



CESM2 release of CAM-SE (& CAM-SE-CSLAM)

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New SE dynamical core for CESM2

Science changes:

- Dry mass vertical coordinates
- Condensate loading in dynamical core (recommended as default)
- Separate physics grid and CSLAM options (not scientifically supported “yet”)
- Eulerian vertical advection no longer supported! Moist vertical coordinates not supported! (to keep code base simpler)

Other:

- Out-of-the-box CESM configurations for idealized setups (Held-Suarez, moist baroclinic wave with Kessler physics, terminator chemistry, ...)
- Performance upgrades from CISL (threading, more efficient SE transport)
- Dynamical core is no longer imported from HOMME (High-Order Methods Modeling Environment) ; code must go through CAM code review
- Cleaned up code base: trunk SE has ~61000 lines of code; new SE has ~39000 lines of code (further cleanup in progress ...)

CAM-SE : dry-mass eta

Consider a ‘moist’ η -coordinate system: The pressure is given by

$$p(\eta) = A(\eta)p_0 + B(\eta)ps,$$

where ps is ‘moist’ surface pressure.

In a floating η -coordinate system, $\dot{\eta} = 0$, the continuity equation for p can be written as

$$\frac{\partial}{\partial t} \left[\left(\frac{\partial p}{\partial \eta} \right) \right] + \nabla \cdot \left[\left(\frac{\partial p}{\partial \eta} \right) \vec{v} \right] = S^p,$$

where $S^p(q_v)$ is the source/sink term for pressure ($q_v \equiv$ specific humidity).

- This source/sink term:
 - makes the handling of tracers more complicated

An inert tracer will have source/sink terms (i.e. if there are moisture changes all “wet” mixing ratios must be changed accordingly)
 - makes it harder to move towards conserving a more comprehensive total energy
 - makes it harder to represent condensate loading in the dynamical core
- Complicates CSLAM-SE coupling in a moist atmosphere

$$\frac{\partial}{\partial t} \left[\left(\frac{\partial p}{\partial \eta} \right) \right] + \nabla \cdot \left[\left(\frac{\partial p}{\partial \eta} \right) \vec{v} \right] = S^p,$$

where $S^p(q_v)$ is the source/sink term for pressure ($q_v \equiv$ specific humidity).

CAM-SE : dry-mass eta

If one uses a dry mass vertical coordinate

$$p(\eta_d) = A(\eta_d)p_0 + B(\eta_d)ps_d,$$

where ps_d is dry surface pressure, then the continuity equation for pressure does not have sources/sinks

$$\frac{\partial}{\partial t} \left[\left(\frac{\partial p_d}{\partial \eta_d} \right) \right] + \nabla \cdot \left[\left(\frac{\partial p_d}{\partial \eta_d} \right) \vec{v} \right] = 0.$$

**Model levels do not move during
physics-dynamics coupling!**

CAM-SE : dry-mass eta

The η_d -coordinate atmospheric primitive equations assuming floating Lagrangian vertical coordinates [Starr, 1945; Lin, 2004] can be written as

$$\frac{\partial \vec{v}}{\partial t} + (\zeta + f) \hat{k} \times \vec{v} + \nabla_{\eta_d} \left(\frac{1}{2} \vec{v}^2 + \Phi \right) + \frac{1}{\rho} \nabla_{\eta_d} p = \nu \nabla^4 \vec{u}, \quad (16)$$

$$\frac{\partial T}{\partial t} + \vec{v} \cdot \nabla_{\eta_d} T - \frac{1}{c_p \rho} \omega = \nu \nabla_{\eta_d}^4 T, \quad (17)$$

$$\frac{\partial}{\partial t} \left(\frac{\partial p_d}{\partial \eta_d} m_i \right) + \nabla_{\eta_d} \cdot \left(\frac{\partial p_d}{\partial \eta_d} m_i \vec{v} \right) = \nu \nabla_{\eta_d}^4 (m_i), \quad i = d, v, cl, ci, \dots \quad (18)$$

where Φ is the geopotential height ($\Phi = g z$, where g is the gravitational constant), c_p is the specific heat constant for dry air, \hat{k} is the unit vector normal to the surface of the sphere, $\zeta = \hat{k} \cdot \nabla \times \vec{v}$ is vorticity, f Coriolis parameter, and $\omega = Dp/Dt$ is the pressure vertical velocity.

CAM-SE : dry-mass eta

The η_d -coordinate atmospheric primitive equations assuming floating Lagrangian vertical coordinates [Starr, 1945; Lin, 2004] can be written as

$$\frac{\partial \vec{v}}{\partial t} + (\zeta + f) \hat{k} \times \vec{v} + \nabla_{\eta_d} \left(\frac{1}{2} \vec{v}^2 + \Phi \right) + \frac{1}{\rho} \nabla_{\eta_d} p = \nu \nabla^4 \vec{u}, \quad (16)$$

$$\frac{\partial T}{\partial t} + \vec{v} \cdot \nabla_{\eta_d} T - \frac{1}{c_p \rho} \omega = \nu \nabla_{\eta_d}^4 T, \quad (17)$$

$$\frac{\partial}{\partial t} \left(\frac{\partial p_d}{\partial \eta_d} m_i \right) + \nabla_{\eta_d} \cdot \left(\frac{\partial p_d}{\partial \eta_d} m_i \vec{v} \right) = \nu \nabla_{\eta_d}^4 (m_i), \quad i = d, v, cl, ci, \dots \quad (18)$$

where Φ is the geopotential height ($\Phi = g z$, where g is the gravitational constant), c_p is the specific heat constant for dry air, \hat{k} is the unit vector normal to the surface of the sphere, $\zeta = \hat{k} \cdot \nabla \times \vec{v}$ is vorticity, f is the Coriolis parameter, and $\omega = Dp/Dt$ is the pressure vertical velocity.

$$\rho = \rho_d \left(\sum_i m_i \right), i = 'd', 'v', 'cl', 'ci'.$$

$$c_p = \frac{\sum_i [m_i c_{pi}]}{\sum_i m_i}$$

dry air 'd', water vapor 'v', cloud liquid 'cl' and cloud ice 'ci'

“Correct” Internal Energy

The total internal energy integrated over the entire atmosphere is given by

$$I_{tot} = \iiint \rho c_p T dz \cos(\theta) r d\lambda d\theta$$

Using the hydrostatic balance this equation can be written as

$$I_{tot} = \sum_i I_i,$$

where I_d is the total internal energy of dry air, I_v the total internal energy of water vapor, etc.:

$$\begin{aligned} I_d &= -\frac{1}{g} \iiint c_{pd} T \left(\frac{\partial p_d}{\partial \eta_d} \right) d\eta_d \cos(\theta) r d\lambda d\theta, \\ I_v &= -\frac{1}{g} \iiint c_{pv} m_v T \left(\frac{\partial p_d}{\partial \eta_d} \right) d\eta_d \cos(\theta) r d\lambda d\theta, \\ I_{cl} &= -\frac{1}{g} \iiint c_{pcl} m_{cl} T \left(\frac{\partial p_d}{\partial \eta_d} \right) d\eta_d \cos(\theta) r d\lambda d\theta, \\ I_{ci} &= -\frac{1}{g} \iiint c_{pci} m_{ci} T \left(\frac{\partial p_d}{\partial \eta_d} \right) d\eta_d \cos(\theta) r d\lambda d\theta. \end{aligned}$$

“Correct” Internal Energy

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Using the hydrostatic balance equation can be written as

$$I_{tot} = \sum_i I_i,$$

where I_d is the total internal energy of dry air, I_v the total internal energy of water vapor, etc.:

The internal energy in CAM physics is defined as

$$I_{tot}^{(CAM)} = -\frac{1}{g} \iiint c_{pd} T (1 + m_v) \left(\frac{\partial p_d}{\partial \eta_d} \right) d\eta_d \cos(\theta) r d\lambda d\theta$$

$$I_{cl} = -\frac{1}{g} \iiint c_{pcl} m_{cl} T \left(\frac{\partial p_d}{\partial \eta_d} \right) d\eta_d \cos(\theta) r d\lambda d\theta,$$

$$I_{ci} = -\frac{1}{g} \iiint c_{pci} m_{ci} T \left(\frac{\partial p_d}{\partial \eta_d} \right) d\eta_d \cos(\theta) r d\lambda d\theta.$$

“Correct” Internal Energy

The total internal energy integrated over the entire atmosphere is given by

$$I_{tot} = \iiint \rho c_p T dz \cos(\theta) r d\lambda d\theta$$

Enforcing the correct energy in CAM physics is NON-TRIVIAL:

- If a parameterization alters water vapor, cloud liquid, and/or cloud ice then internal energy (and kinetic energy) changes
- The assumption that pressure levels stay fixed during physics updates is violated unless we switch to dry pressure levels

$$I_v = -\frac{1}{g} \iiint c_{pv} m_v T \left(\frac{\partial p_d}{\partial \eta_d} \right) d\eta_d \cos(\theta) r d\lambda d\theta,$$
$$I_{cl} = -\frac{1}{g} \iiint c_{pcl} m_{cl} T \left(\frac{\partial p_d}{\partial \eta_d} \right) d\eta_d \cos(\theta) r d\lambda d\theta,$$
$$I_{ci} = -\frac{1}{g} \iiint c_{pci} m_{ci} T \left(\frac{\partial p_d}{\partial \eta_d} \right) d\eta_d \cos(\theta) r d\lambda d\theta.$$

Two SE configurations

controlled with namelist: `qsize_condensate_loading = 1,3`

1. Set $c_p = c_{pd}$ and $\rho = \rho_d + \rho_v$
2. Use ‘correct’ c_p and $\rho = \rho_d + \rho_v + \rho_{ci} + \rho_{cl}$

In dynamics-physics coupling we pass $\Delta p = \Delta p_d(1 + m_v)$ to remain consistent with the physics definition of total energy.

Both configurations have pros and cons:

- 1a. Only difference between #1 and trunk is vertical coordinate
- 1b. The continuous equations of motion conserve the same “wrong” energy as CAM physics.
- 1c. The adiabatic momentum equations and thermodynamic equations do not “feel” the condensates
- 2a. The continuous equations of motion conserve the “correct” energy but CAM physics will (through the energy fixer) enforce the “wrong” energy
- 2b. The adiabatic momentum equations and thermodynamic equations “feel” the condensates (a.k.a. condensate loading; may be significant at higher resolution)

Energy budgets in CAM-FV (0.9x0.9)

CAM5.3 physics; 10 year run

pEFIX = dE/dt energy fixer
pDMEA = dE/dt dme_adjust

(pBP-pBF) : -0.8103 W/m²
(pAM-pAP) : 0.2636 W/m²

Energy fixer fixes dme_adjust (pDMEA), lack of energy conservation in adiabatic dynamical core (dADIA) and energy lost/gained in physics-dynamics coupling (dPDC):

$$-pEFIX = pDMEA + dADIA + dPDC$$

CAM-FV uses updated state (no “dripping” of tendencies) from physics so dPDC=0

⇒ dADIA = dE/dt adiabatic dynamical core = -pEFIX - pDMEA : -1.0739 W/m²

Aside:

At 2 degree horizontal resolution dE/dt adiabatic dynamical : -1.2738 W/m²

Energy budgets in CAM-SE configuration 1

CAM5.9999 physics; 6 year run; qsize_condensate_loading = 1
http://webext.cgd.ucar.edu/FCLIMO/f.e20.F2000_DEV.ne30_ne30.physgrid25_cam5_4_96_se_qsize_1/atm/

pEFIX = dE/dt energy fixer
pDMEA = dE/dt dme_adjust

(pBP-pBF) : -0.1913 W/m²
(pAM-pAP) : 0.3064 W/m²

Energy fixer fixes dme_adjust (pDMEA), lack of energy conservation in adiabatic dynamical core (dADIA) and energy lost/gained in physics-dynamics coupling (dPDC):

$$-pEFIX = pDMEA + dADIA + dPDC$$

dADIA = dE/dt adiabatic dynamical core = -0.0732 W/m²
(-0.1604 W/m² vertical remapping, 0.0872 W/m² Lagrangian dyn, hypervis V added as heating = 0.7110 W/m²)

dPDC = dE/dt physics-dynamics coupling (ftype=0) = -0.0419 W/m²

dADIA (SE configuration 1) = -0.0732 W/m² << dADIA (FV) = -1.0739 W/m²
pDMEA is about the same for CAM-FV and CAM-SE

Energy budgets in CAM-SE configuration 2

CAM5.9999 physics; 6 year run; qsize_condensate_loading = 3
http://webext.cgd.ucar.edu/FCLIMO/f.e20.F2000_DEV.ne30_ne30.physgrid25_cam5_4_96_se_qsize_3/atm/

pEFIX = dE/dt energy fixer
pDMEA = dE/dt dme_adjust

(pBP-pBF) : -0.7070 W/m²
(pAM-pAP) : 0.3102 W/m²

Energy computations in dynamical use “correct” energy formula
Energy computations in physics use “wrong” energy formula

=> We can not mix computations done in physics and dynamics but

dADIA = dE/dt adiabatic dynamical core = -0.0928 W/m²
pEFIX configuration 2 – pEFIX configuration 1 = 0.5257 W/m²
i.e. the inconsistency in energy formula is ca. 0.5 W/m² consistent with
Mark Taylors findings.

Even with this inconsistency the energy fixer fixes less than for CAM-FV
=> I recommend configuration 2 as default in CESM2



CAM-SE-CSLAM without moisture

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Monthly Weather Review

CAM-SE-CSLAM: Consistent Coupling of a Conservative Semi-Lagrangian Finite-Volume Method with Spectral Element Dynamics

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CAM-SE-CSLAM without moisture

MARCH 2017

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(b) Local iteration problem

CAM-SE-C
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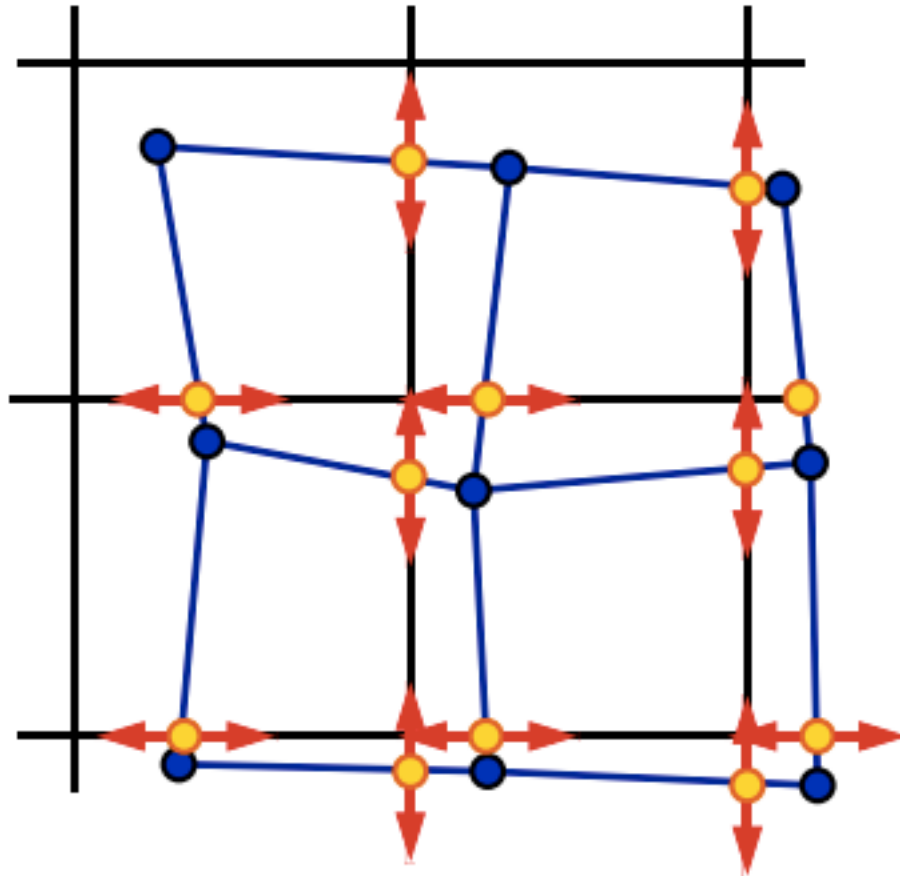
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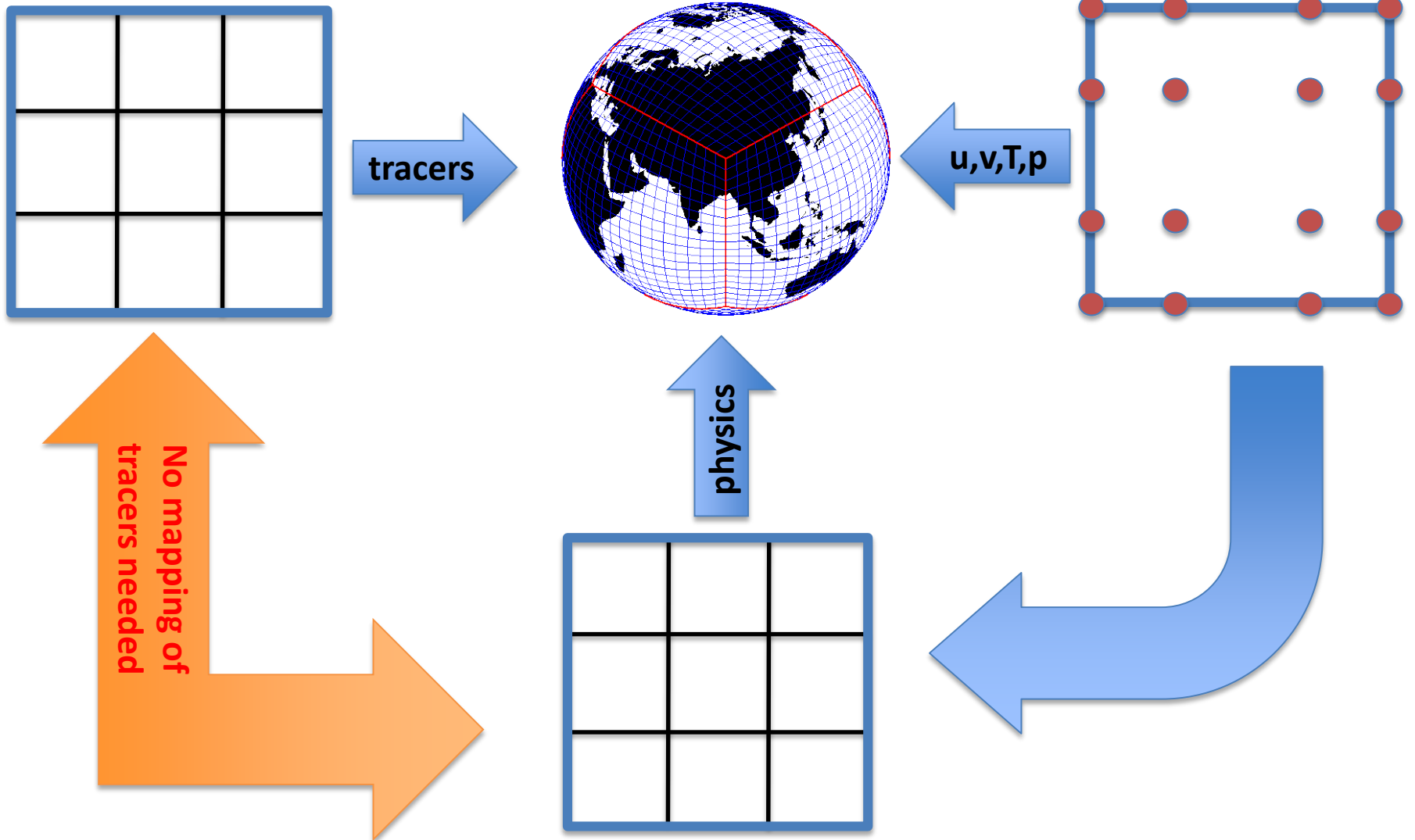
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CAM-SE-CSLAM with moisture

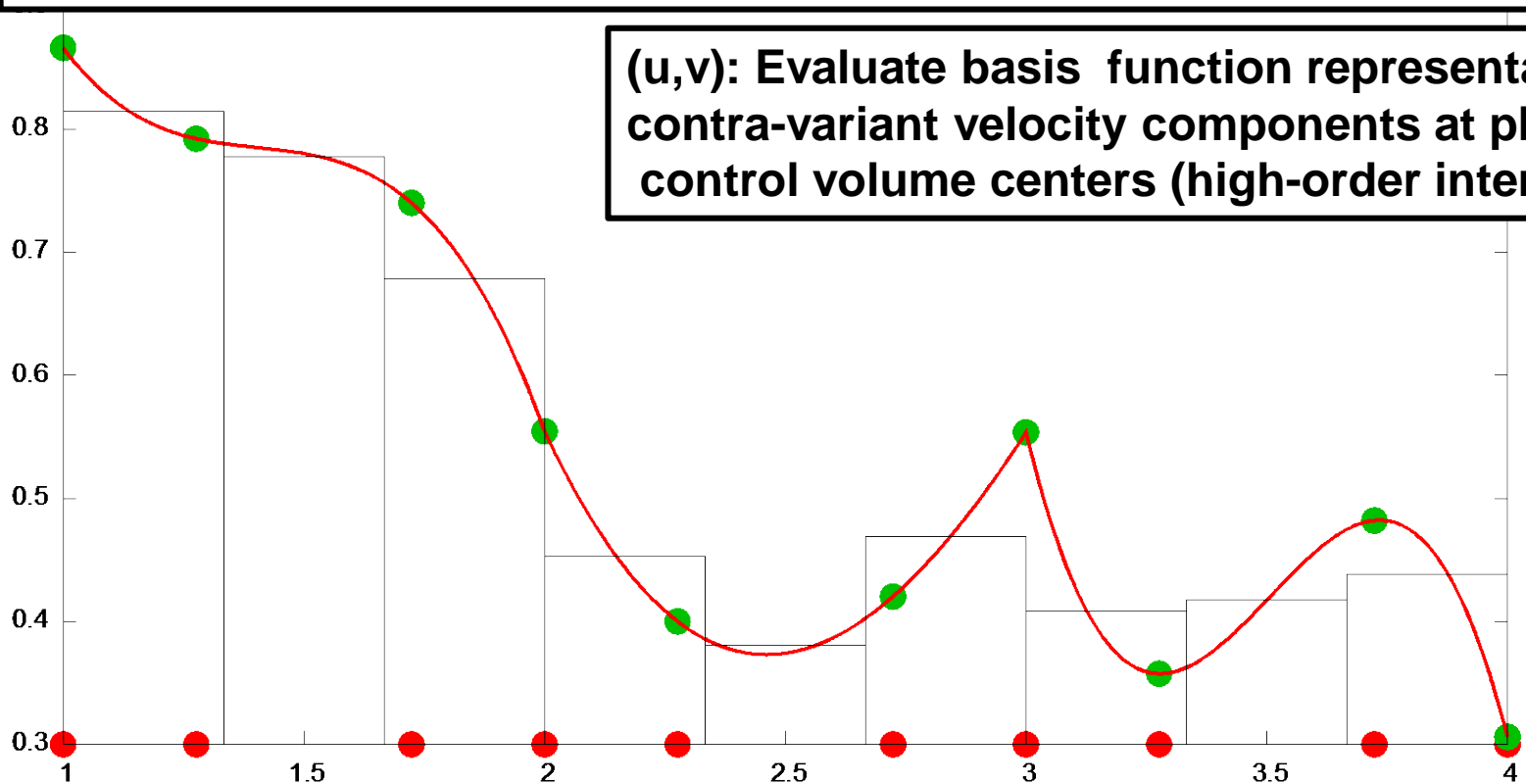
“This is where the fun begins!” – Staniforth et al. (2006)



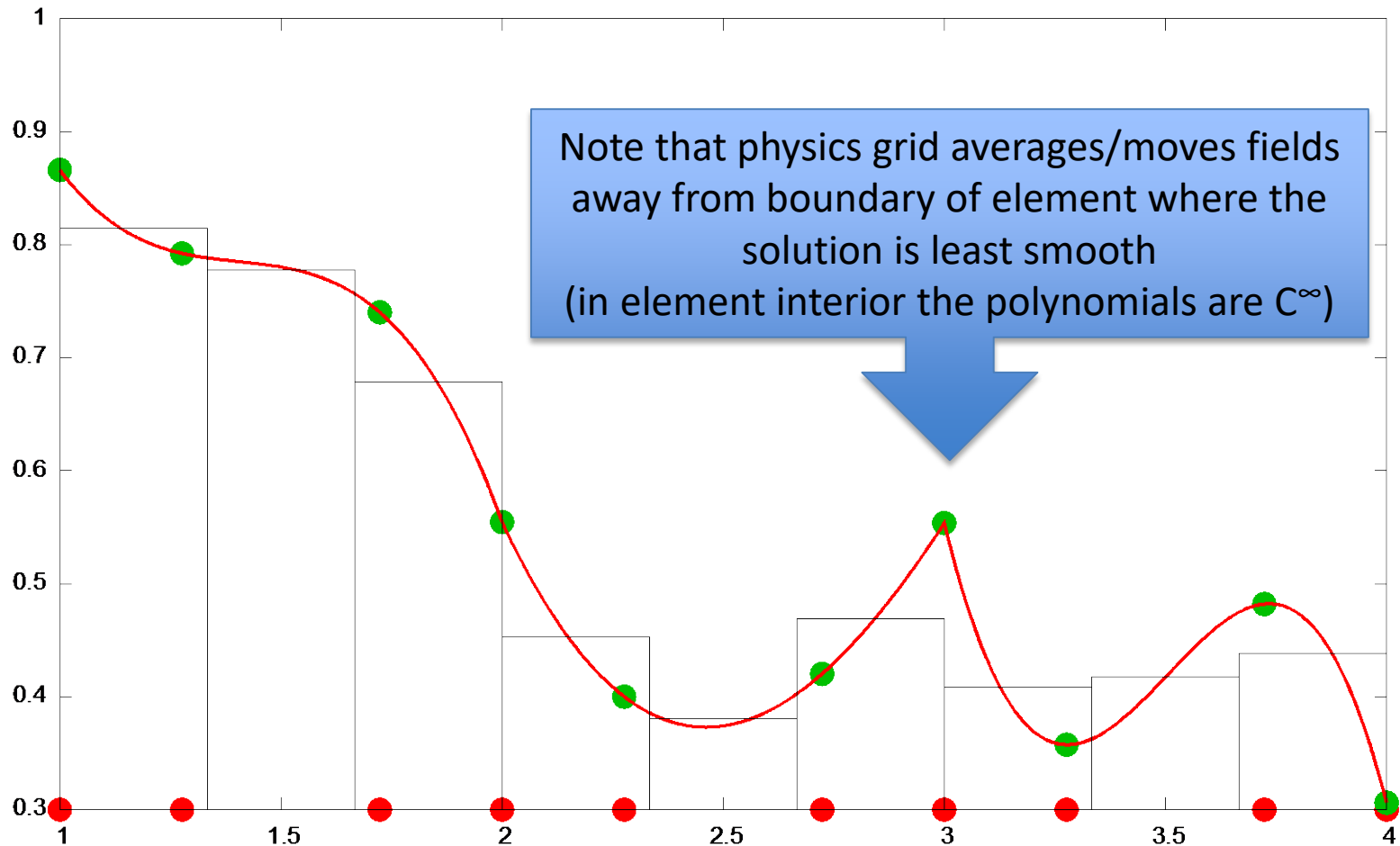
Mapping u, v, T, ω from dynamics grids (GLL) to finite-volume (CSLAM) grid

Temperature (& omega): Integrate basis function representation of dp^*T over physics grid control volumes (high-order remapping; conserves internal energy)

(u, v): Evaluate basis function representation of contra-variant velocity components at physics control volume centers (high-order interpolation)



Mapping u,v, T, omega from dynamics grids (GLL) to finite-volume (CSLAM) grid



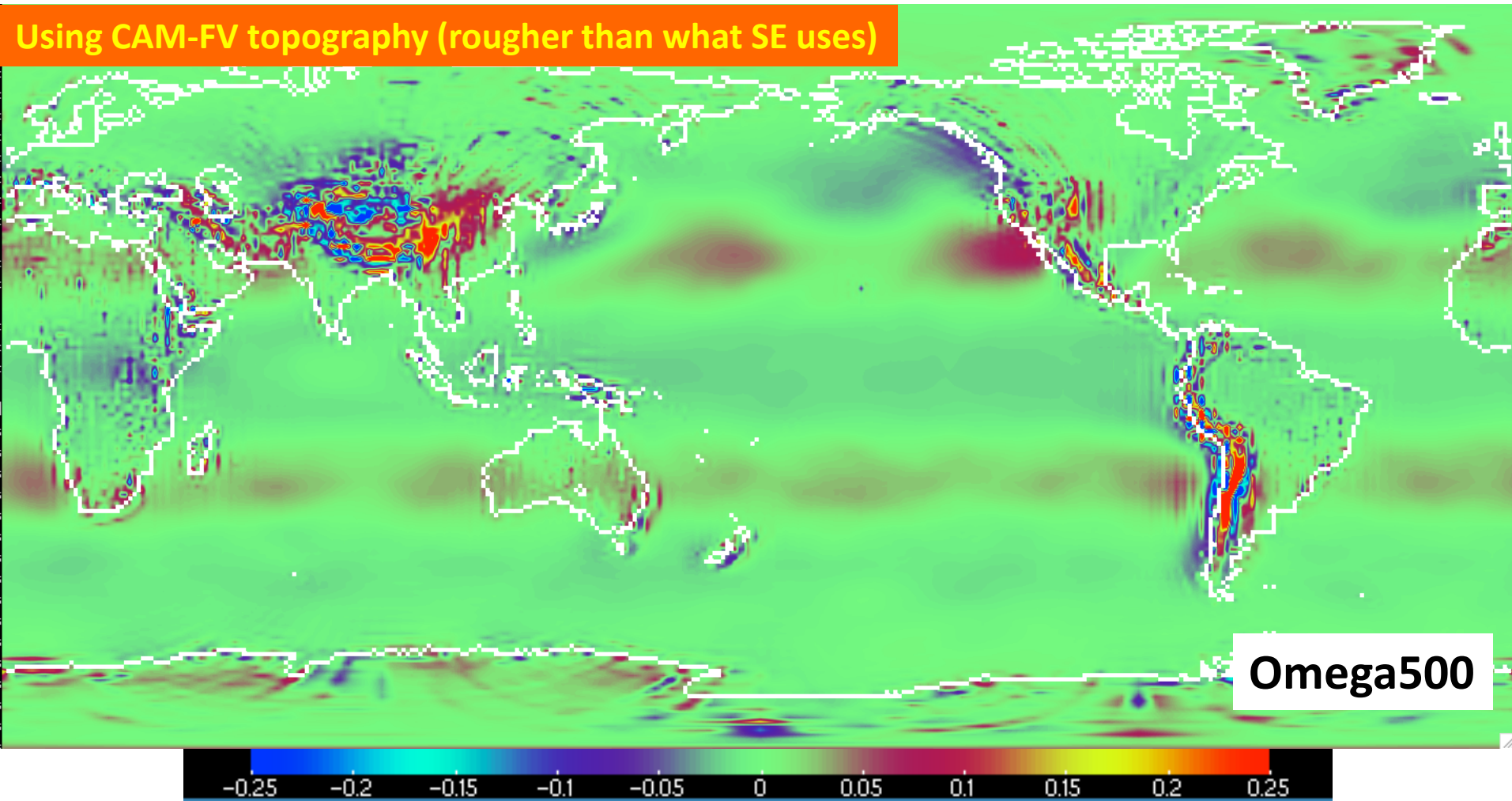
CAM-SE with “rougher” topography

Held-Suarez forcing with real-world topography (6 months spin-up; 2 years and 9 months average)

Note: dry test so no moist physics feedback

bnd_topo = '/home/pel/run-scripts/topo/ne30np4_nc3000_Nsw042_Nrs008_Co060_Fi001_ZR_test_vX_111416.nc'

Using CAM-FV topography (rougher than what SE uses)



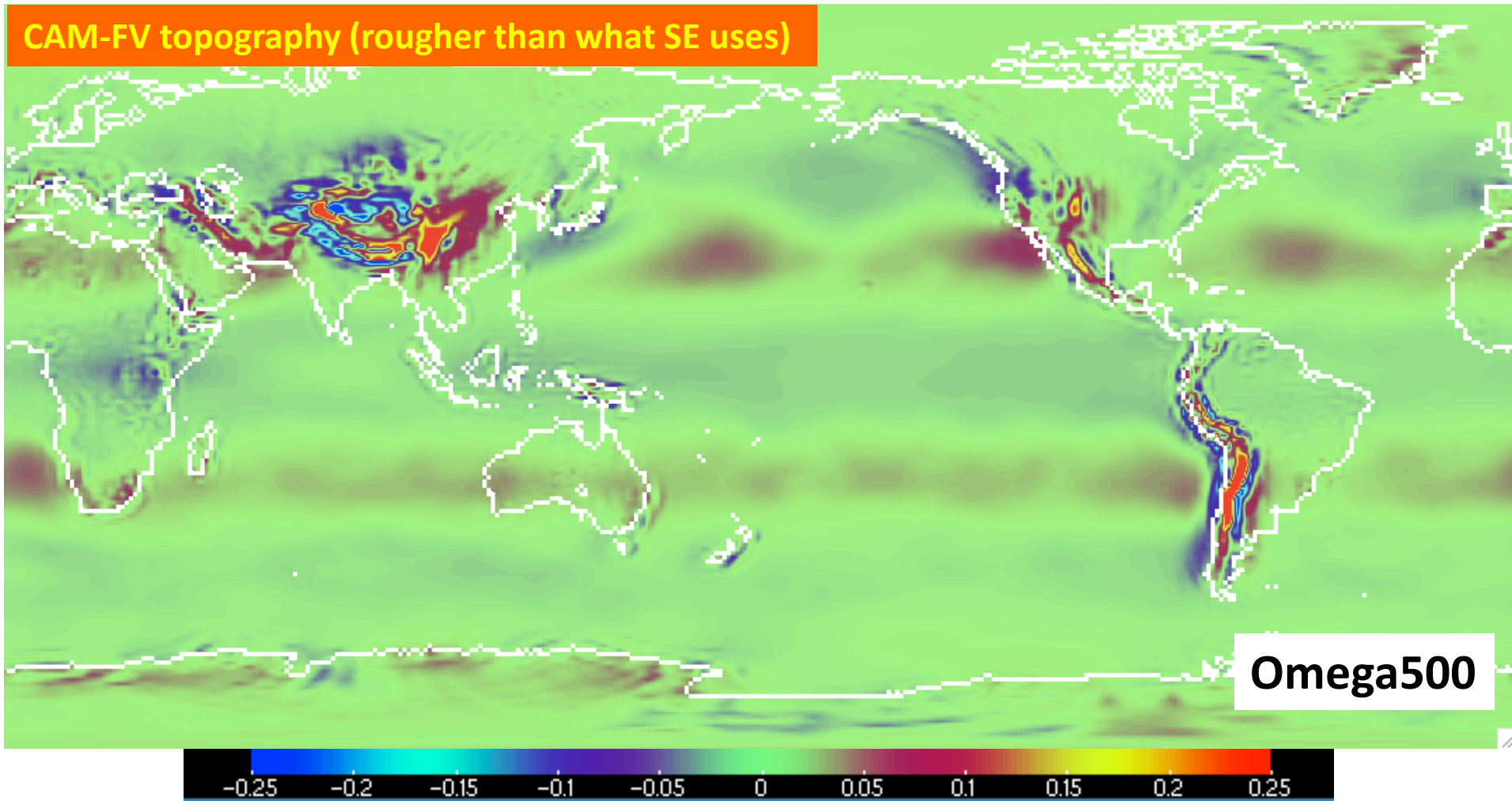
CAM-SE-CSLAM

Held-Suarez forcing with real-world topography (6 months spin-up; 2 years and 9 months average)

Note: dry test so no moist physics feedback

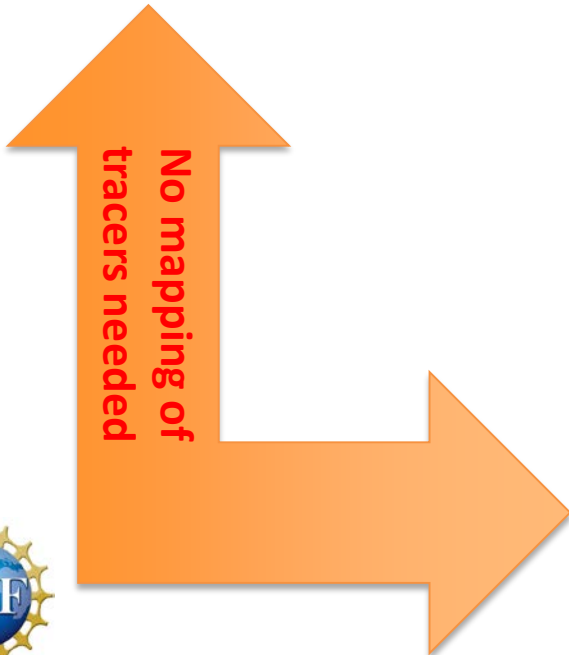
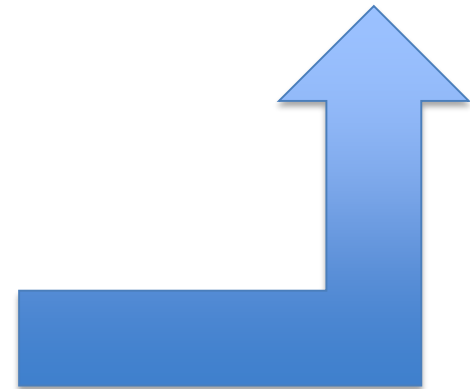
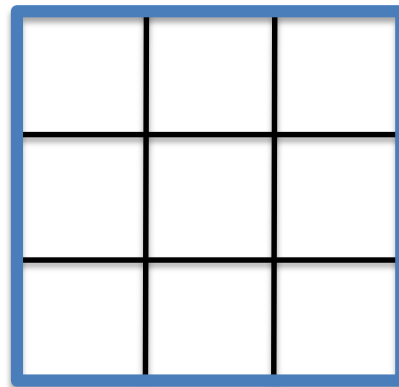
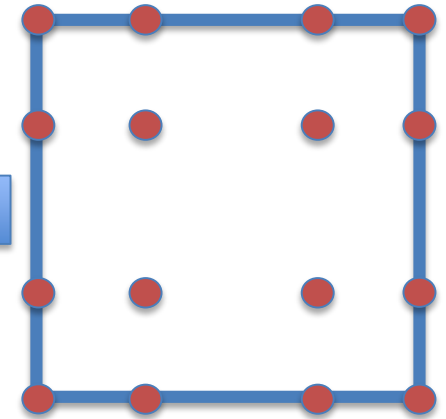
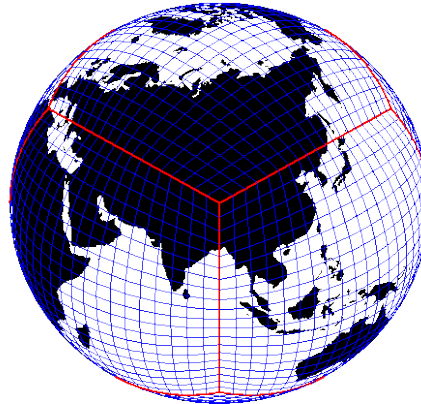
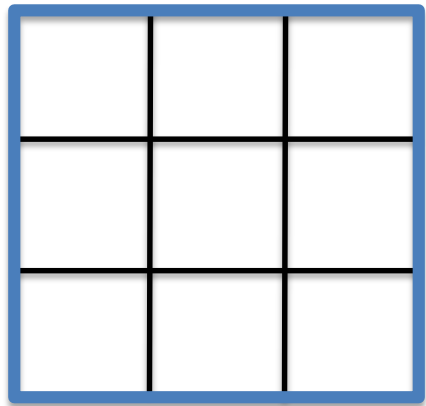
bnd_topo = '/home/pel/run-scripts/topo/ne30np4_nc3000_Nsw042_Nrs008_Co060_Fi001_ZR_test_vX_111416.nc'

CAM-FV topography (rougher than what SE uses)



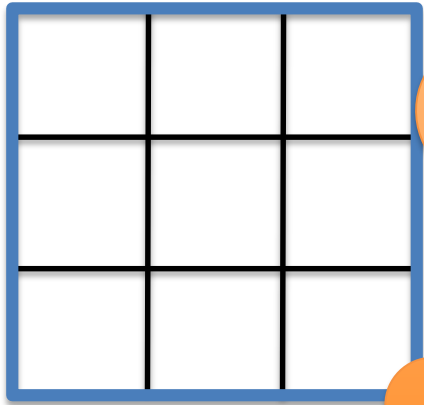


CAM-SE-CSLAM configuration



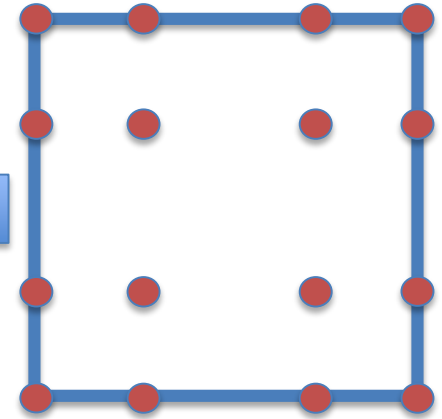


CAM-SE-CSLAM configuration

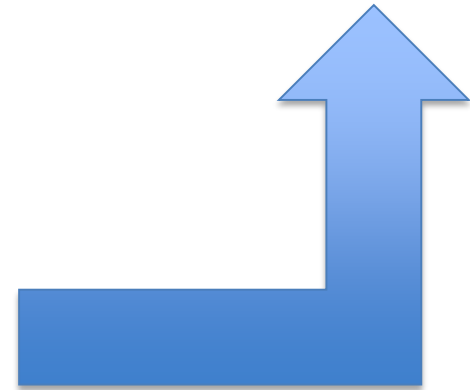
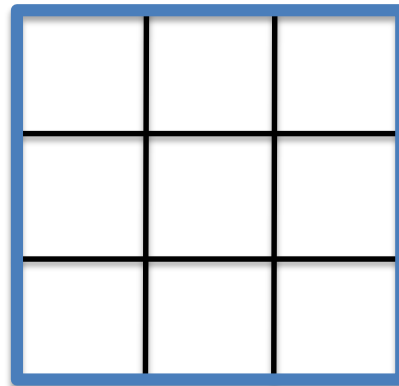


“Tendencies from physics parameterizations are low order anyway so I can just use low order mapping ...”

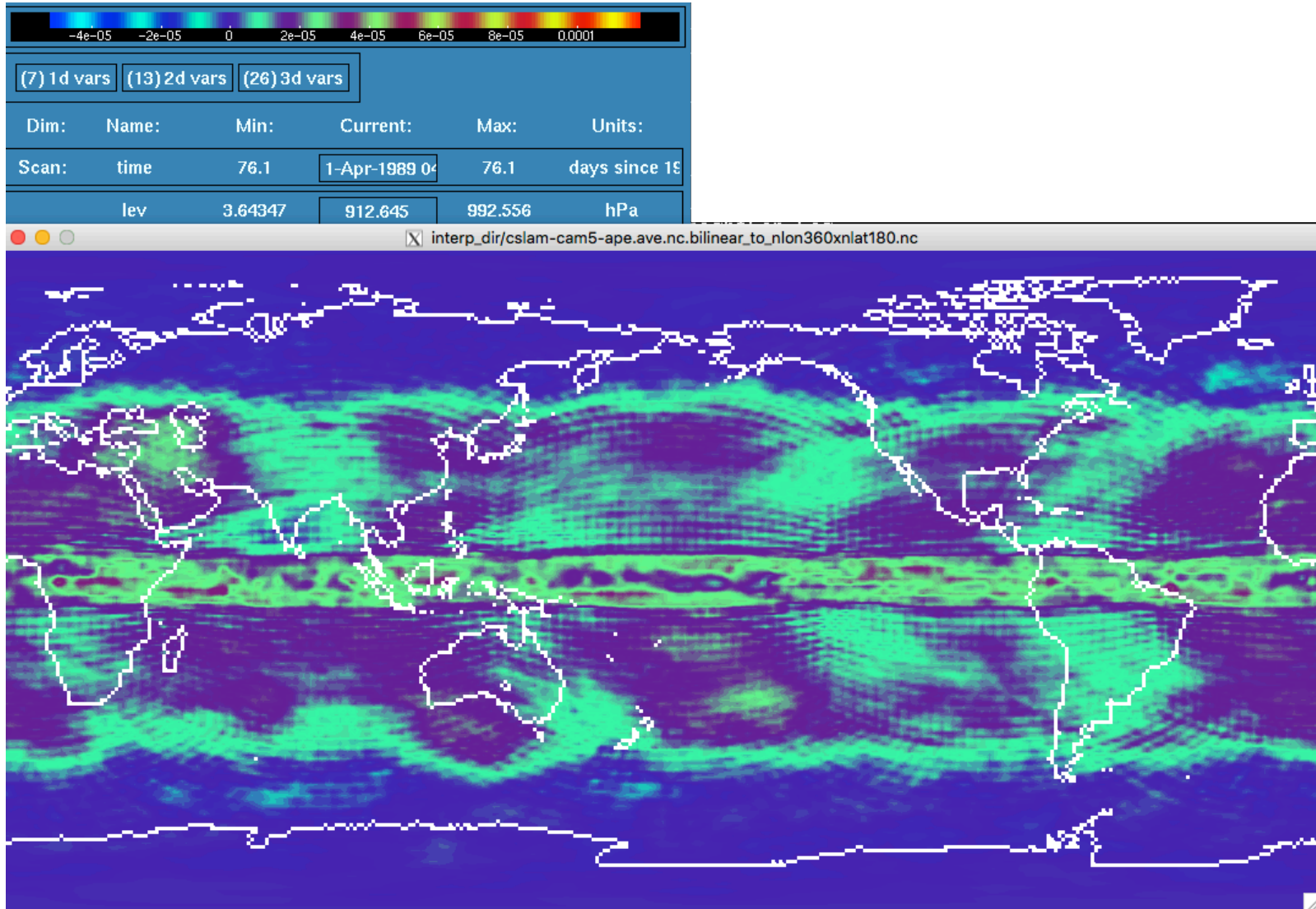
u, v, T, p



physics

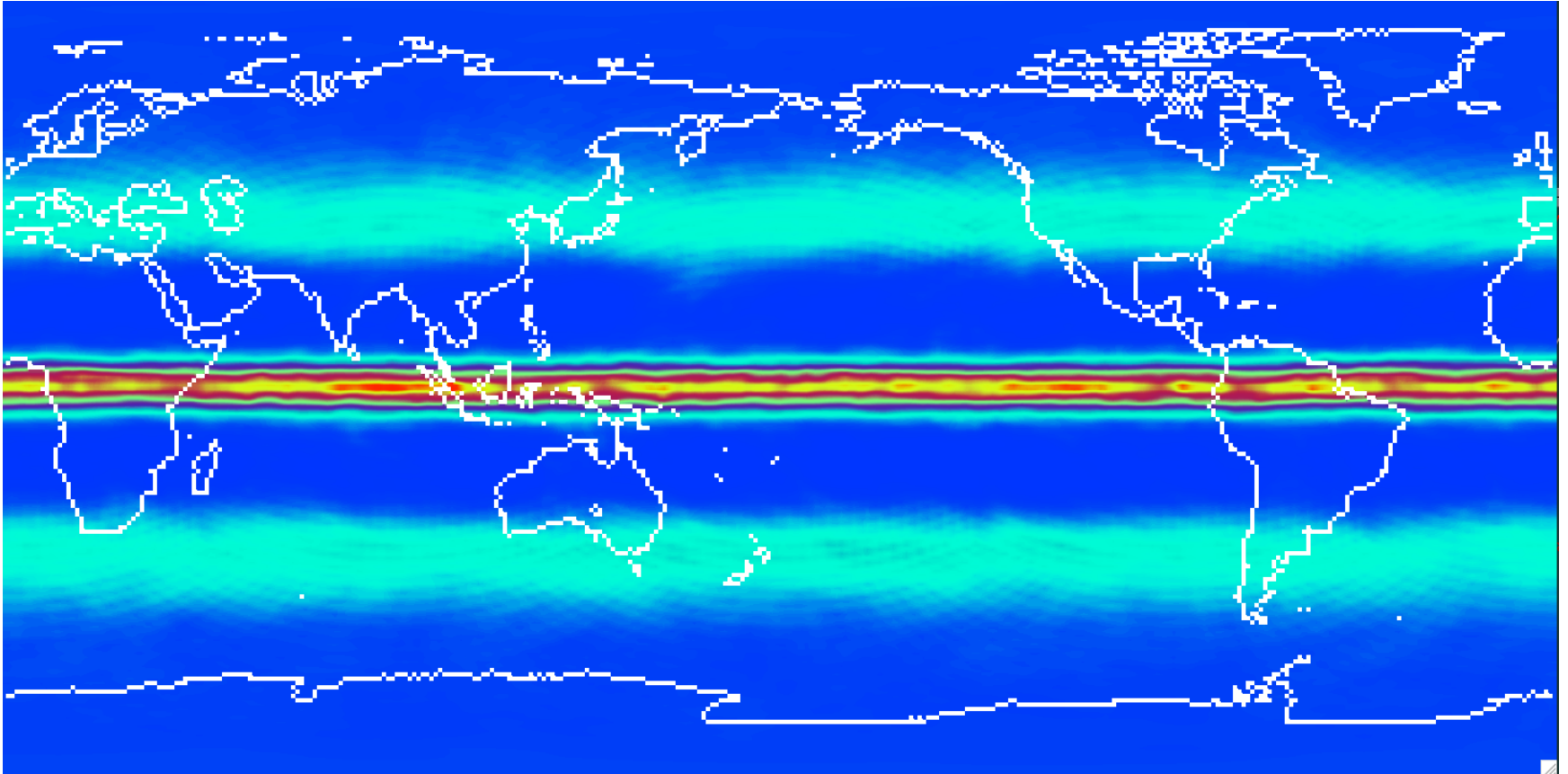


Temperature tendency: FT



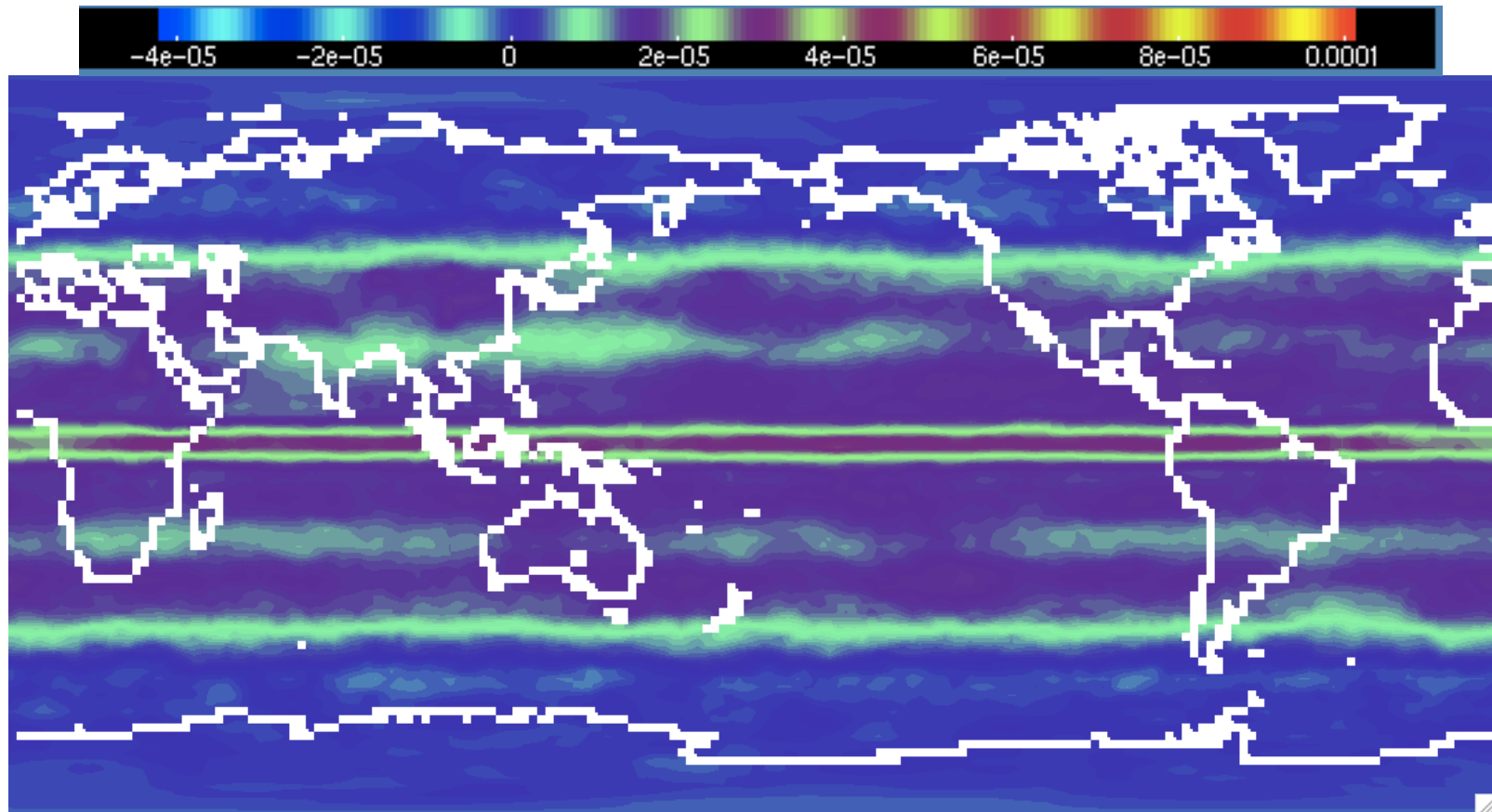
CAM-SE-CSLAM with linear interpolation from phys to dyn: 5 month average

PRECT



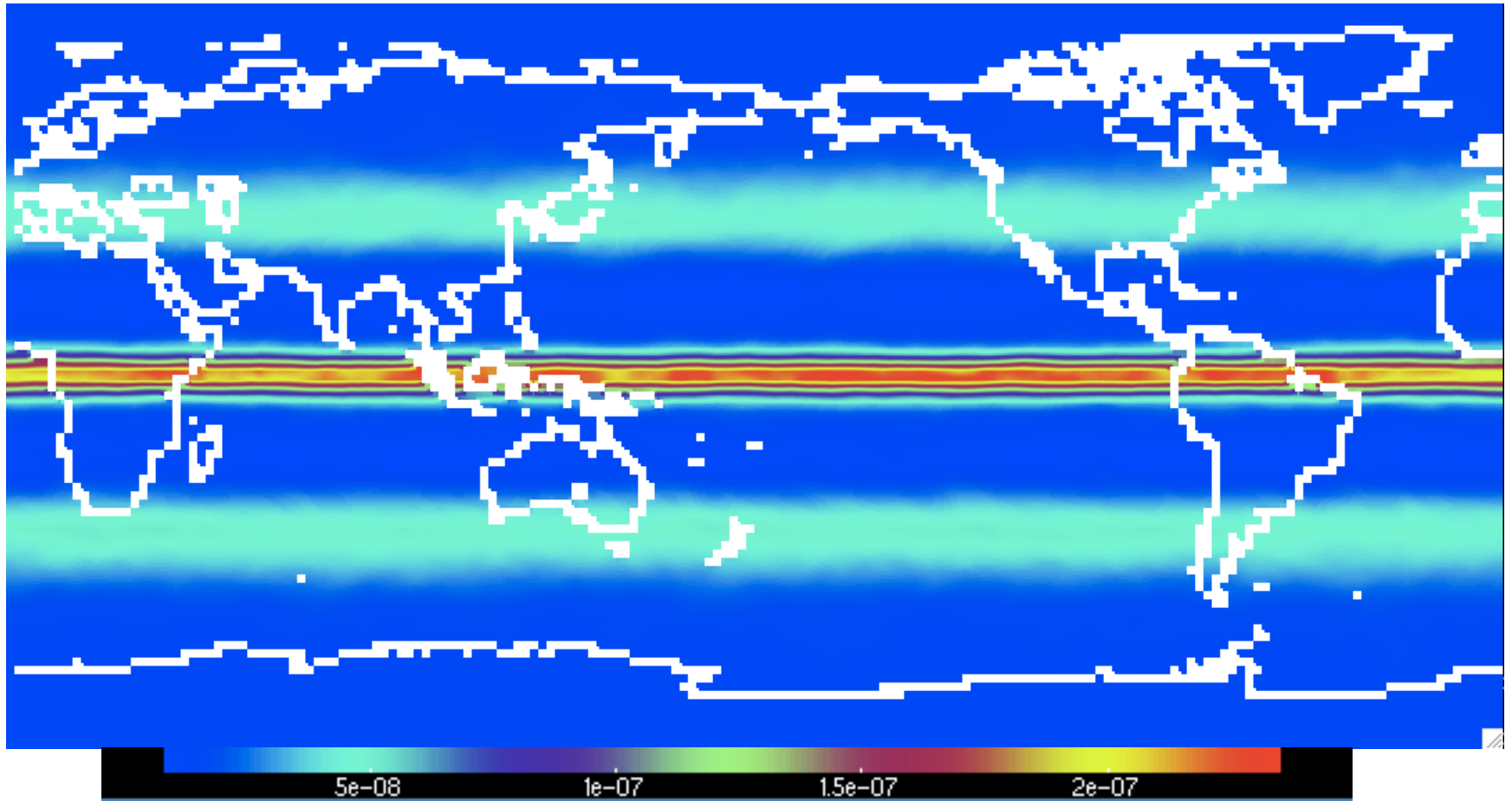
CAM4 SE-CSLAM-physgrid: linear interpolation phys to dyn: 5 month average
Plot looks similar for standard SE (but less noisy)

Temperature tendency: FT



**CAM-SE-CSLAM with cubic tensor product interpolation from phys to dyn:
18 month average**

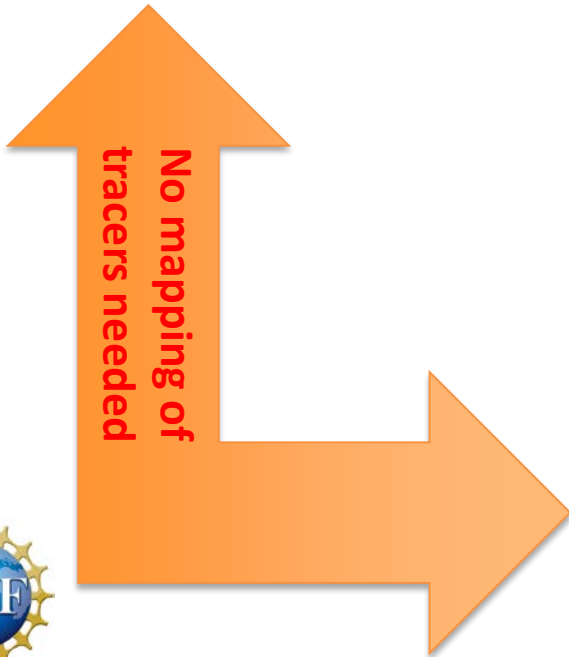
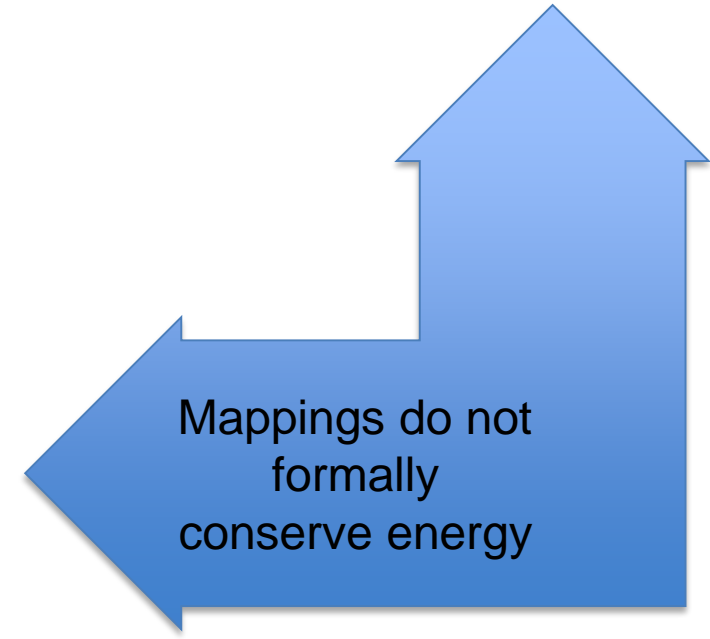
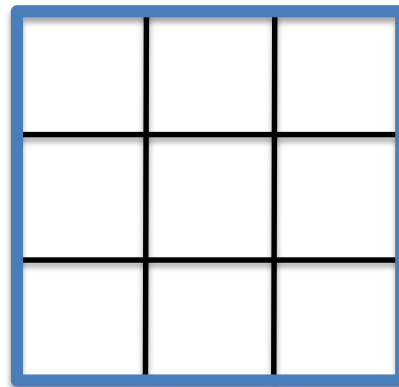
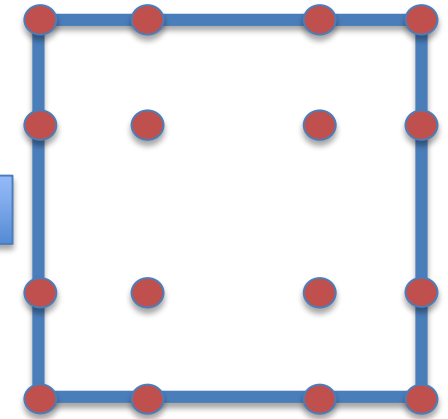
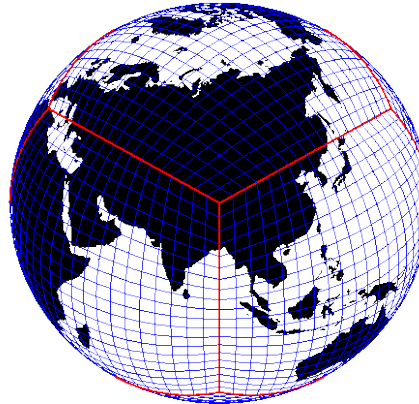
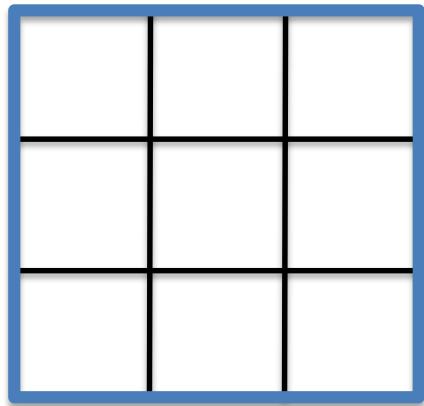
PRECT



**CAM-SE-CSLAM with cubic tensor product interpolation from phys to dyn:
18 month average**

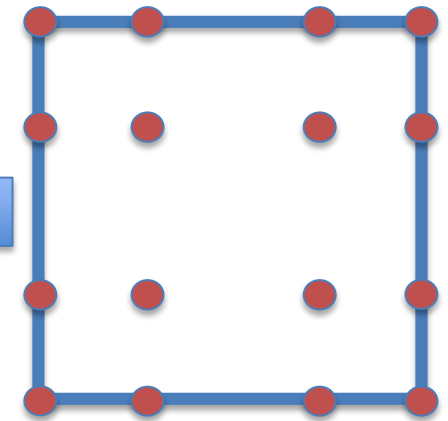
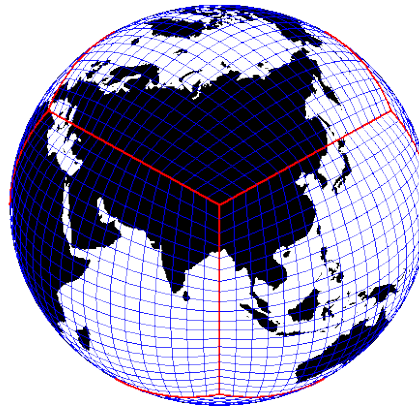
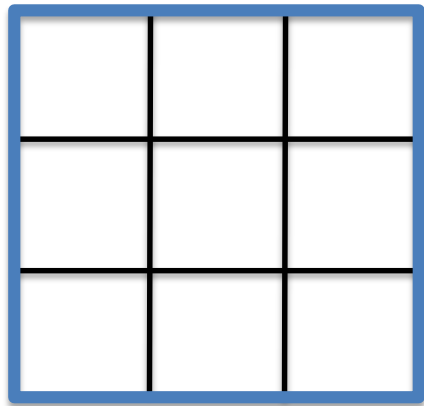


CAM-SE-CSLAM configuration



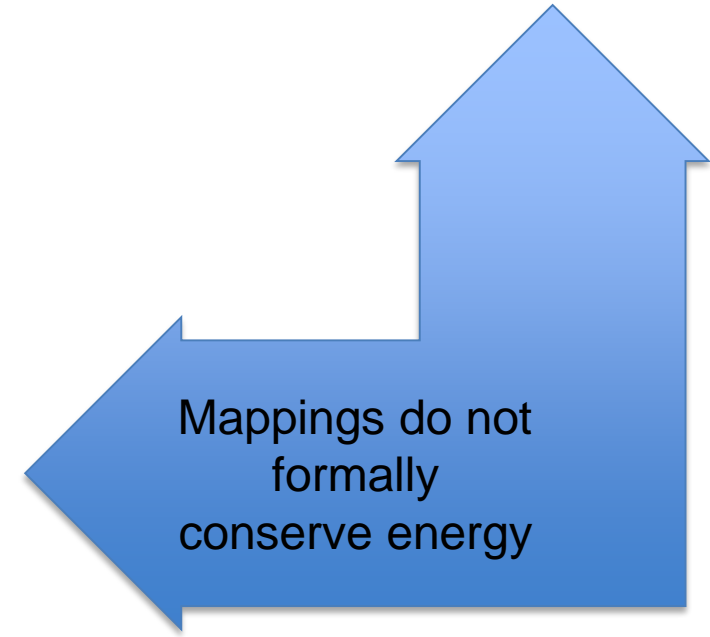
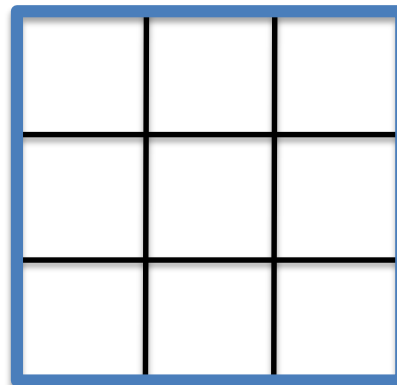


CAM-SE-CSLAM configuration



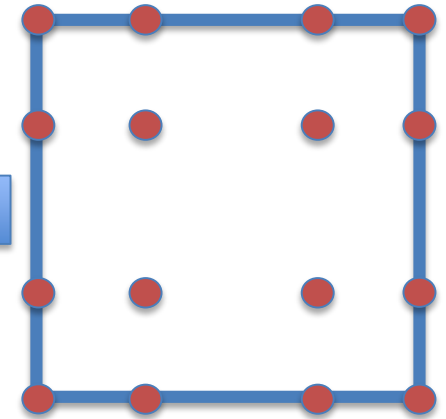
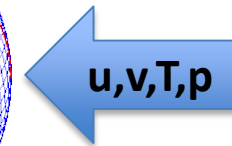
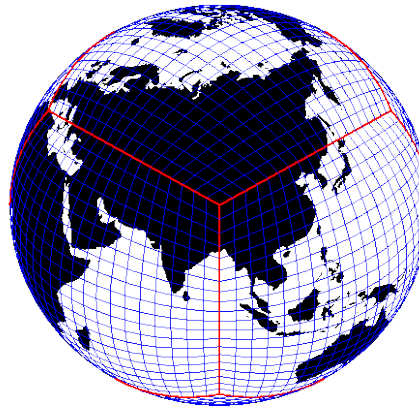
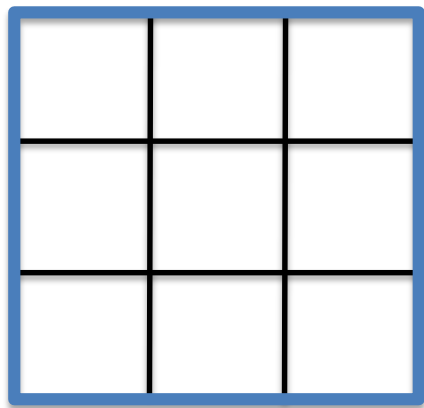
dE/dt energy fixer: -0.50 W/m^2
dE/dt dme_adjust : 0.27 W/m^2

(from CAM4 aqua-planet simulations)

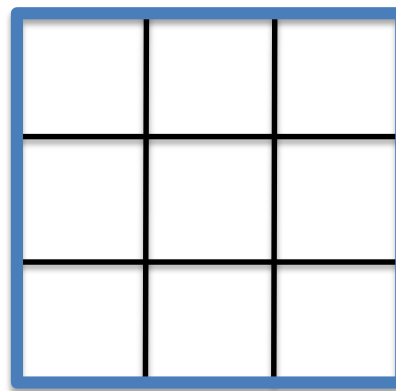
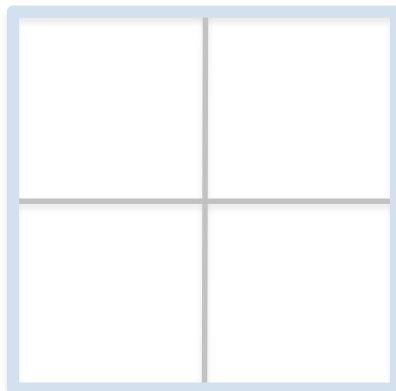




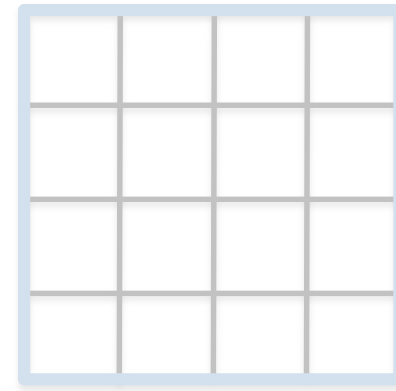
CAM-SE-CSLAM configuration



Coarser physics grid



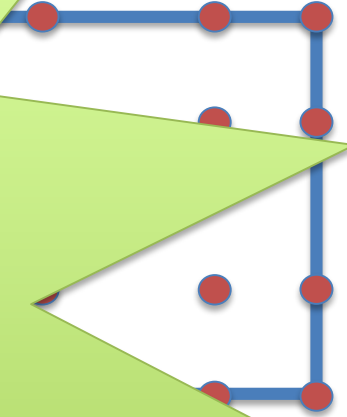
Finer physics grid





CAM-SE-~~CSLAM~~ configuration

We are getting very close to finally do science (beyond numerical methods research) with CAM-SE-~~CSLAM~~



Coa

Flu

