Thermal and silicate fronts in the Southern Ocean using the CESM LE simulations

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Antarctic Polar Front is a biogeochemical from

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°C Freeman & Lovenduski (2016)

Antarctic Polar Front is a biogeochemical from

Subduction, transition zone

South of Polar Front:

- salinity-dominated stratification
- high [macronutrients]
- low [micronutrients]

60E 120E 2002-2014

Implications: local physical, biogeochemical processes

°C Freeman & Lovenduski (2016)

Southern Ocean change

Observed change

- warmed, freshened
- carbon sink strengthened
- westerlies strengthened, shifted poleward

Predicted change

- continued poleward shift in westerlies
- ACC system to mirror westerlies

How has the Antarctic Polar Front (PF) changed over this period of substantial Southern Ocean change?

Has the PF moved?

SSH mean latitude -50 -52 -54 1995 2010

Sokolov & Rintoul (2009); Sallee et al. (2008)

Has the PF moved?



Sokolov & Rintoul (2009); Sallee et al. (2008)

Has the PF moved?



Sokolov & Rintoul (2009); Sallee et al. (2008)

Has the PF moved? SST SSH mean latitude mean latitude -54.4 no! -54.8 -50 -55.2 ves! mean intensity (°C/100 km) 0.04 -52 0 but, intensified... -0.04 -54 2014 2002 1995 2010

Sokolov & Rintoul (2009); Sallee et al. (2008)

What does this mean for BGC?



Sokolov & Rintoul (2009); Sallee et al. (2008)

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Freeman et al. (2016)

Sokolov & Rintoul (2009); Sallee et al. (2008)

The PF and the Silicate Front (SF)

World Ocean Atlas: annual mean surface silicate

sharp meridional Si gradients at the PF

relationship important in terms of **biogeography**



phytoplankton drawings: Sally Bensusen, NASA EOS Project Science Office

F-BGC relationship apparent in Drake Passag

MODIS Aqua [PIC] (December 2002)



PF-BGC relationship apparent in Drake Passag

60° S

75° S

MODIS Aqua [PIC] (December 2002)

x10⁻⁴

8

6

2

nol / m³

Freeman et al. (in prep)

Particulate Inorganic Carbon (PIC; calcite) concentration empirically derived from satellite

sharp PIC gradient at PF

Has the associated Si boundary shifted? (invoke model here)

Exploring BGC fronts in CESM LE

Thanks to CESM LE Community Project (Kay et al., 2015) & NSF/CISL/Yellowstone

Exploring BGC fronts in CESM LE

∆Si (2002-2014)

diatoms generally take up nutrients at Si:N = 1 ratio under adequate light, <u>nutrient conditions</u> regional coincidence: Si:N = 1 Δ Si PF (Δ SST)

mean PF (observed)
±1σ monthly mean PF
mean Si:N = 1 (modeled)



SF shifts poleward

mean **latitude where Si:N = 1** at the surface displaced southward both in historical record and in future climate scenario (RCP8.5)



Silicate Front (SF) shifts poleward

mean **latitude where Si:N = 1** displaced southward both in historical record and in future climate scenario (RCP8.5)



Defining long-term change

Present: 2000s 2000-2009 decadal average

Future: **2090s** *2090-2099 decadal average*

Long-term change: **2090s – 2000s** *Future – Present*

Ensemble Member spread: 1σ Significant change in Ens. Mean: long-term change > 1σ





















Conclusions



Observed frontal variability results depend on methods

Motivated by the observed biogeochemical significance of the PF

CESM LE suggests a poleward contraction of the "SF"

Thanks for your attention.

Questions? Feedback?

Contact me at natalie.freeman@colorado.edu

Exploring BGC fronts in CESM LE





SF seasonal cycle



Silica cycle

Surface waters receive considerable amounts of silicic acid from the rest of the world ocean through the upwelling of the CDW, fed by contributions of deep waters of the Atlantic, Indian, and Pacific Oceans.

The SO exports a considerable flux of the silicic acid that is not used by diatoms in surface waters through the northward pathways of the SAMW, AAIW, and ABW.

**The question of the steady state of the Si cycle in the modern ocean remains an open question.

Location Relative to Polar Front South North

High	Big, well-silicified diatoms outcompete nano-, pico-phytopl., coccos. and dinos; Low Calcification; CO ₂ drawdown (NCP>NCC); High export flux; Low transfer eff.;(e.g. South of Polar Front near islands)	Coccos outcompete diatoms, nano-, pico-phytopl., dinos ; High calcification; CO ₂ source (NCC>NCP); Moderate export flux; High transfer eff. (e.g. Patagonian Shelf; SAF [Indian sector], south of ARC, near islands)	Strong shear; enhanced Ekman pumping
Low	Small diatoms (e.g. <i>Fragillariopsis</i> <i>sp.</i>); few coccos; nano-, pico- phytopl., dinos coexist and covary; No net impact on pCO_2 (moderate NCP:NCC); Moderate export flux; Low-to-moderate transfer eff.; (e.g. south of Polar front, Indian sector, away from islands)	Coccos, dinos, nano-, pico- phytopl.coexist and covary; Low to moderate calcification; Neutral to slight pCO_2 source (NCC=>NCP); Moderate export flux; Moderate transfer eff.; (e.g. north of ARC and STF, Atlantic and Indian sectors, away from islands)	Low shear; reduced Ekman pumping

ron Availabilit

Negative Zero to Positive RNPG (μ_N-μ_{Si}; d⁻¹)

Figure 11. Conceptual model for balance of coccolithophores and noncalcifying phytoplankton growth in GCB versus diatom growth (determined as RNPG) and effects on CO₂ source/sink dynamics and sinking particle fluxes. Abbreviations used in table: coccos = coccolithophores; dinos = dinoflagellates; nanophytopl. = nanophytoplankton; picophytopl. = picophytoplankton; eff. = efficiency; ARC = Agulhas Return Current.

$(Si:N = 1) \sim (Si^*)$



CESM LE vs. WOA13





τ_x : present-day mean and long-term change



Silicate Front shifts poleward (1920-2100)



Silicate Front shifts poleward (2002-2014)



CESM LE: PF vs. SF at 200 m 0 subsurfa - subsurfa bottom depth (km) 3 6



NPP

changes in marine NPP can thus result only from (i) changes in the phytoplankton growth rate and (ii) changes in phytoplankton biomass