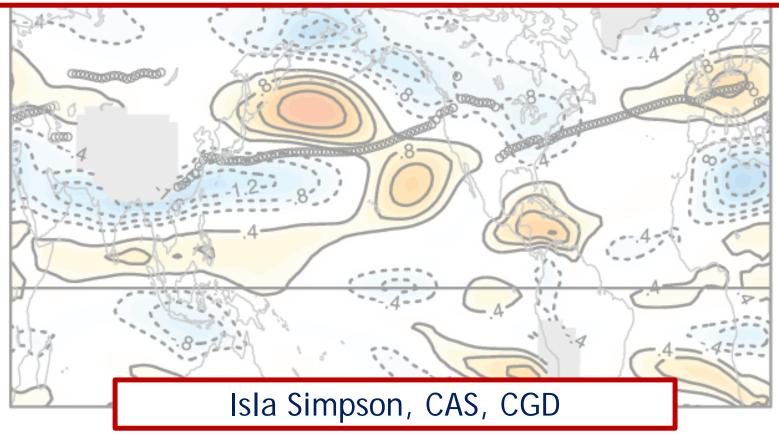
Model Diversity in Future Predicted Pacific Circulation Change







Two circulation changes that are commonly considered to occur under global warming:

(1) A poleward shifting of the mid-latitude westerlies

(2) A weakening of the tropical divergent circulation

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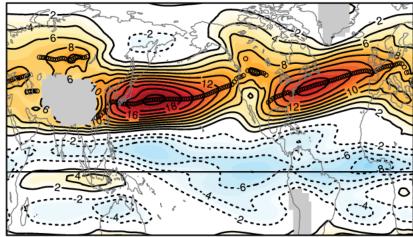
Focus: These circulation responses in the wintertime North Pacific (DJF) in the CMIP5 models.

(2070-2099) RCP8.5 - (1979-2005) historical

35 different models

(1) A poleward shifting of the mid-latitude westerlies

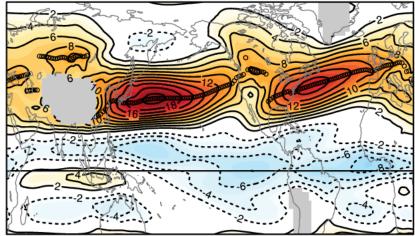




700hPa zonal wind Multi-model mean

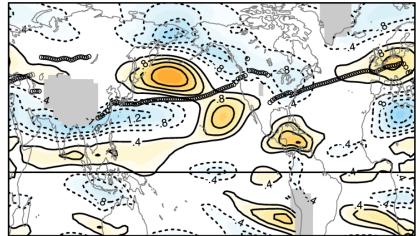
(1) A poleward shifting of the mid-latitude westerlies

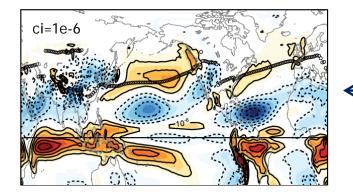




700hPa zonal wind Multi-model mean

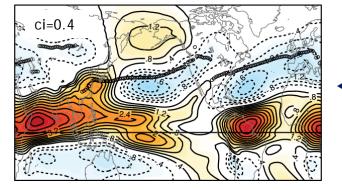




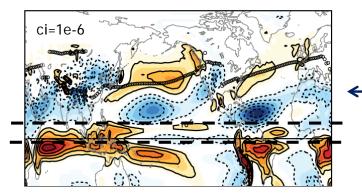


____250hPa Divergence

Past Climatology

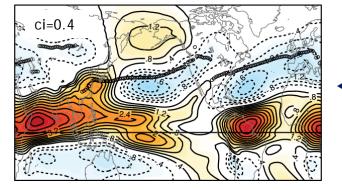


250hPa Meridional ←Component of the Divergent Flow

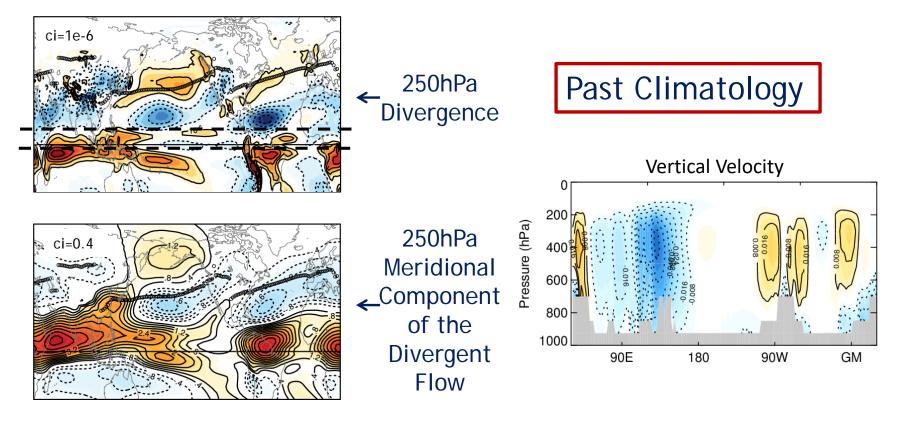


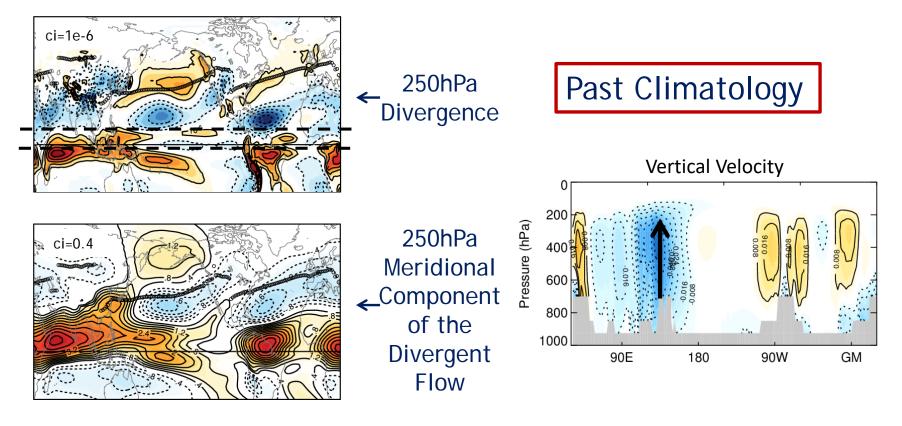
____250hPa Divergence

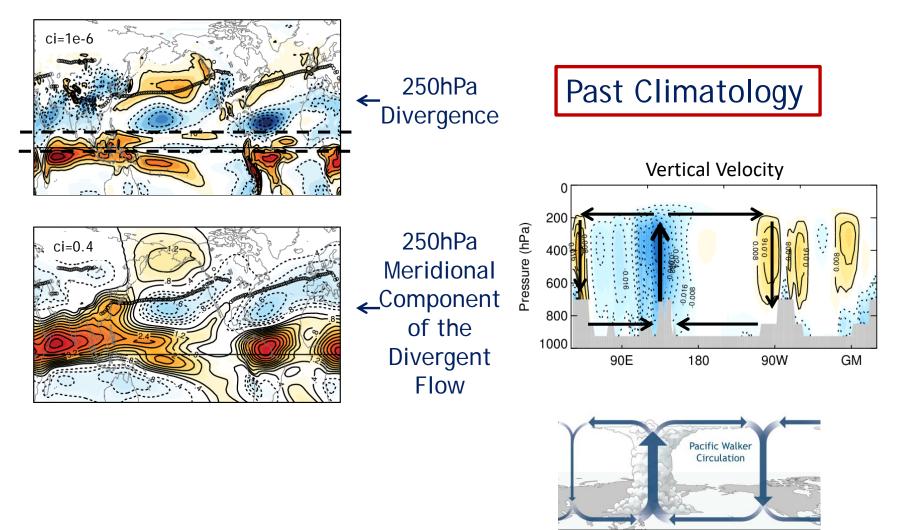
Past Climatology



250hPa Meridional ←Component of the Divergent Flow

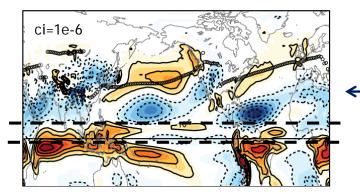






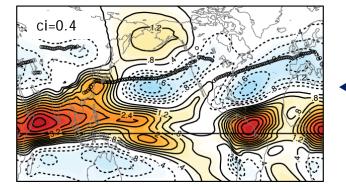
www.noaa..gov

equator



____250hPa Divergence

Past Climatology

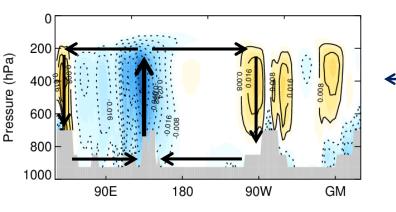


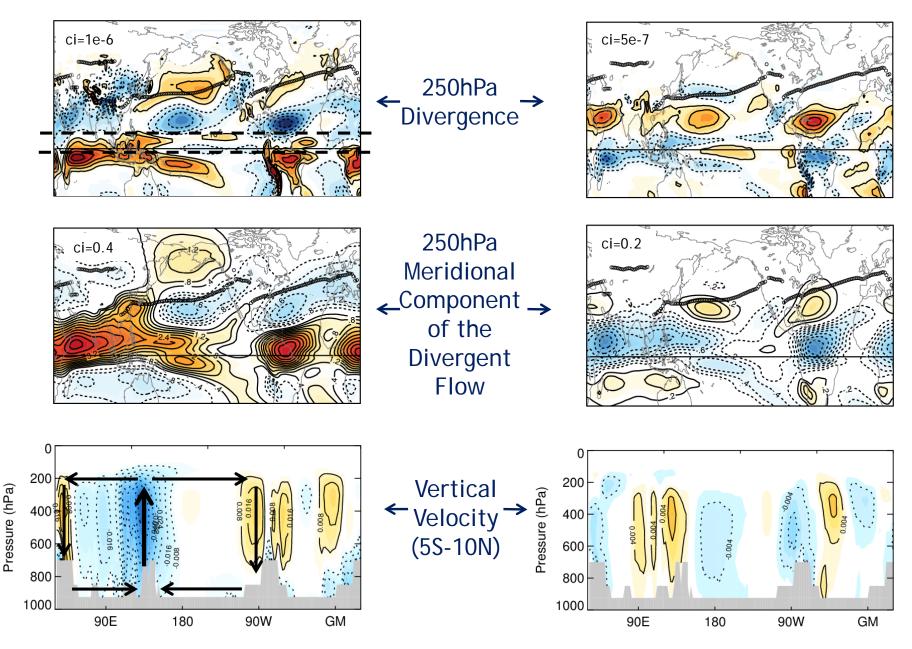
250hPa Meridional Component of the Divergent Flow

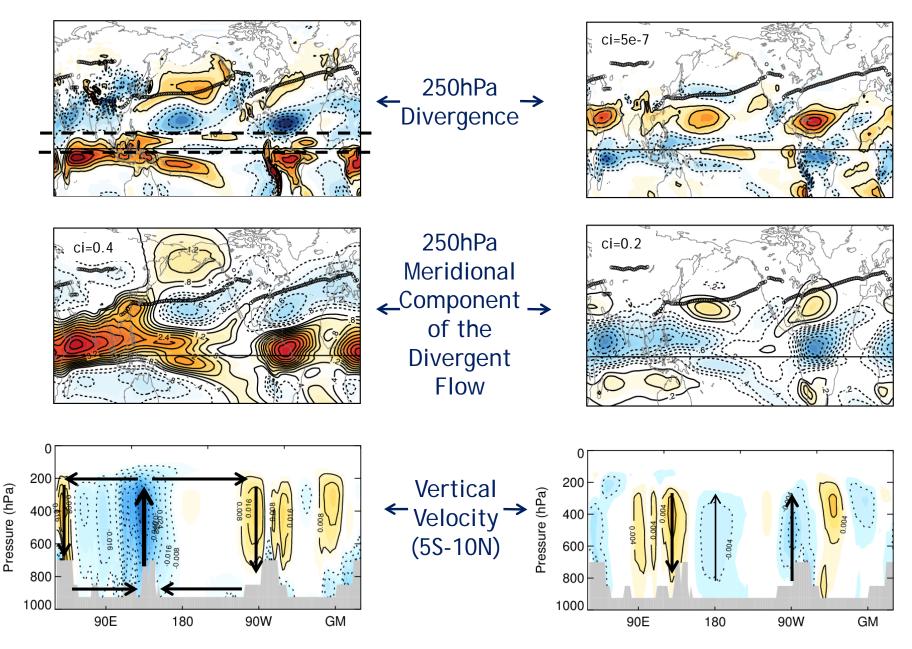
Vertical

Velocity

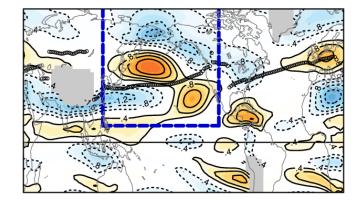
(5S-10N)



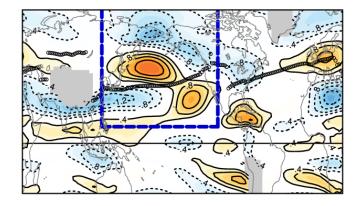




BUT, there is considerable spread among the models in both these aspects, and they are connected.

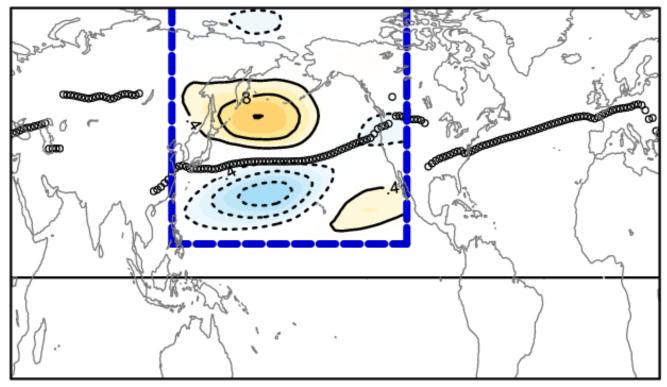


EOF analysis (across models) of zonal wind response



EOF analysis (across models) of zonal wind response

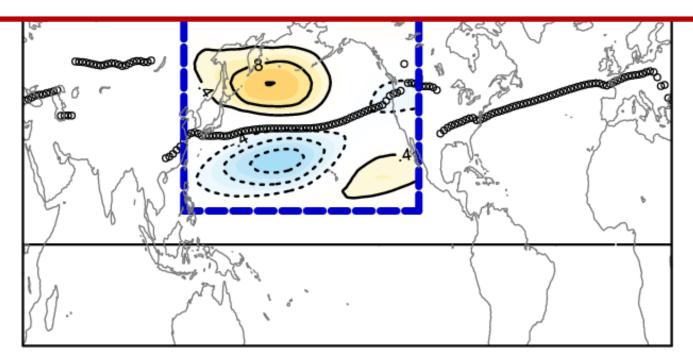
EOF1, 37.5% of variance explained



Use the EOF Index as a measure of a models behaviour in the extra-tropical west Pacific

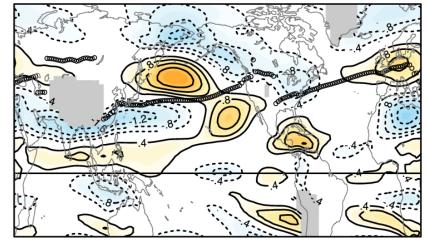
25% most positive EOF1 = Poleward Models

25% most negative EOF1 = Not Poleward Models (9 models in each group)



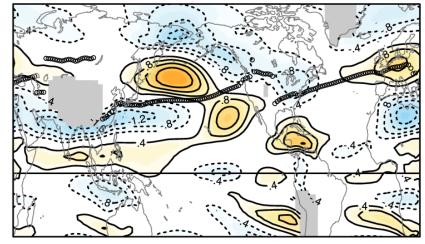
700hPa zonal wind, Future - Past difference

Multi-Model Mean

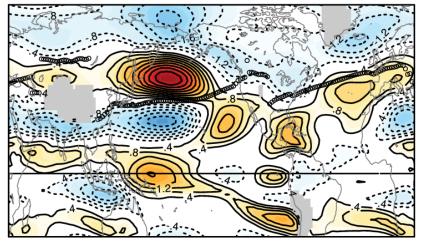


700hPa zonal wind, Future - Past difference

Multi-Model Mean

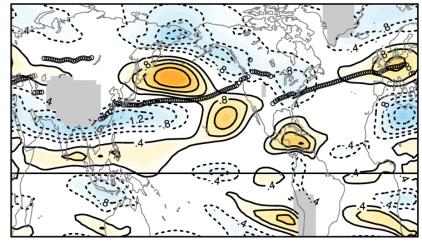


Poleward Models

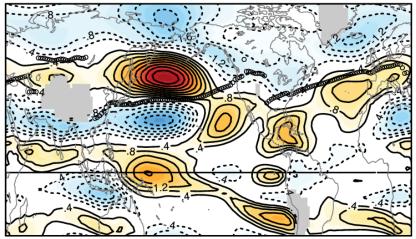


700hPa zonal wind, Future - Past difference

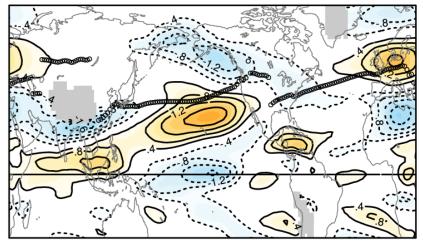
Multi-Model Mean



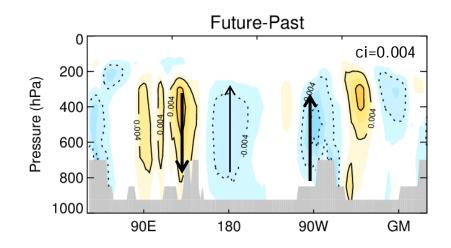
Poleward Models

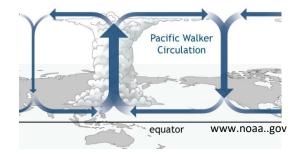


Not Poleward Models

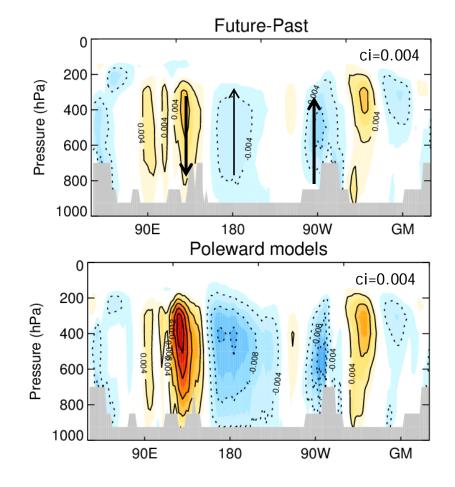


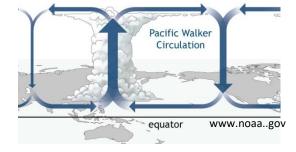
Walker Circulation



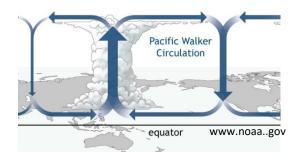


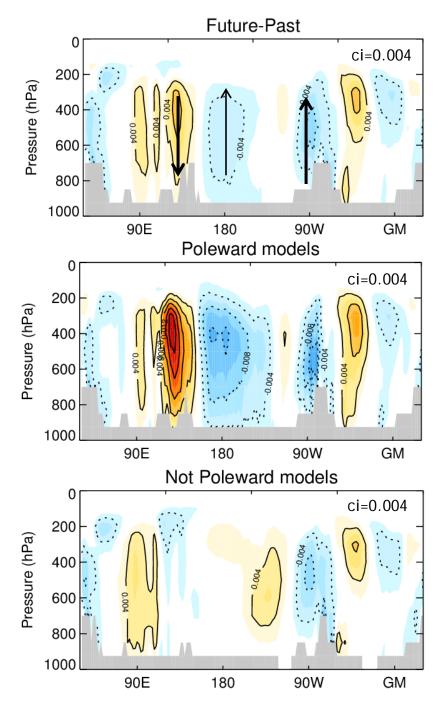
Walker Circulation

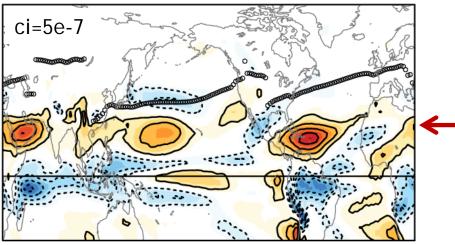




Walker Circulation

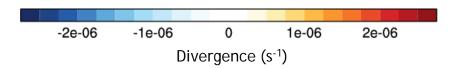


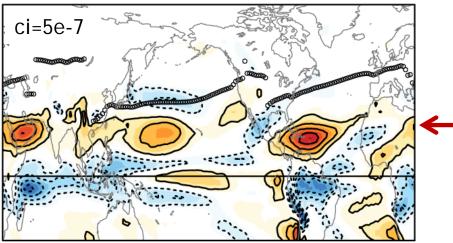




250hPa Divergence

Multi-Model Mean (Future-Past)

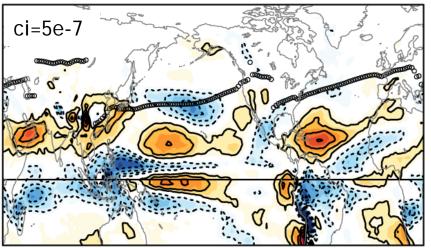


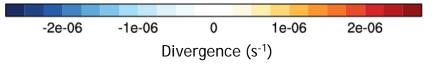


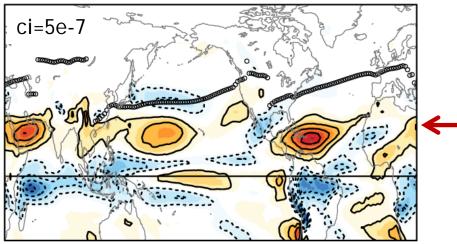
250hPa Divergence

Multi-Model Mean (Future-Past)

Poleward models



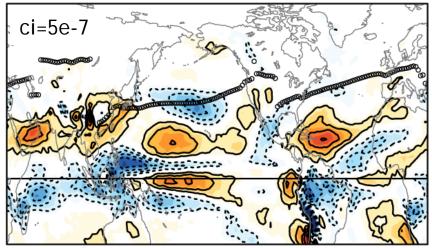




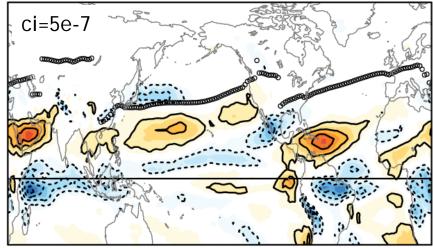
250hPa Divergence

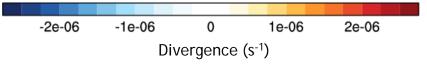
Multi-Model Mean (Future-Past)

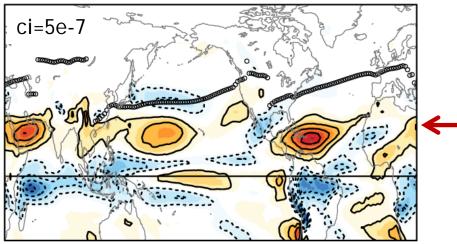
Poleward models



Not Poleward models



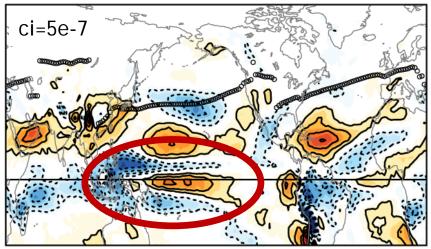




250hPa Divergence

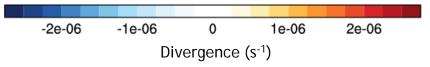
Multi-Model Mean (Future-Past)

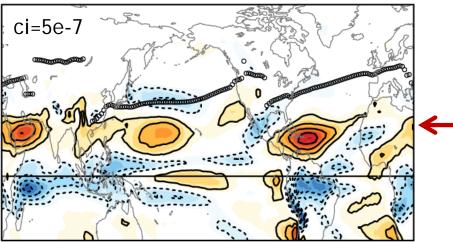
Poleward models



ci=5e-7

Not Poleward models

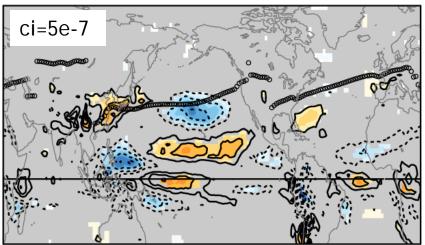




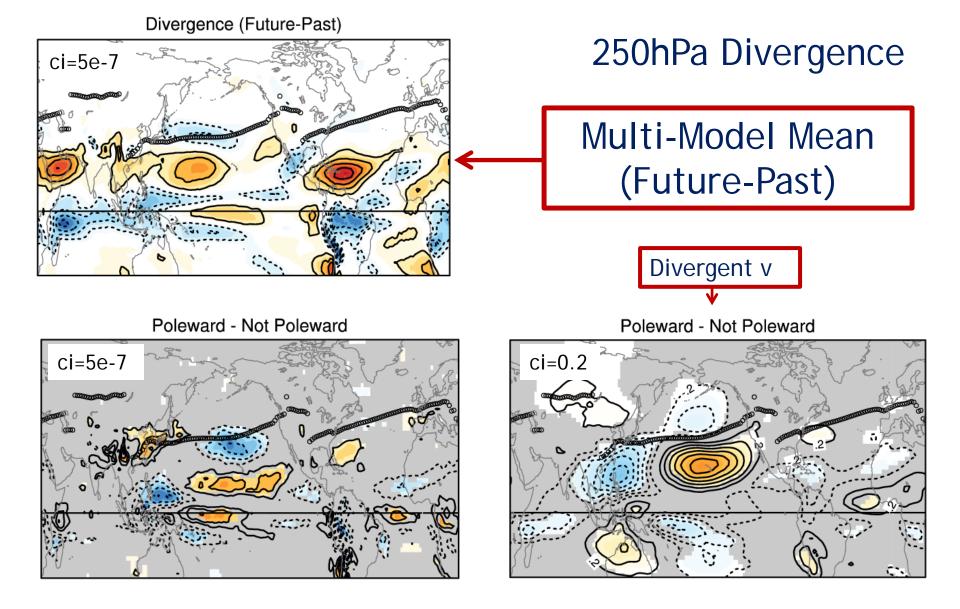
250hPa Divergence

Multi-Model Mean (Future-Past)

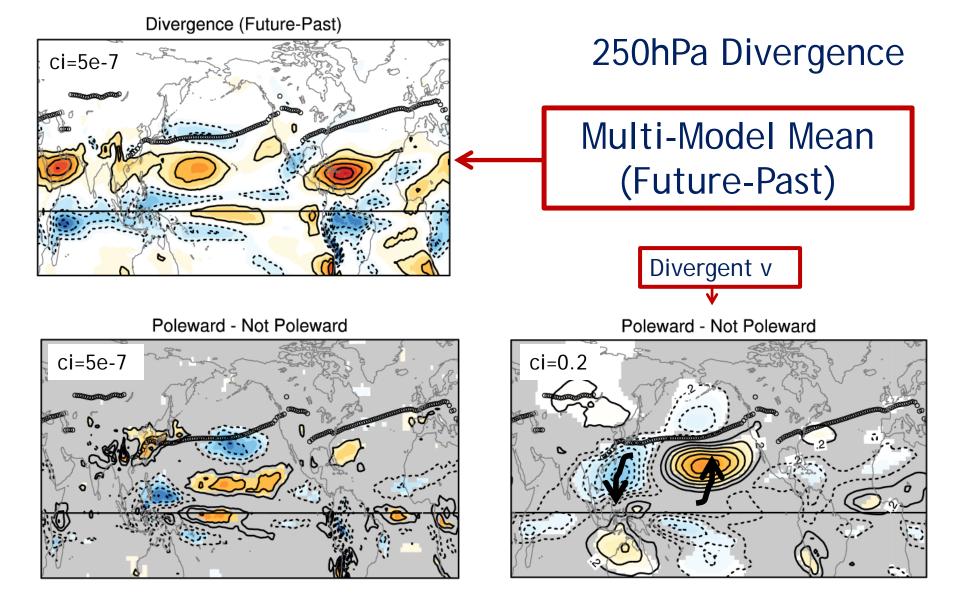
Poleward - Not Poleward



Grey = Not statistically significant



Grey = Not statistically significant

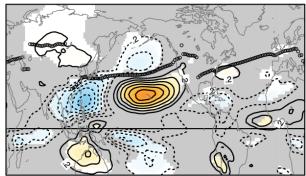


Grey = Not statistically significant

There is a very strong connection between the poleward shifting of the westerlies in the extra-tropics and the weakening of the divergent circulation in the tropics.

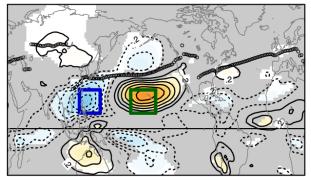
Divergent Meridional Wind Response

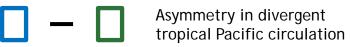
Poleward - Not Poleward



Divergent Meridional Wind Response

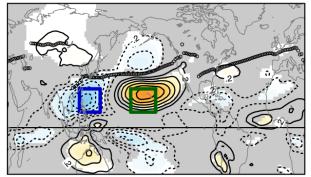
Poleward - Not Poleward





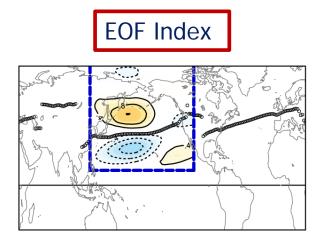
Divergent Meridional Wind Response

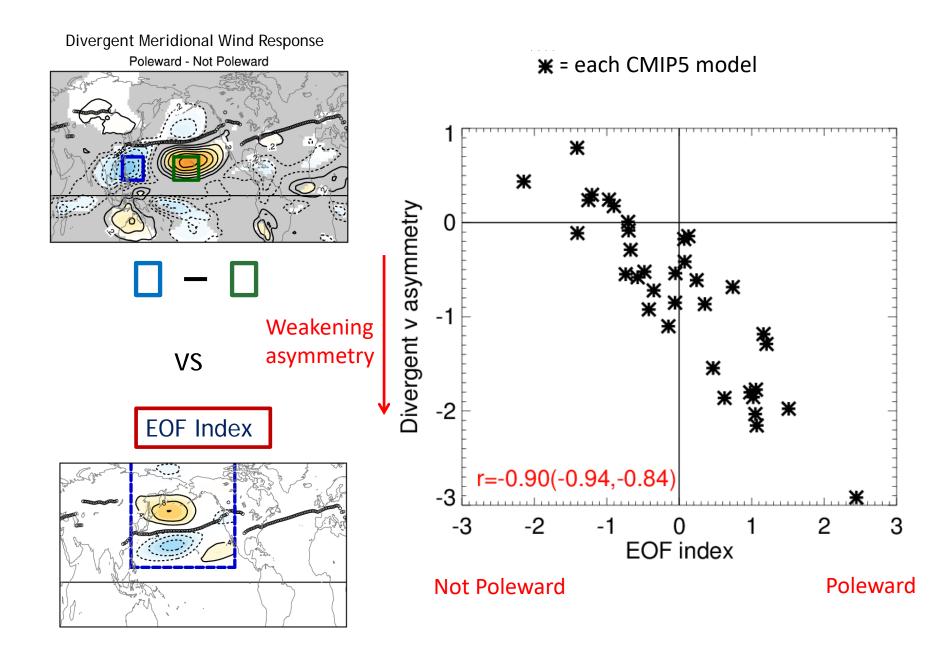
Poleward - Not Poleward



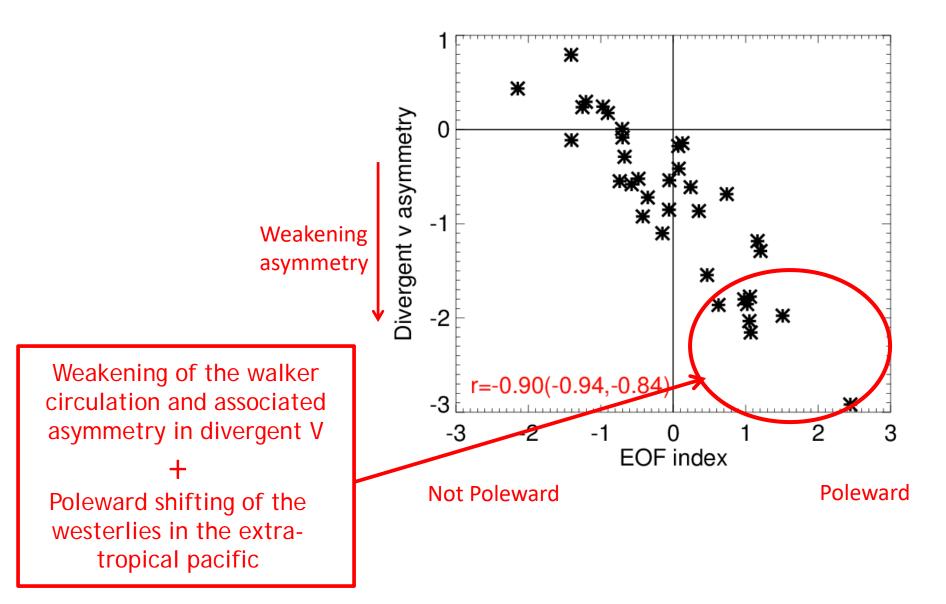
- Asymmetry in divergent tropical Pacific circulation

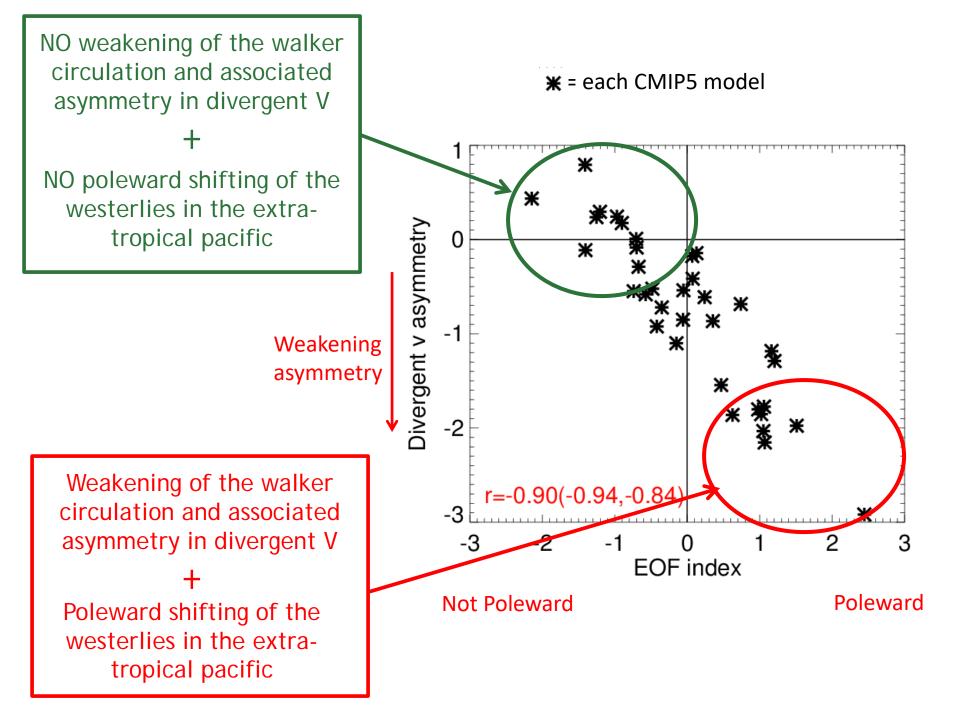
VS

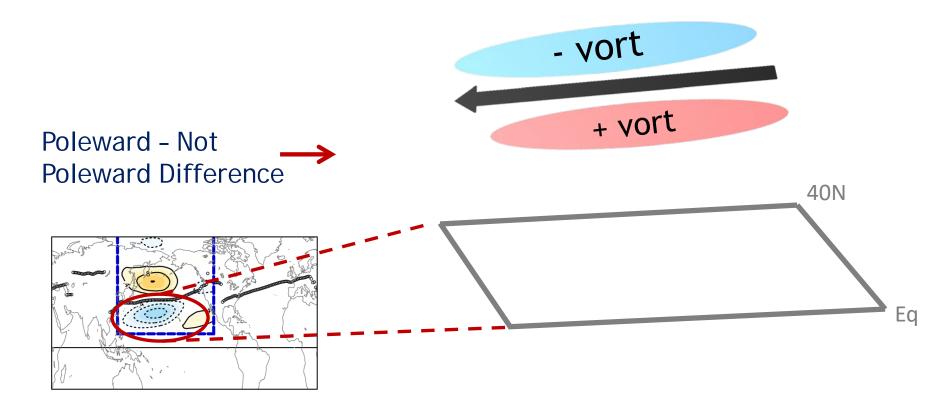


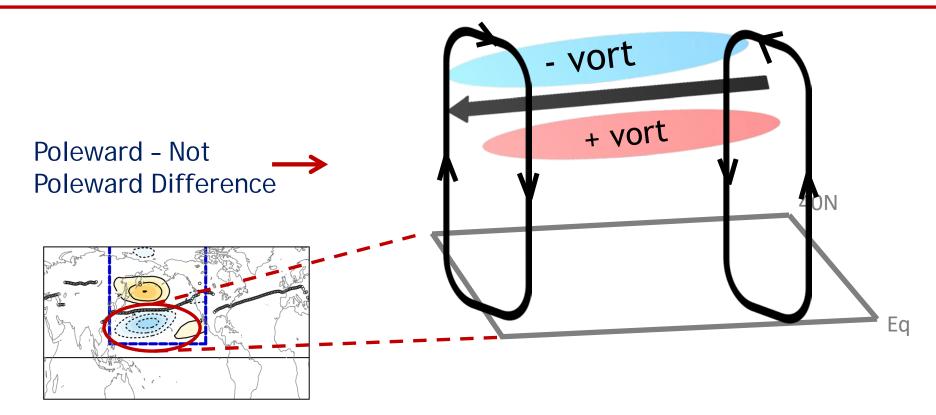


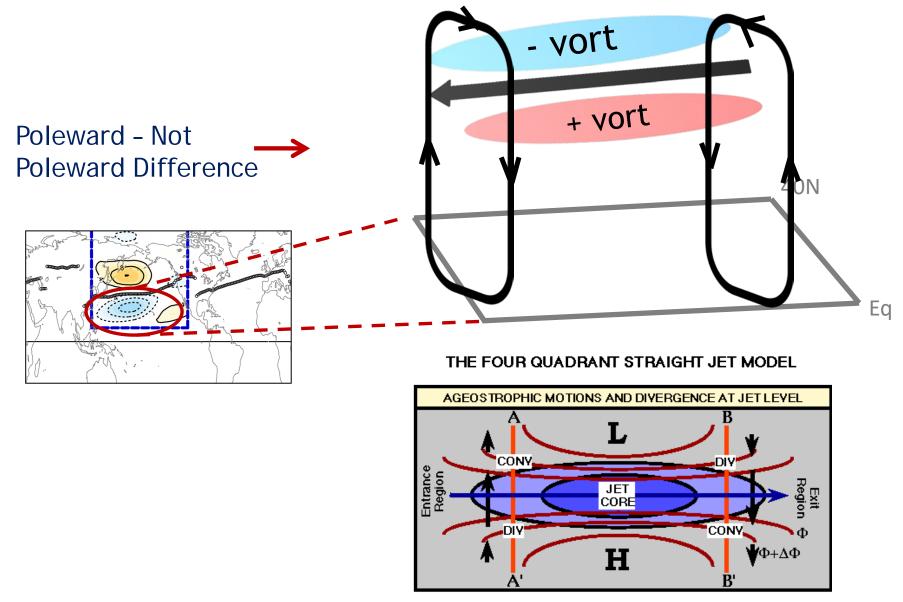
***** = each CMIP5 model



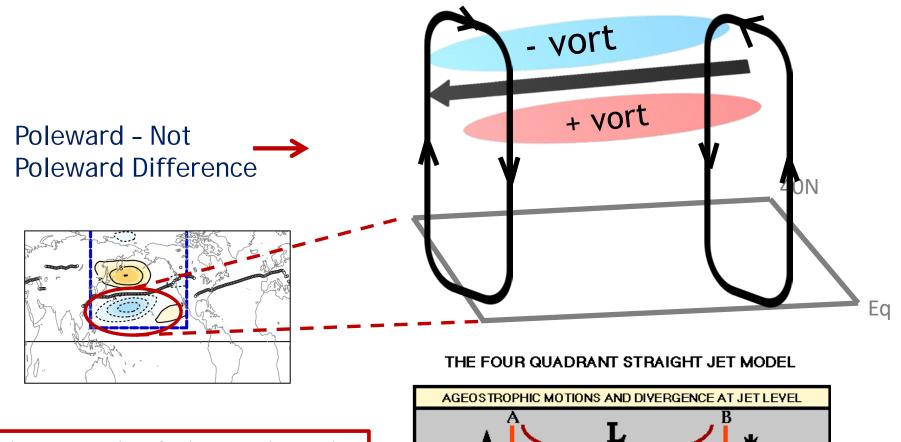




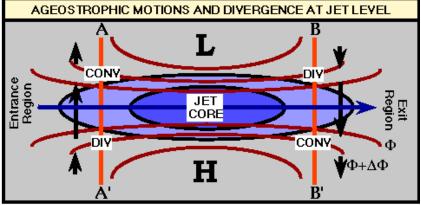




http://www4.ncsu.edu/~nwsfo/storage/training/jets/straightjet.html

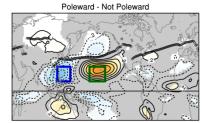


Divergent circulations at jet exit and entrance are predicted by quasi-geostrophic theory (omega equation) (Hoskins et al 1978)

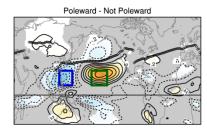


http://www4.ncsu.edu/~nwsfo/storage/training/jets/straightjet.html

Relationship between Future-Past difference and Past Climatologies



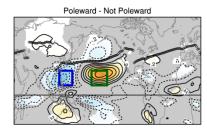
0 - 0



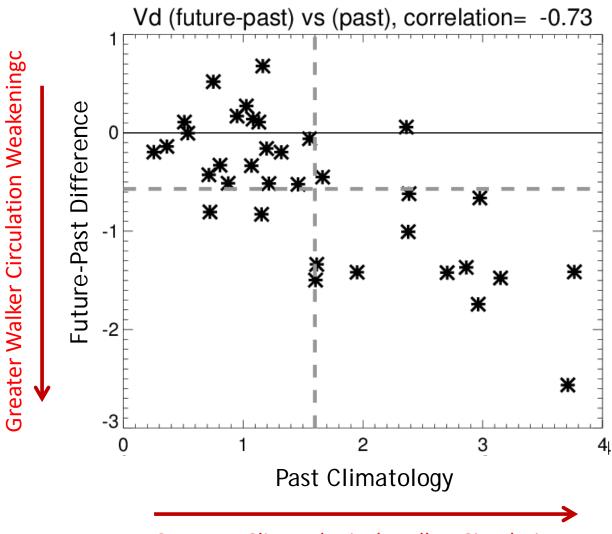
- 🛛

Vd (future-past) vs (past), correlation= -0.73 Greater Walker Circulation Weakeningc ж ж Future-Past Difference ¥ 0 ₩ ж ж ж ж ж ж -1 ** жЖ ж ж ж ж -2 ж -3 2 0 3 1 4 Past Climatology

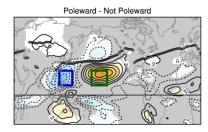
Stronger Climatological Walker Circulation



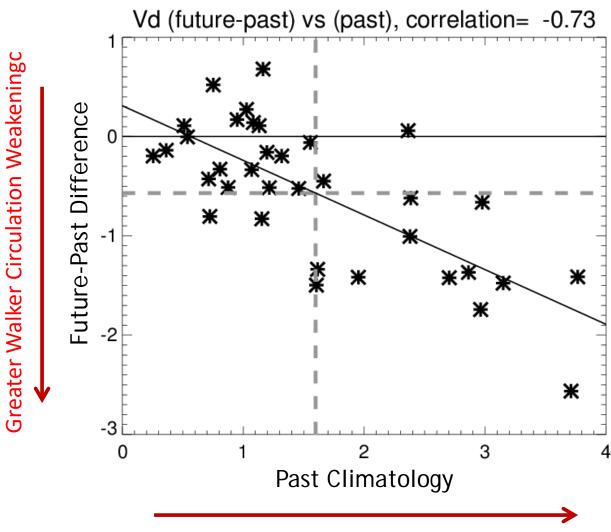
- 🛛



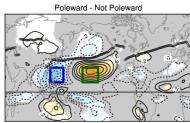
Stronger Climatological Walker Circulation



- 0

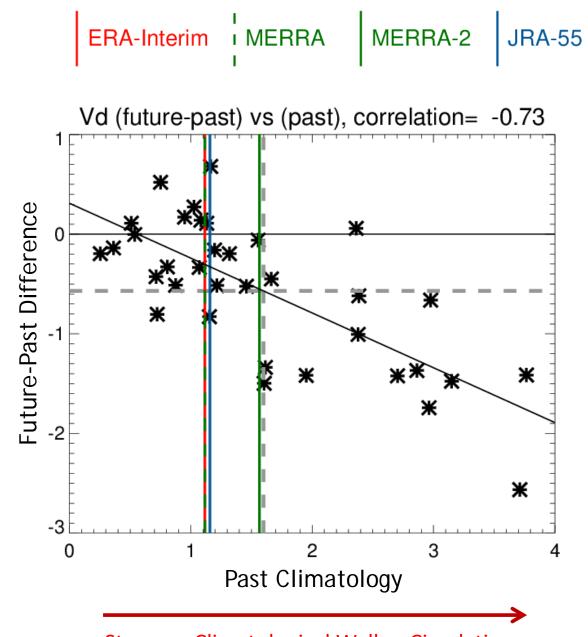


Stronger Climatological Walker Circulation

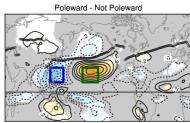


- 🛛



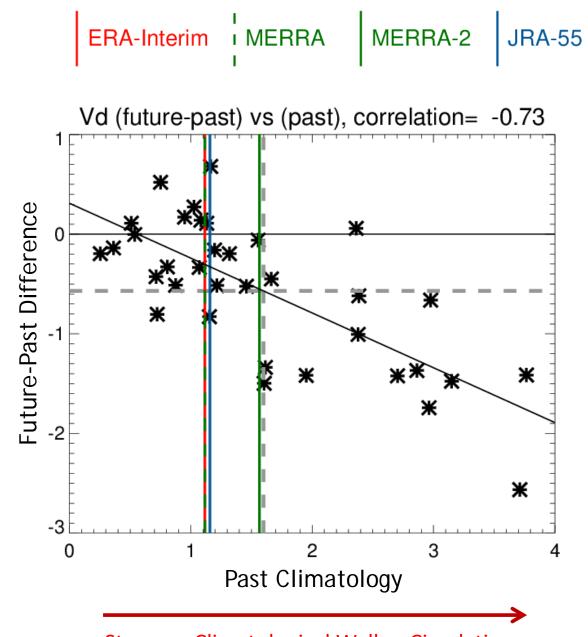


Stronger Climatological Walker Circulation



- 🛛





Stronger Climatological Walker Circulation



But Why?

Where should we start in trying to gain this understanding?

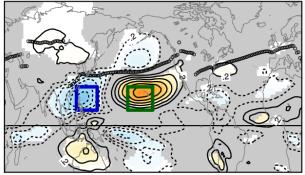
The tropics or the extra-tropics?

A natural variability analogue in CESM

- 2600y long pre-industrial control simulation (CAM5, 1deg)
- Prescribed Climatological SSTs from coupled piControl run
- Inter-annual variability (DJF averages)

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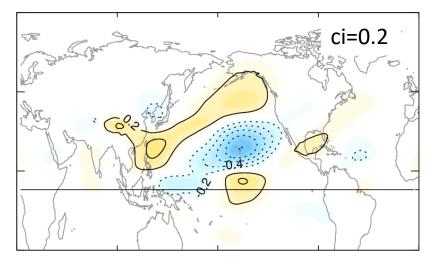
Poleward - Not Poleward



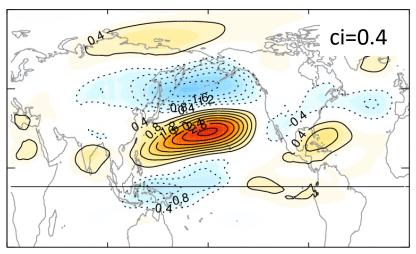
Asymmetry in divergent tropical Pacific circulation

Regress fields onto the difference in divergent v between the blue box and the green box

250hPa divergent v

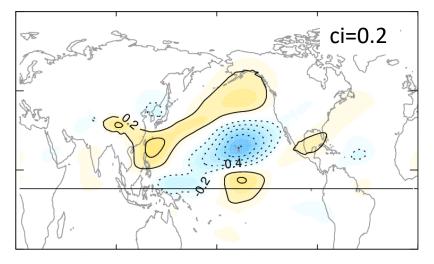


700hPa u

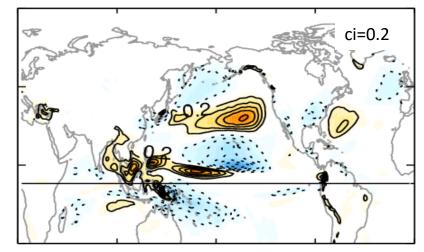


Regression onto interannual variability in divergent v index in the natural variability of CESM

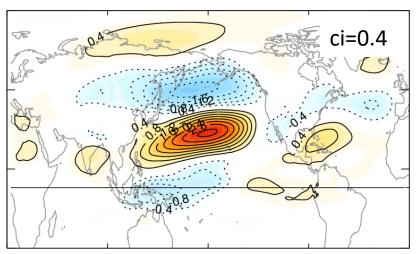
250hPa divergent v



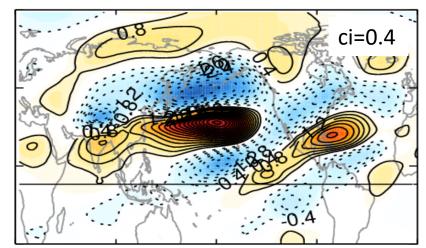
Precipitation (mm/day)



700hPa u



250hPa u



Idealized experiments within CESM to identify the leader in these tropical/mid-latitude connections

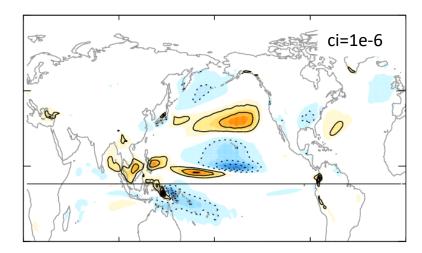
Hypothesis 1: The tropics are the leader.

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e.g., some organization of tropical precipitation and diabatic heating, forcing extra-tropical circulation anomalies

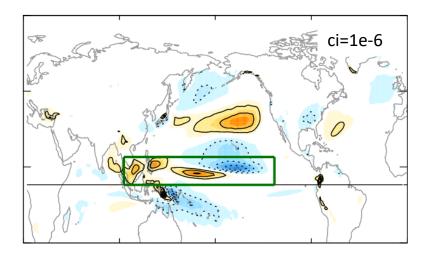
Heating anomalies (from moist processes) associated with interannual variability

Vertically Integrated Condensational Heating (Ks⁻¹) from regression



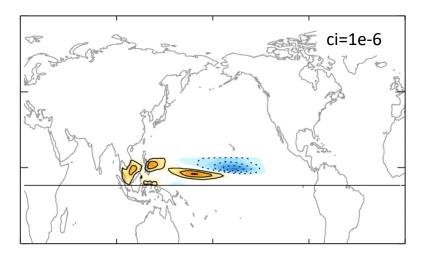
Heating anomalies (from moist processes) associated with interannual variability

Vertically Integrated Condensational Heating (Ks⁻¹) from regression



Heating anomalies (from moist processes) associated with interannual variability

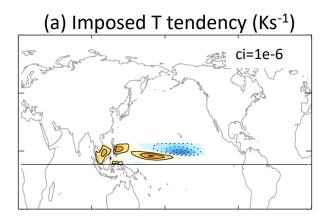
Vertically Integrated Condensational Heating (Ks⁻¹) from regression

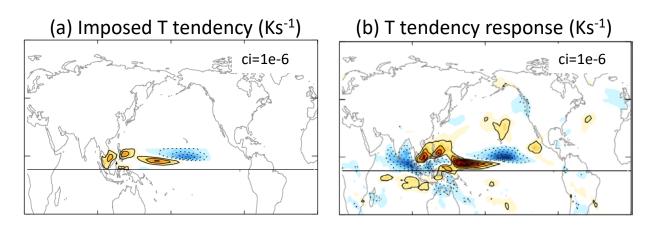


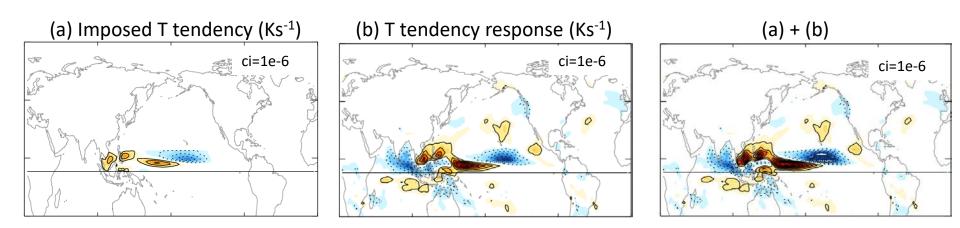
Heating anomalies (from moist processes) associated with interannual variability

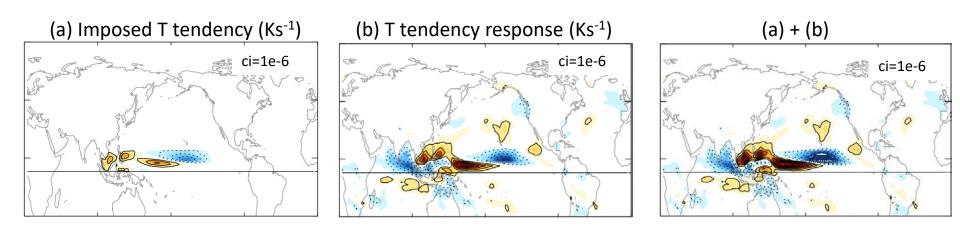
ci=1e-6 Vertically Integrated **Condensational Heating** (Ks⁻¹) from regression $\frac{\partial T}{\partial t}$ Impose a temperature tendency $\dots + T_{tend}$ within the model (CAM4, 2deg)

Thanks to Patrick Callaghan



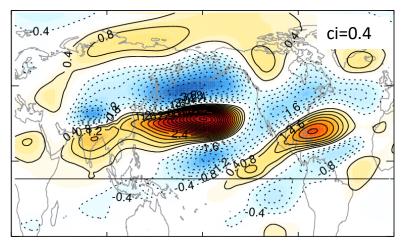


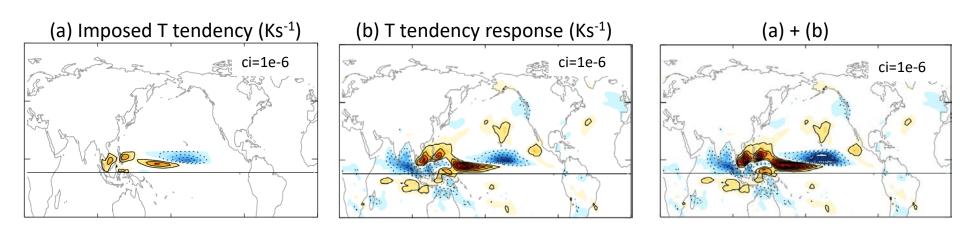




Can we reproduce this?

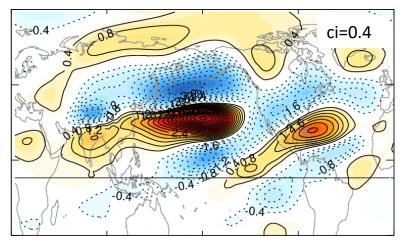
250hPa u (interannual variability regression)



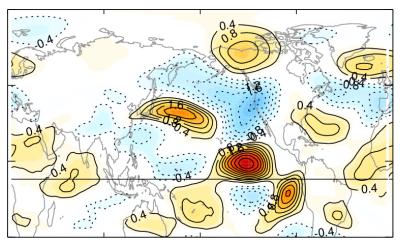


Can we reproduce this?

250hPa u (interannual variability regression)



250hPa u (response to heating perturbation)

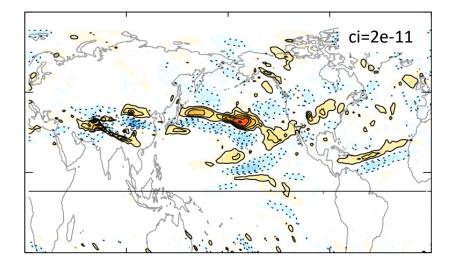


Hypothesis 2: The mid-latitudes are the leader.

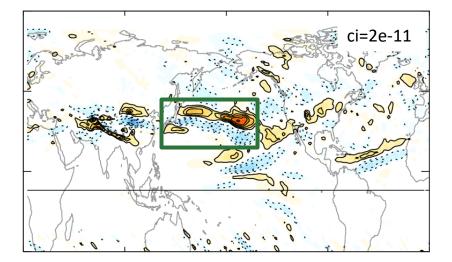
Hypothesis 2: The mid-latitudes are the leader.

e.g., a transient eddy driven shift in the westerlies with impacts on the tropical circulation

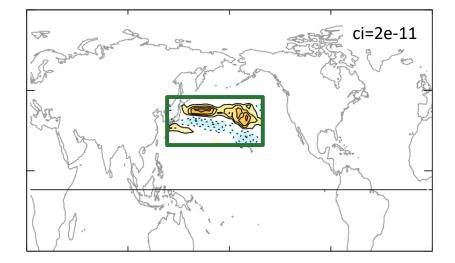
250hPa transient vorticity flux convergence regression onto divergent v index (s⁻¹)

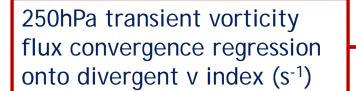


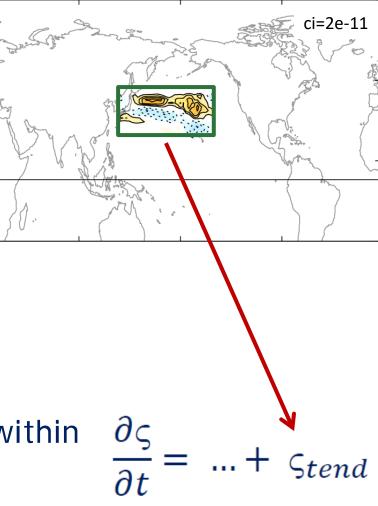
250hPa transient vorticity flux convergence regression onto divergent v index (s⁻¹)



250hPa transient vorticity flux convergence regression onto divergent v index (s⁻¹)

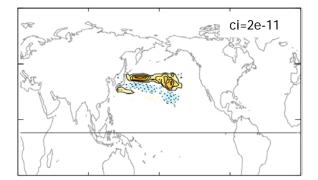


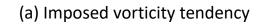


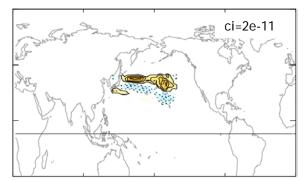


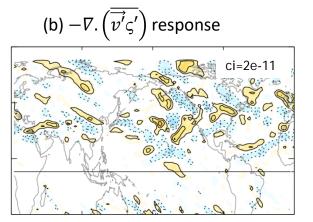
Impose a vorticity tendency within the model (CAM4, 2deg)

(a) Imposed vorticity tendency

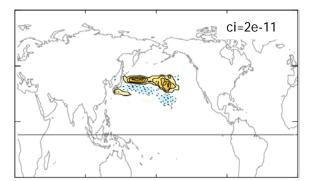


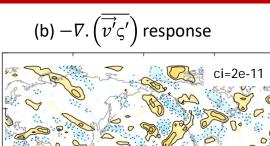


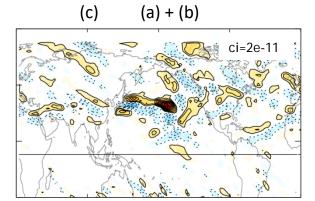




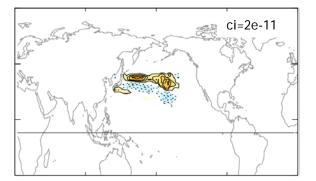
(a) Imposed vorticity tendency



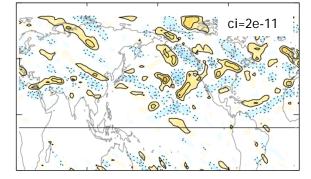




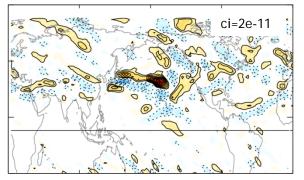
(a) Imposed vorticity tendency



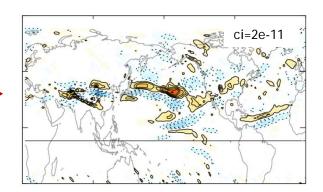
(b) $-\nabla . \left(\overrightarrow{\overline{v'} \varsigma'} \right)$ response



(c) (a) + (b)

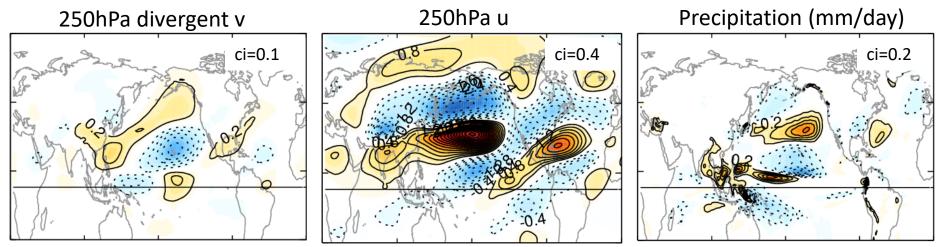


Transient vorticity flux convergence regression onto divergent v index



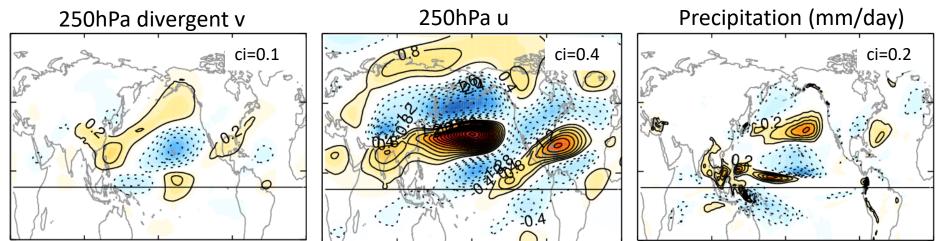
Can we reproduce this?

(regressions onto divergent v index)



Can we reproduce this?

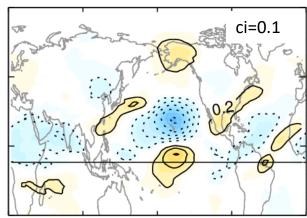
(regressions onto divergent v index)

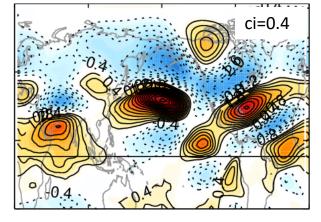


Response to vorticity tendency

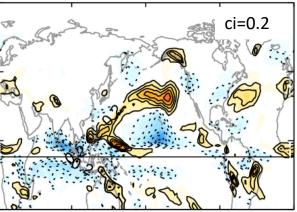
250hPa divergent v







Precipitation (mm/day)



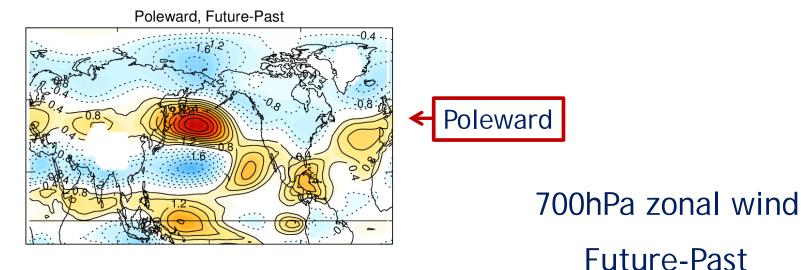
The origin of the different behavior among the models likely lies in the extra-tropics.



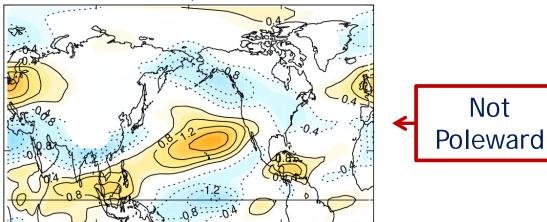
Why?

What I'm trying to do now...

CCSM4 is a "not poleward" model

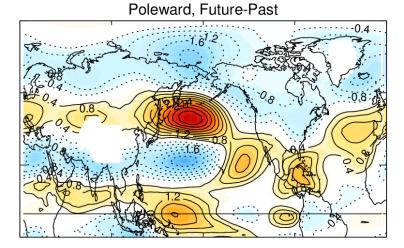


Not Poleward, Future-Past

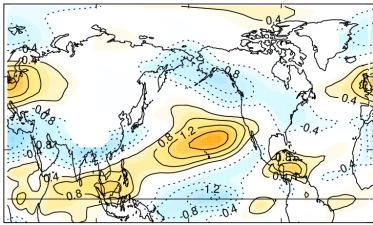


What I'm trying to do now...

CCSM4 is a "not poleward" model

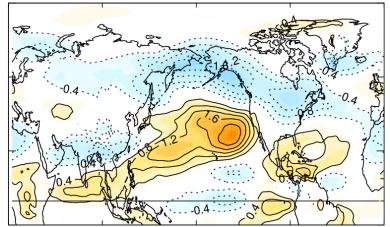


Not Poleward, Future-Past

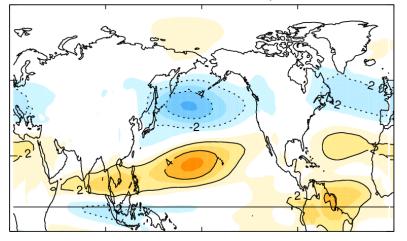




CCSM4, Future-Past

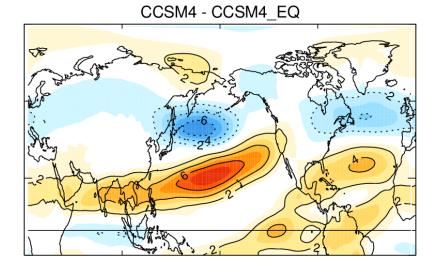


Poleward - Not Poleward, Past

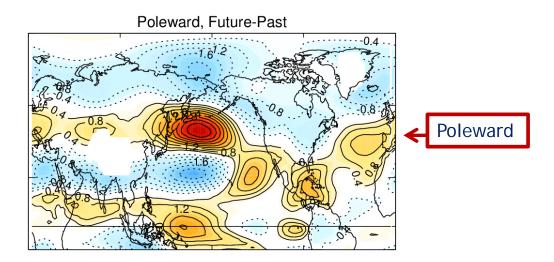


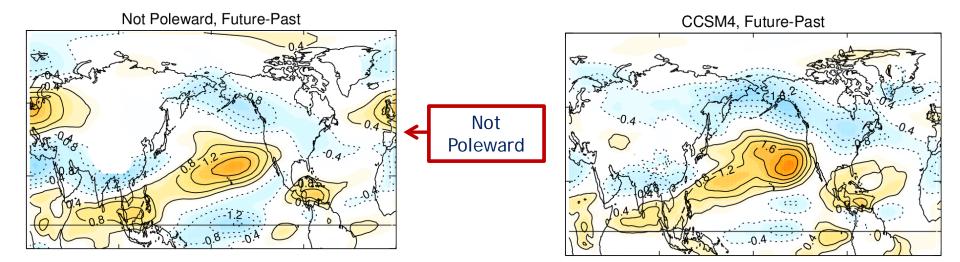
Past Climatology Poleward - Not Poleward

Force CCSM4 to have more equatorward westerlies via an imposed vorticity tendency (CCSM4_EQ)

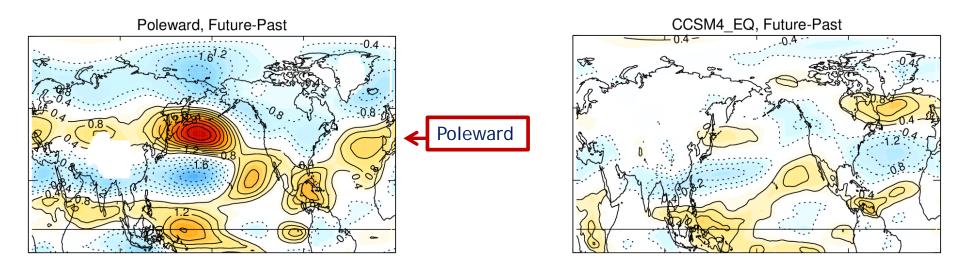


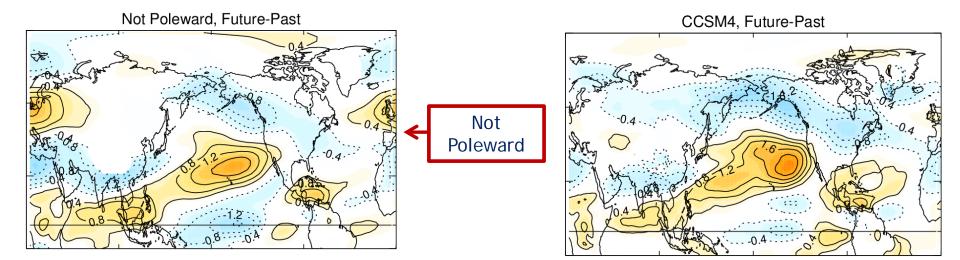
Future - Past Differences





Future - Past Differences

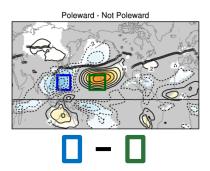




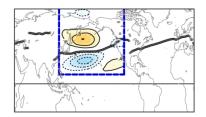
Conclusions

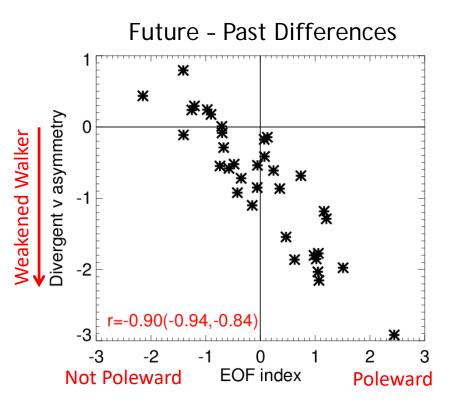
- Models exhibit diverse behavior in the tropical and extra-tropical Pacific circulation response to climate change in DJF.
- The tropical and the extra-tropical responses are highly correlated.
- A model's response is also strongly related to its climatological circulation.
- Models with a more equatorward westerly jet stream and stronger Walker circulation in their present day climatologies exhibit stronger poleward shifting of the westerlies and weakening of the Walker circulation.
- Hypothesis: The climatology of the westerly jets stream governs the models extra-tropical westerly jet stream response to climate change. How the extra-tropical westerlies change has an impact on the tropics.

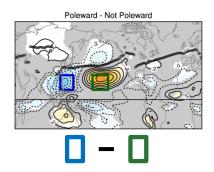
We already have mechanisms proposed for why jet latitude affects the strength of eddy feedbacks/forced responses (Barnes et al 2010, Simpson et al 2012). Hopefully further analysis of CCSM4 runs will indicate whether these, or other mechanisms apply here.



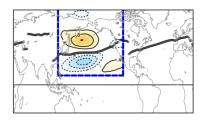
versus







versus



More

Equatorward

Climatology

Future - Past Differences (Anomalies from the multimodel mean) Weaker Climatological Walker Circulation ж *** *** Ж * * Divergent v asymmetry Divergent v asymmetry 3 0 Ж × Ж Weakened Walker Ж ж 2 Ж ¥ ж ж ¥ -2 * * r=-0.78(-0.85,-0.70) r=-0.90(-0.94,-0.84) -3 0 -3 -2 2 3 -3 0 2 -1 0 -2 -1 1 1 EOF index **Not Poleward** EOF index Poleward

Past Climatologies

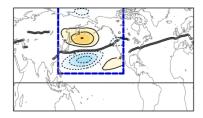
3

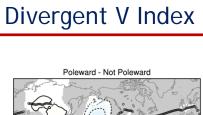
More

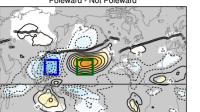
Poleward

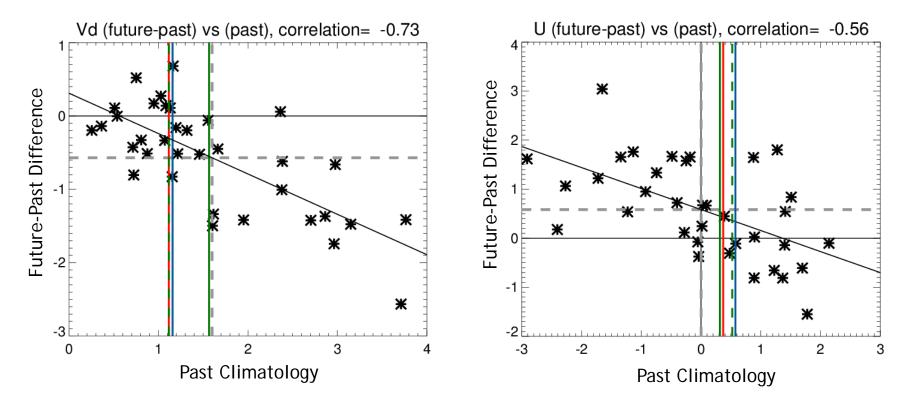
Climatology

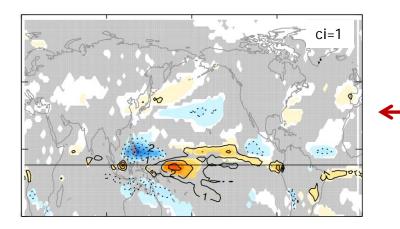
U Index



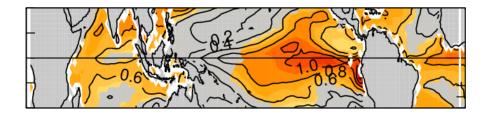








Poleward – Not Poleward precipitation anomalies (mm/day)

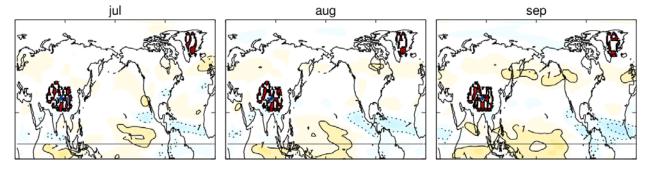


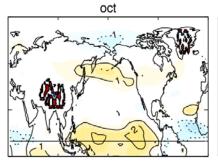
Poleward – Not Poleward SST anomalies (K)

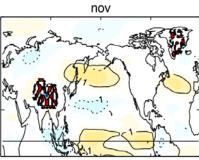
Not Poleward Models

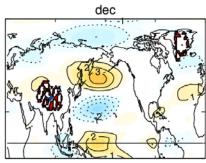
CMCC-CMS CMCC-CMS CSIRO-Mk3-6-0 GISS-E2-R IPSL-CM5A-MR MIROC5 MPI-ESM-LR MPI-ESM-MR MRI-CGCM3

BNU-ESM CanESM2 CCSM4 CNRM-CM5 FIO-ESM GFDL-ESM2G GFDL-ESM2M inmcm4 NorESM1-ME





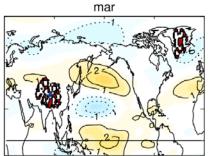




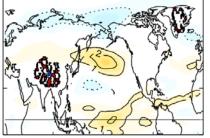
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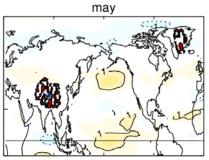
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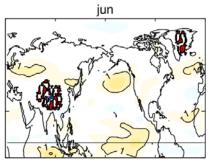
feb . .1 00

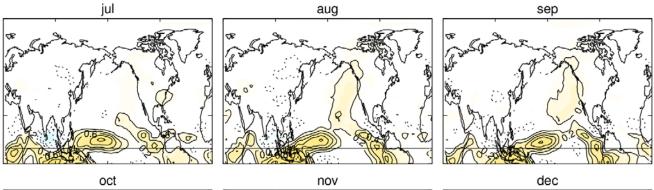


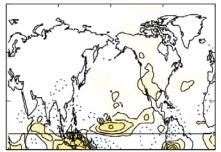


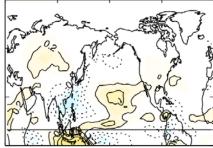


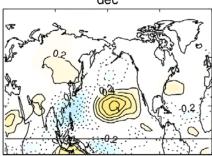








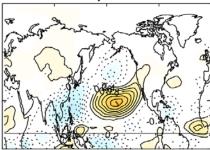


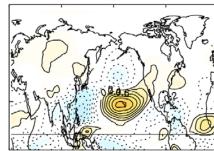


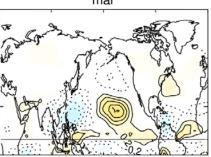








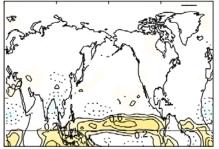


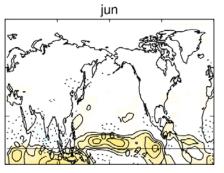




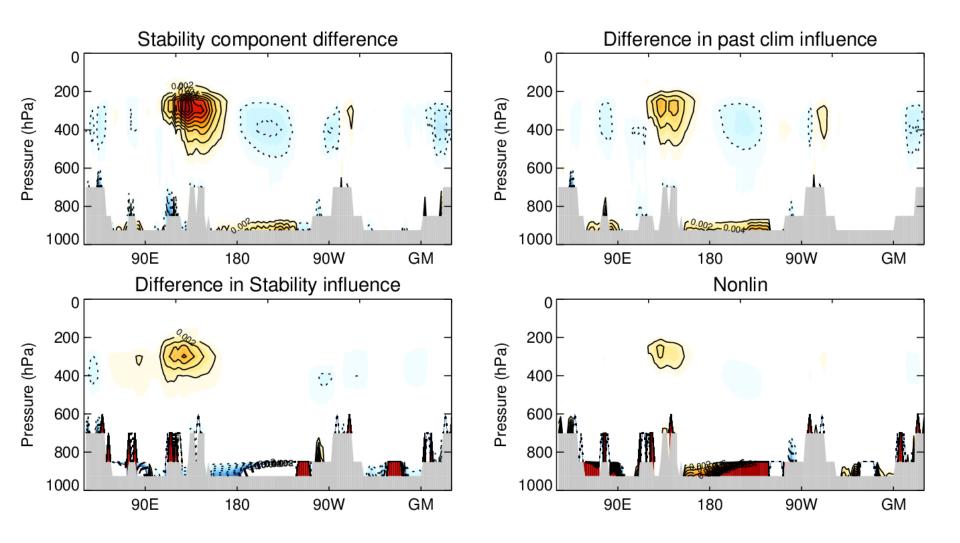


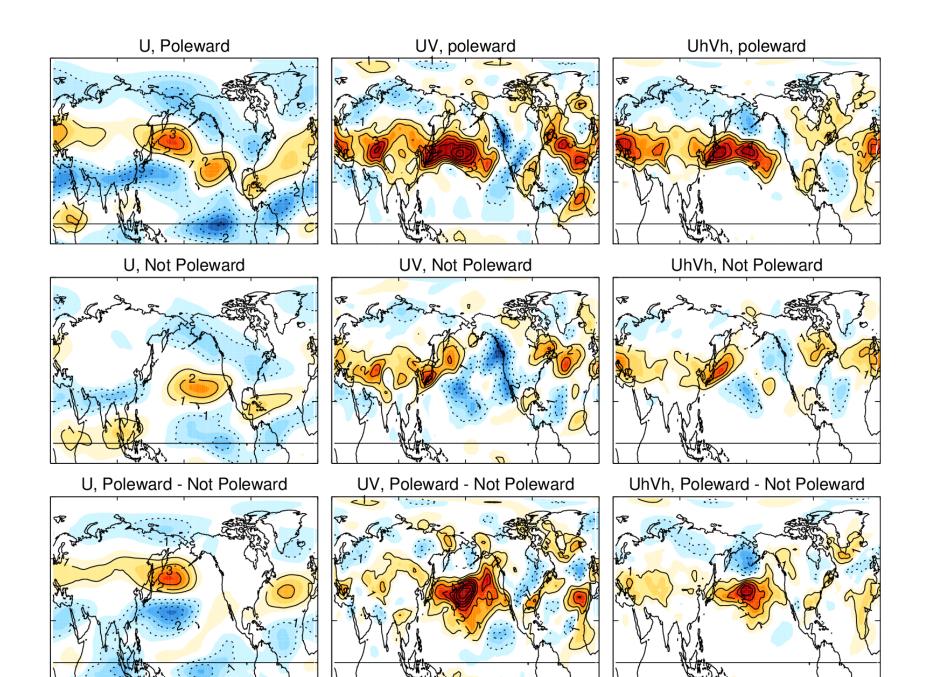




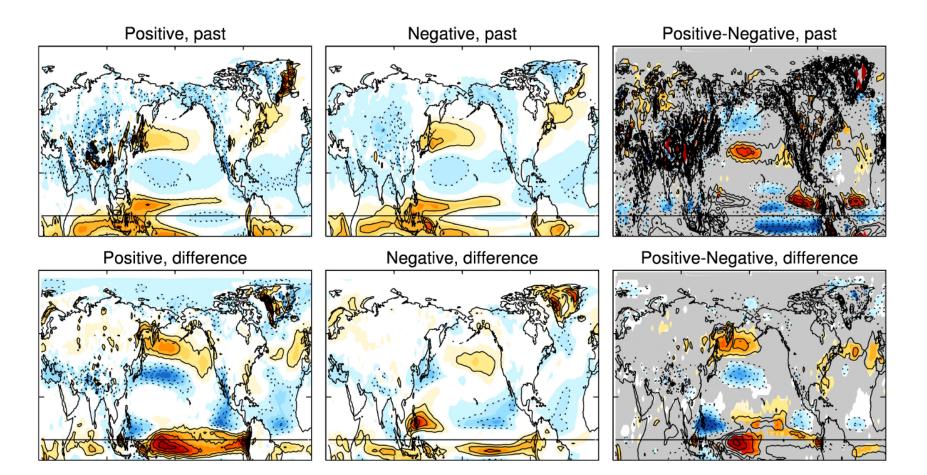


$$-\left(\frac{\partial\theta}{\partial P}\right)_{p}^{-1}\omega_{p}\Delta\left(\frac{\partial\theta}{\partial P}\right)$$

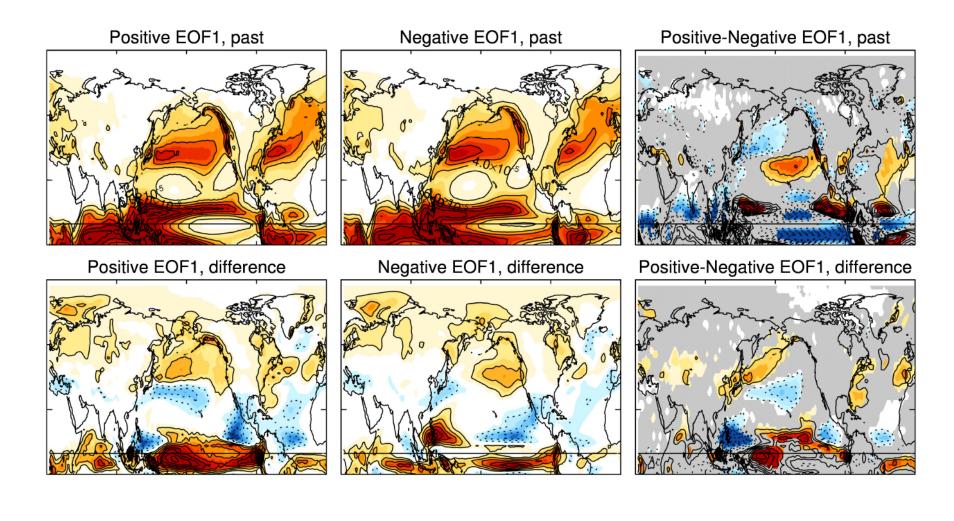


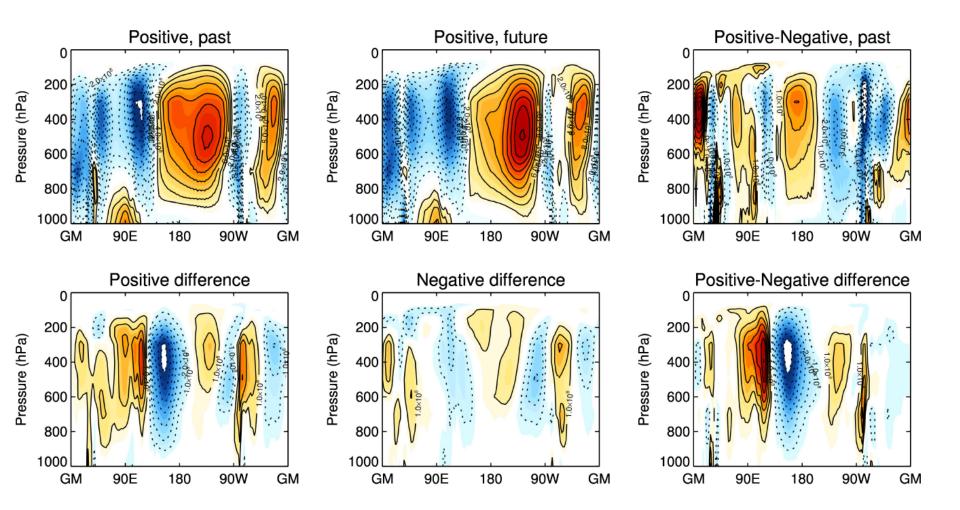


Vertically integrated diabatic heating (top) past, (bottom) future-past

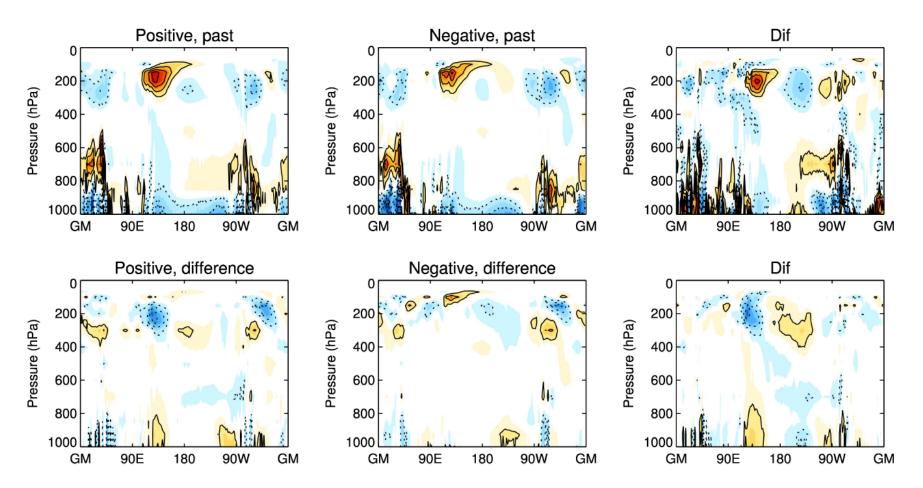


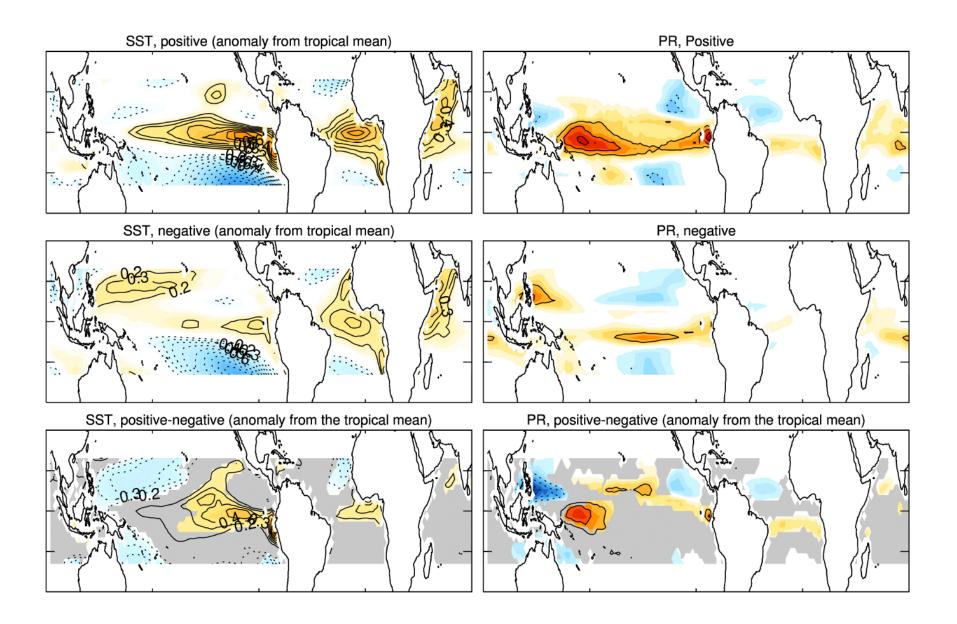
Precipitation (top) past, (bottom) future-past



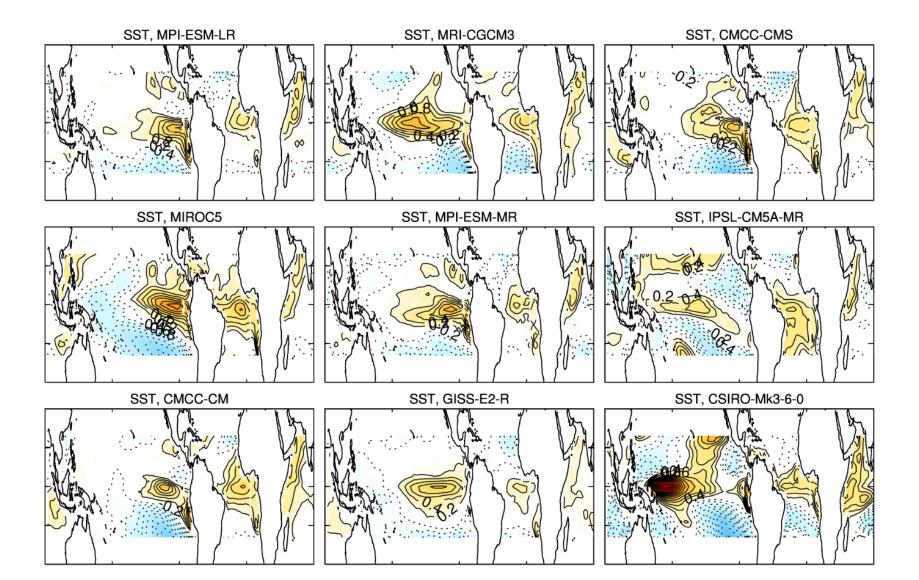


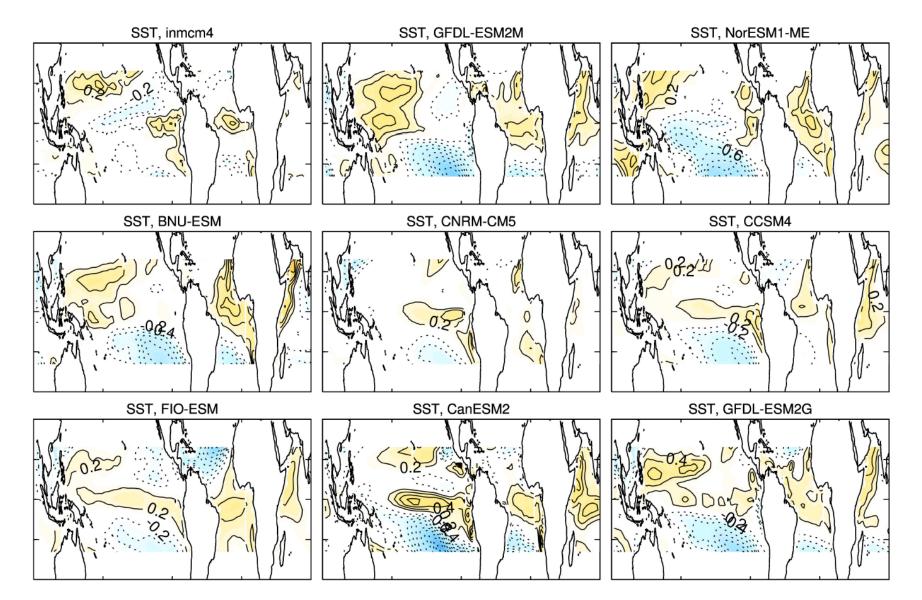
Divergence (top) Past, (bottom) Future-past





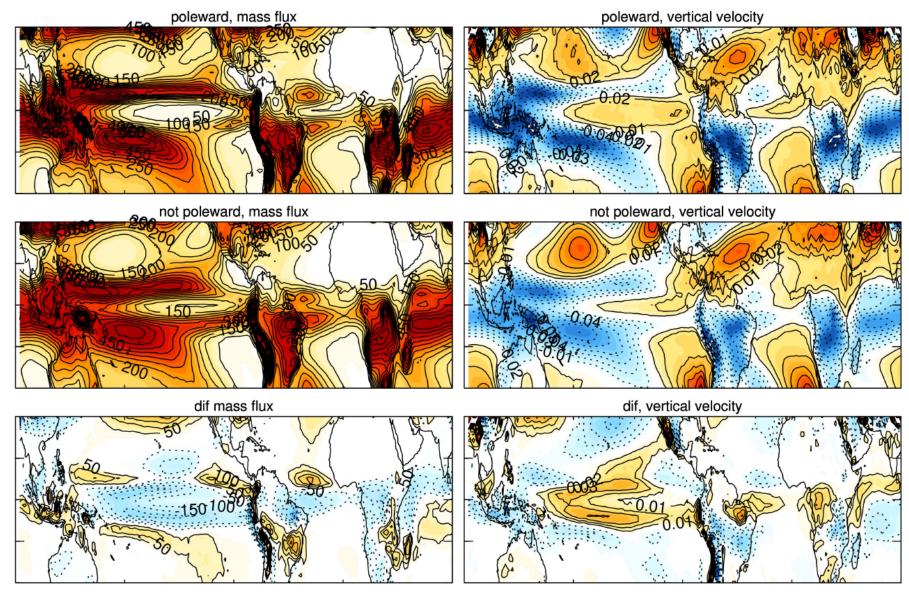
SST response, anomalies from the tropical mean, Poleward models.





SST response, anomalies from the tropical mean, Poleward models.

Convective mass flux and vertical velocity



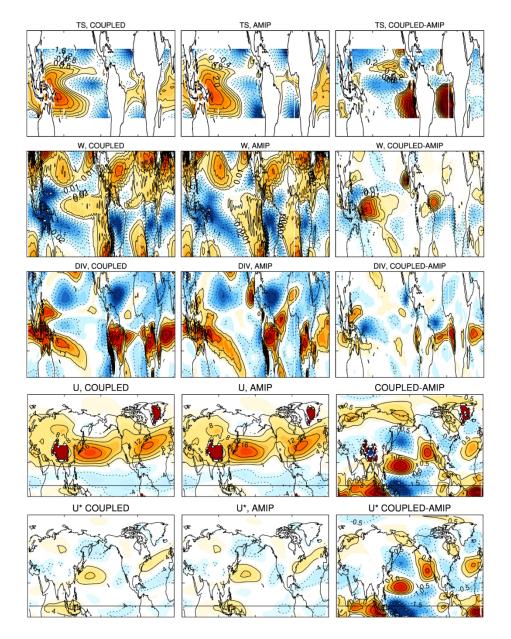


Figure 40: For the average of four models positive models (MPI-ESM-LR, MPI-ESM-MR, MIROC5, MRI-CGCM3) comparing the AMIP climatology with the coupled climatology (top) sst anomaly from the tropical mean, (2nd) 500hPa omega, (3rd) 250hPa divergence, (4th) 700hPa zonal wind, (5th) 700hPa eddy zonal wind.

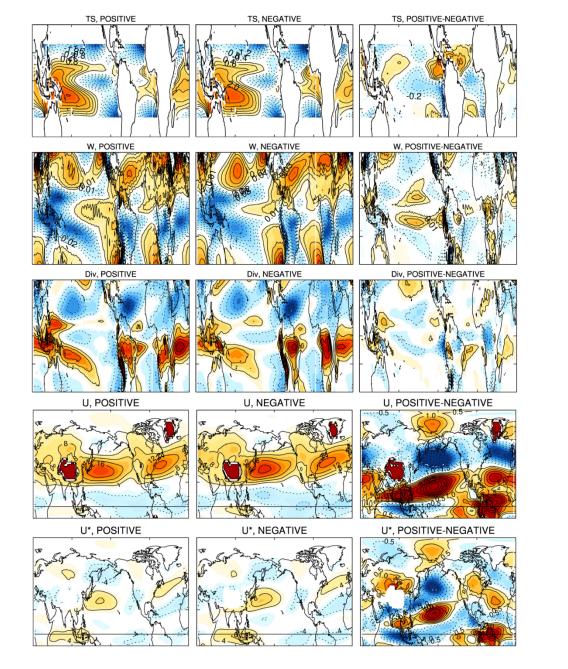
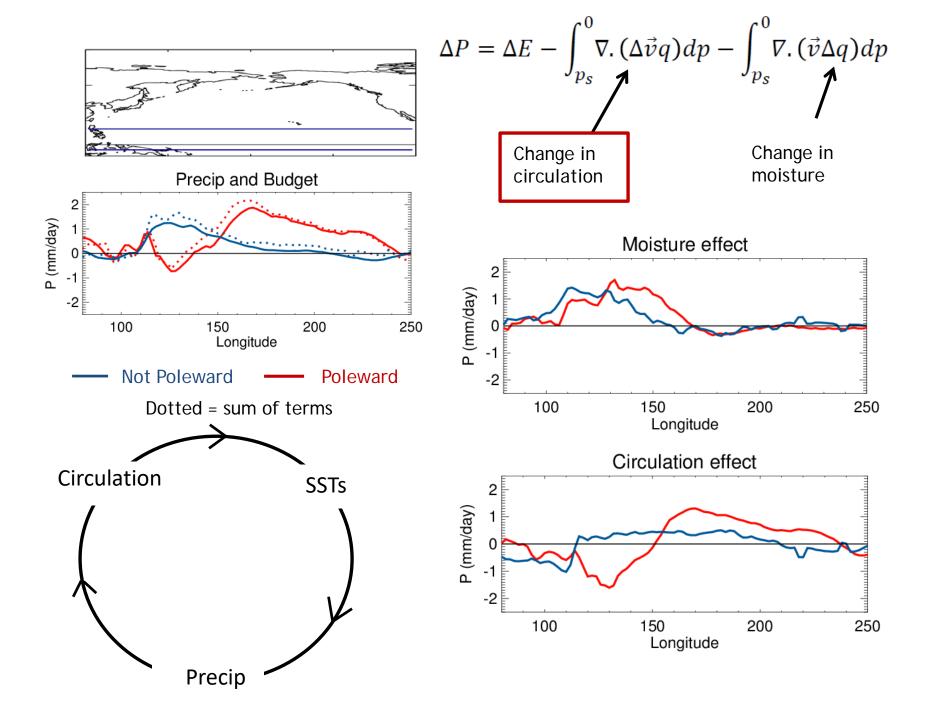
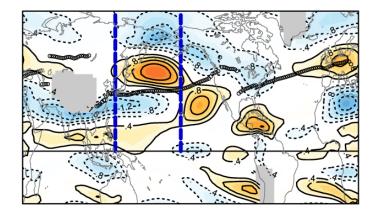


Figure 41: As Fig. 40 but for the positive and negative model composites (for comparison with Fig. 40 $\,$





West Pacific Jet Shifts

Distribution of jet shifts between pairs of 30 year samples taken at random from the piControl simulation (500y from CCSM4, 1000y from MPI-ESM-LR)

