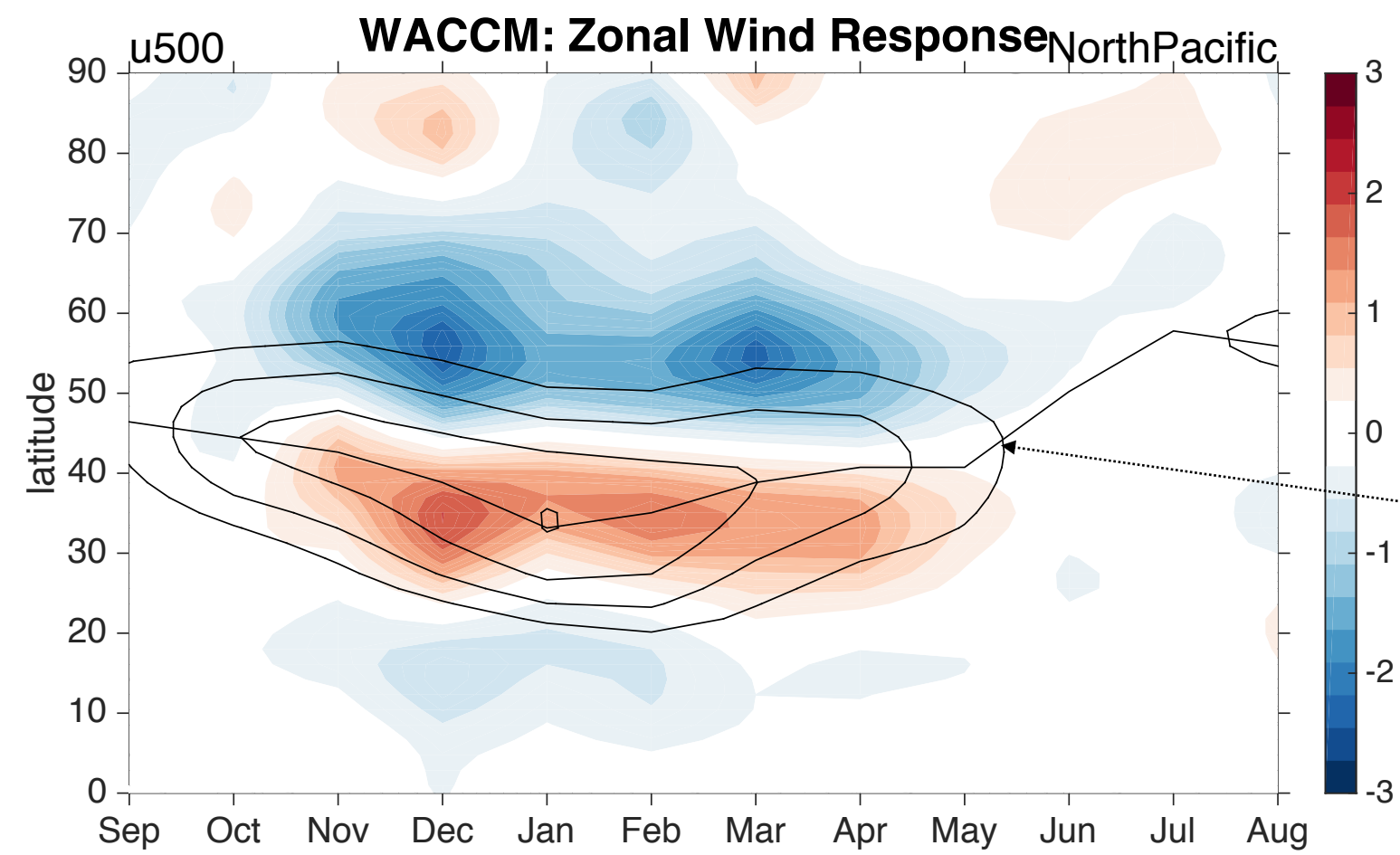
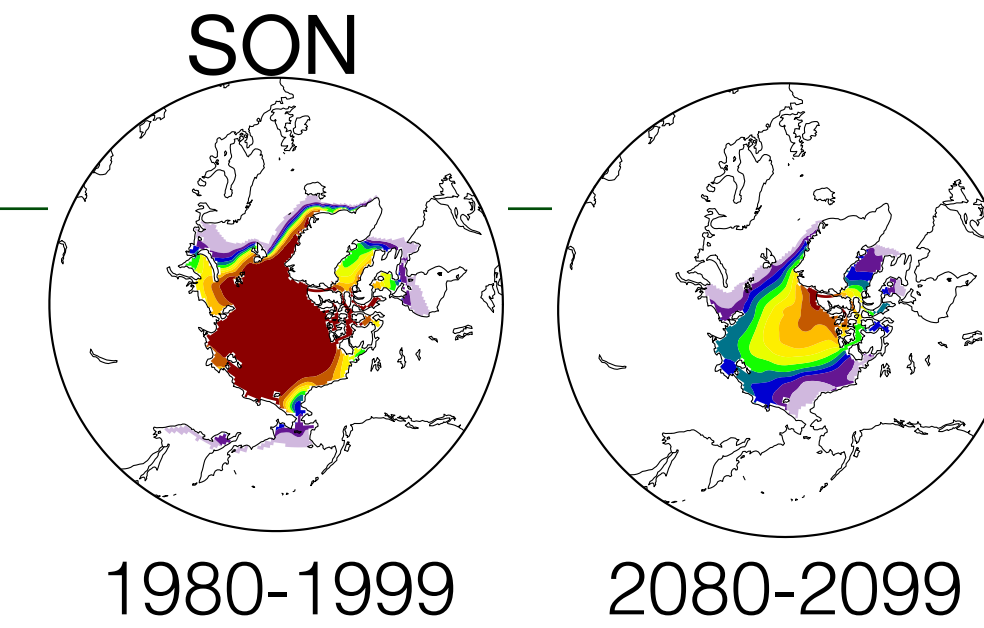
*based on 160-year WACCM simulations of
Sun et al. (2015; JCLI)*

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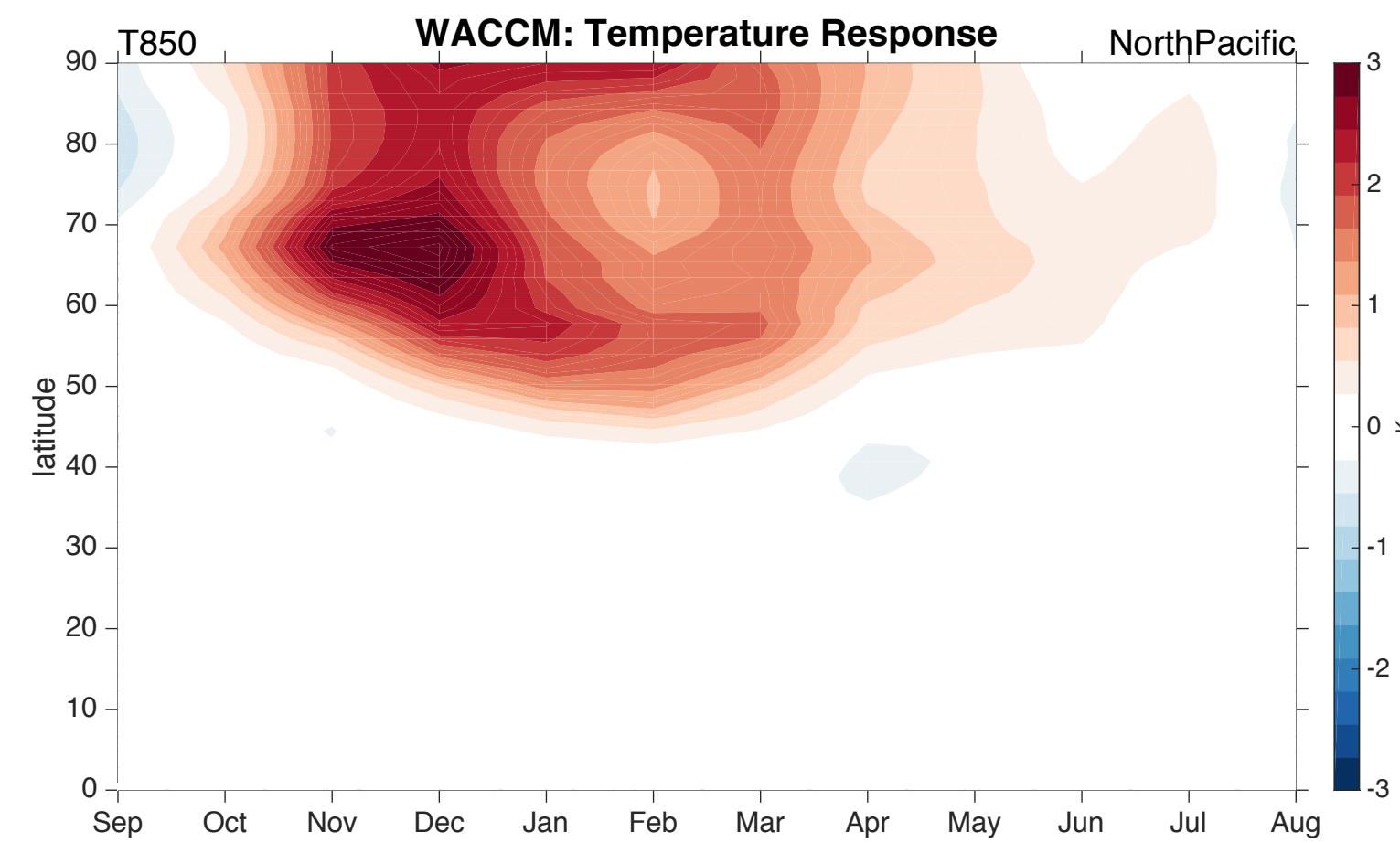


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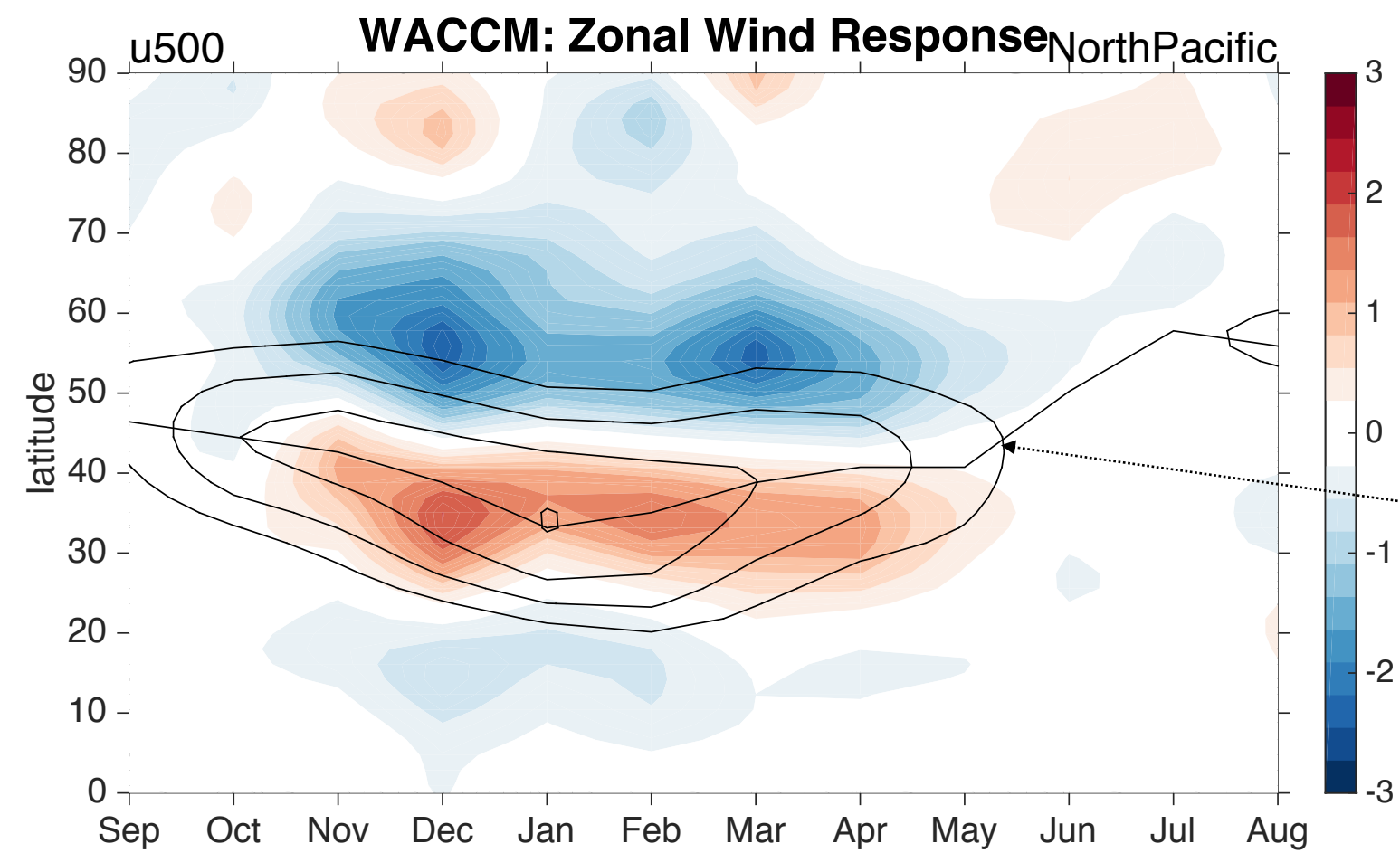
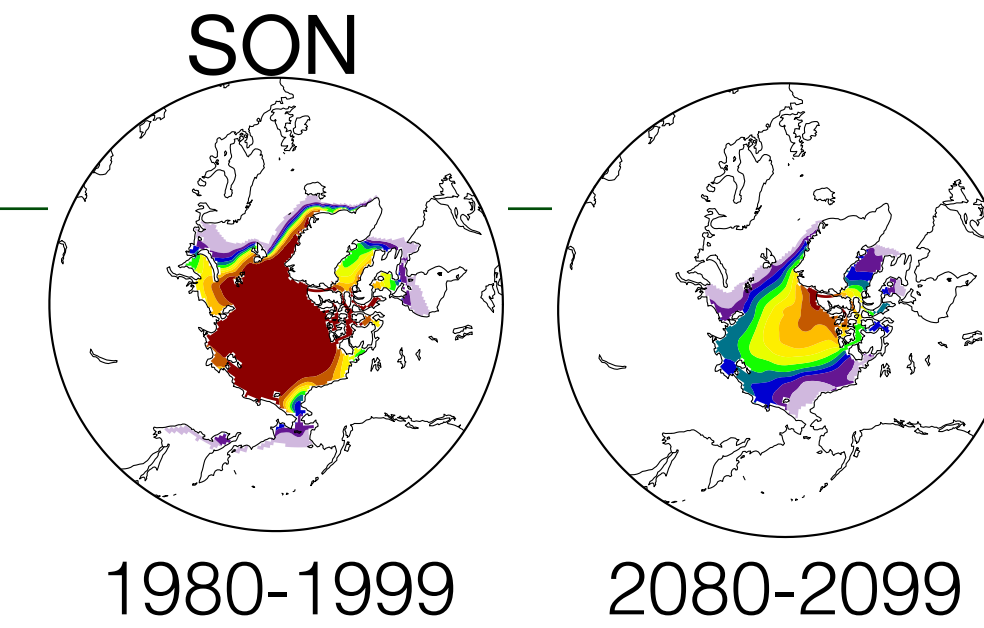


seasonality of jet shifts in-and-out of the sea-ice forced anomalies

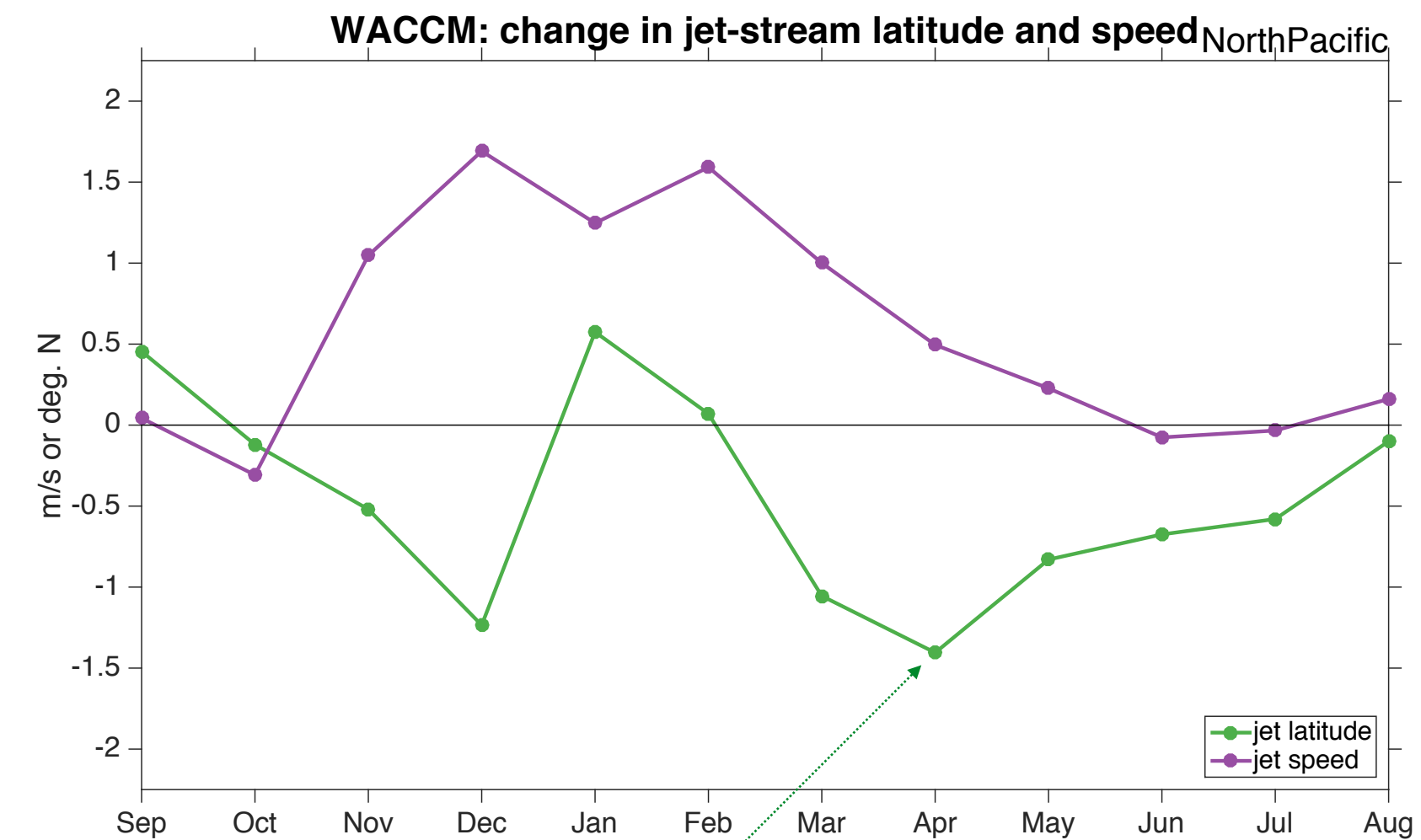
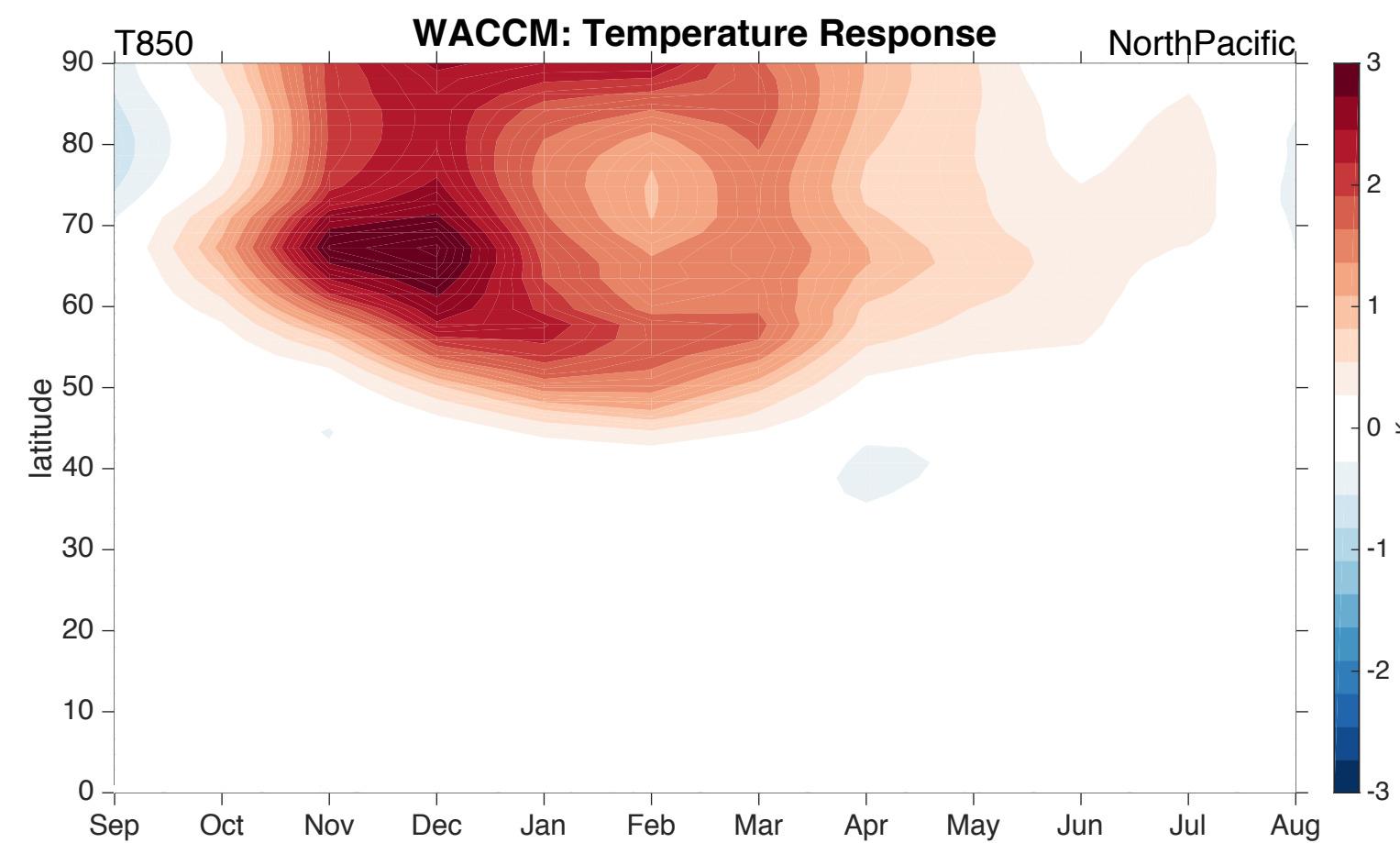


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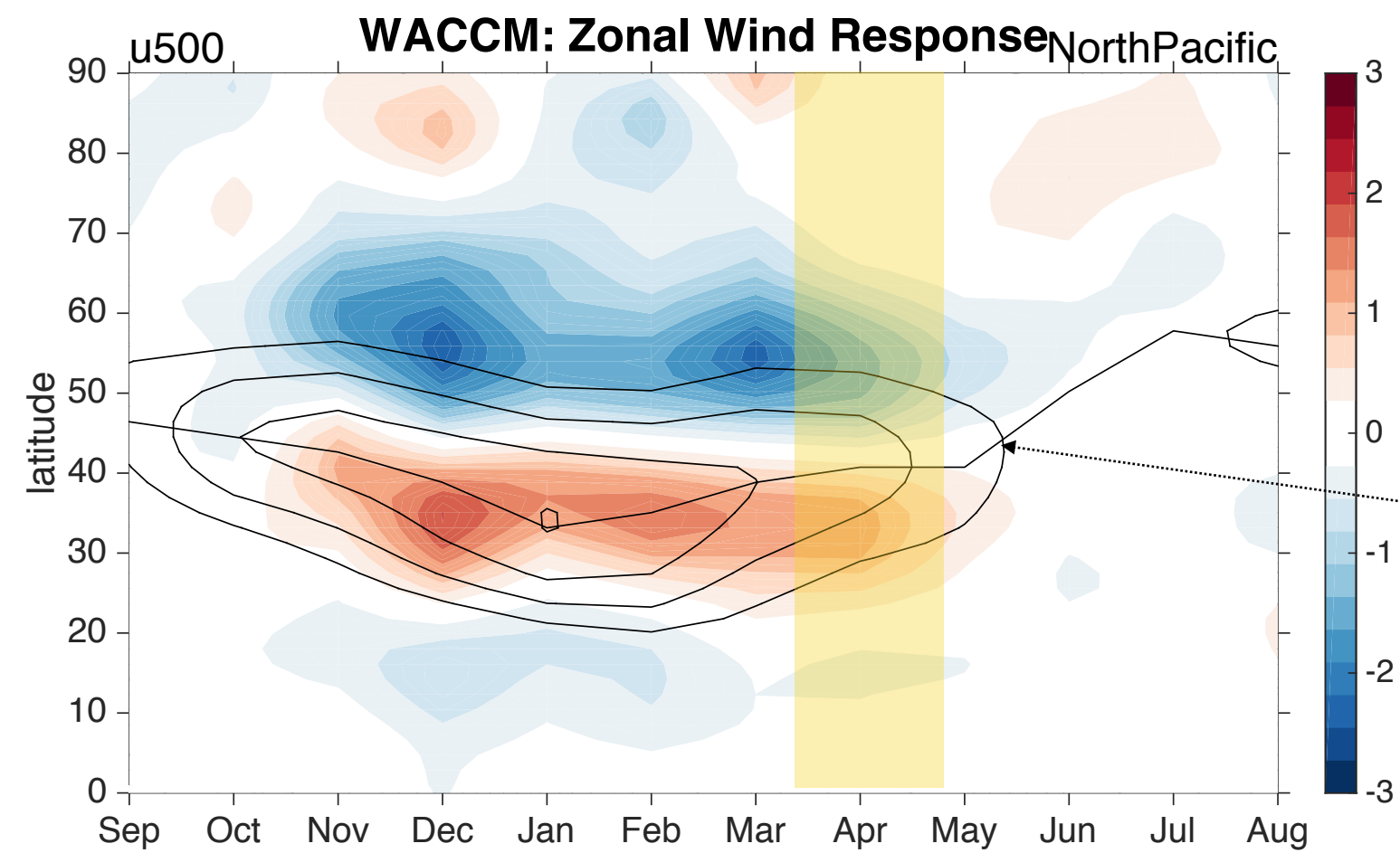
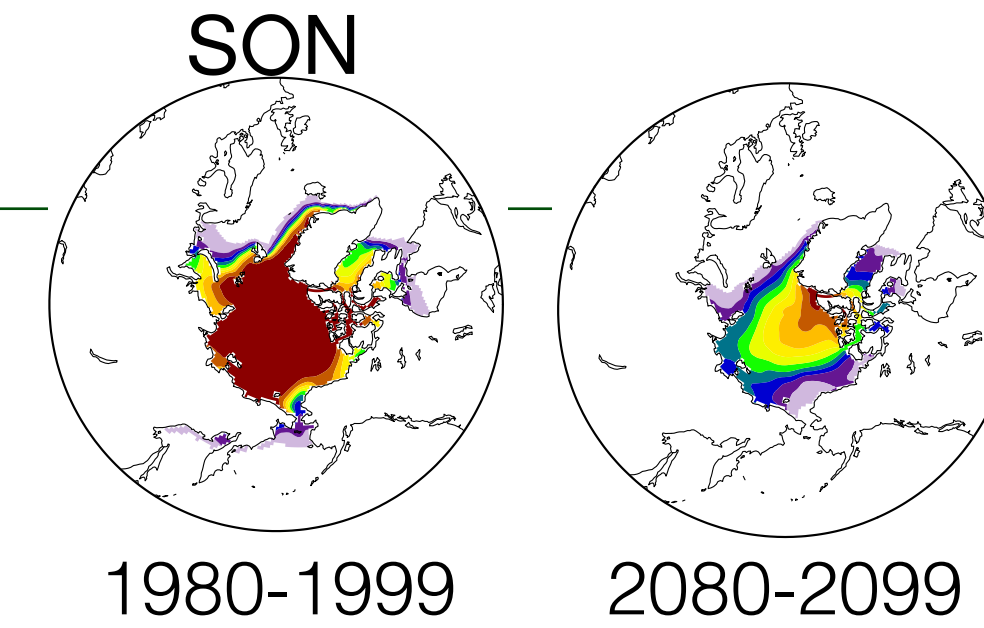
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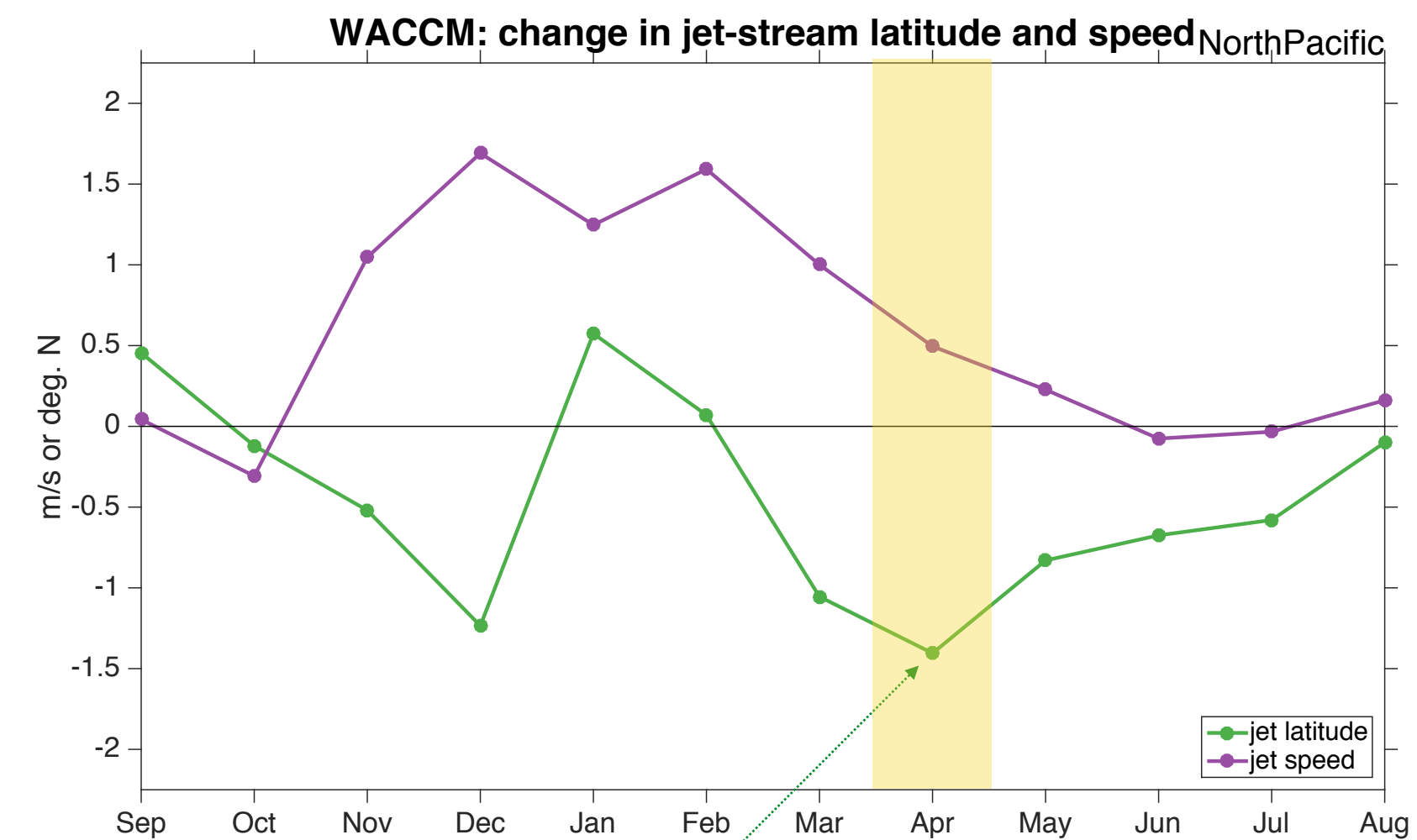
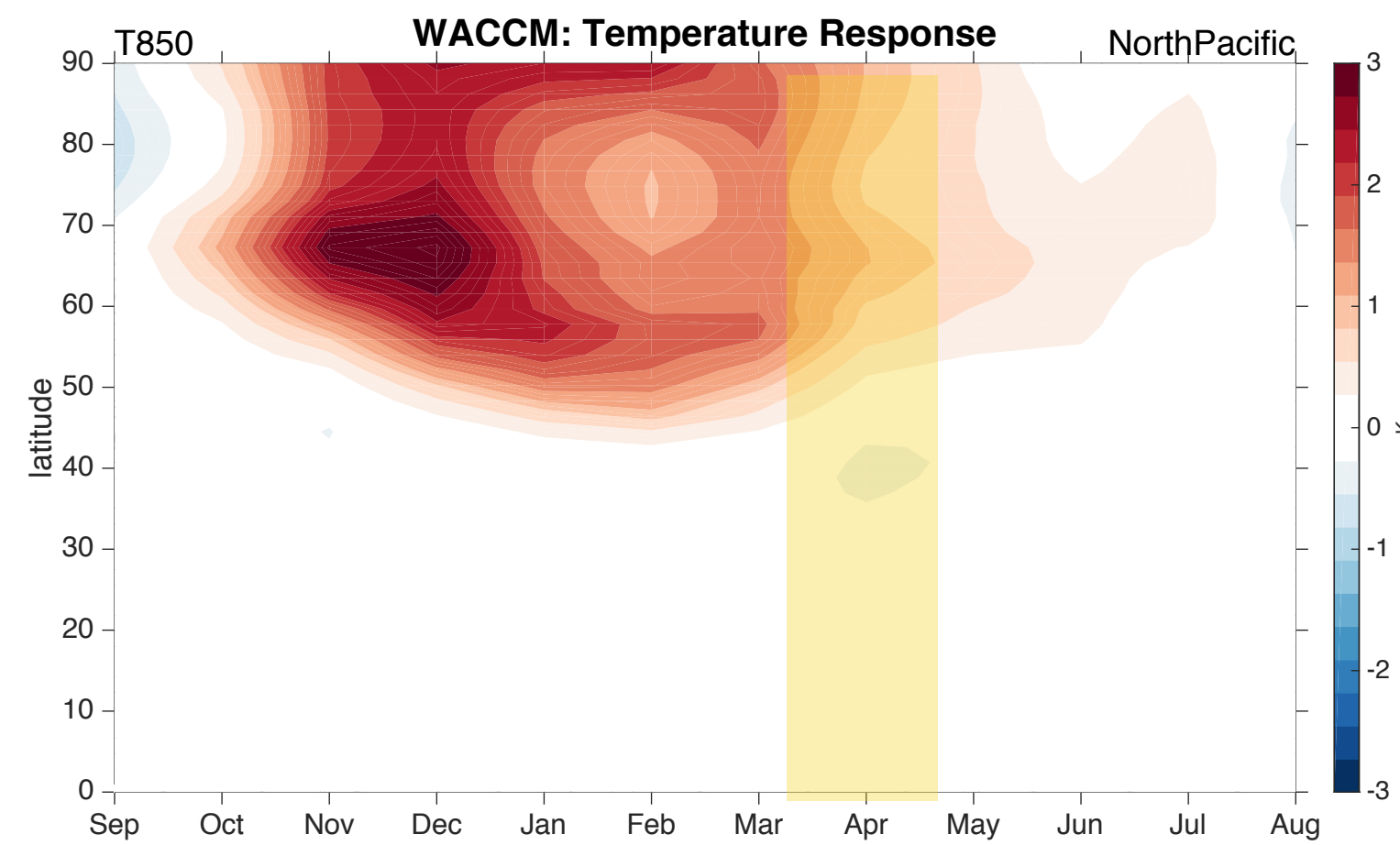
jet shift largest in Mar.-Apr.

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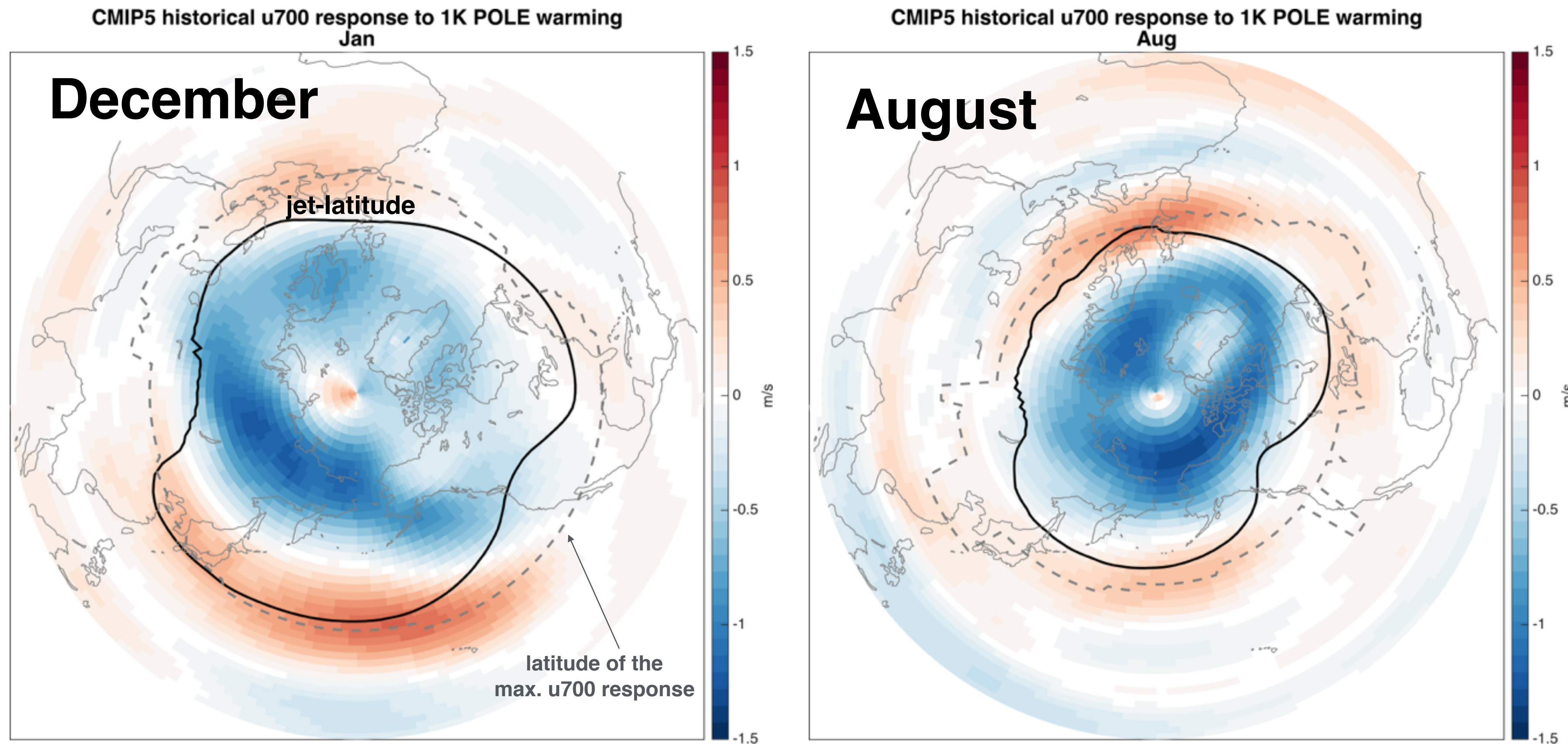
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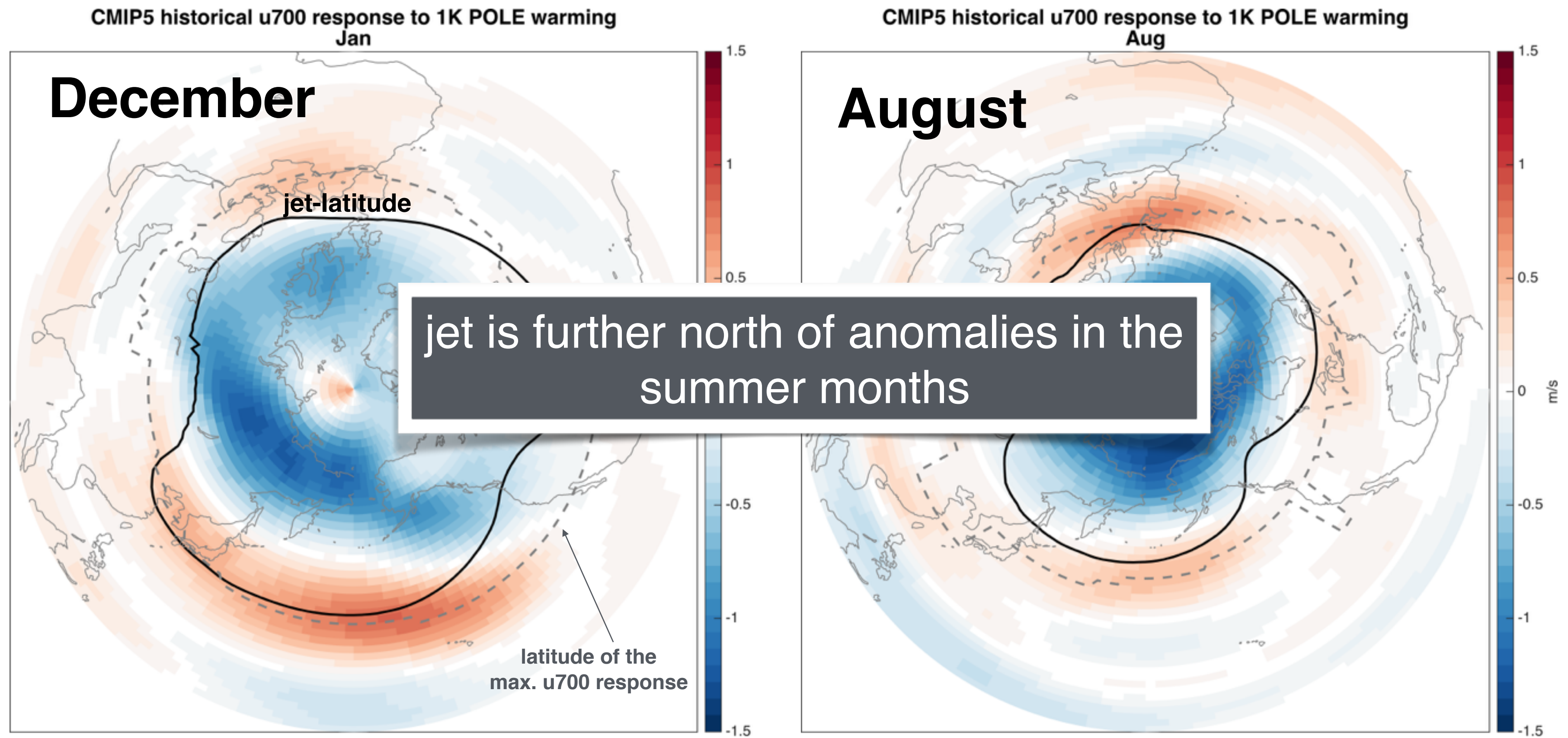
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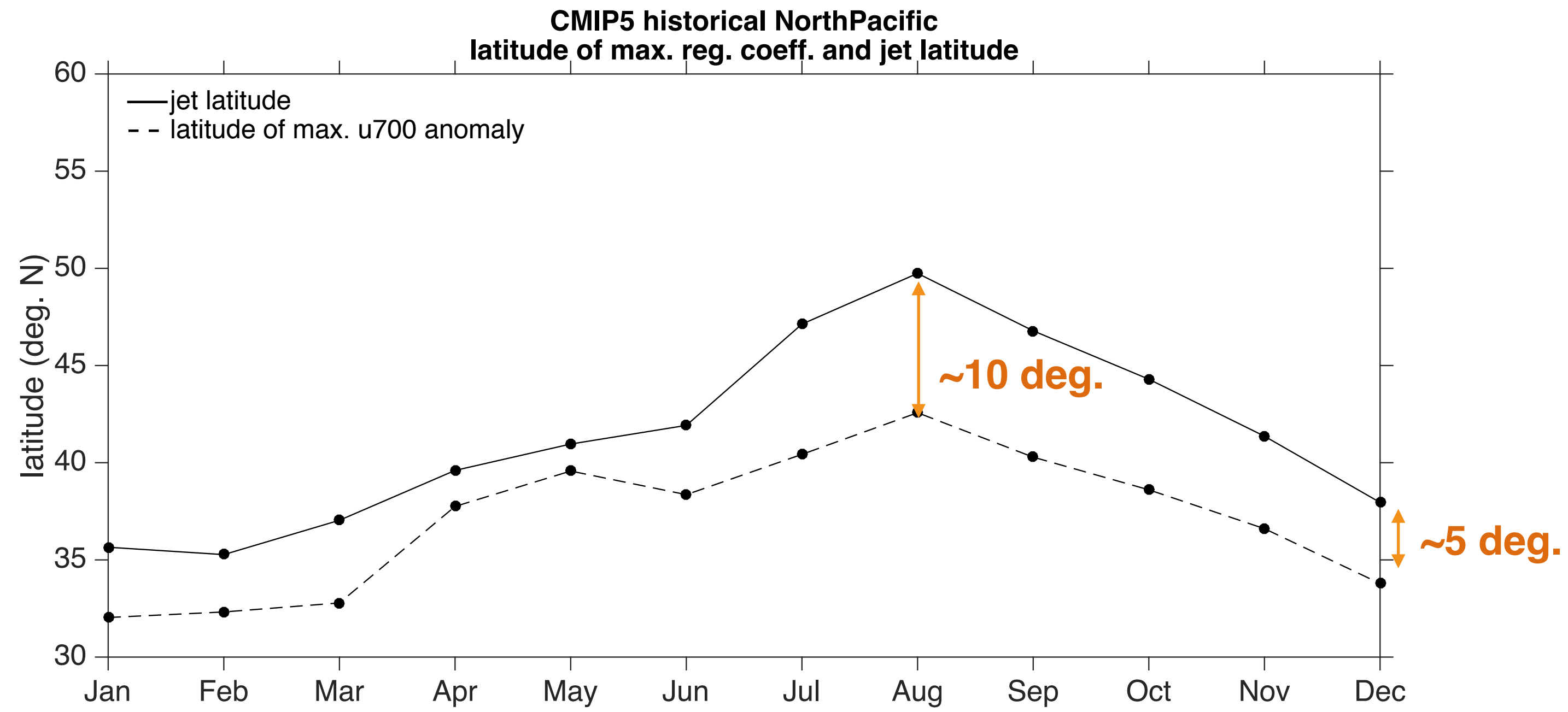
Summer vs. Winter: regression maps of u700



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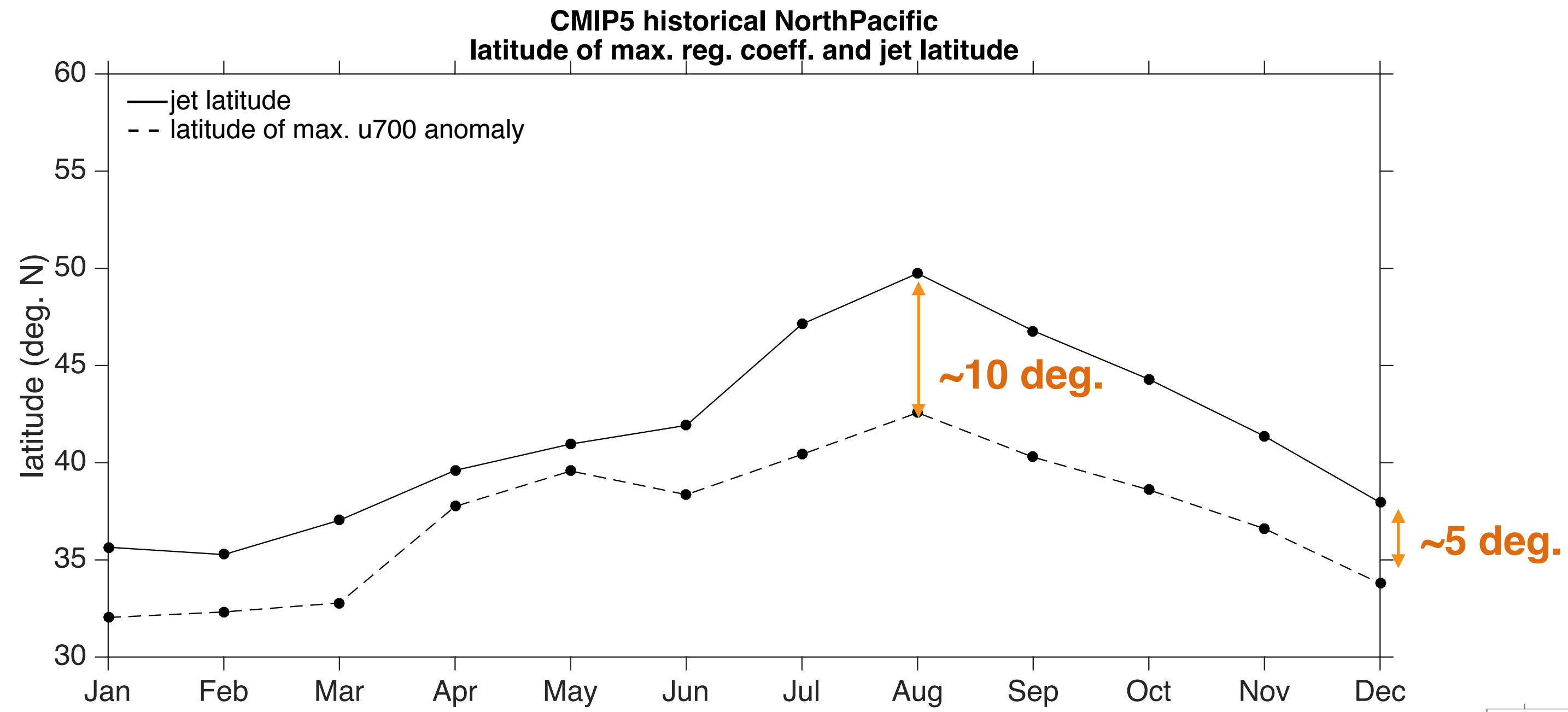
Relative location of the forcing matters for the jet shift



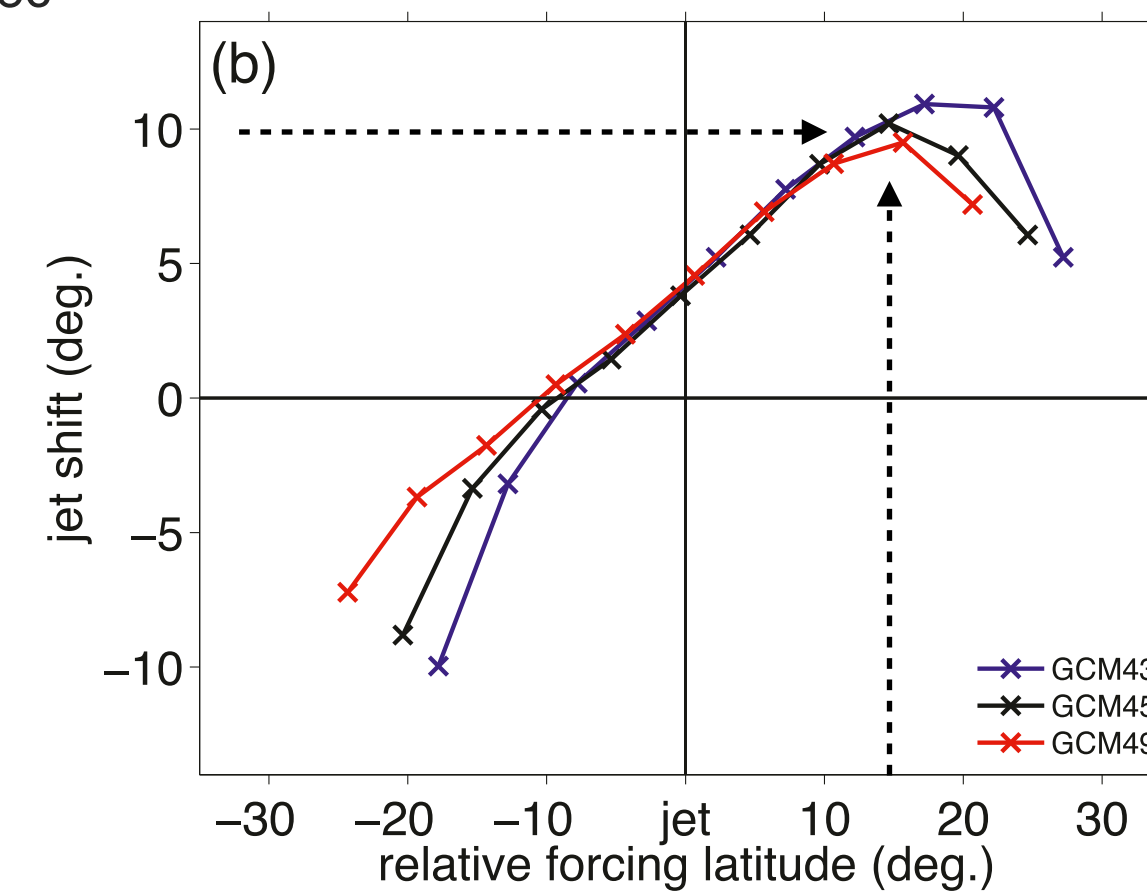
also explored in detail by Ring & Plumb (2007; JAS)

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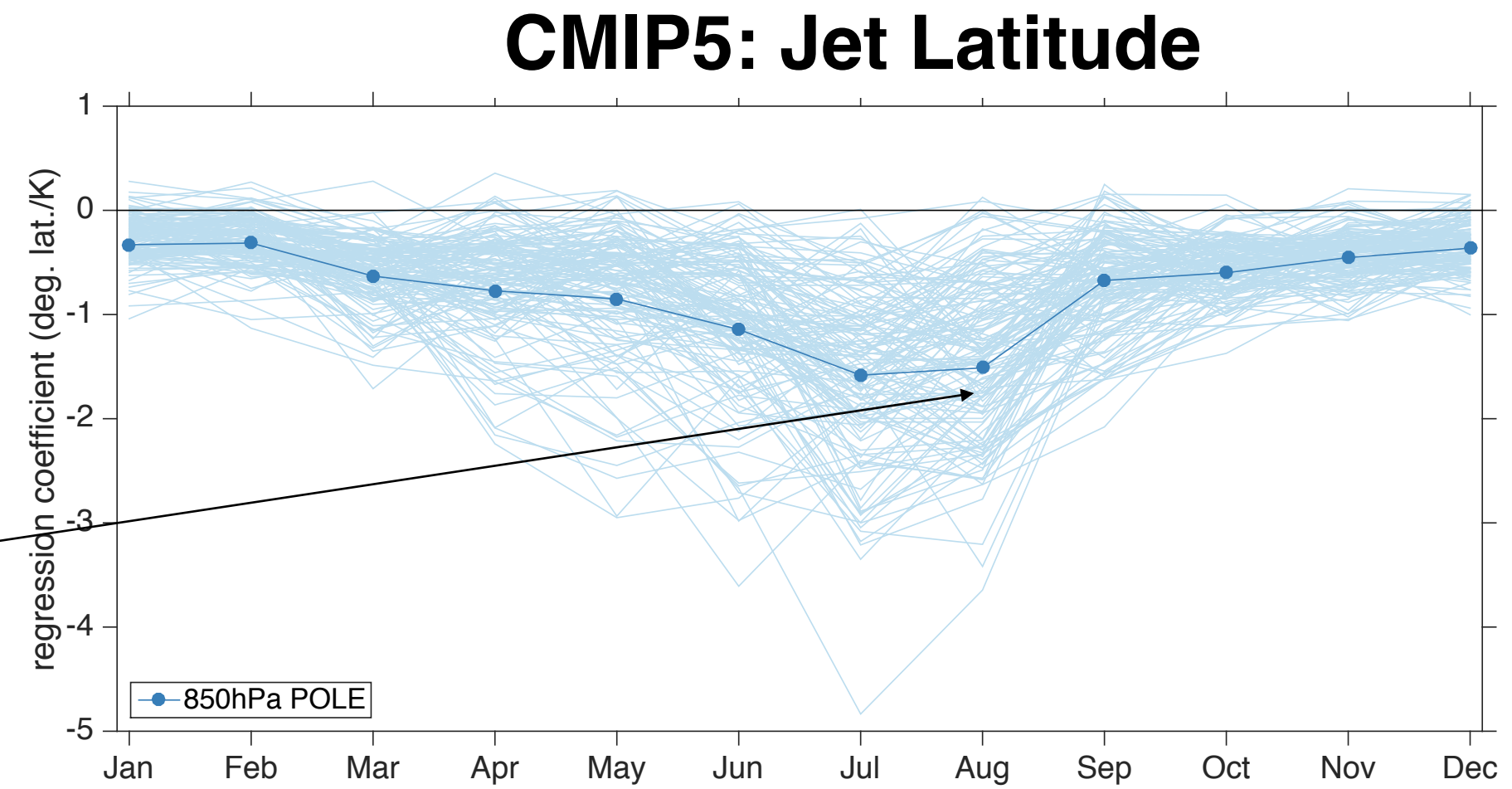
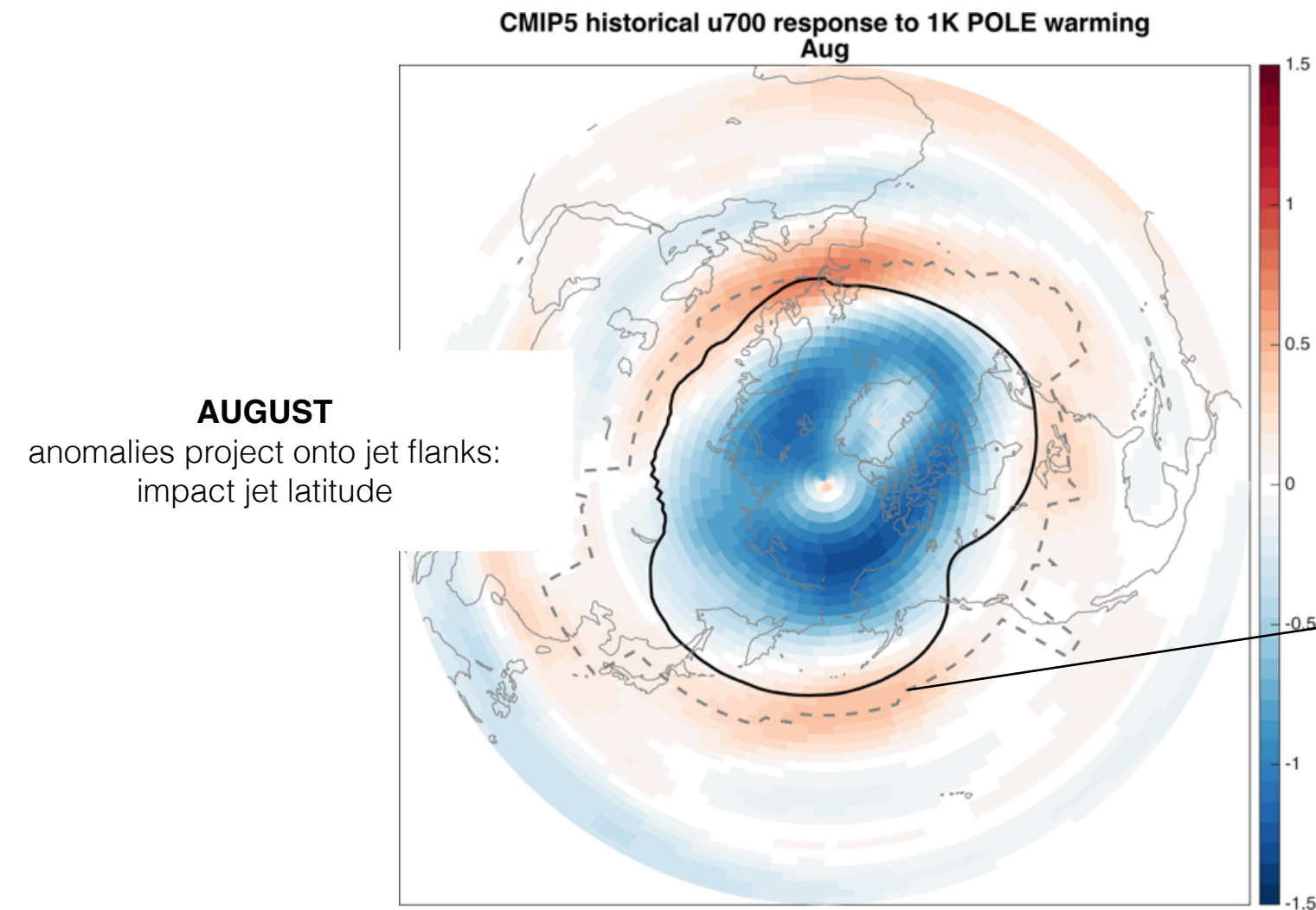
largest jet-shift when model is forced approximately 10-15 degrees off of its mean jet position



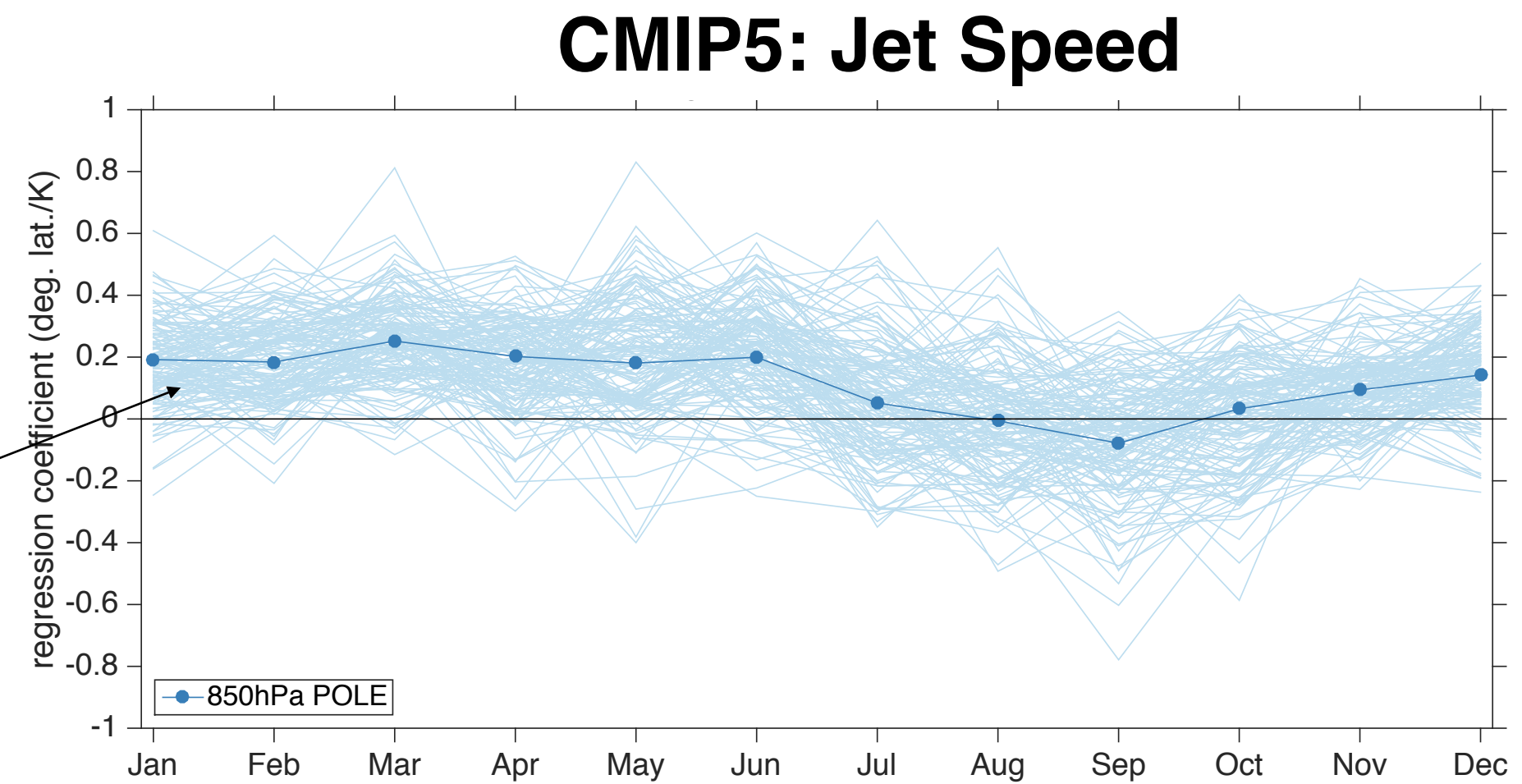
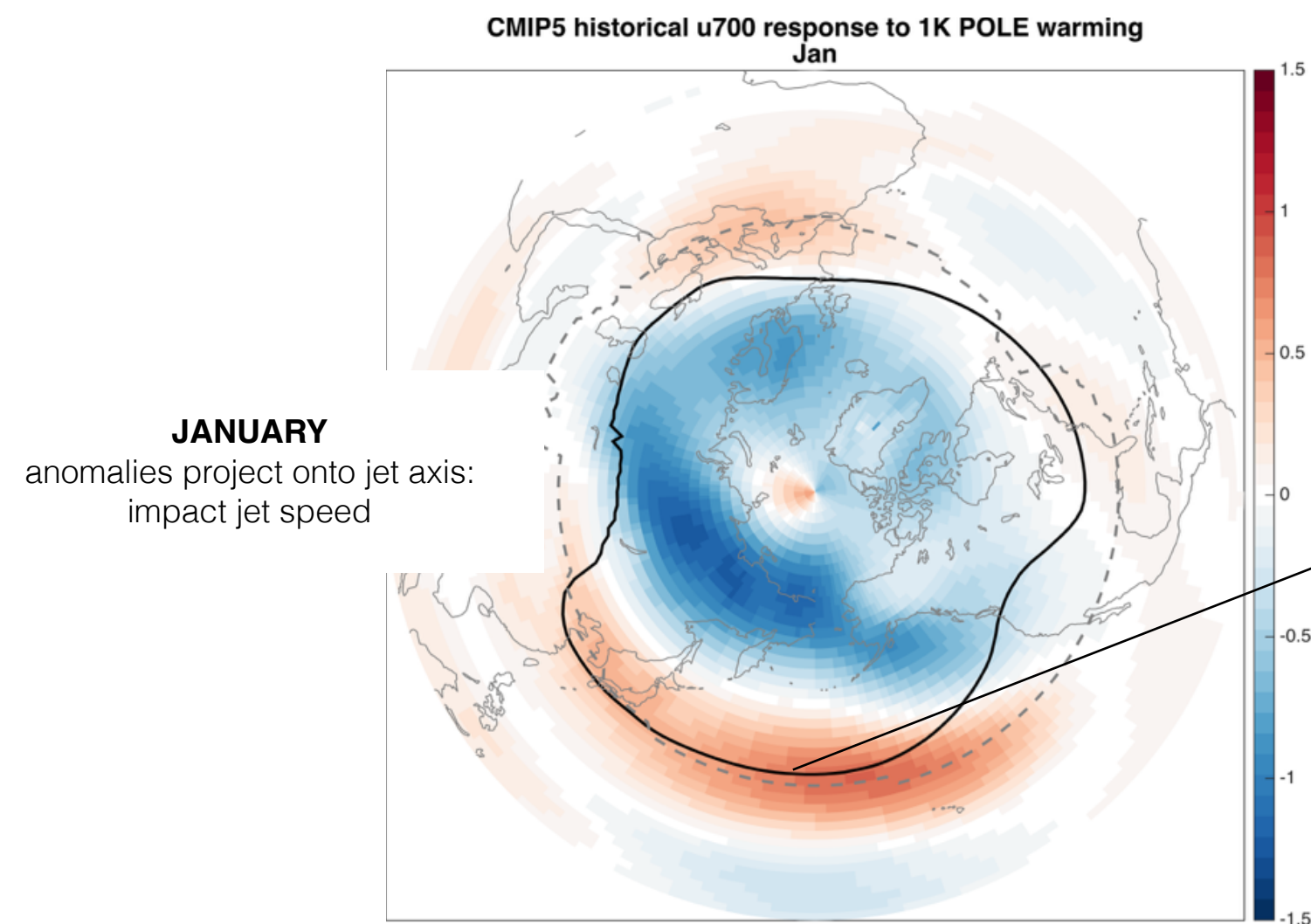
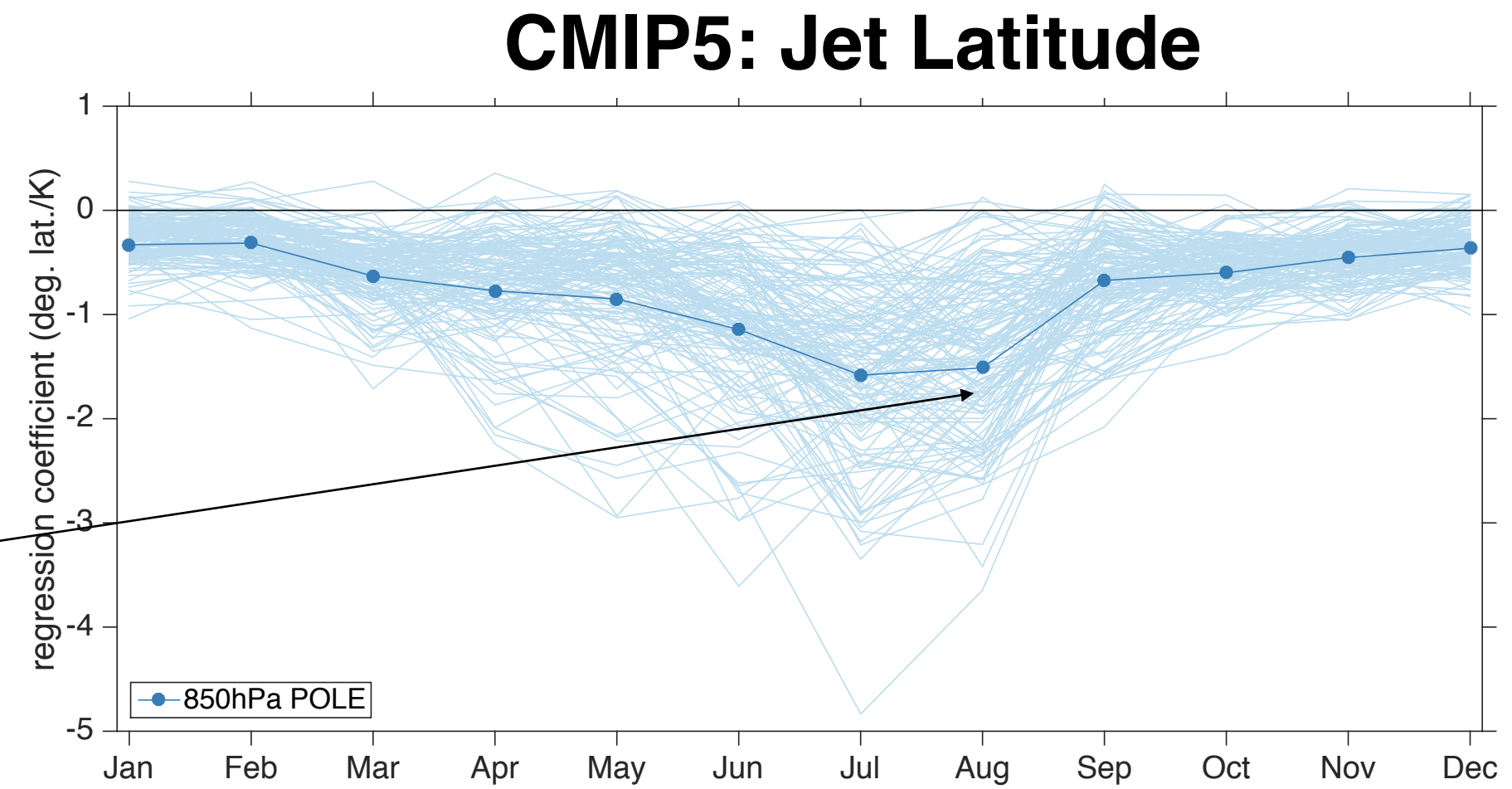
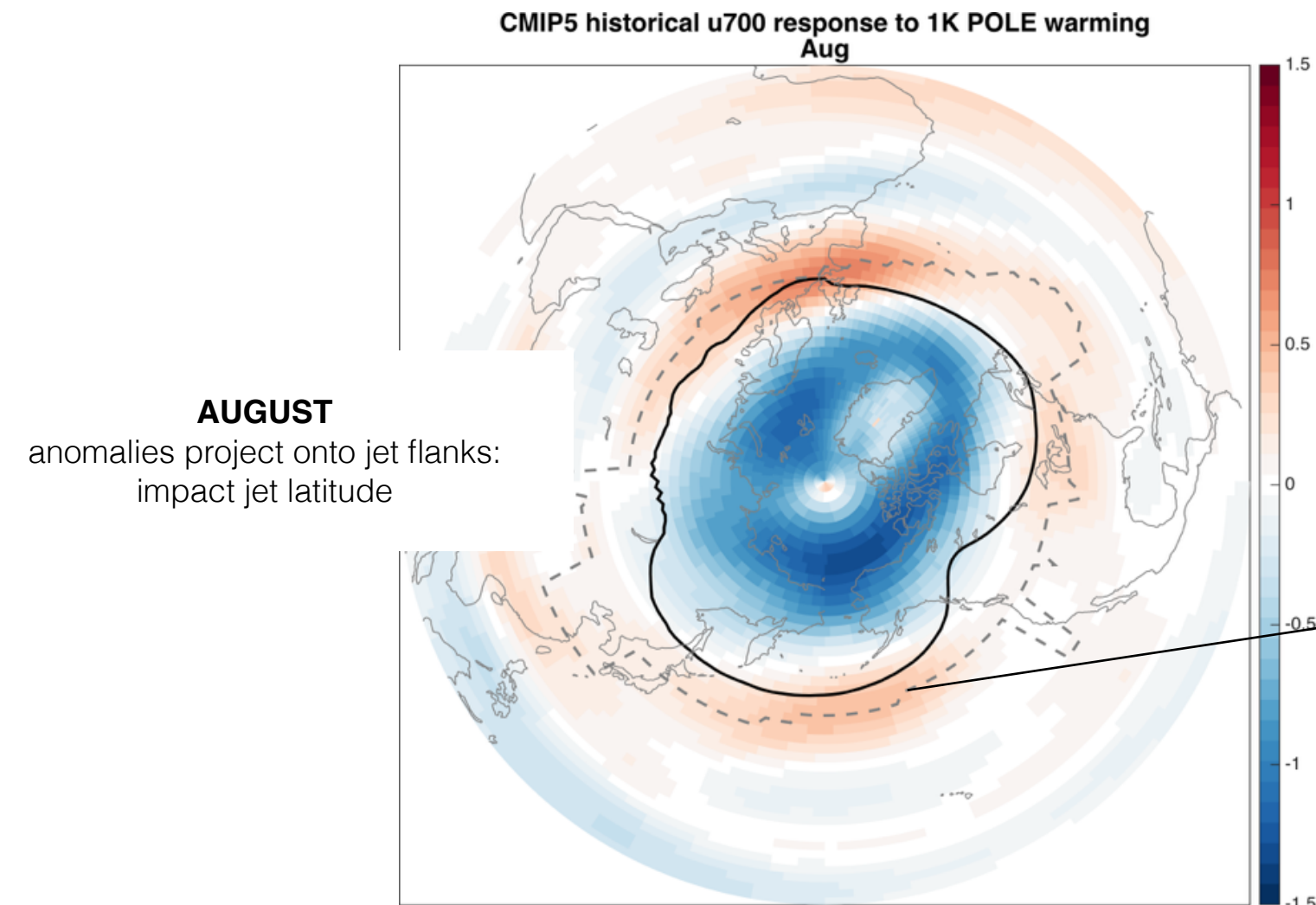
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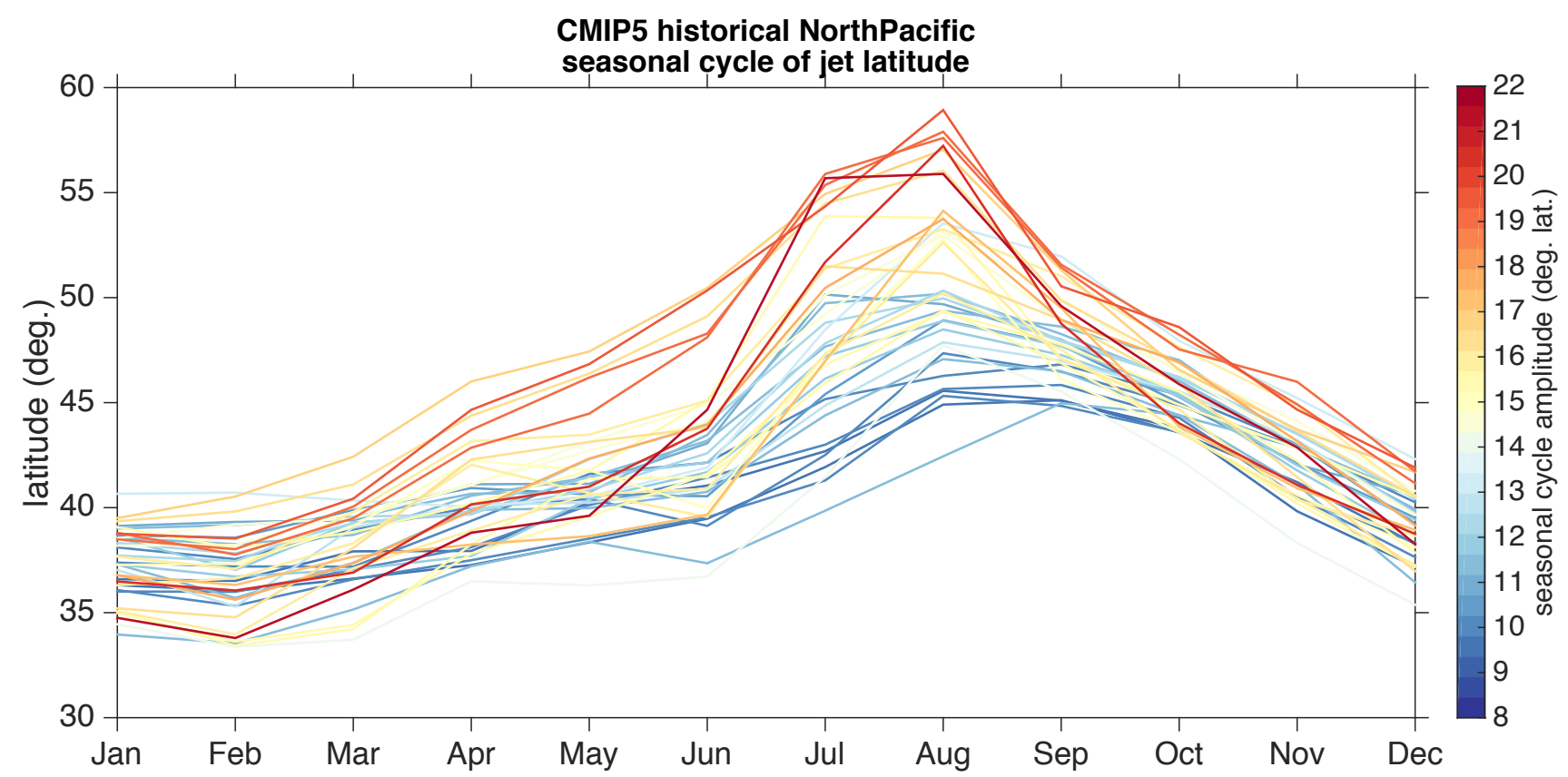
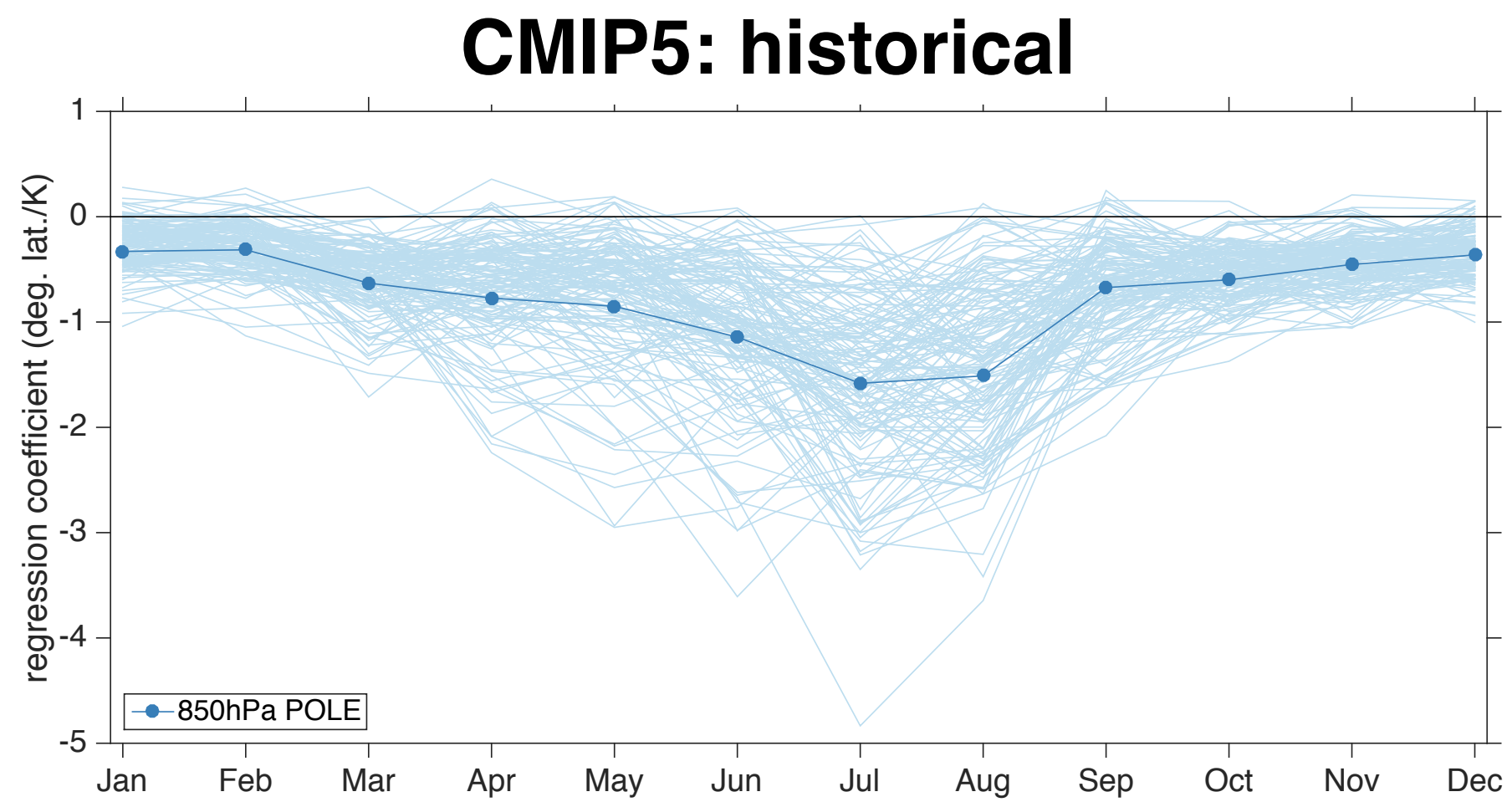
Summer vs. Winter: jet latitude vs. jet speed



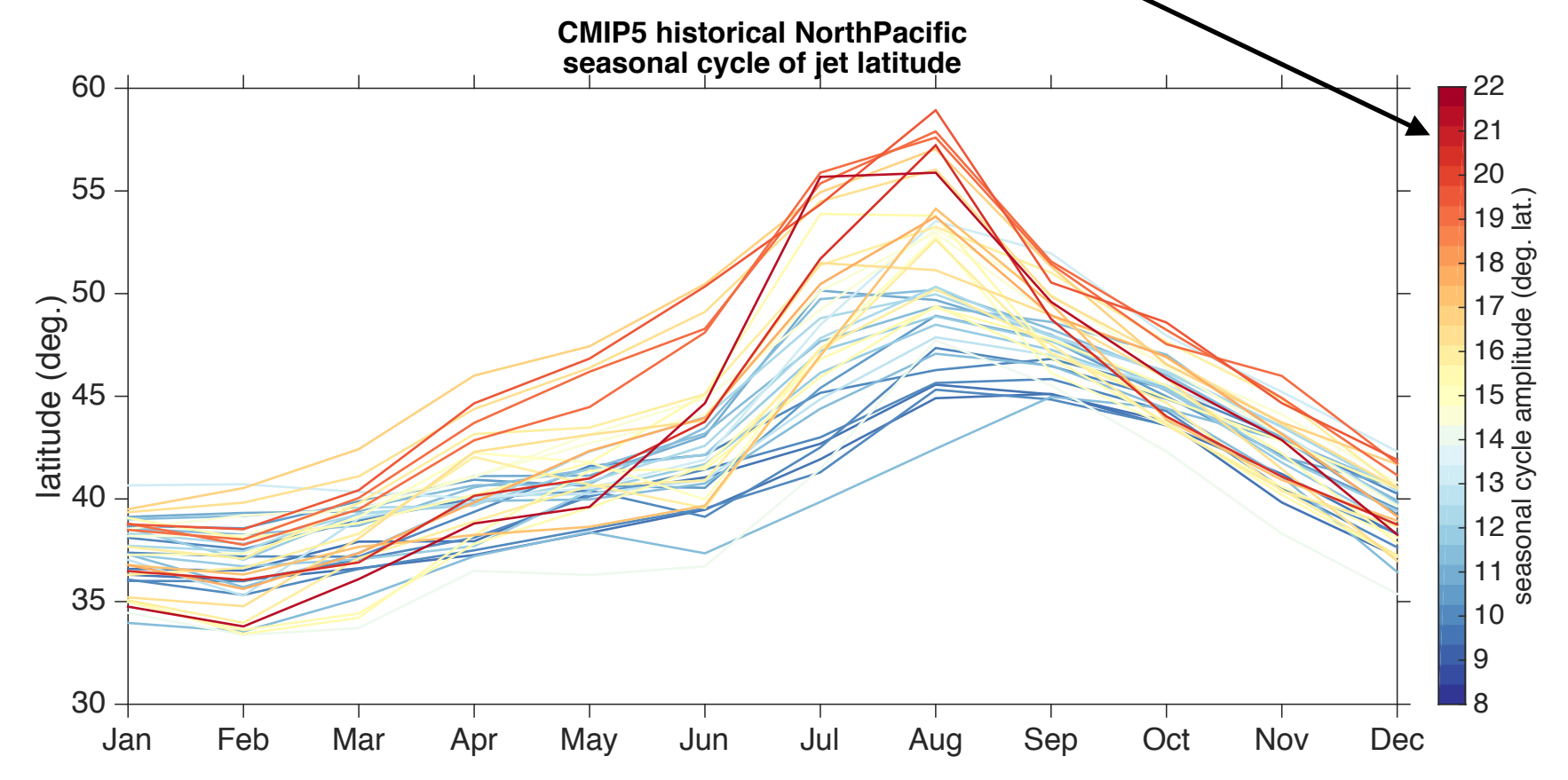
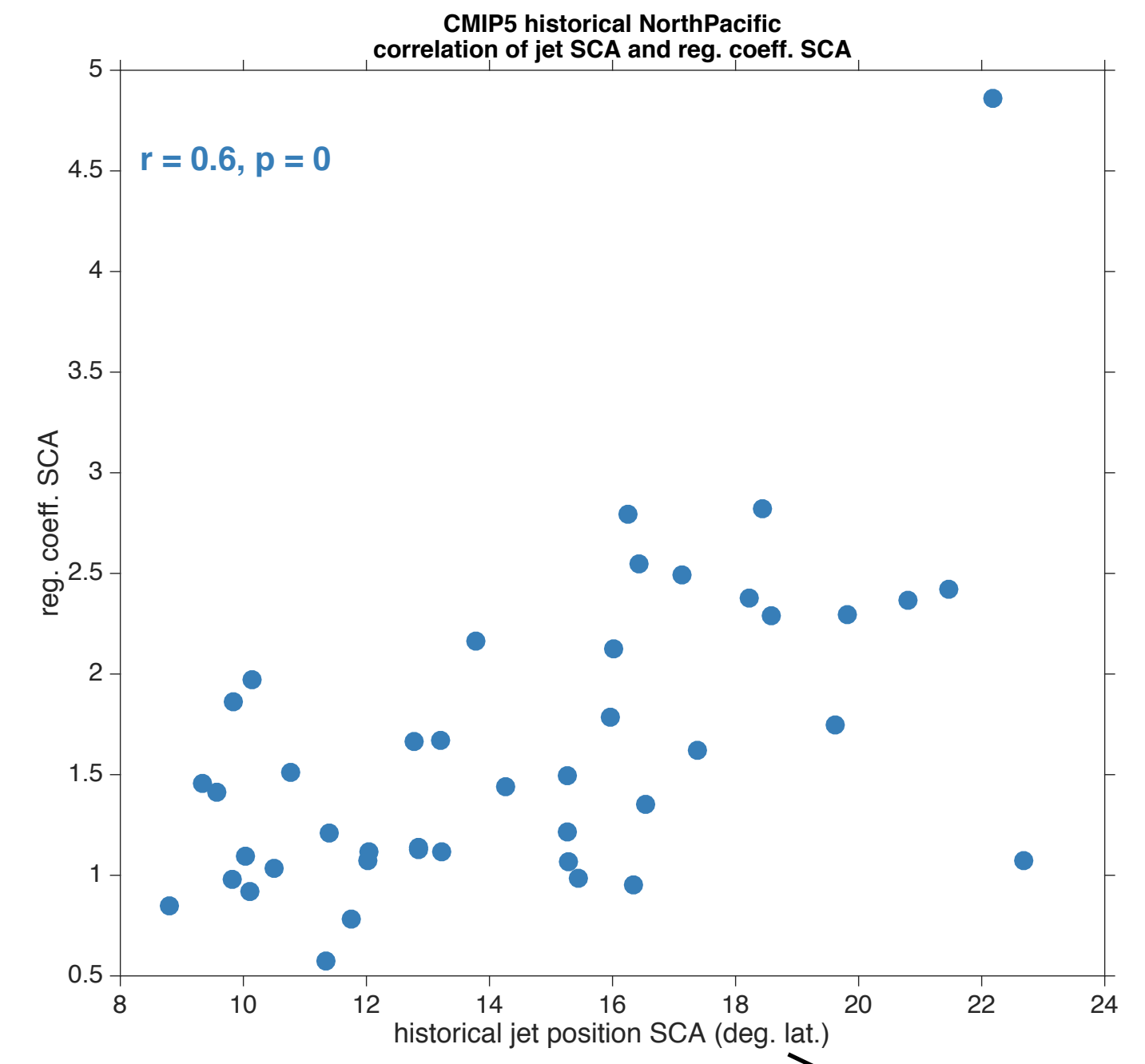
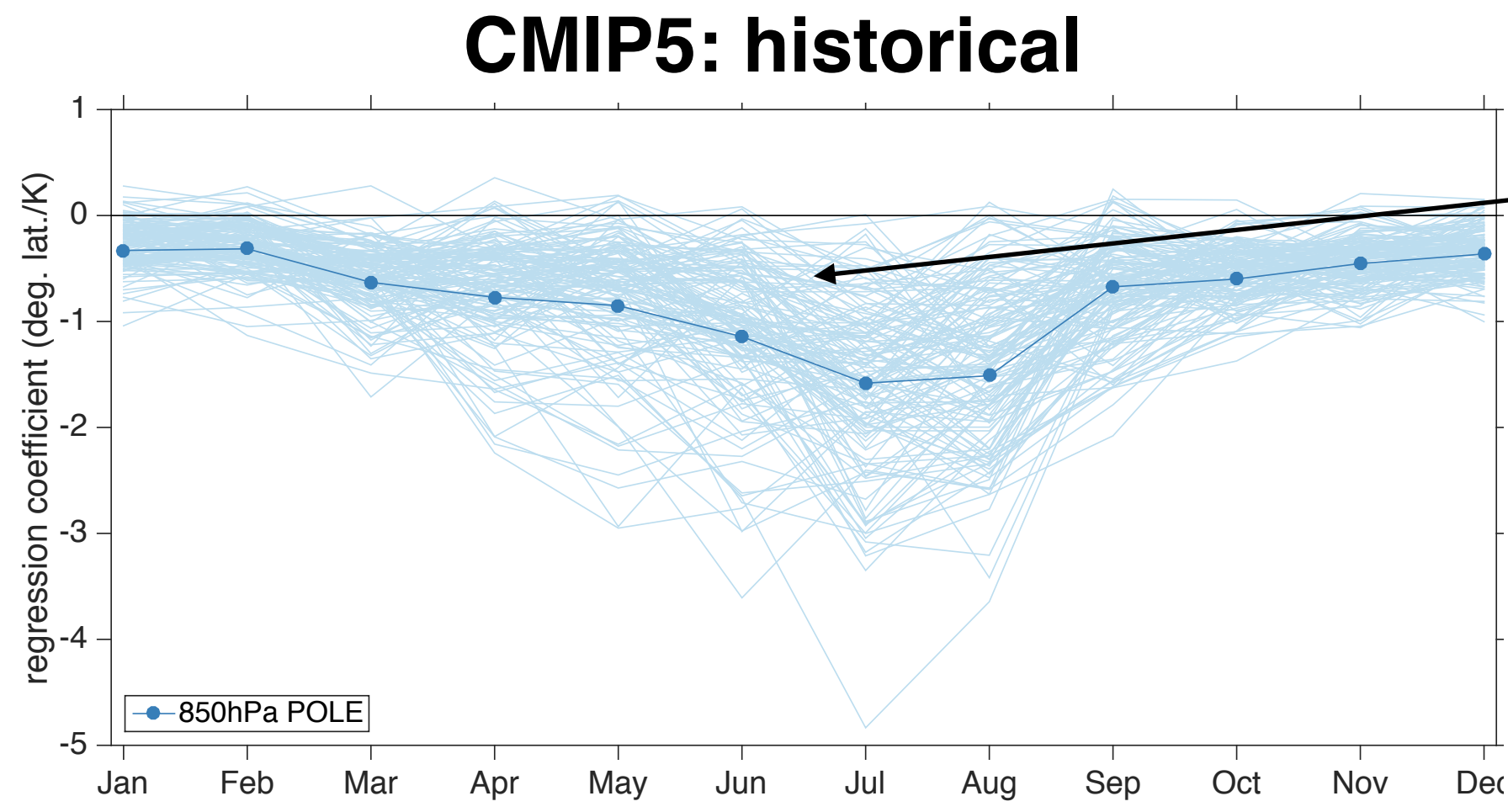
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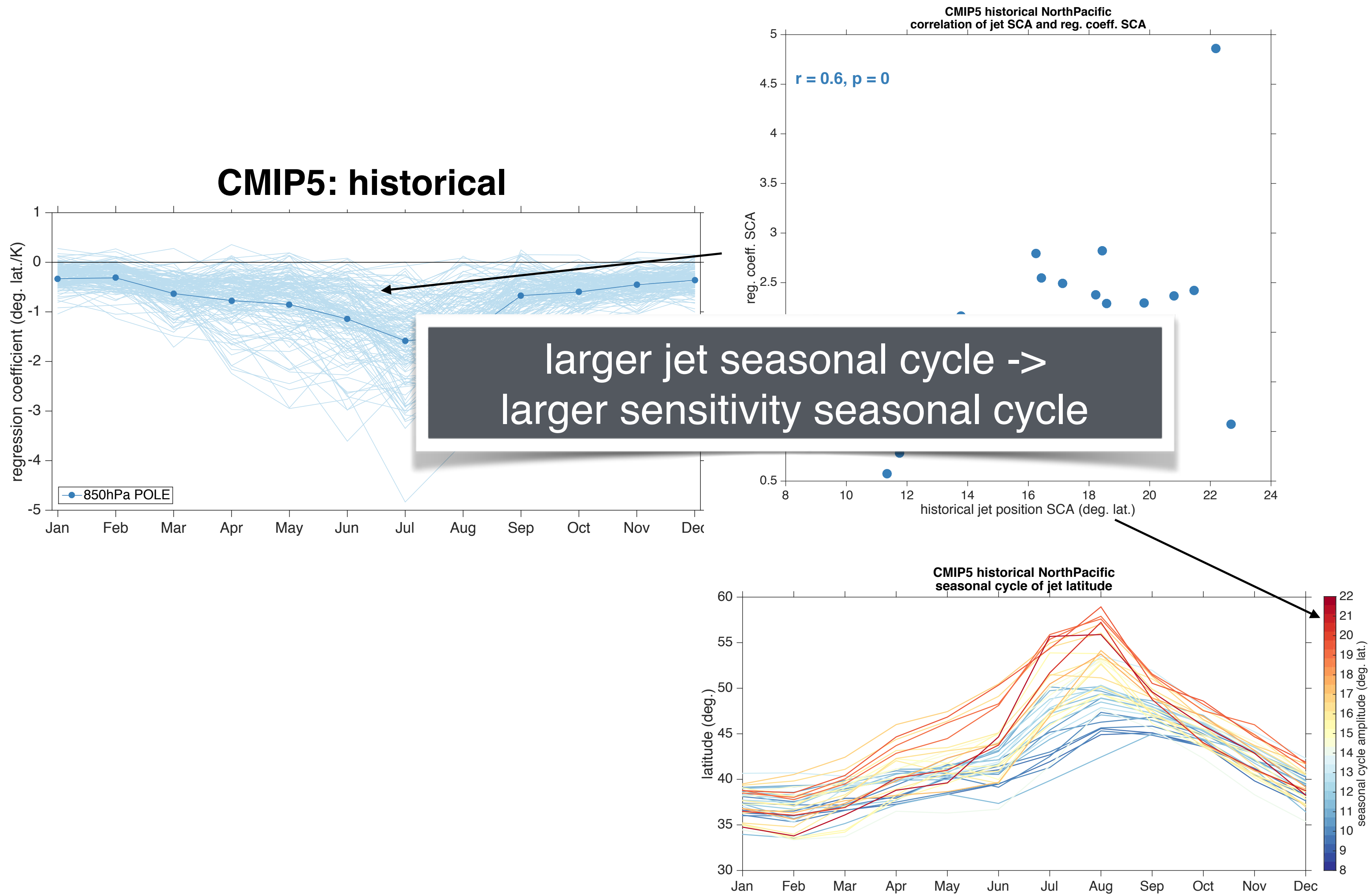
CMIP5 model biases in the seasonal cycle



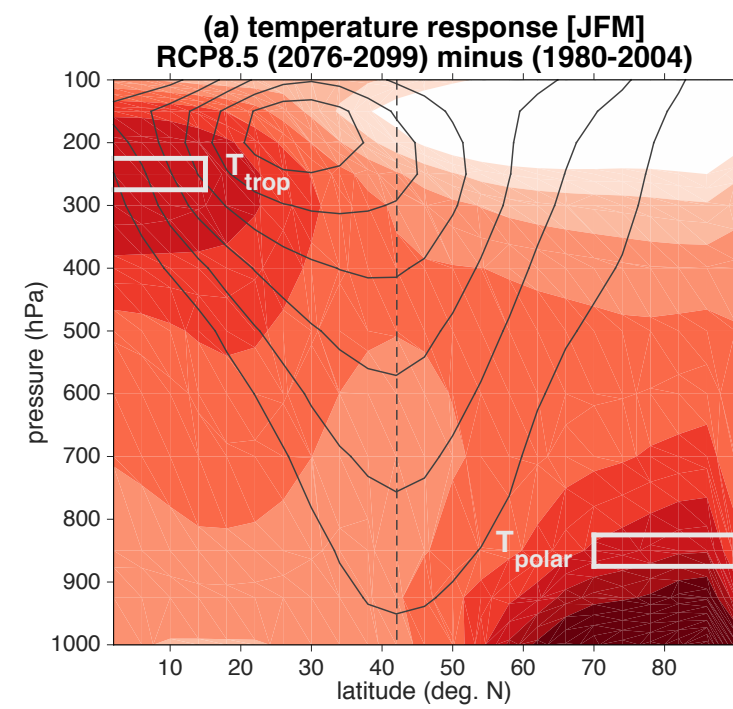
CMIP5 model biases in the seasonal cycle



CMIP5 model biases in the seasonal cycle



Next...

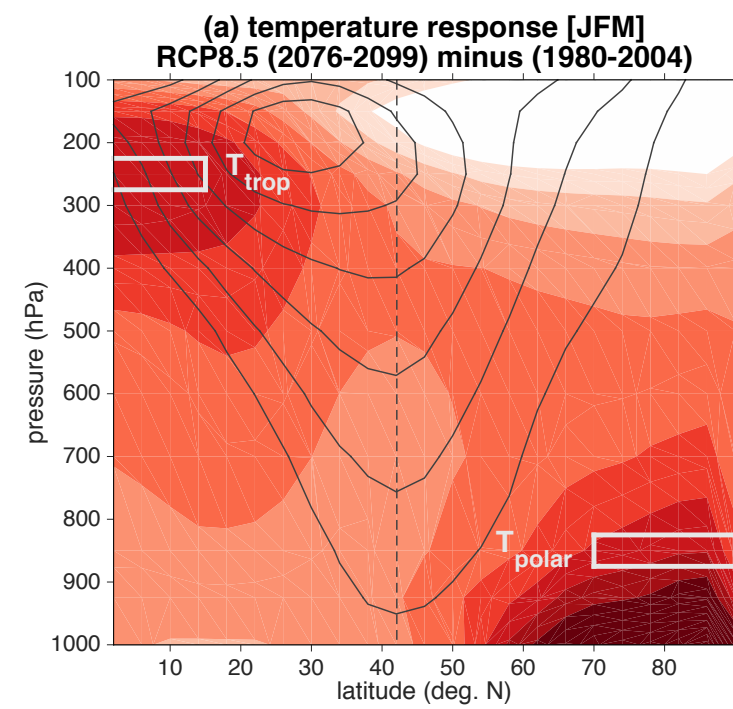


Can we find evidence of seasonal sensitivity of the jet-streams in comprehensive GCMs?

Data

- CESM1 Large Ensemble (LENS; 40 simulations)
- CMIP5 models (historical & RCP8.5)

Next...



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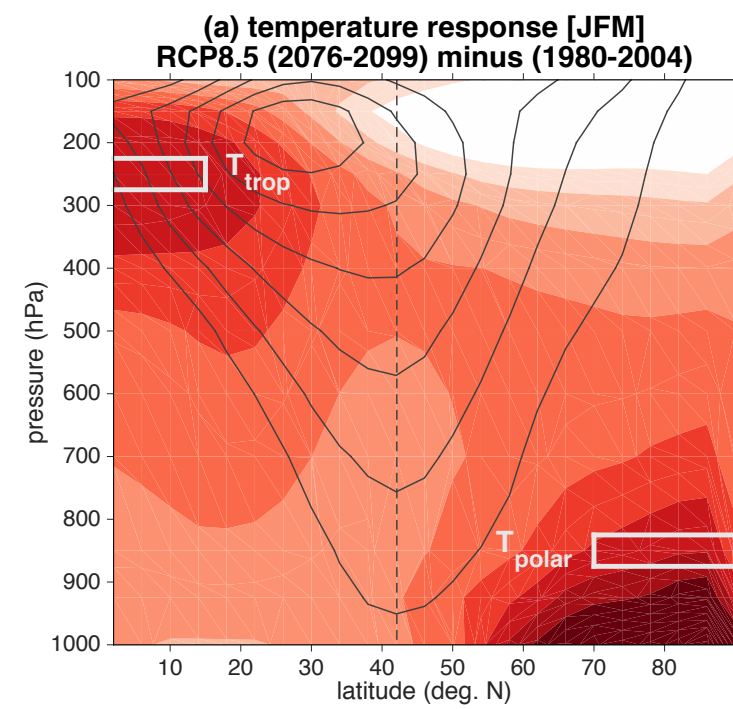
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$$\Delta\theta_{jet} = r_{polar}\Delta T_{polar} + r_{trop}\Delta T_{trop}$$

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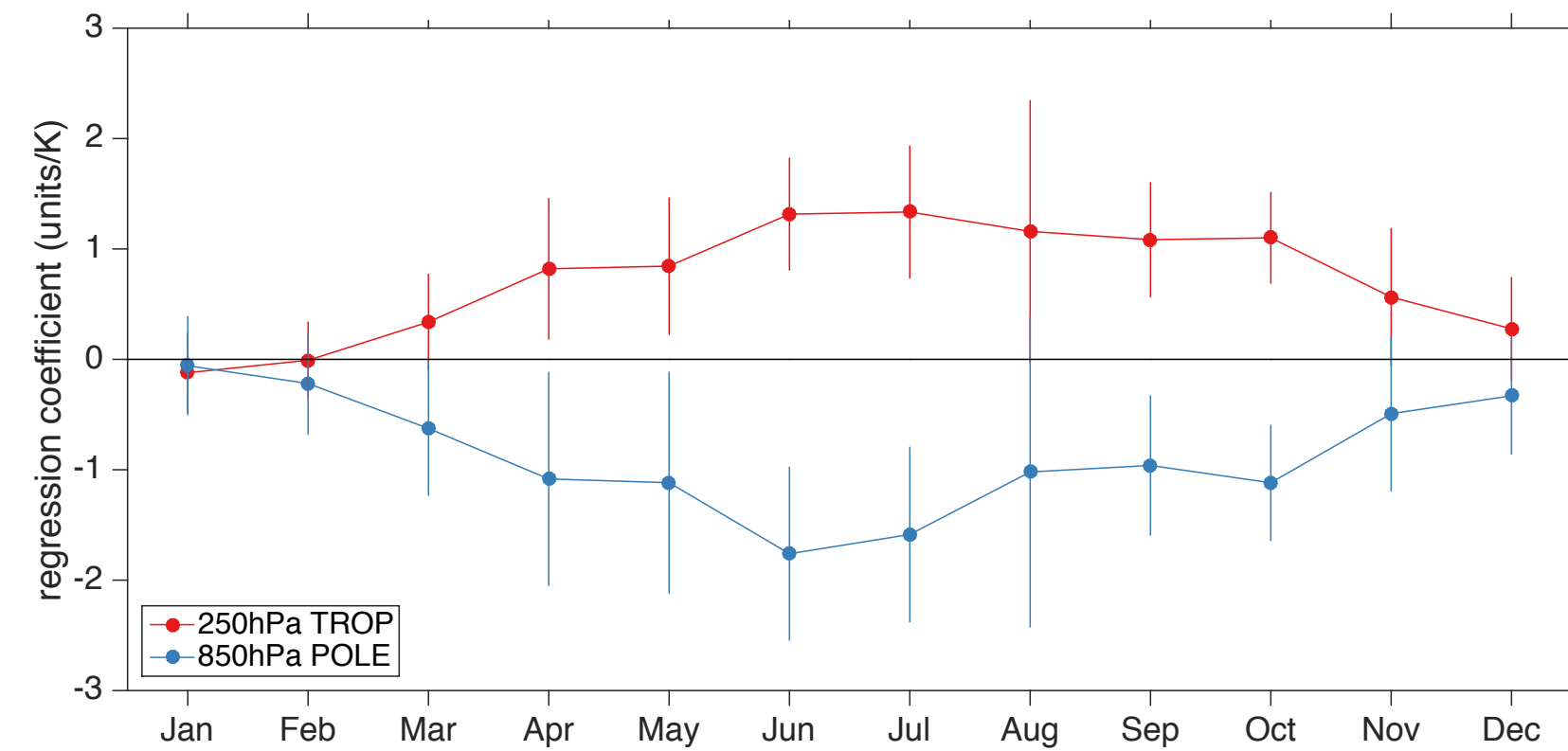
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Time evolution within simulations

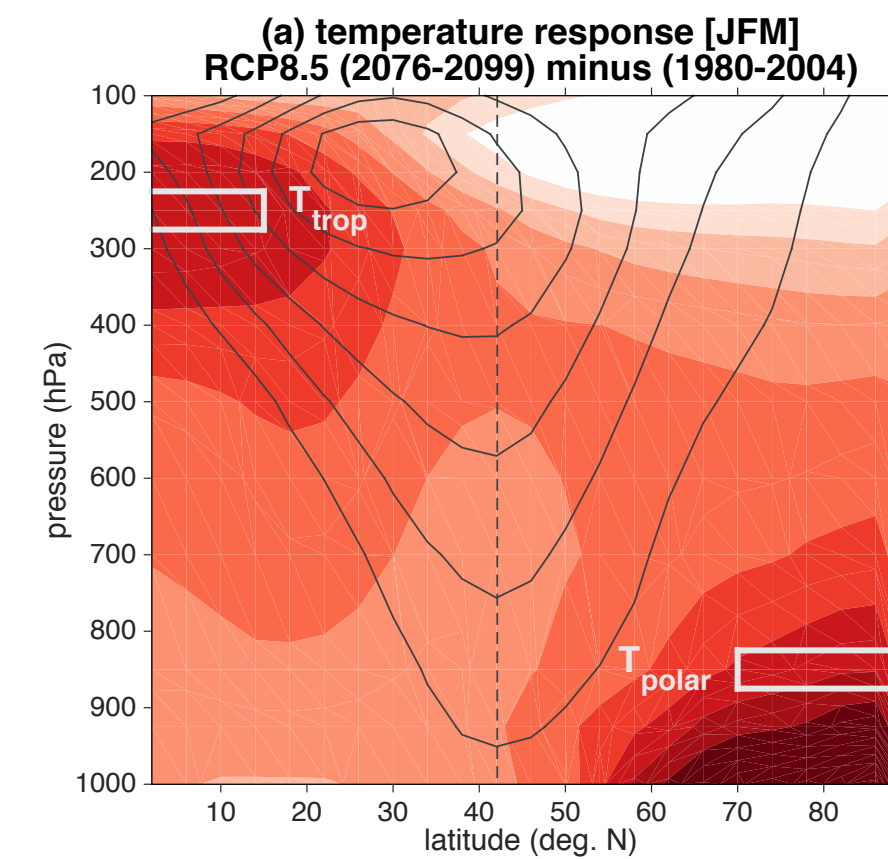
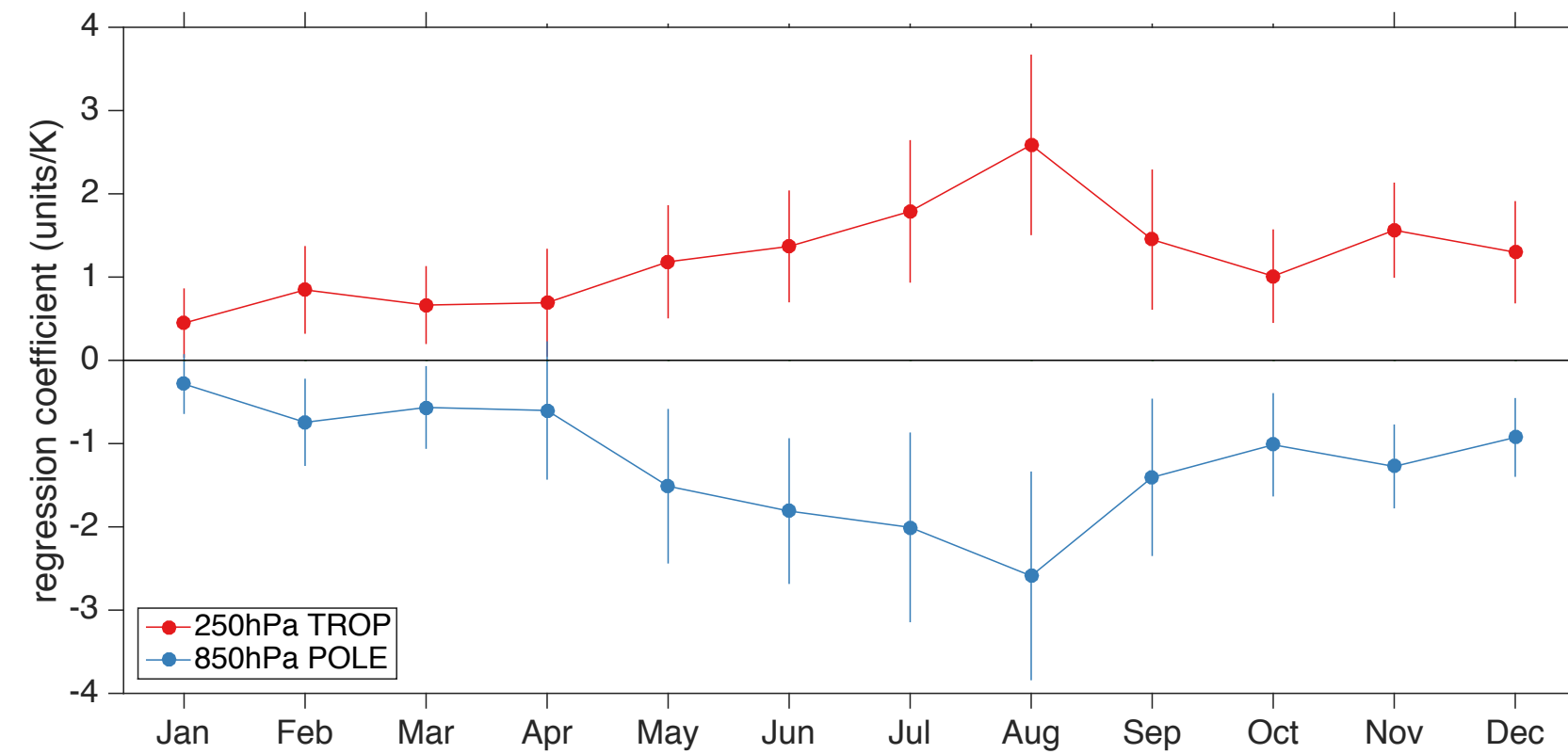
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Model spread: seasonal sensitivity of jet position

LENS: North Atlantic



LENS: North Pacific



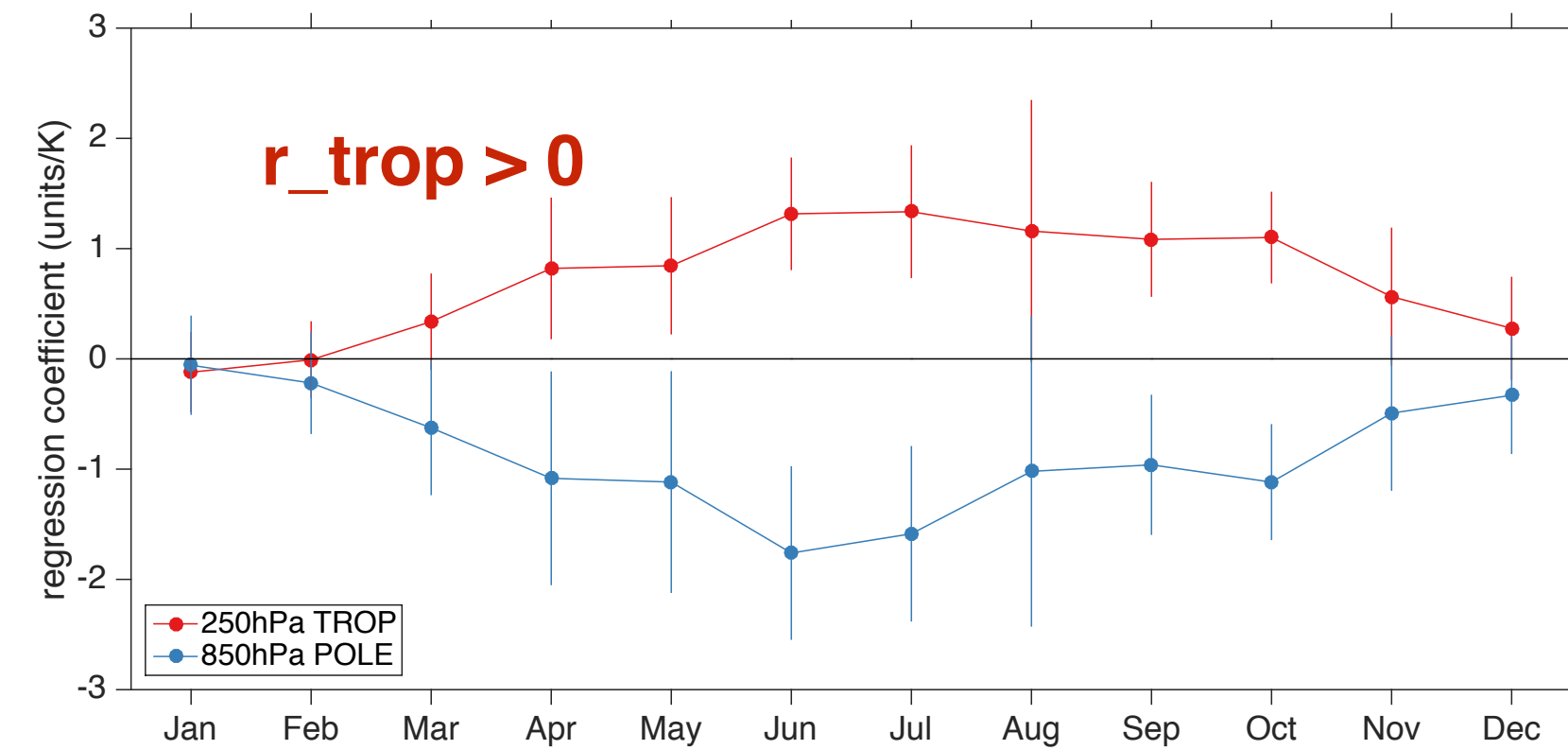
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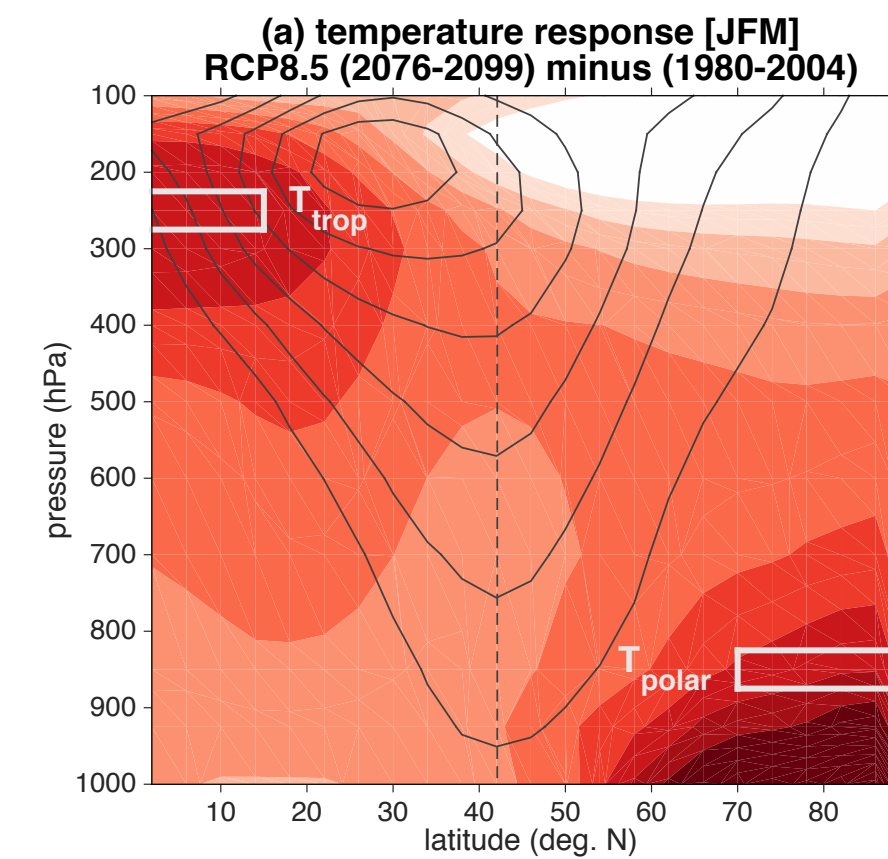
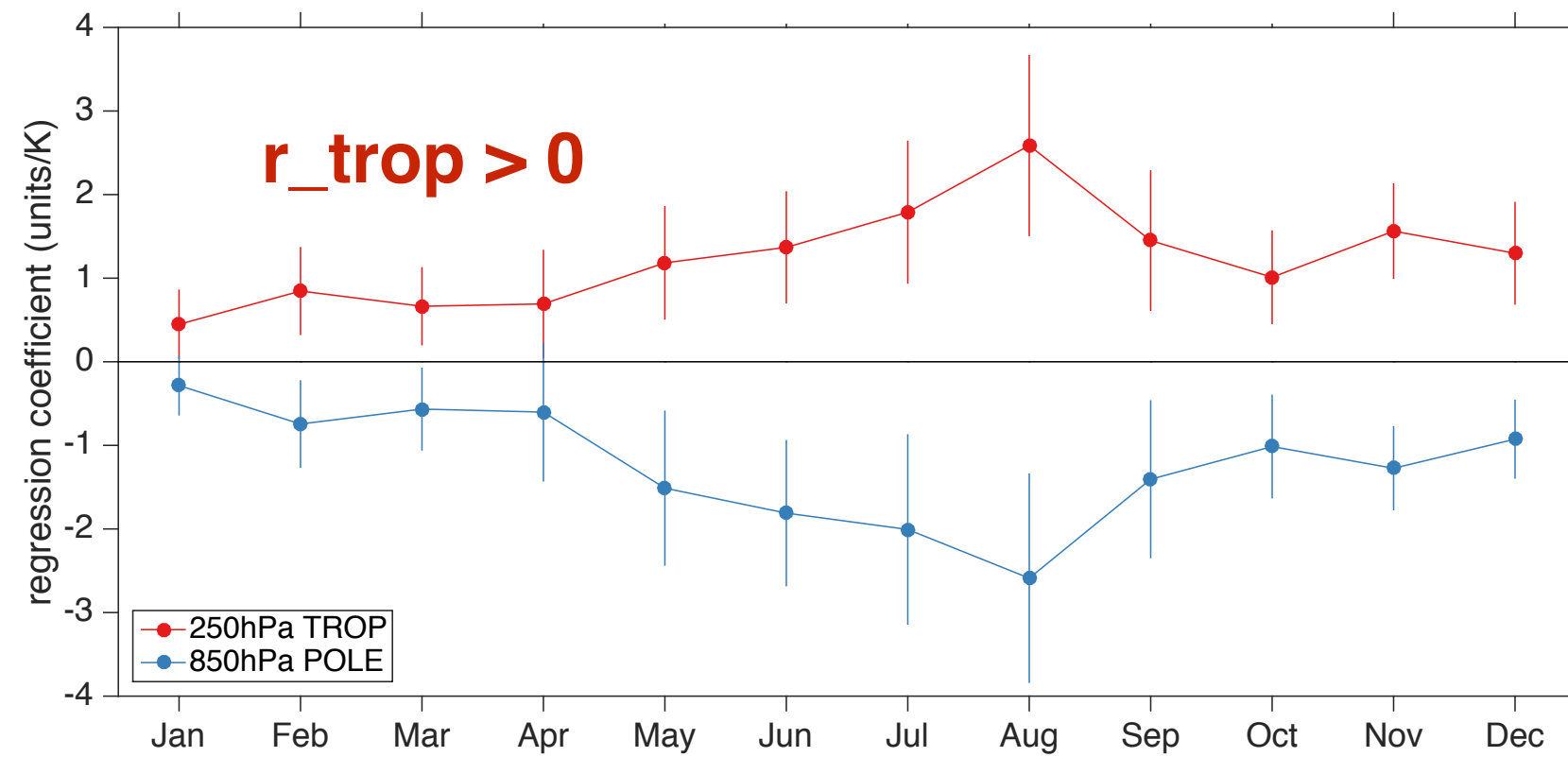
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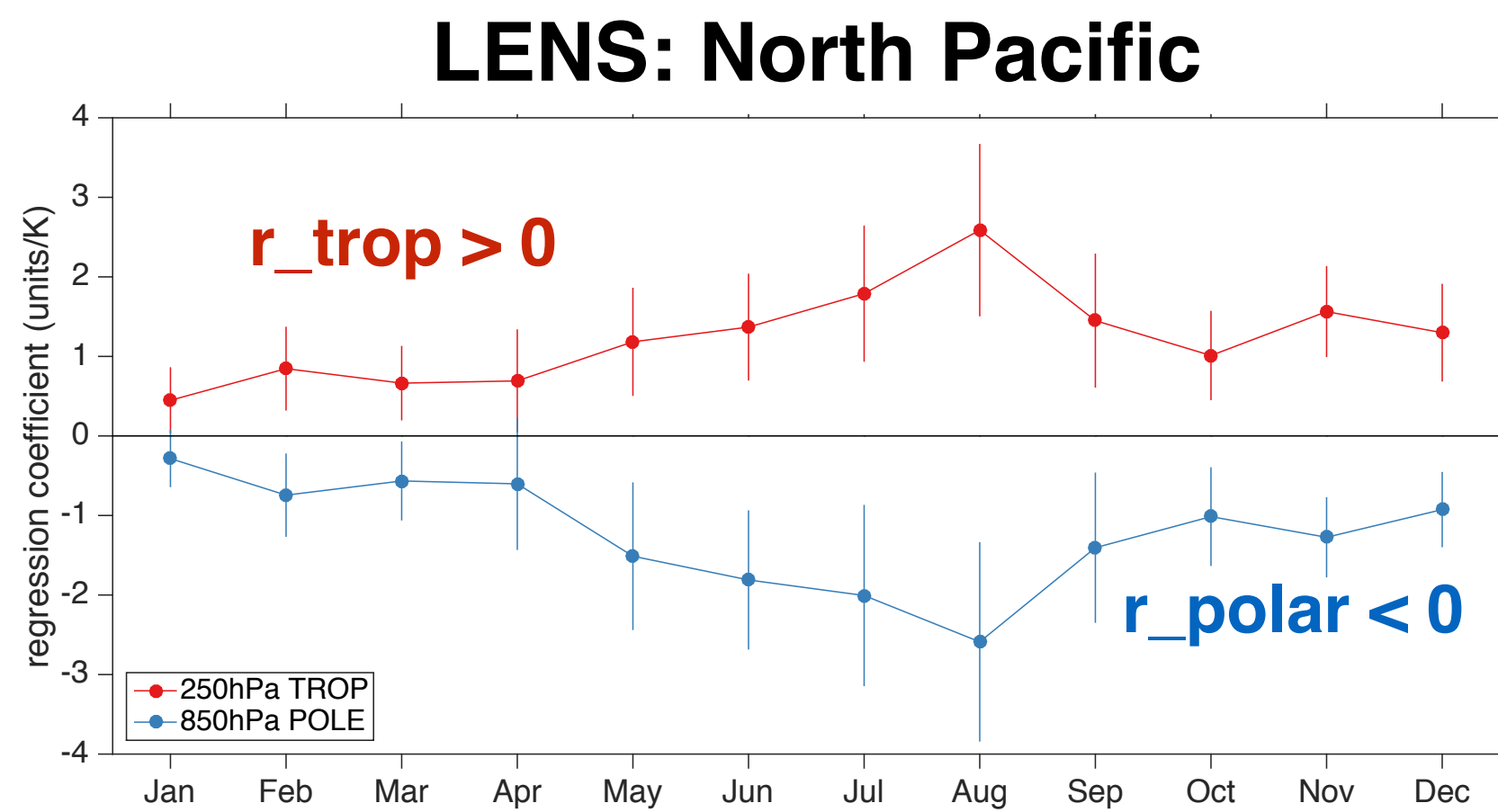
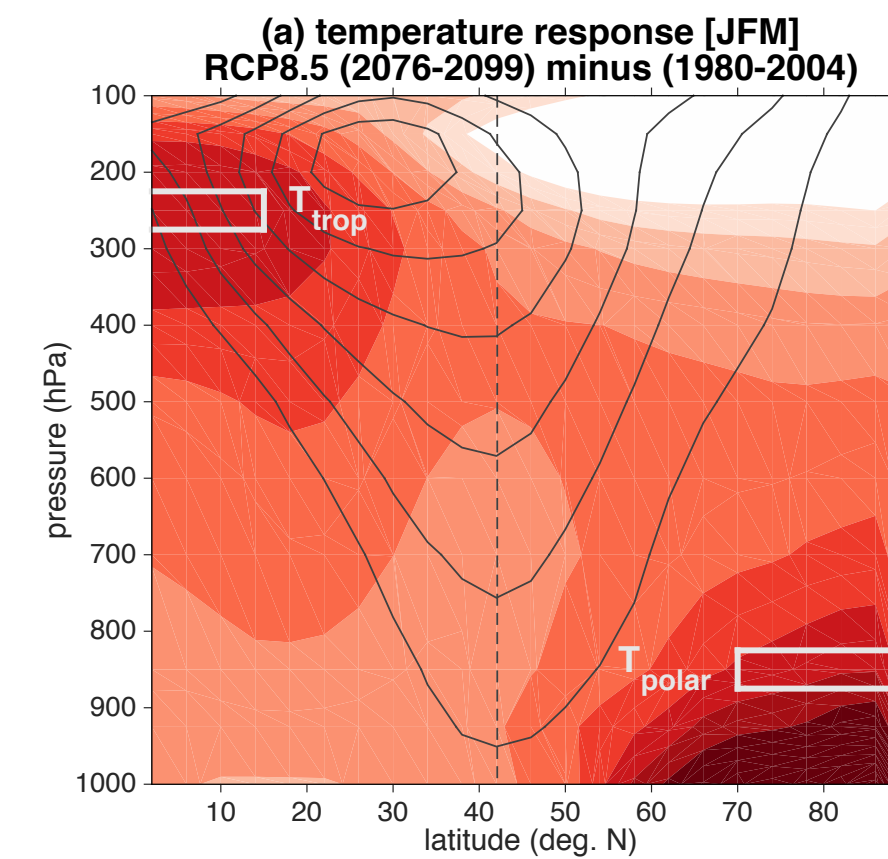
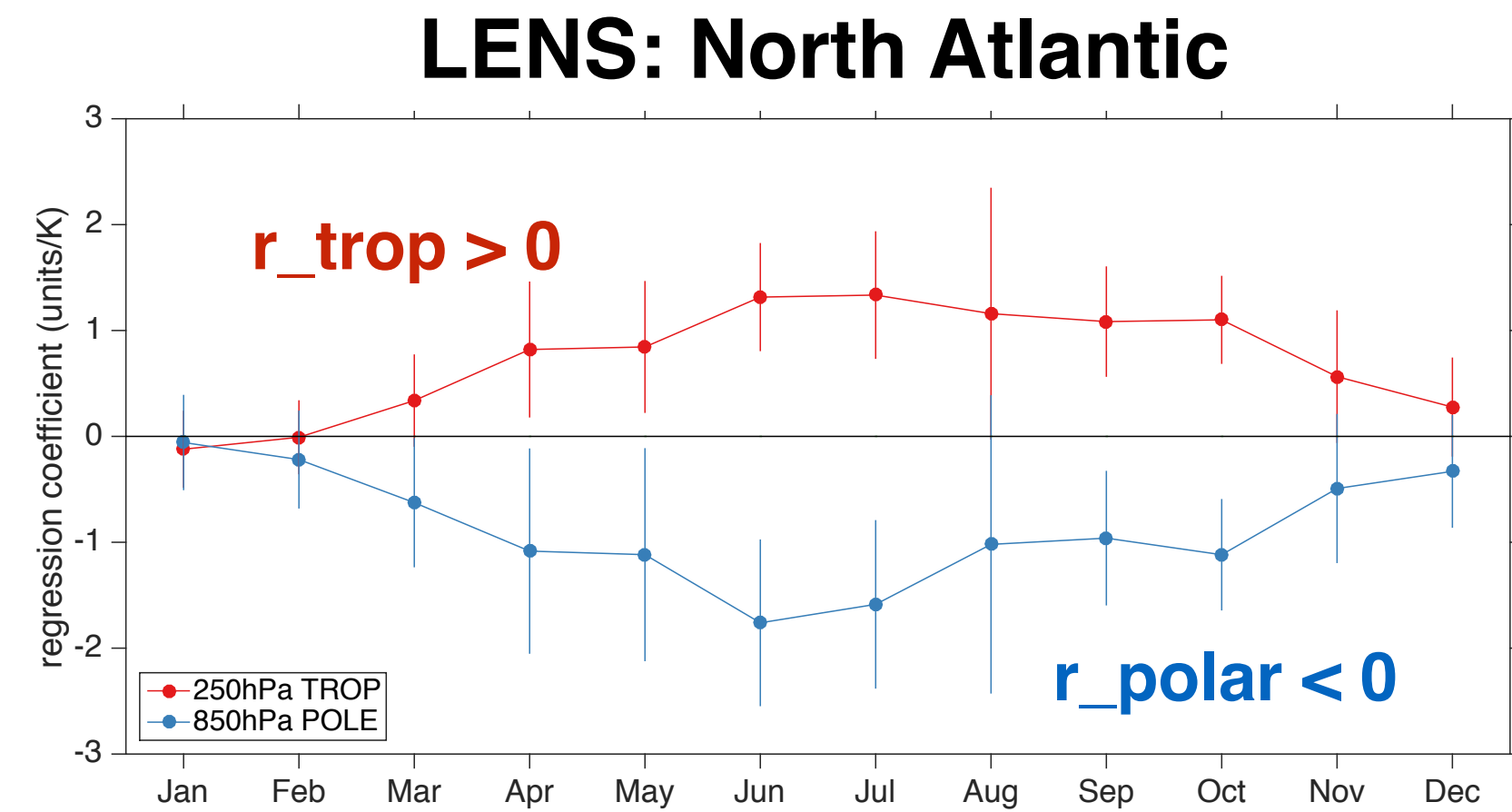


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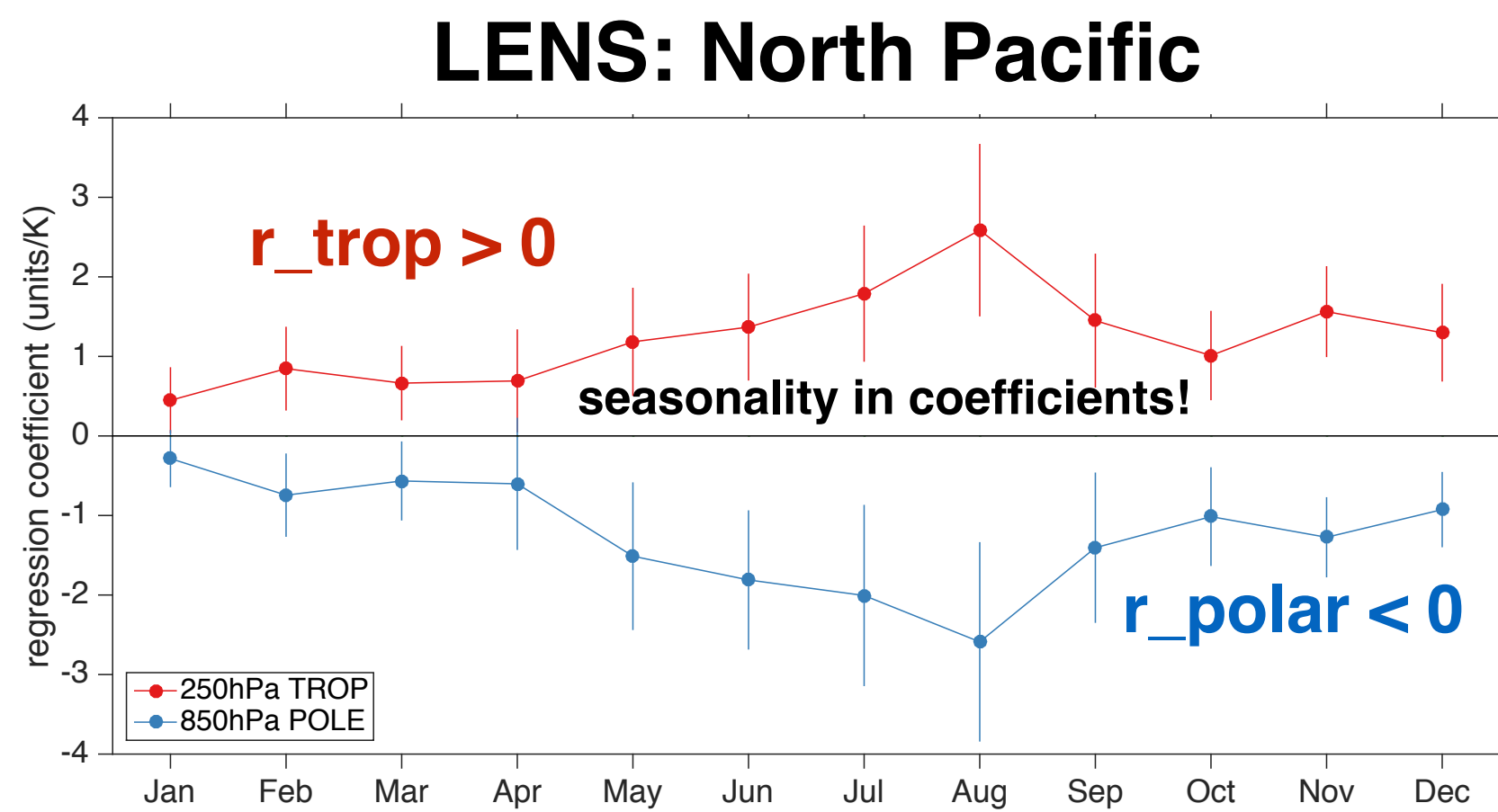
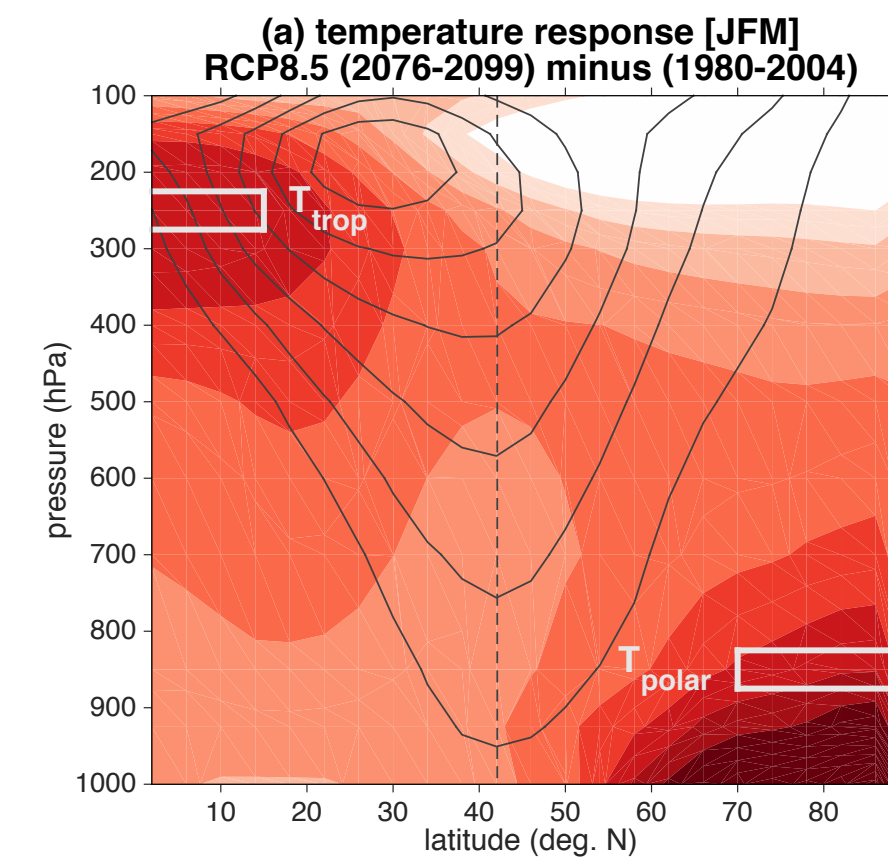
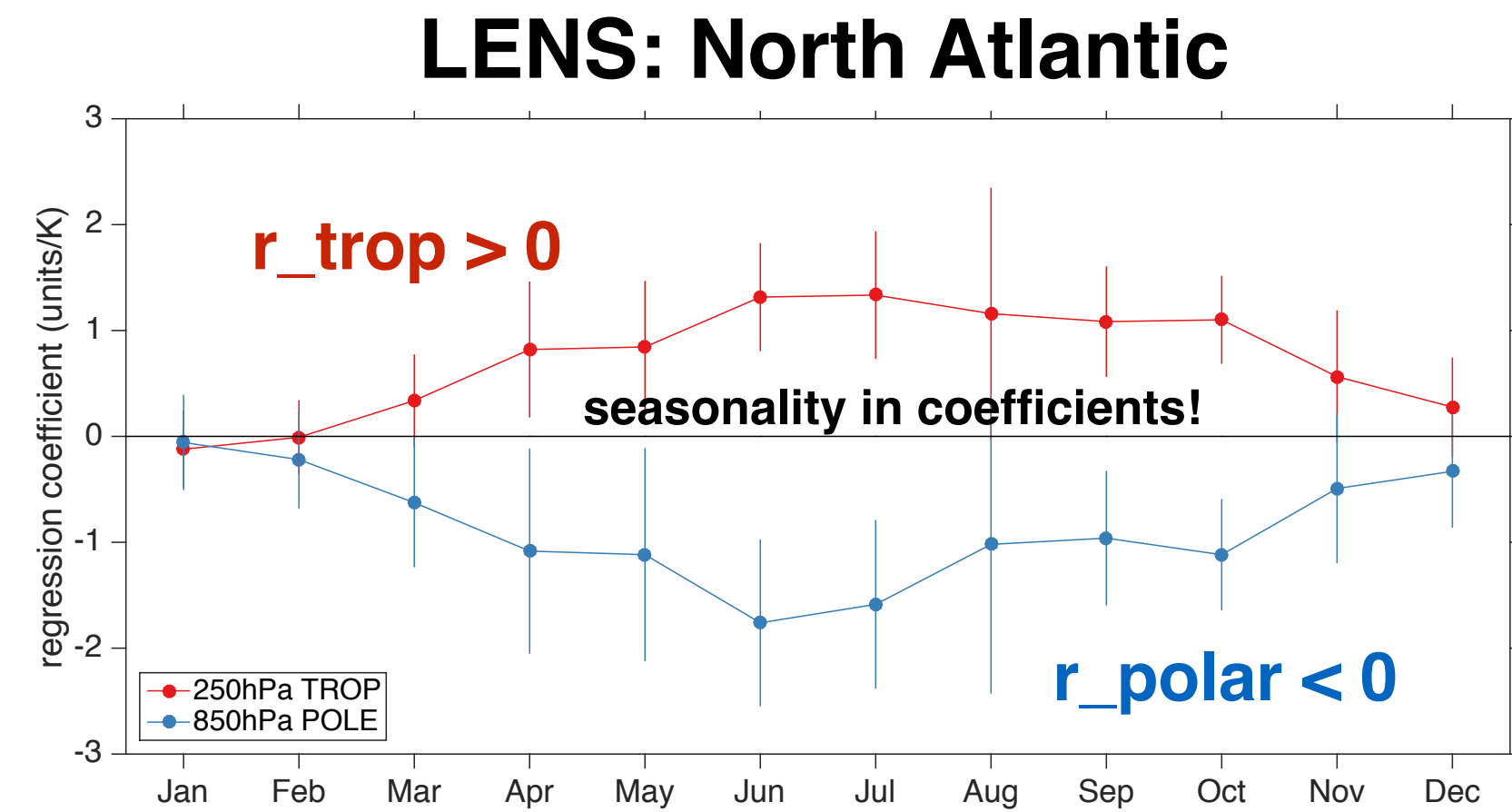


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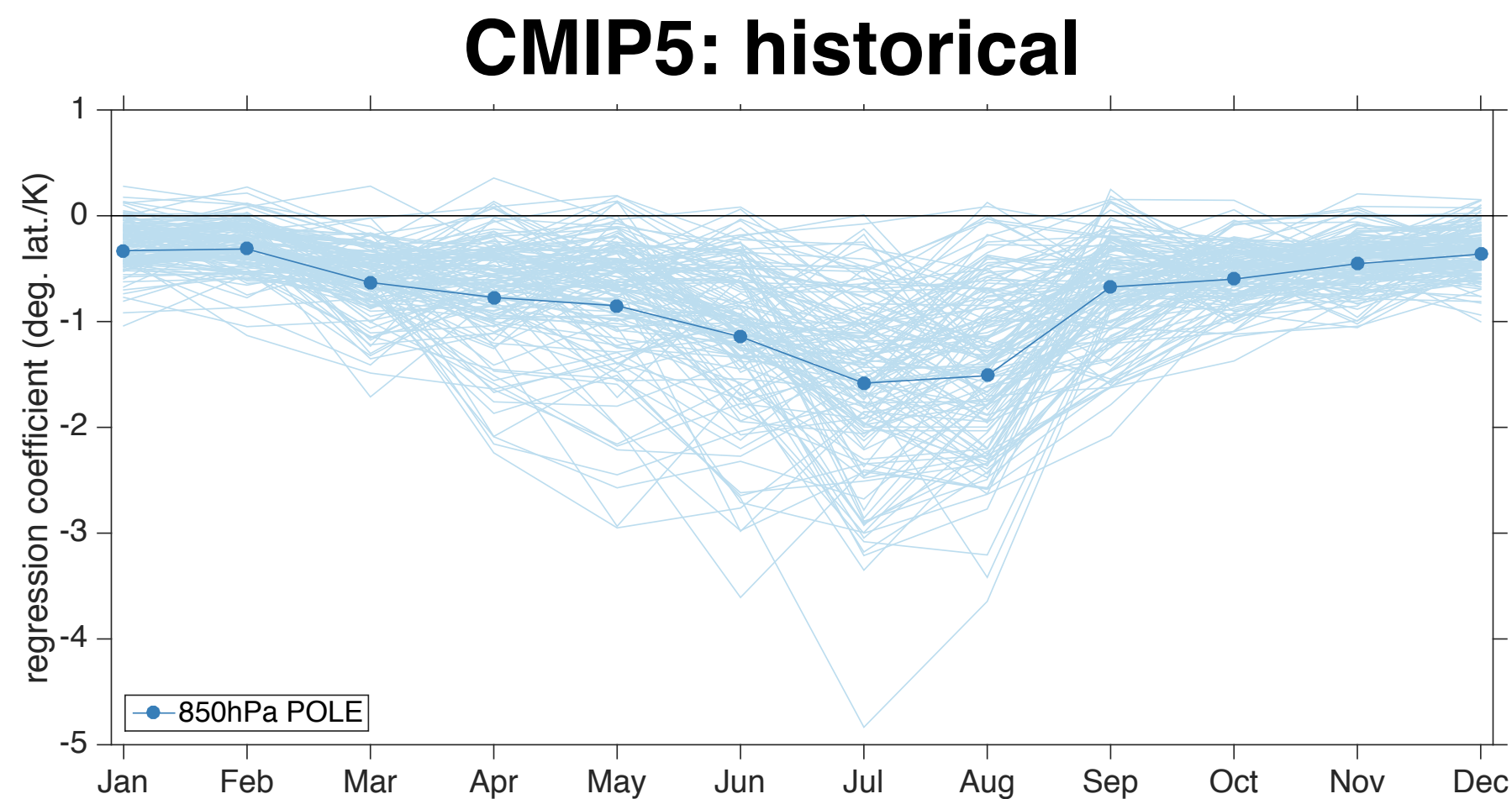
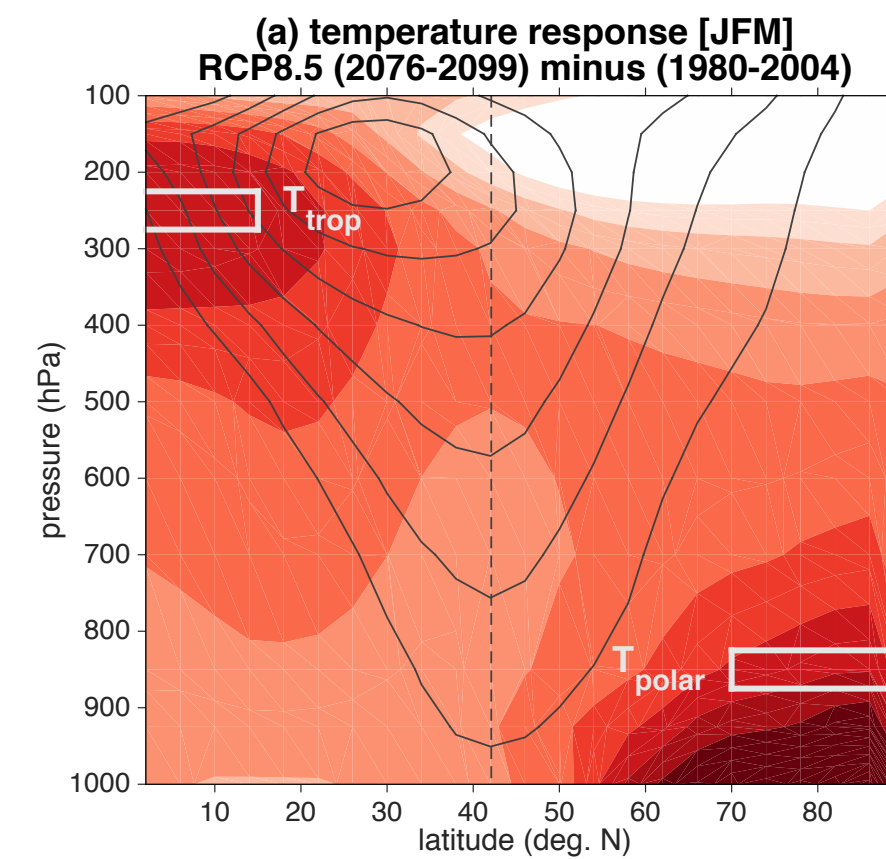
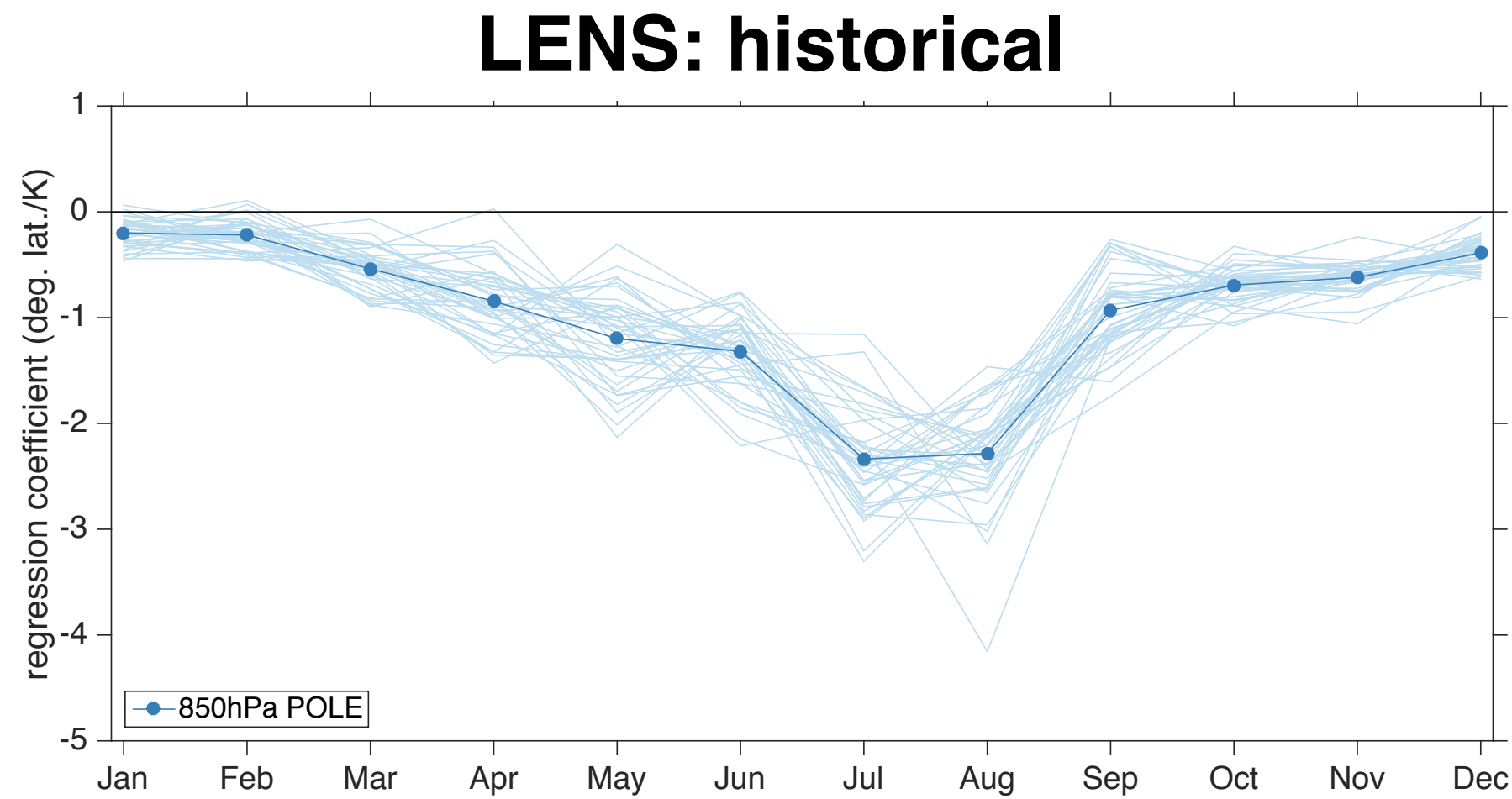


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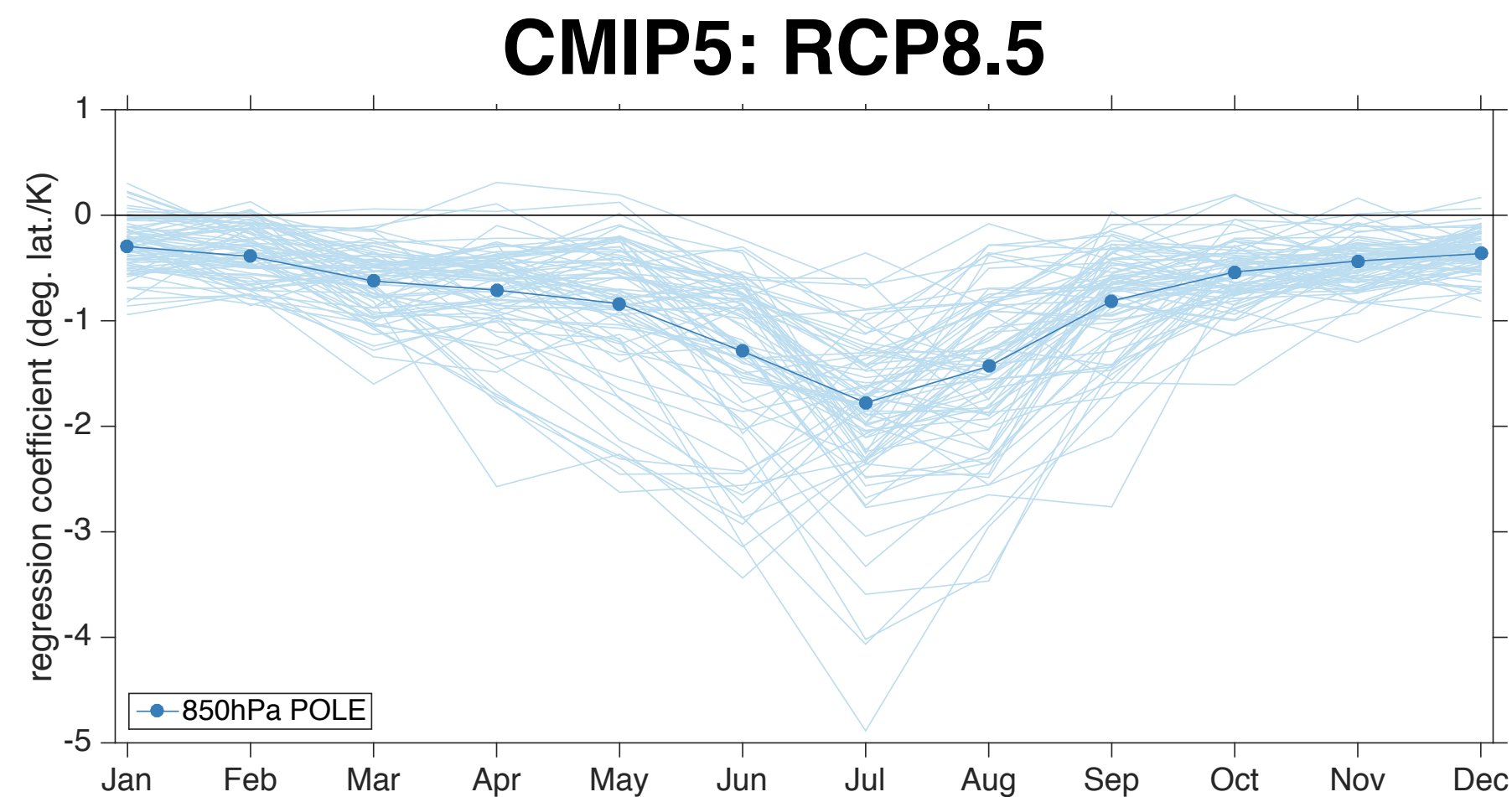
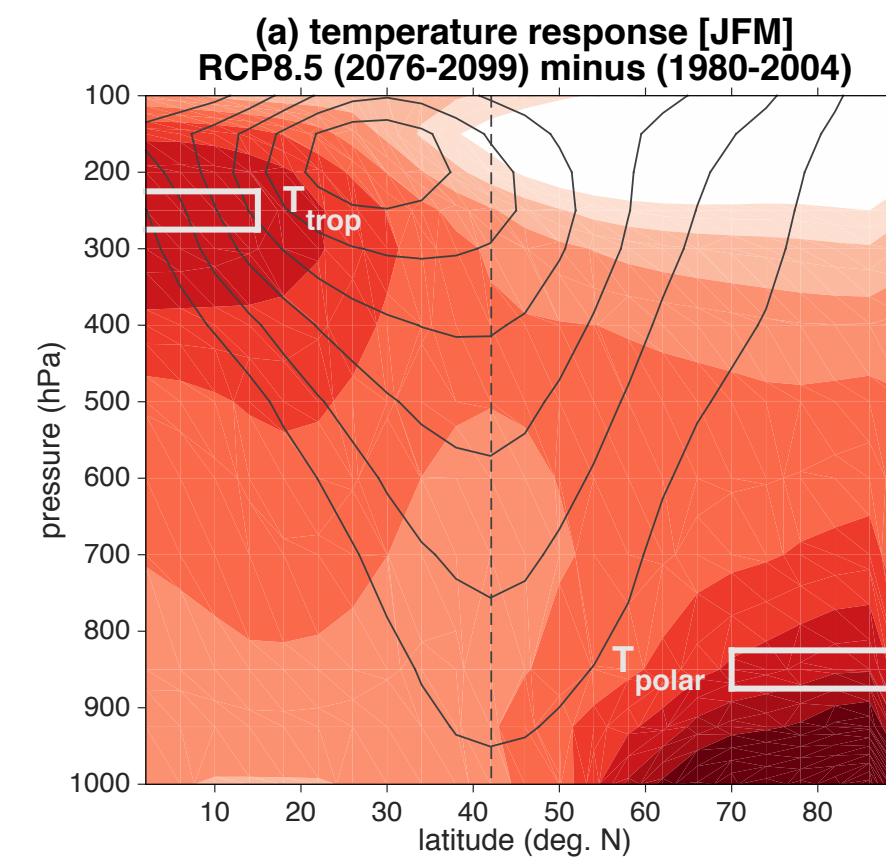
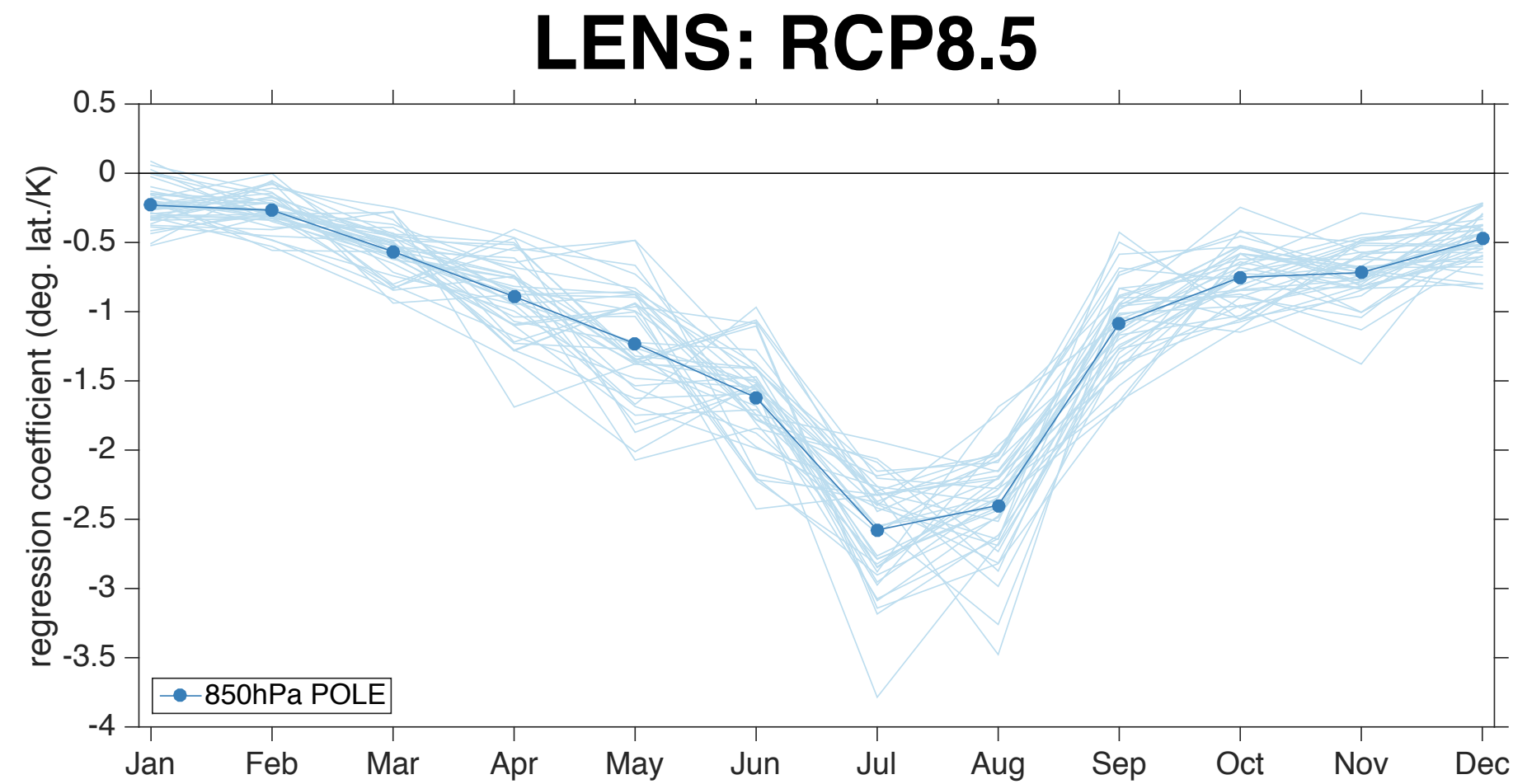
North Pacific: seasonal sensitivity of jet position



Time evolution within simulations

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Aside: collinearity

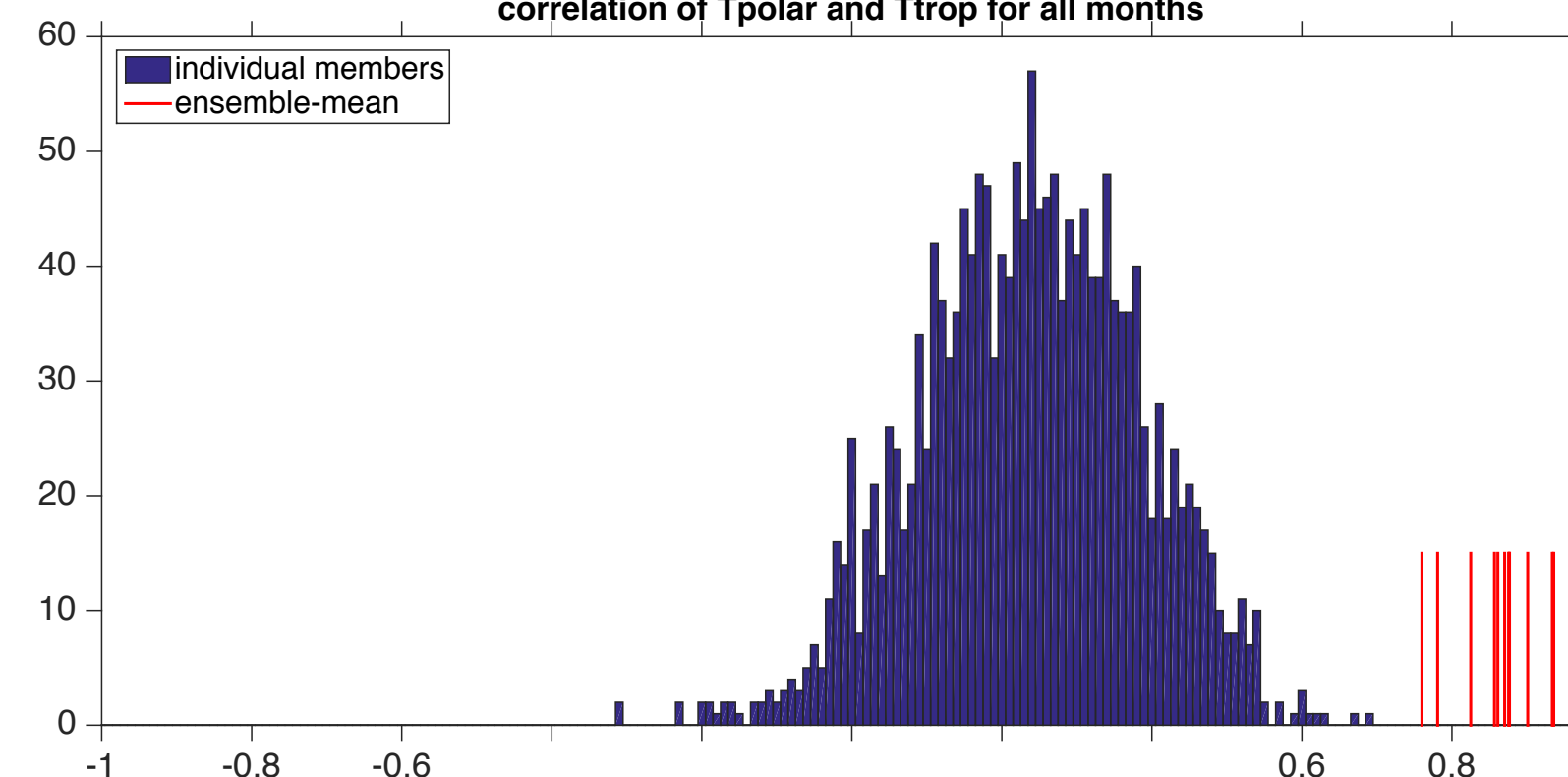
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- T_trop and T_polar are correlated in the future simulations, which can be a major problem for linear regression
- So far, similar results are obtained in different data sets, not detrended vs. detrended, LENS vs. CMIP5, and historical simulations

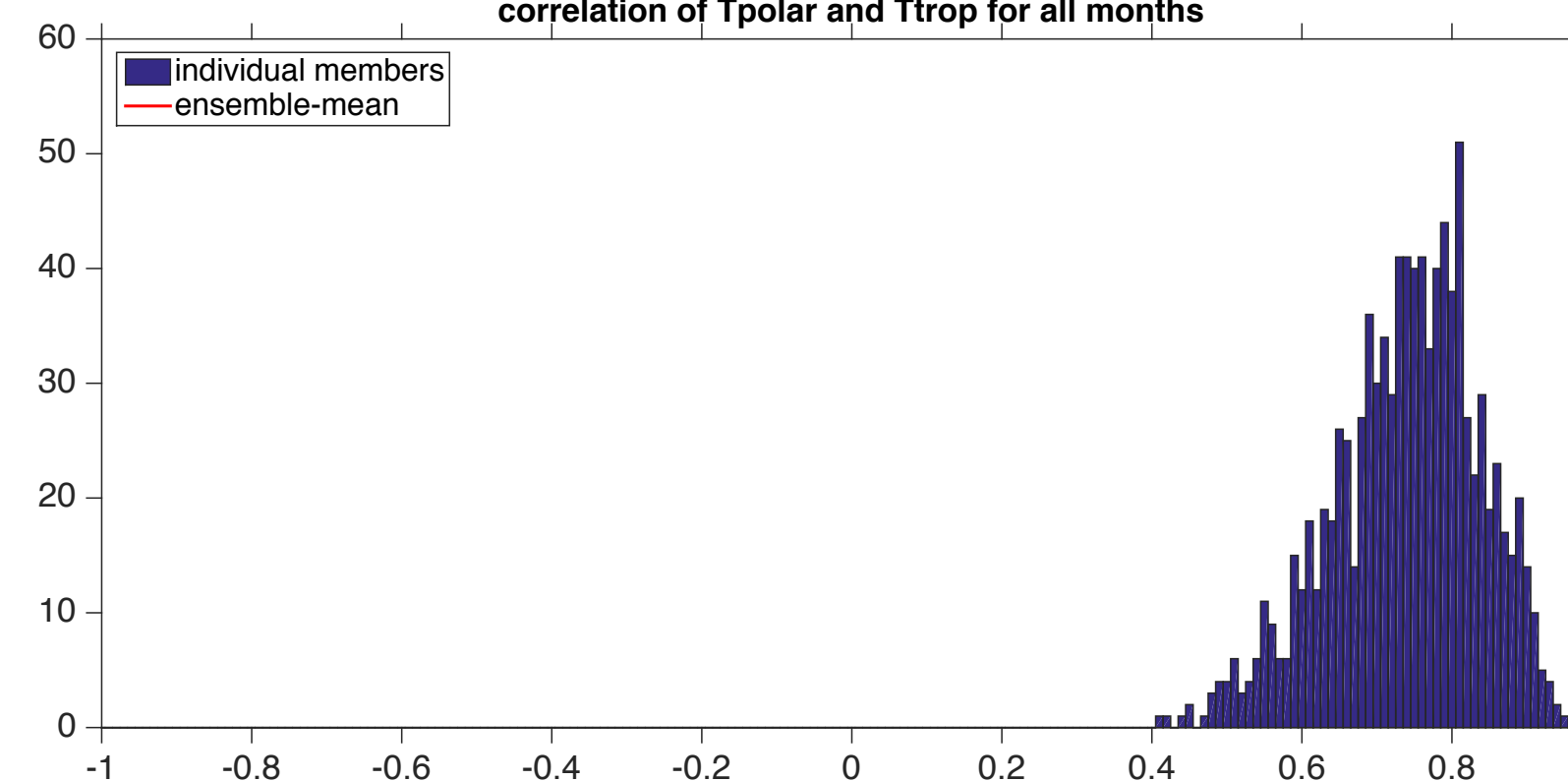
CMIP5: historical

correlation of T_{polar} and T_{trop} for all months

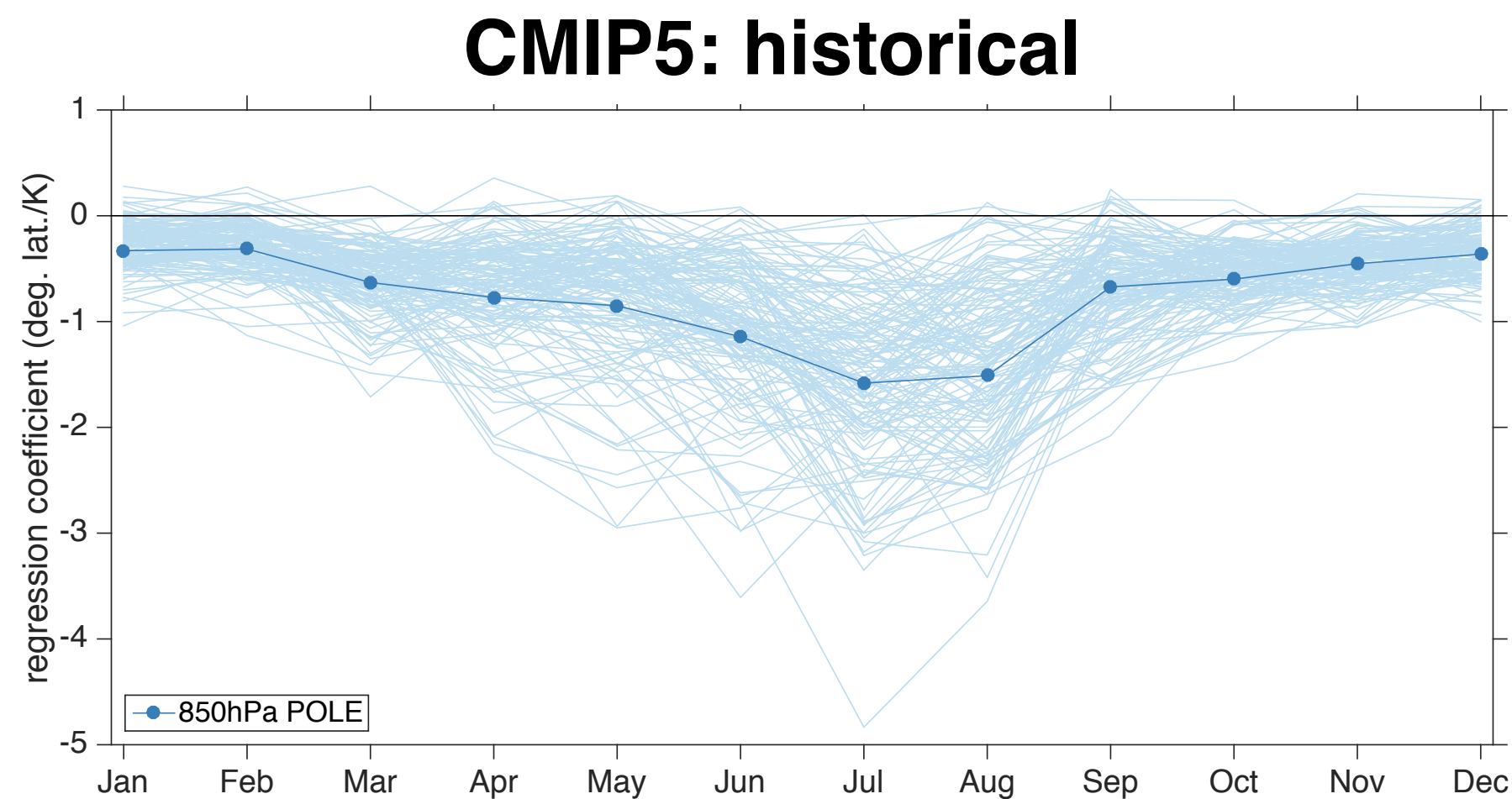
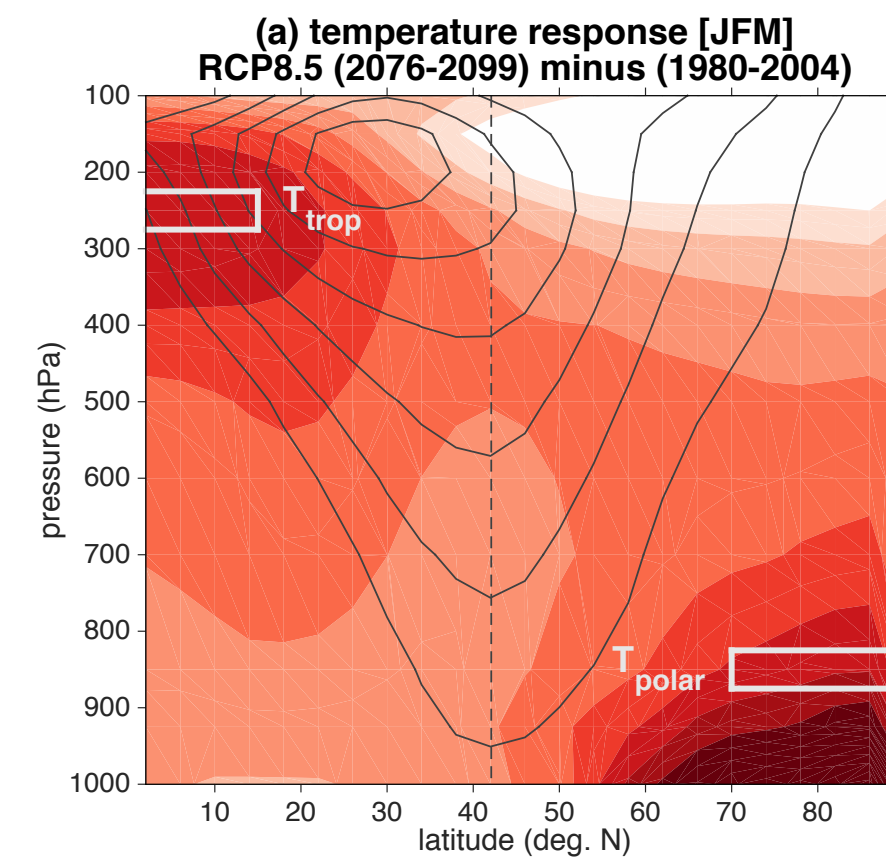
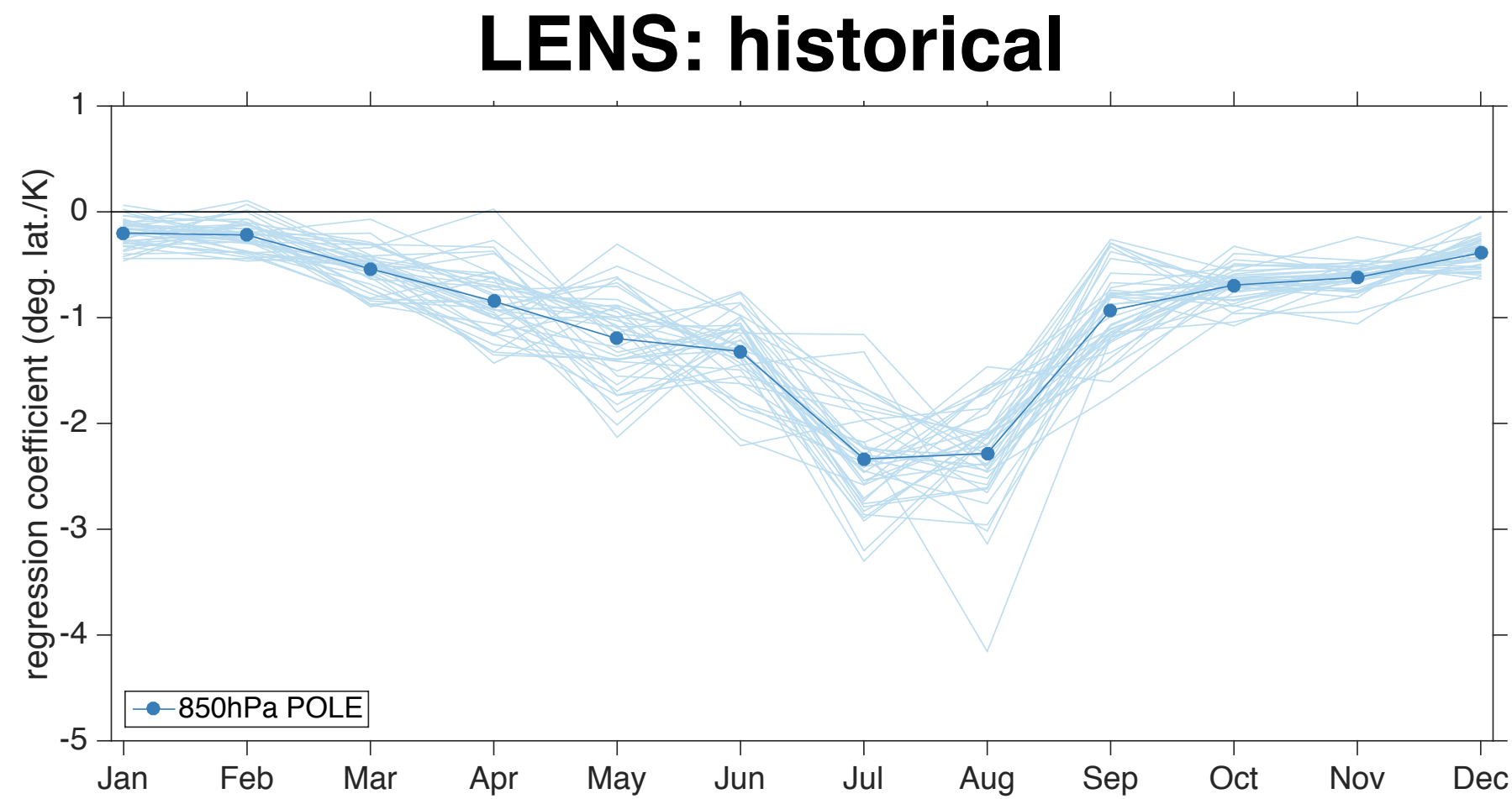


CMIP5: RCP8.5

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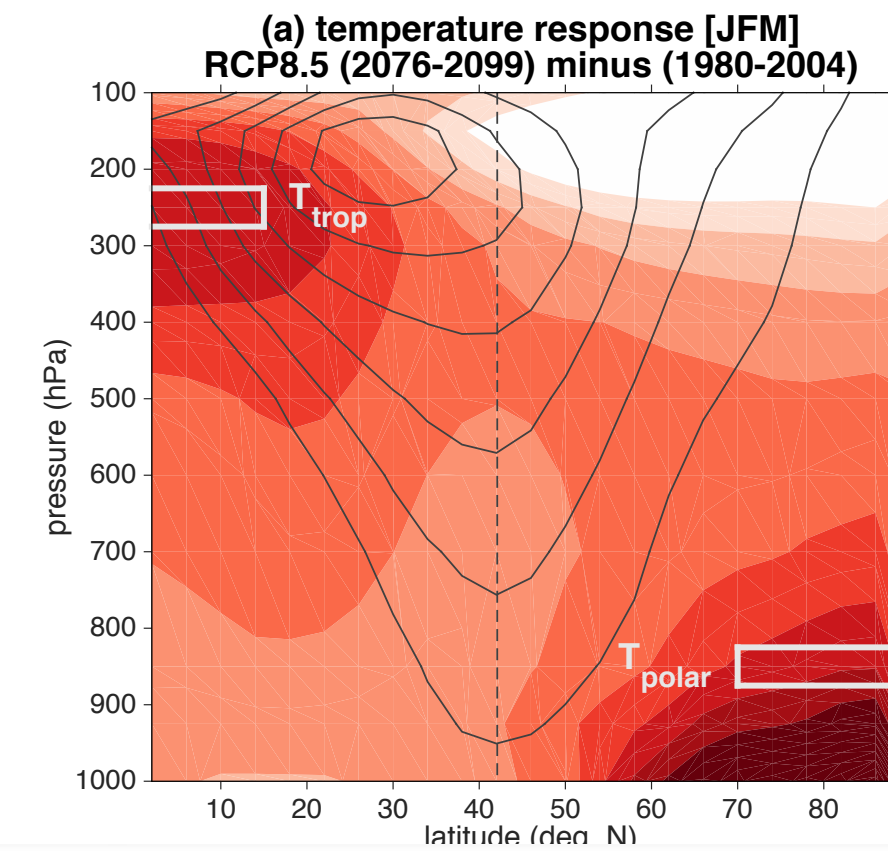
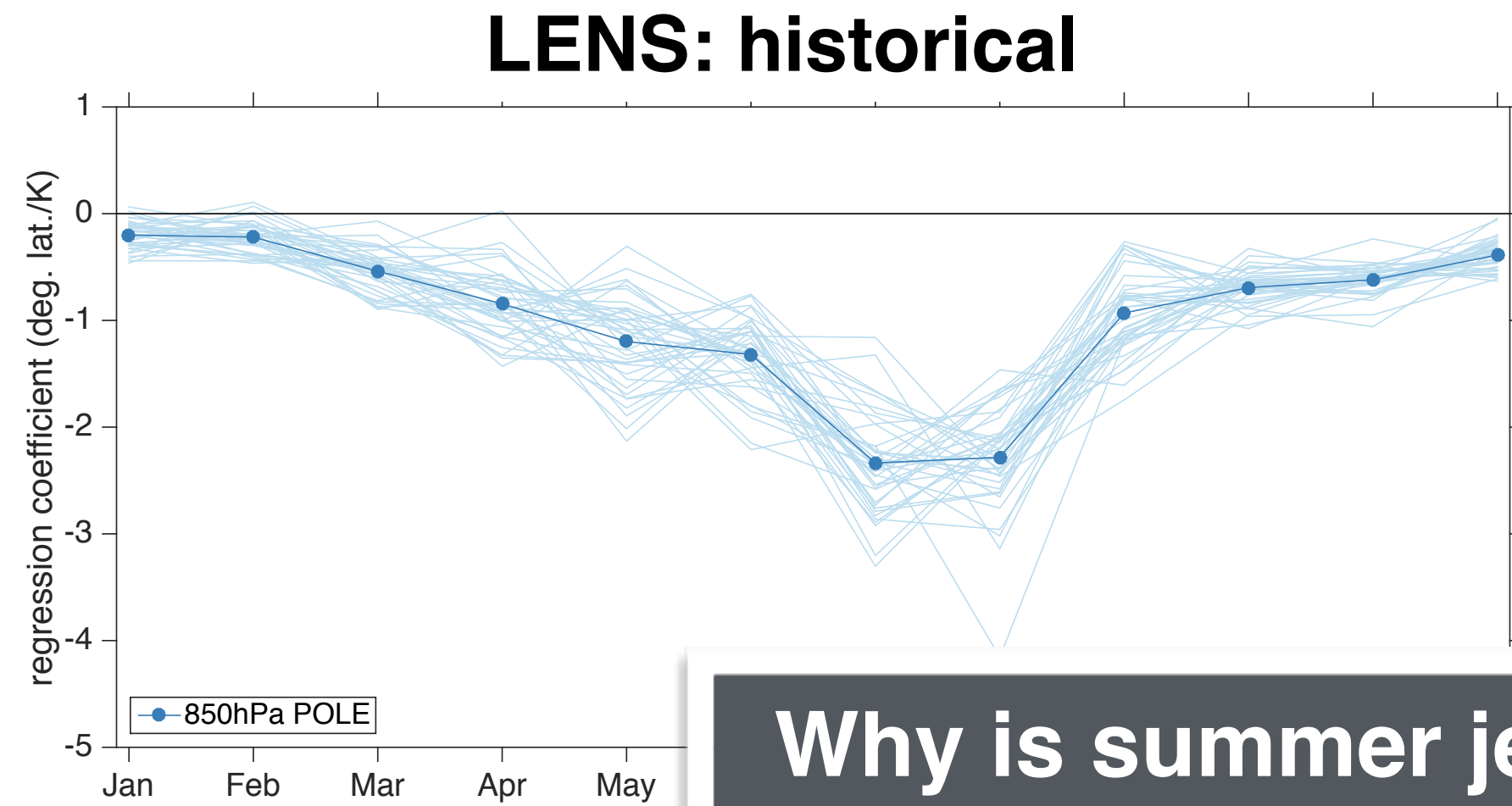
North Pacific: seasonal sensitivity of jet position



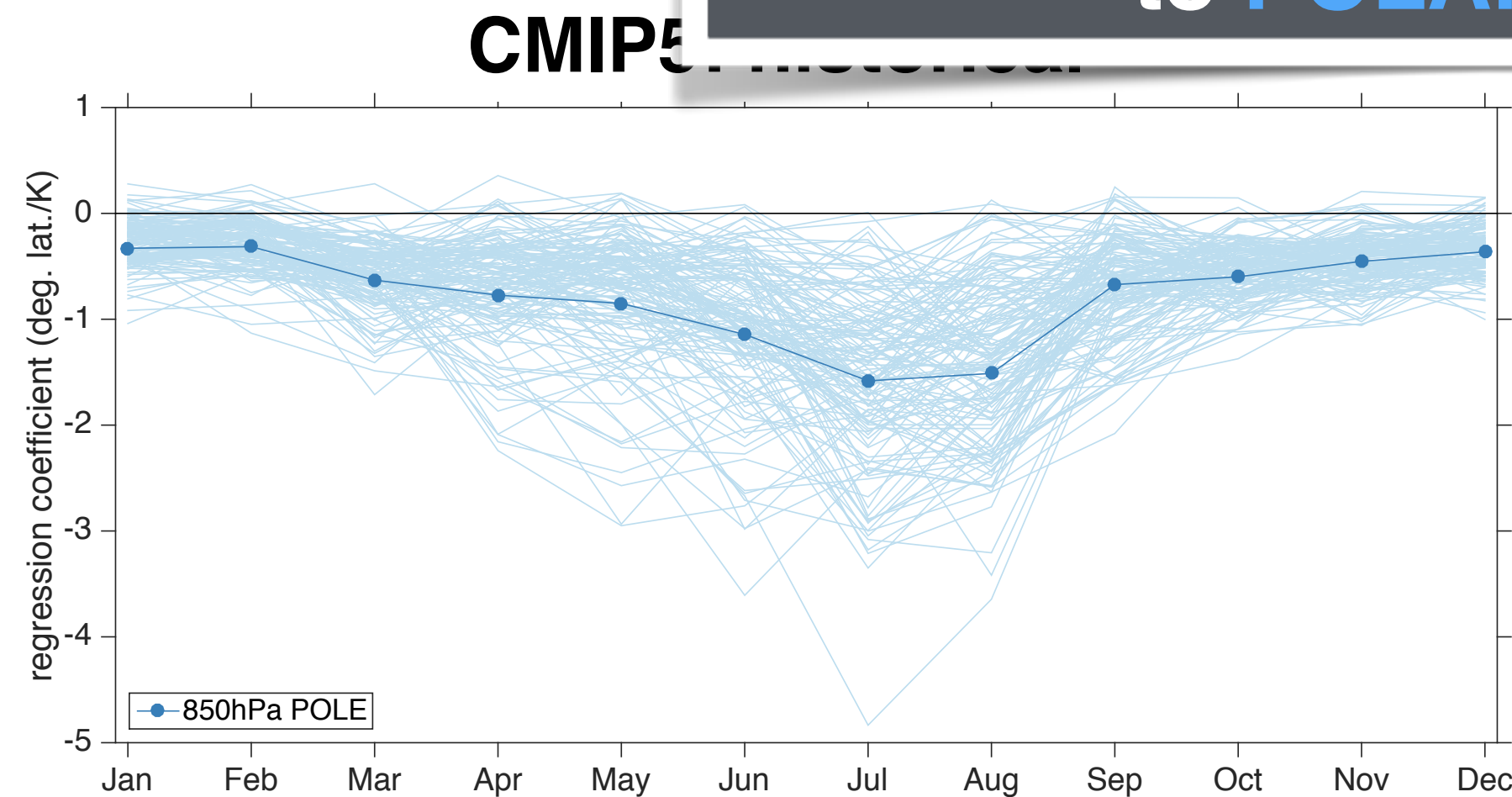
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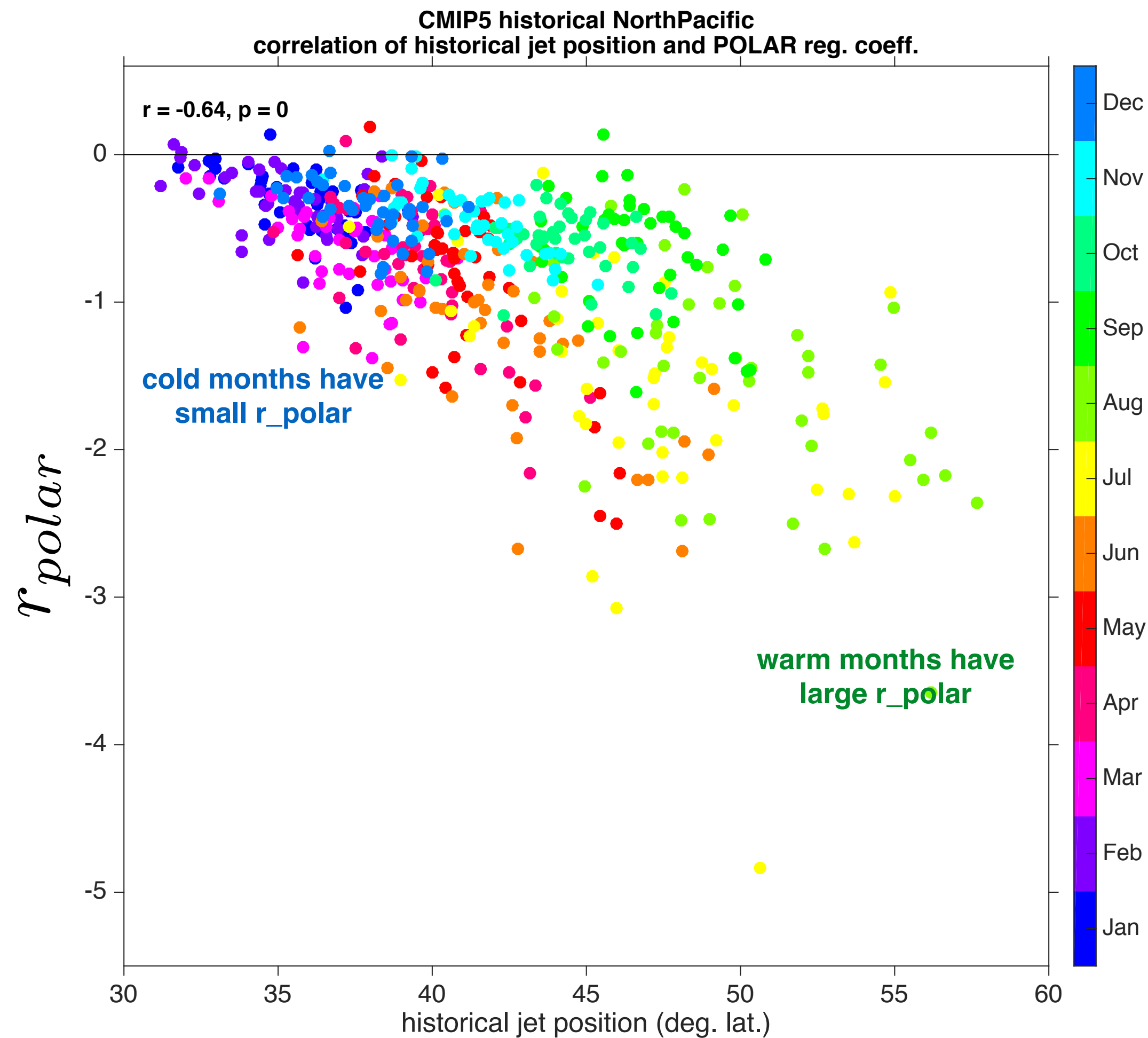
Why is summer jet the most sensitive to **POLAR** warming?



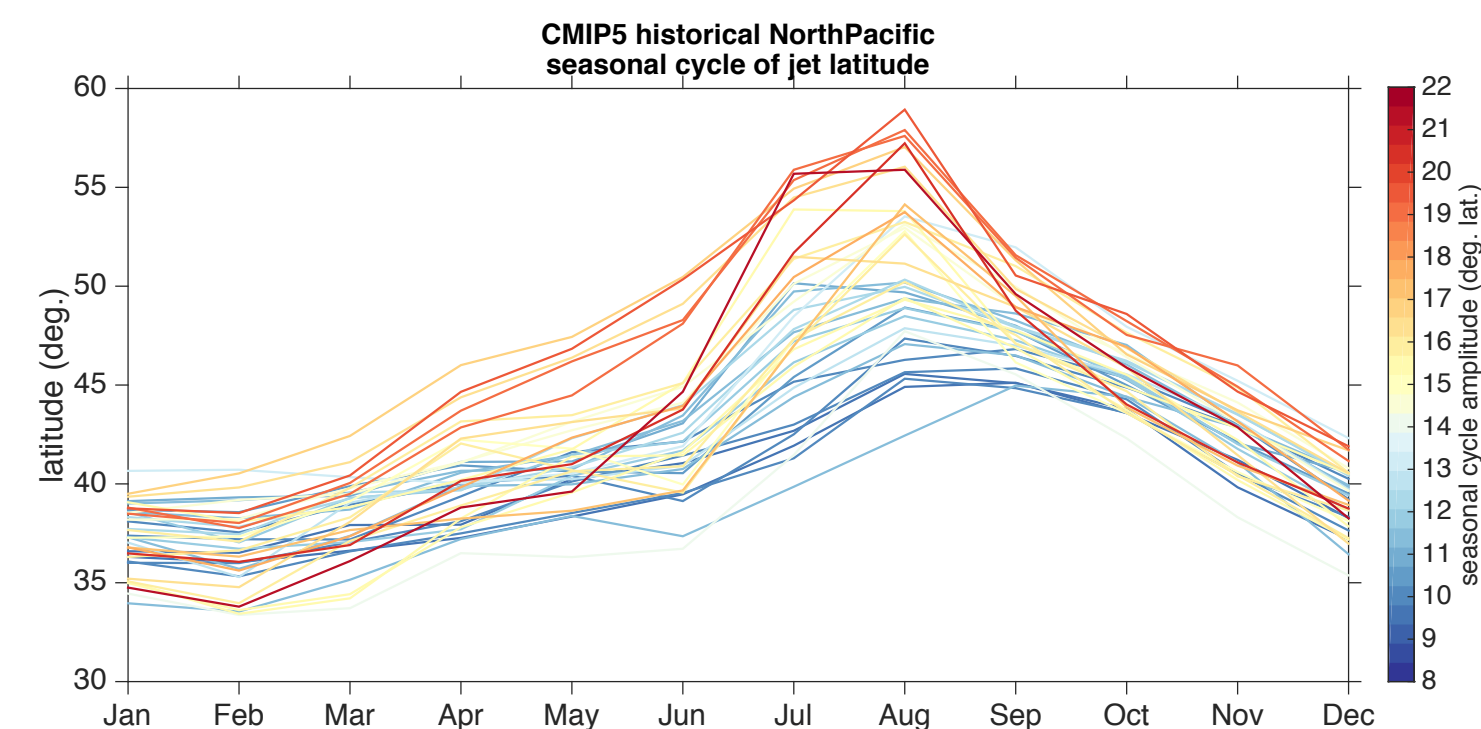
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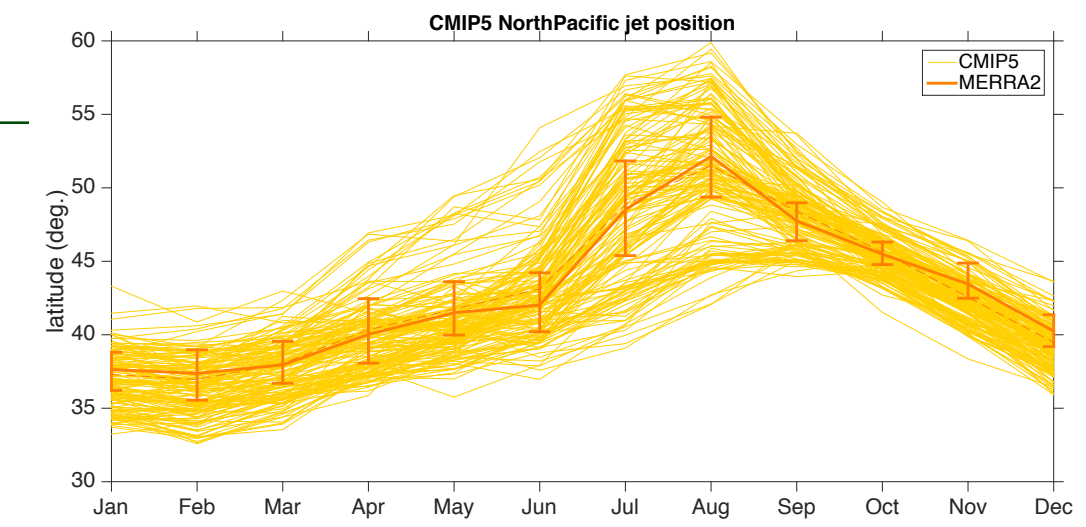
Seasonality of the jet shift w/ POLAR warming



The seasonal migration of the jet impacts the jet shift's seasonal sensitivity



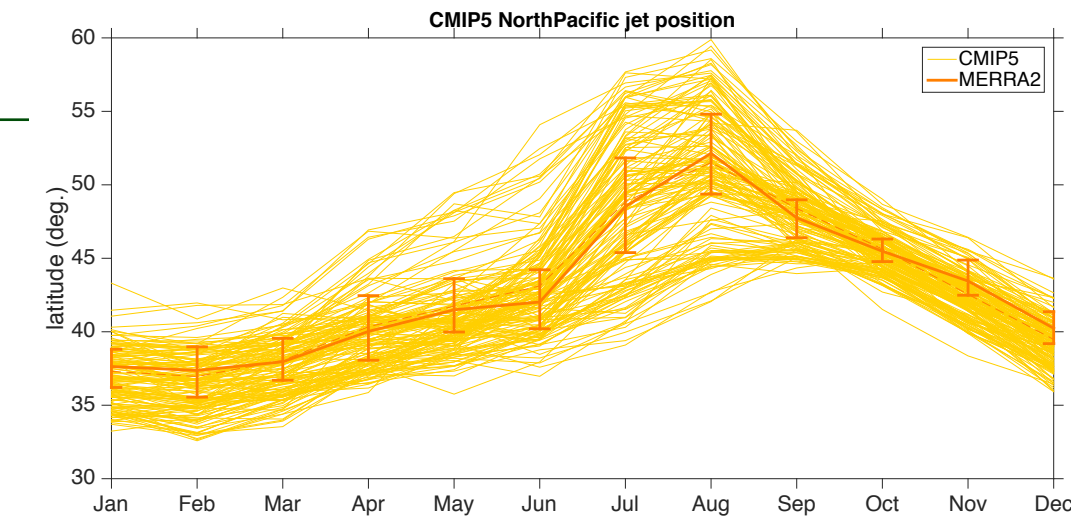
Why might there be seasonal sensitivity?



hypotheses and framing based on multiple studies:

e.g. Garfinkel et al. 2013; Simpson et al. 2010; O'Rourke & Vallis 2013; Peng et al. 1995, 1997; Newman & Sardeshmukh 1998 ...

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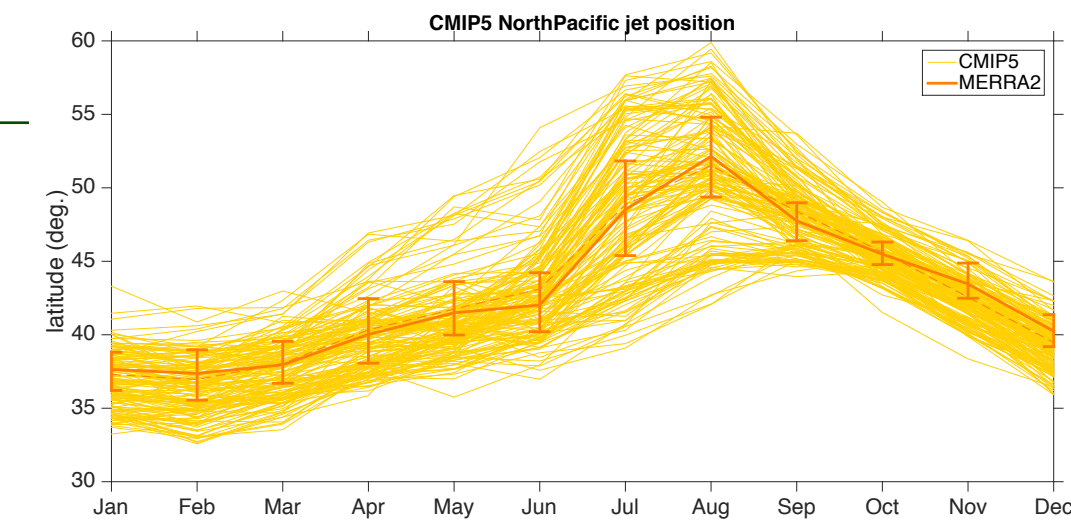


Hypothesis 1: The climatological jet-stream is the furthest poleward in summer and the furthest equatorward in winter. Thus, the jet is most sensitive to polar warming in summer and tropical warming in winter because it is closer to the warming in these seasons.

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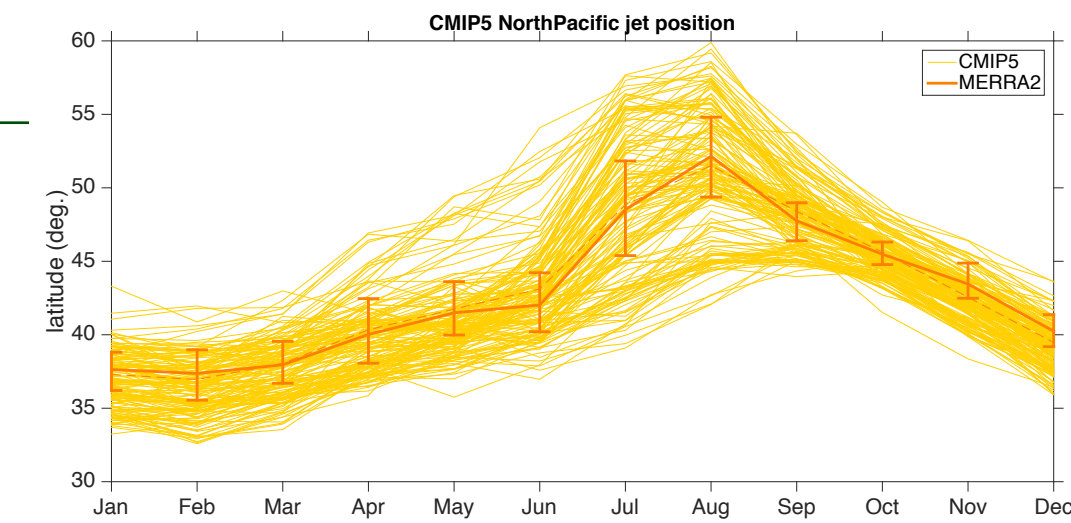
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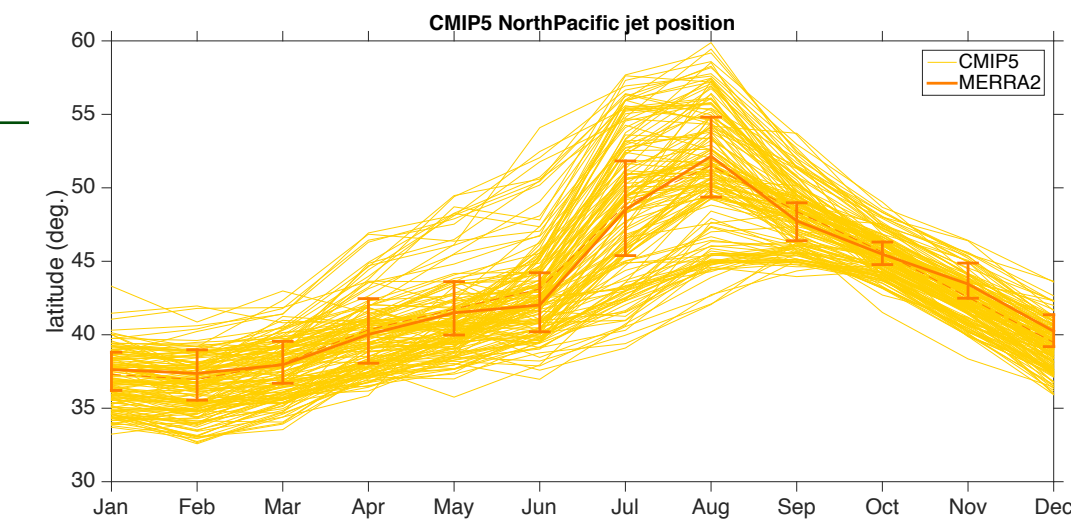
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Hypothesis 4: The shoulder seasons mark the transition between a high-latitude summer regime and low-latitude winter regime, and thus, the jet is most sensitive to forcing in spring and fall when the jet can move over a wider range of latitudes.

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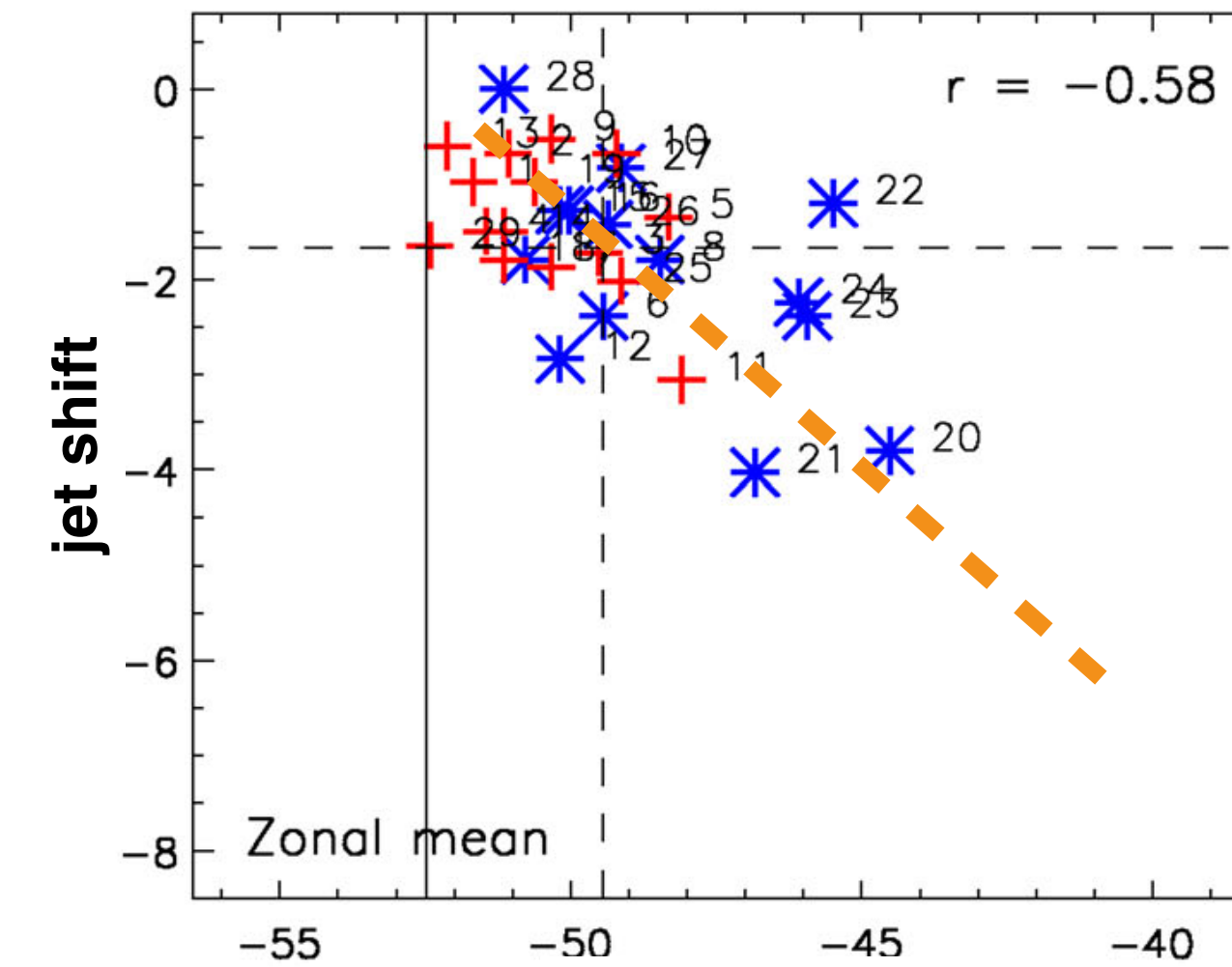
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Historical jet position vs jet shift

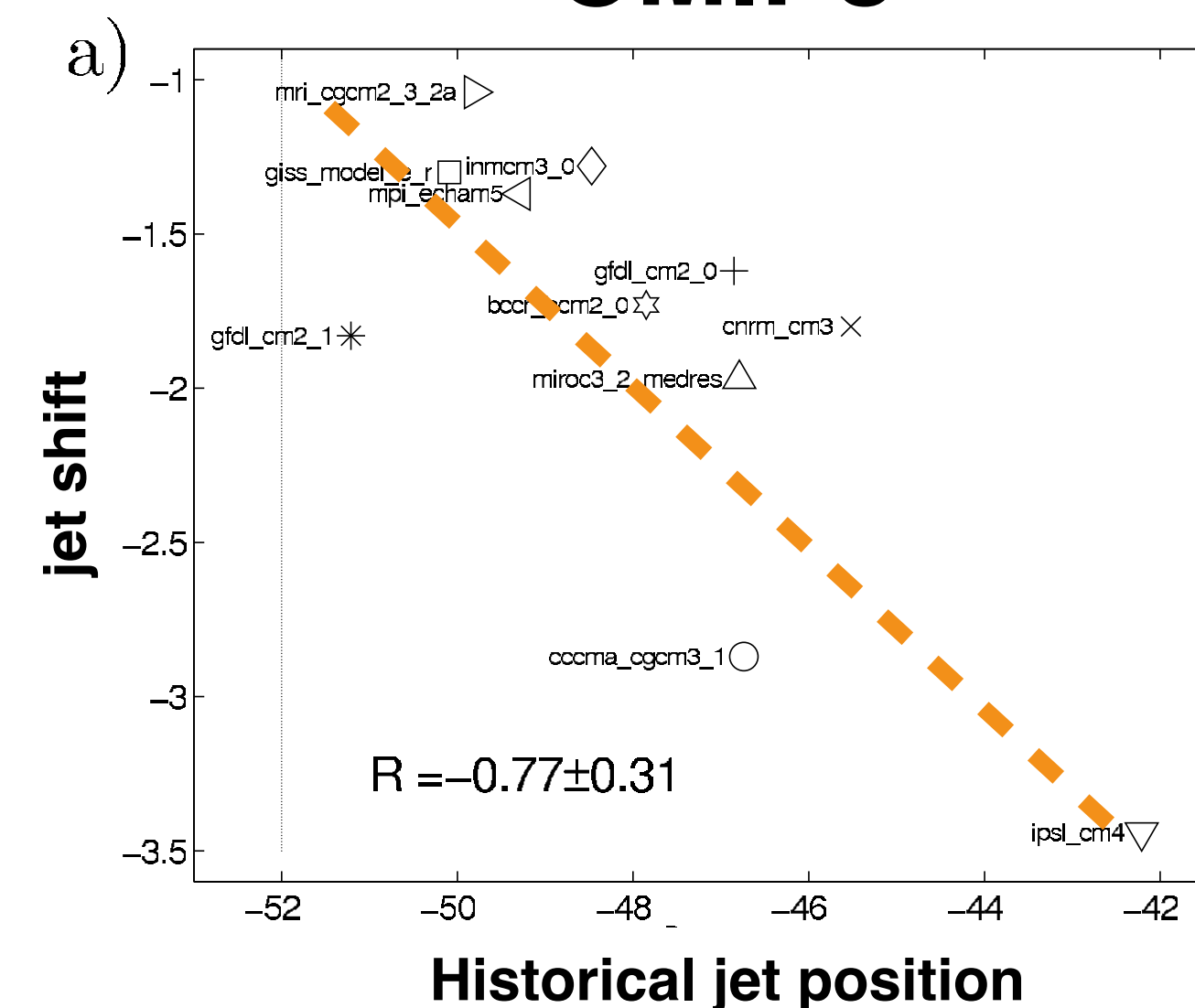
Southern Hemisphere

- idealized and comprehensive GCM evidence that higher latitude jets shift less
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- typically this relationship has been investigated for the annual mean or the winter season...

CMIP5



CMIP3



Bracegirdle et al. (2013); JGR
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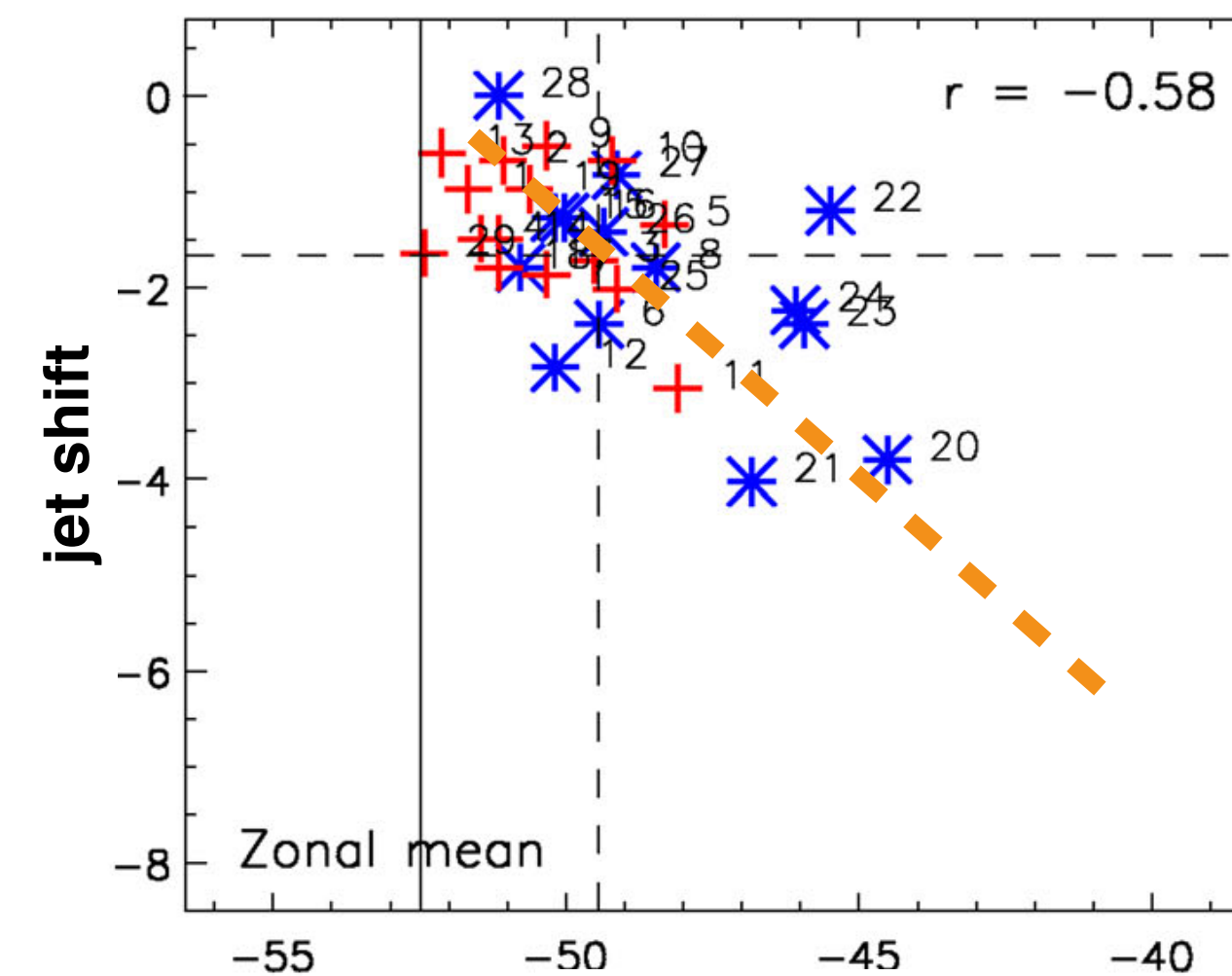
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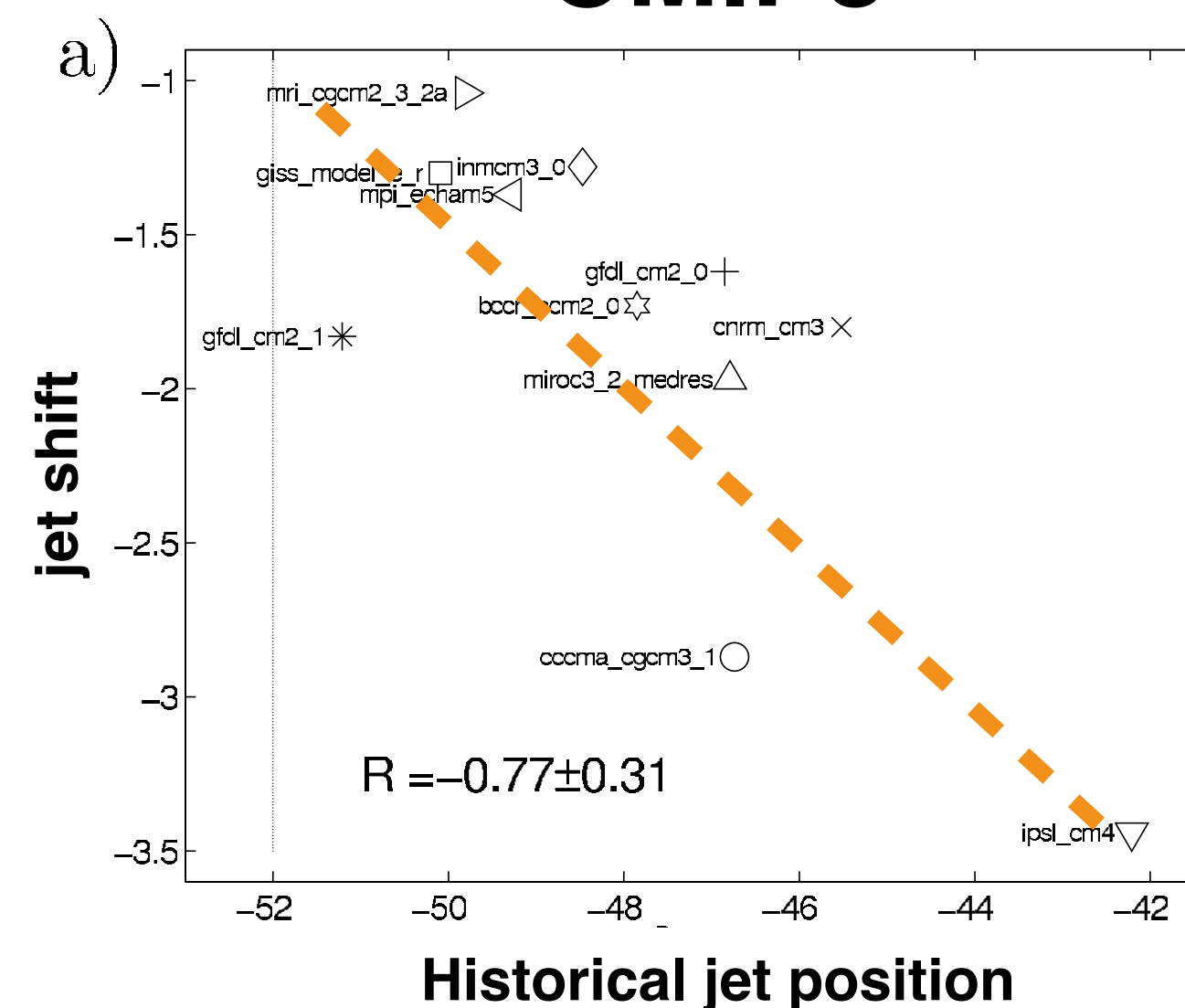
what about other seasons?

Bracegirdle et al. (2013); JGR
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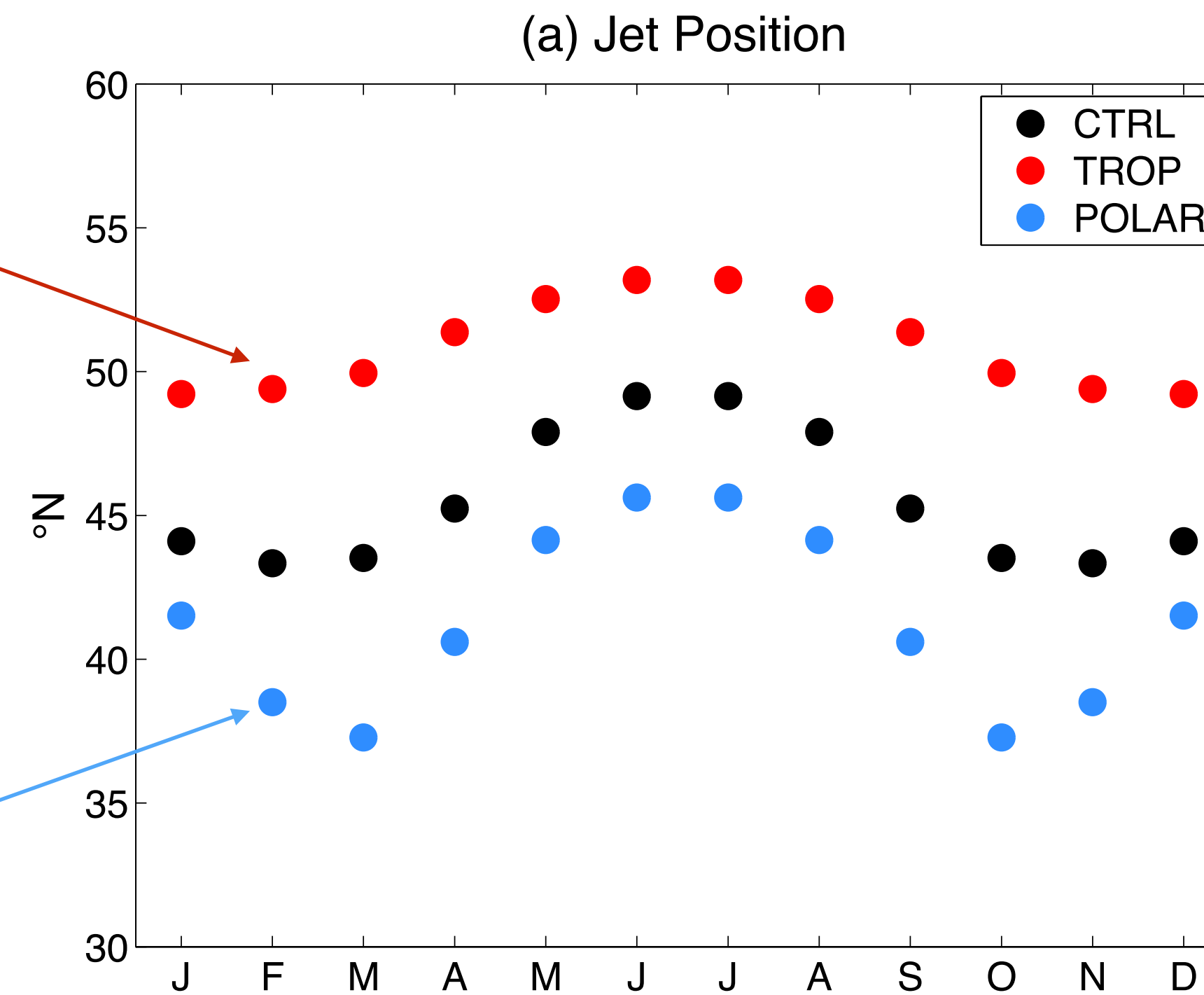
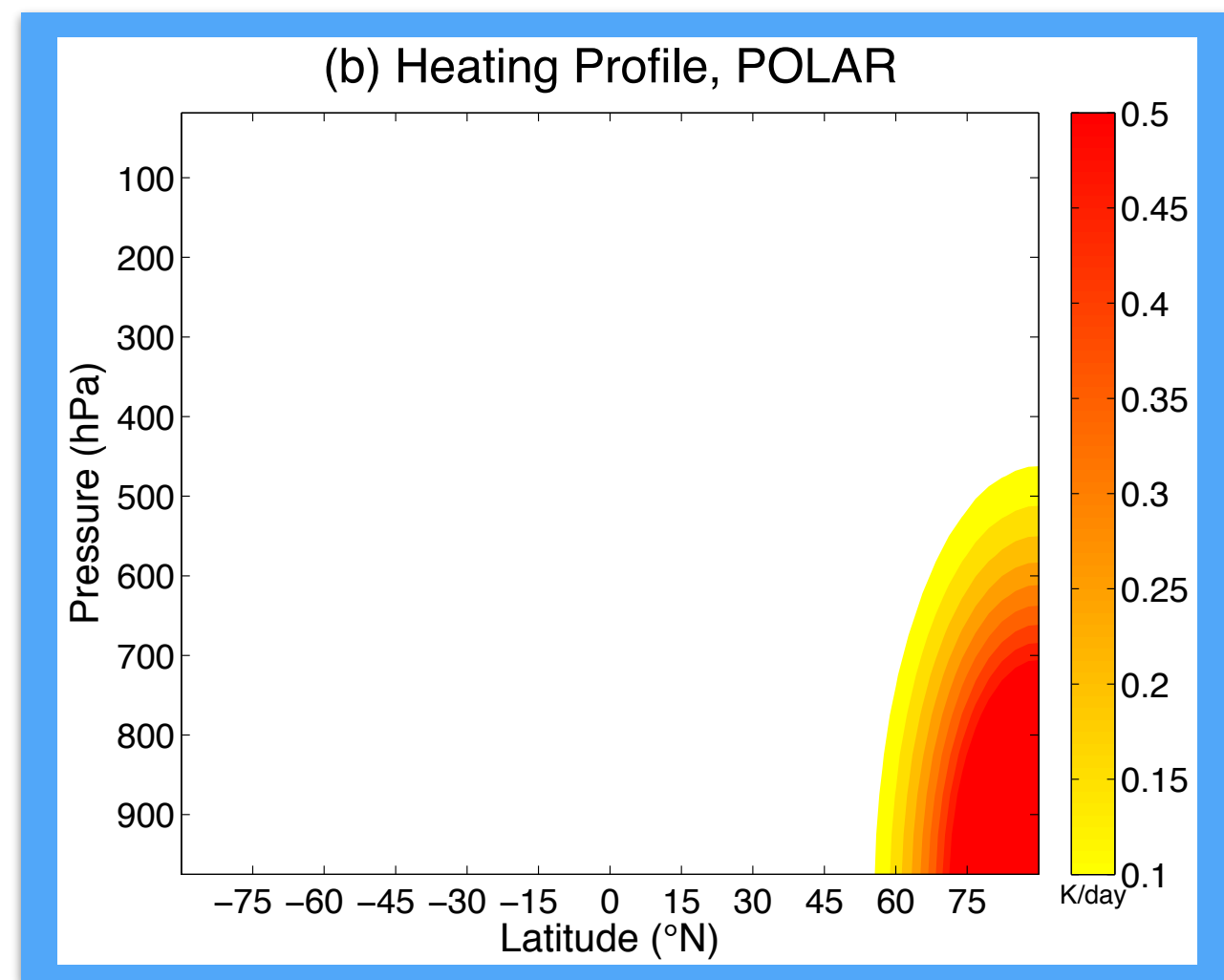
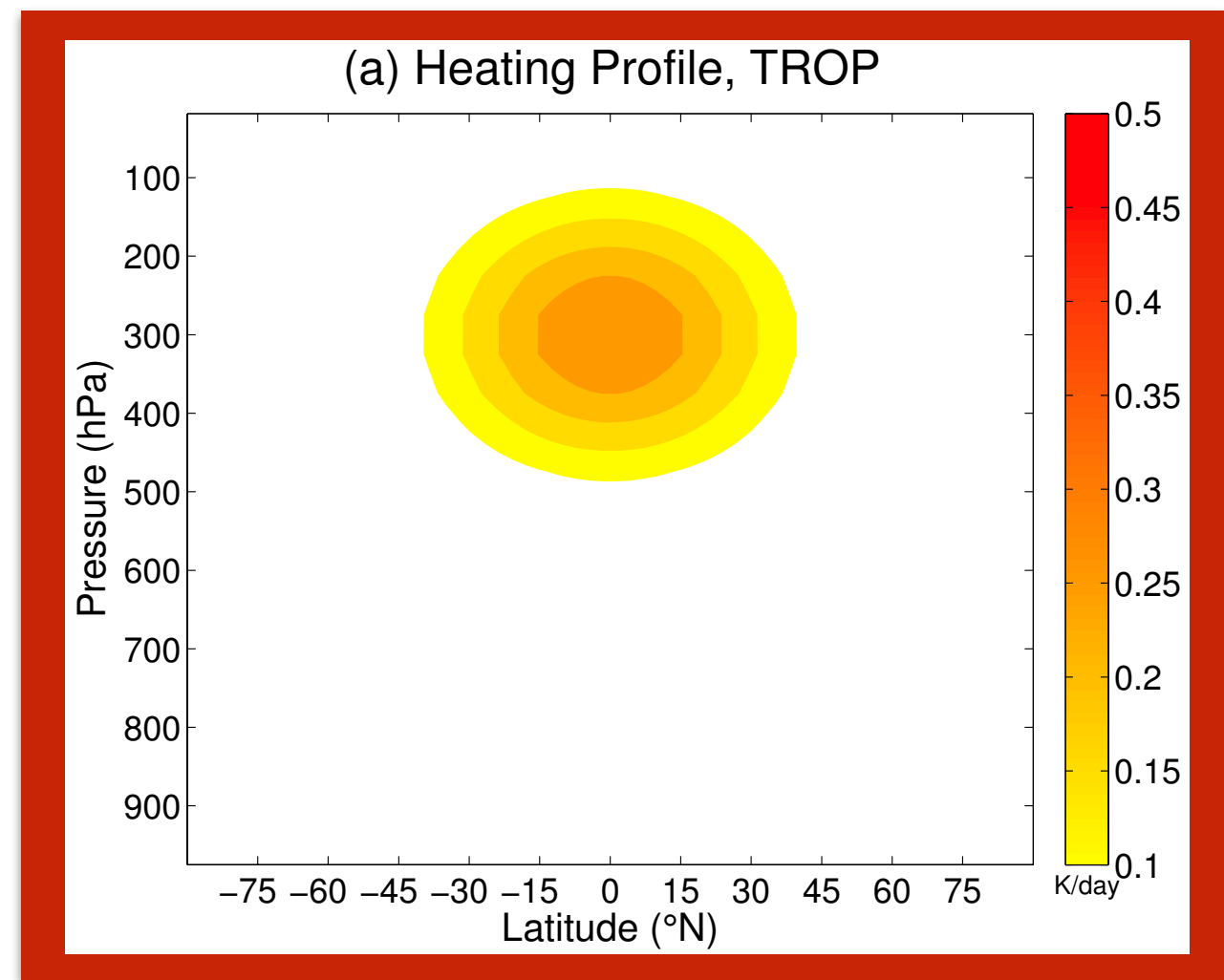
CMIP5



CMIP3



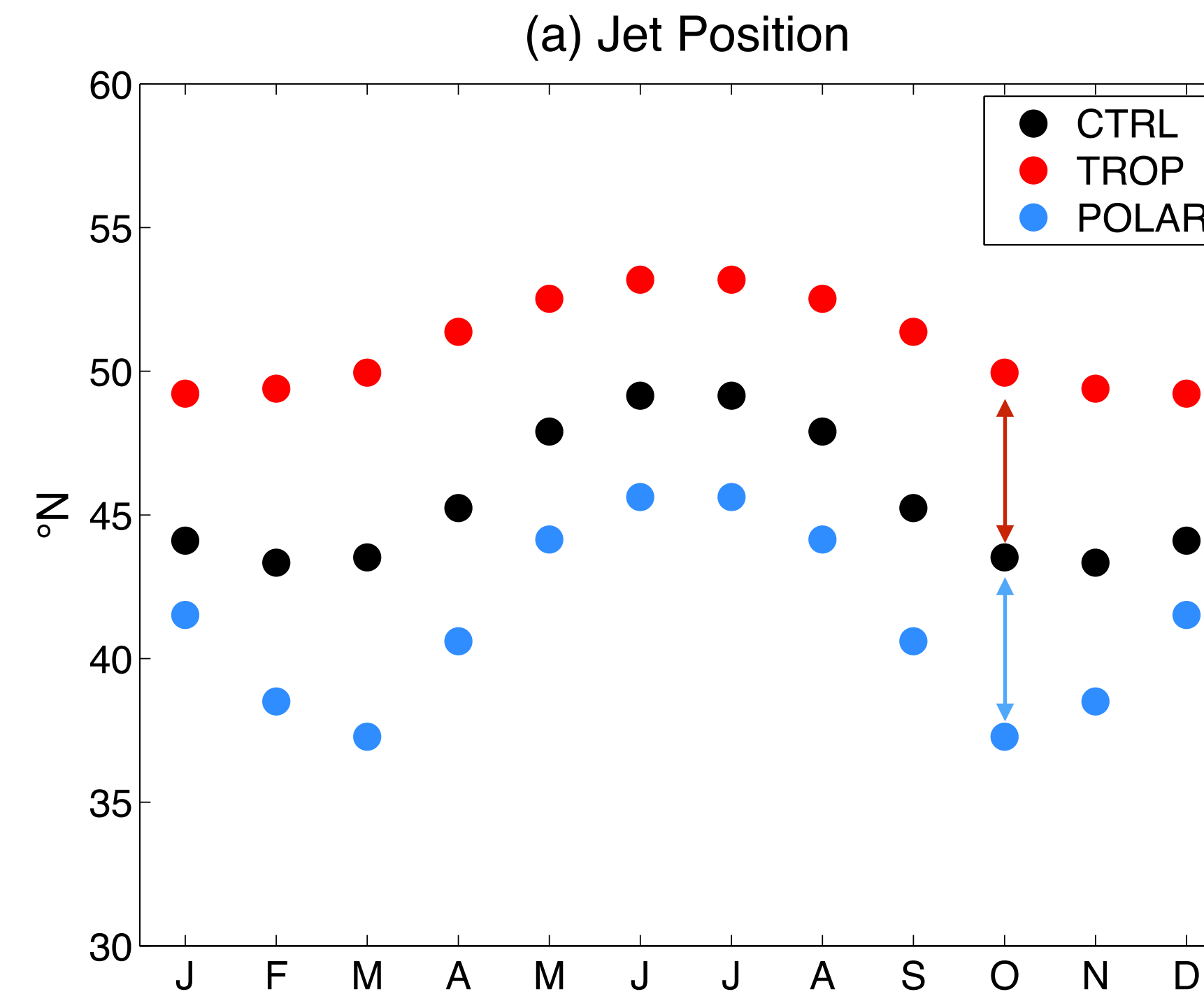
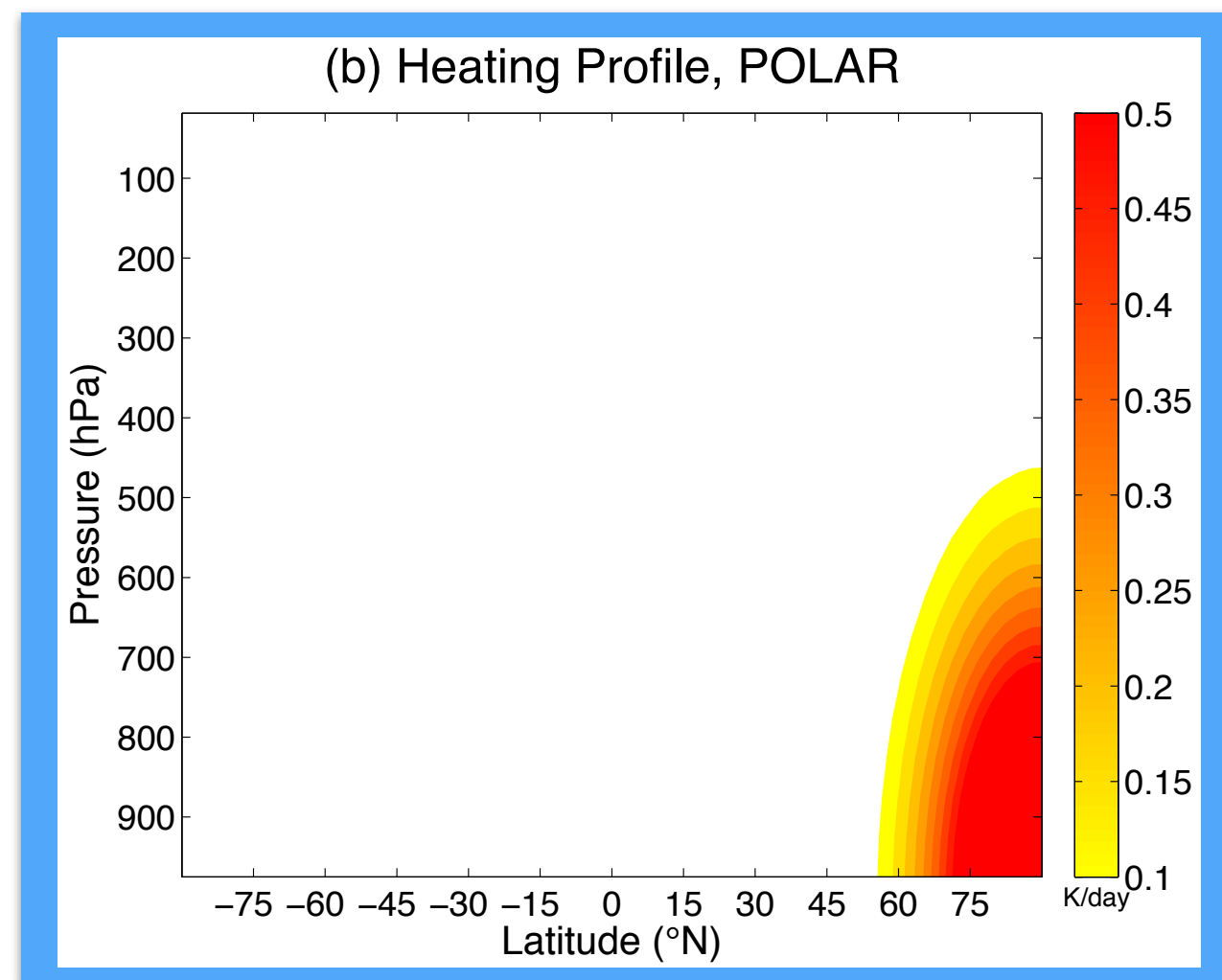
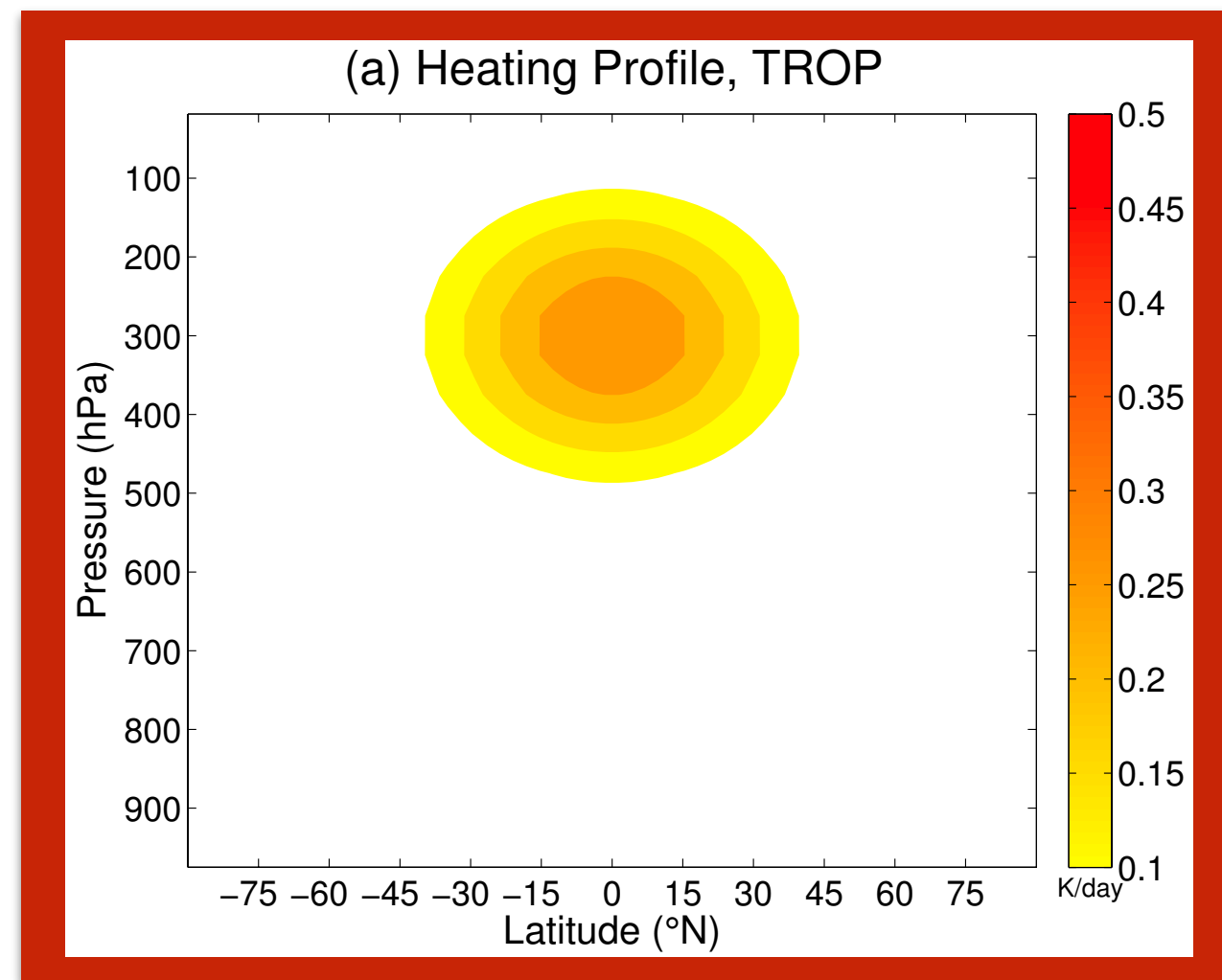
GFDL Dynamical Core simulations



- forced by relaxing temperatures to an equilibrium profile based on Held & Suarez (1994)
- run under perpetual climate conditions (each month is separate)
- no topography
- zonally symmetric
- no well-resolved stratosphere

McGraw & Barnes (2016); in press

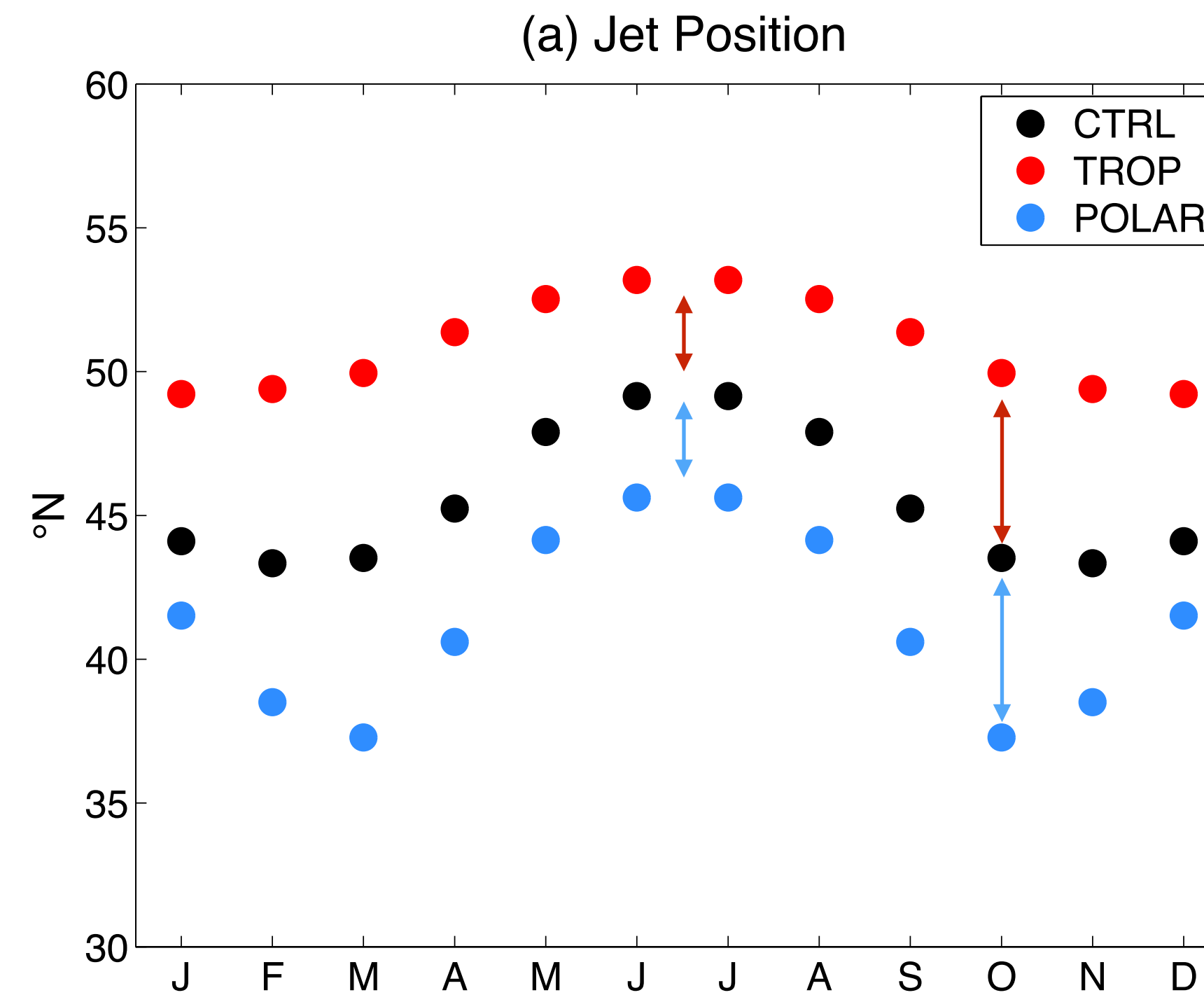
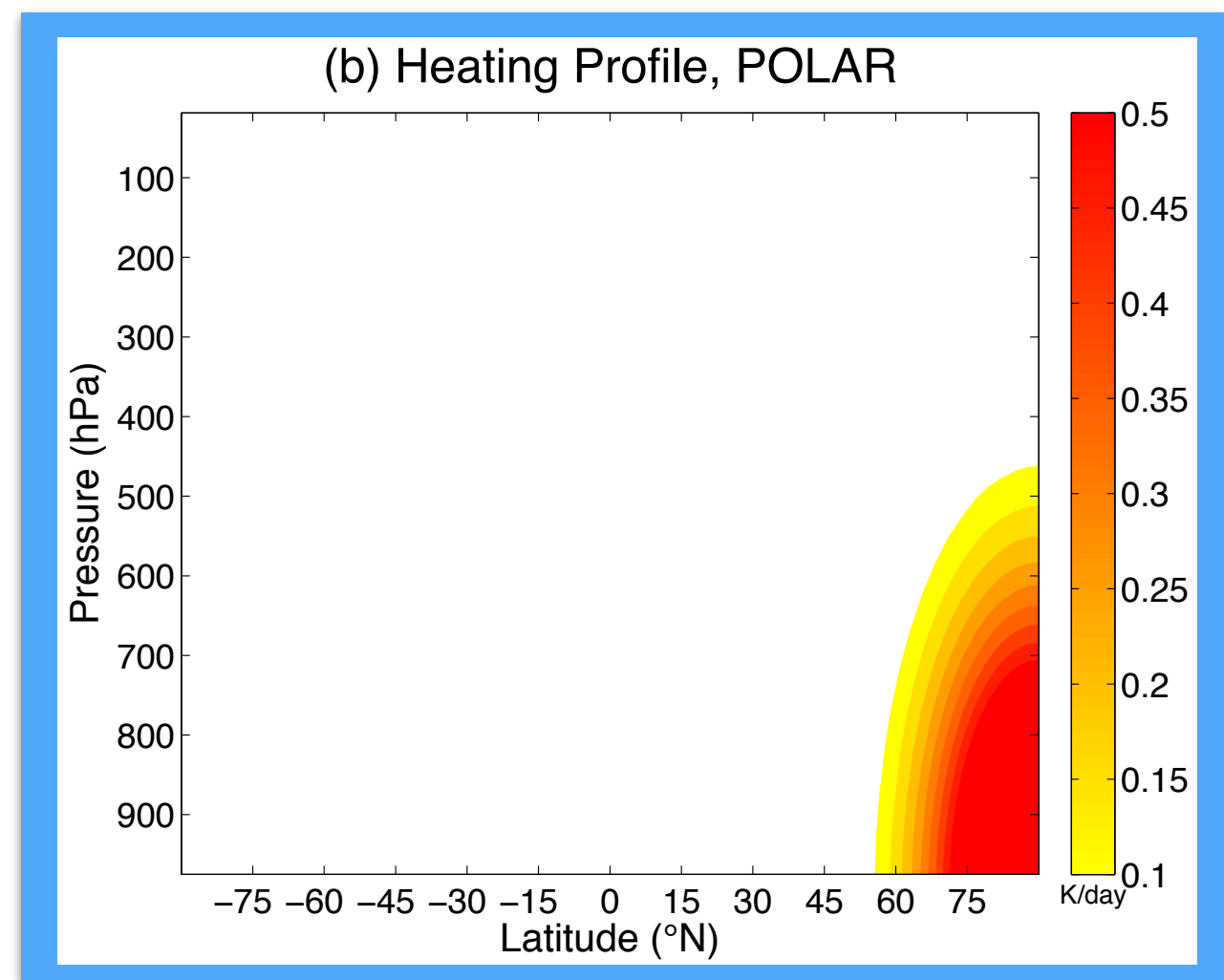
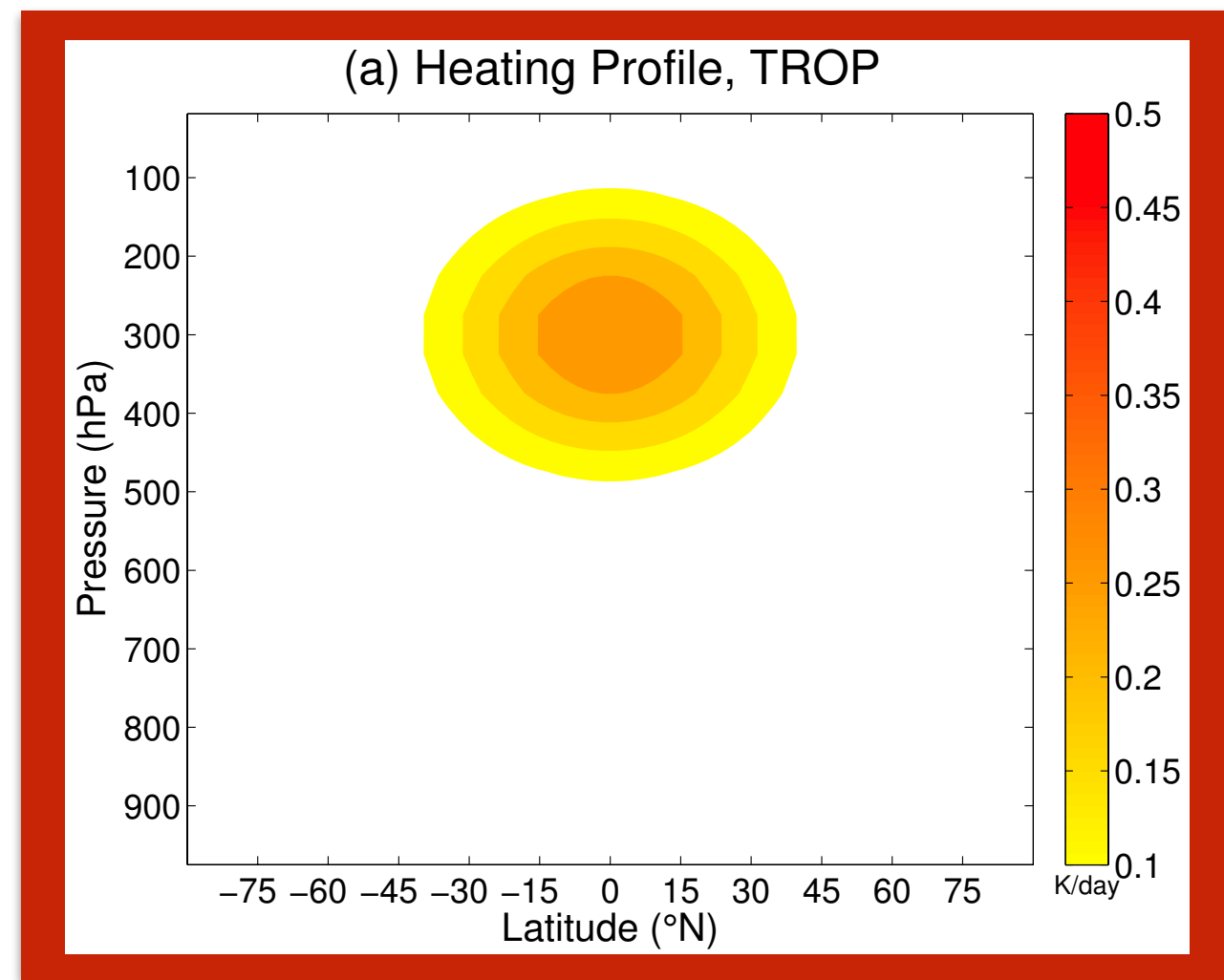
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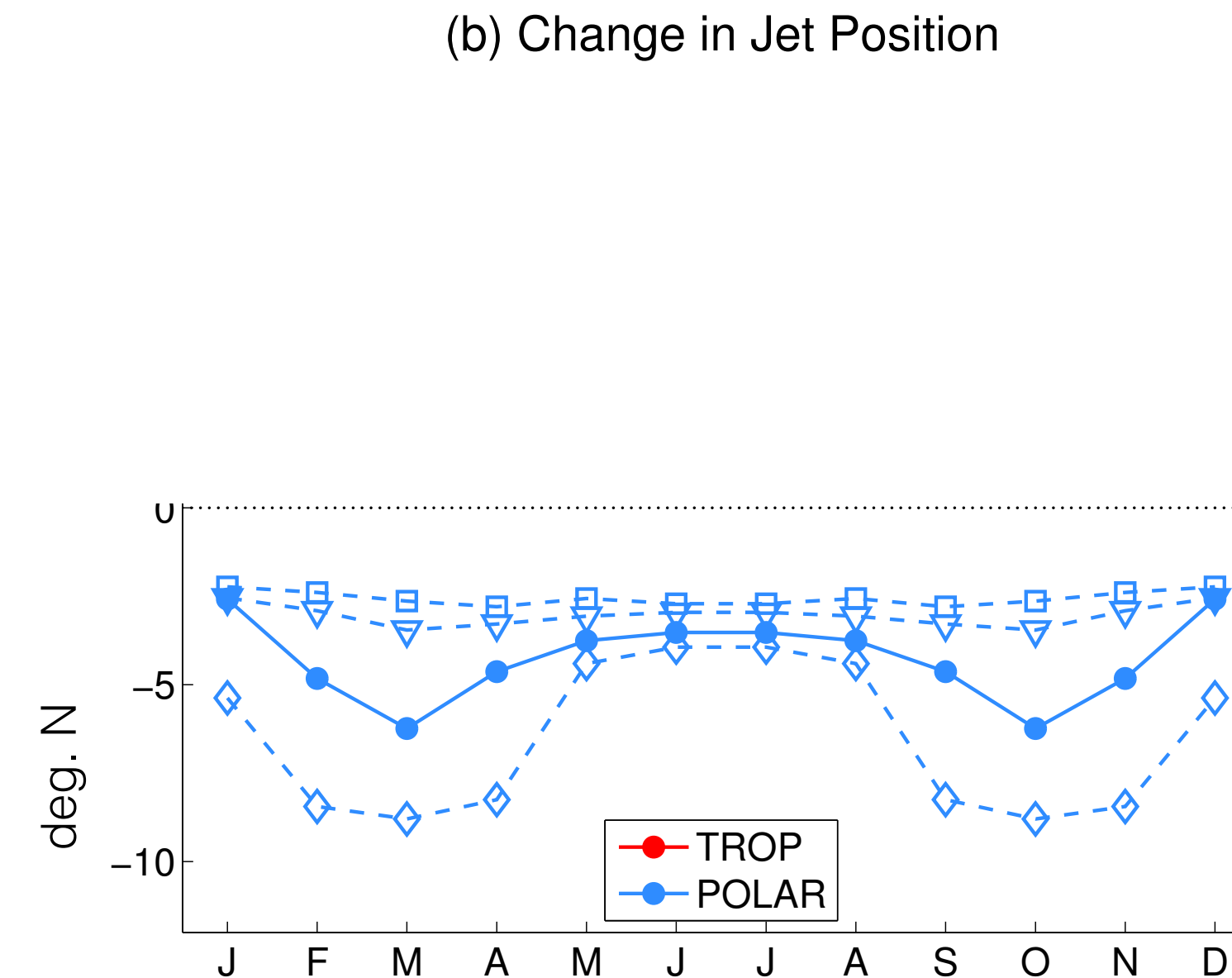
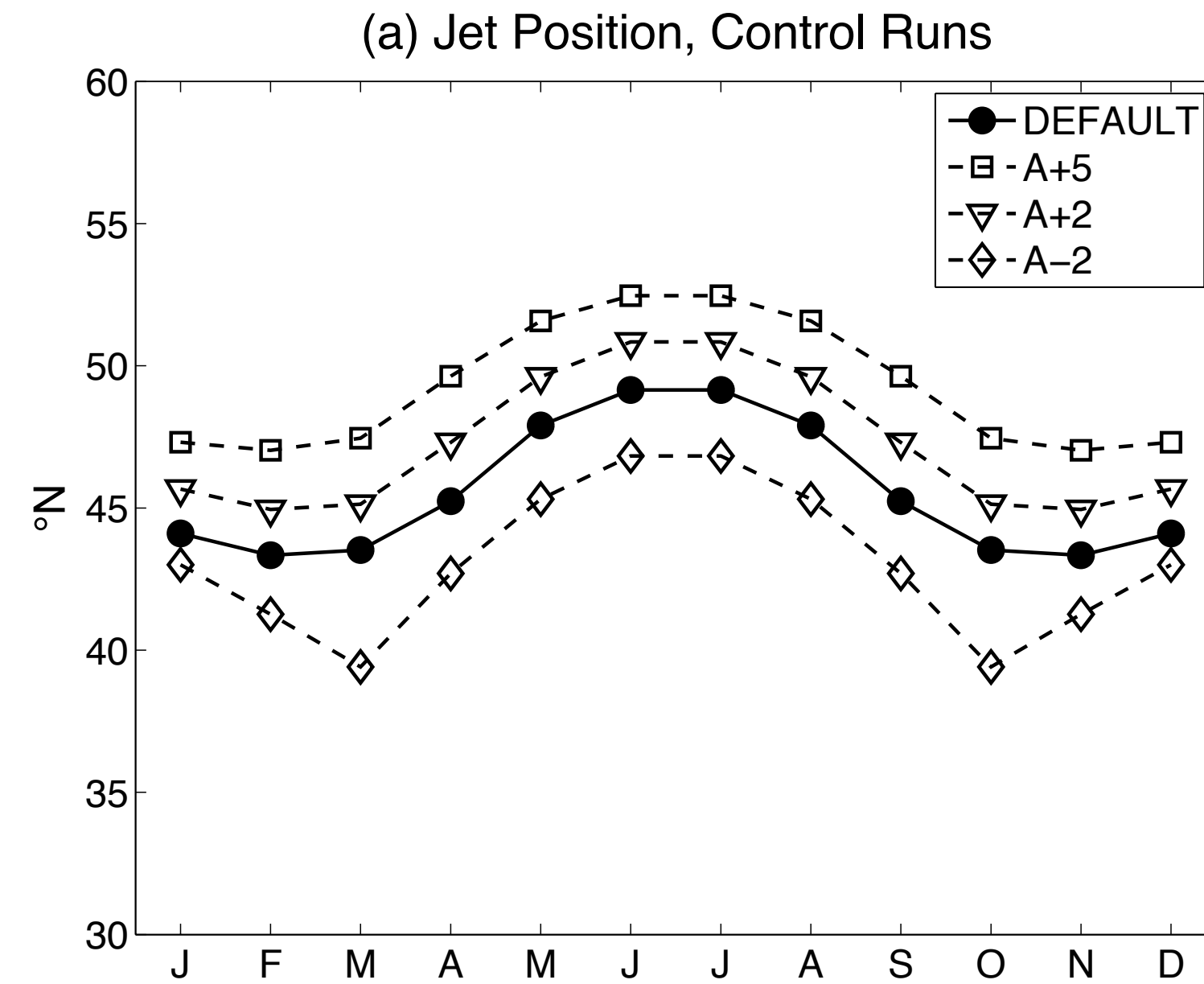
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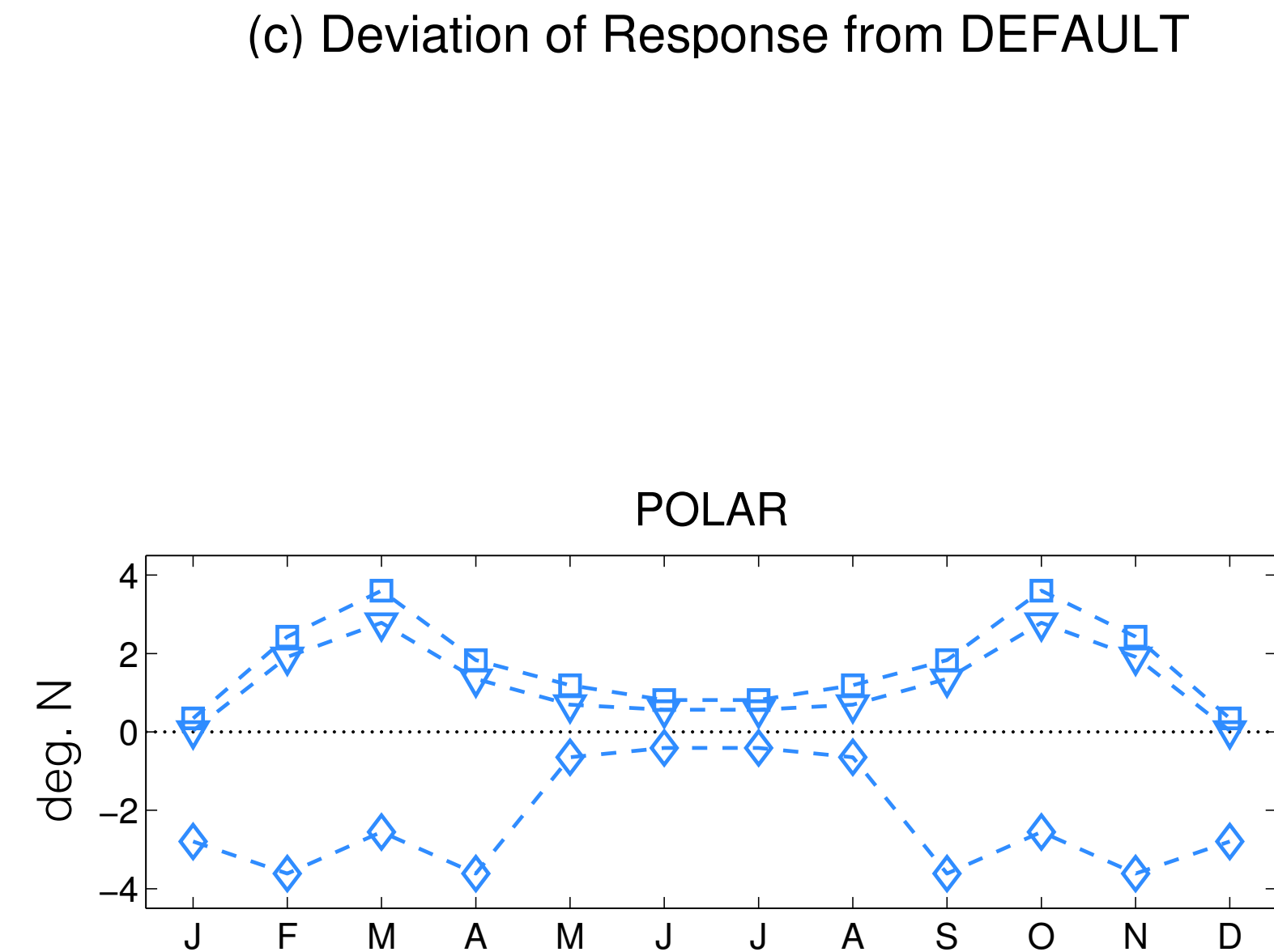
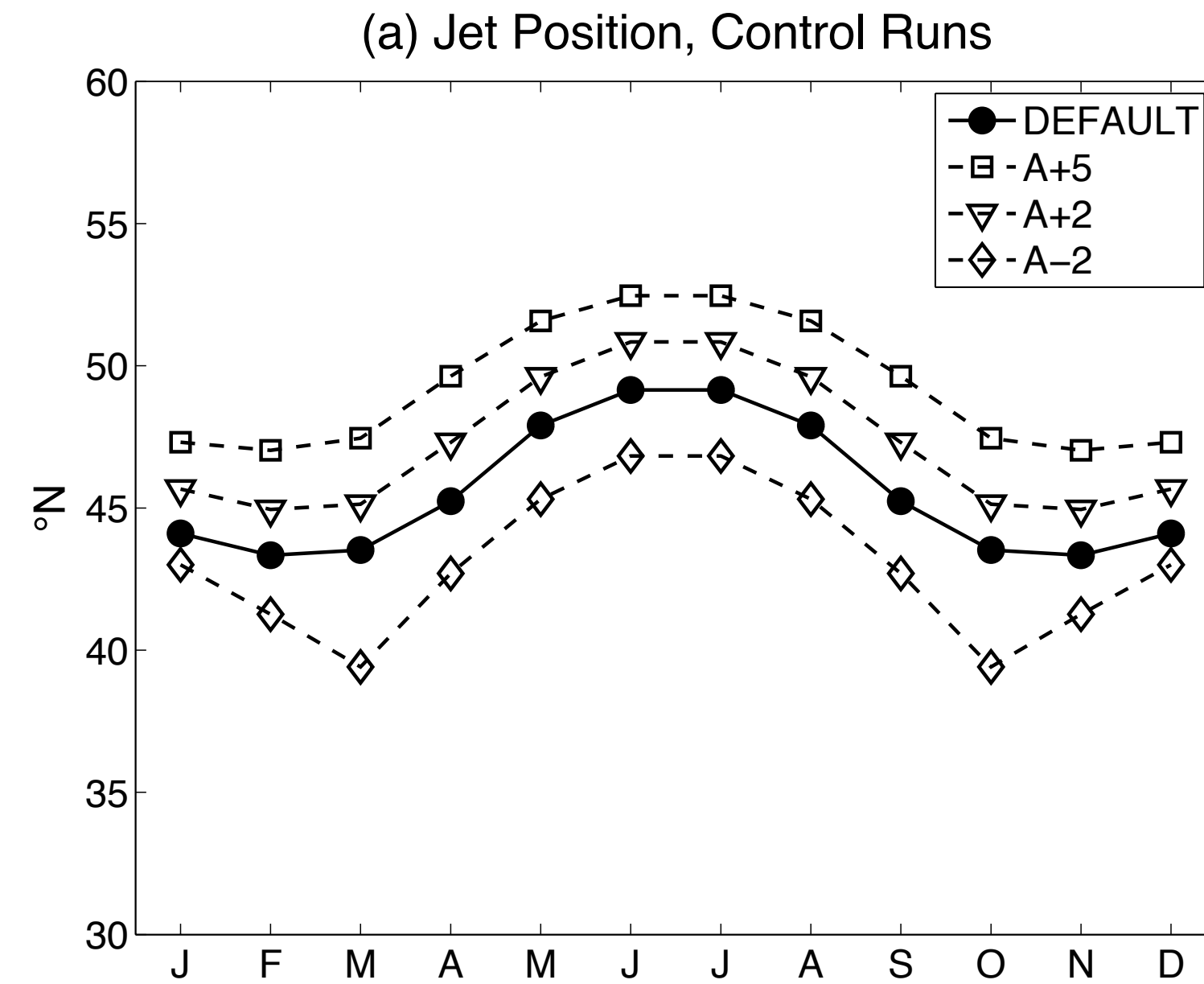
GFDL Dynamical Core simulations



dependence on the basic state changes throughout the year:

- matters most in the shoulder seasons
 - for POLAR warming, Dec/Jan & Jun/Jul are independent of basic state
 - invoke a fluctuation-dissipation argument for why

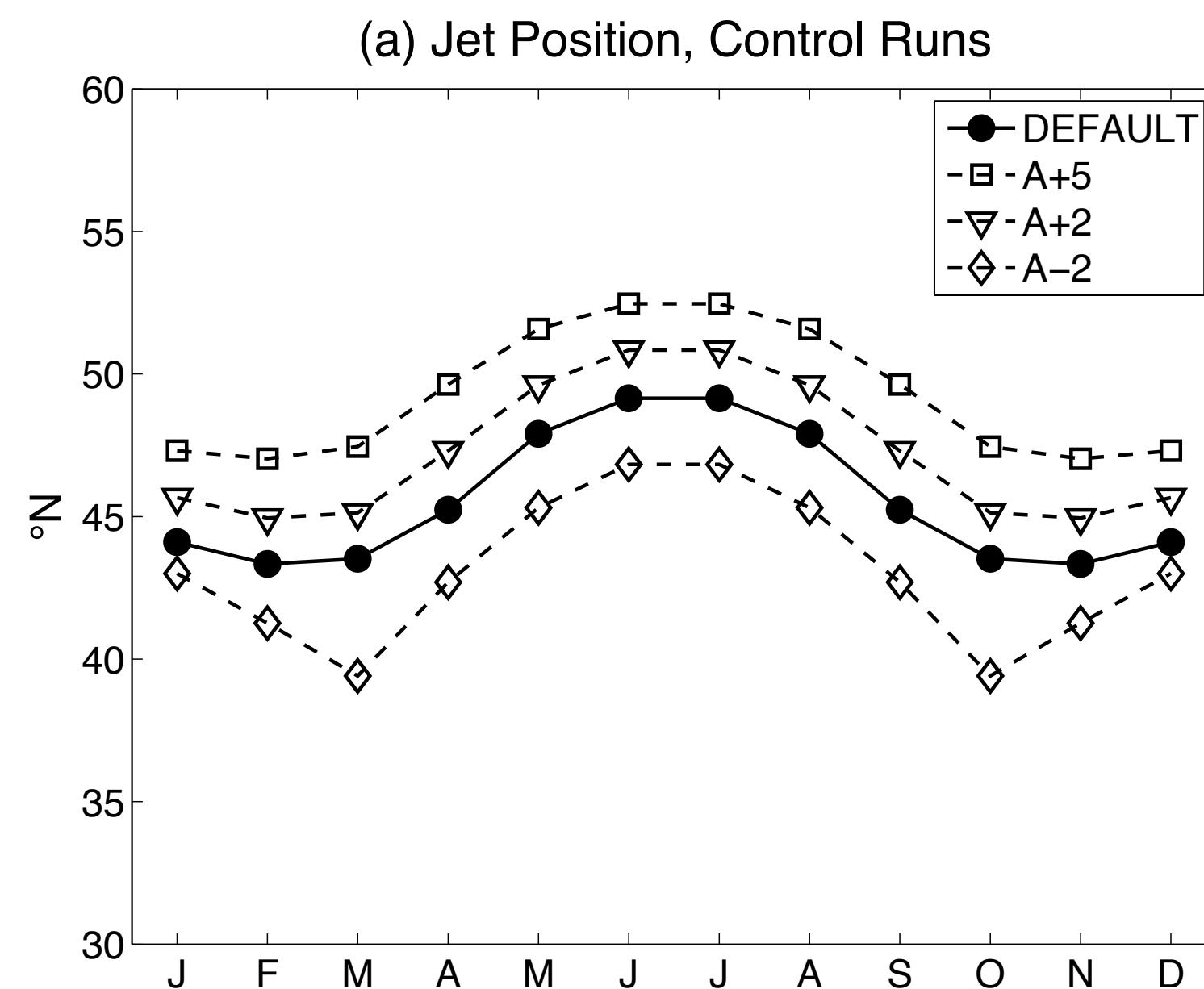
GFDL Dynamical Core simulations



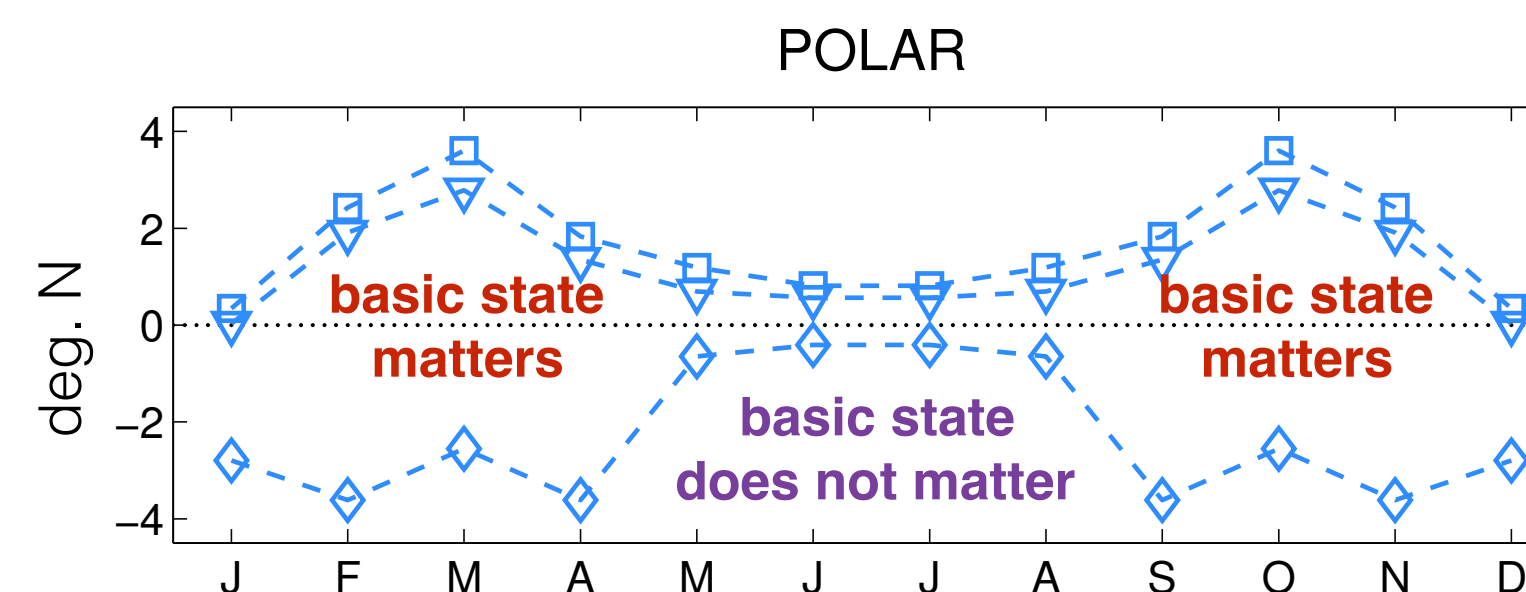
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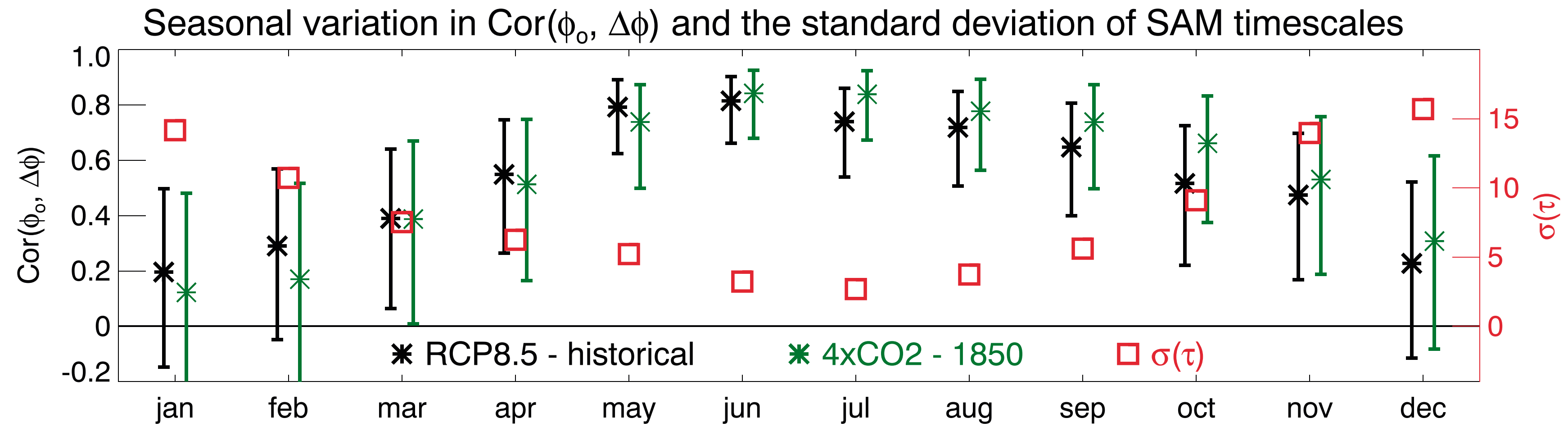


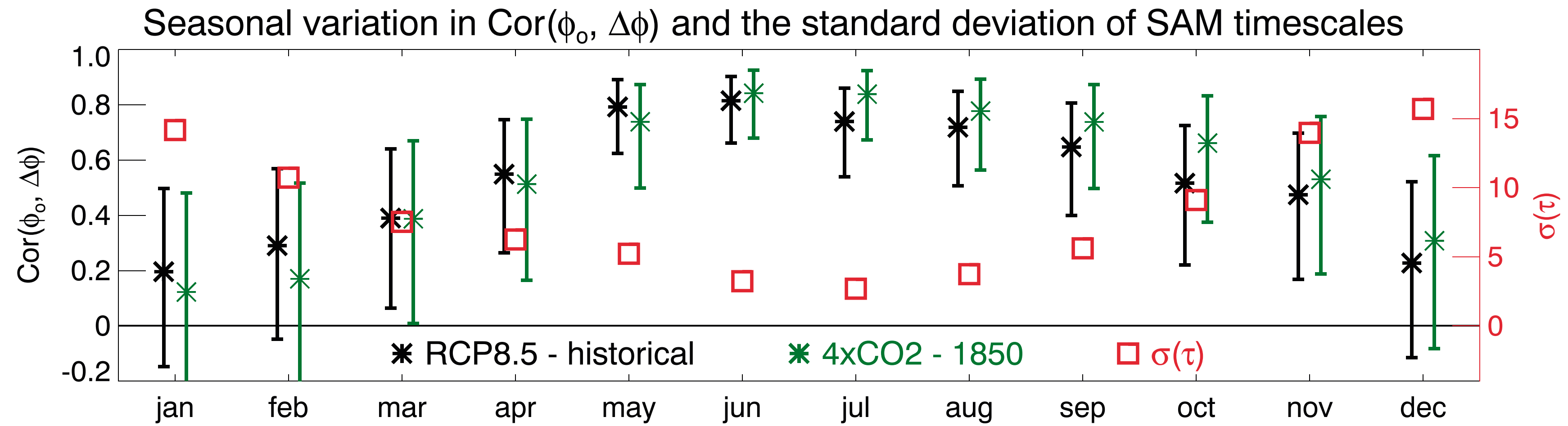
(c) Deviation of Response from DEFAULT



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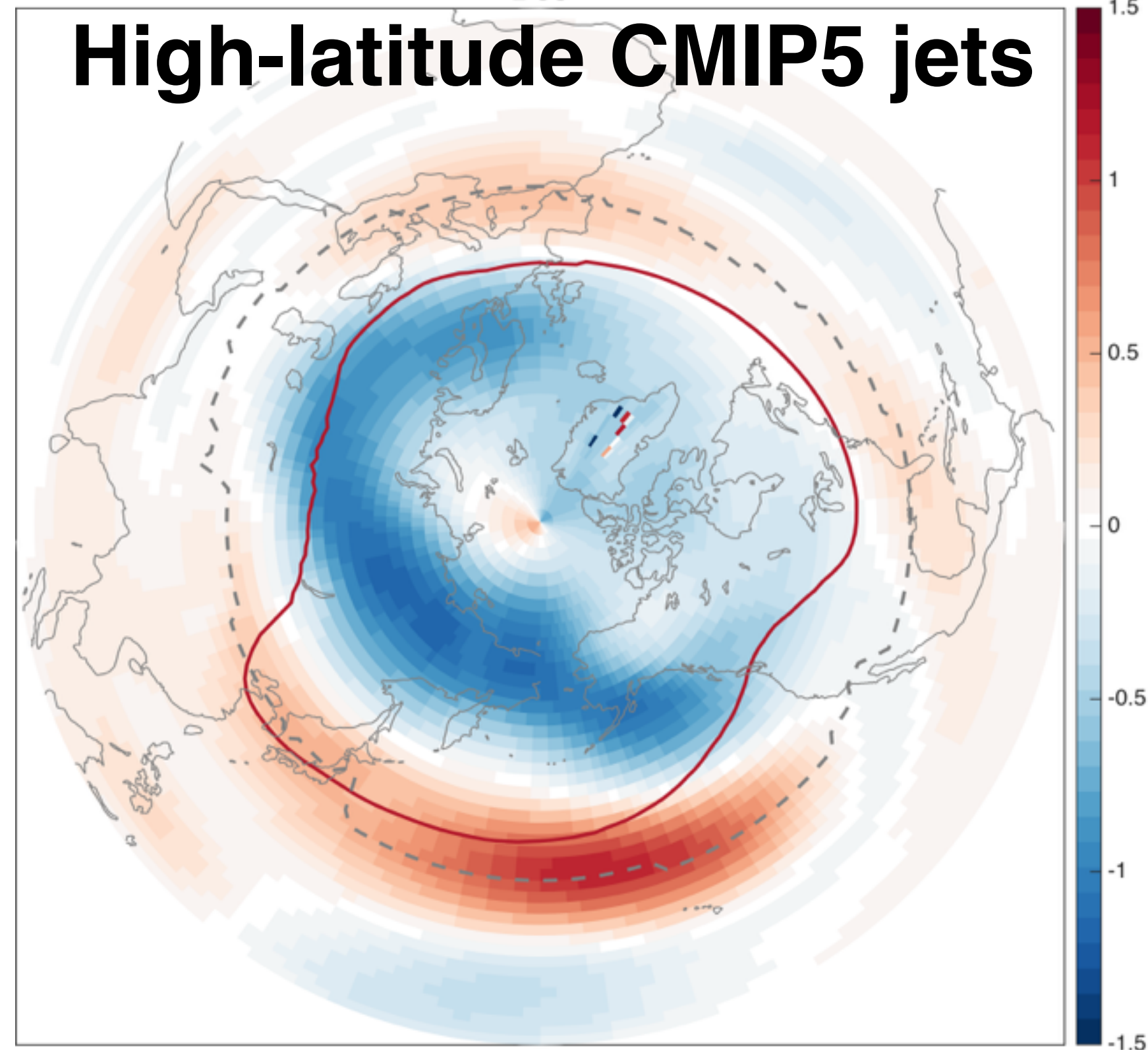




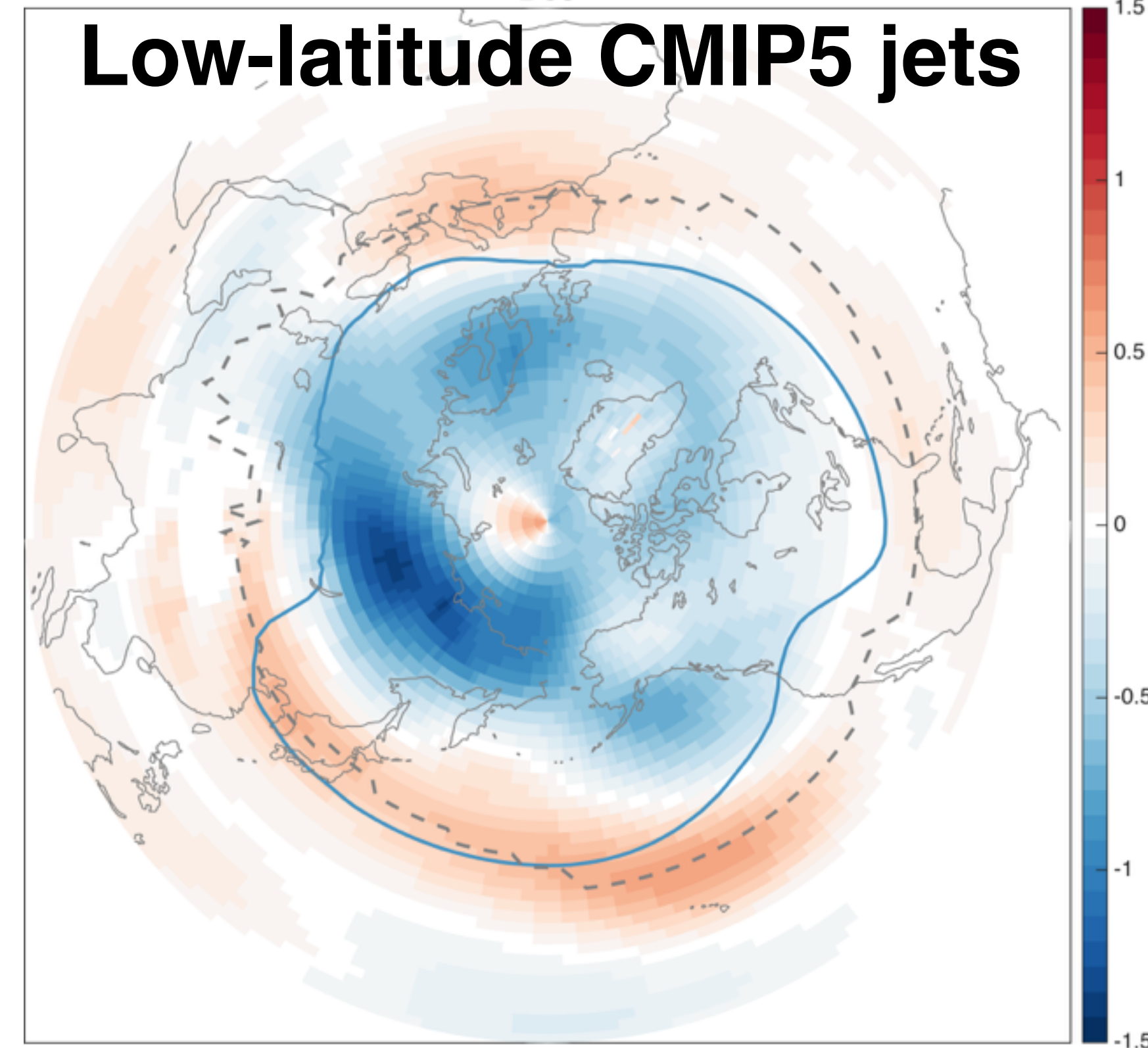
correlation between Southern Hemisphere jet position and jet shift largest in winter...
when timescale is the *shortest*

Regression maps: full wind field perspective

CMIP5 HIGH jetlat historical u700 response to 1K POLE warming
Dec



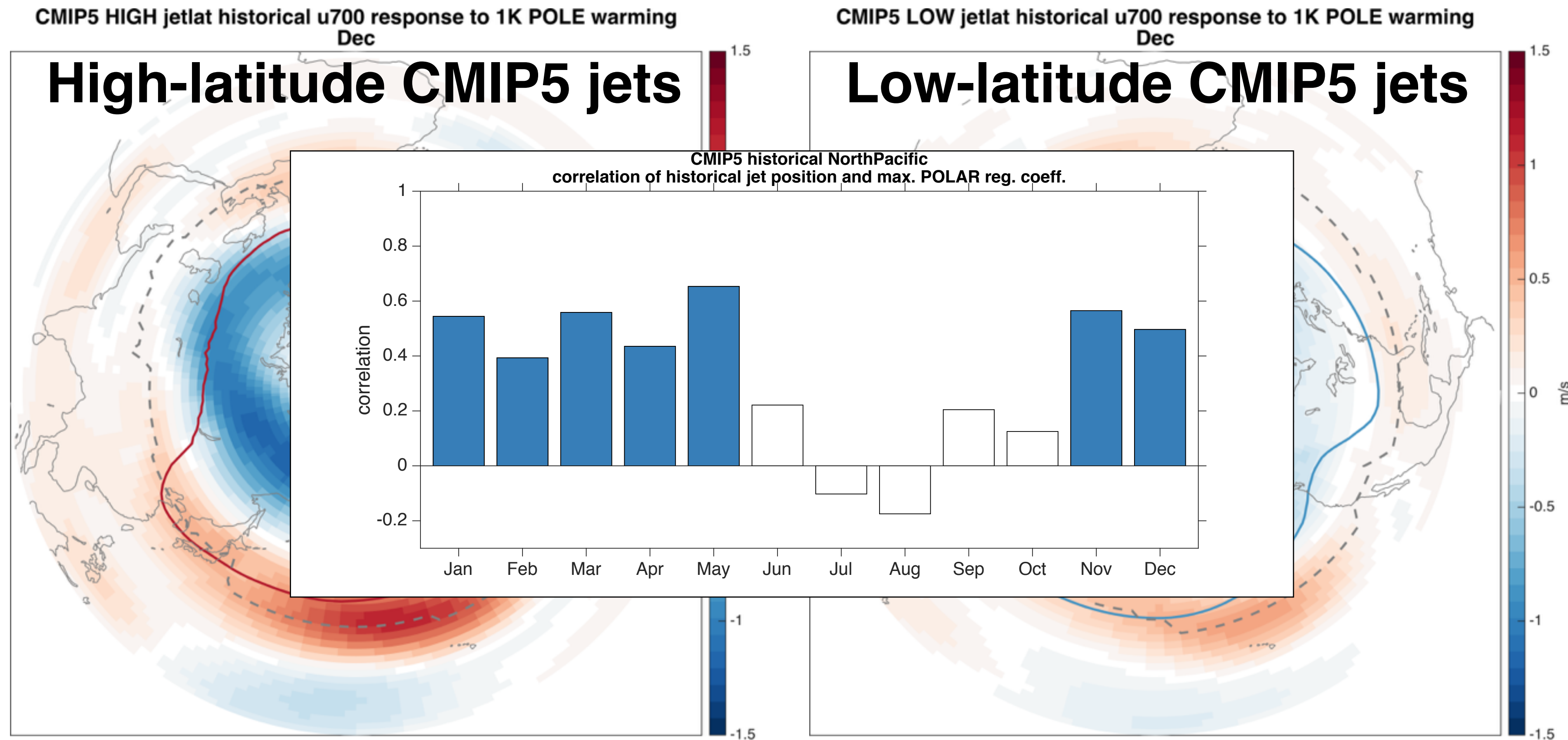
CMIP5 LOW jetlat historical u700 response to 1K POLE warming
Dec



25% most equatorward CMIP5 jet-streams
25% most poleward CMIP5 jet-streams
latitude of maximum regression coefficient

We're still missing part of the story though...
eddy feedbacks are likely at play here.

Regression maps: full wind field perspective



25% most equatorward CMIP5 jet-streams
 25% most poleward CMIP5 jet-streams

latitude of maximum regression coefficient

We're still missing part of the story though...
eddy feedbacks are likely at play here.

CMIP5: Improving climate change detection

Improving Climate Change Detection through Optimal Seasonal Averaging: The Case of the North Atlantic Jet and European Precipitation

GIUSEPPE ZAPPA, BRIAN J. HOSKINS, AND THEODORE G. SHEPHERD

Department of Meteorology, University of Reading, Reading, United Kingdom

(Manuscript received 4 December 2014, in final form 15 April 2015)

ABSTRACT

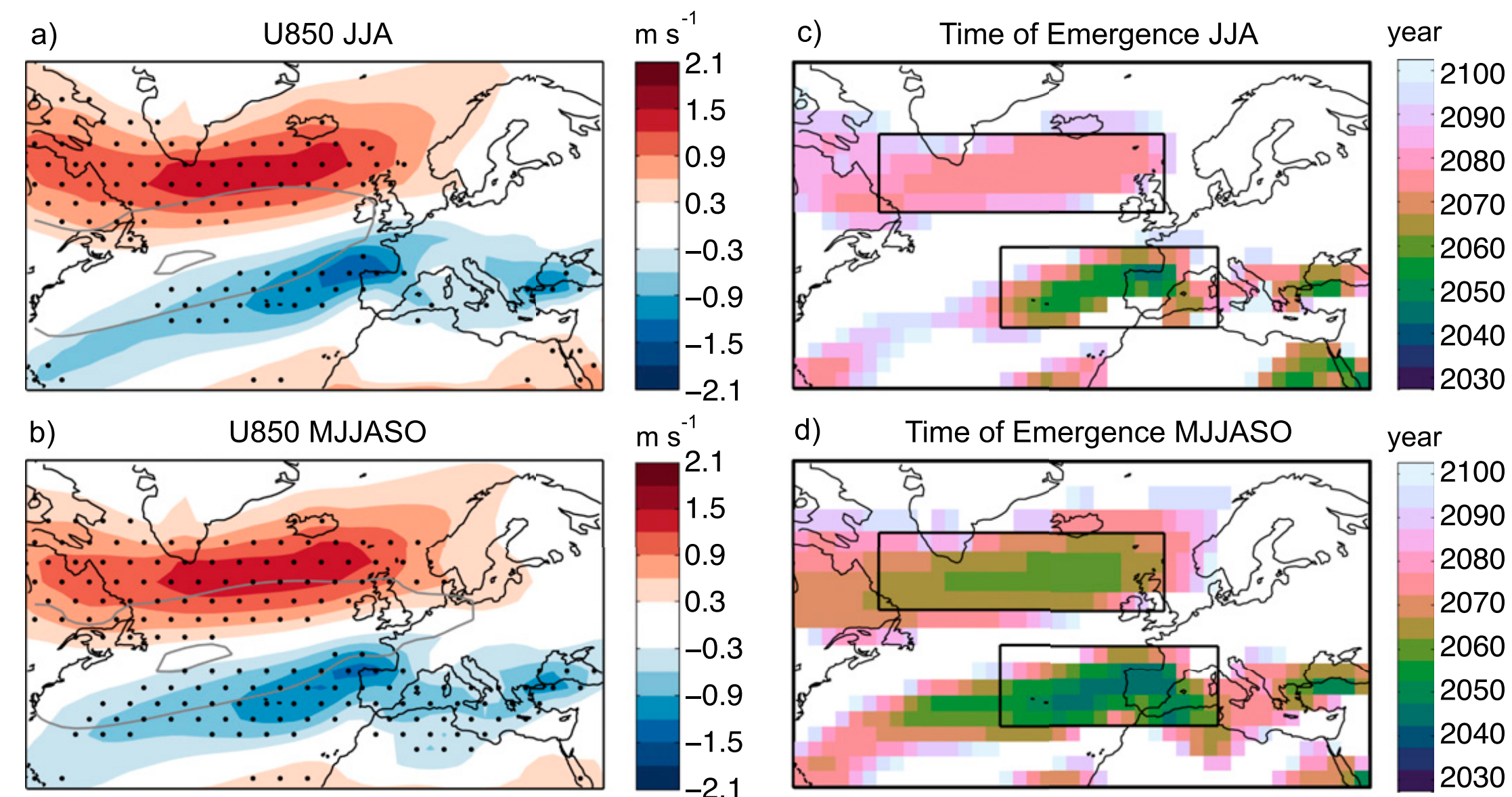
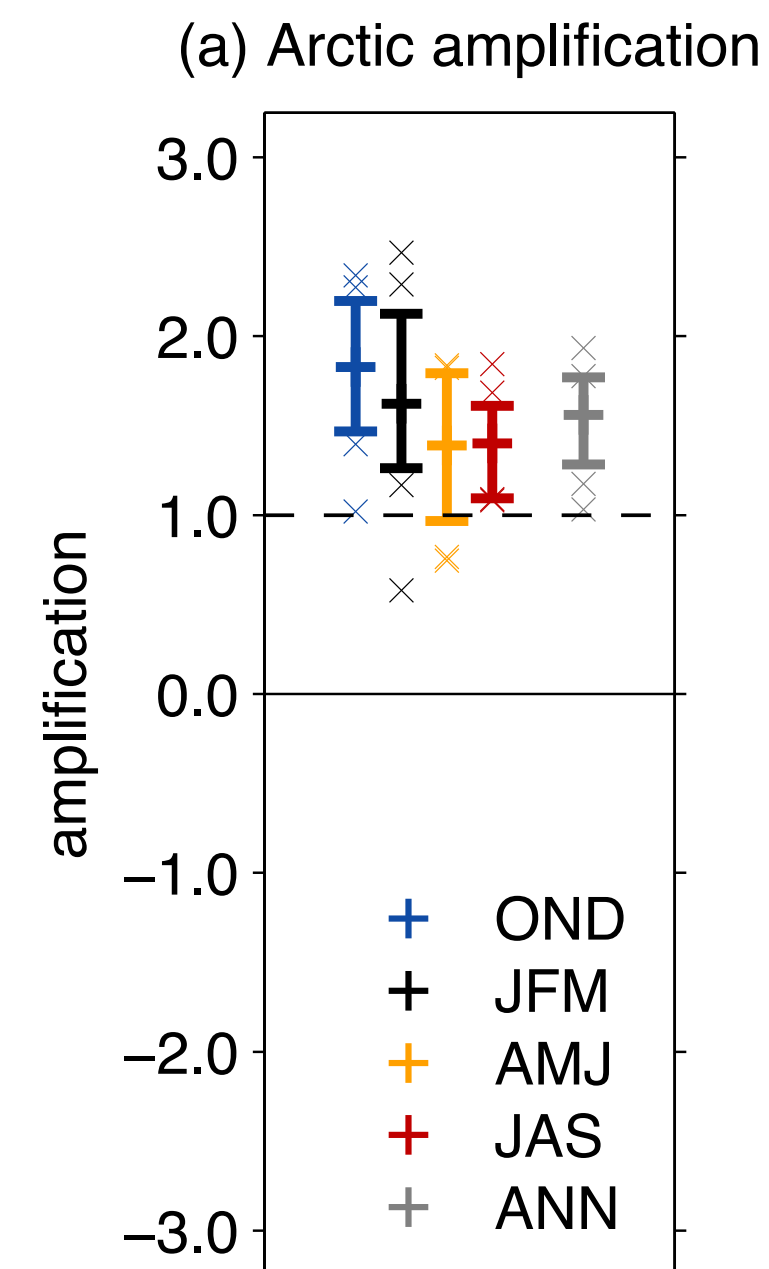
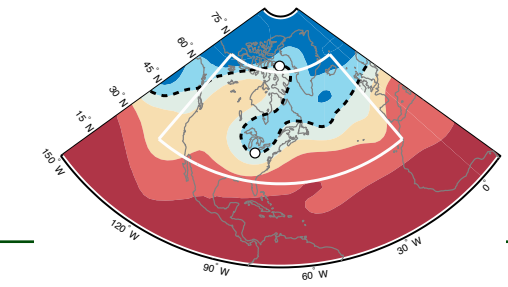


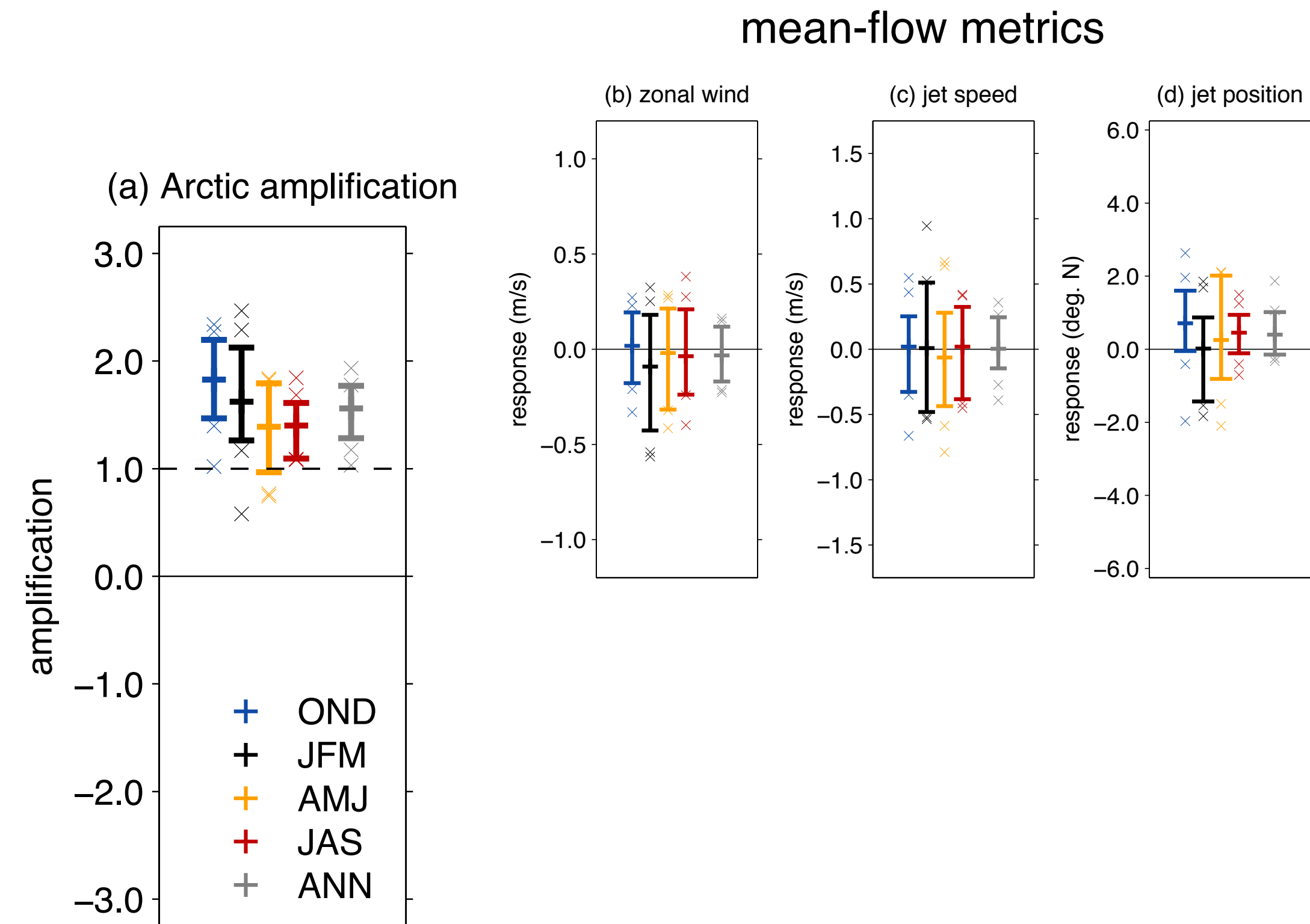
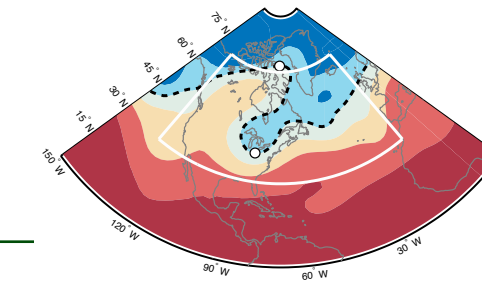
FIG. 2. Multimodel mean end-of-century U850 response separately computed for the (a) meteorological summer (JJA) and (b) extended summer (MJJASO) time averages. (c),(d) The time of emergence of the U850 response evaluated for the time periods in (a) and (b), respectively. In (a) and (b), stippling is applied where at least 90% of the models show a response of the same sign for the end-of-century climate change response, and the gray contours correspond to the 4 (outer) and 8 (inner) m s^{-1} isotachs of U850 in the historical period in the multimodel mean.

Near-term projections over N. Atlantic/N. America (2020-2044) - (1980-2004)



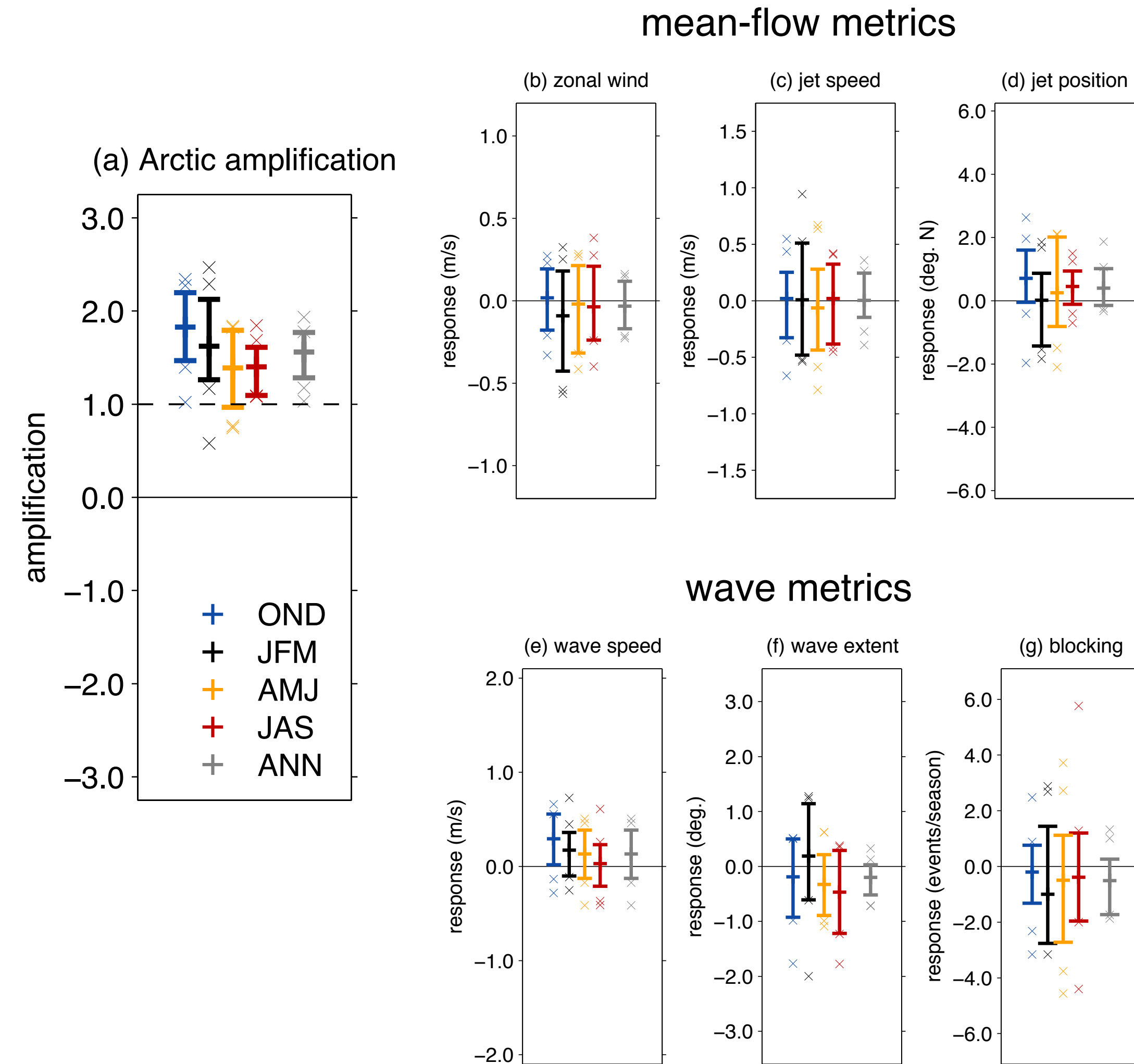
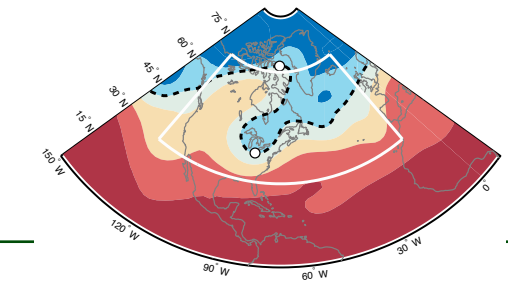
27 CMIP5 GCMs
Barnes & Polvani (2015; JCLI)

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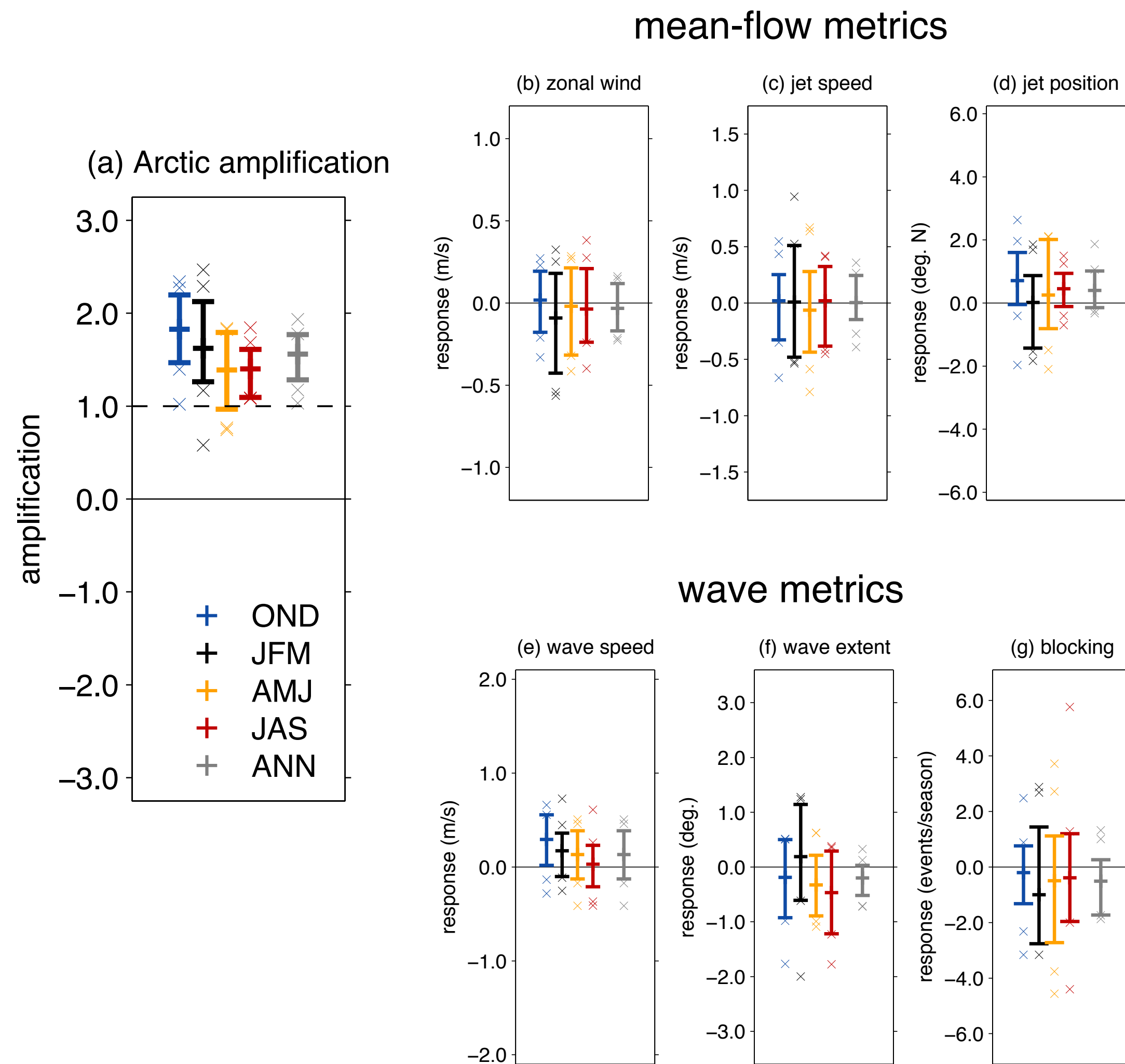
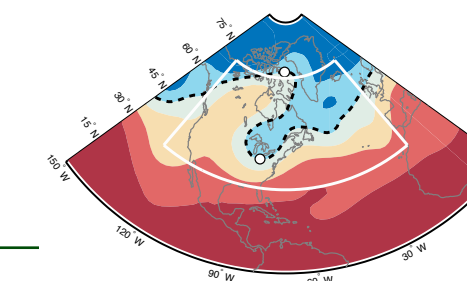


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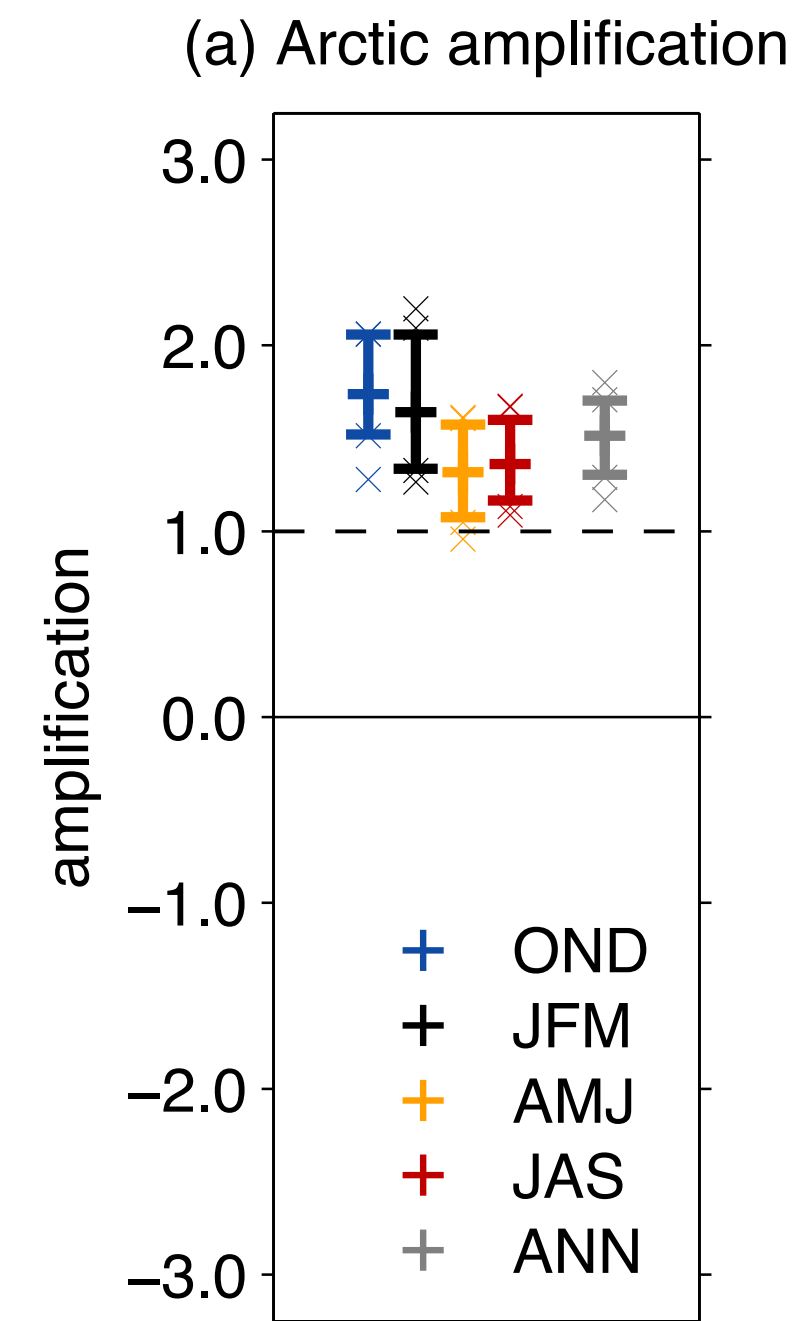
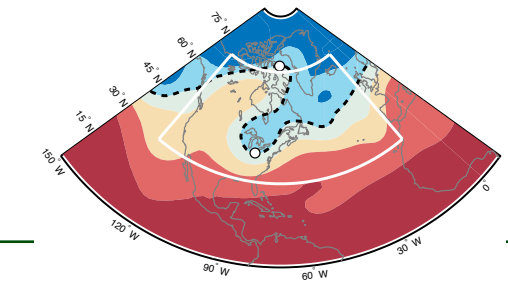
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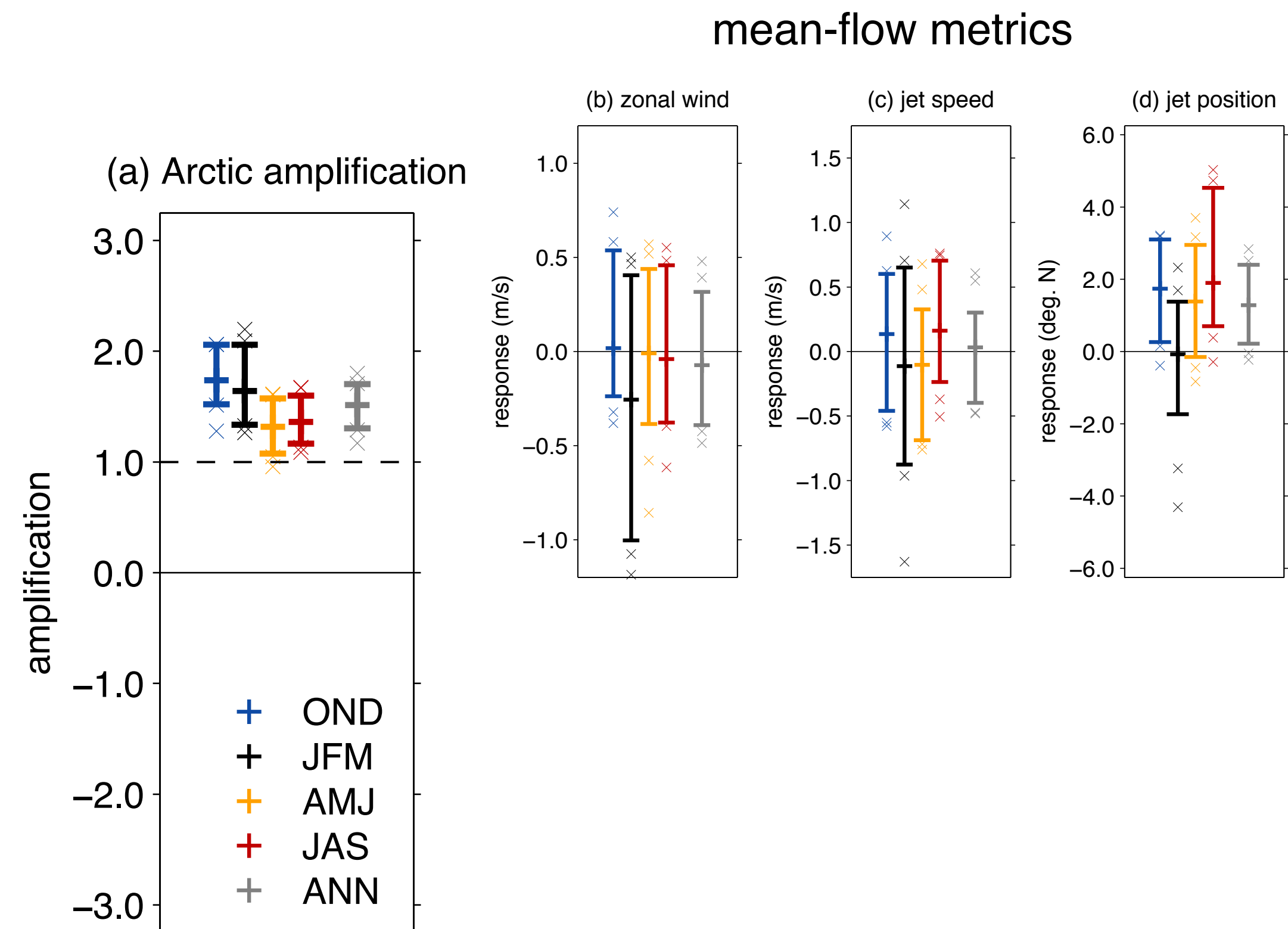
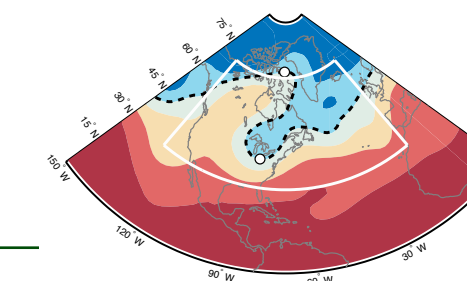
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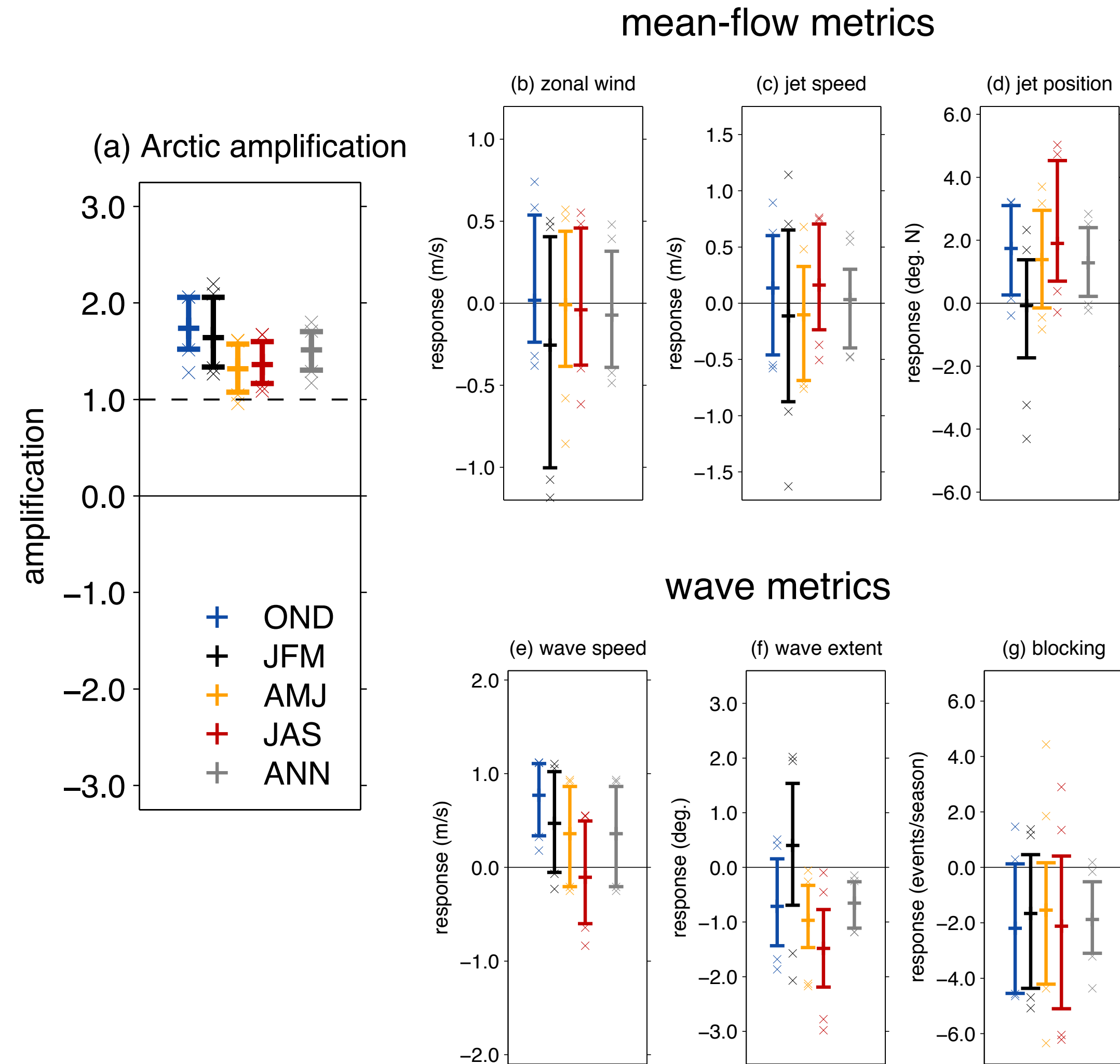
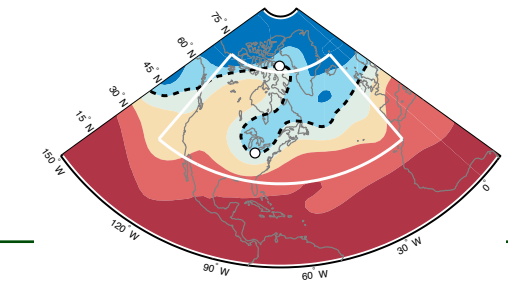


No consensus in the circulation response

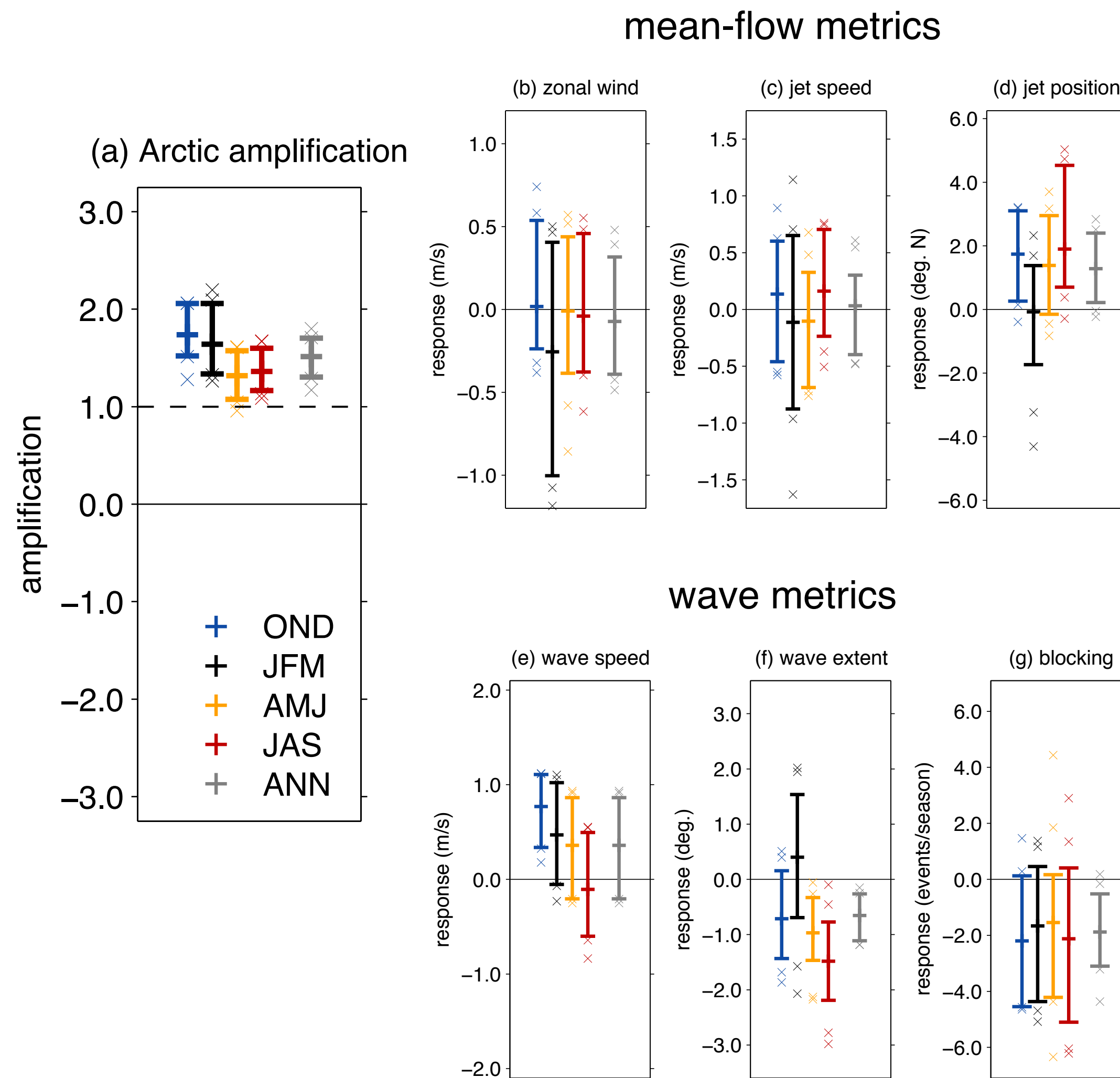
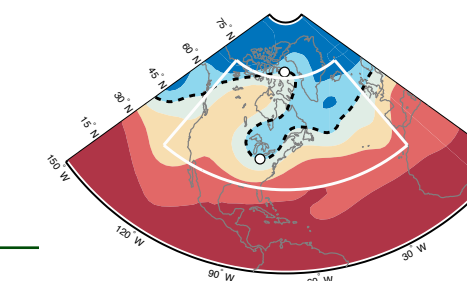


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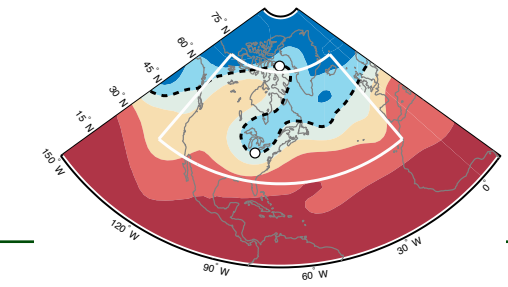




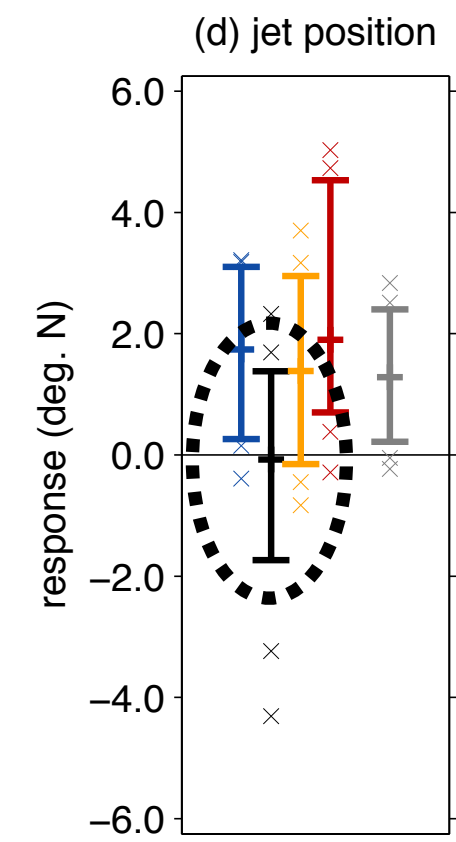
27 CMIP5 GCMs
Barnes & Polvani (2015; JCLI)



No consensus in the circulation response or response is of the opposite sign to that hypothesized

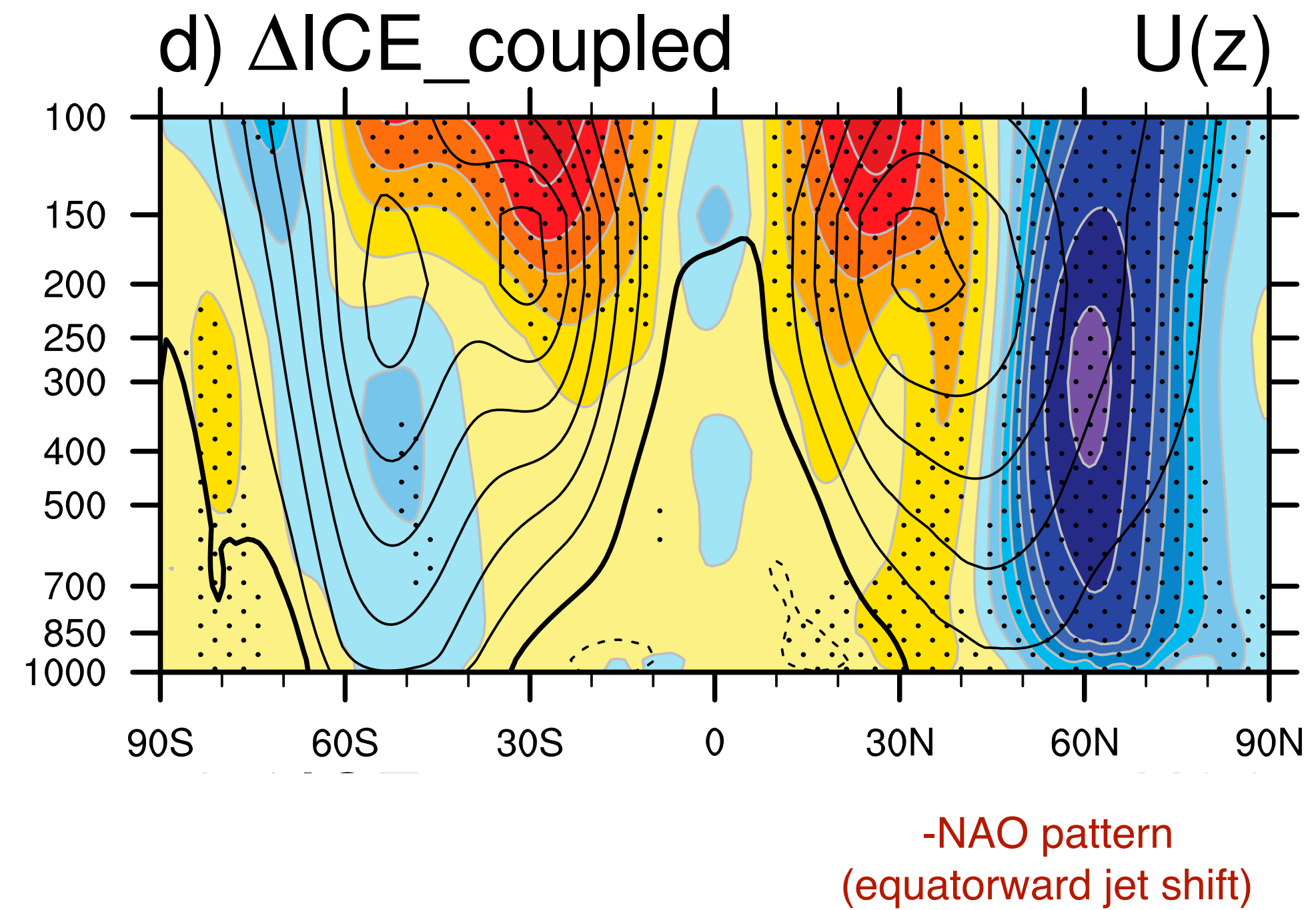
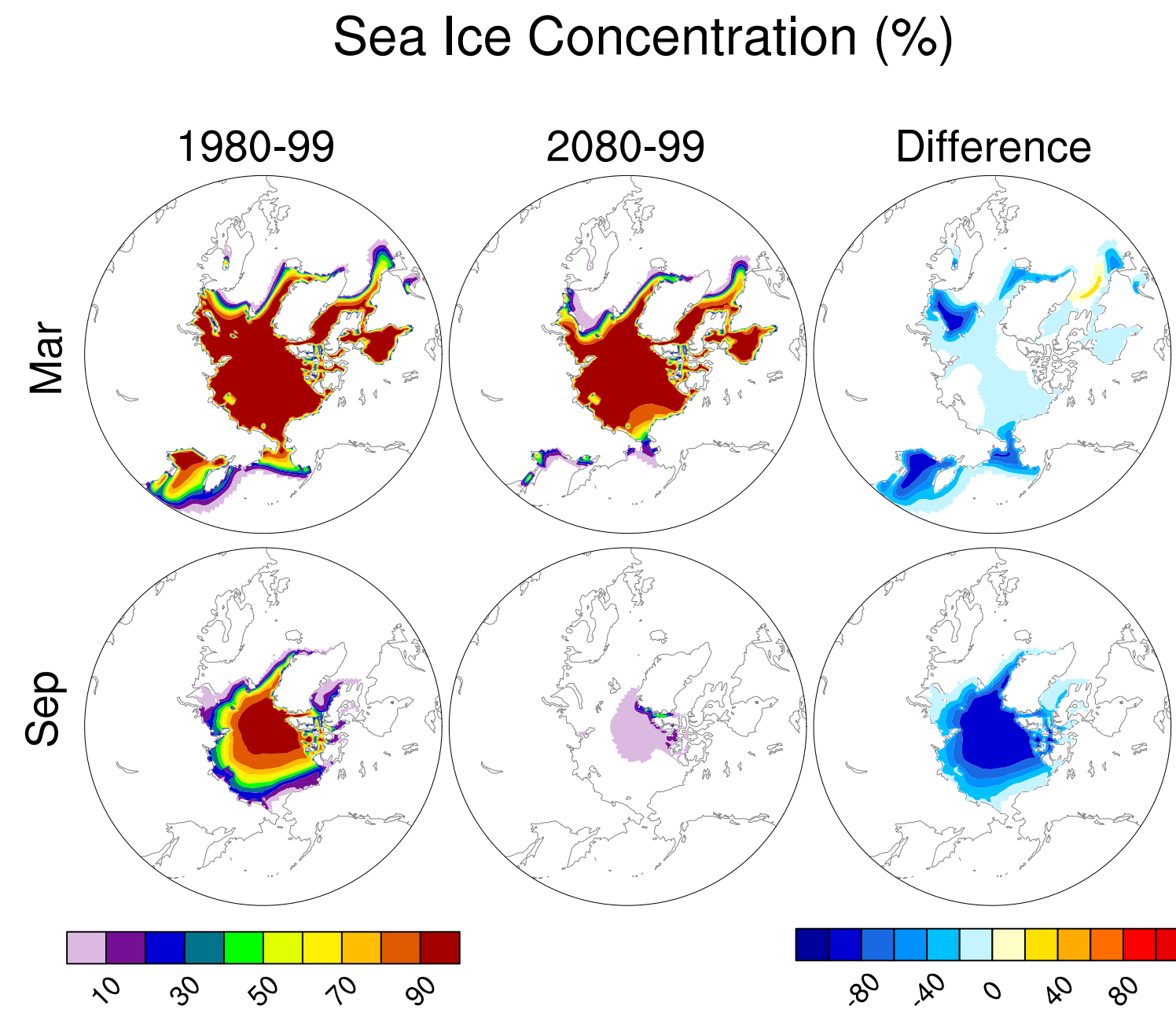


mean-flow metrics



What is different about winter?

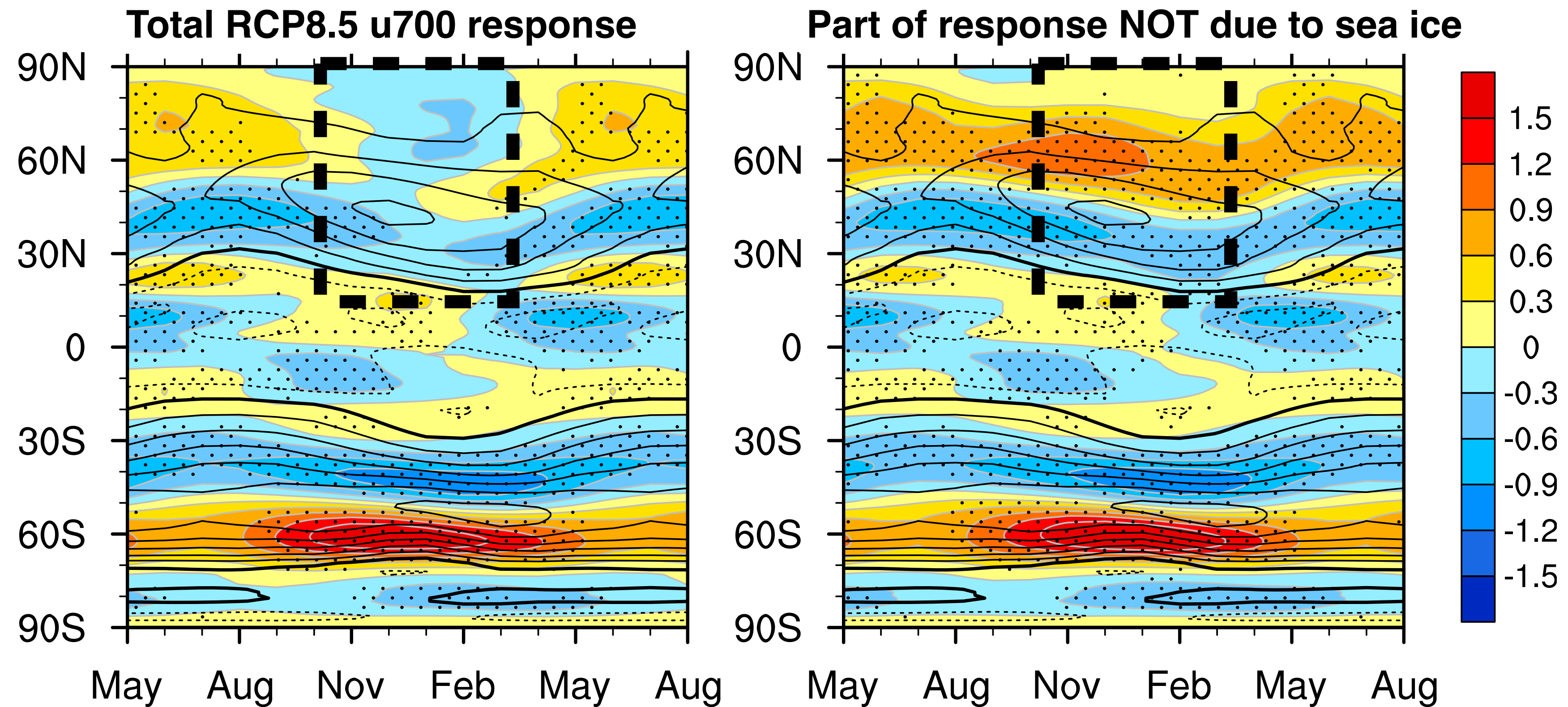
Can it? Modeling evidence



recent coupled GCM experiments also demonstrate a midlatitude response

*coupled CCSM4 simulations
with additional long wave radiative fluxes in the ice model
Deser, Tomas, et al. (2015; JCLI)*

Tug-of-war



- In CCSM4, the sea ice loss effects appears to cancel the poleward shift of the jet
- In other CMIP5 models, the poleward shift “wins”, and others the equatorward shift “wins”

*coupled CCSM4 simulations
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Response to GHG independent of season in SH?

Delayed Southern Hemisphere Climate Change Induced by Stratospheric Ozone Recovery, as Projected by the CMIP5 Models

ELIZABETH A. BARNES

Lamont-Doherty Earth Observatory, Columbia University, Palisades, New York

NICHOLAS W. BARNES

Department of Computer Science and Engineering, University of Minnesota, Twin Cities, Minneapolis, Minnesota

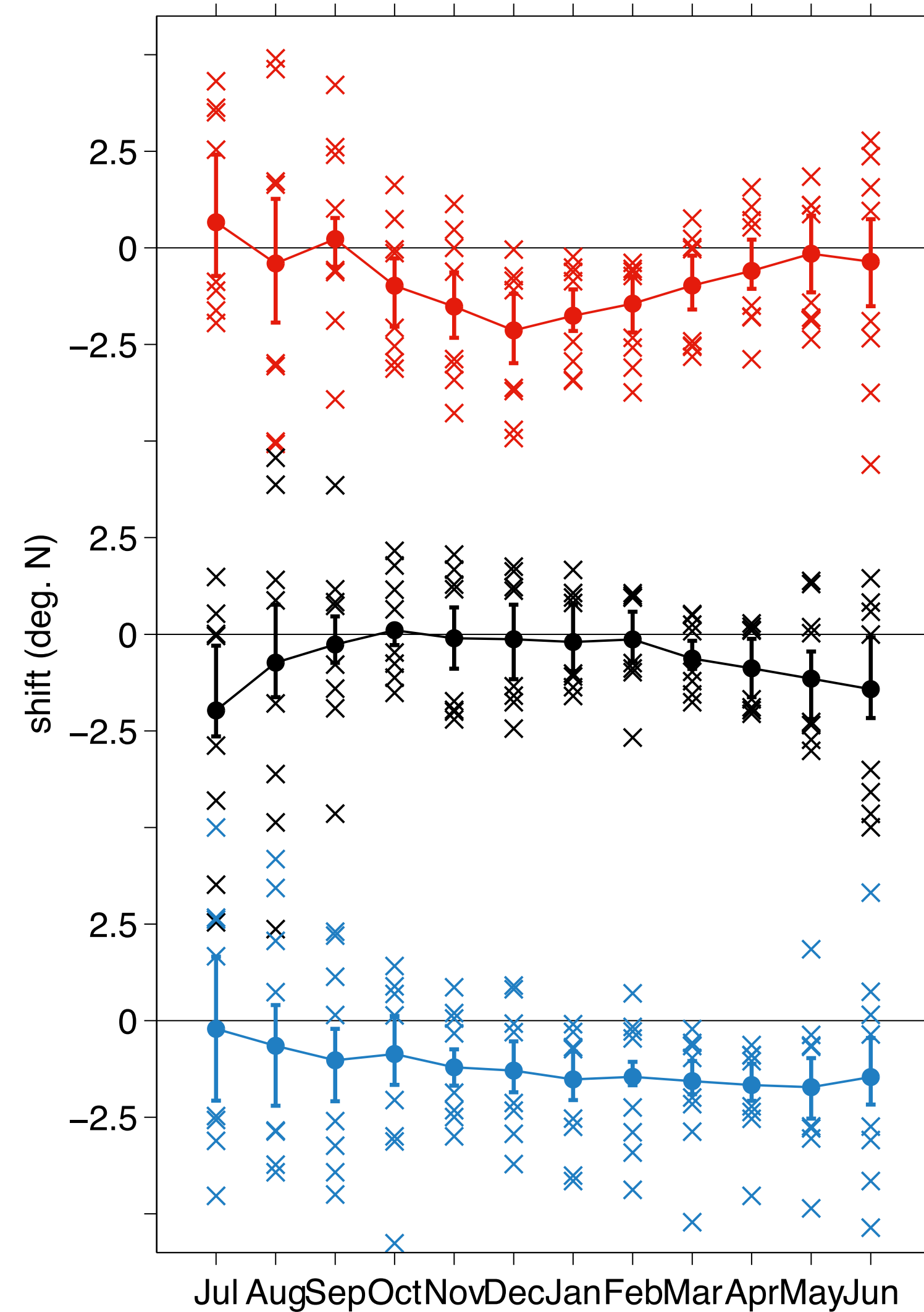
LORENZO M. POLVANI

Lamont-Doherty Earth Observatory, Columbia University, Palisades, and Department of Applied Physics and Applied Math, Columbia University, New York, New York

(Manuscript received 20 April 2013, in final form 5 August 2013)

O3DEPL: (2000-2010) - (1960-1970)
O3RCVR: (2040-2050) - (2000-2010)
FUTR: (2089-2099) - (2040-2050)

(a) Jet shift by month



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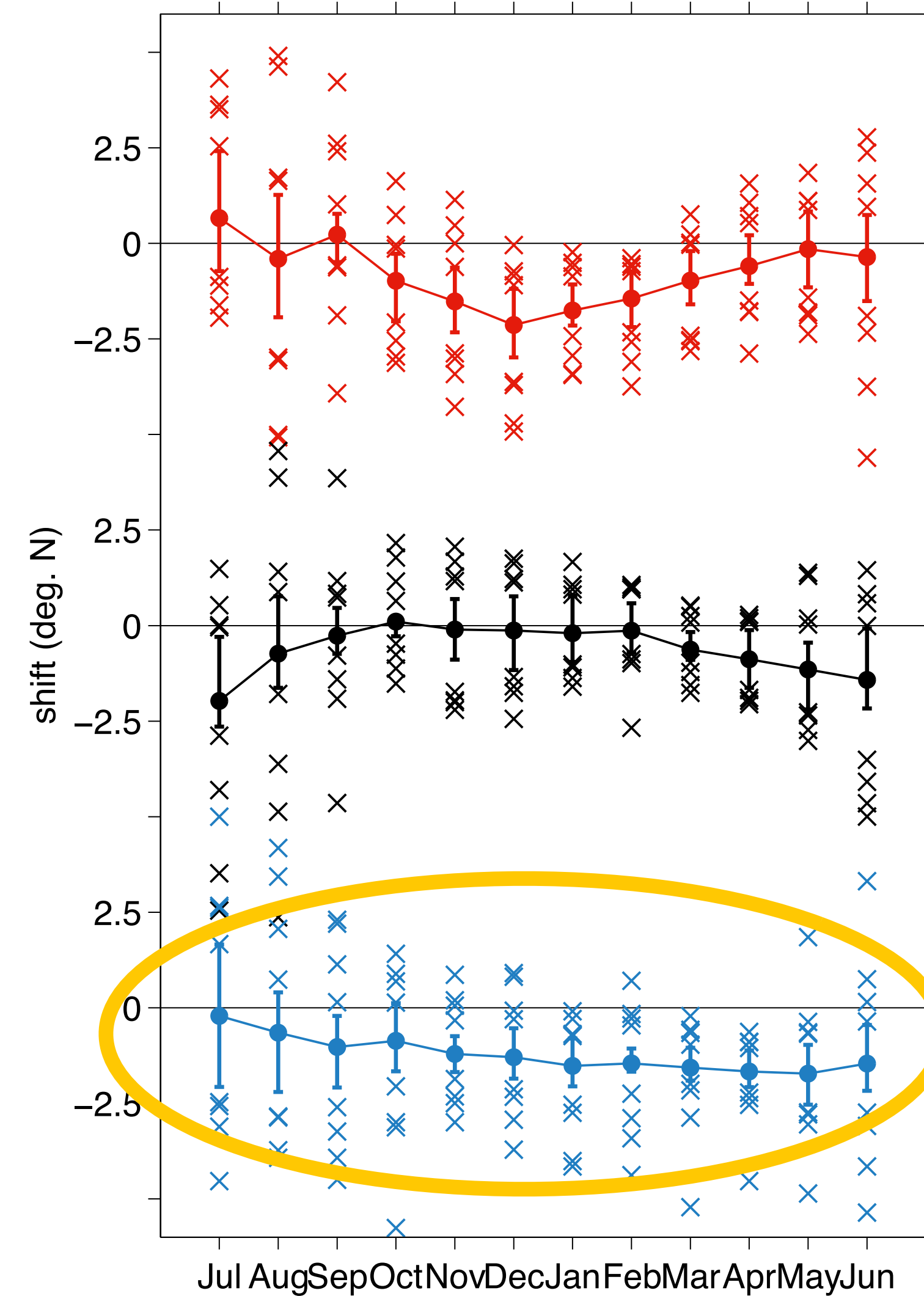
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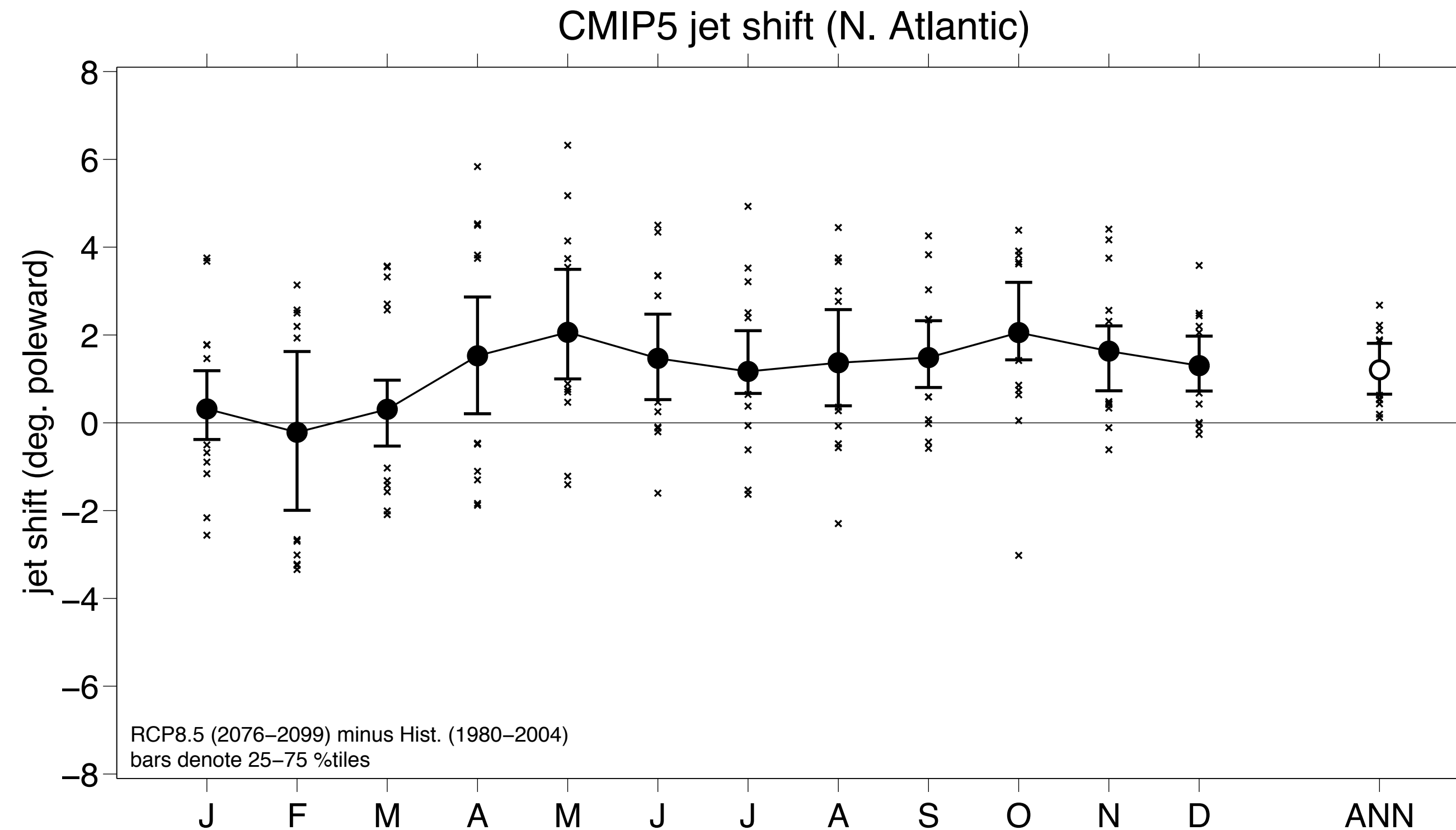
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(a) Jet shift by month

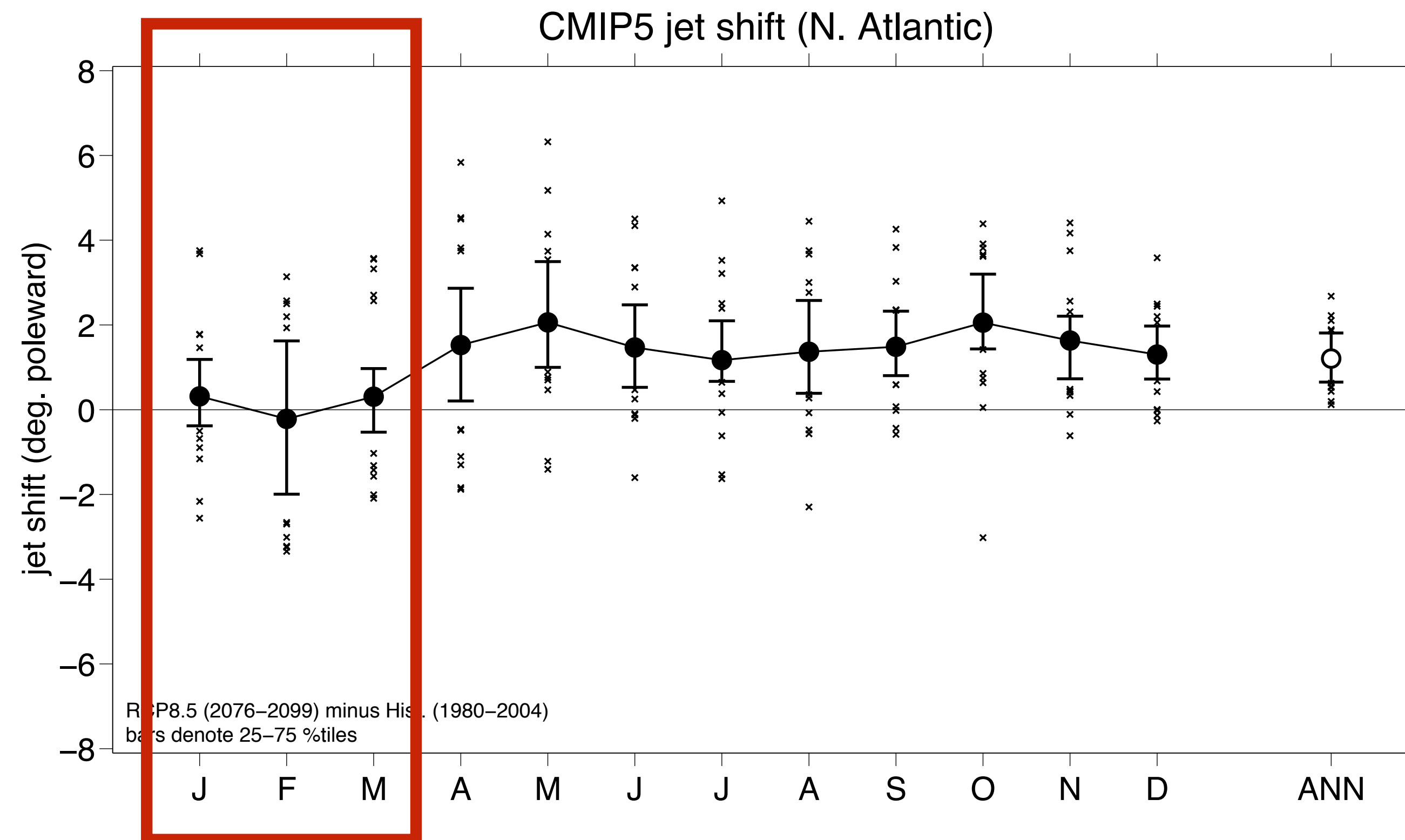


Shifts of the North Atlantic jet-stream by 2100



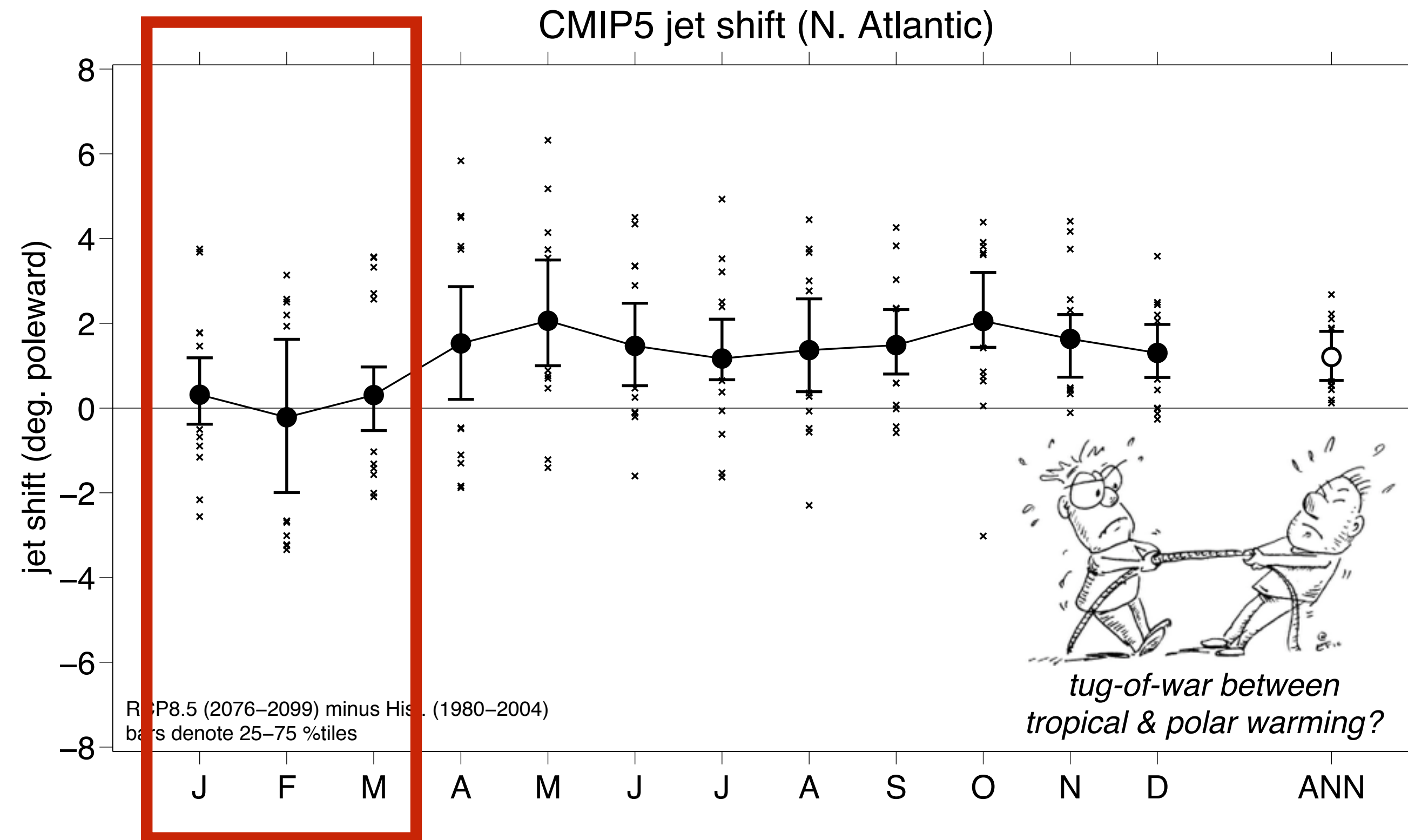
- jet-stream shifts poleward in most months of the year but not in winter
- interplay between high- and low latitude warming? (see Held (1993; BAMS), Harvey, Shaffrey et al. (2013), Cattiaux & Cassou (2013))

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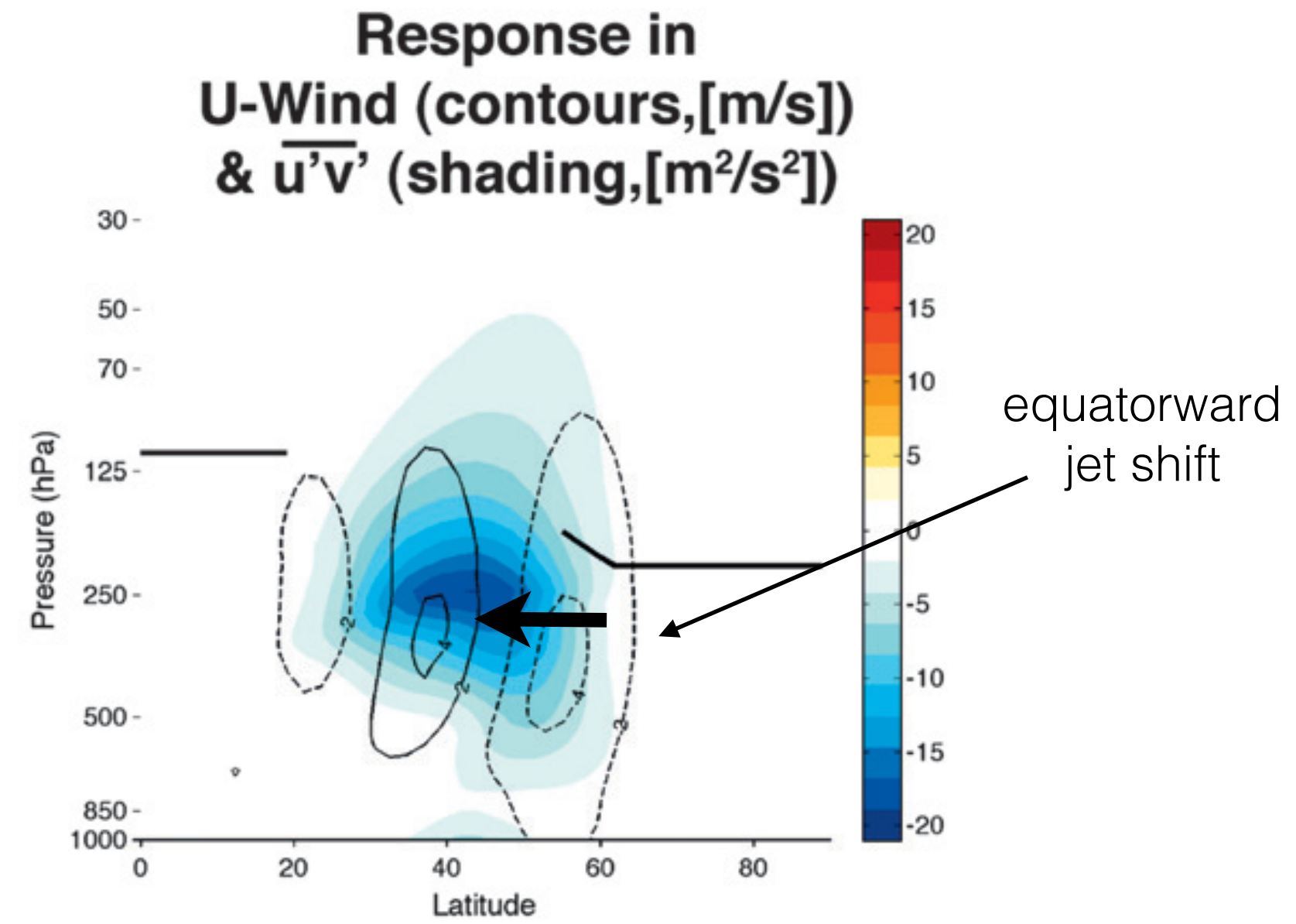
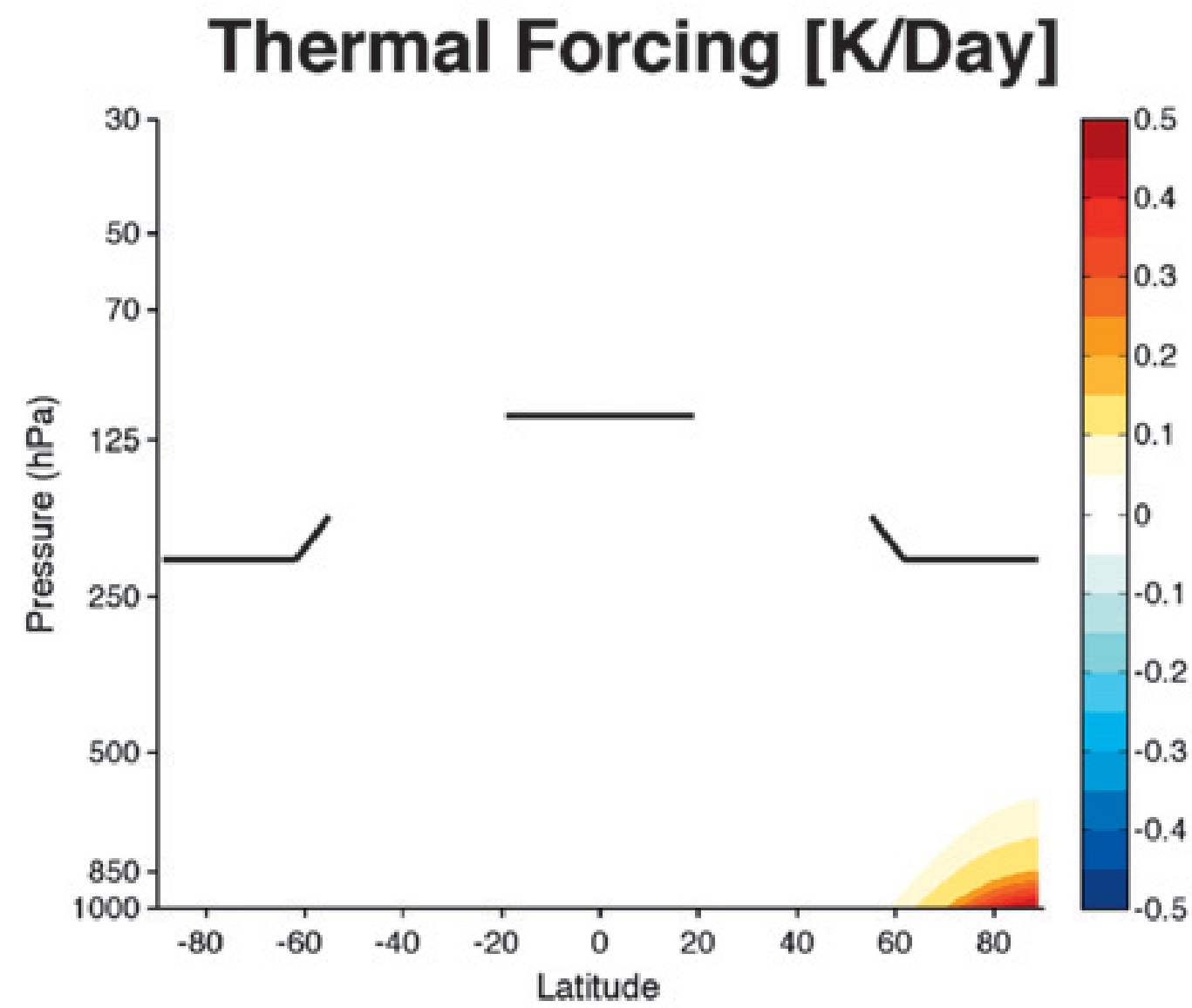
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Modeling evidence



simulations of a dry, dynamical core with imposed polar surface heating under perpetual equinox conditions

Butler, Thompson et al. (2010)

idealized modeling studies with polar heating show an equatorward jet shift when polar cap is heated

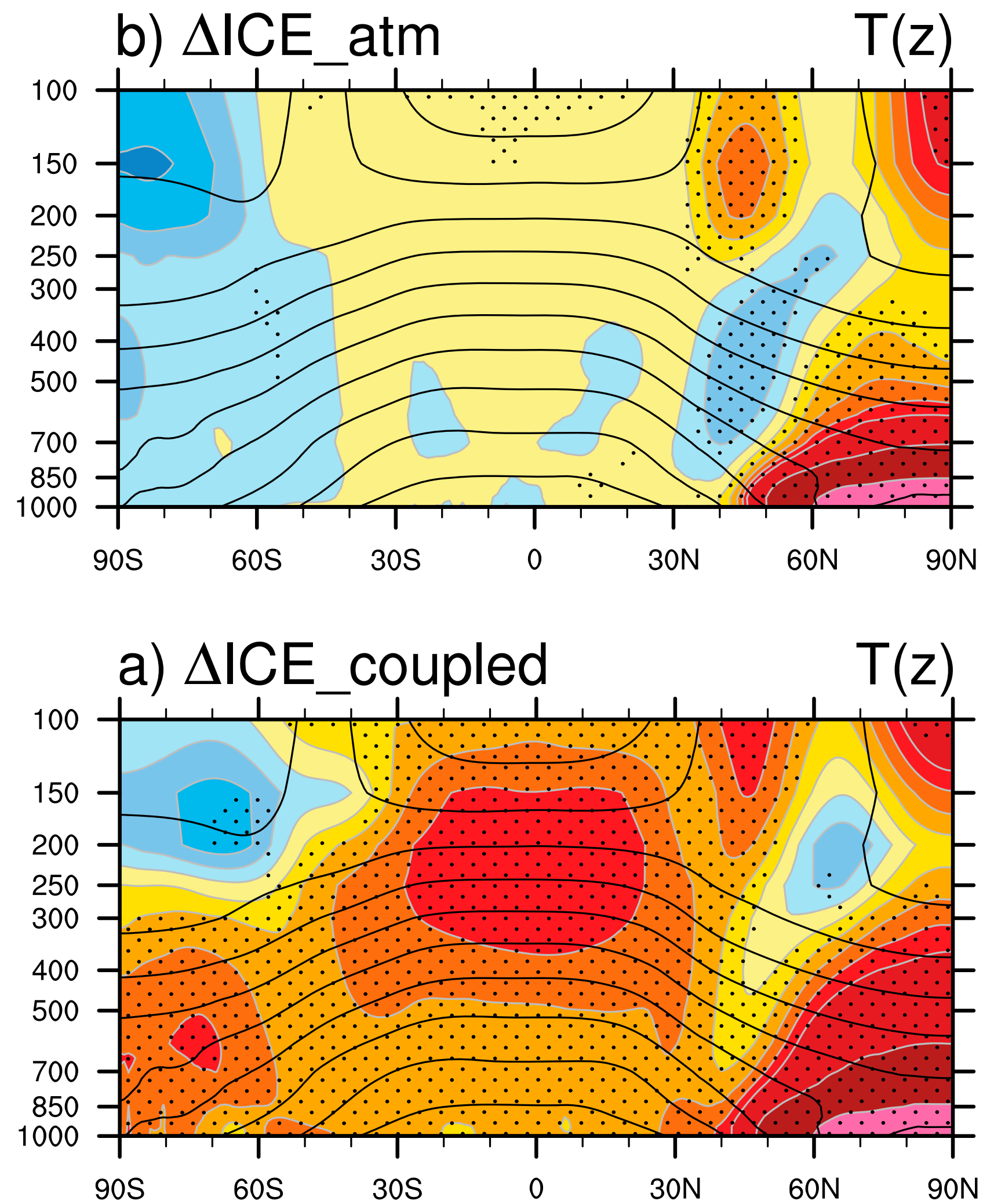
TABLE 1. Values of χ from equation 2 for every month.

Month	χ	Month	χ
JAN	+0.9659	JUL	-0.9659
FEB	+0.7071	AUG	-0.7071
MAR	+0.2588	SEPT	-0.2588
APR	-0.2588	OCT	+0.2588
MAY	-0.7071	NOV	+0.7071
JUN	-0.9659	DEC	+0.9659

$$T_{eq}^{trop}(p, \phi) = \max \left[200 \text{ K}, (T_0 - \delta T_{HS94}) \left(\frac{p}{p_0} \right)^\kappa \right],$$

$$\delta T_{HS94} = (\Delta T)_y \sin \phi^2 + \epsilon \chi \sin \phi + (\Delta T)_z \log \left(\frac{p}{p_0} \right) \cos \phi^2,$$

New results from Deser et al. (2015)

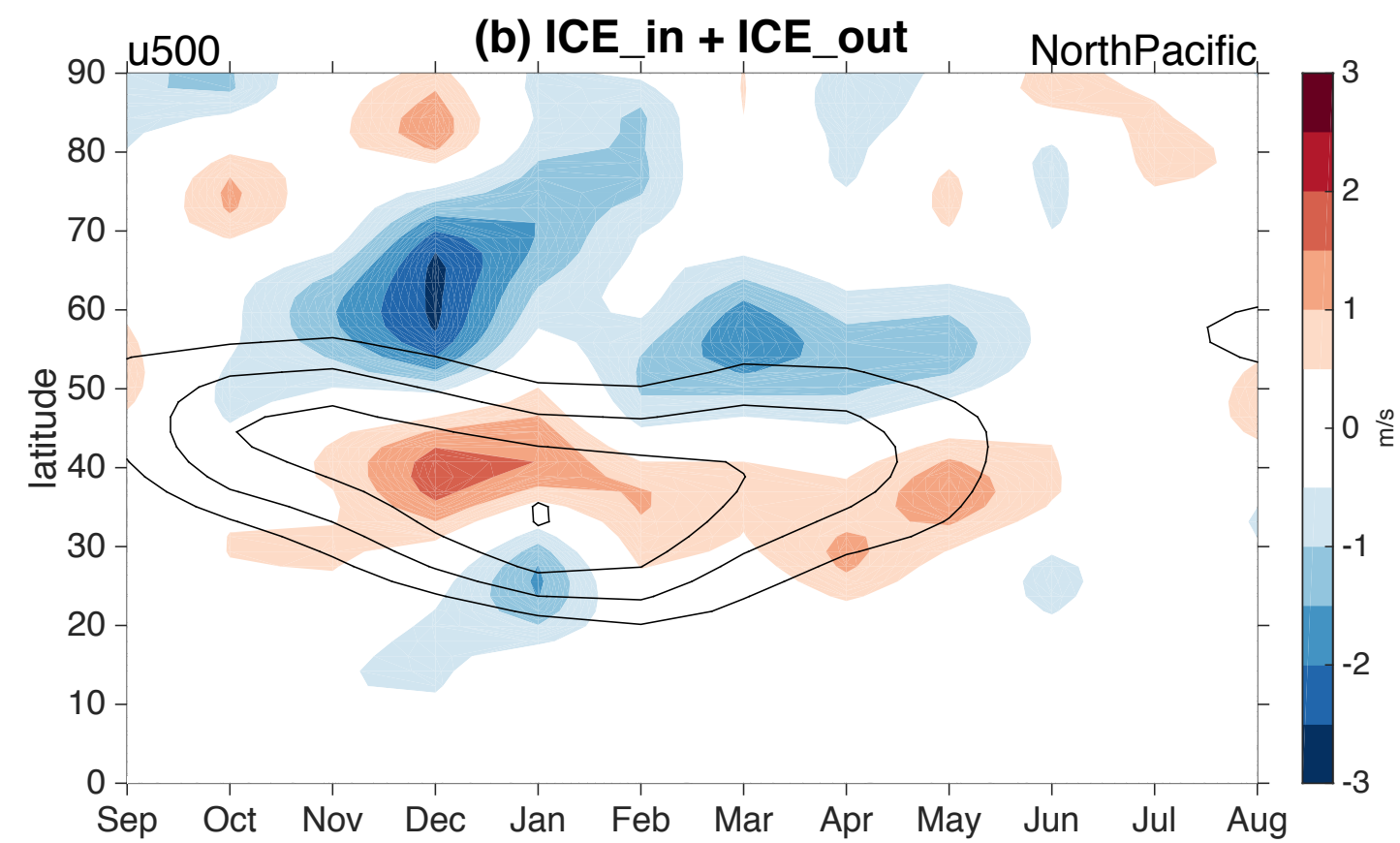
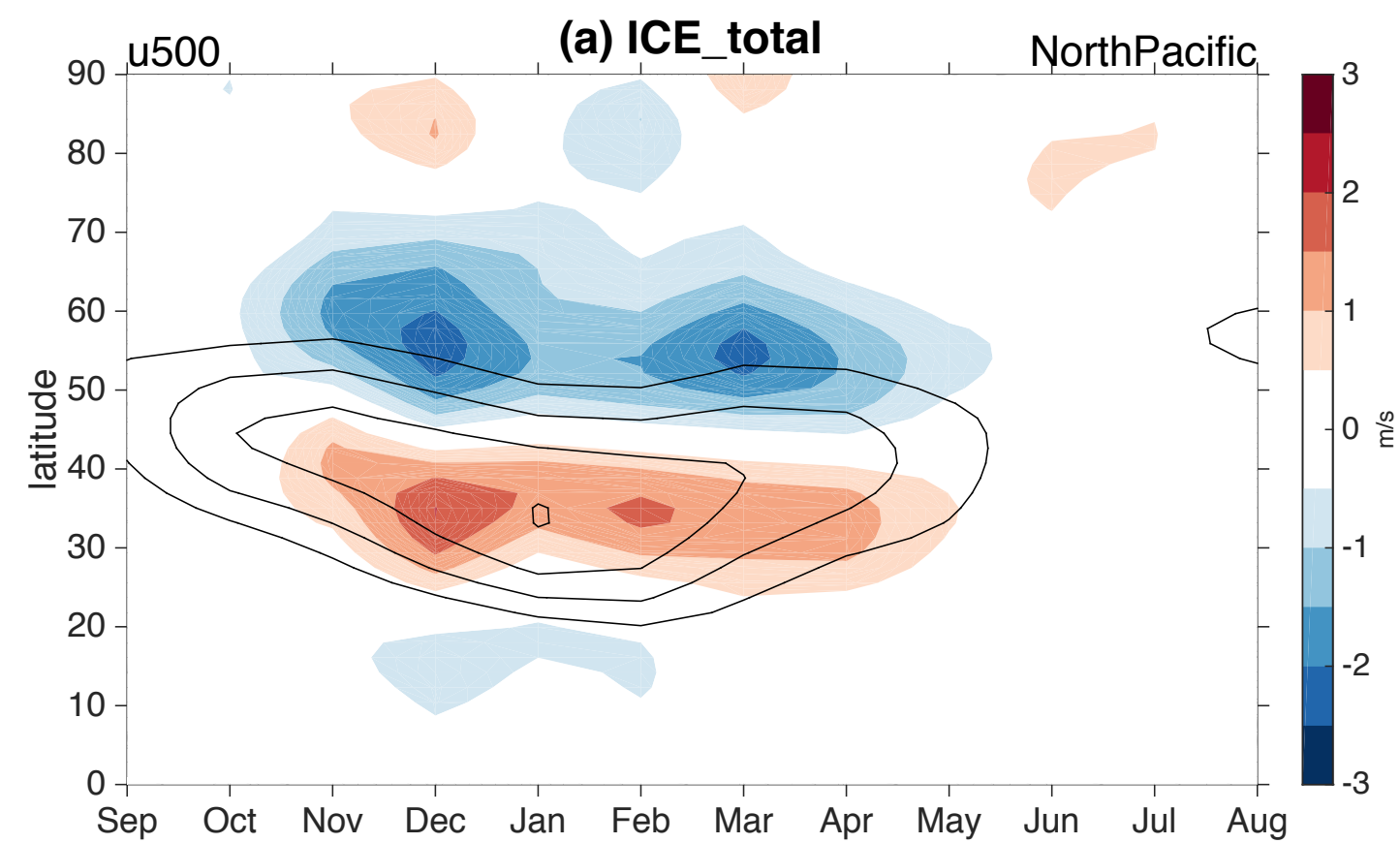
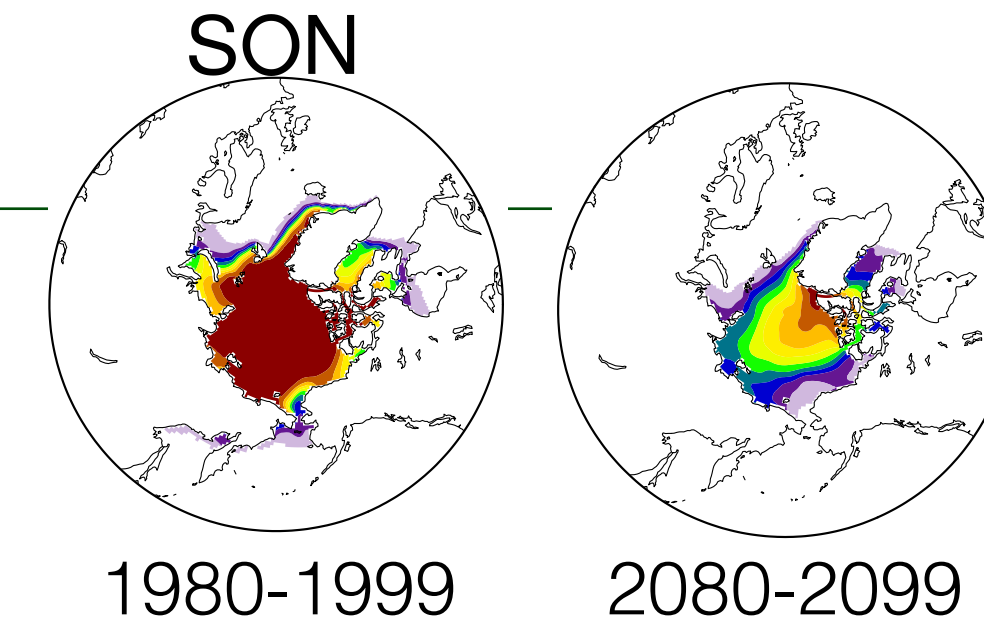


AMIP simulations may underestimate sea ice-induced warming compared to coupled simulations

*coupled CCSM4 simulations
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Deser, Tomas, et al. (2015; JCLI)*

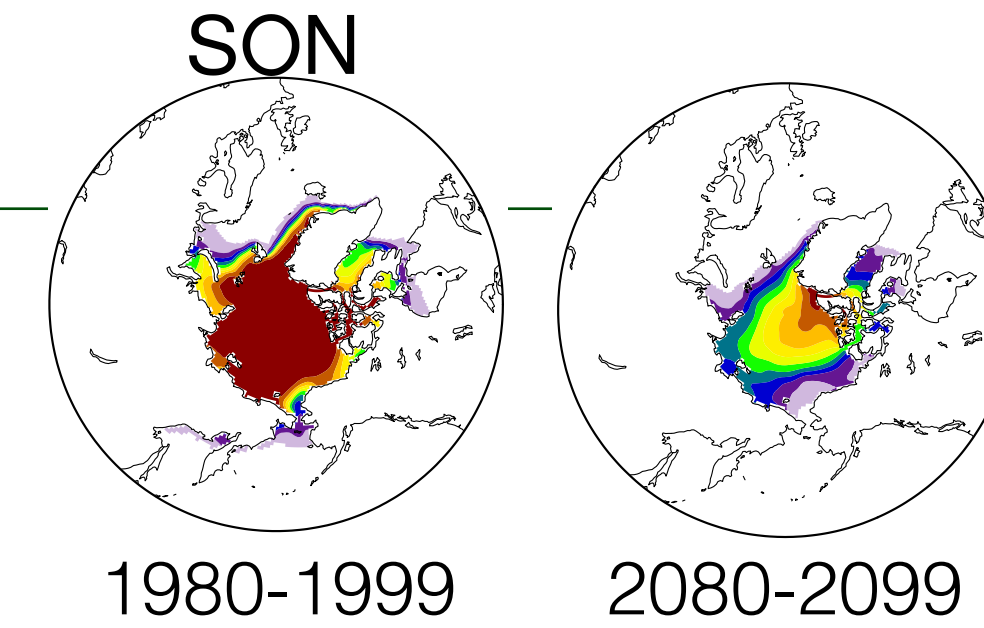
Nonlinear response

North Pacific jet response to sea ice loss **inside** and **outside** of the Arctic circle.

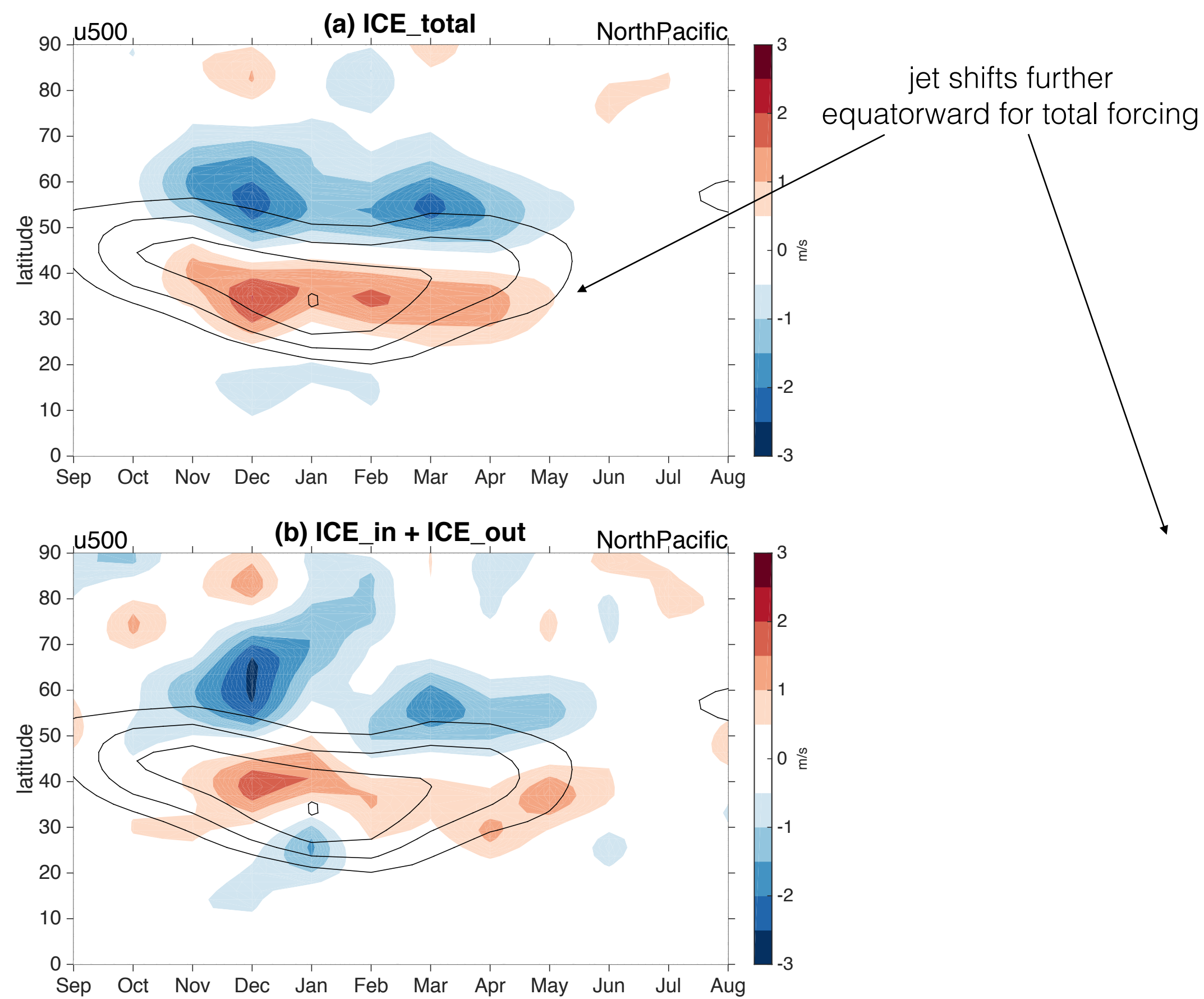


based on 160-year WACCM simulations of
Sun et al. (2015; JCLI)

Nonlinear response

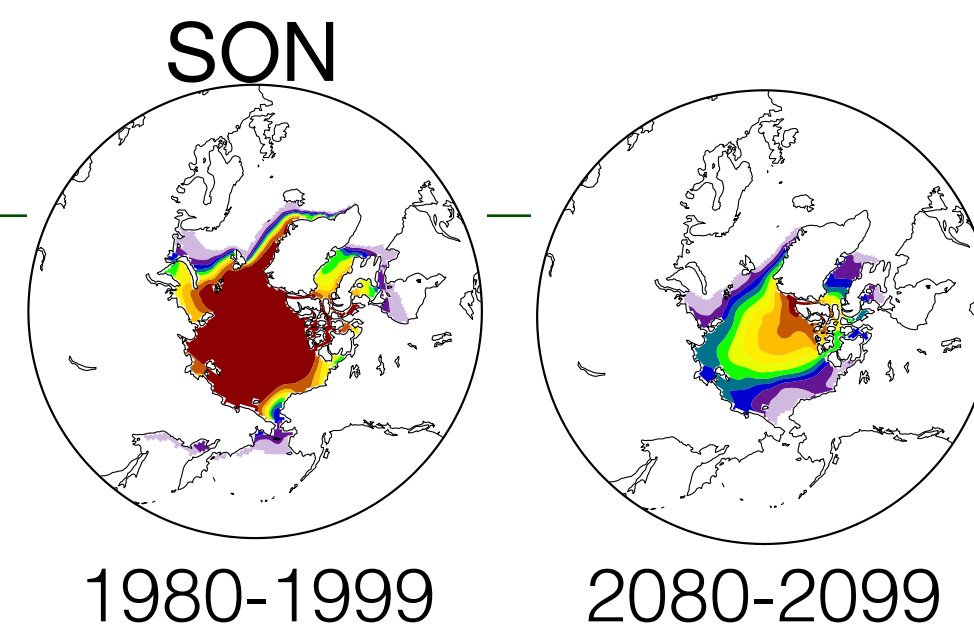


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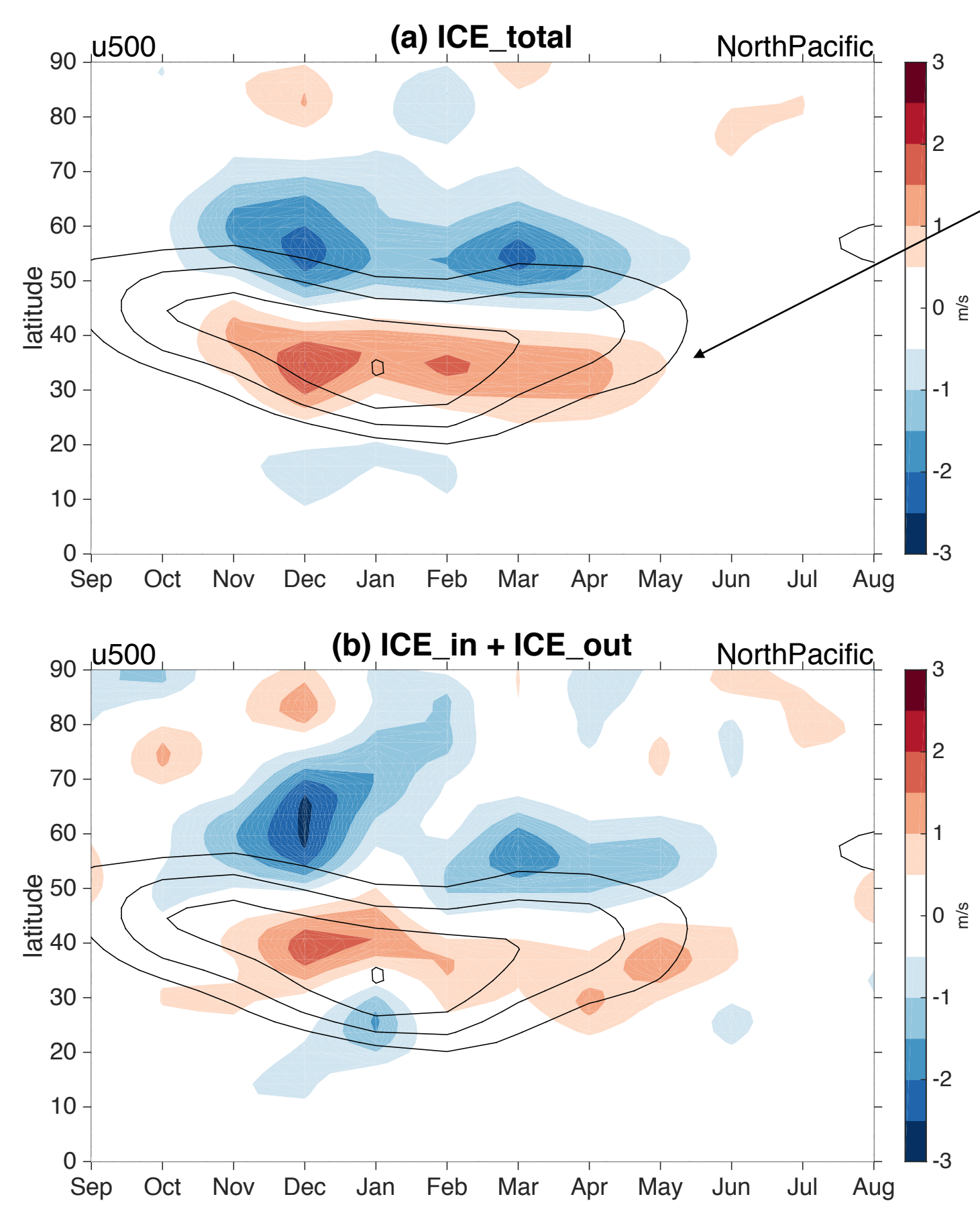


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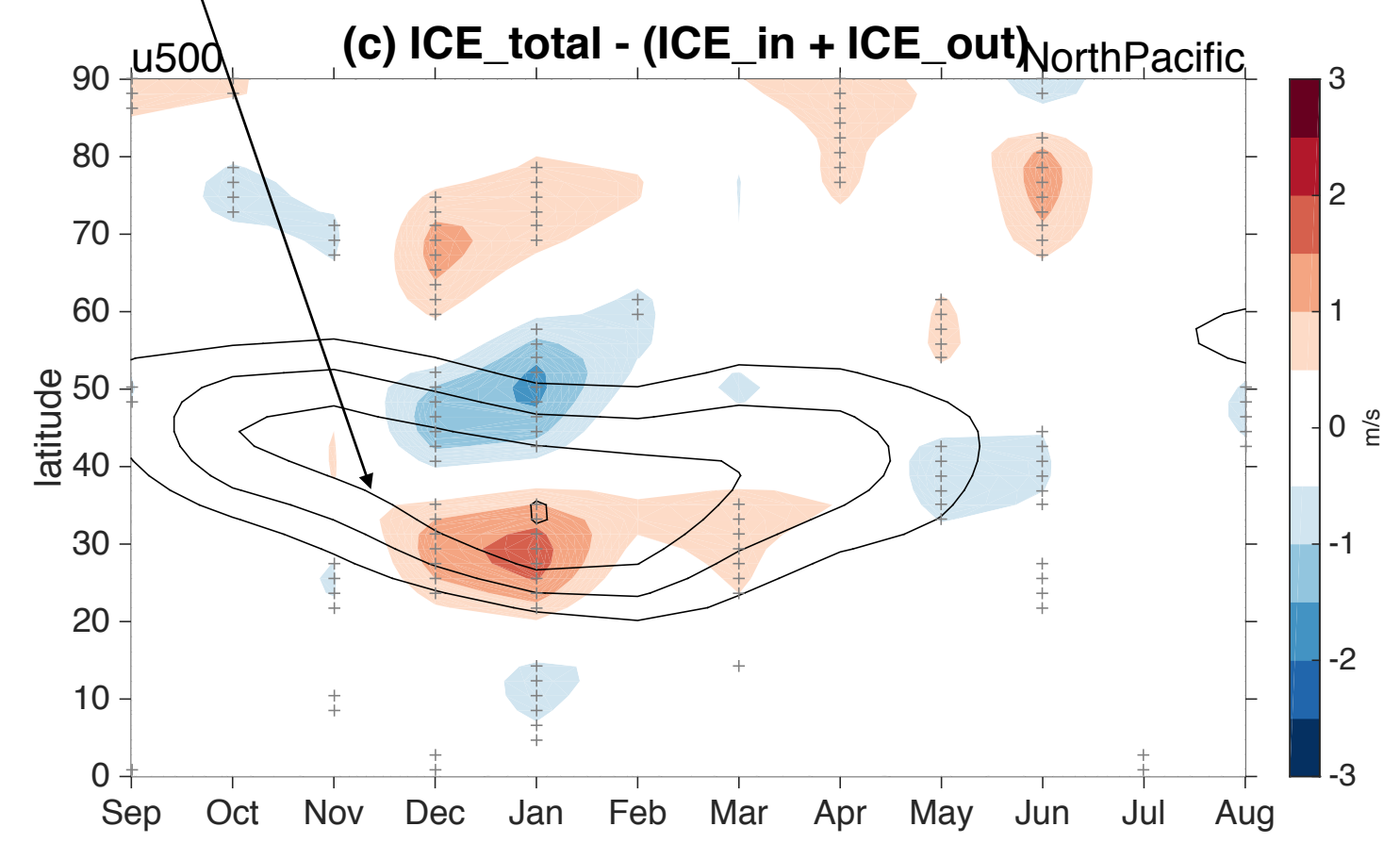
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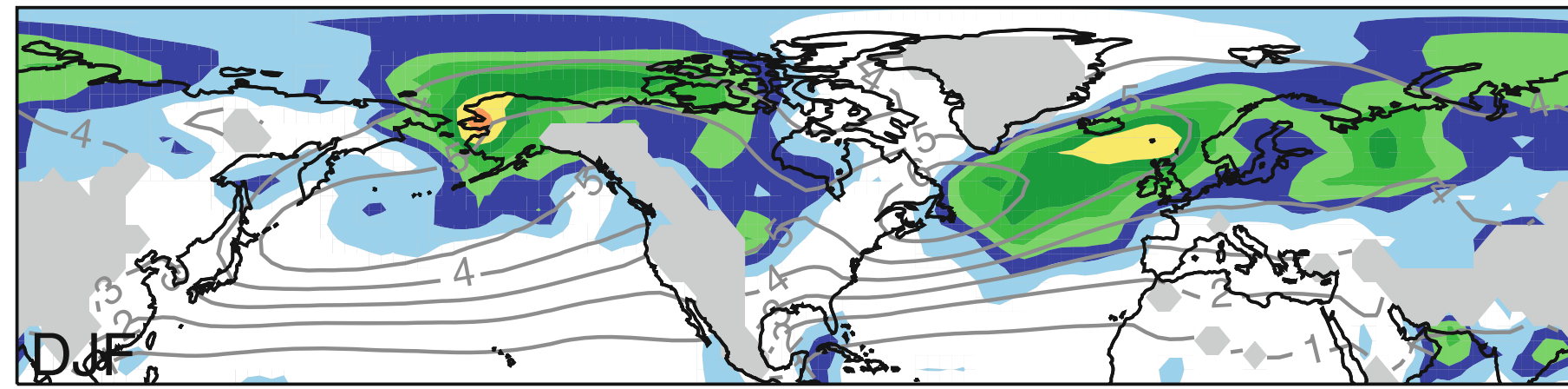
jet shifts further equatorward for total forcing



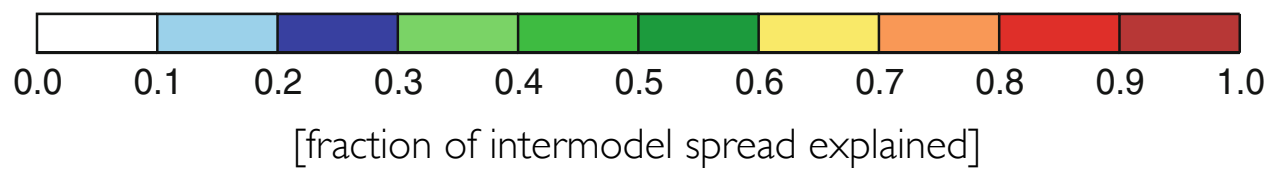
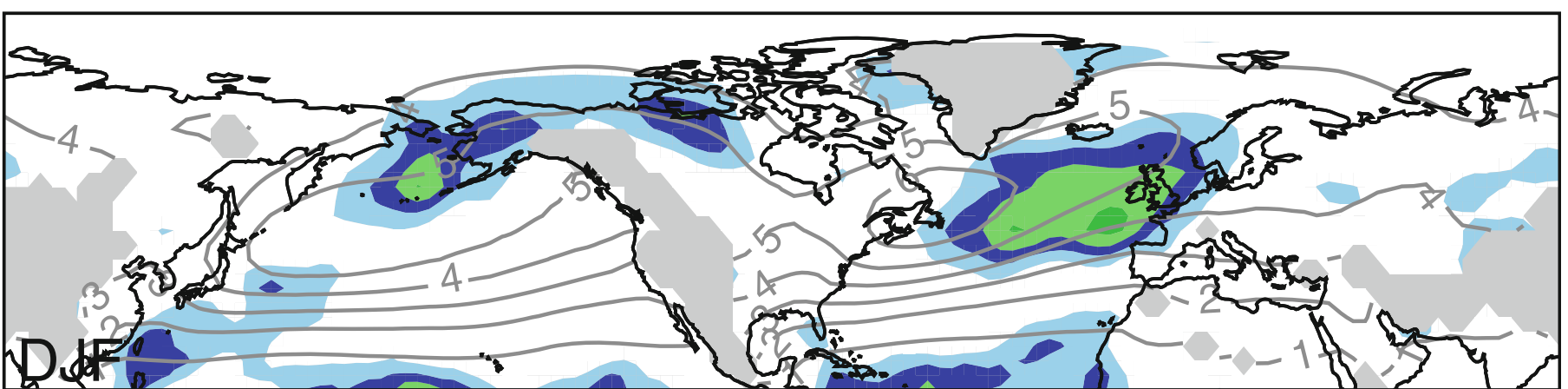
based on 160-year WACCM simulations of Sun et al. (2015; JCLI)

Will it modulate?

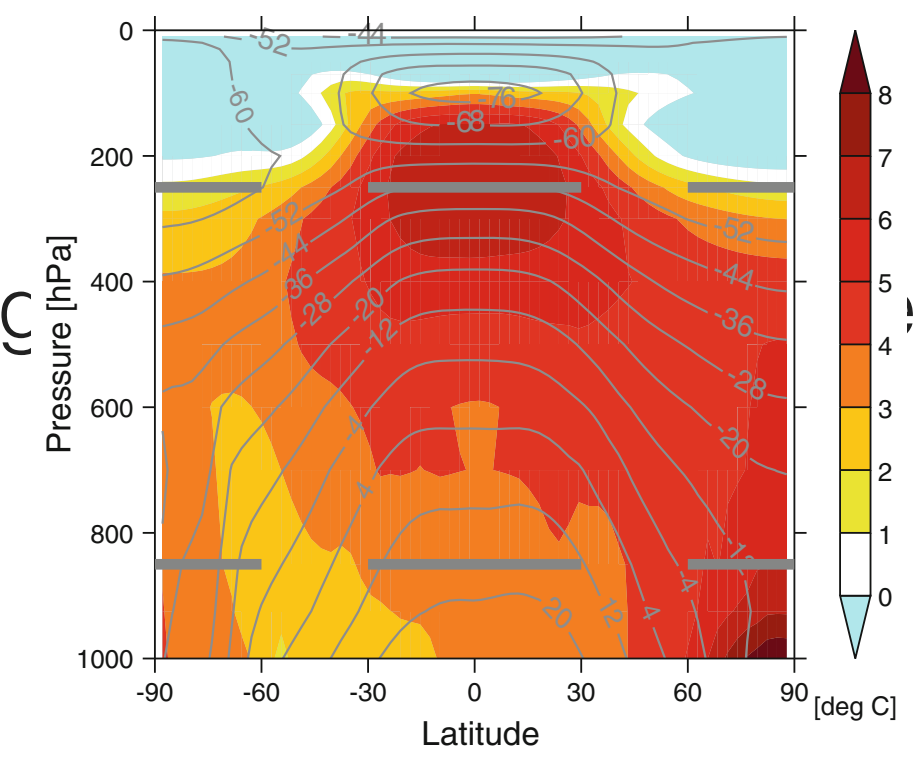
- In the N. Atlantic, 850 hPa and 250 hPa responses in winter
- Variance explained of storm track response by N. Atlantic temperature gradients



250 hPa



850 hPa and 250 hPa changes



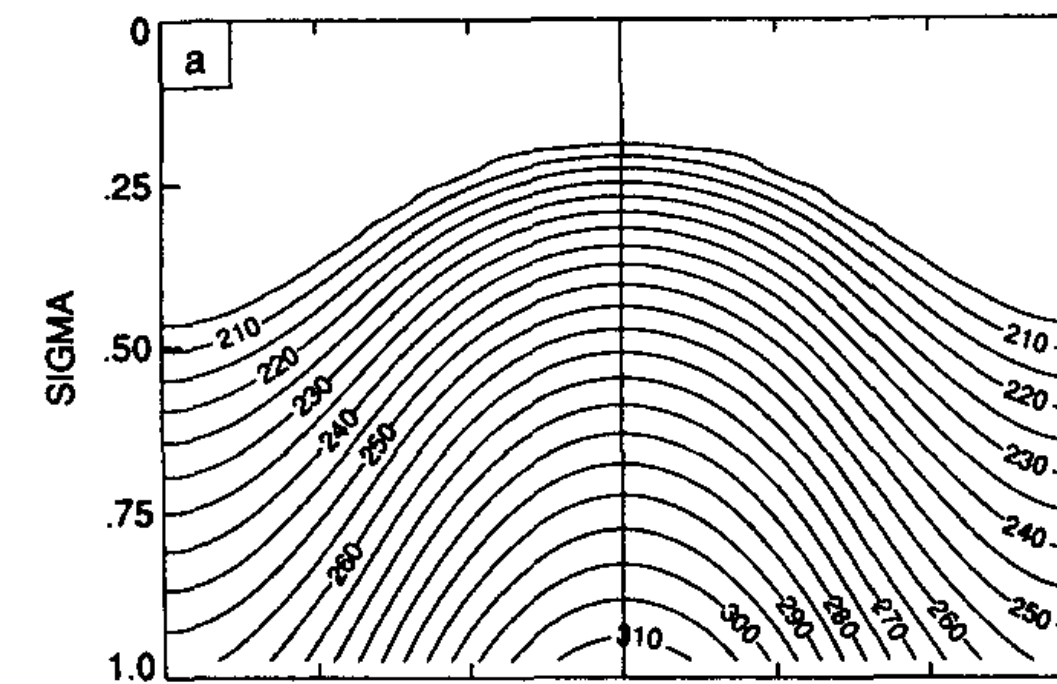
temperature response under RCP8.5 (1976-2005) - (2070-2099)

CMIP5 model analysis
storm track = variance of 2-6 dy. SLP
Harvey, Shaffrey et al. (2013)

GFDL Dynamical Core simulations

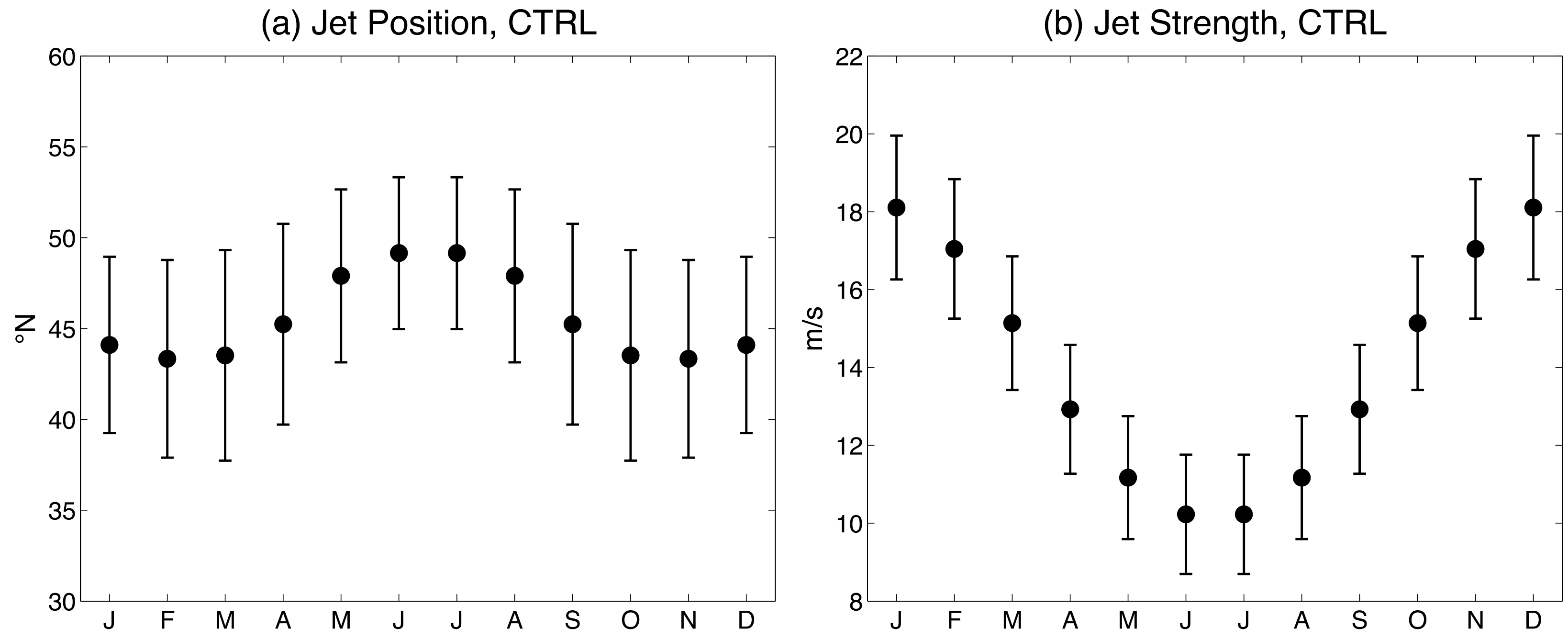
- driven by Newtonian relaxation to an equilibrium temperature profile
- simulation 360-day seasonal cycle by varying the equilibrium temperature profile (as in Polvani & Kushner, 2002)
- no well-resolved stratosphere
- zonally-symmetric
- each month is run under perpetual conditions (e.g. perpetual January, perpetual February...)

equilibrium
temperature profile

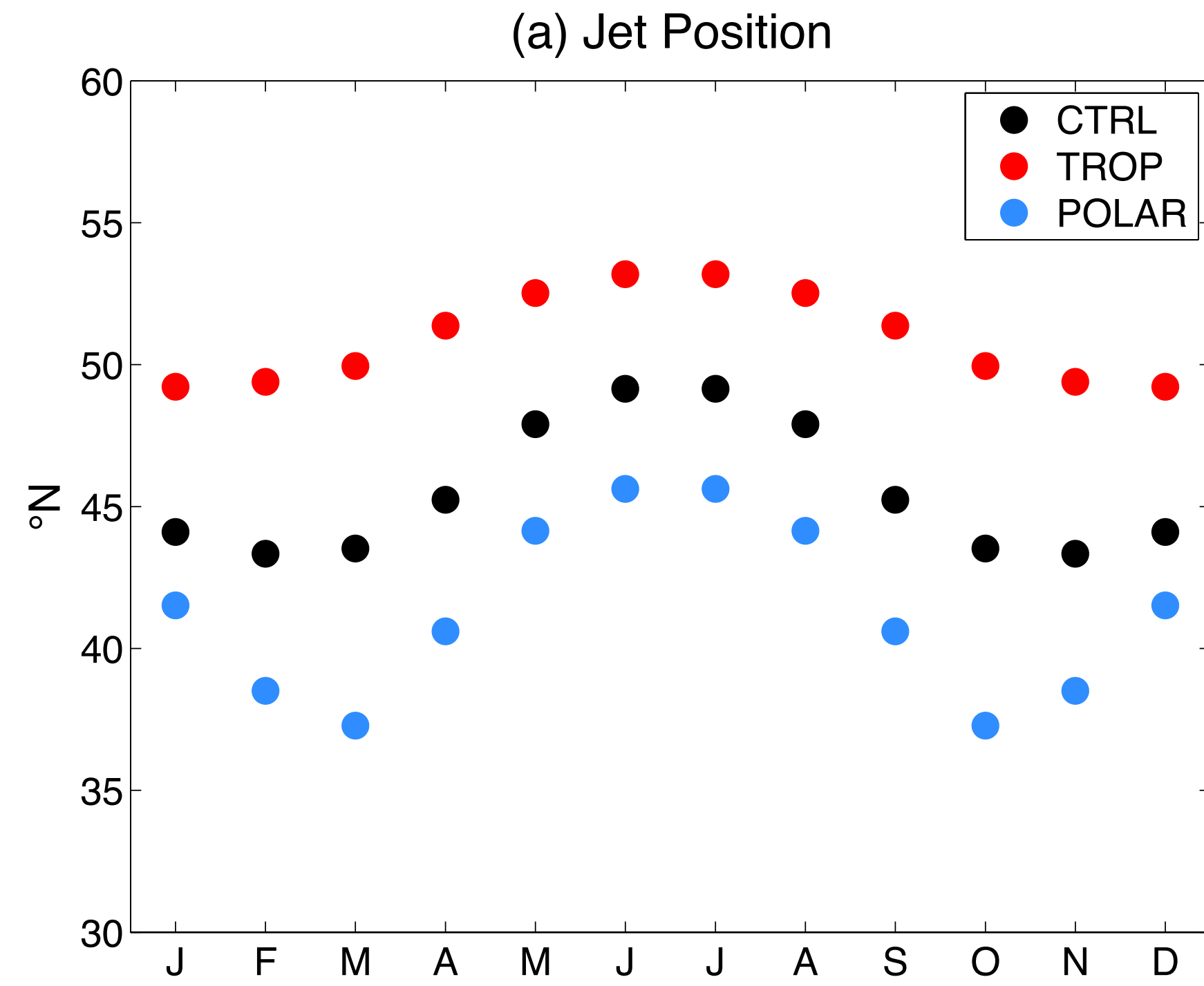
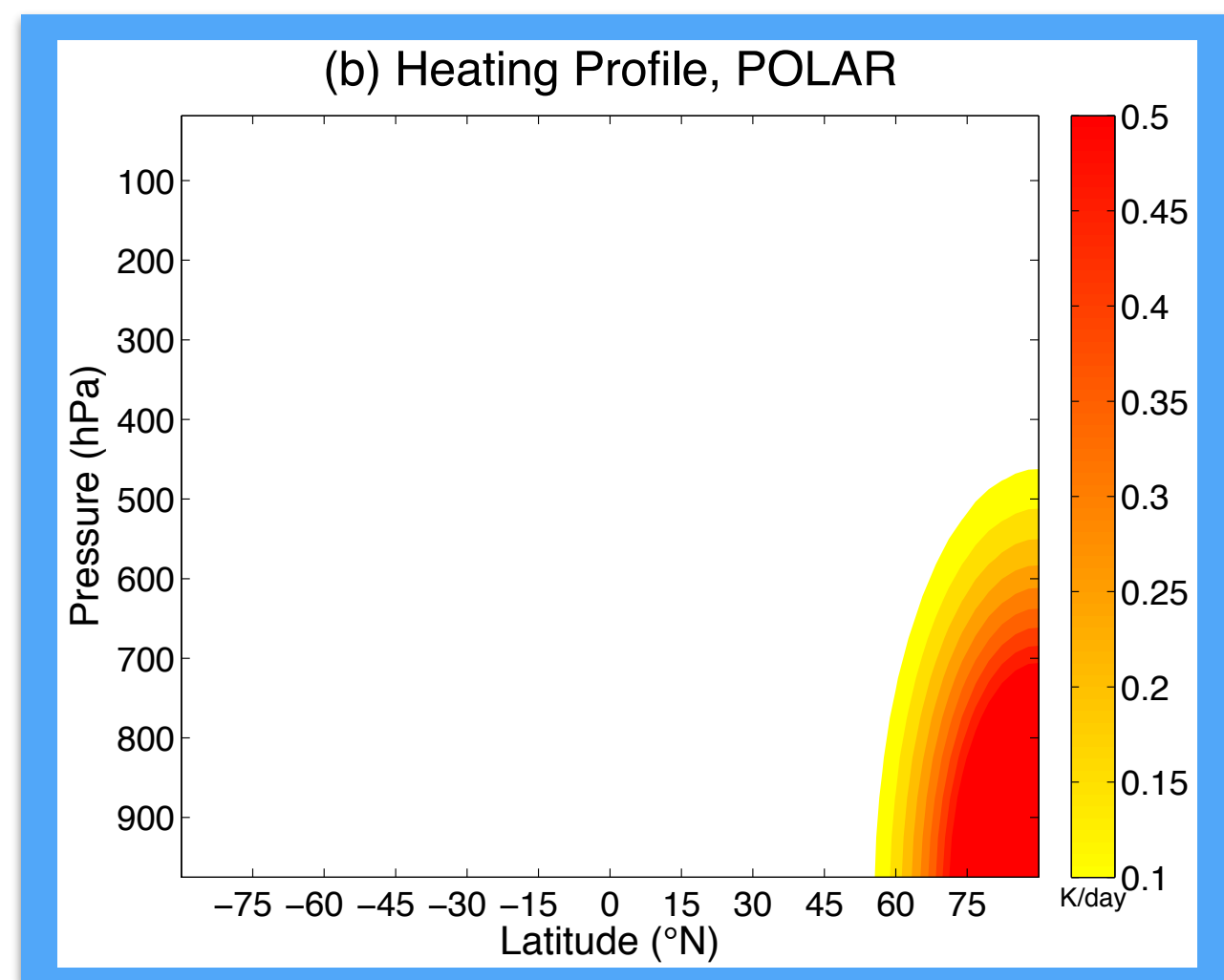
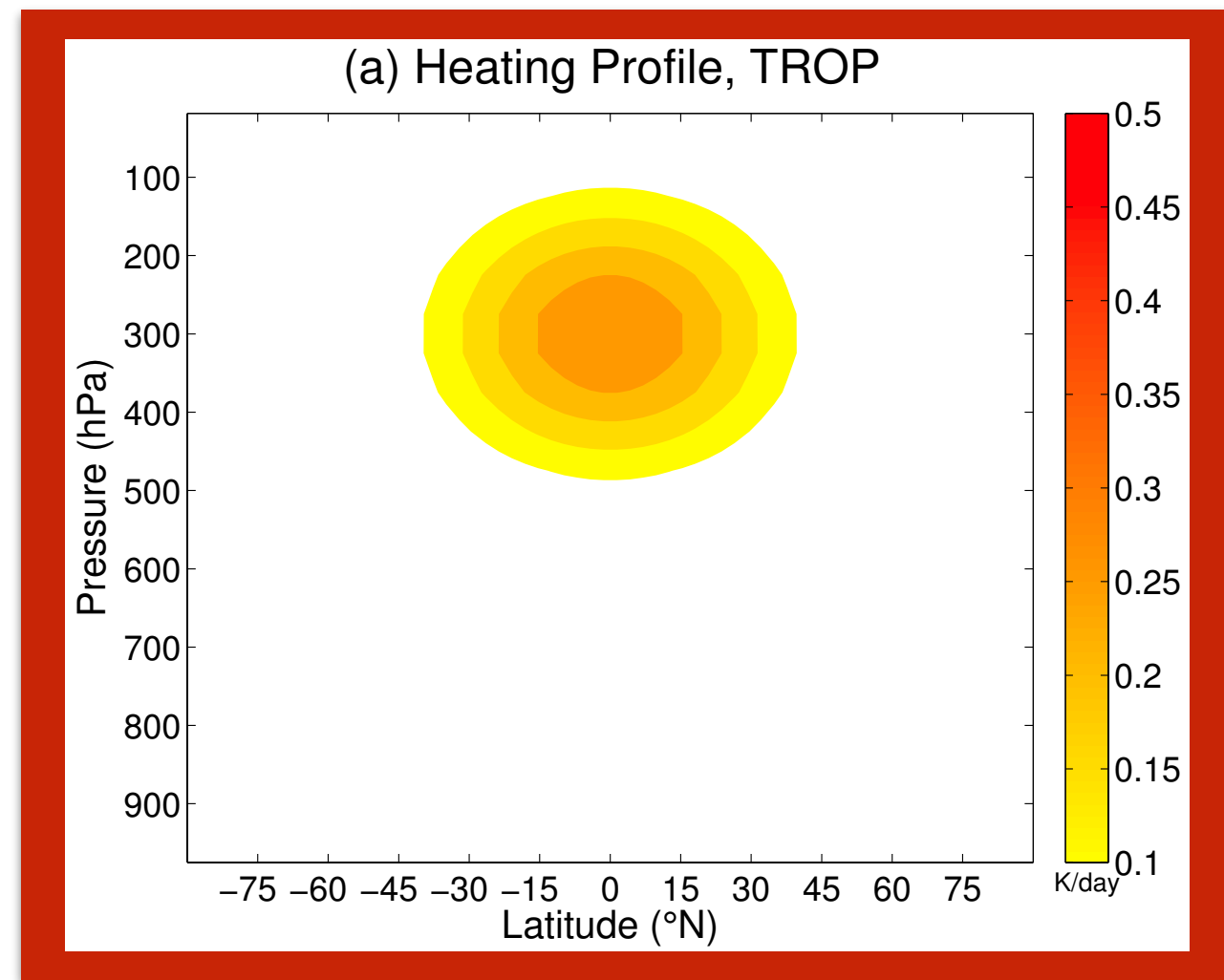


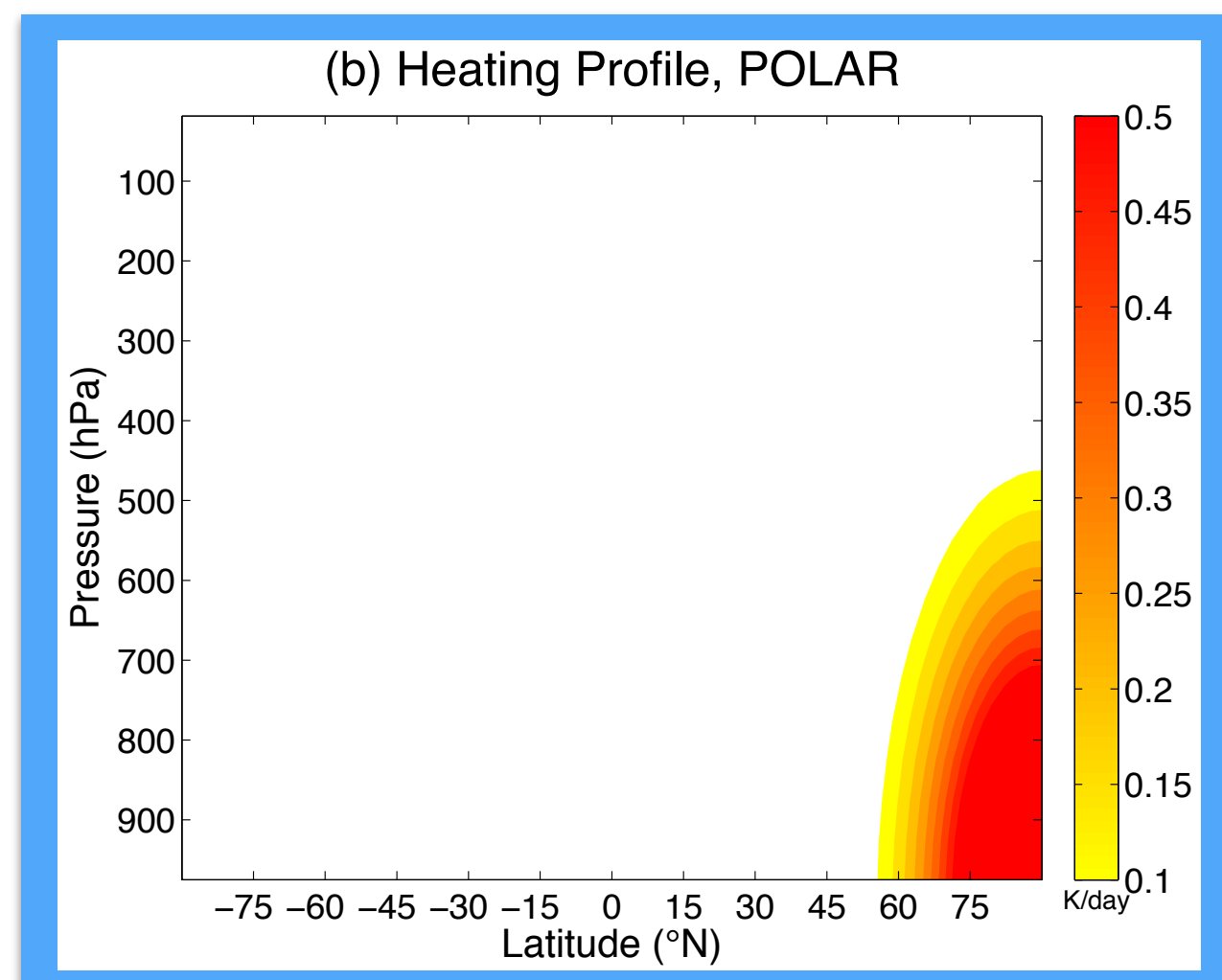
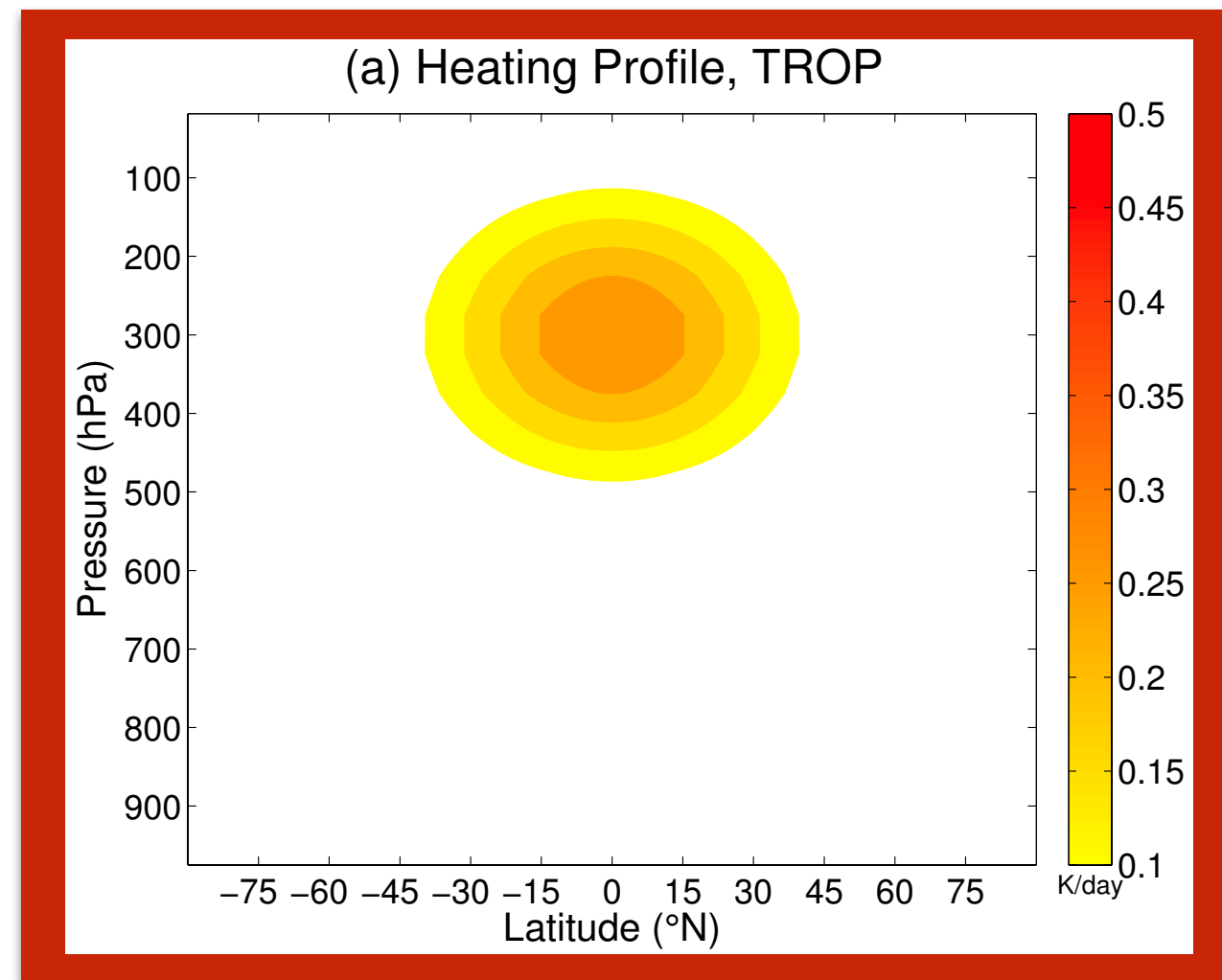
Held & Suarez (1994)
Polvani and Kushner (2002)

Control simulations

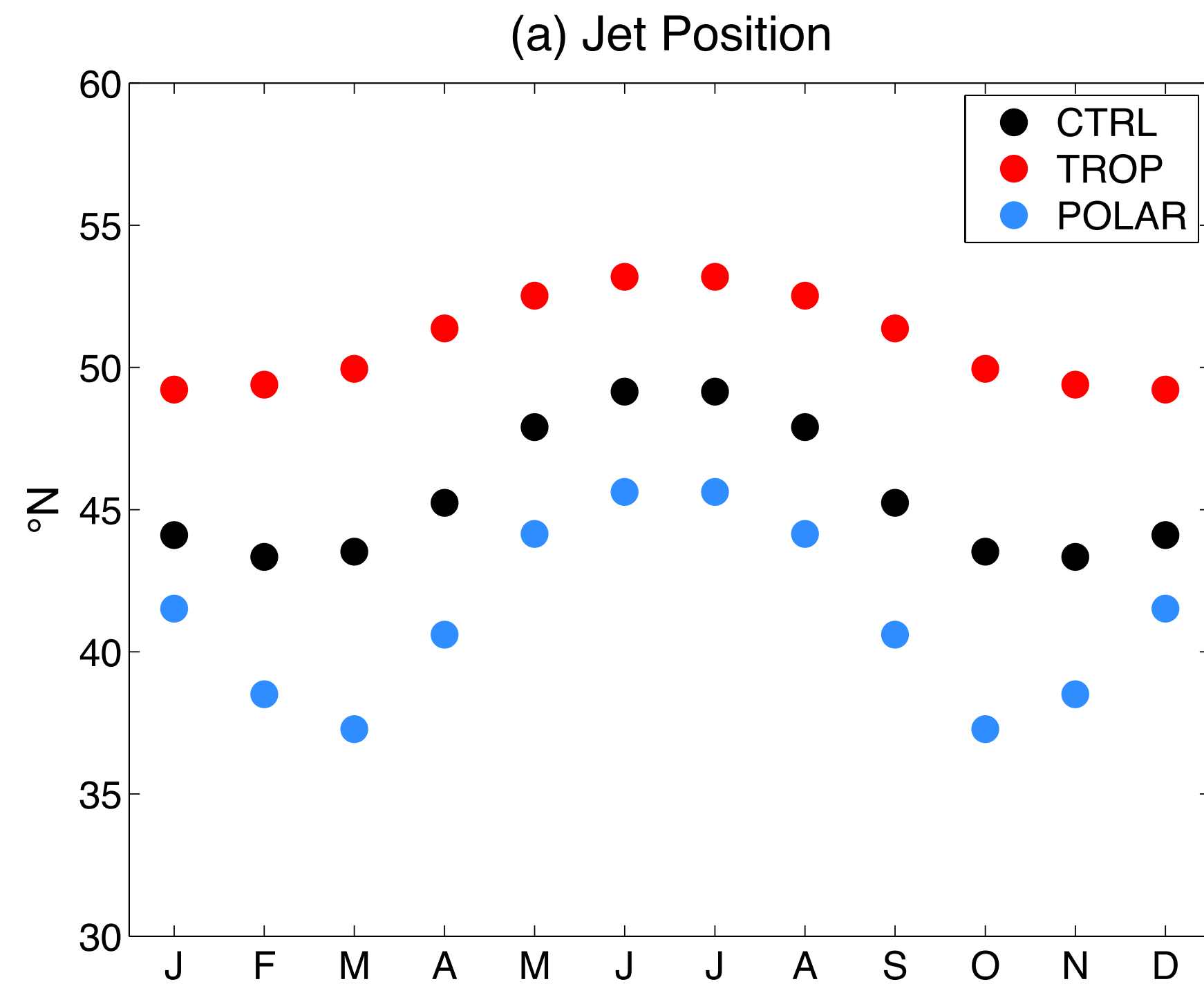


McGraw & Barnes (2016); in review



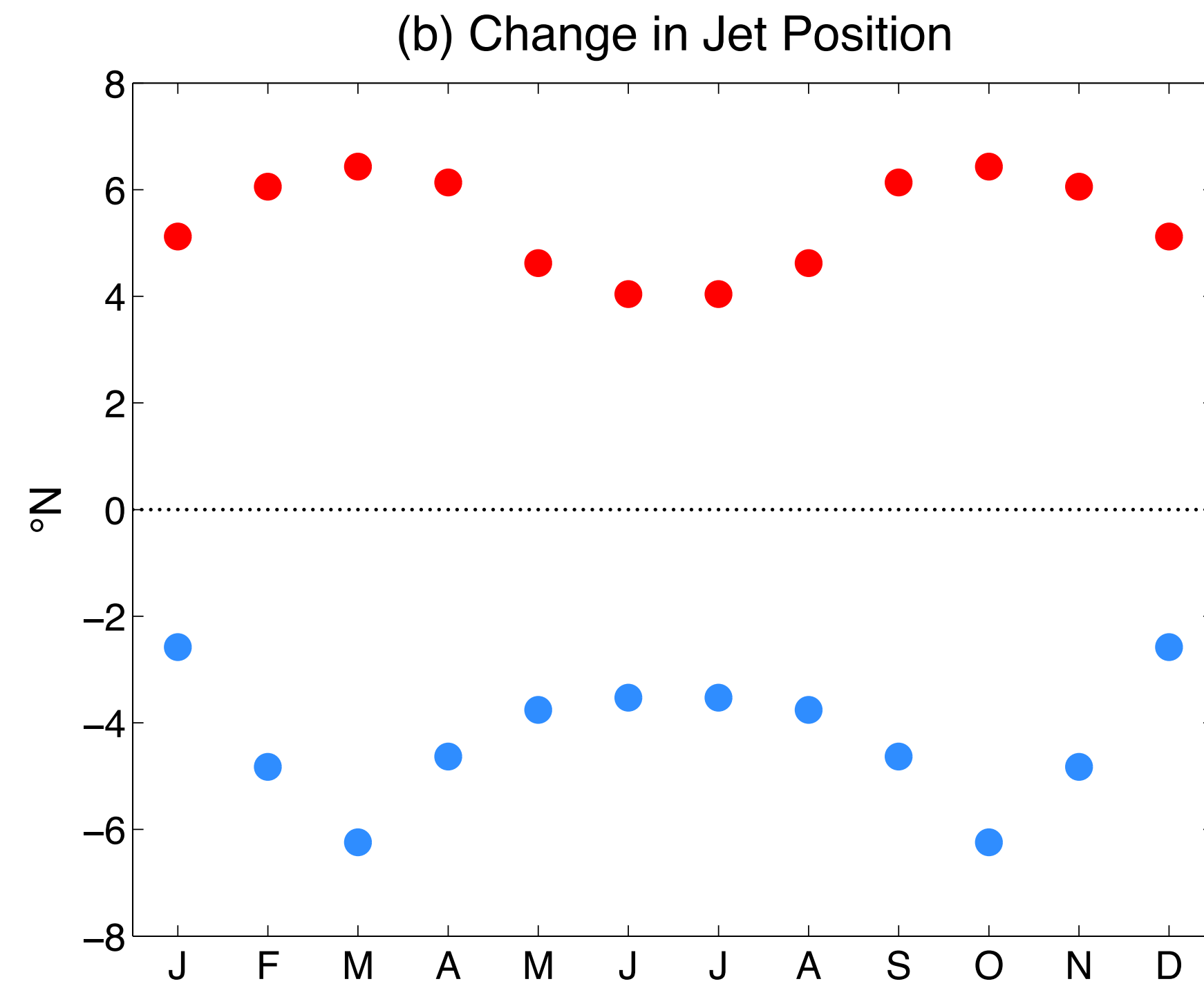


Dry core: Jet-stream response



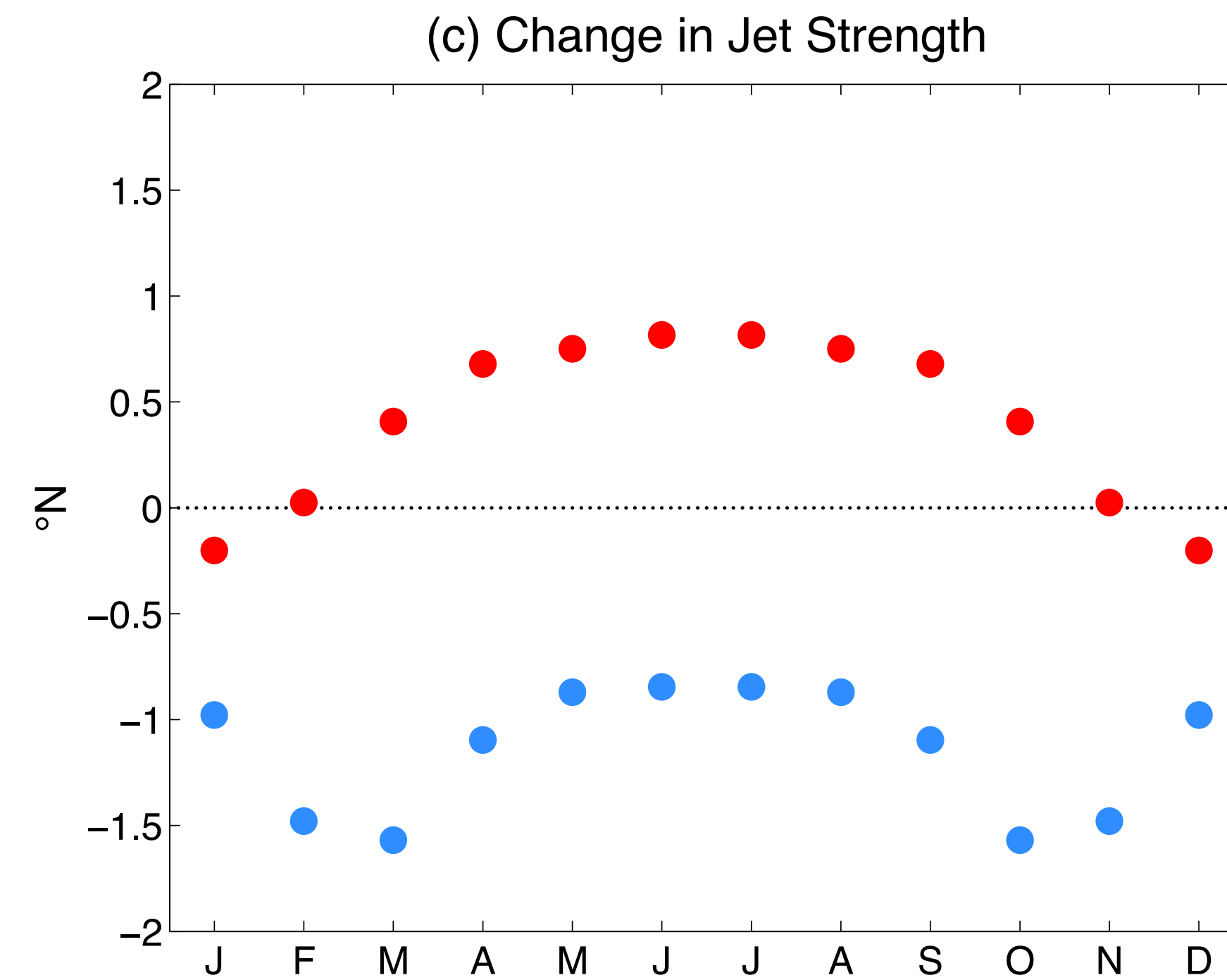
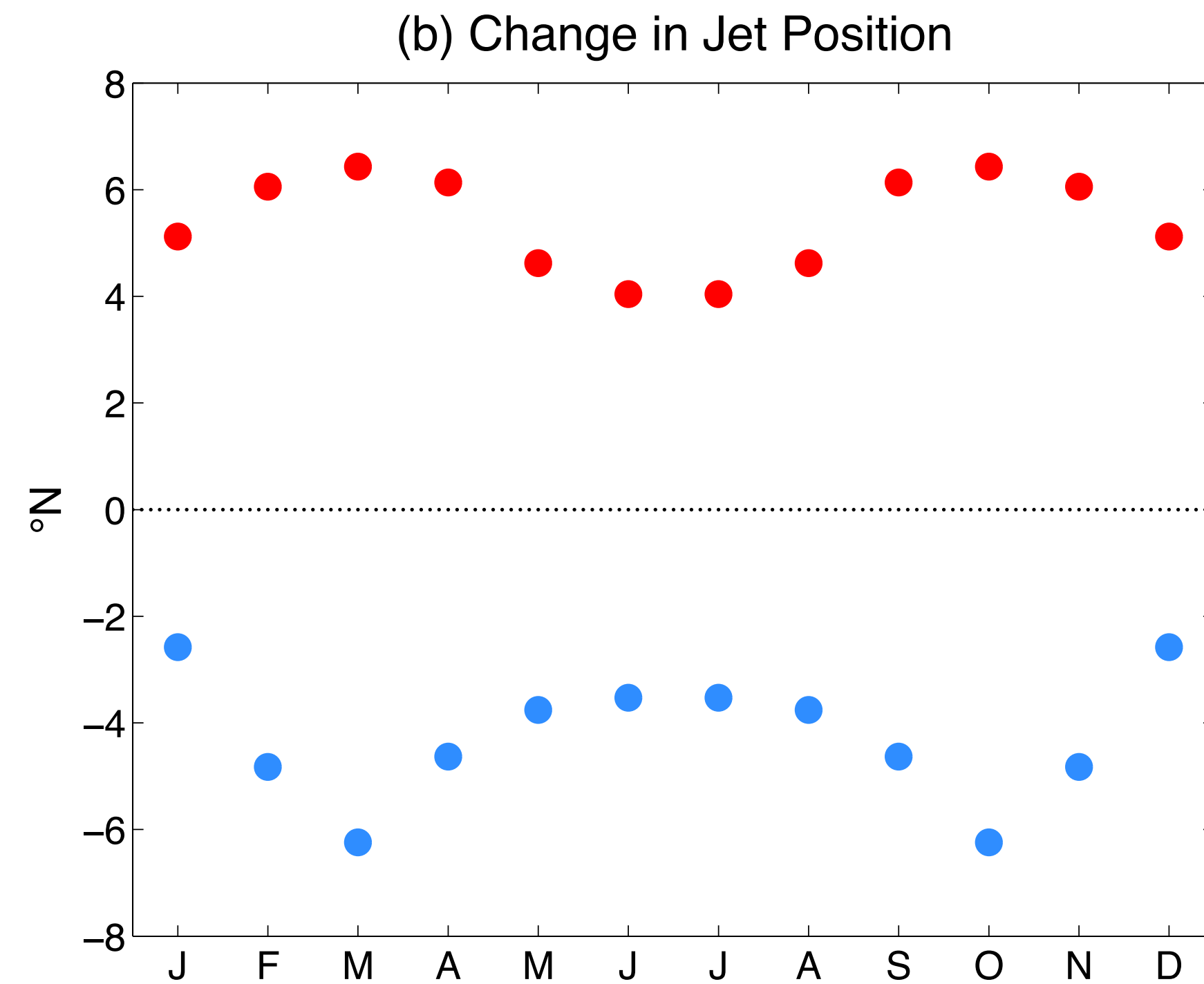
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Dry core: Jet-stream response



McGraw & Barnes (2016); in review

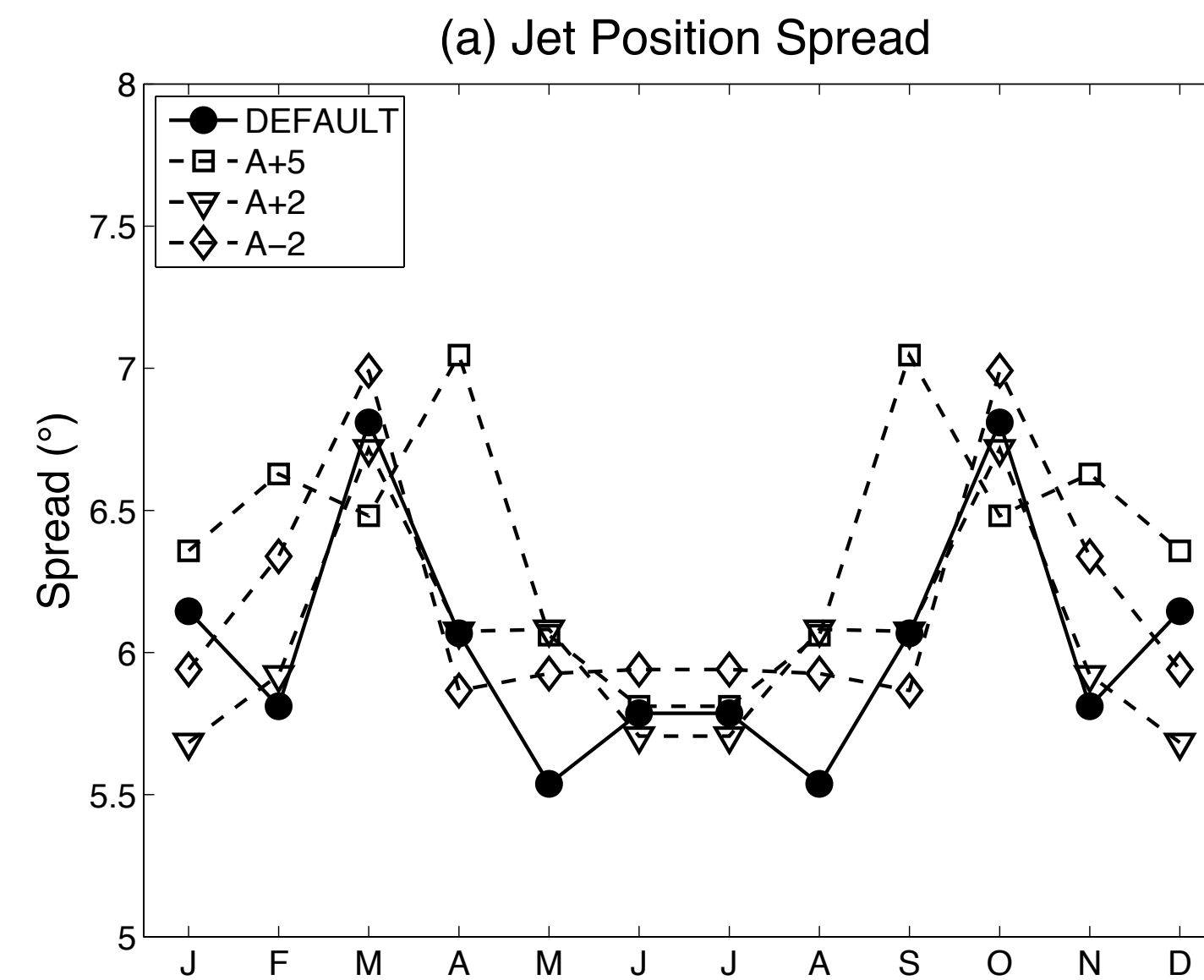
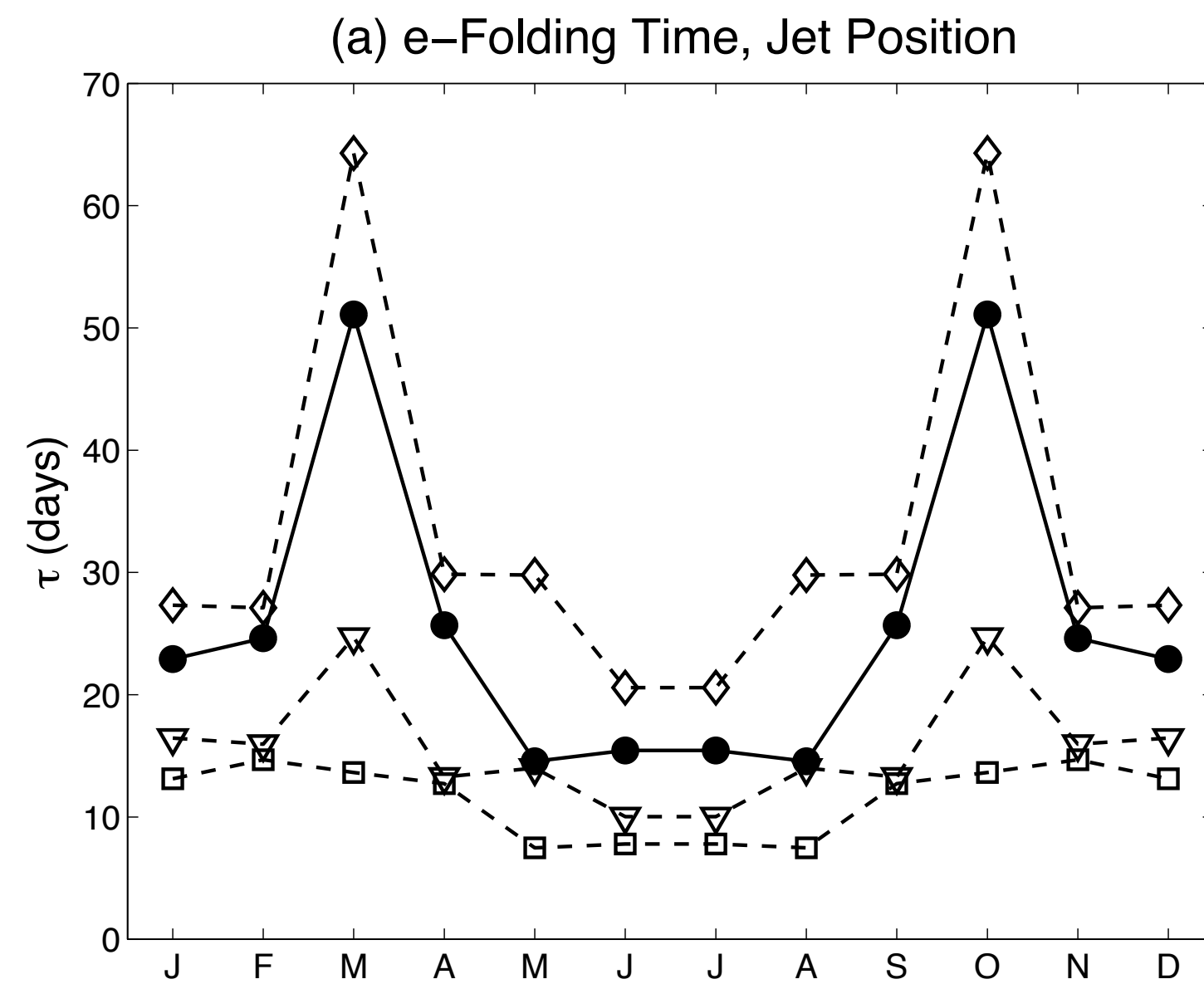
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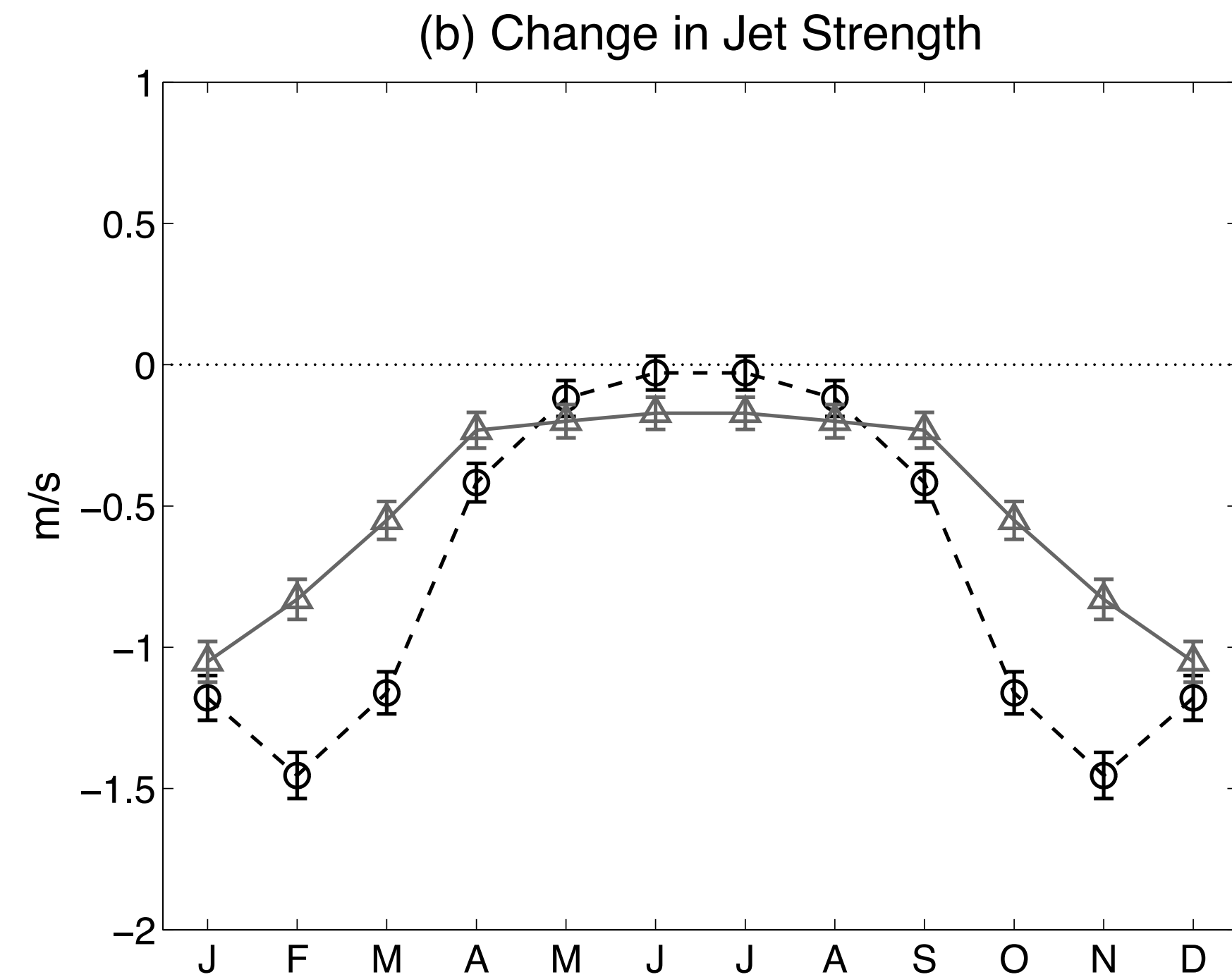
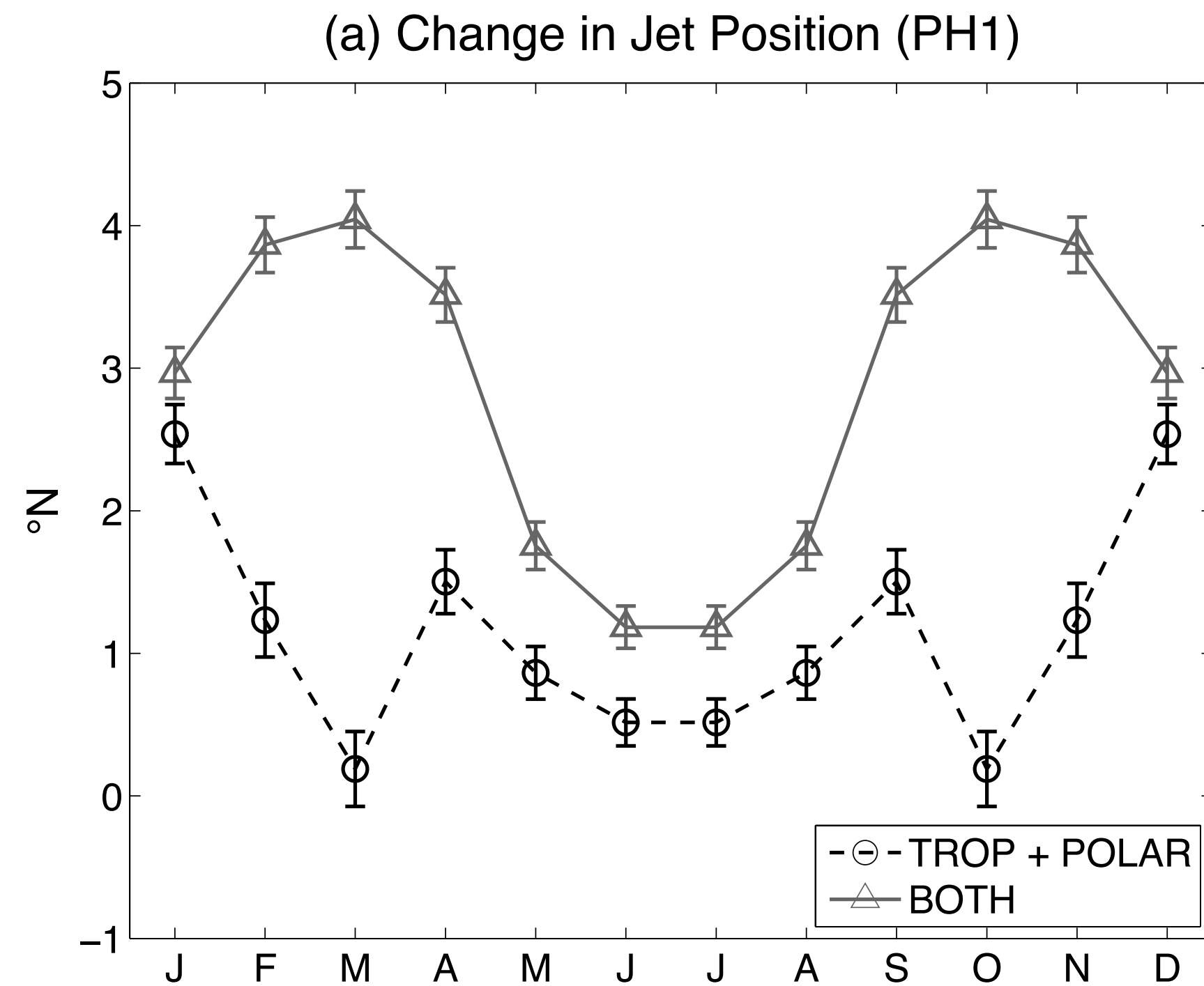
McGraw & Barnes (2016); in review

Dry core: Fluctuation-dissipation?

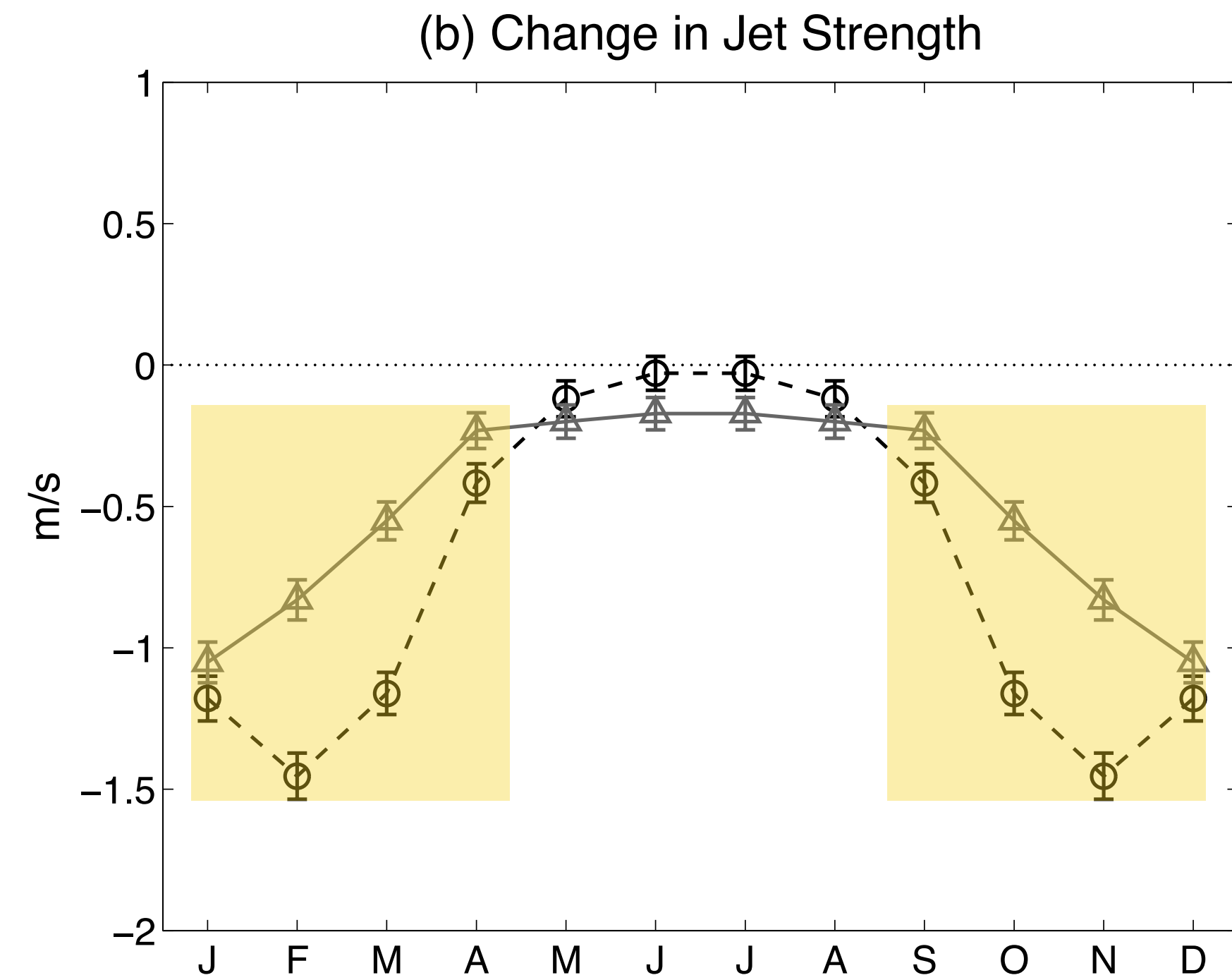
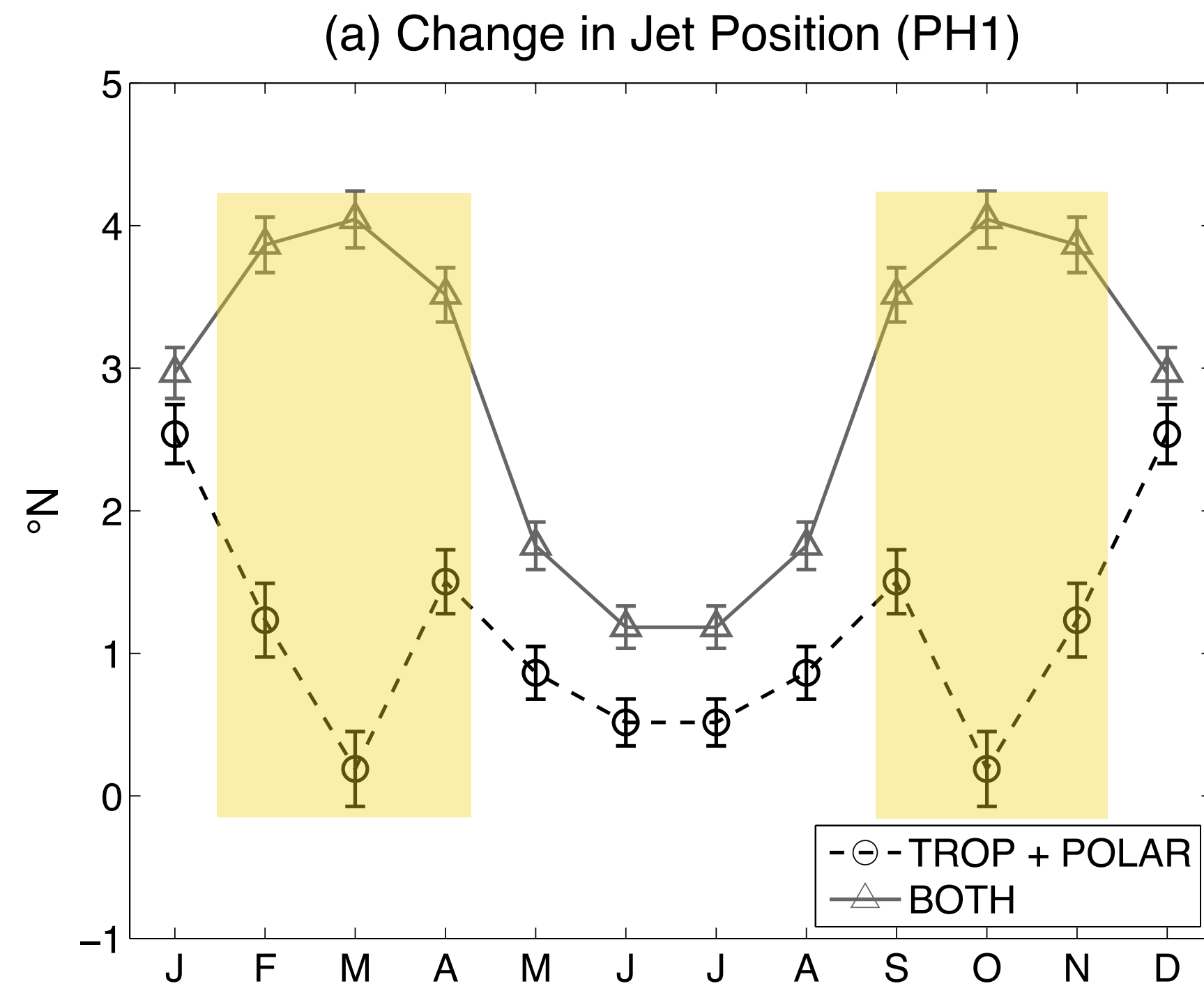
- shoulder seasons have the largest internal variability
- these seasons also exhibited the largest response



Dry core: Nonlinearity of response

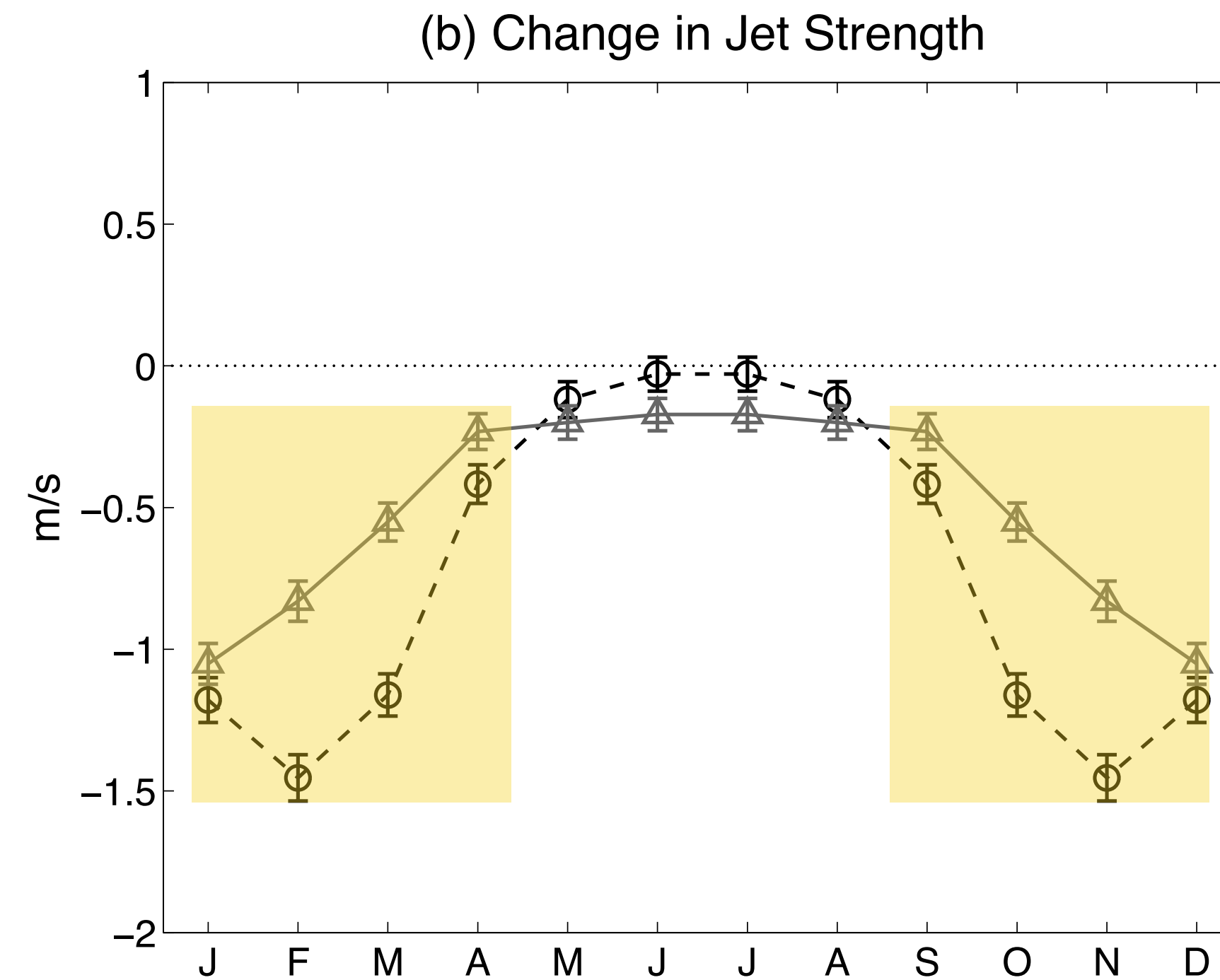
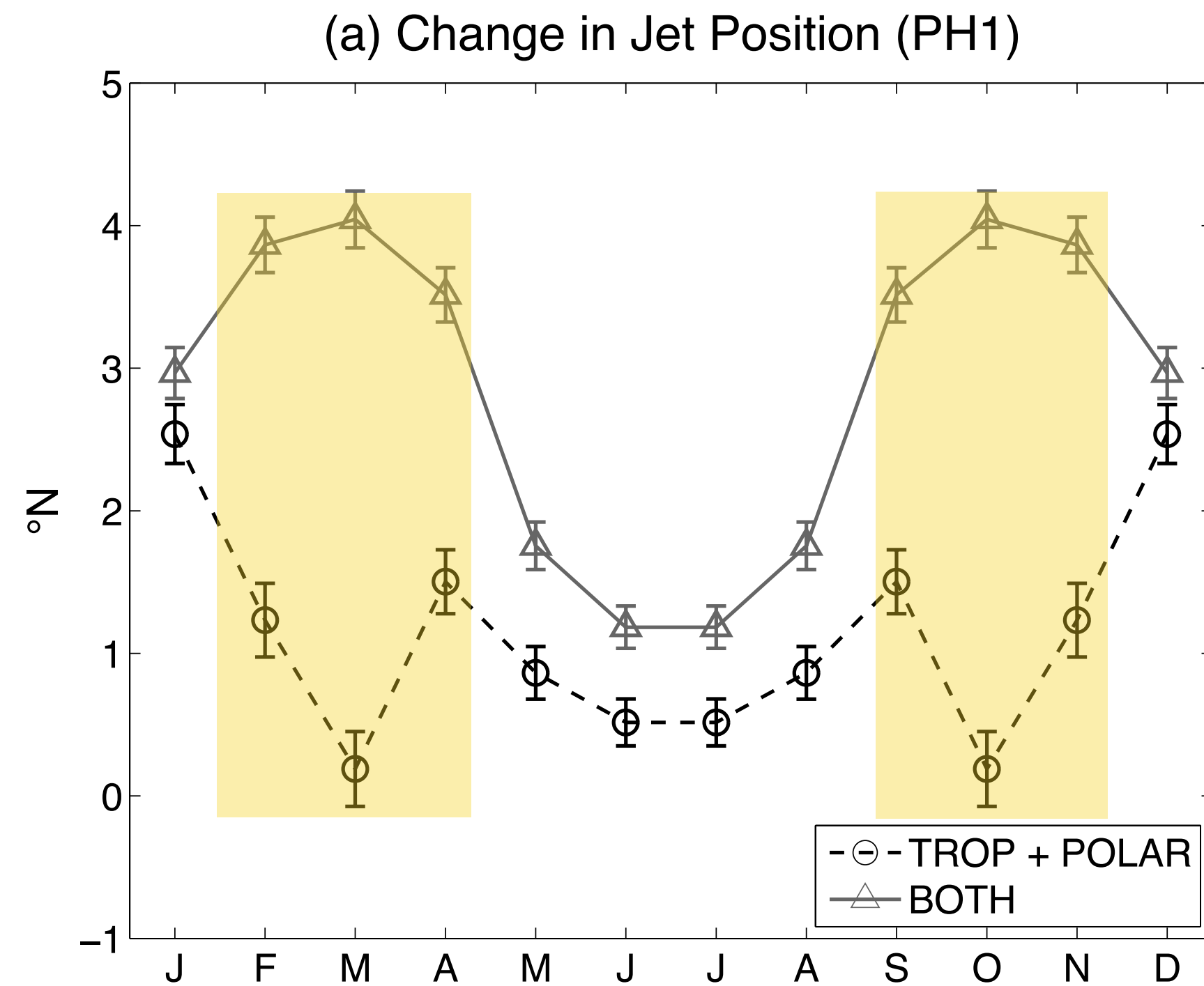


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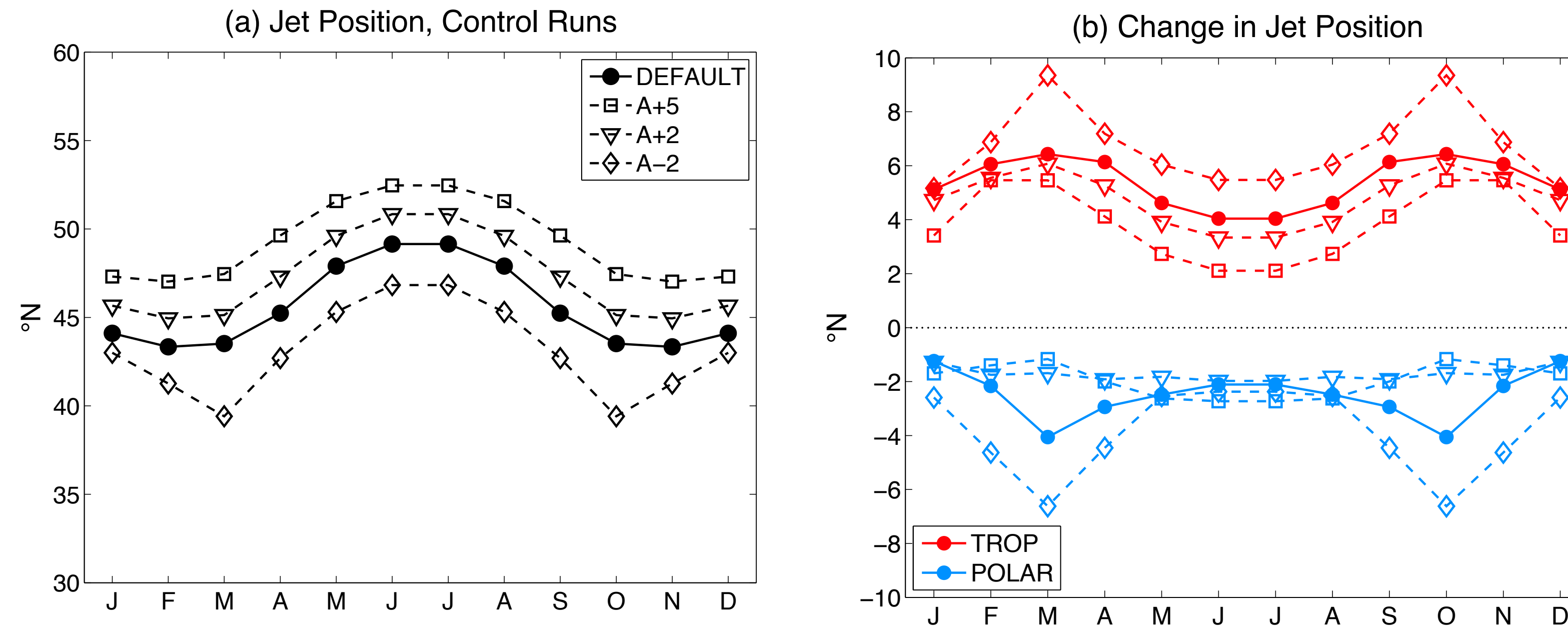
McGraw & Barnes (2016); in review

Dry core: Nonlinearity of response



Nonlinearity present in shoulder seasons:
POLAR matters less when both forcings are simulated at the same time

Dry core: initial jet position important in all seasons?

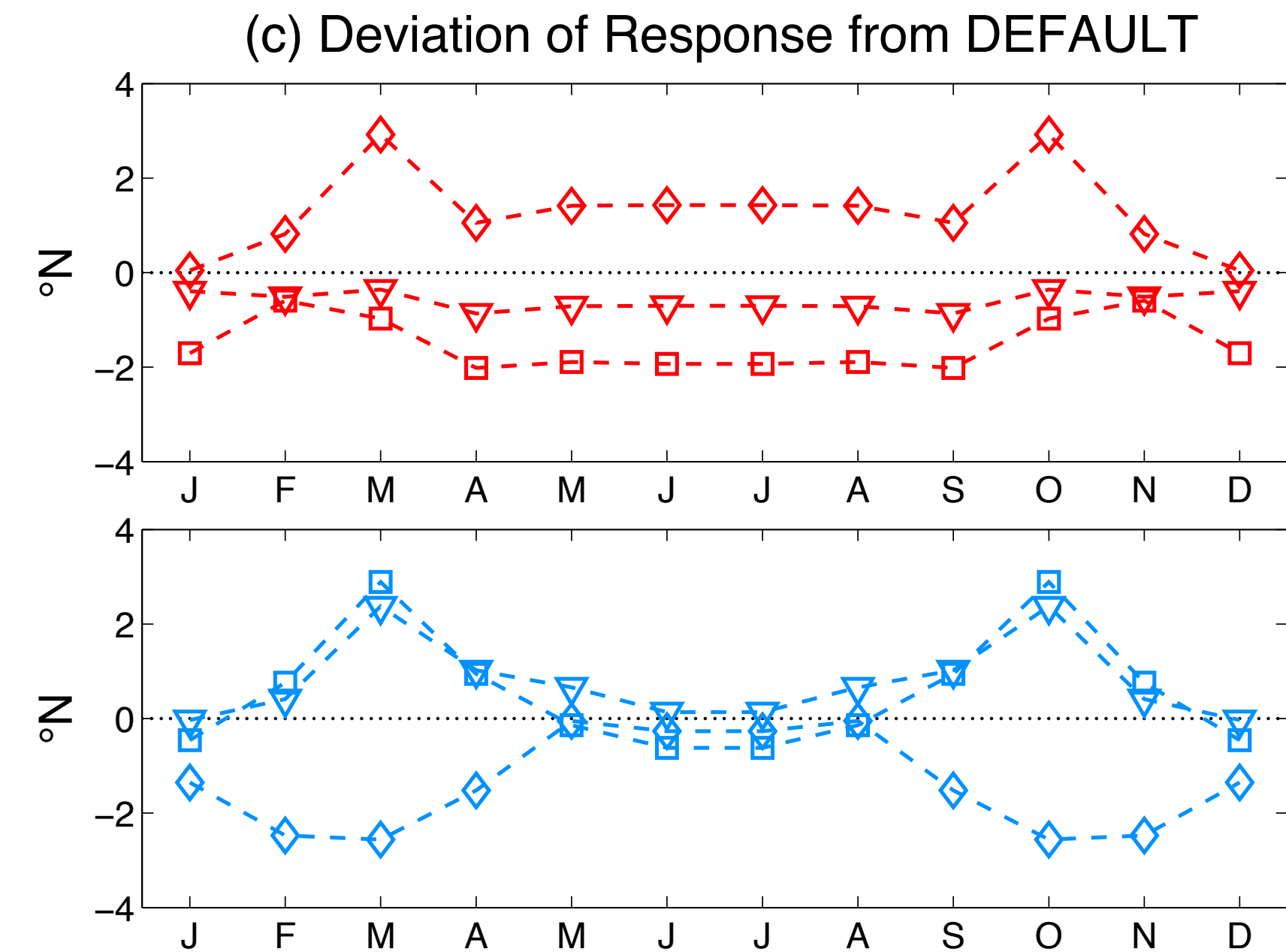
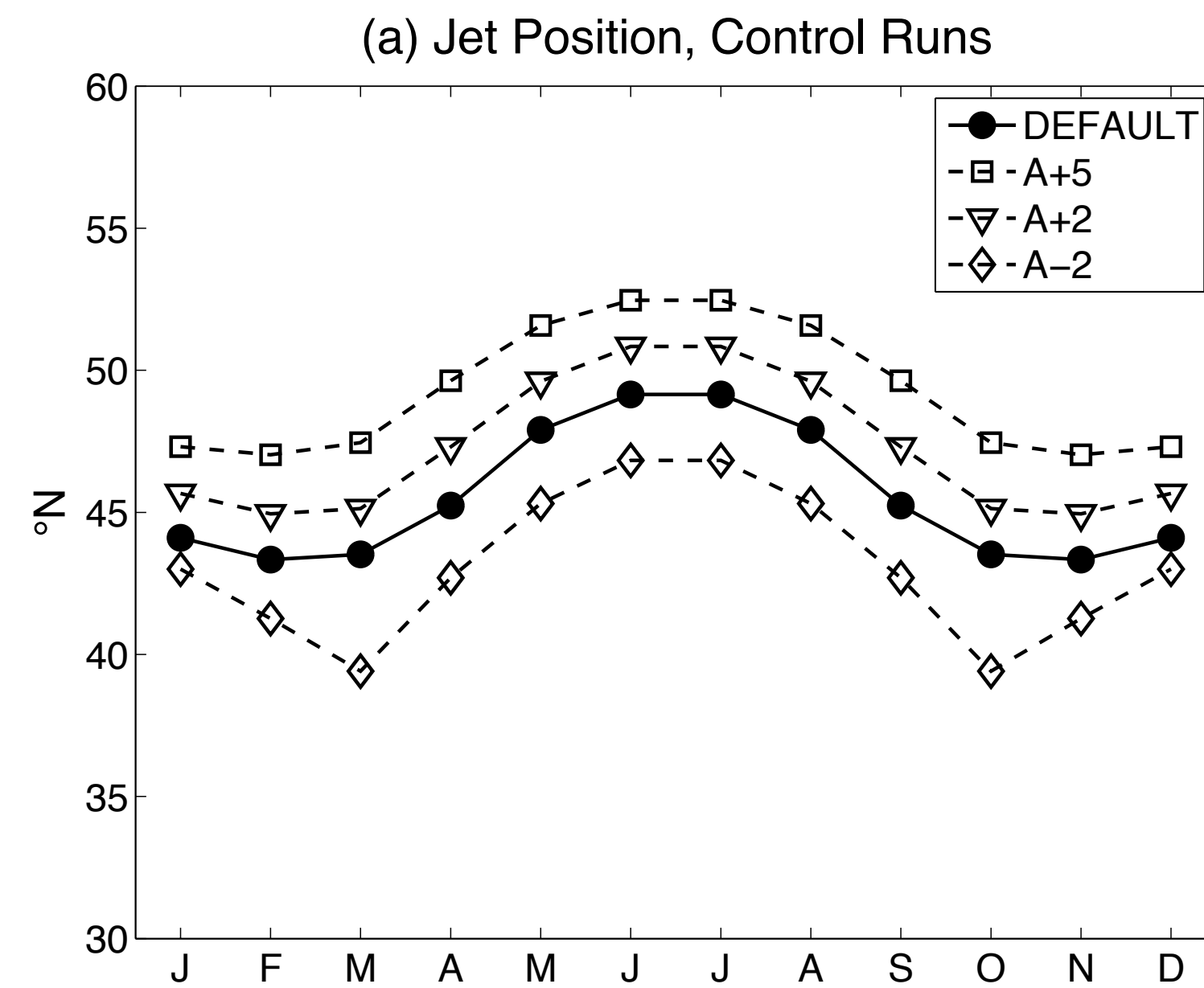


dependence on the basic state changes throughout the year:

- matters most in the shoulder seasons
- for POLAR warming, Dec/Jan & Jun/Jul are independent of basic state

*POLAR heating weaker and elevated off of the surface for these simulations
McGraw & Barnes (2016); in review*

Dry core: initial jet position important in all seasons?



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- for POLAR warming, Dec/Jan & Jun/Jul are independent of basic state

*POLAR heating weaker and elevated off of the surface for these simulations
McGraw & Barnes (2016); in review*

Seasonality of SH response to GHG

- forced GCM with increased GHGs
- found largest SH response in the zonal winds in summer/fall

Southern Hemisphere Atmospheric Circulation Response to Global Warming

PAUL J. KUSHNER, ISAAC M. HELD, AND THOMAS L. DELWORTH

NOAA/Geophysical Fluid Dynamics Laboratory, Princeton, New Jersey

(Manuscript received 24 February 2000, in final form 1 August 2000)

ABSTRACT

The response of the Southern Hemisphere (SH), extratropical, atmospheric general circulation to transient, anthropogenic, greenhouse warming is investigated in a coupled climate model. The extratropical circulation response consists of a SH summer half-year poleward shift of the westerly jet and a year-round positive wind anomaly in the stratosphere and the tropical upper troposphere. Along with the poleward shift of the jet, there is a poleward shift of several related fields, including the belt of eddy momentum-flux convergence and the mean meridional overturning in the atmosphere and in the ocean. The tropospheric wind response projects strongly onto the model's "Southern Annular Mode" (also known as the "Antarctic oscillation"), which is the leading pattern of variability of the extratropical zonal winds.

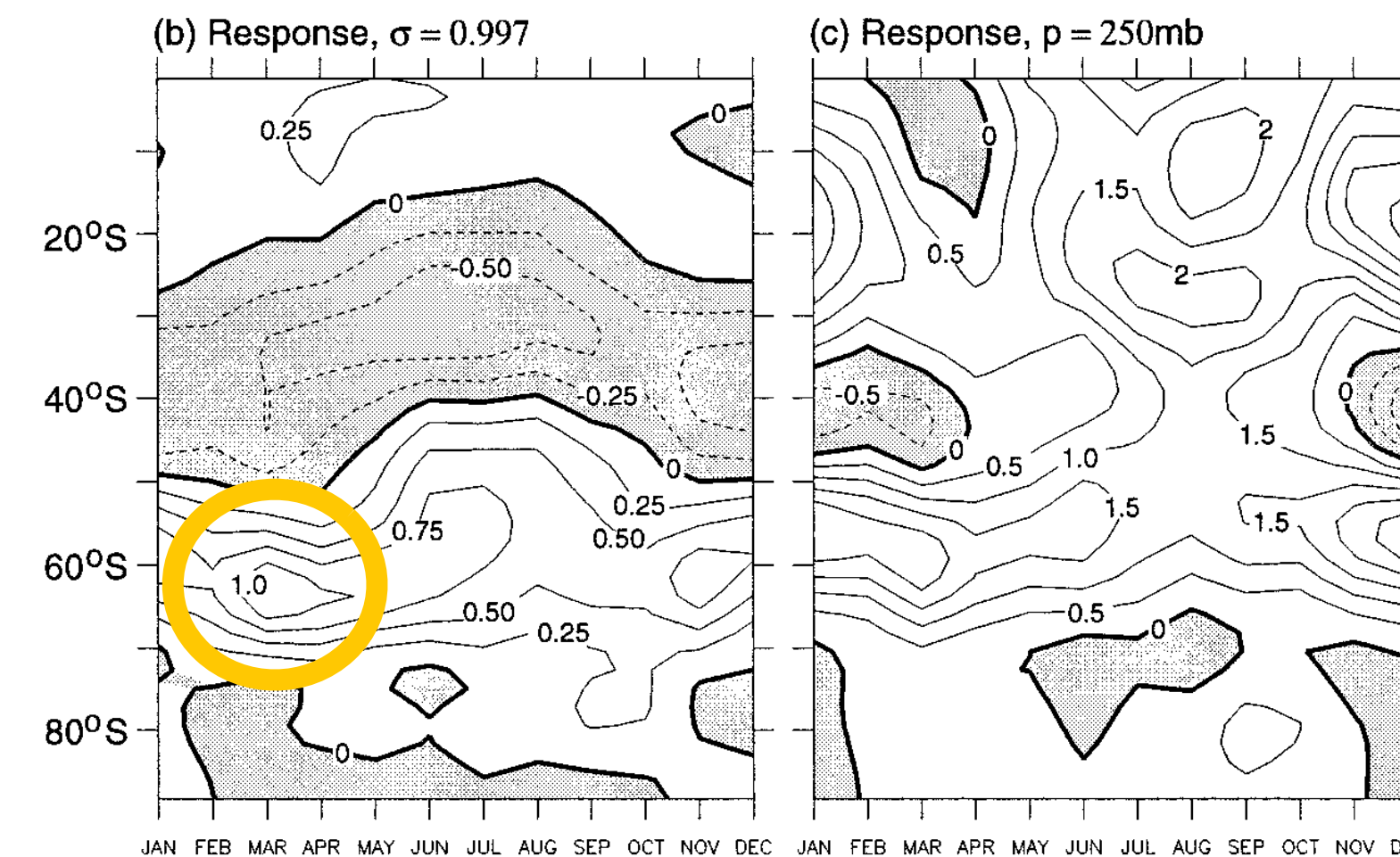


FIG. 4. The seasonal cycle of the climatological surface zonal-mean zonal wind for (a) the 800-yr time mean of the control integration, and (b) the ensemble mean response, years 2065–89. (c), (d) As in (a) and (b), but at 250 mb. Shading and dashed contours indicate negative values. Contour interval: (a) 2 m s^{-1} ; (b) 0.25 m s^{-1} ; (c) 5 m s^{-1} ; (d) 0.5 m s^{-1} .

Seasonality of SH response to GHG

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We will investigate the seasonal sensitivity of the circulation under constant forcing

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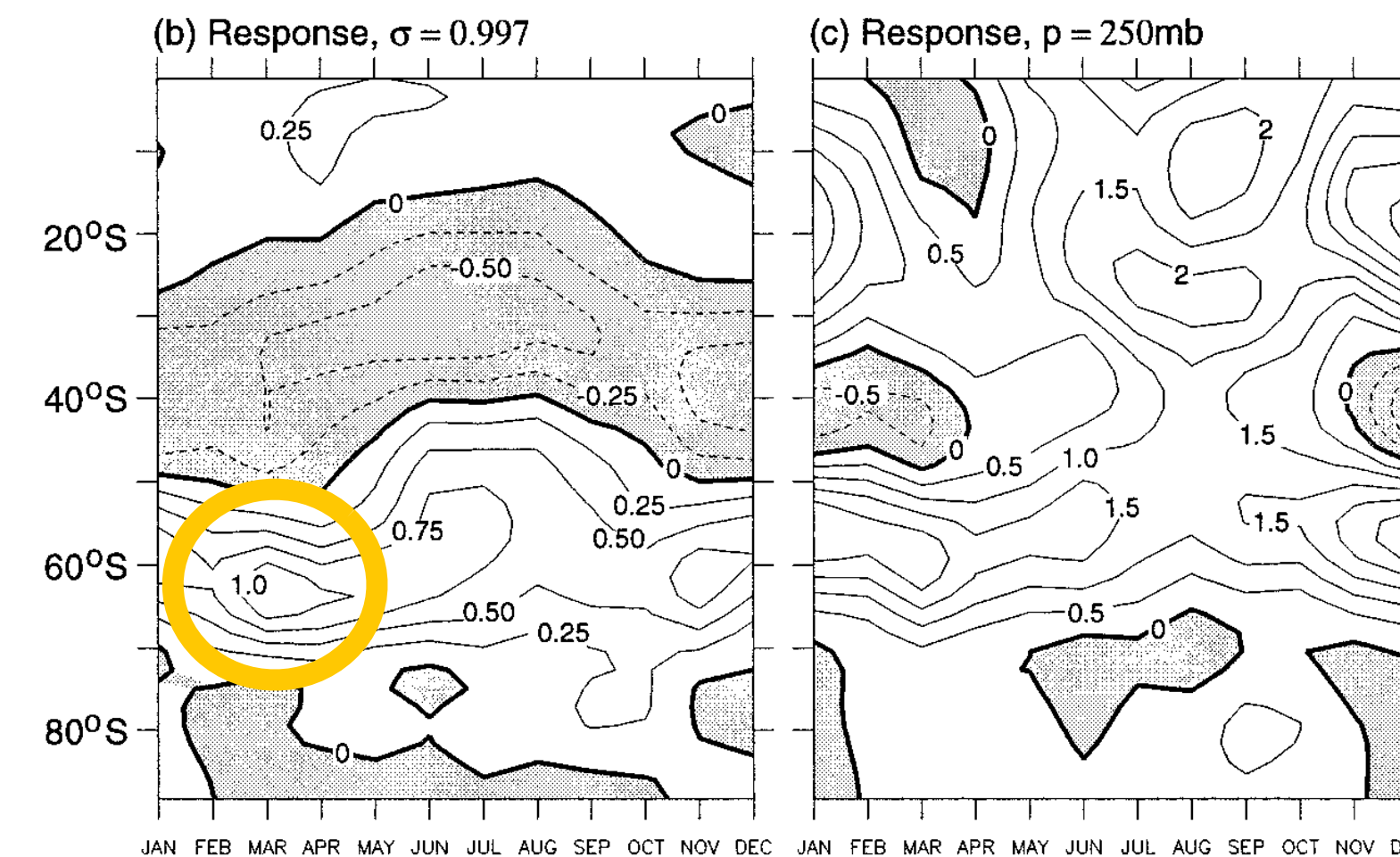
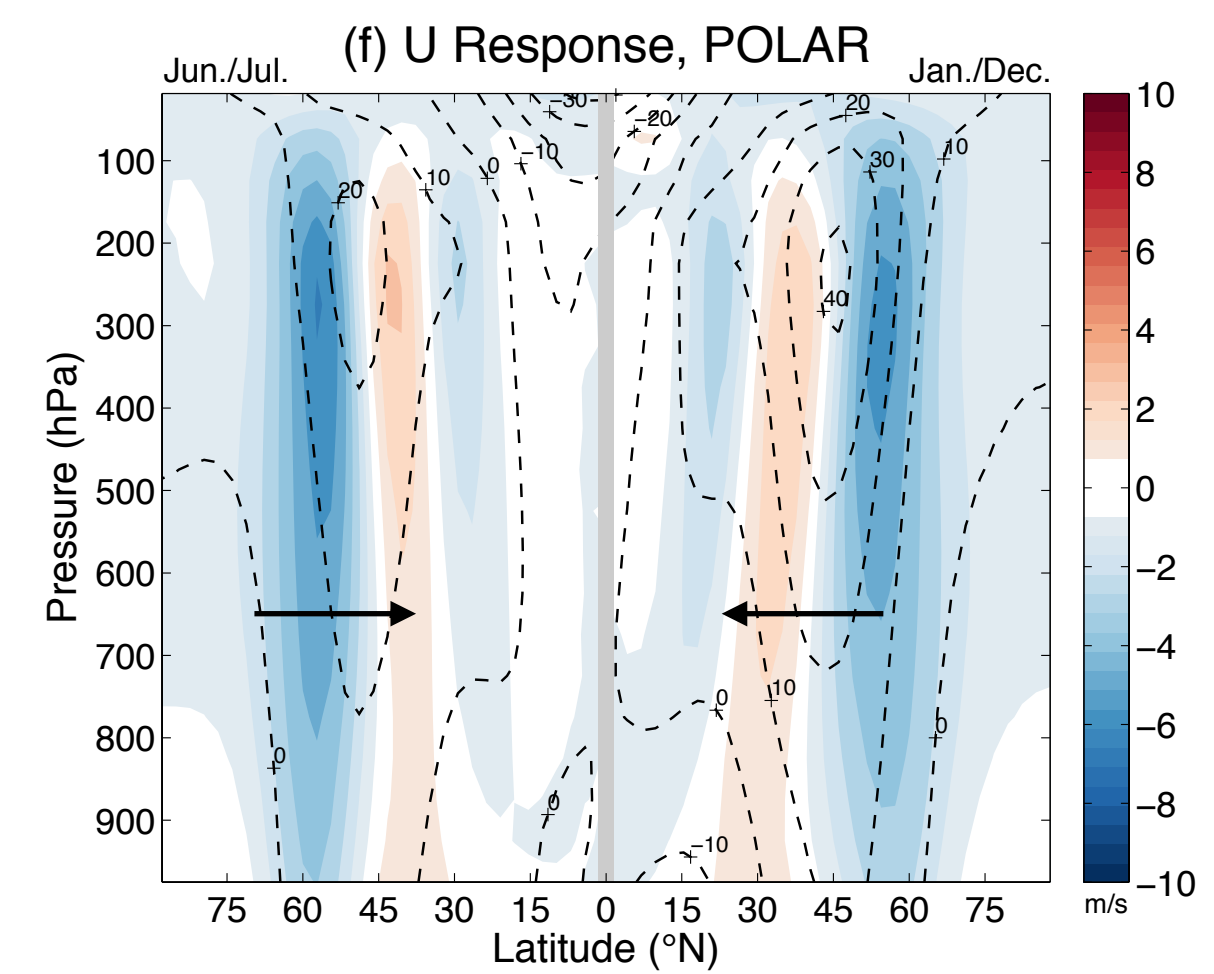
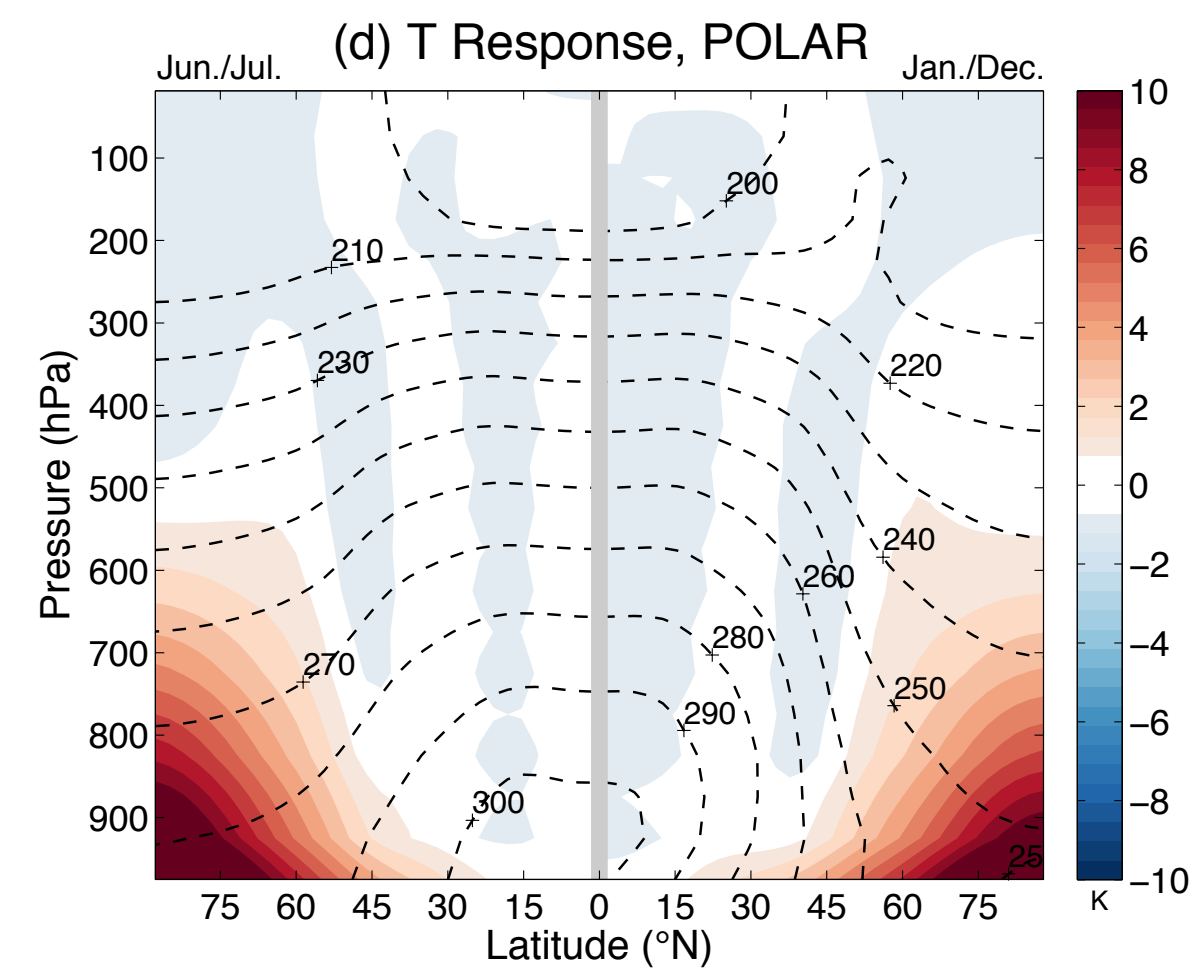
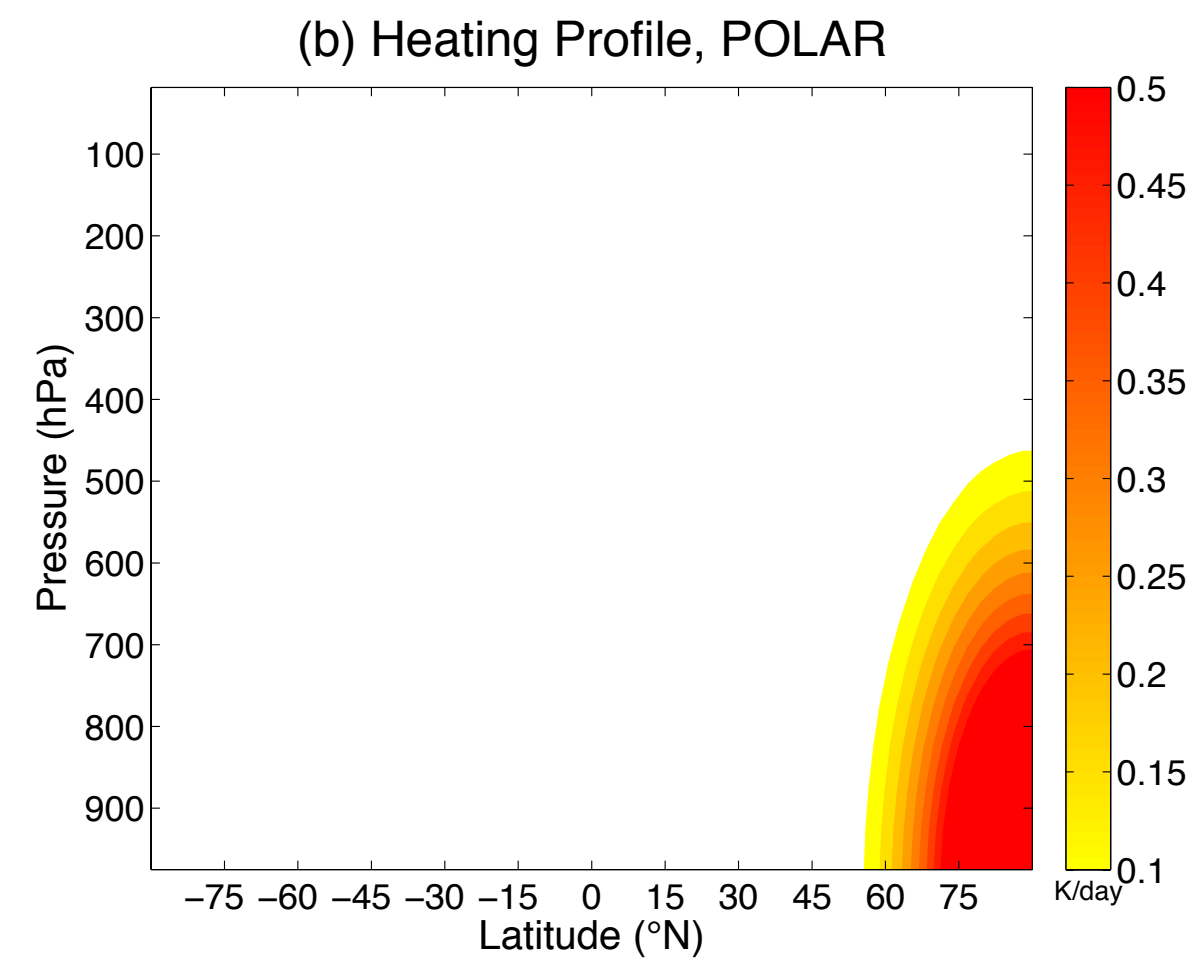
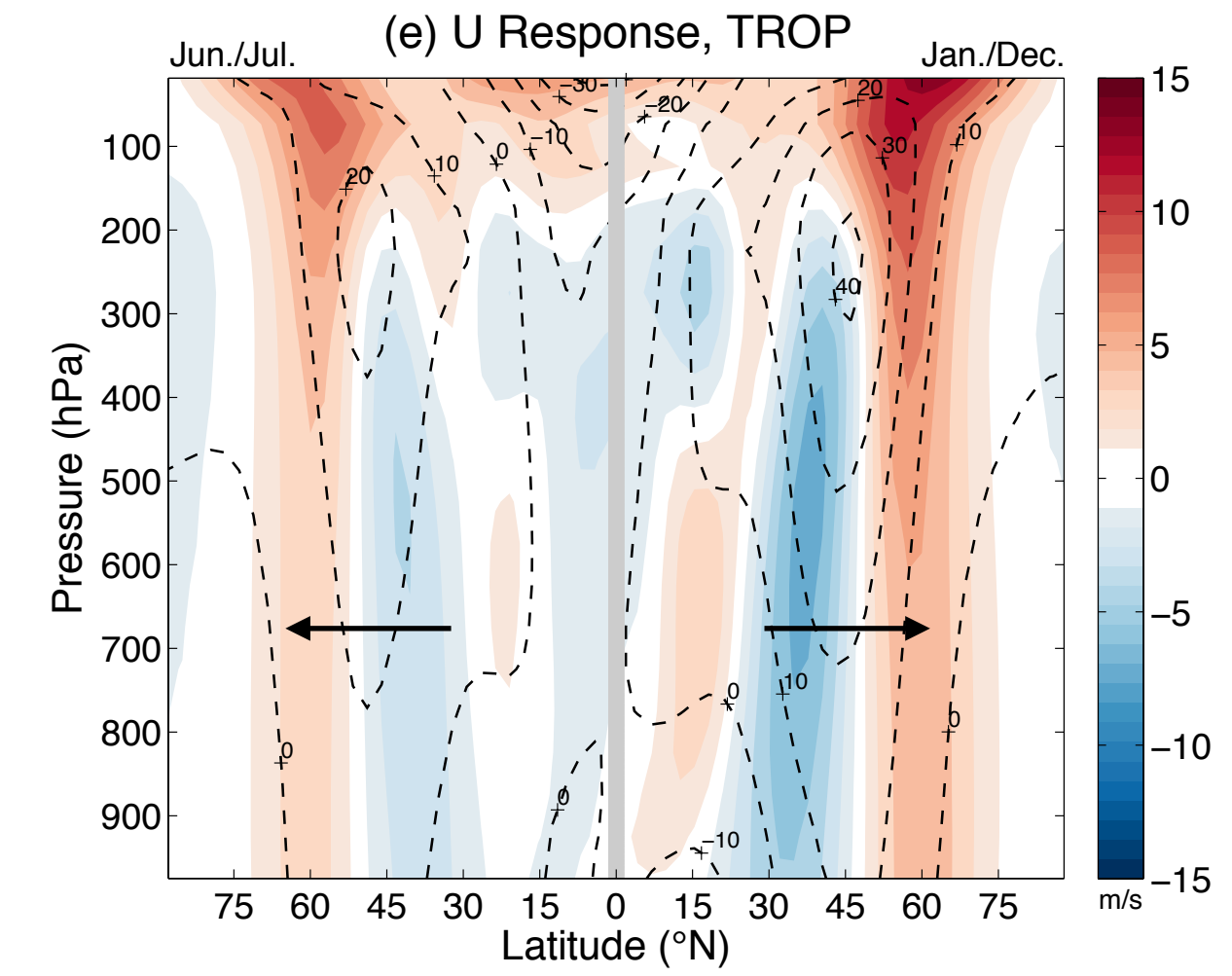
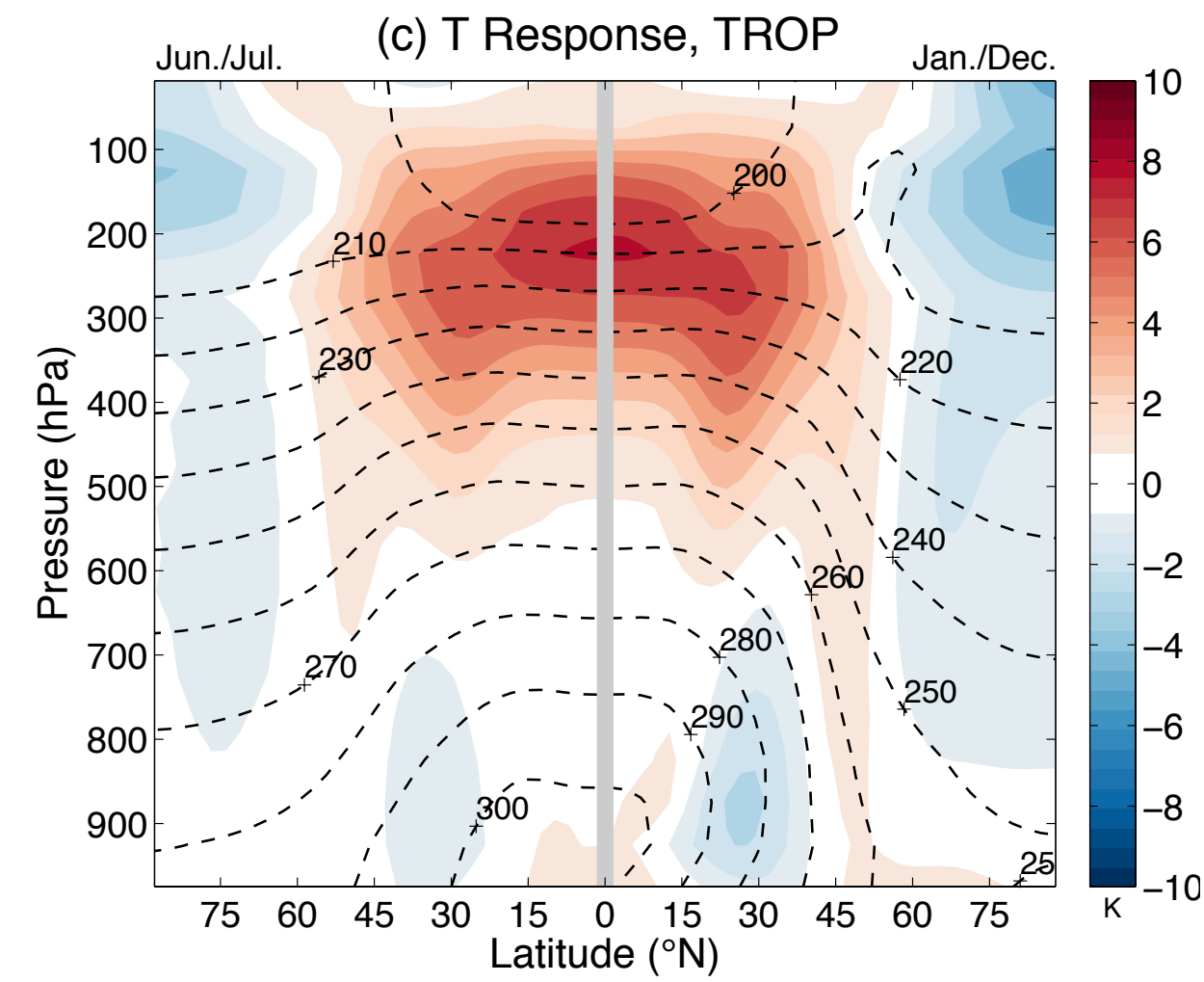
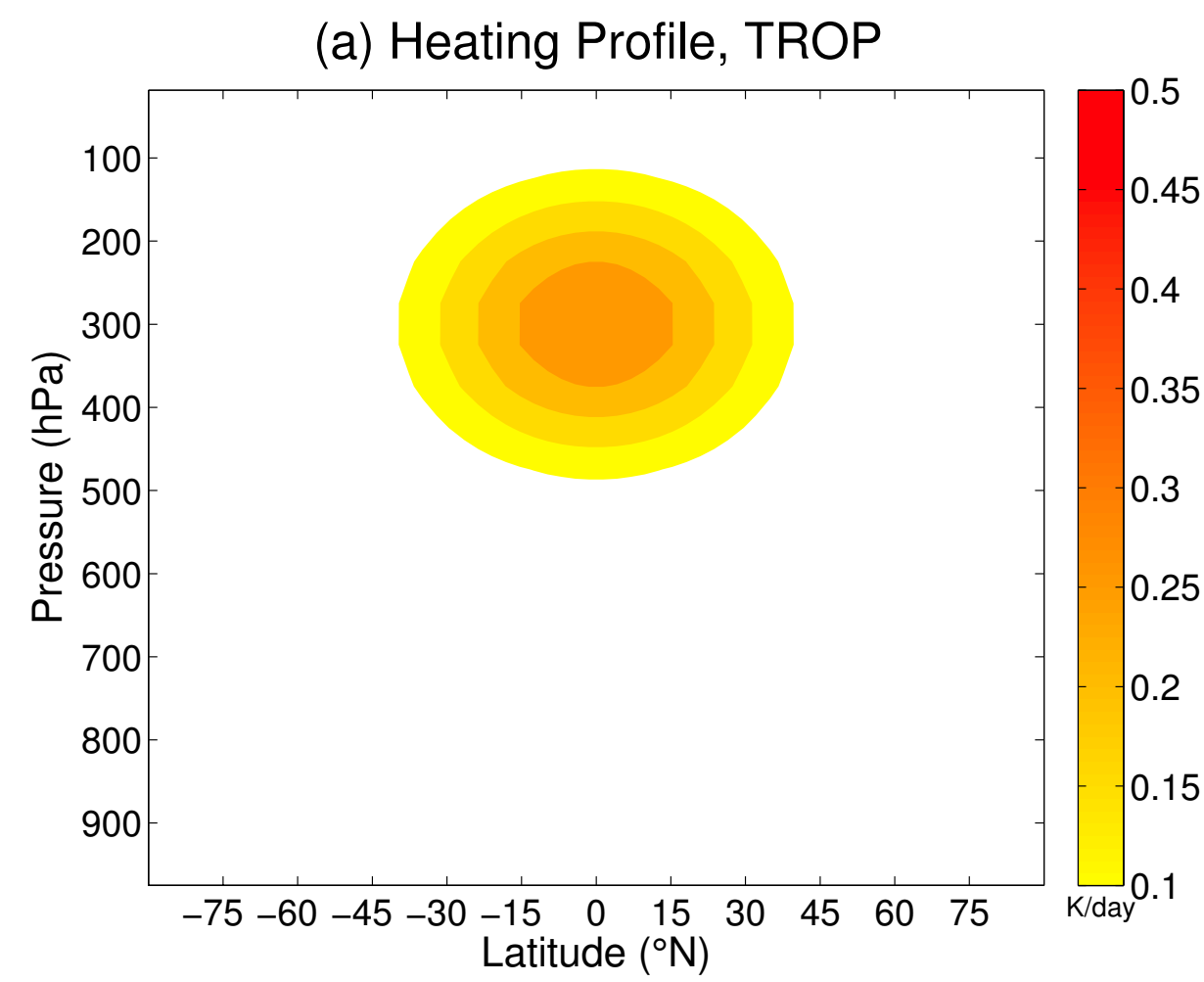


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TROP heating



McGraw & Barnes (2016); in press