

Rapid expansion of Greenland's low-permeability ice slabs through the 21st Century

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Refrozen ice inhibiting runoff in Greenland

► Firn not acting as the “buffer” it once was

NATURE CLIMATE CHANGE | LETTER



Greenland meltwater storage in firn limited by near-surface ice formation

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Approximately half of Greenland's current annual mass loss is attributed to runoff from surface melt¹. At higher elevations, however, melt does not necessarily equal runoff, because meltwater can refreeze in the porous near-surface snow and firn². Two recent studies suggest that all³ or most^{3, 4} of Greenland's firn pore space is available for meltwater storage, making the firn an important buffer against contribution to sea level rise for decades to come³. Here, we employ *in situ* observations and historical legacy data

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Research article

30 May 2016

Extraordinary runoff from the Greenland ice sheet in 2012 amplified by hypsometry and depleted firn retention

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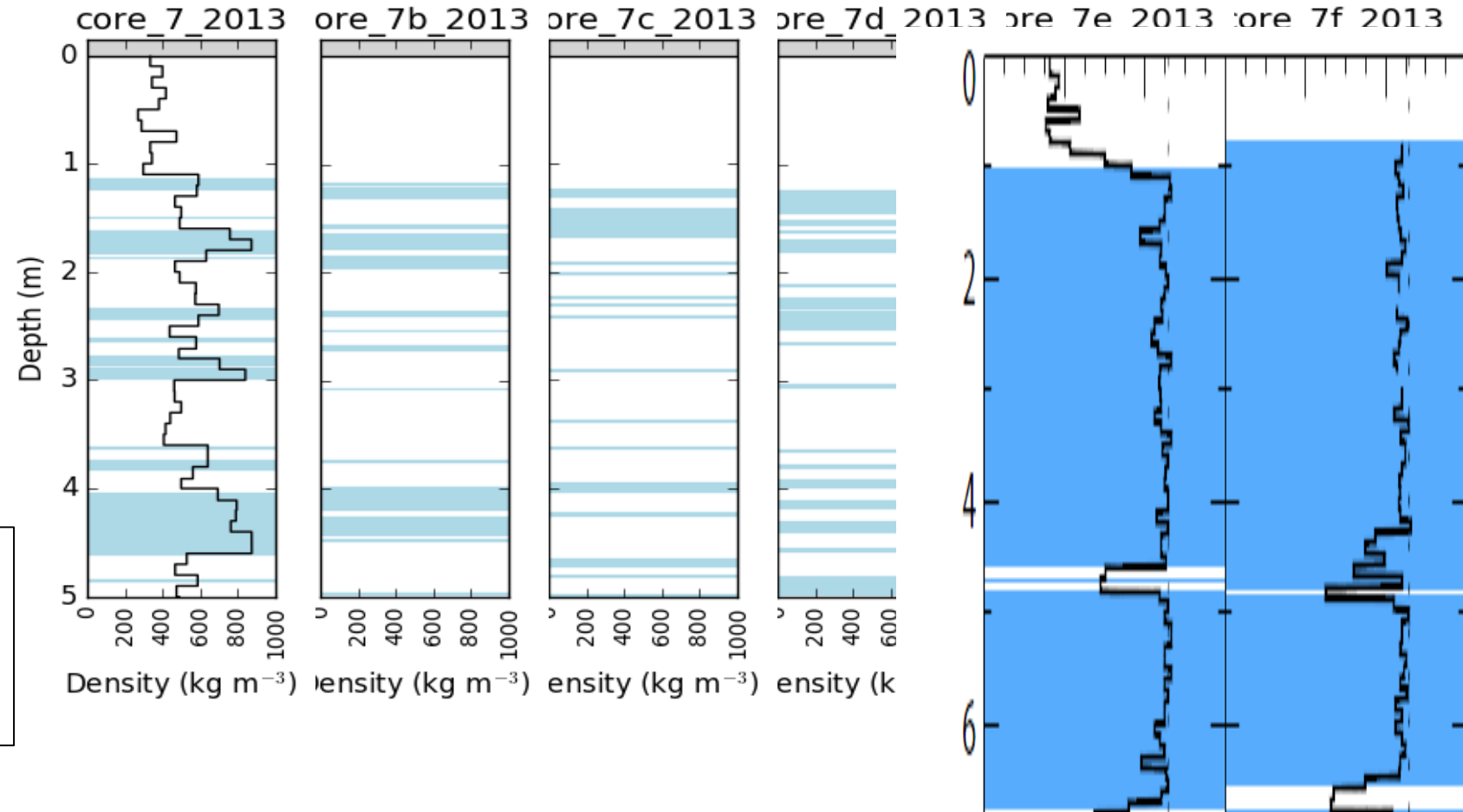


Ice "lenses" versus "slabs"

"Site 7", 2100 m a.s.l.

KAN-U, 1850 m a.s.l.

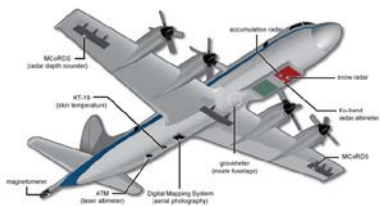
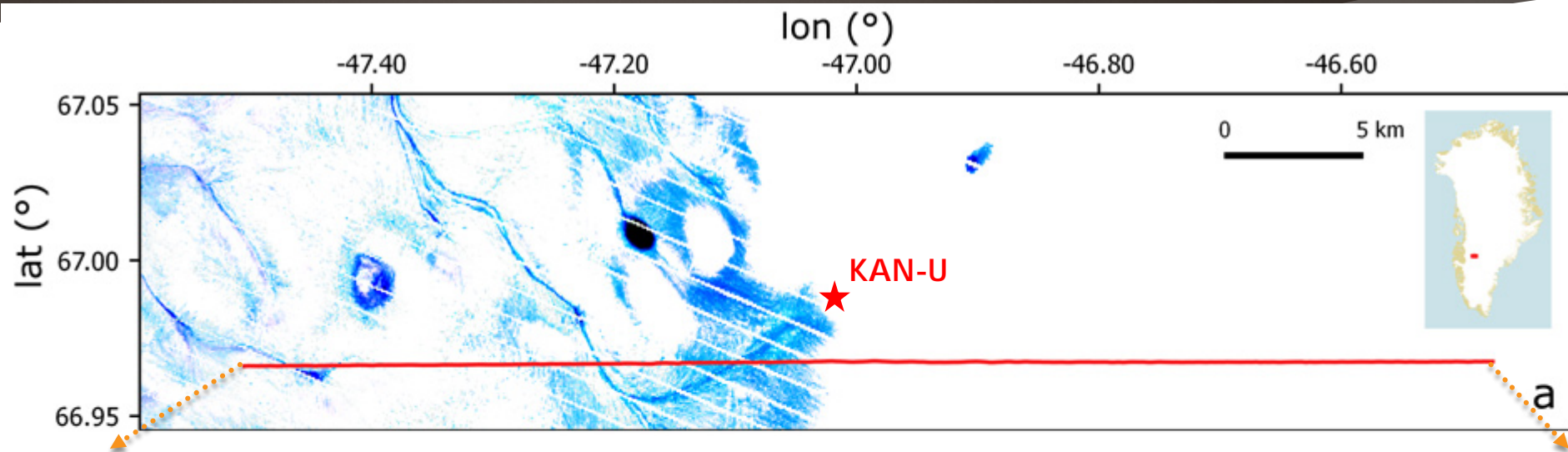
- ▶ Lenses <30-50 cm are spatially heterogeneous
- ▶ Runoff **only** yet seen over >1 m multi-annual "ice slabs"



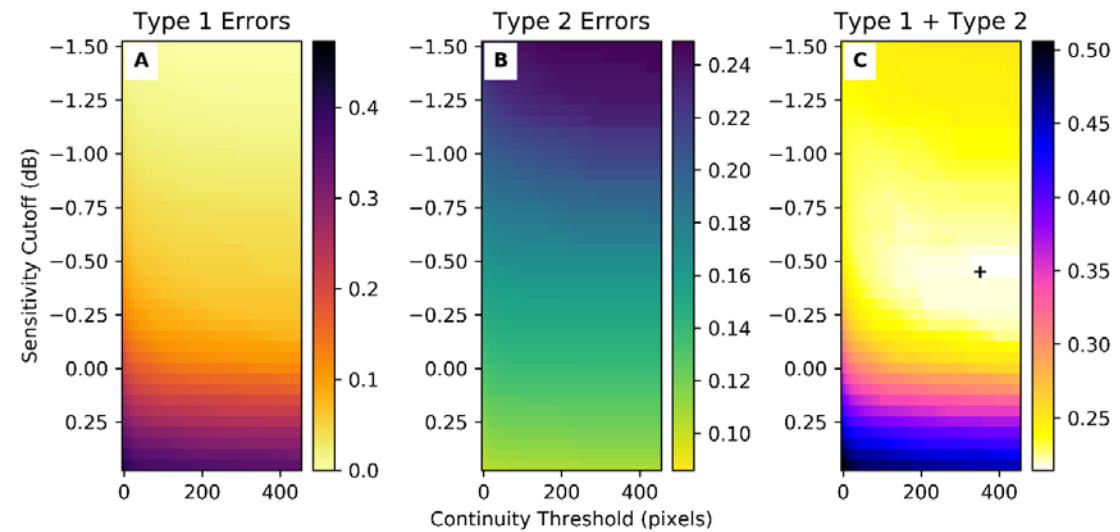
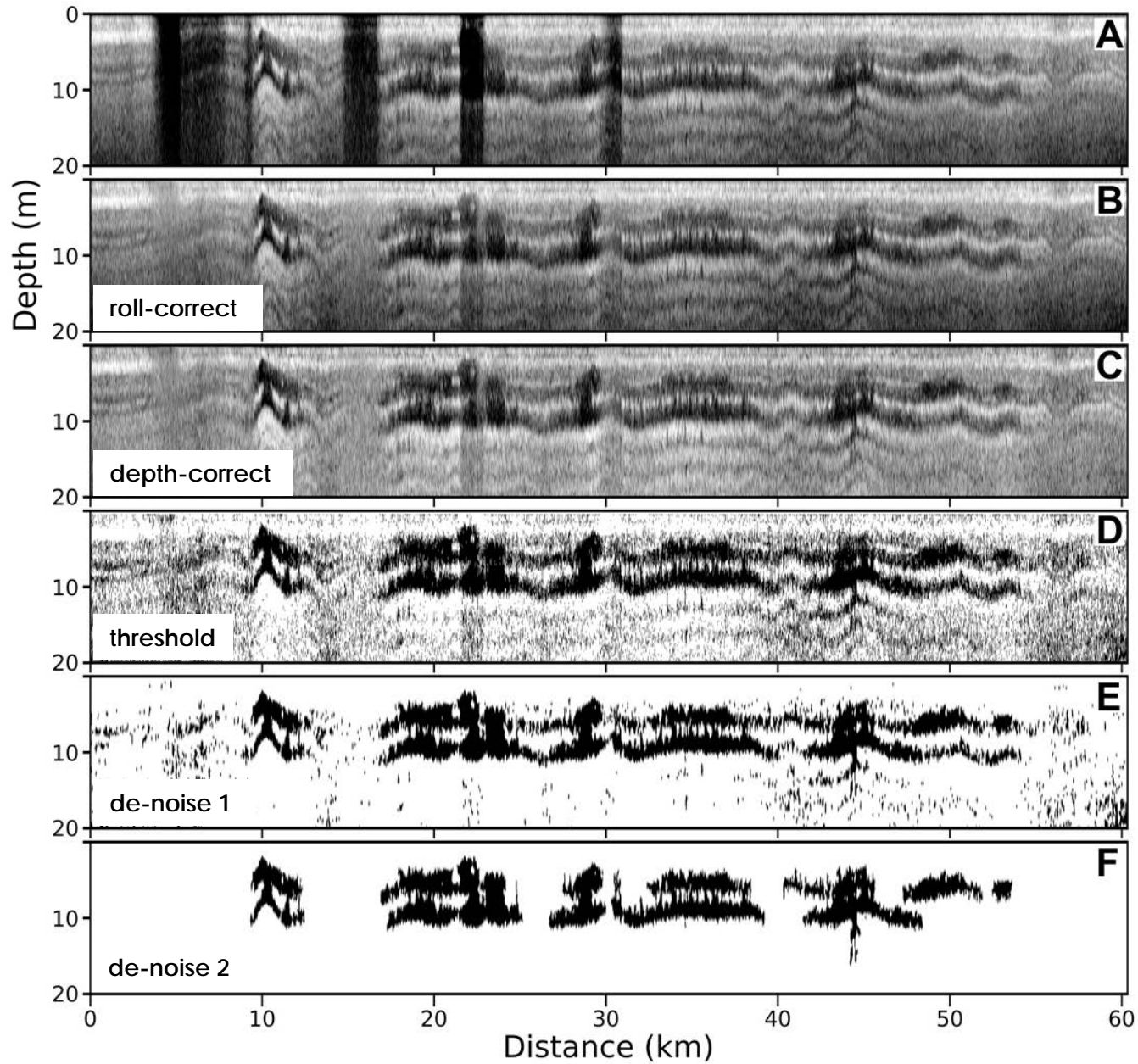
ICE SLABS:

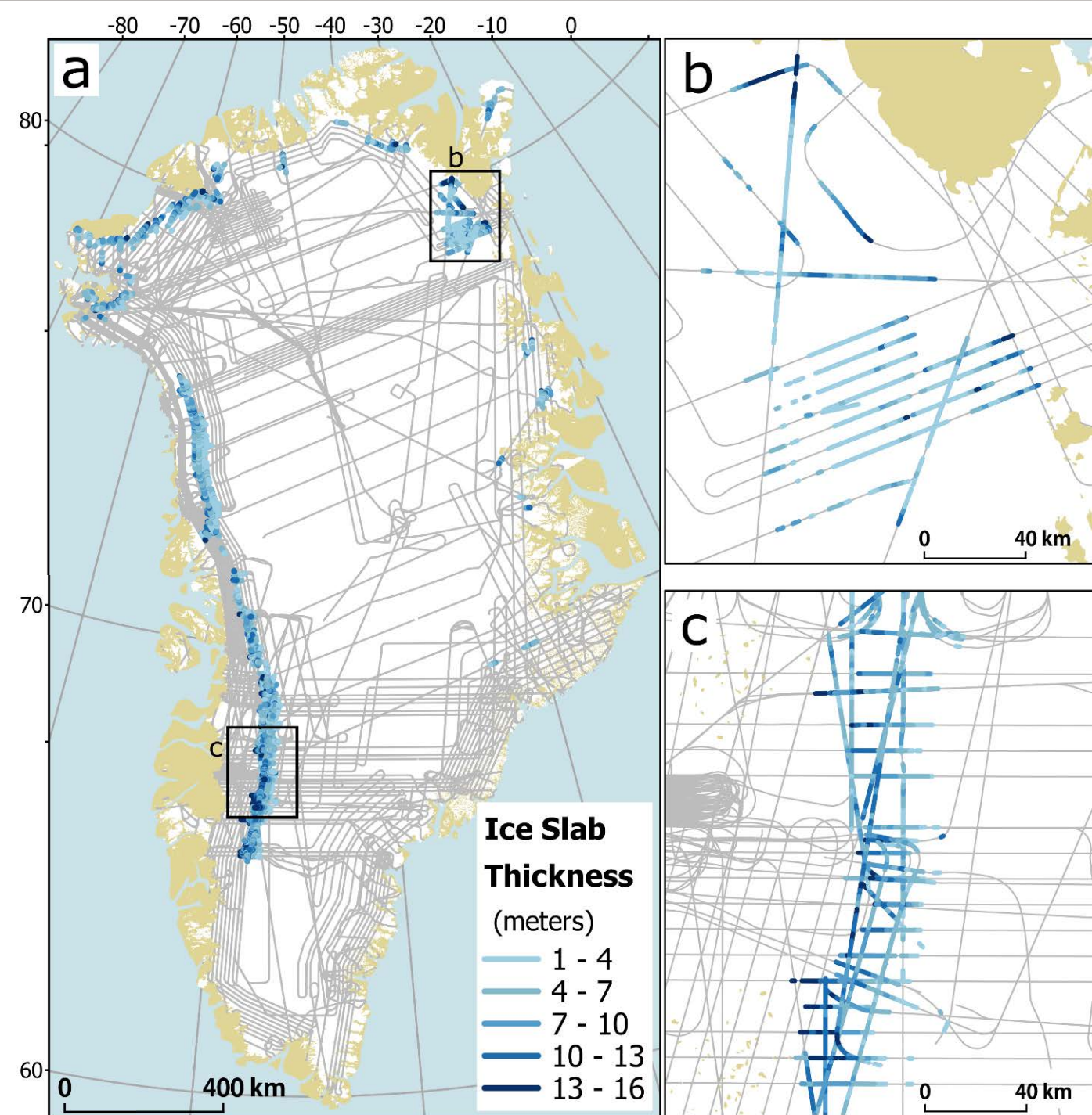
Layers of refrozen ice within firn ≥ 1 m thick and continuous for ≥ 1 kilometer.

In Situ and IceBridge Radar



OIB Radar – Data Processing

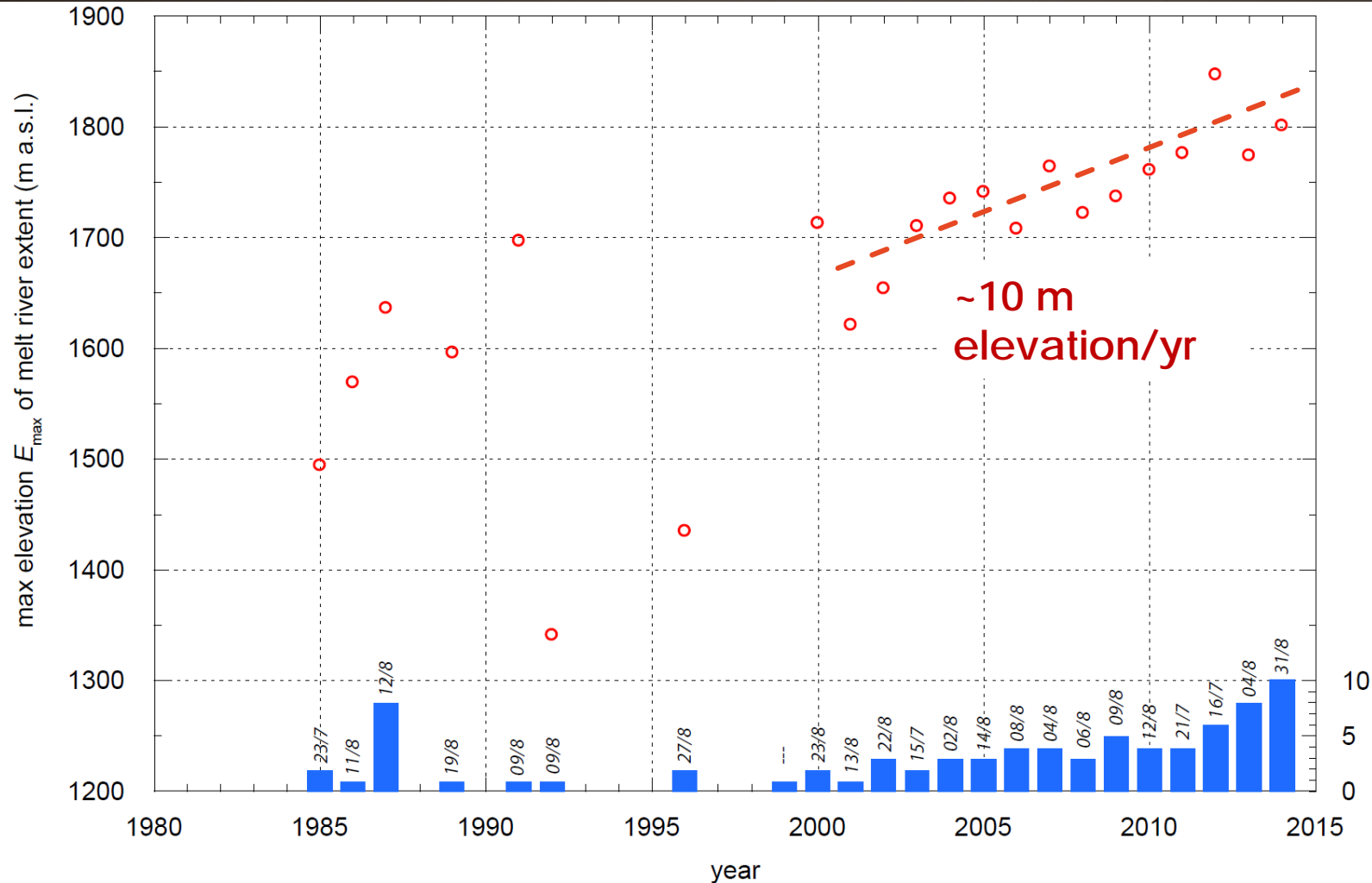




Ice Slabs in Greenland

- ▶ >69,000 km² of Greenland's firn (~5 % of Greenland)
- ▶ ~15-25% increase in runoff zone size
- ▶ Enhance runoff in WARM melt years

Uphill Migration of the K-Transect Runoff Line

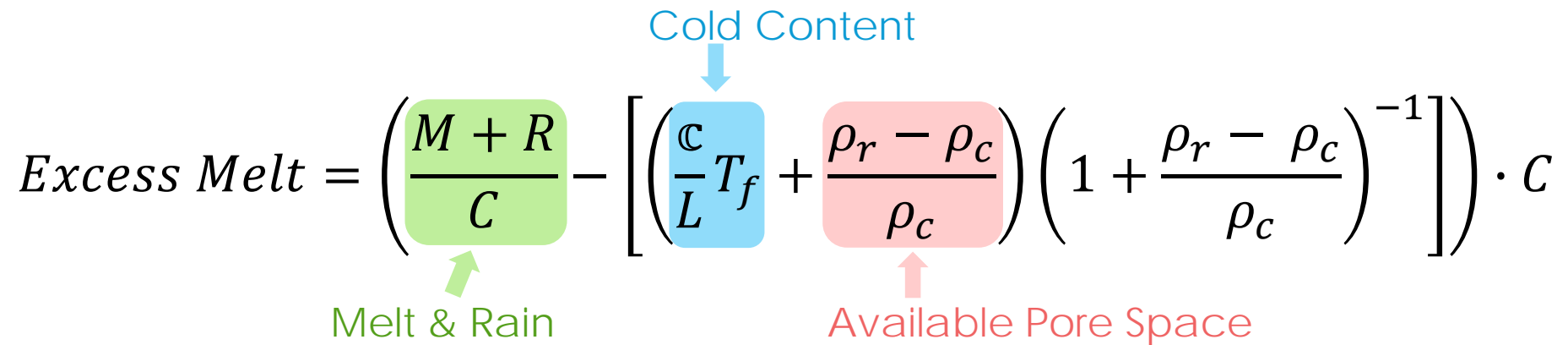


Machguth, H., MacFerrin, M., van As, D., Box, J. E., Charalampidis, C., Colgan, W., ... van de Wal, R. S. W. (2016). Greenland meltwater storage in firn limited by near-surface ice formation. *Nature Climate Change*, 6(4), 390–393.

Excess Melt

Modified from Pfeffer, Meier & Illangasekare (1991)

$$\text{Excess Melt} = \left(\frac{M + R}{C} - \left[\left(\frac{c}{L} T_f + \frac{\rho_r - \rho_c}{\rho_c} \right) \left(1 + \frac{\rho_r - \rho_c}{\rho_c} \right)^{-1} \right] \right) \cdot C$$

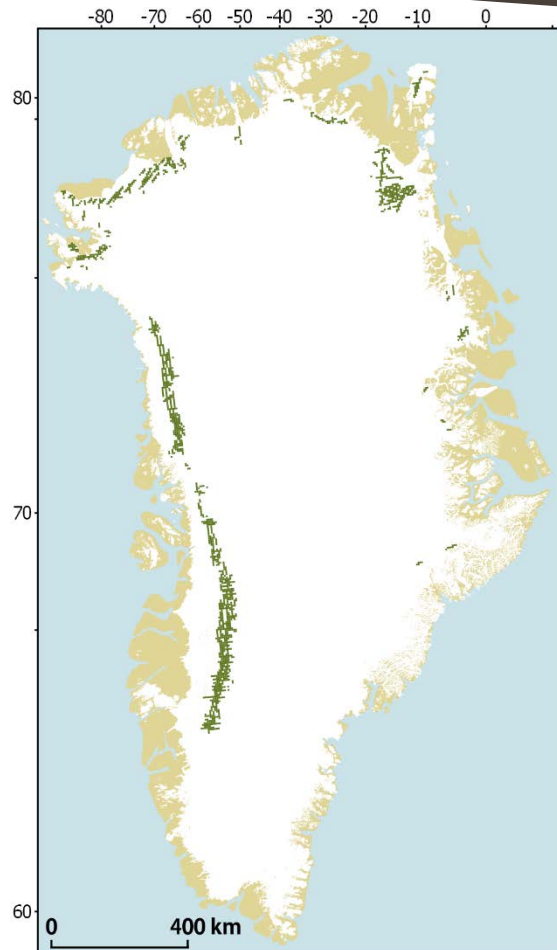


- M = melt (kg m^{-2})
- C = snowfall (kg m^{-2})
- R = rain (kg m^{-2})
- L = latent heat of fusion of ice (J kg^{-1})
- c = heat capacity of water ($\text{J kg}^{-1} \text{ } ^\circ\text{C}^{-1}$)
- T_f = firm temperature ($^\circ\text{C}$ below freezing)
- ρ_r = density of refrozen ice (kg m^{-3})
- ρ_c = density of fresh snow (kg m^{-3})

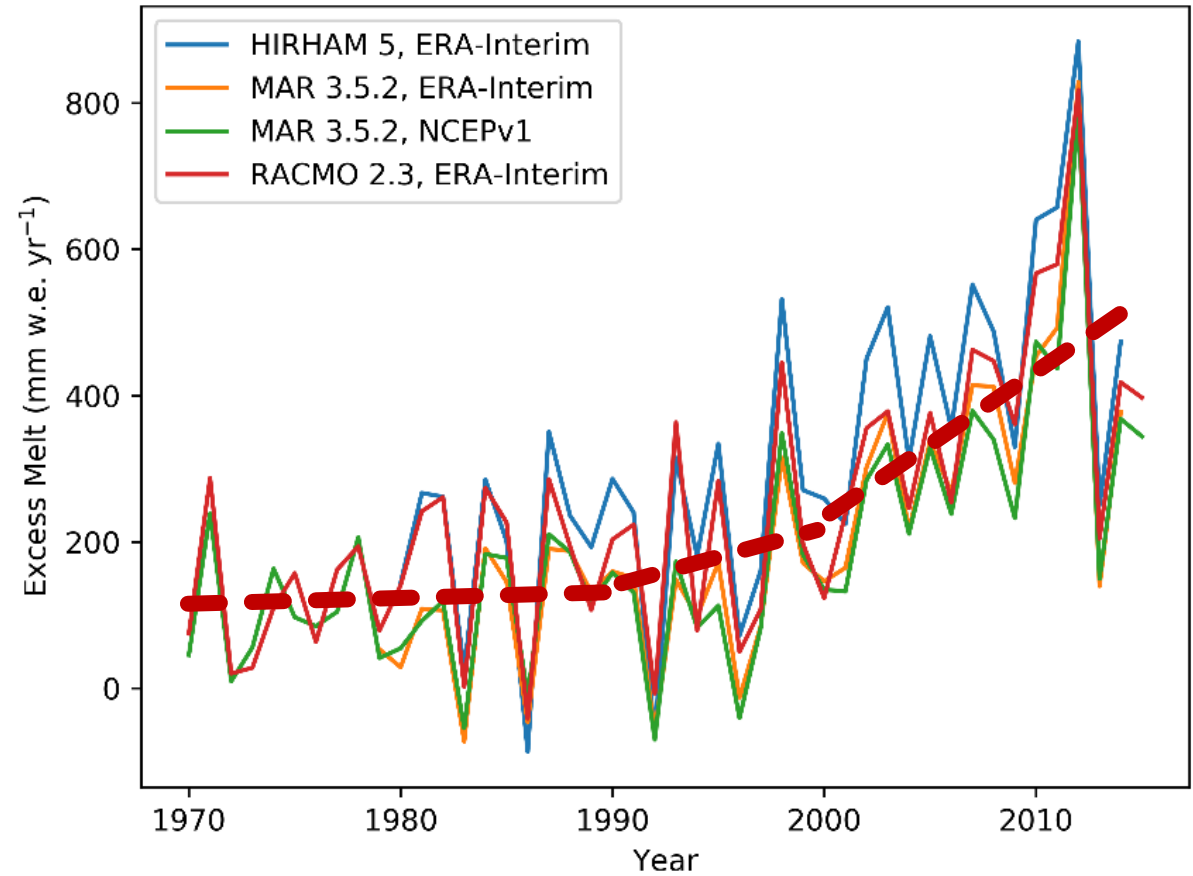
Excess Melt < 0 **BUFFERING** ❄️

Excess Melt > 0 **SATURATING** 💧

Excess Melt Trends



- ▶ Slight increases after 1990
- ▶ Larger increases after 2000



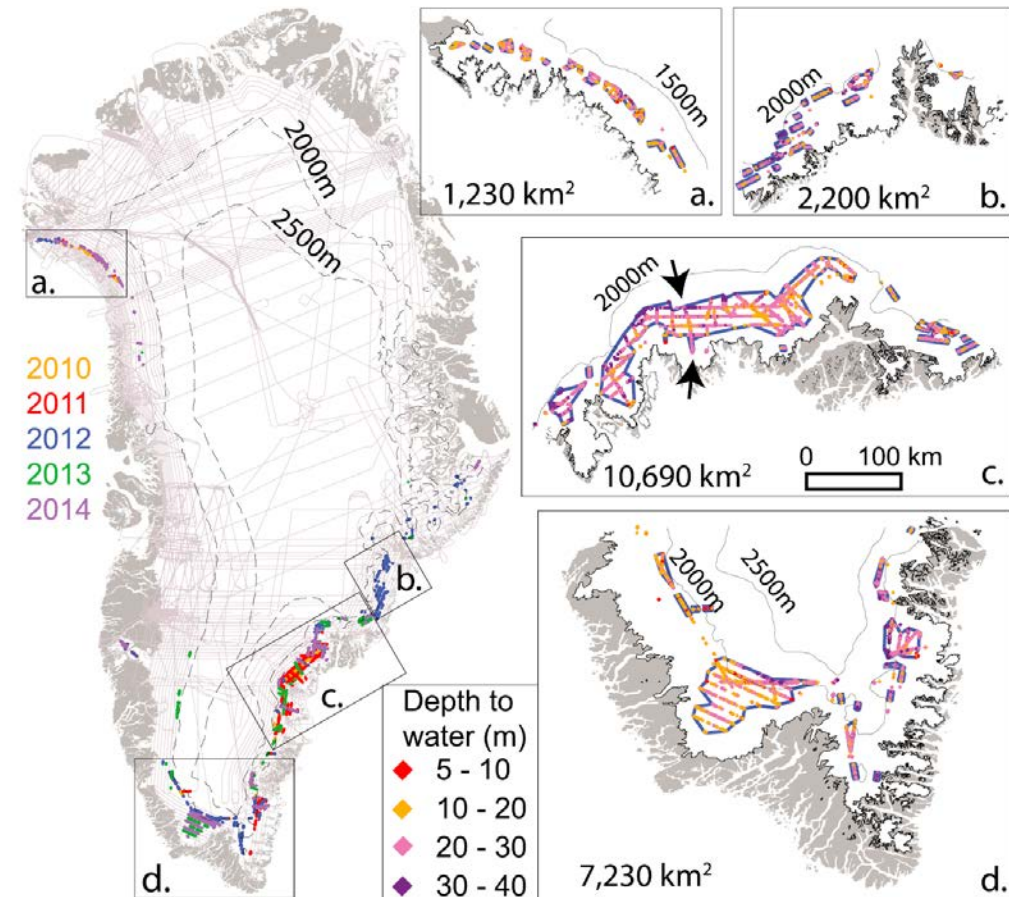
Accumulation Cutoff

- ▶ Perennial Firn Aquifers (PFAs) form in **high**-accumulation areas

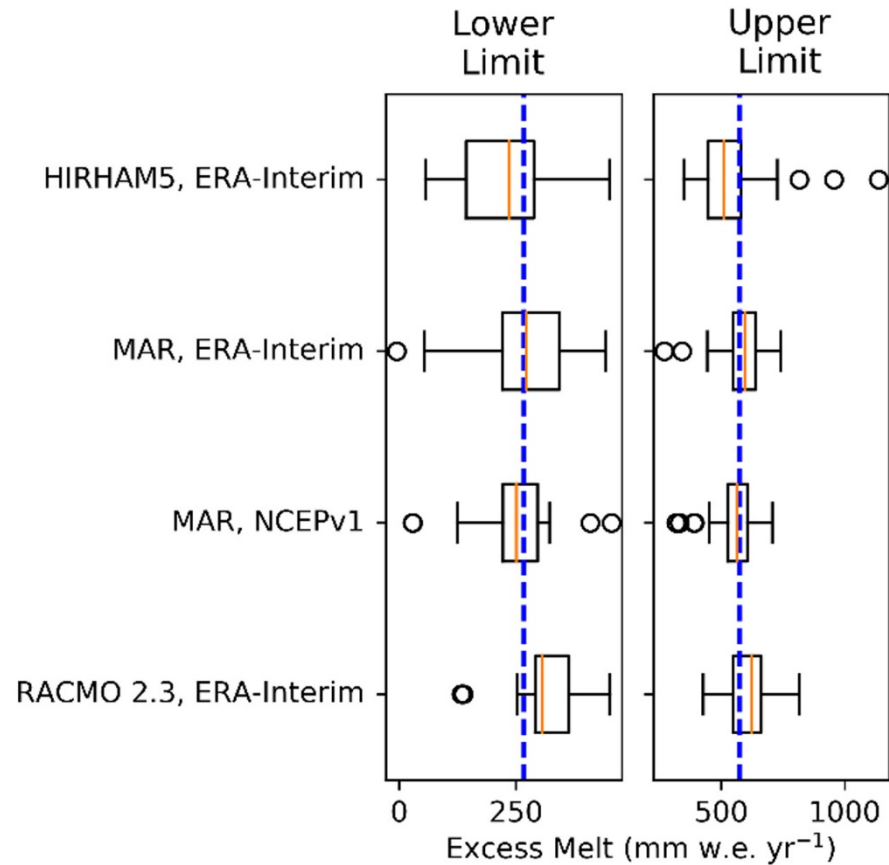
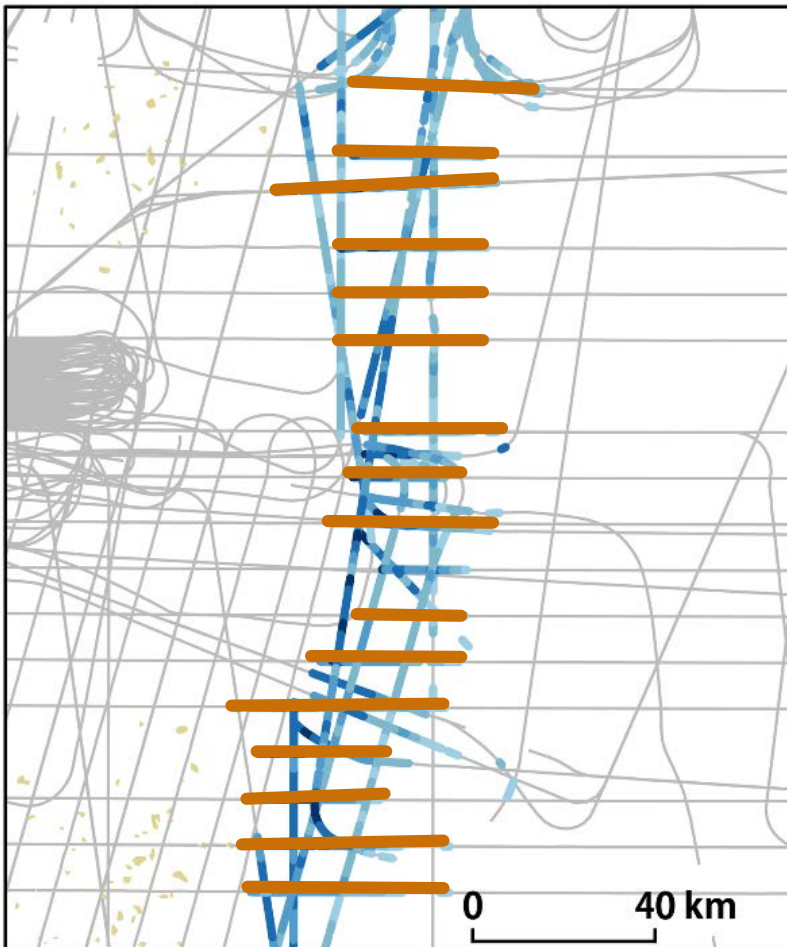
Forster, et al. (2014), Koenig, et al. (2013), Miege, et al (2016)

- ▶ Ice Slabs only form in **low**-accumulation areas

≤ 572 mm w.e. / year (~ 1.8 m snow)



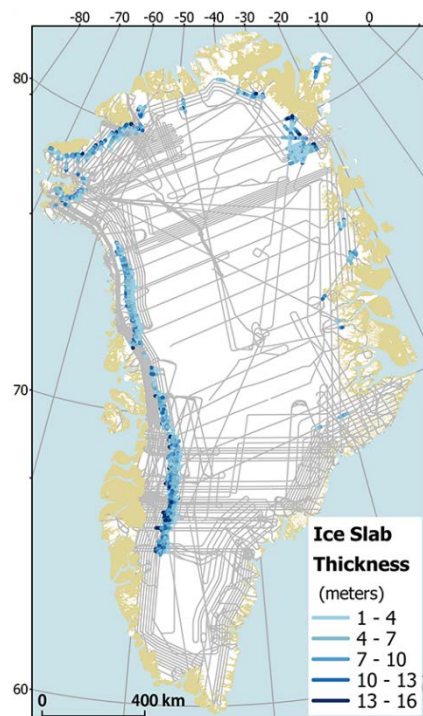
How much Excess Melt causes ice slabs?



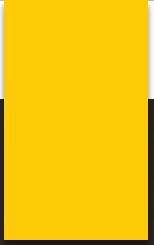
~266 – 573 mm w.e. / yr
for 10 years or more
has caused ice slabs to
form in Greenland

Mapping with RCMs & Reanalysis

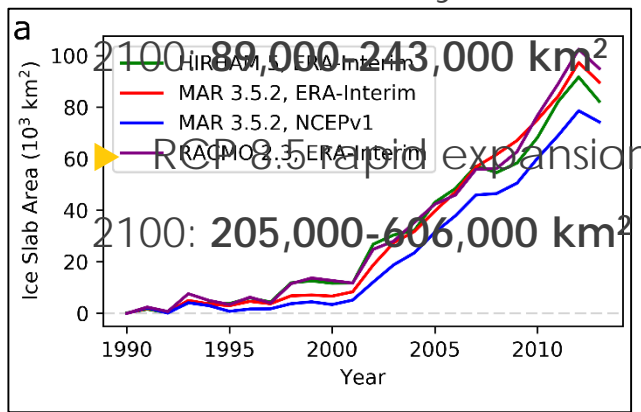
▶ 2014 extent 74,000-95,000 km²



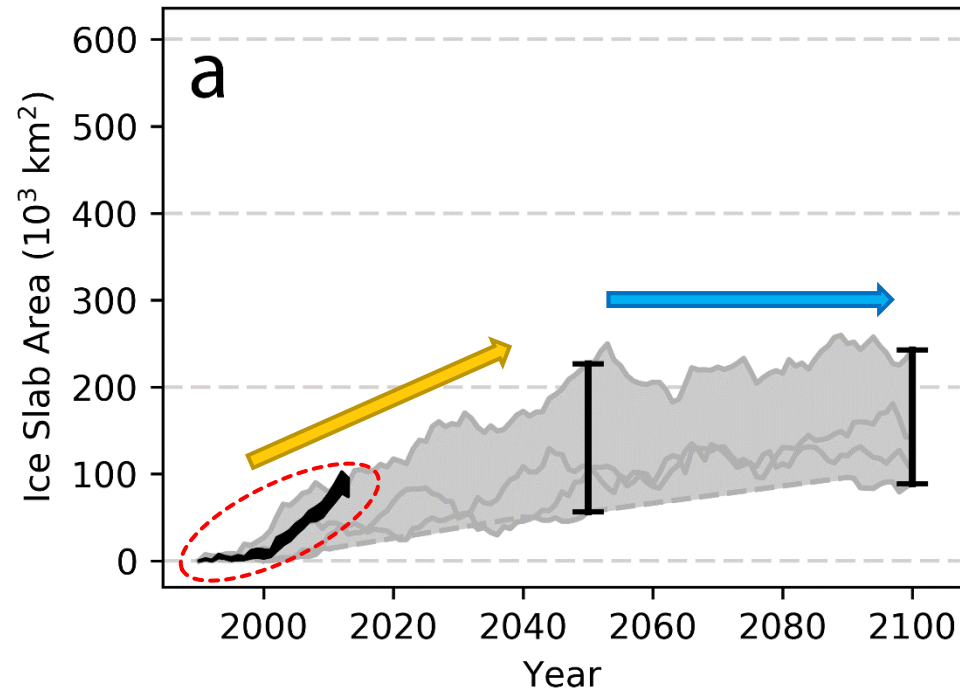
Modeling ice slabs with GCMs



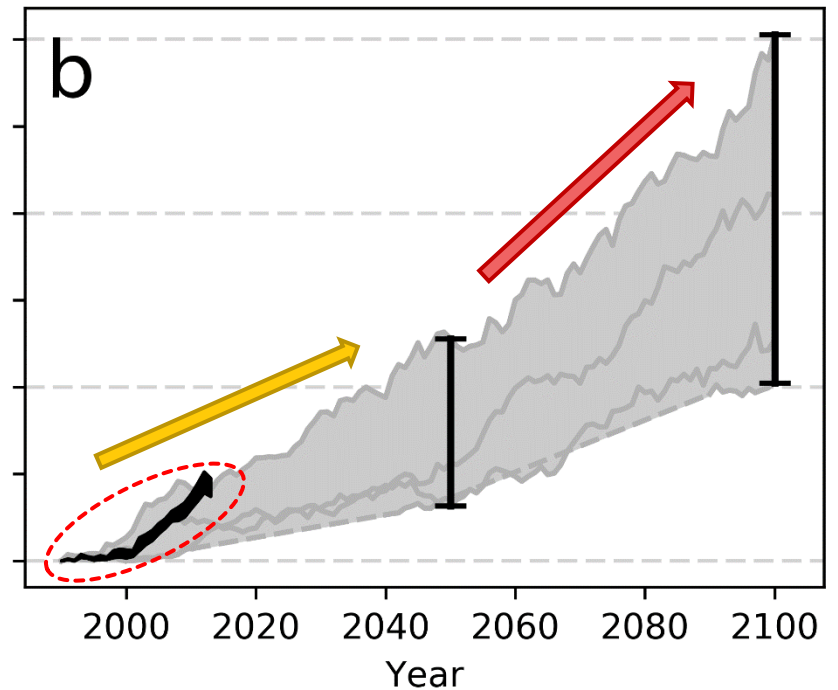
- ▶ 1990-2050: ECEarth
- ▶ HIRHAM 5, CanESM2
- ▶ MAR 3.5.2, CanESM2
- ▶ MAR 3.5.2, MIROC5
- ▶ MAR 3.5.2, NorESM1
- ▶ 2050-2100: RACMO 2.1, HadGEM2
- ▶ RCP 4.5 steady



“Moderate” emissions
RCP 4.5



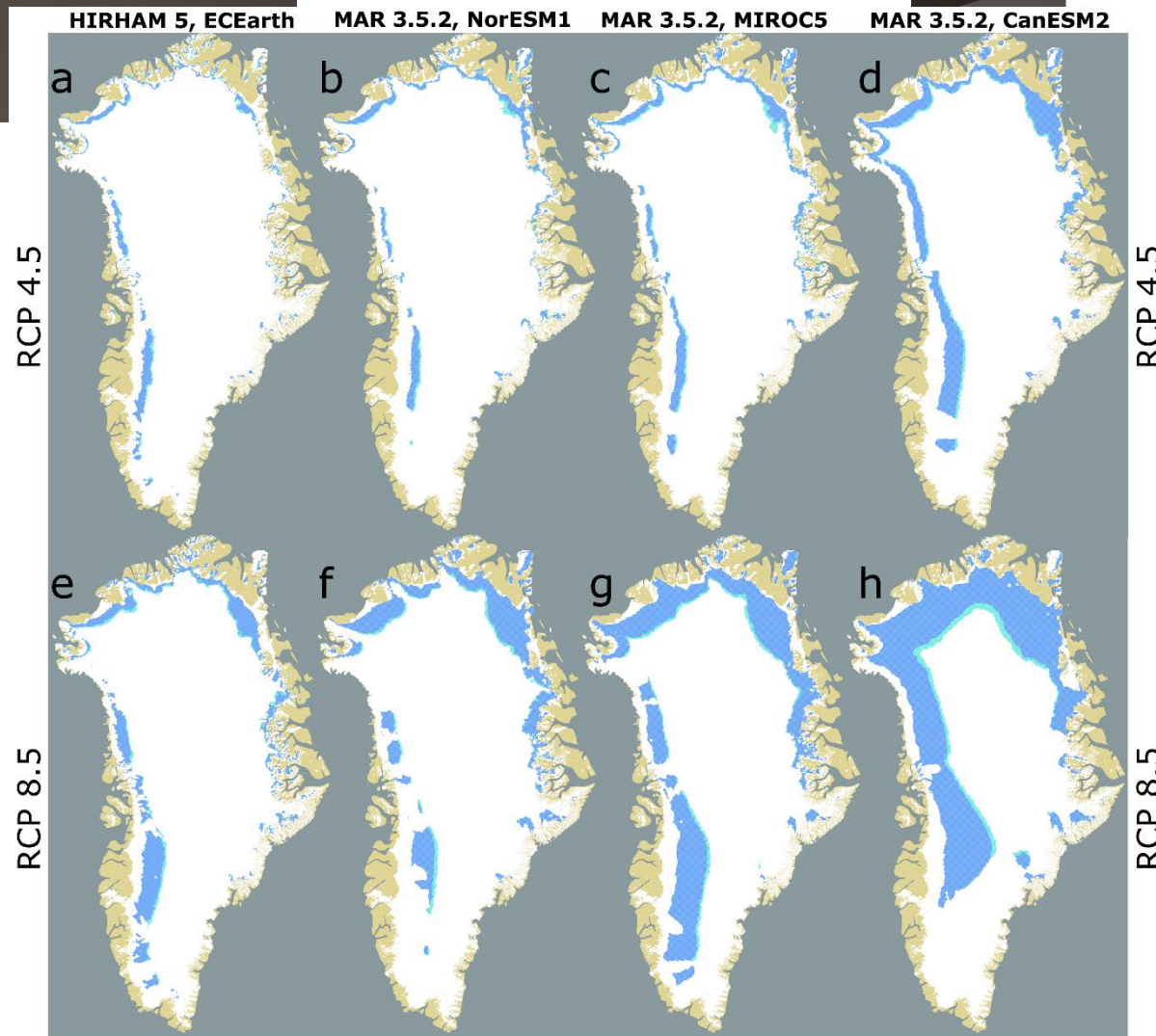
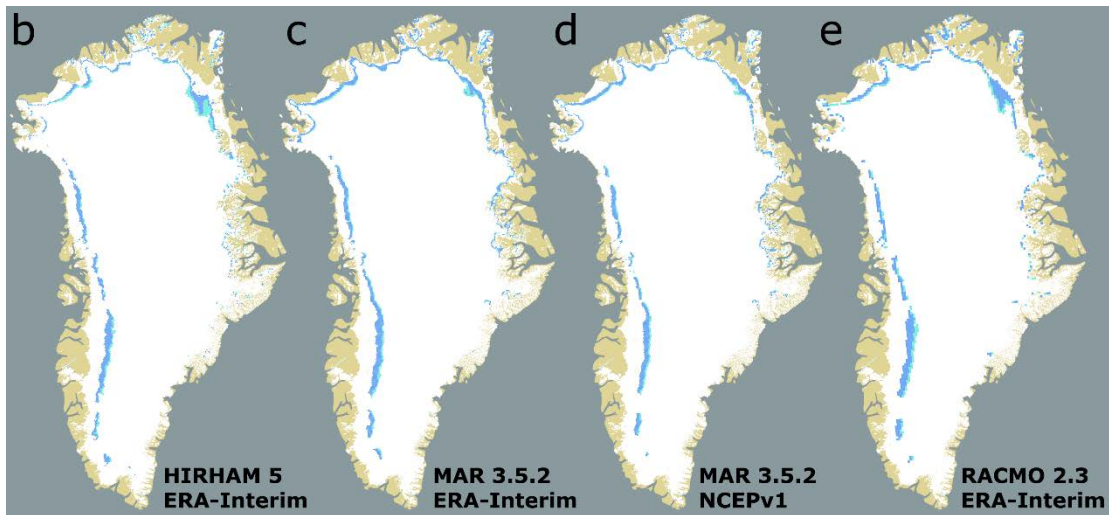
“High” emissions
RCP 8.5



Boundary forcing is critical

Bounded by **GCMs** (2100)

Bounded by **Reanalysis** (2013)



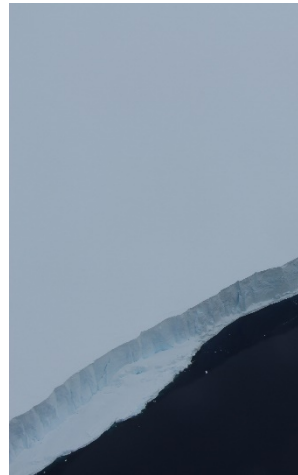
Ice Slabs in Antarctica?

▶ Larsen C ice shelf

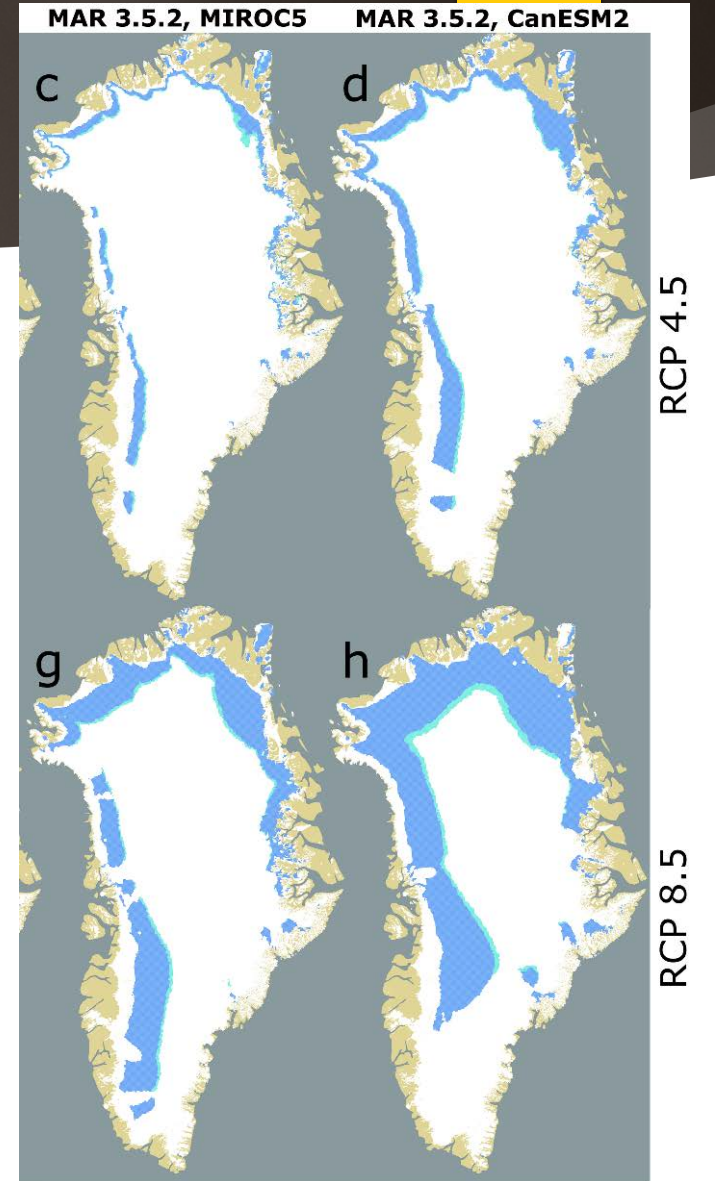
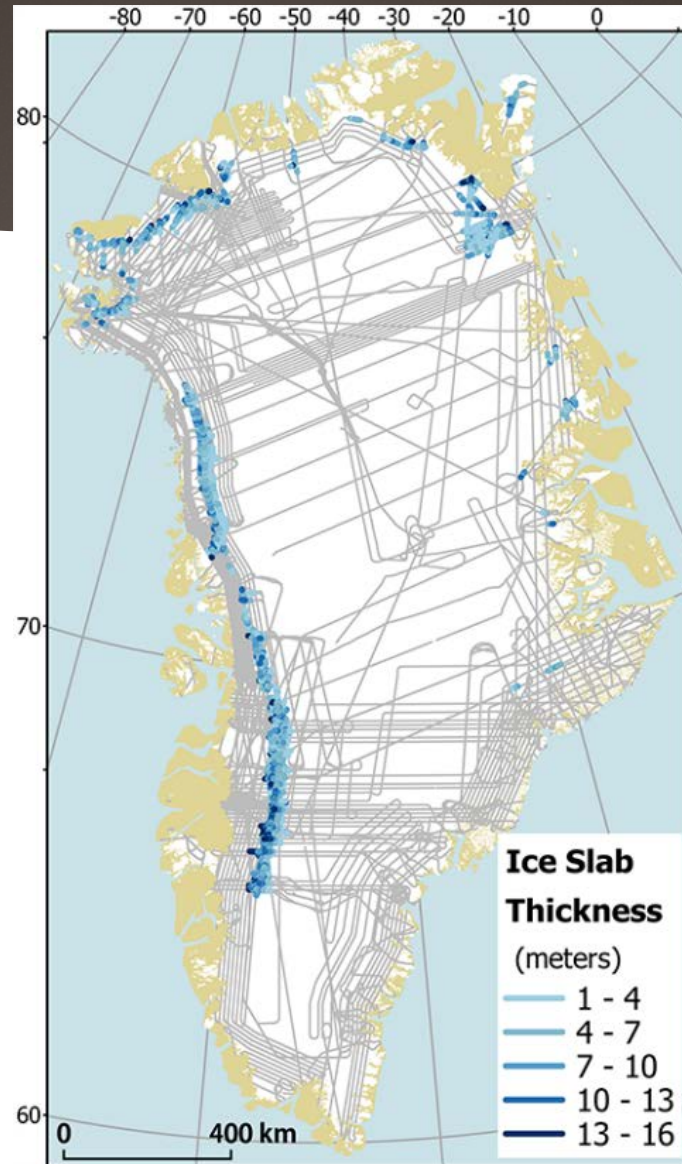
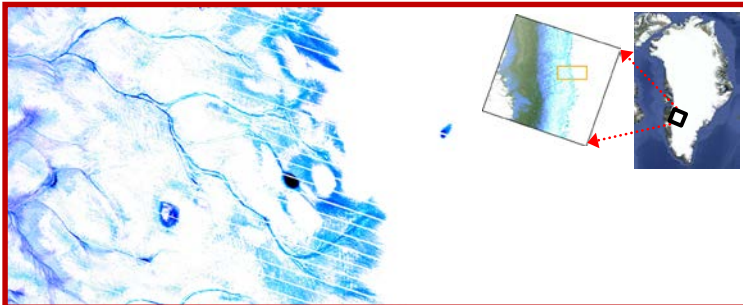
Scambos, 2006



Sonntag, 2017



Questions?



Slabs growing thicker: KAN-U, 2009-2017

