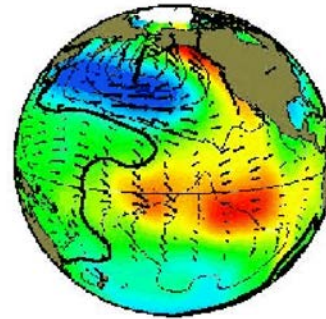
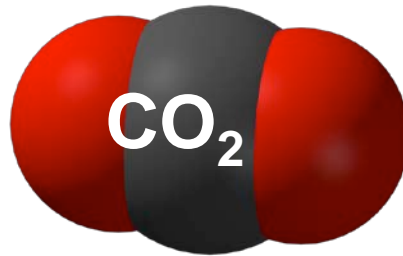
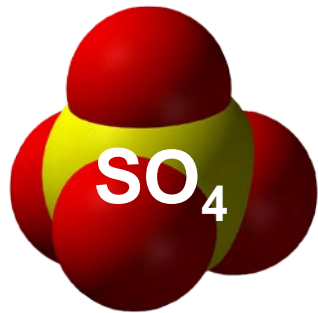
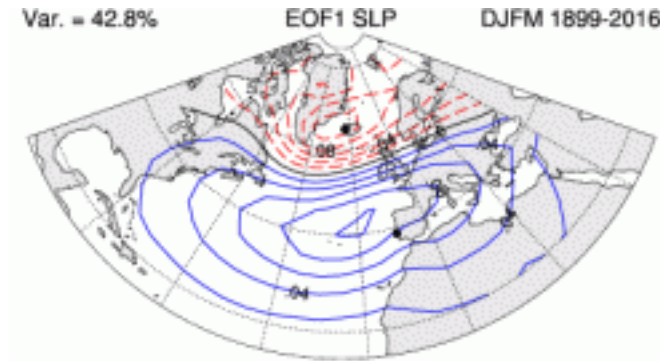


Towards an improved climate event attribution framework

Laurent Terray, Cerfacs/CNRS and CAS/CGD



PDV/IPV



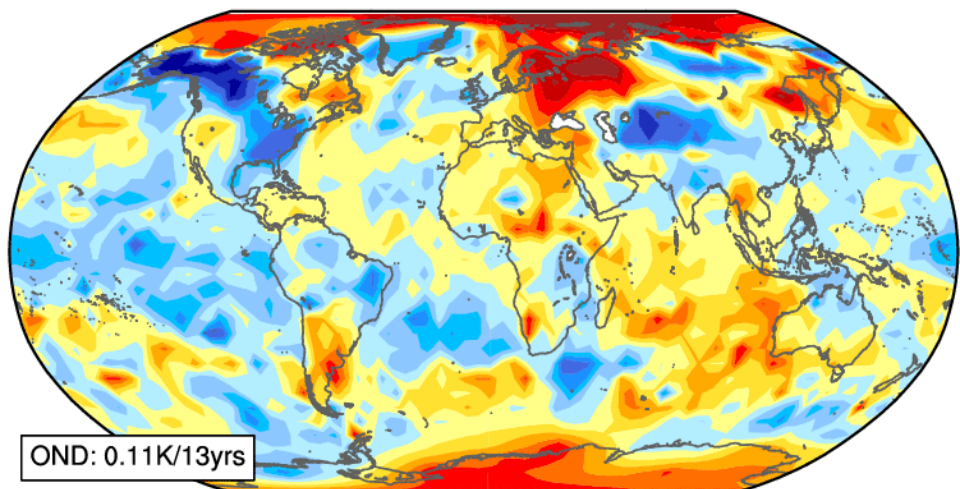
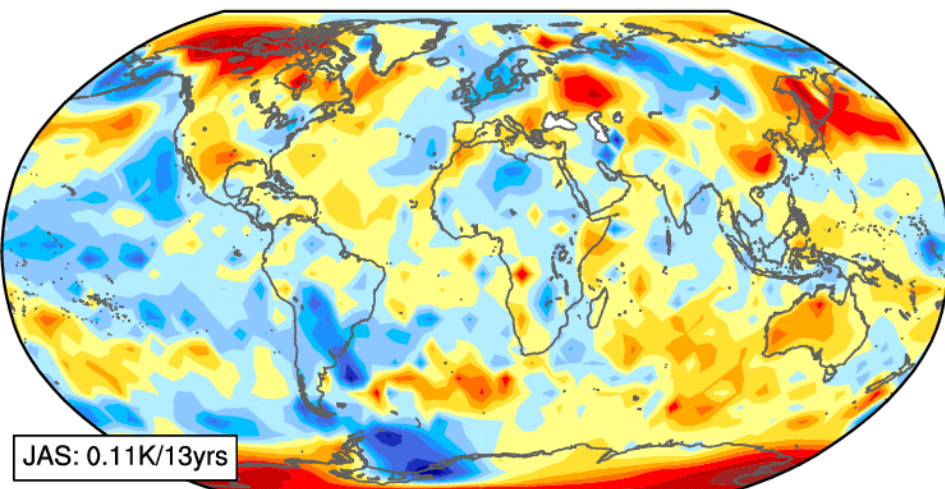
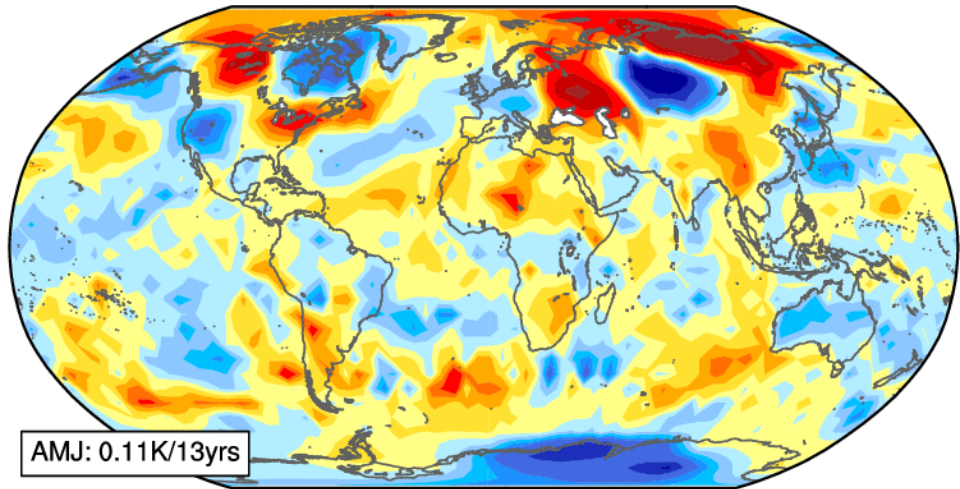
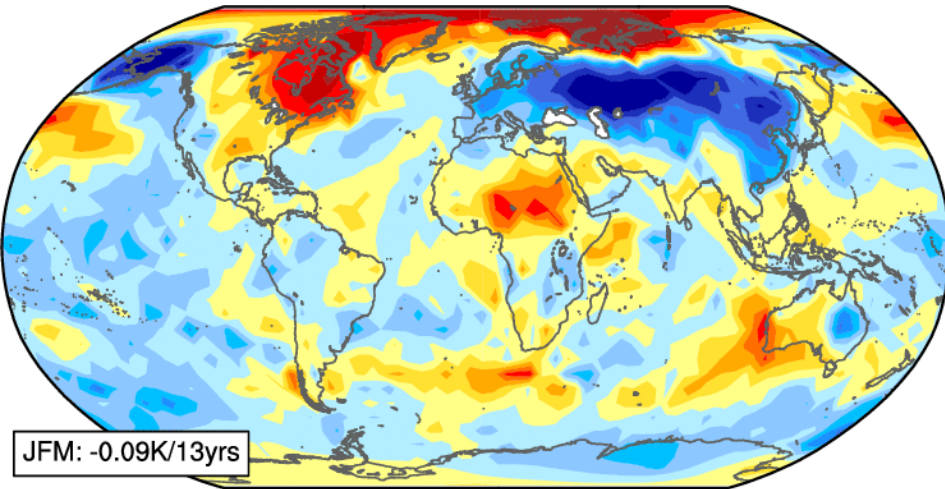
NAO/AO

The early 2000s Eurasian Winter Cooling

Thanks to all CAS members and to T. Sheperd

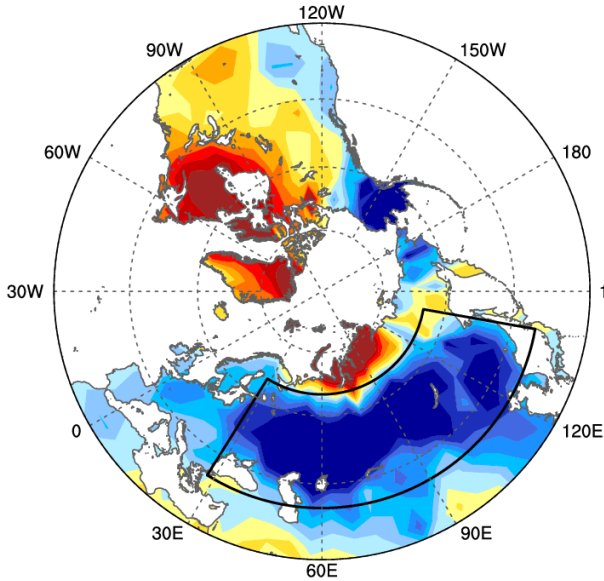
The early 2000s « Hiatus »

Cowan&Way 2001-2013 linear trend

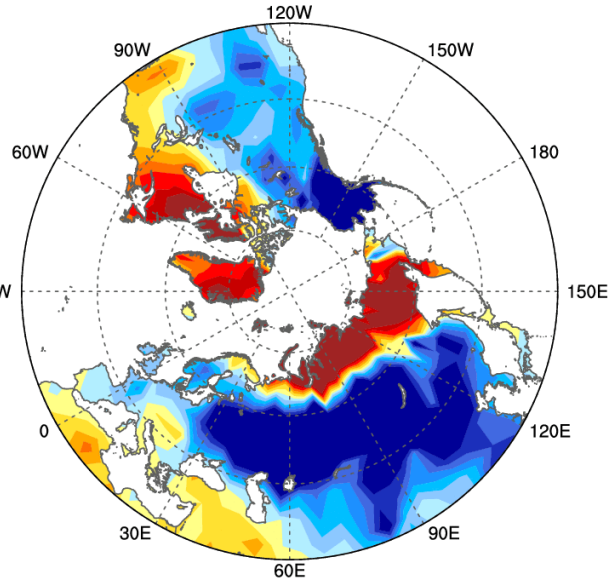


Observed Winter T2m trend 2001-2013

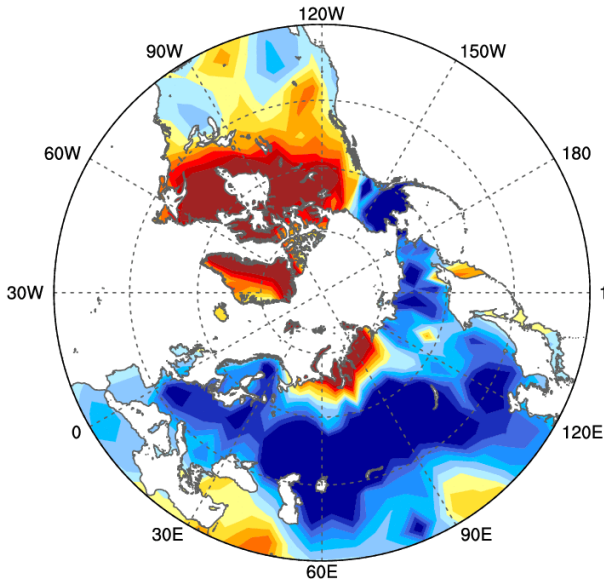
**Winter
JFM**



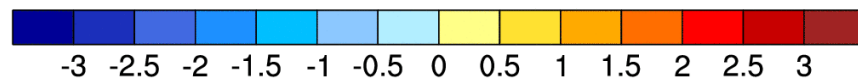
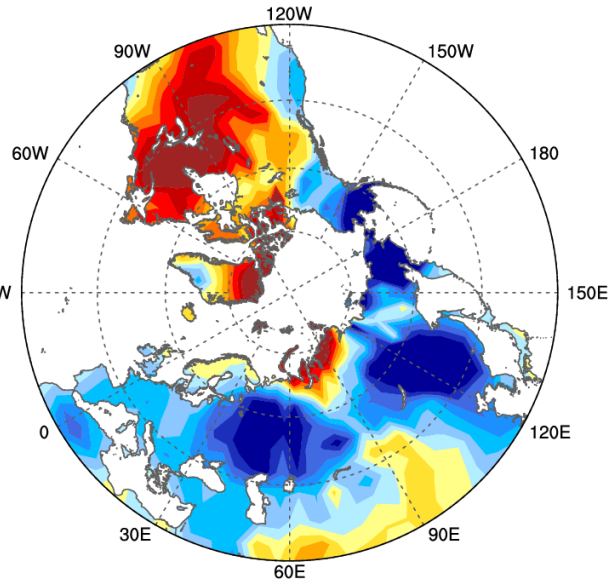
Jan.



Feb.

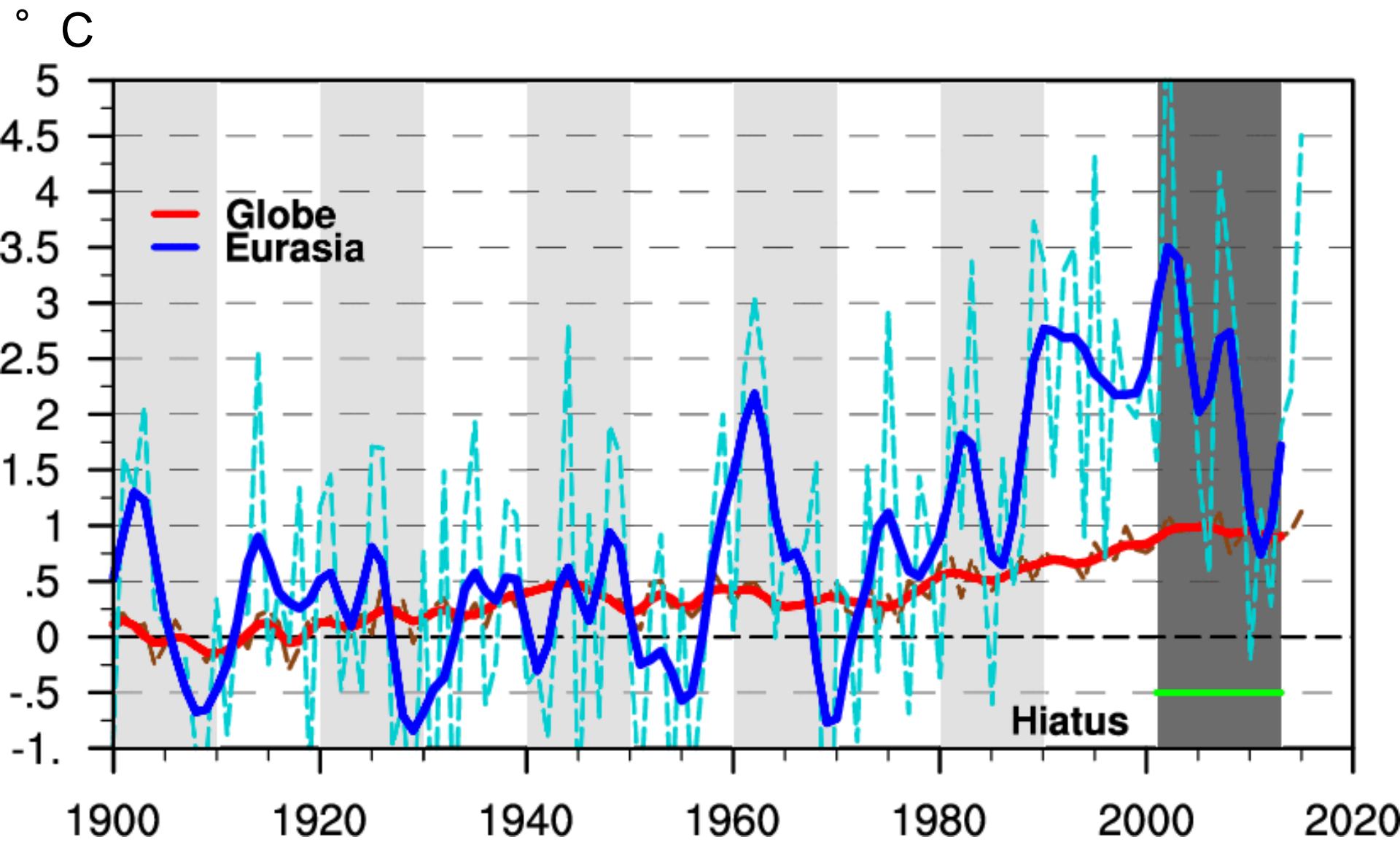


March



$^{\circ}$ C / 13 years

Observed T2m in winter (JFM)



Science current status ?

- (2014): « **tropical Pacific forcing** of the atmosphere such as that associated with a negative phase of the PDO produces **many of the pronounced atmospheric circulation anomalies** observed globally during the hiatus »
- (2014): « Here we use a 100-member ensemble of simulations ... **to show that as a result of sea-ice reduction in the Barents–Kara Sea, the probability of severe winters has more than doubled in central Eurasia.** »
- (2015): « Our experimental results suggest that the **Arctic sea ice loss does not drive systematic changes in the Northern Hemisphere large-scale circulation in the past decades.**
- (2016): « In our atmospheric-only simulations, **we find no evidence of Barents and Kara seas sea-ice loss having impacted Eurasian surface temperature** »
- (2016): « The findings confirm that **sea-ice concentrations in Autumn in the Barents and Kara seas are an important driver of winter circulation in the midlatitudes.** »
- External forcing (volcanic & solar), atmospheric internal variability ...

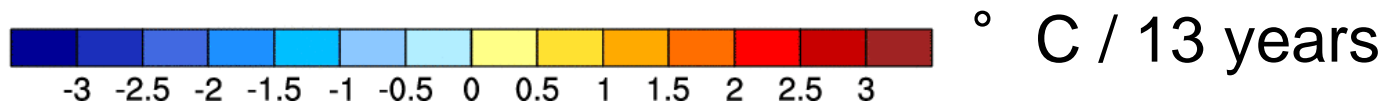
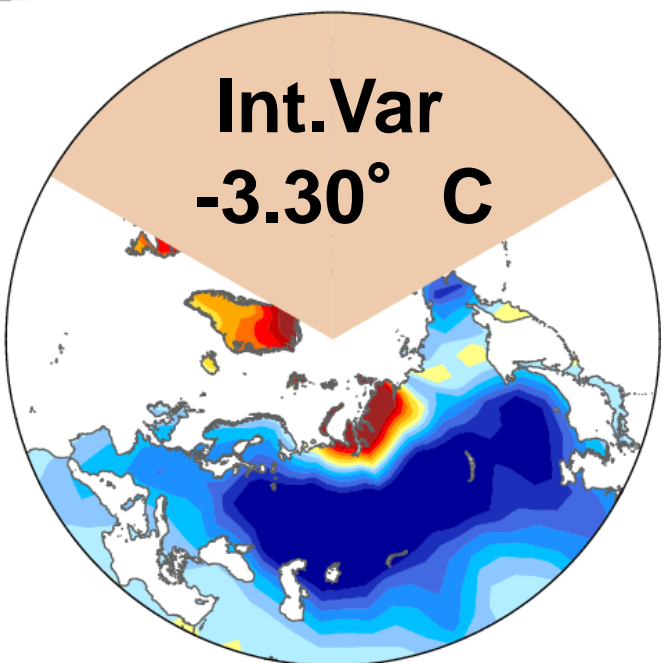
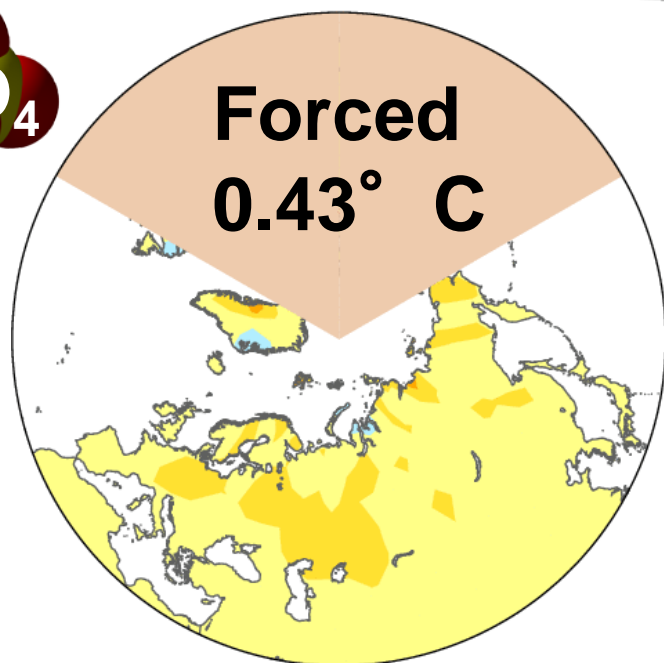
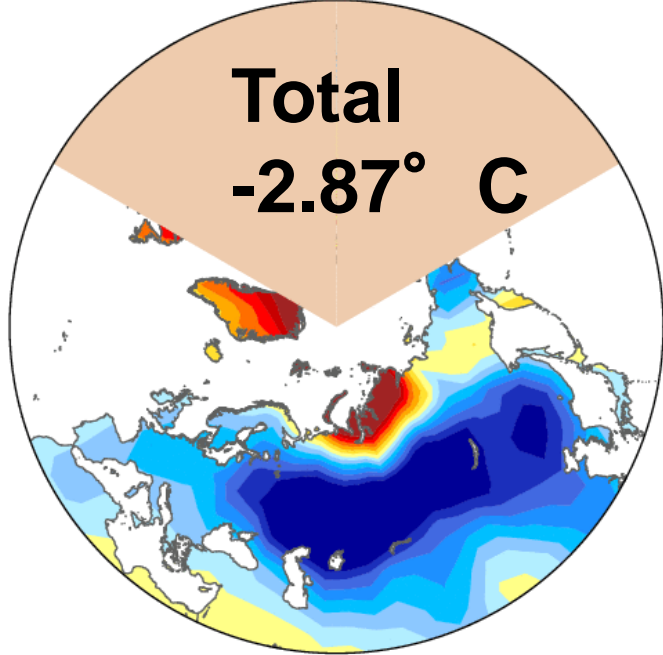
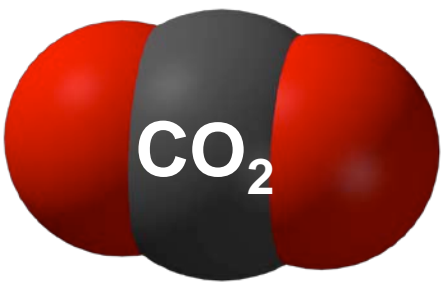
2001-2013 Eurasian cooling: -2.87° C/13
years

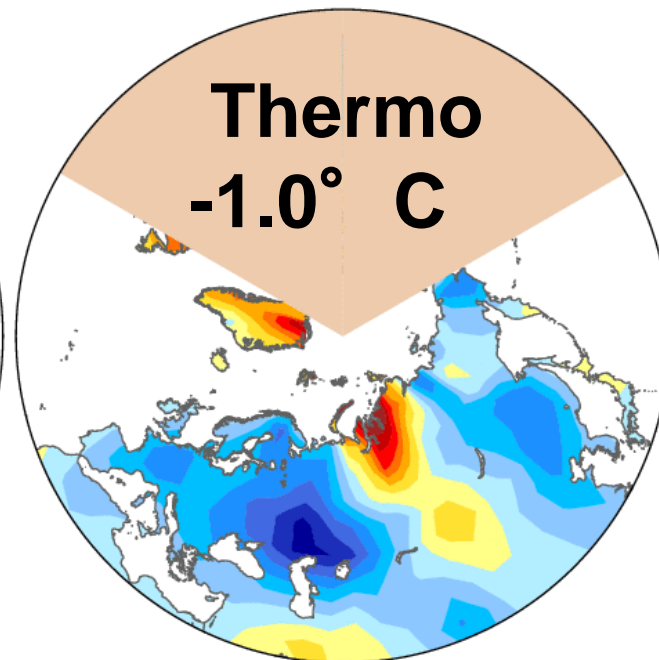
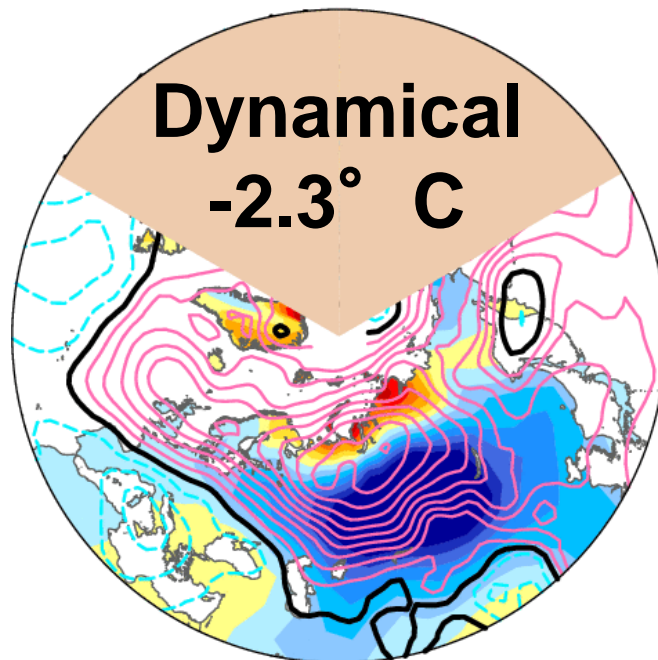
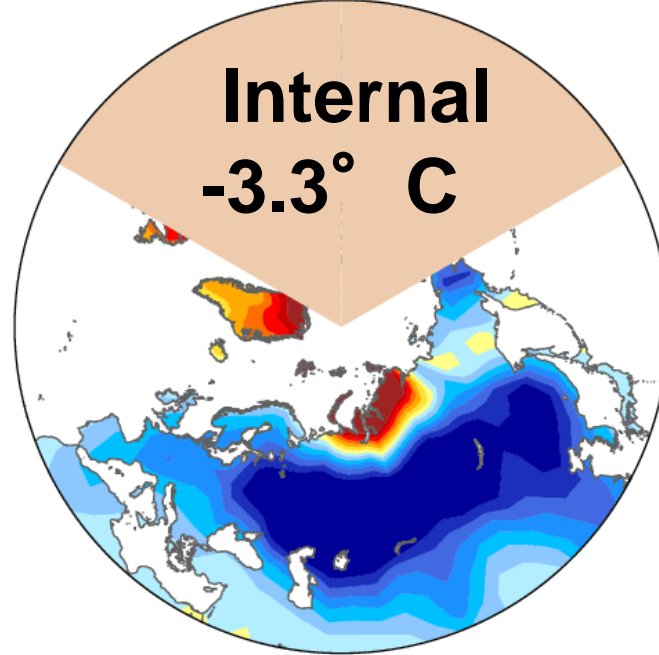
| Processes/ Causes | Thermodynamical | Dynamical |
|------------------------------|------------------------|------------------|
| Forced | ??? ° C | ??? ° C |
| Free | ??? ° C | ??? ° C |

1. **Quantify** the different factors with their mechanisms
2. Give **uncertainty** estimates !

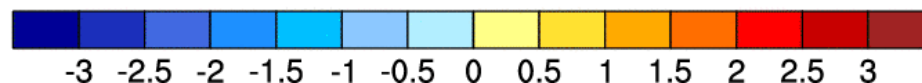
Methods

- External forcing: use Ensemble Empirical Mode Decomposition (EEMD) to derive the forced response
- Dynamical adjustment: use a non-linear regression method (gradient boosting regression trees) to predict temperature from sea level pressure (20CR)
- Use leave-one-out method to derive circulation-related temperature for each month





Cl: 1 hPa



° C / 13 years

2001-2013 Eurasian cooling: -2.87° C/13 years

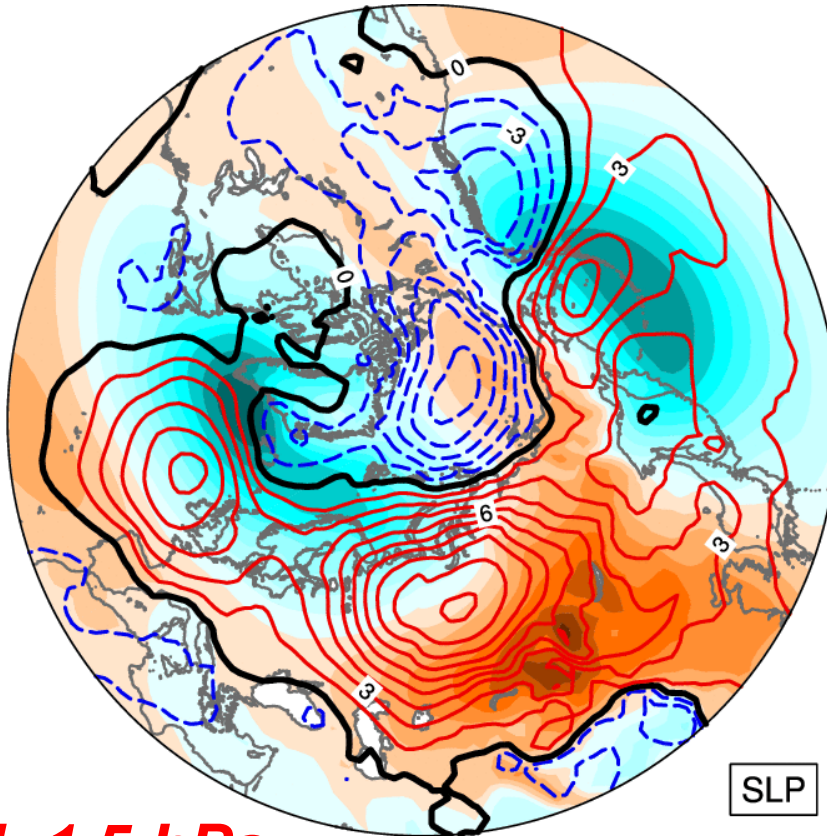
| Processes/ Causes | Thermodynamical | Dynamical |
|----------------------|--------------------------|---------------------------|
| Forced | $[0.29, 0.43]^{\circ}$ C | $\sim 0^{\circ}$ C |
| Free | $[-1., -0.75]^{\circ}$ C | $[-2.41; -2.3]^{\circ}$ C |

1. Origin of the anomalous circulation pattern ?
2. Mechanisms of the thermodynamical part ?

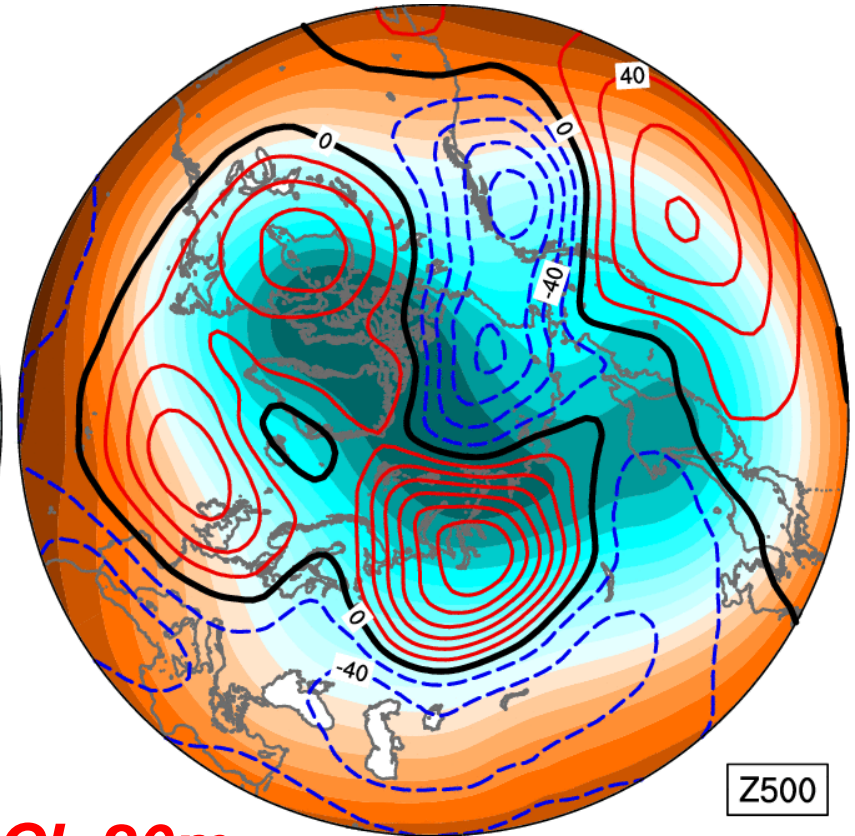
Siberian High: westward shift and intensification

Sea level pressure

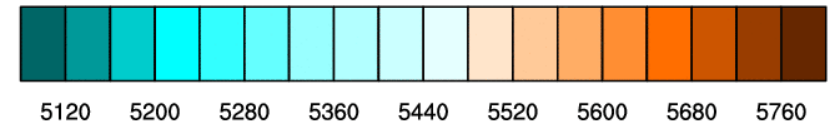
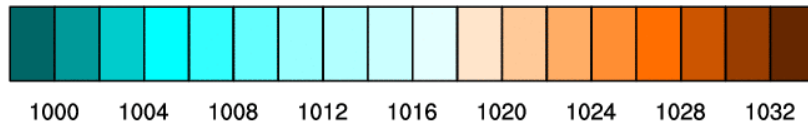
500 hPa Geopotential Height



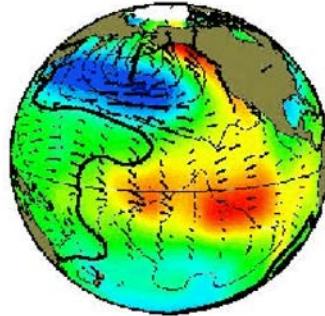
CI: 1.5 hPa



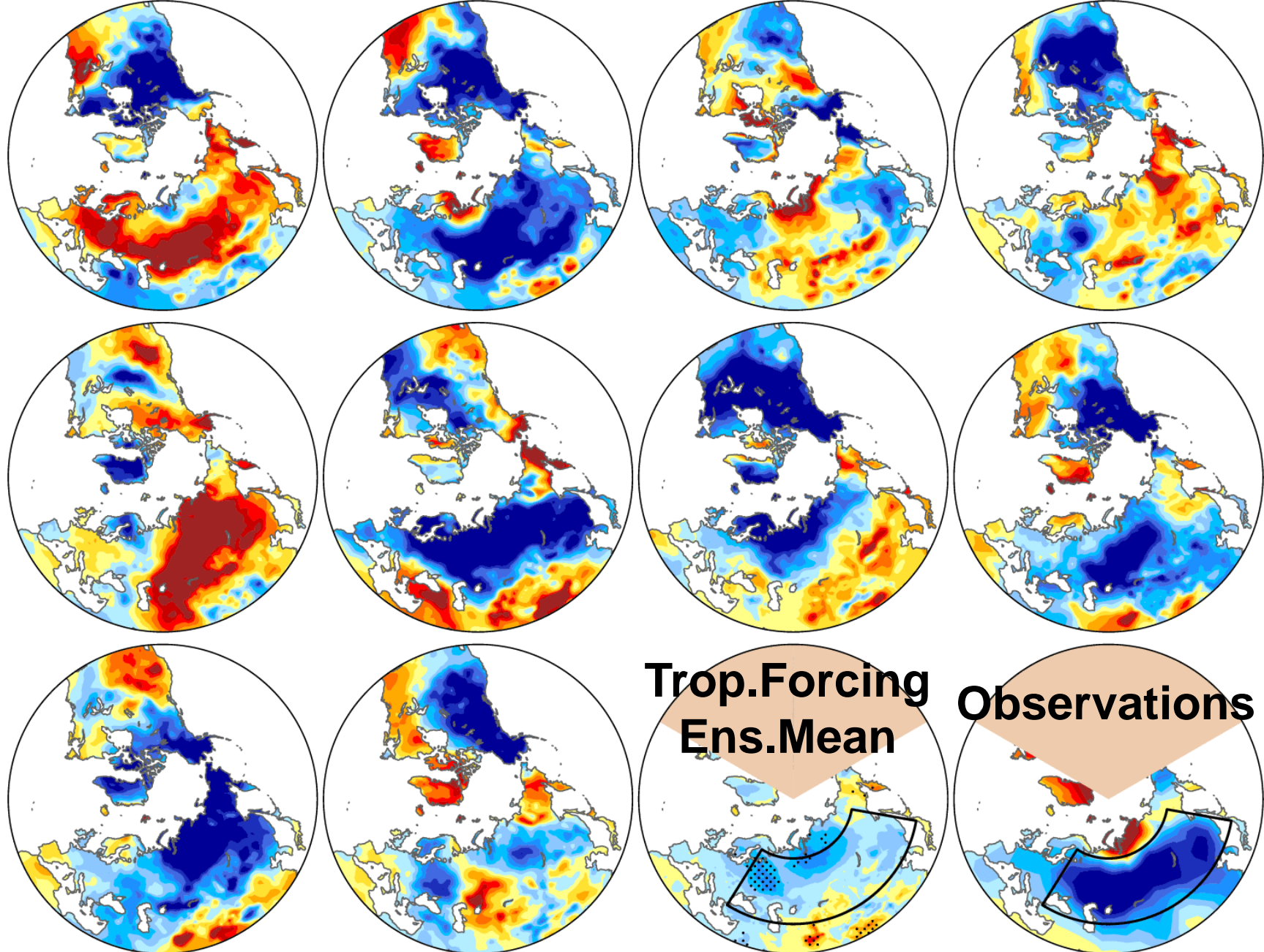
CI: 20m

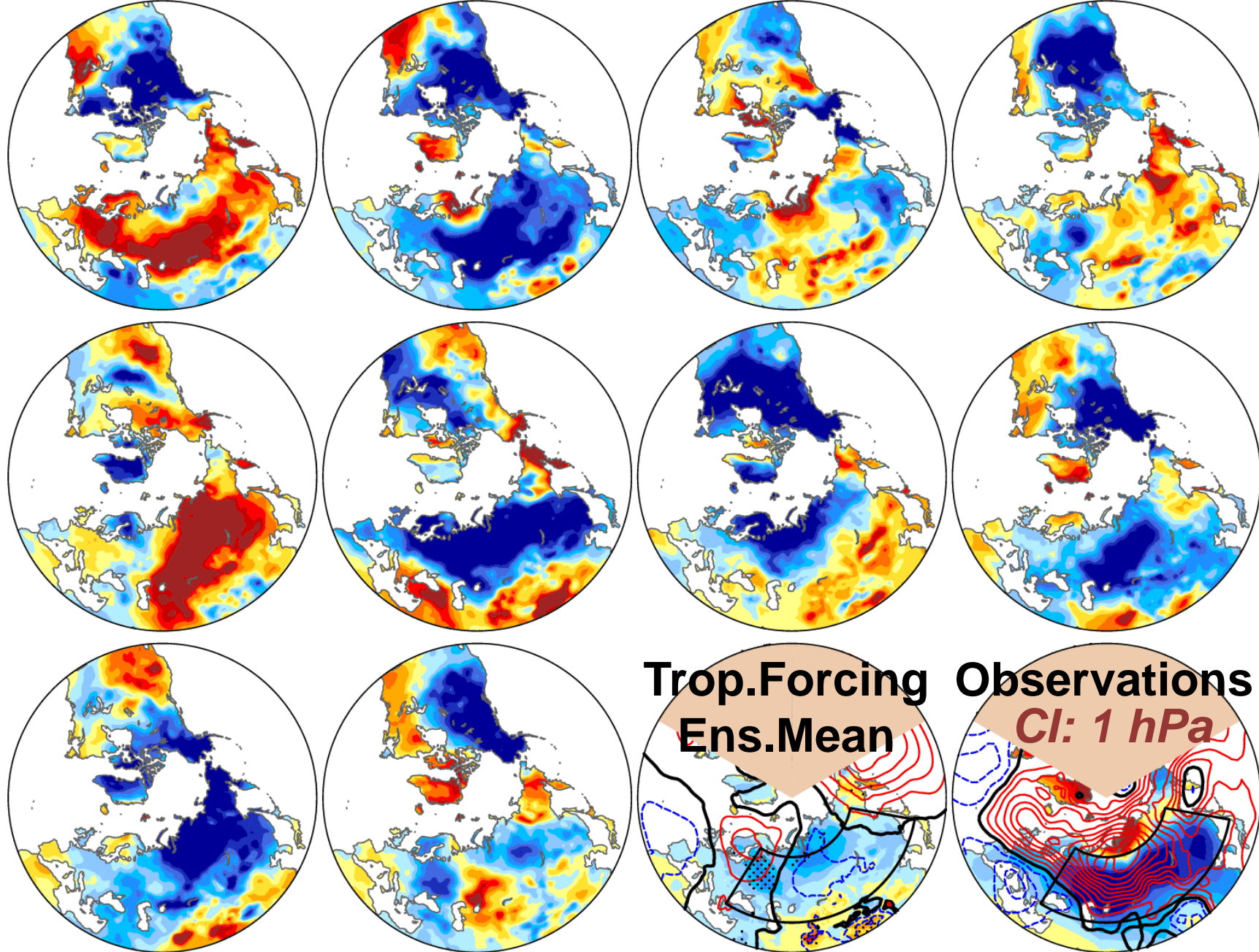


The Tropical forcing hypothesis



- Can be tested with the CESM LENS and Pacemaker (PCMK) ensembles
- Tropical forcing $\sim E_{\text{mean}}(\text{PCMK}) - E_{\text{mean}}(\text{LENS})$
- Observations and role of IPV

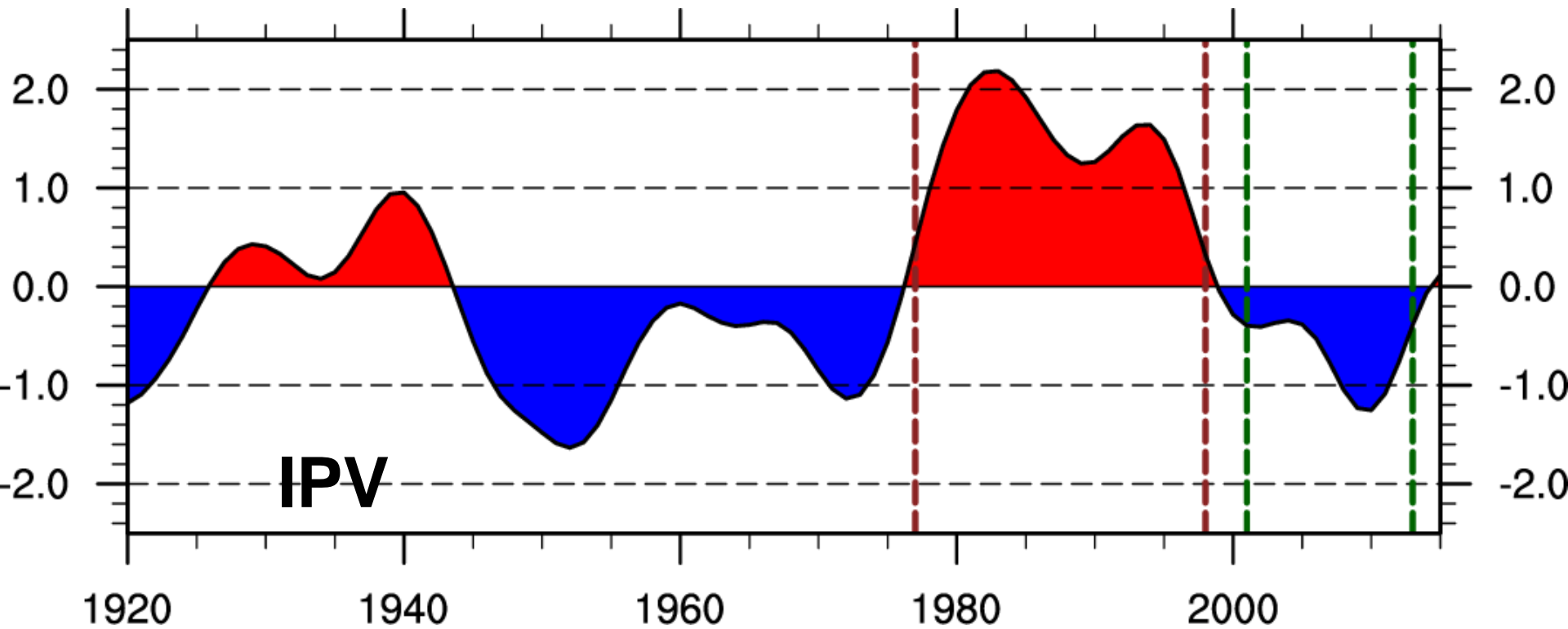




Trop. Forcing **Observations**
Ens. Mean *Cl: 1 hPa*

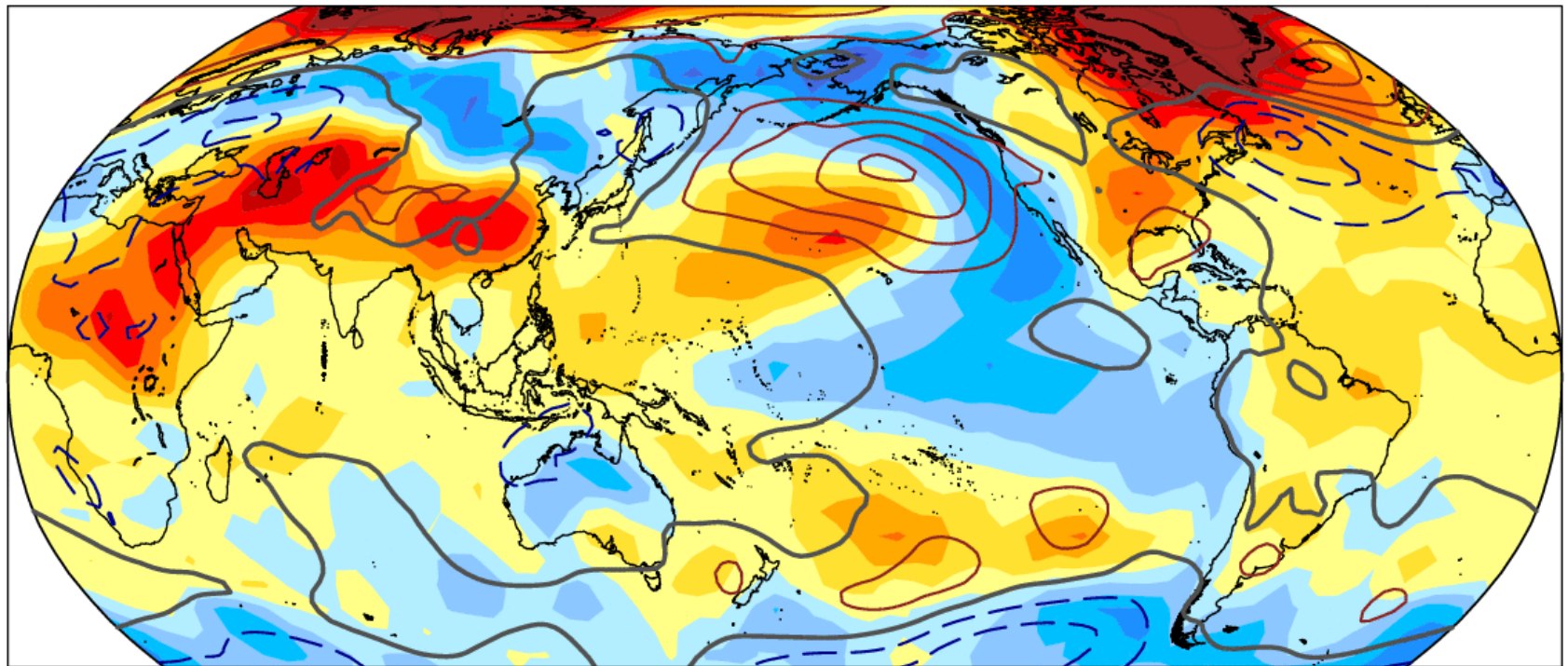


Observed IPV: contrast negative and positive phase composites

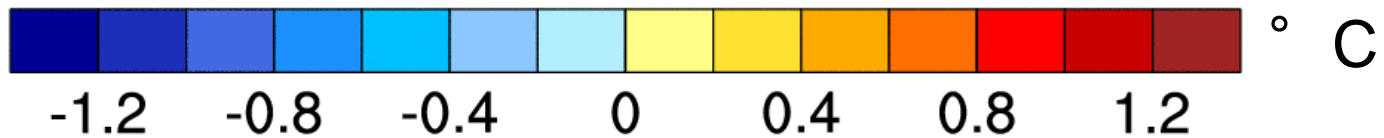


IPV influence: **Cold**(2001-2013) – **Warm**(1977-1998)

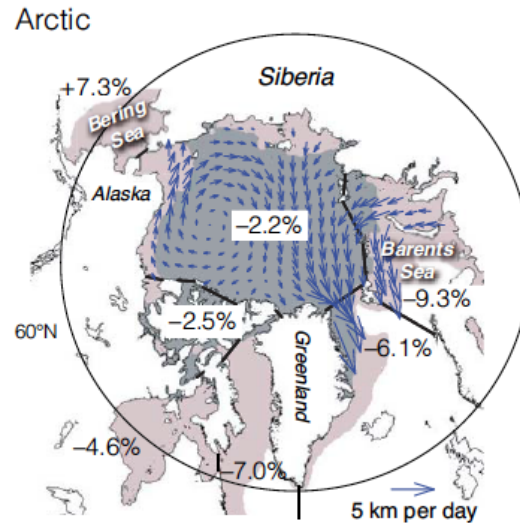
IPV influence: **Cold**(2001-2013) – **Warm**(1977-1998)
Temp. (Cowtan&Way) SLP (20CR)



Cl: 1 hPa

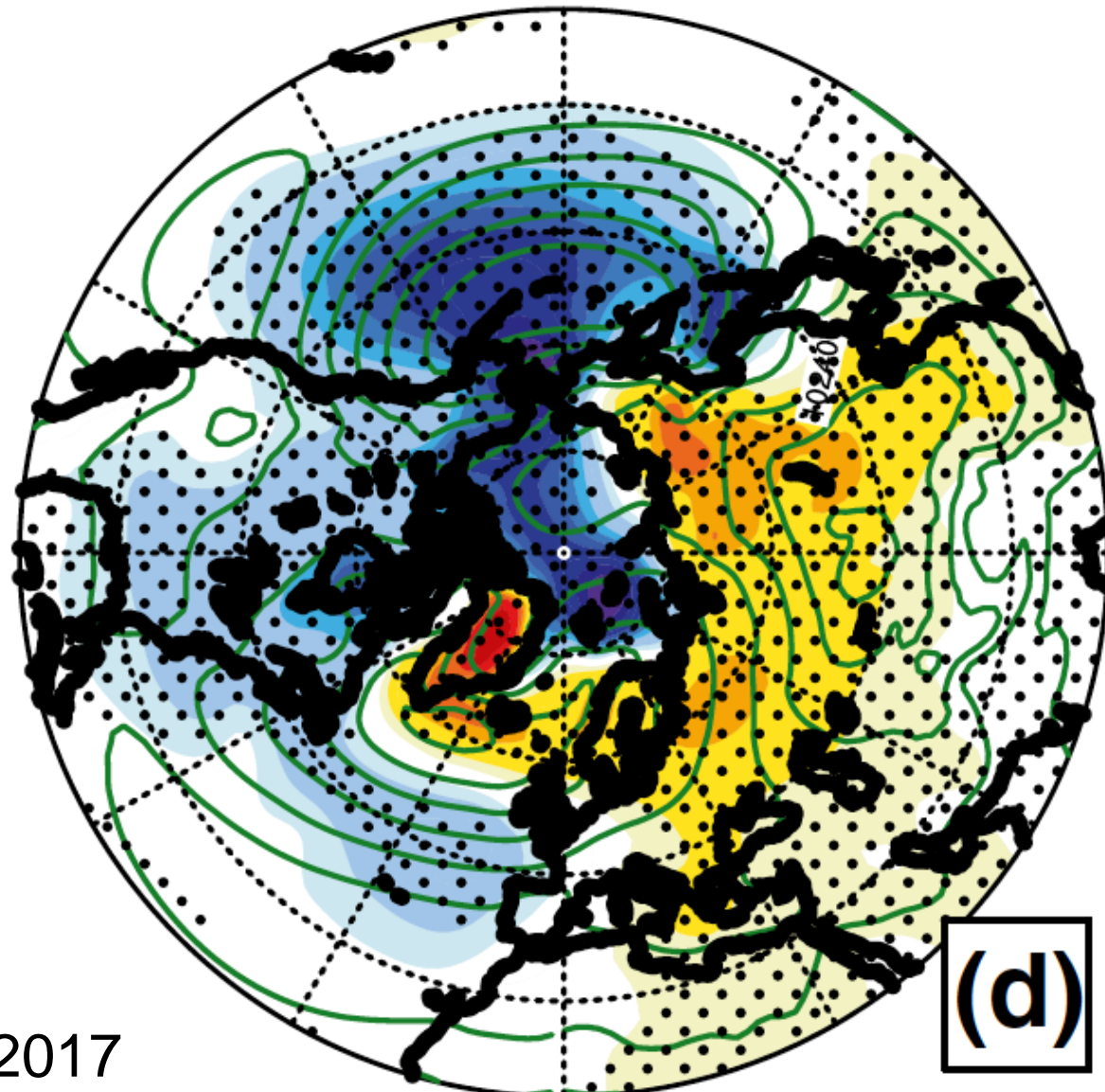


The Barents-Kara sea-ice hypothesis

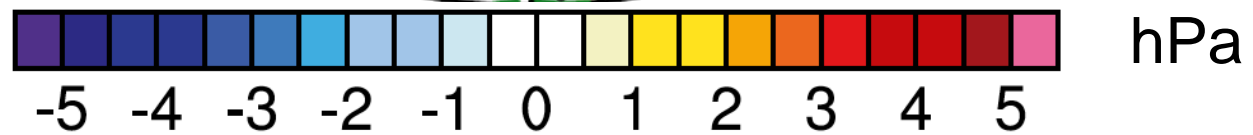


- Anomalous circulation pattern related to sea-ice decline ?
- Contribution of Barents-Kara seas sea-ice decline to thermodynamic component ?

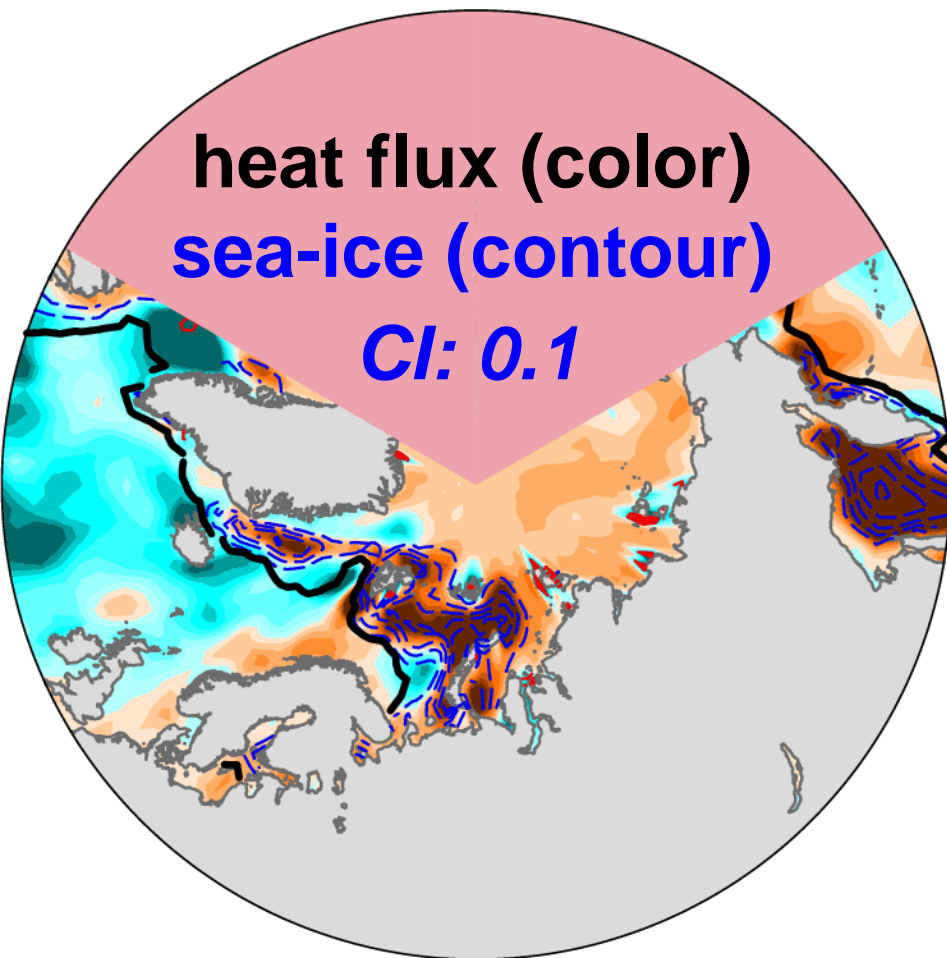
Winter coupled response to Arctic sea ice loss



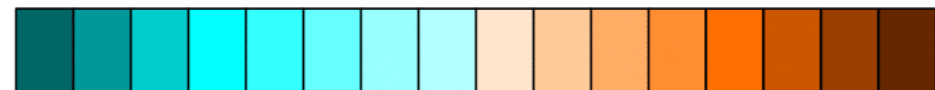
Oudar et al. 2017
ClimDyn



heat flux (color)
sea-ice (contour)
Cl: 0.1

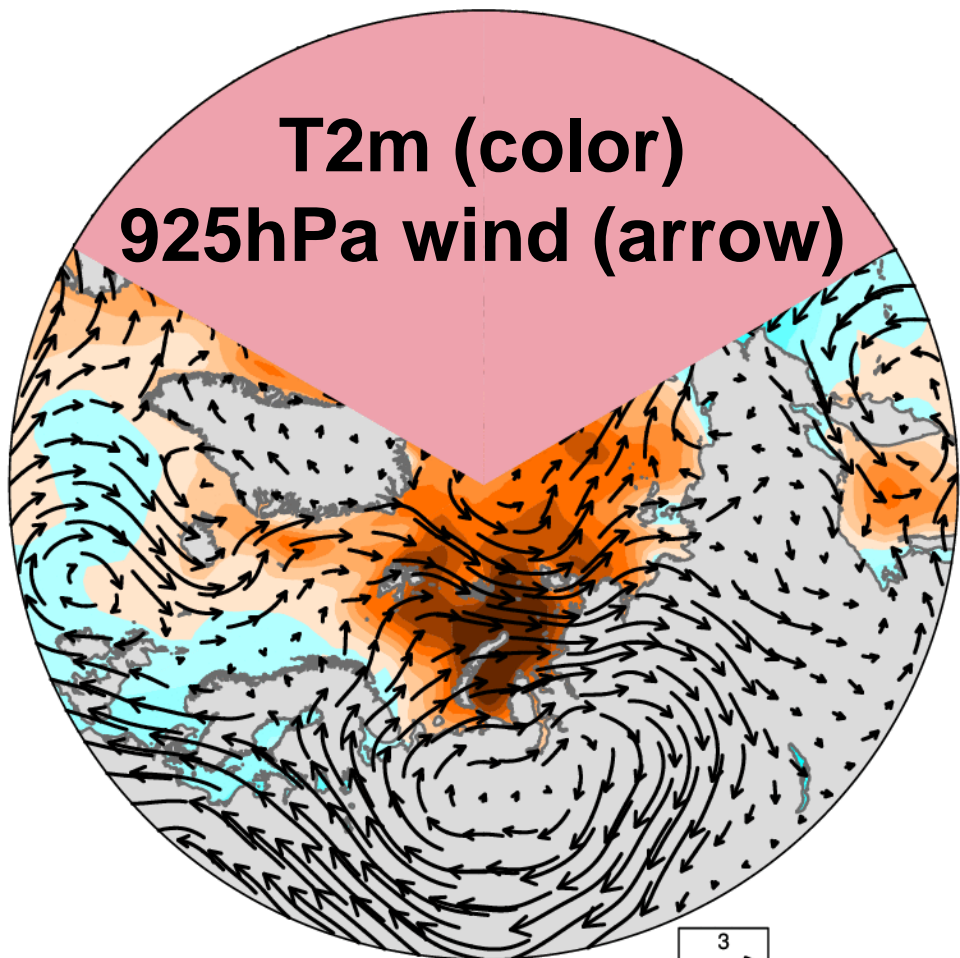


W.m⁻² / 13 years

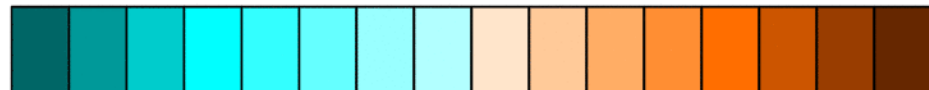


-35 -30 -25 -20 -15 -10 -5 0 5 10 15 20 25 30 35

T2m (color)
925hPa wind (arrow)



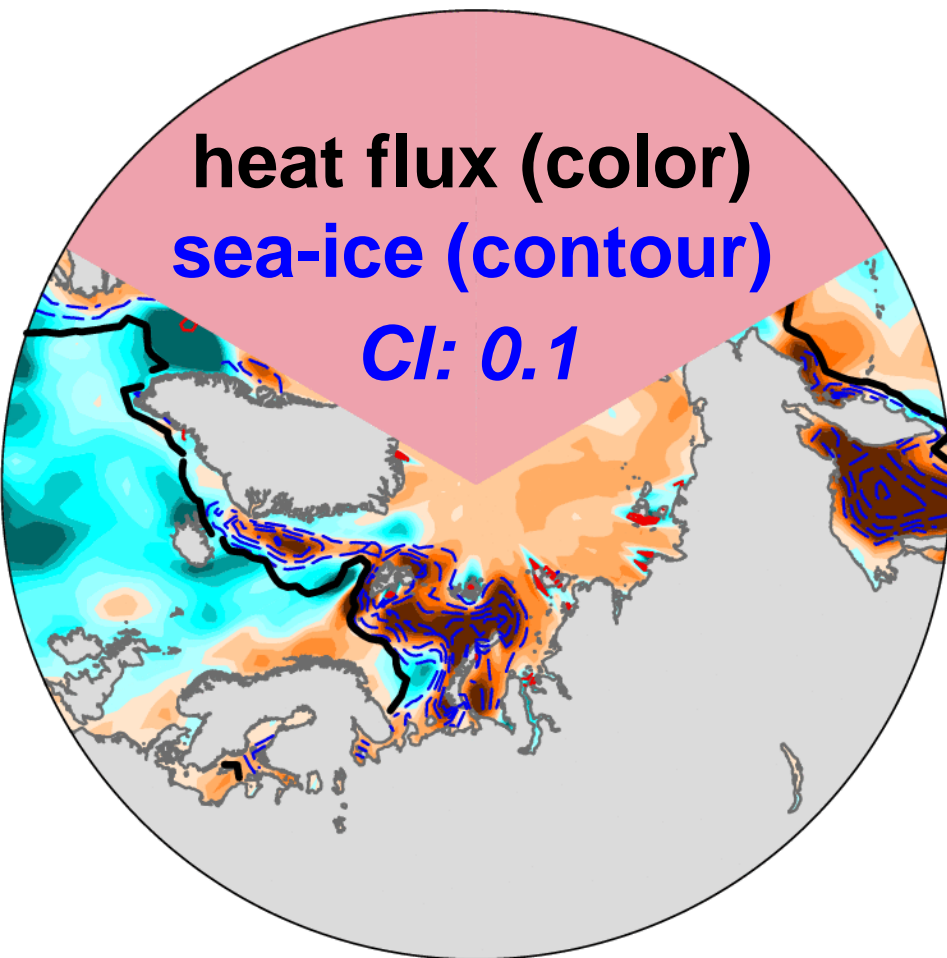
° C / 13 years



-7 -6 -5 -4 -3 -2 -1 0 1 2 3 4 5 6 7

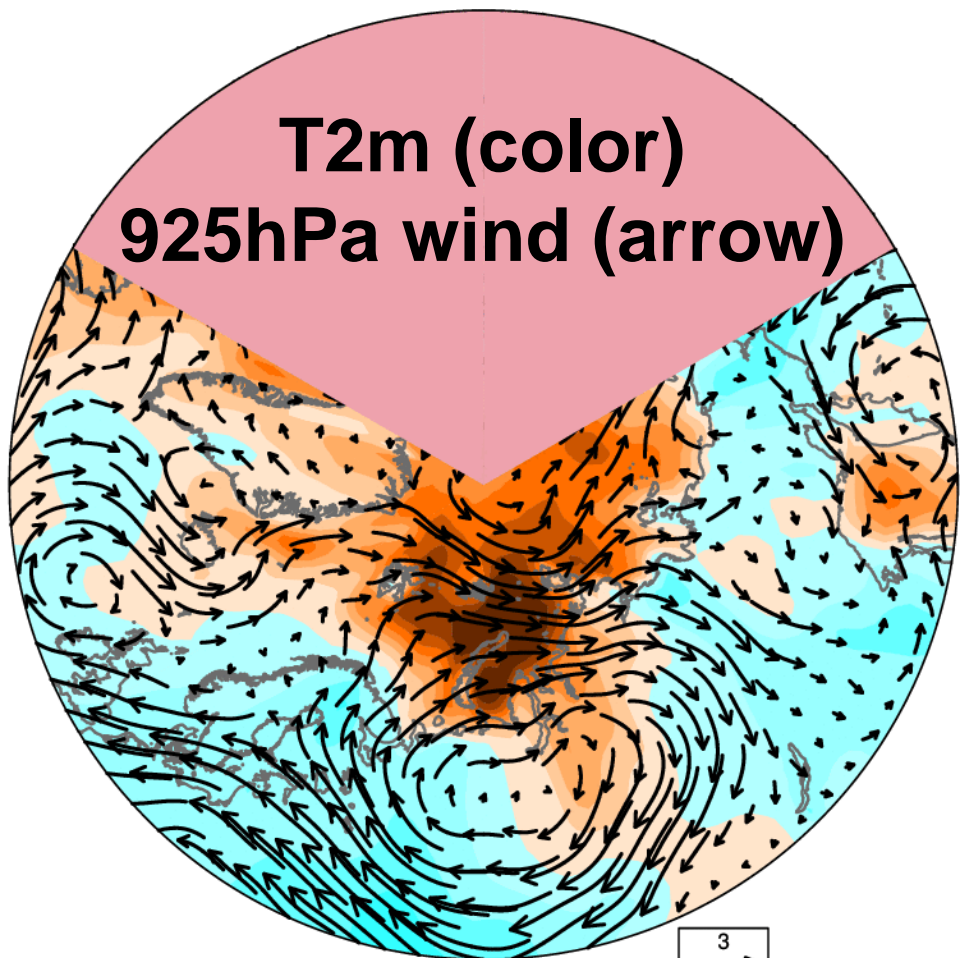


heat flux (color)
sea-ice (contour)
Cl: 0.1



W.m⁻² / 13 years

T2m (color)
925hPa wind (arrow)



° C / 13 years



-35 -30 -25 -20 -15 -10 -5 0 5 10 15 20 25 30 35

-7 -6 -5 -4 -3 -2 -1 0 1 2 3 4 5 6 7

Summary

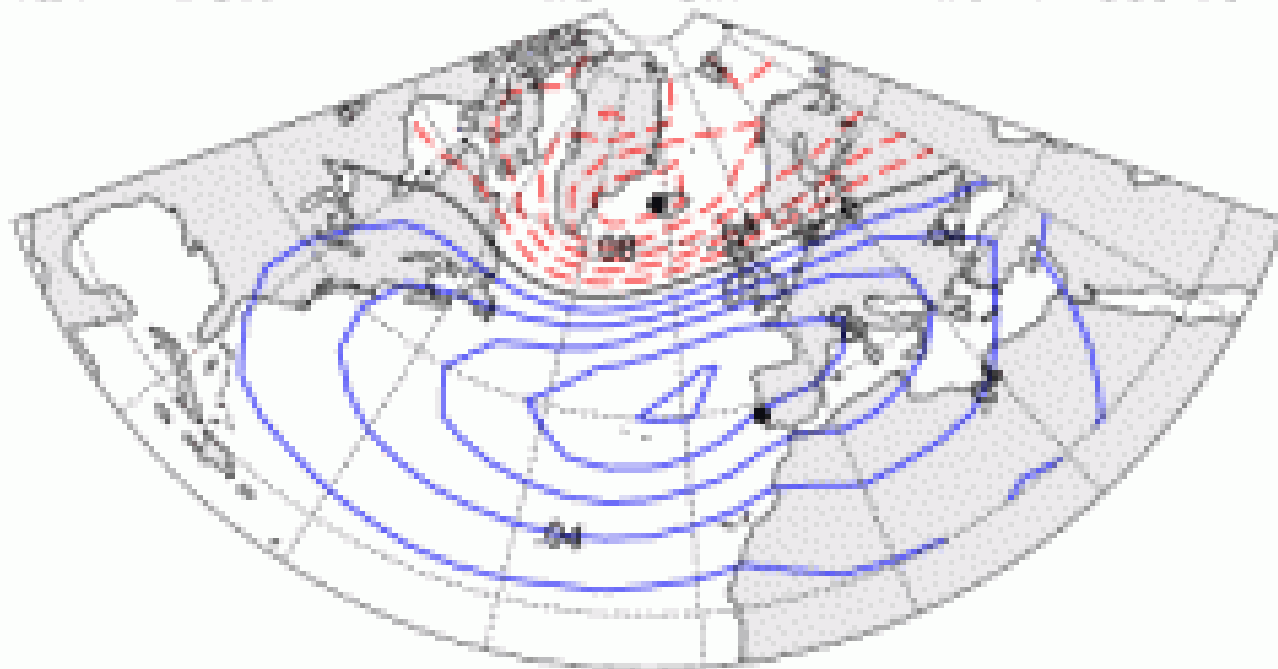
- Causes of a climate event: need to quantify contributing factors and uncertainties (as in D&A studies)
- Combine a purely data-driven approach with well-designed (multi-)model experiments
- Dynamical adjustment can be used to separate large-scale circulation effects from land/ocean surface forcing contributions
- Neither tropical forcing nor Barents-sea ice retreat seem to have significantly contributed to the anomalous Siberian High
- The early 2000s Eurasian Cooling is mostly due to a chaotic, non predictable atmospheric fluctuation

Thanks,
Questions ?

Var. = 42.8%

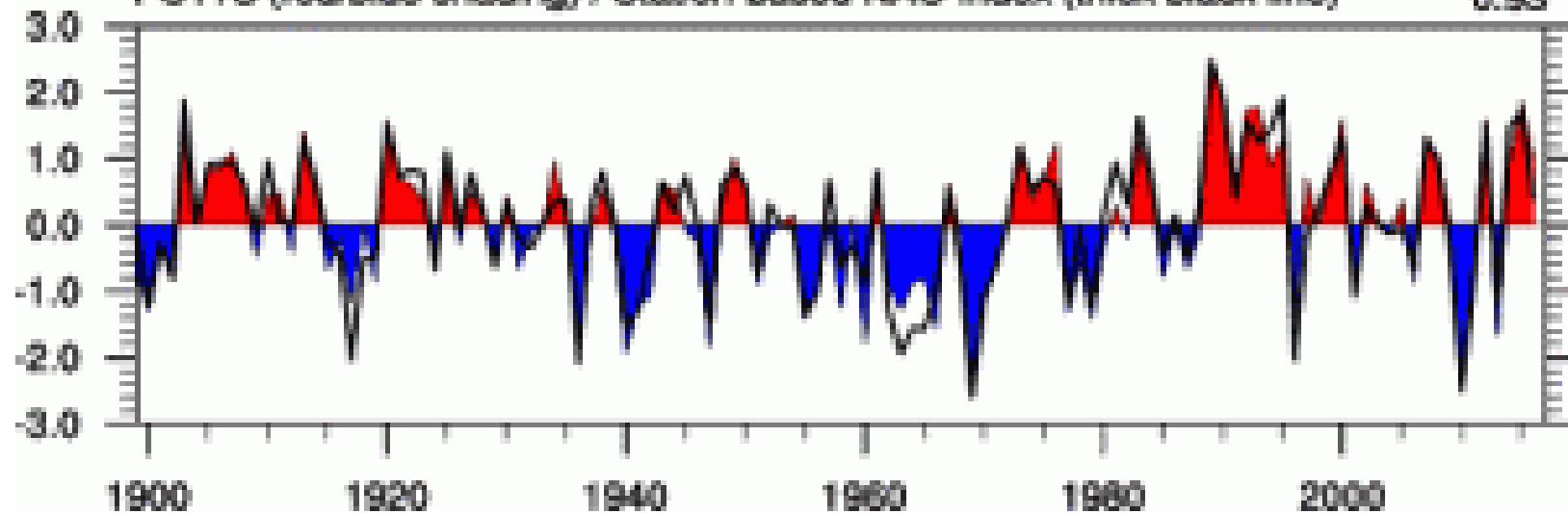
EOF1 SLP

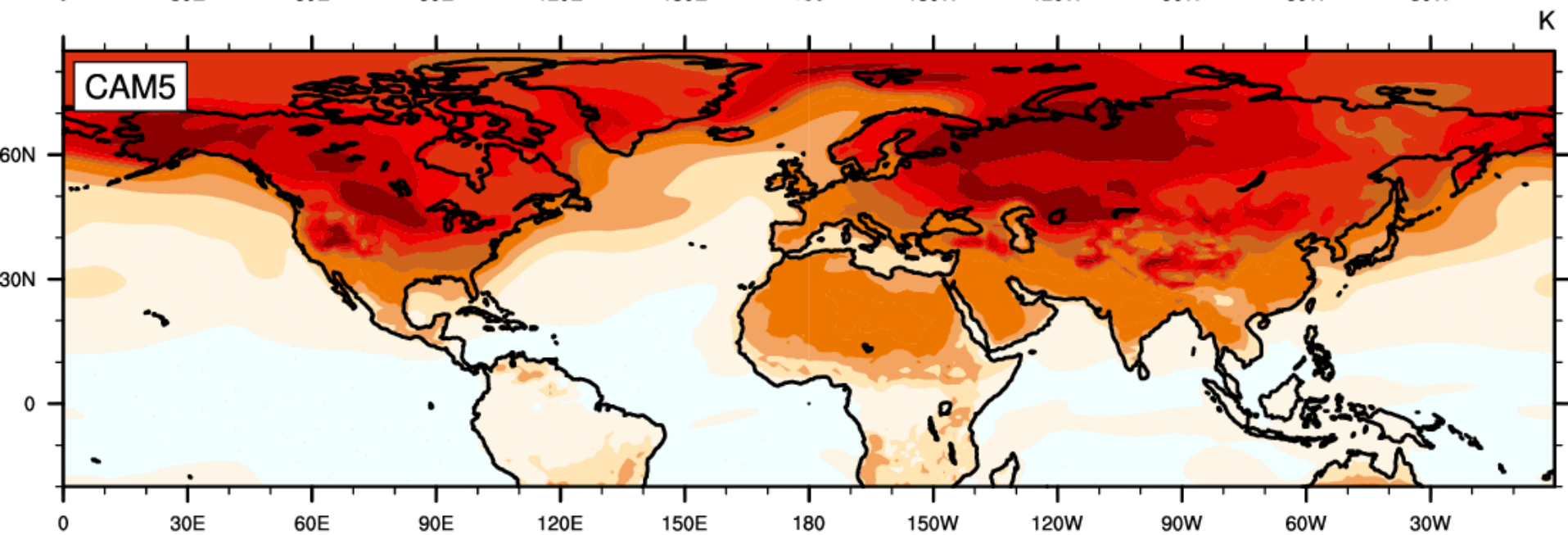
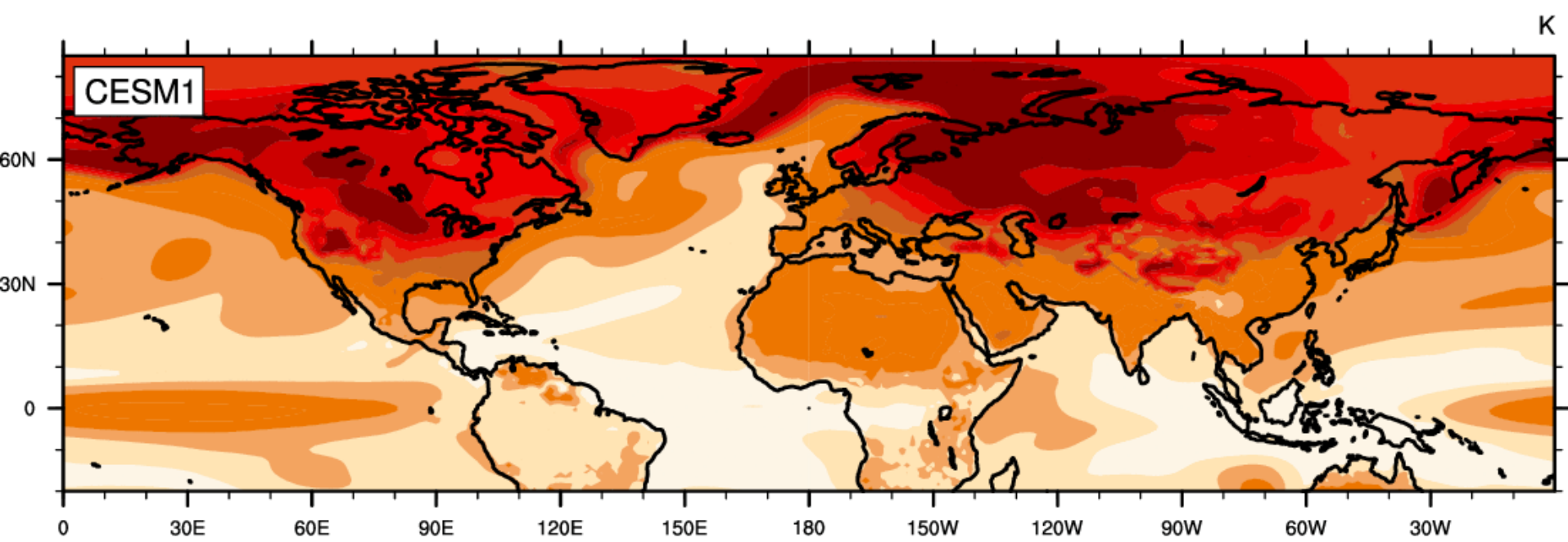
DJFM 1899-2016



PC1TS (red/blue shading) / Station-based NAO Index (thick black line)

0.93

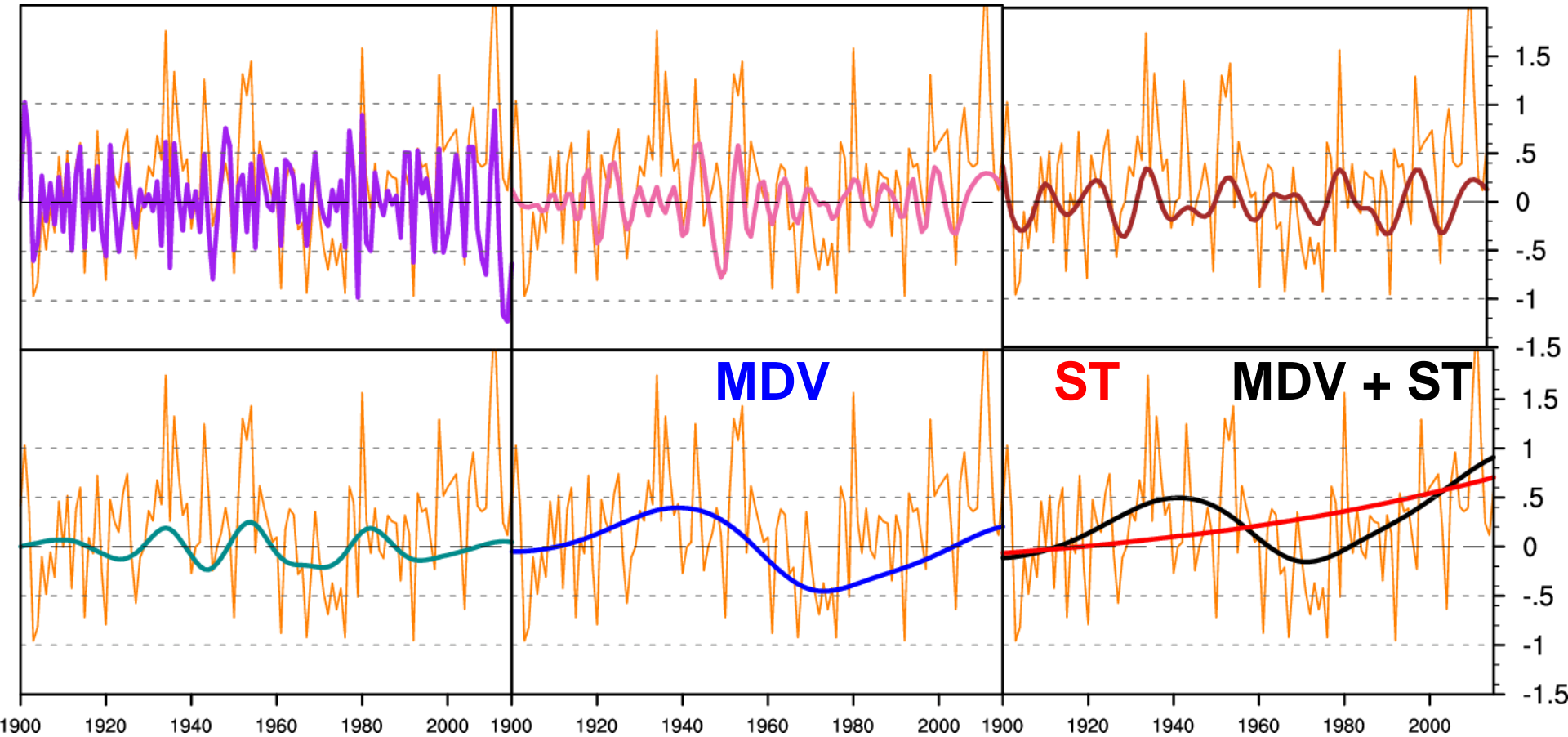


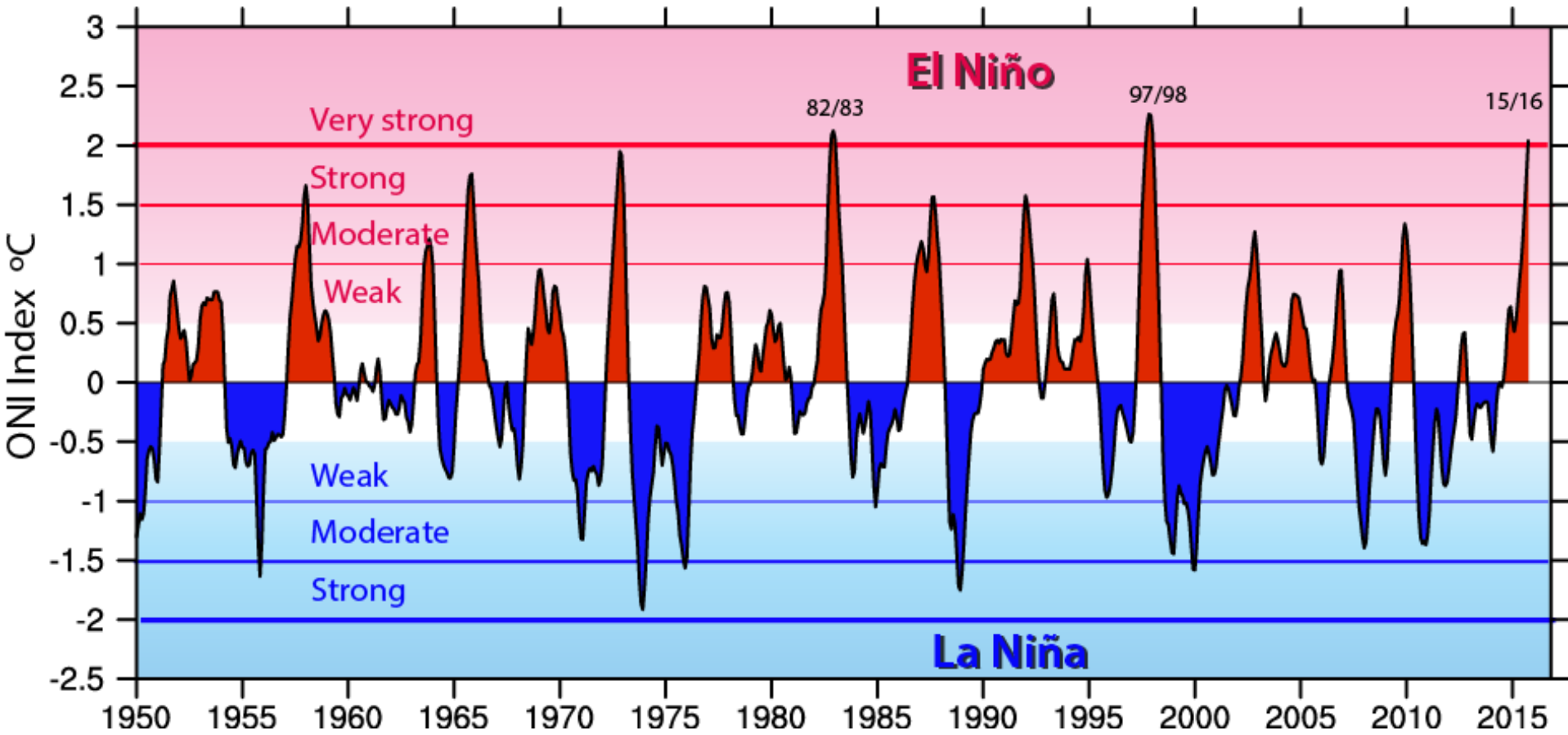


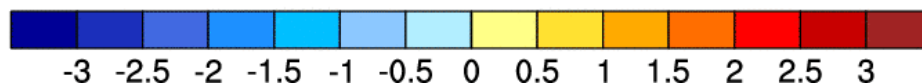
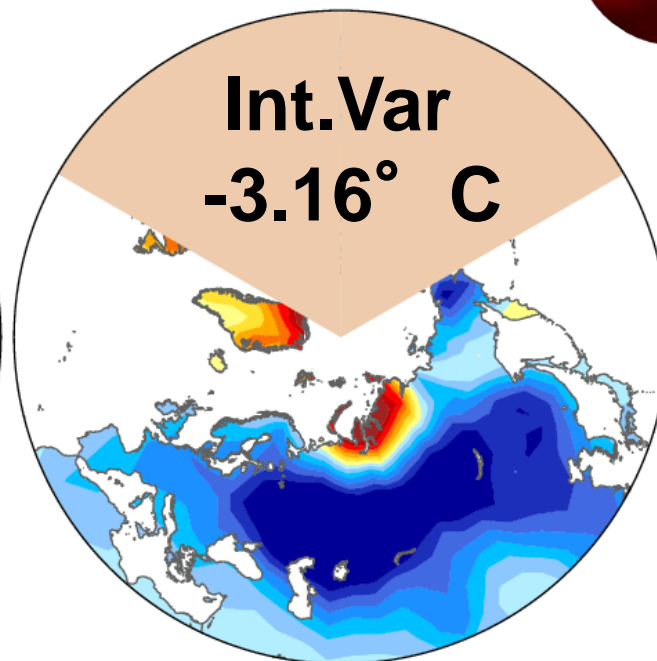
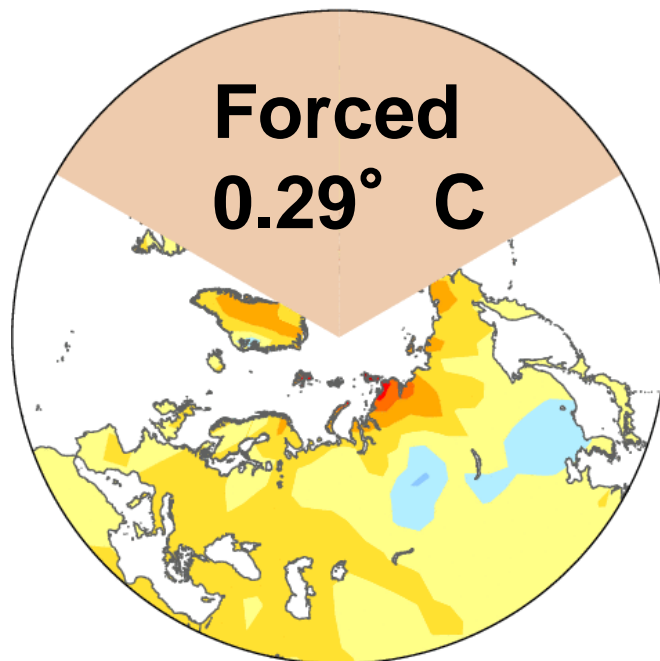
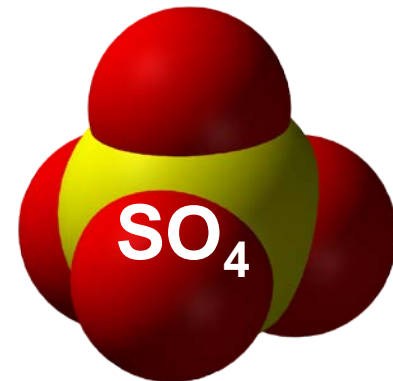
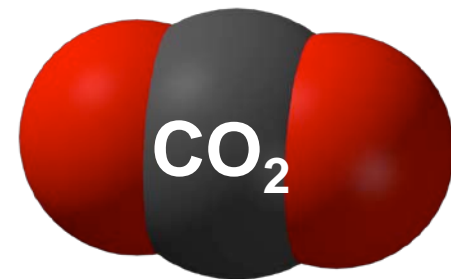
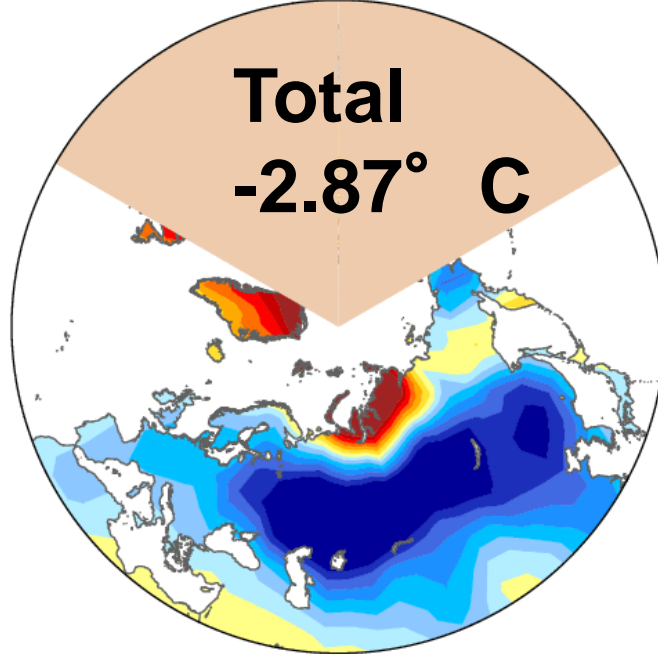
The External Forcing contribution

- › Can be estimated with models or observations
- › **Observations:** use a purely data-driven and non parametric approach, the *Ensemble Empirical Mode Decomposition* (EEMD) technique
- › EEMD decomposes any time series in a suite of amplitude-frequency modulated oscillatory components and a secular (non linear) trend
- › Apply EEMD to the Cowtan&Way data set

EEMD decomposition of temperature







° C / 13 years