

The semi-annual oscillation in WACCM and other models participating in QBOi

plots from: Anne Smith, Laura Holt, Rolando Garcia

framework

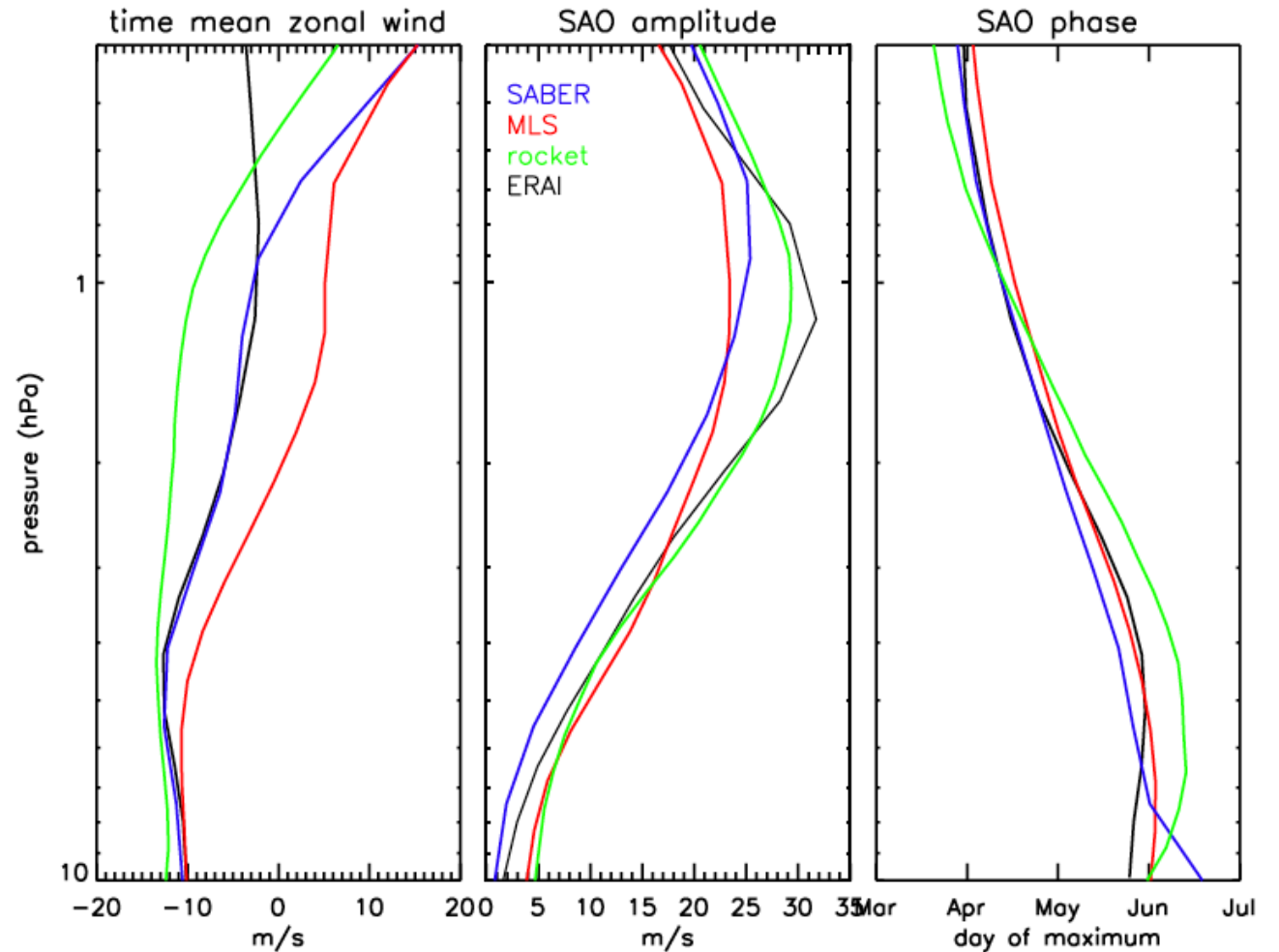
- QBOi: Quasi-Biennial Oscillation initiative
- Goal is an intercomparison of the QBO in global models that simulate a self-generated QBO
- Includes WACCM and CAM (special versions with higher vertical resolution up to the middle stratosphere).
- Most of the models have upper boundaries in or above the mesosphere.
- How well do they simulate the semi-annual oscillation (SAO)?

limitations on evaluating the SAO

- model output has been interpolated to a common grid; highest level 0.4 hPa (~55 km)
- Not all model groups supplied diagnostic products such as EP flux divergence, parameterized gravity wave drag, etc.
- Observations for validation of the SAO and the processes that drive it are scarce.

SAO

- variation of equatorial zonal mean wind extending from upper stratosphere to upper mesosphere
- poorly constrained by observations and reanalyses
- traditionally divided into stratopause SAO (focus of current analysis) and mesopause SAO



SABER: wind derived from geopotential (2002-2016)

MLS: wind derived from geopotential (2004-2016)

rocket: tropical measurements from 1960s and 1970s

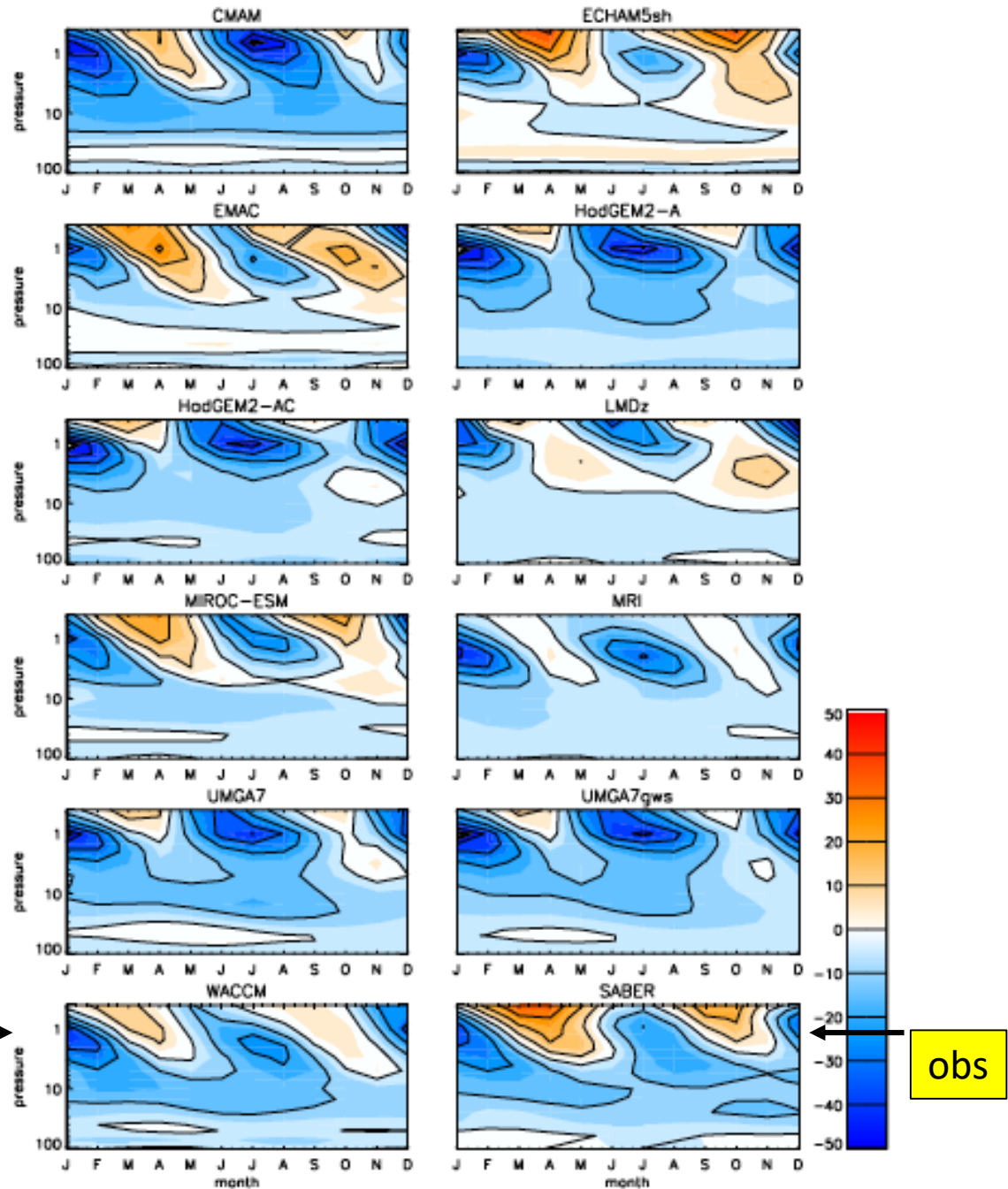
ERAi: reanalysis

Seasonal wind in QBOi models (30-year average)

10 other
models or
model versions

All models simulate an
SAO. At the stratopause,
there are easterly winds
at solstices, westerly (or
weaker easterly) winds at
equinoxes.

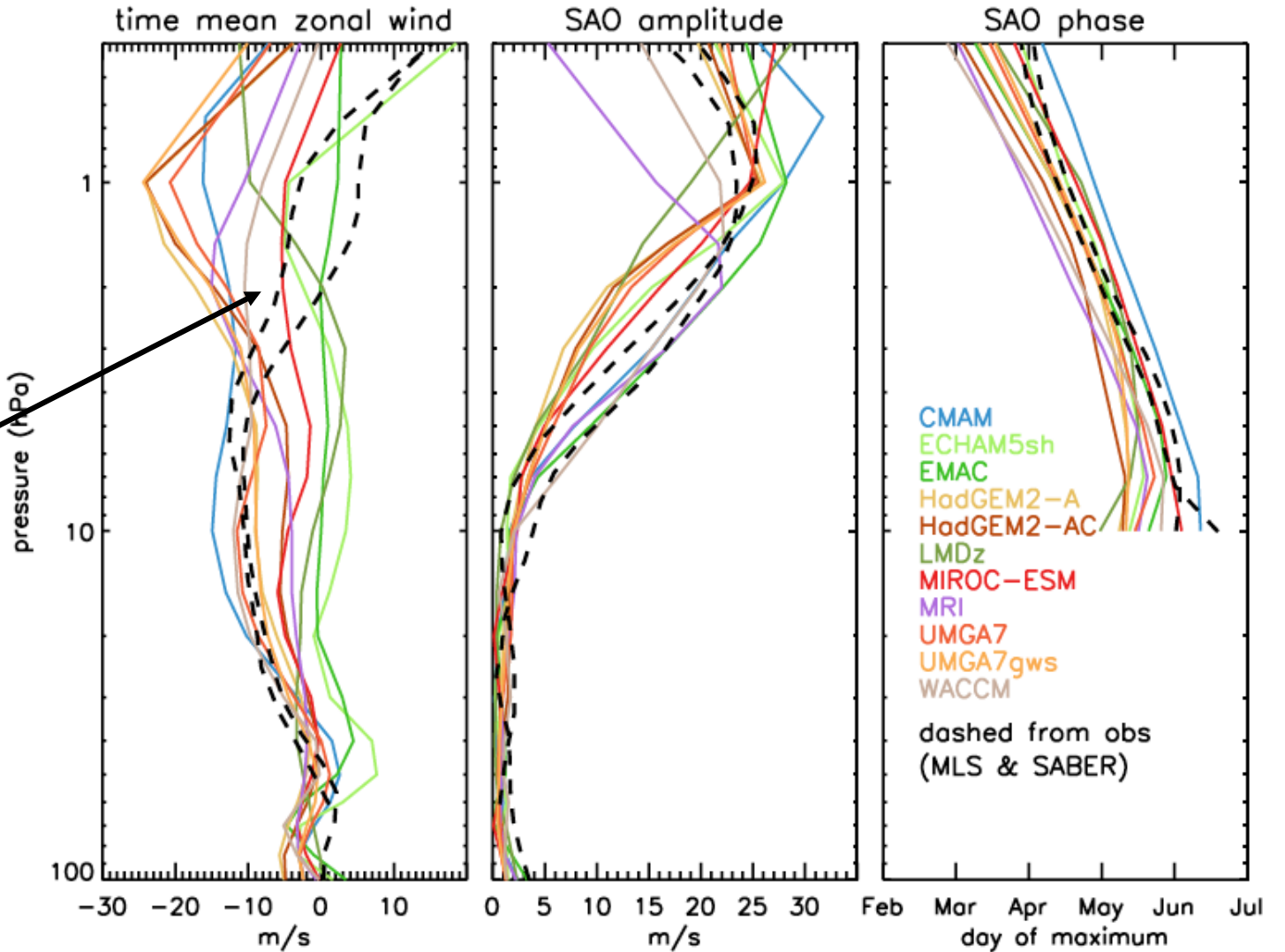
WACCM



SAO and time-mean wind

Models do a better job with the SAO than with the time-mean wind.

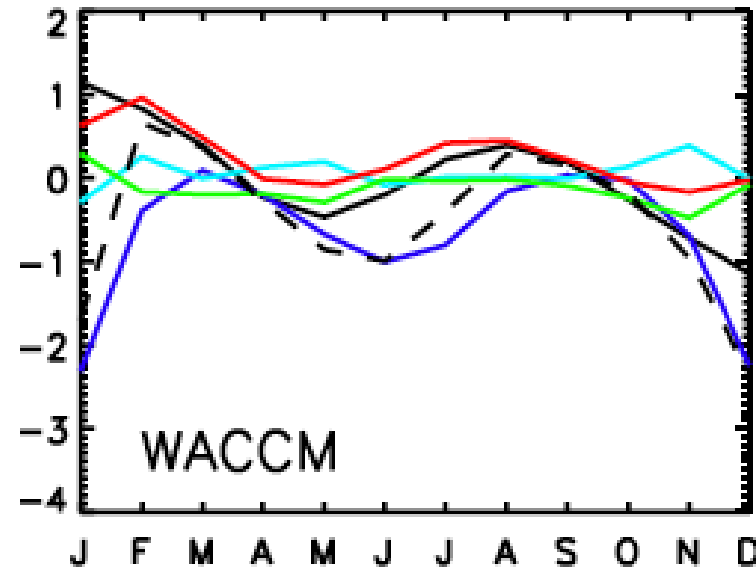
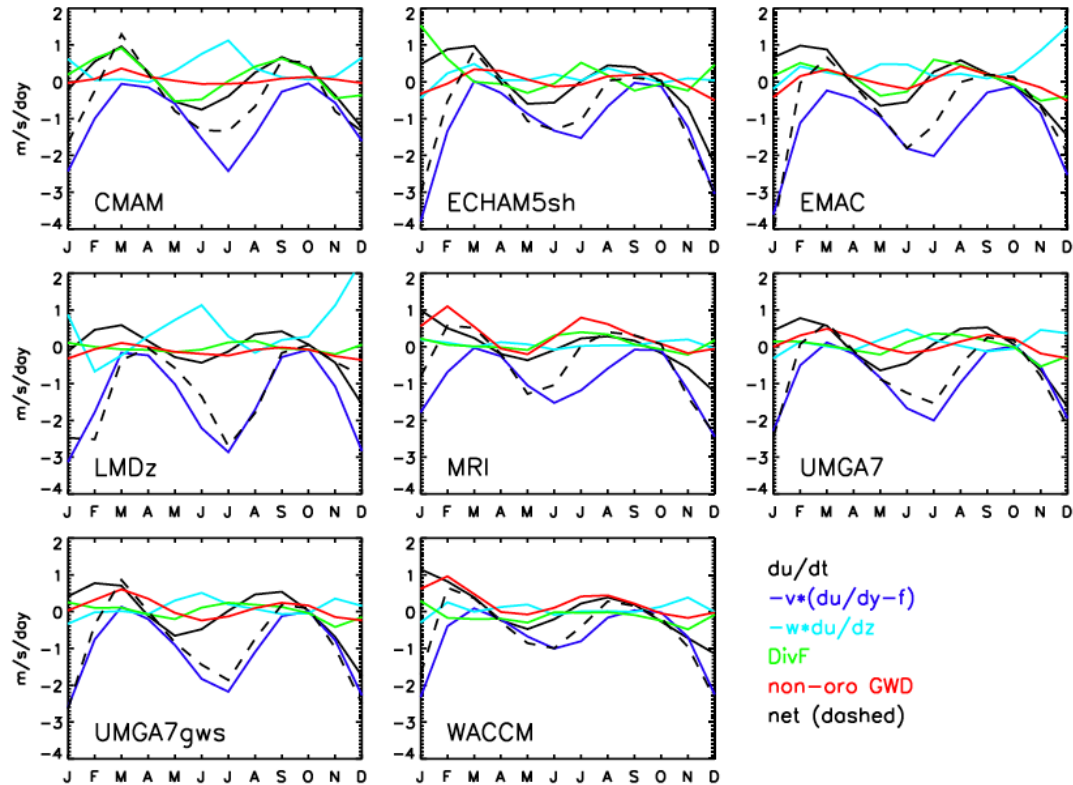
Particularly, the westerly shear in the upper stratosphere is lacking.



processes driving the SAO in transformed mean

$$\frac{\partial \bar{u}}{\partial t} = \underbrace{-\bar{v}^* \left[\frac{1}{a \cos \varphi} \frac{\partial \bar{u} \cos \varphi}{\partial \varphi} - f \right]}_{\text{advection by BDC}} - \underbrace{\bar{w}^* \frac{\partial \bar{u}}{\partial z}}_{\text{resolved waves}} + \underbrace{\frac{1}{\rho_0 a \cos \varphi} \nabla \cdot \mathbf{F}}_{\text{parameterized waves}} + \bar{X}$$

seasonal variation at 1 hPa



SAO timing controlled by BDC advection.

Net forcing by resolved waves is weak.

GW forcing is stronger but still weaker than BDC.

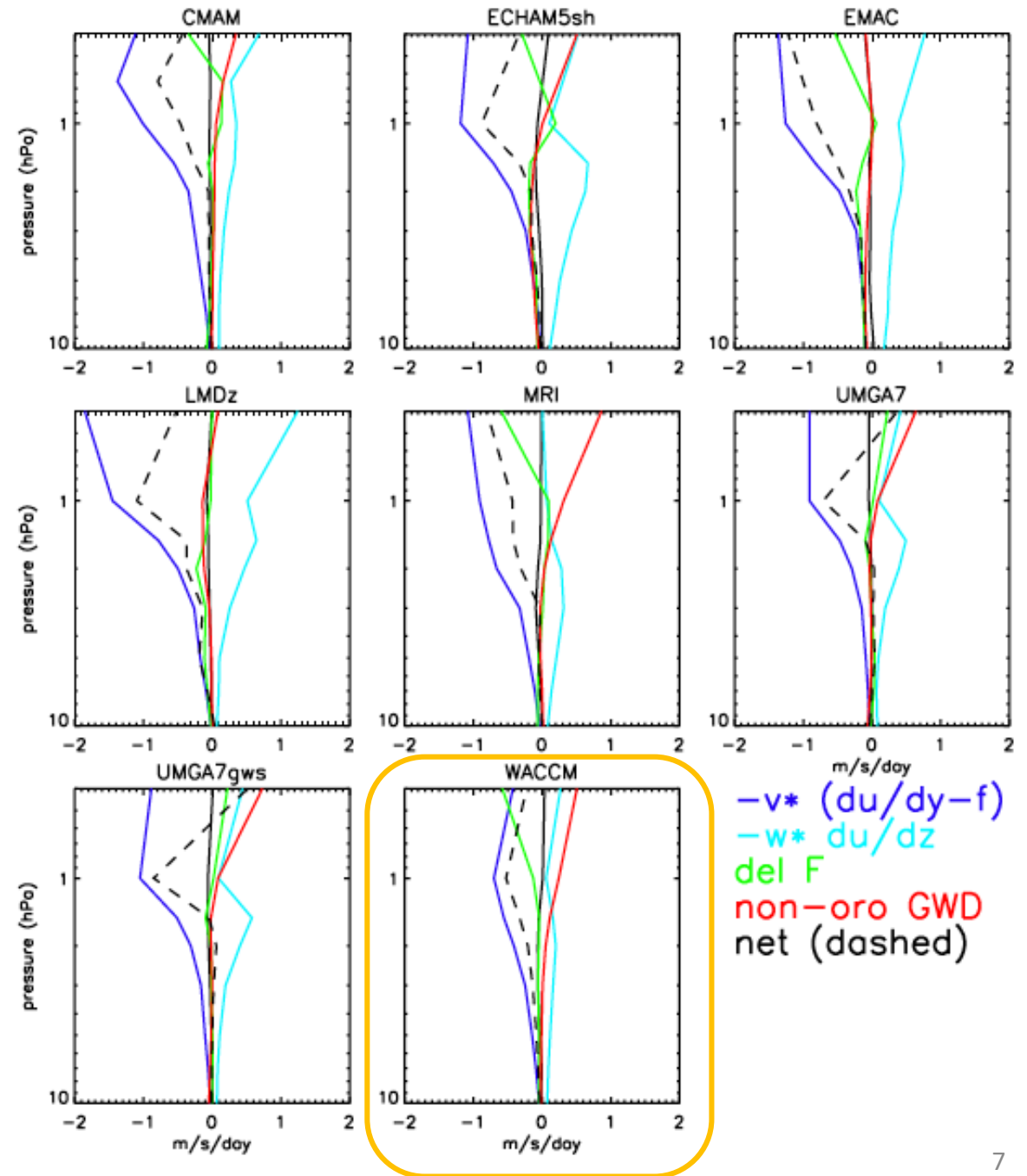
time mean forcing

imbalance due to:

- missing forcing (diffusion)?
- approximation due to interpolation to QBOi vertical grid?
- use of monthly mean for calculating BDC terms?

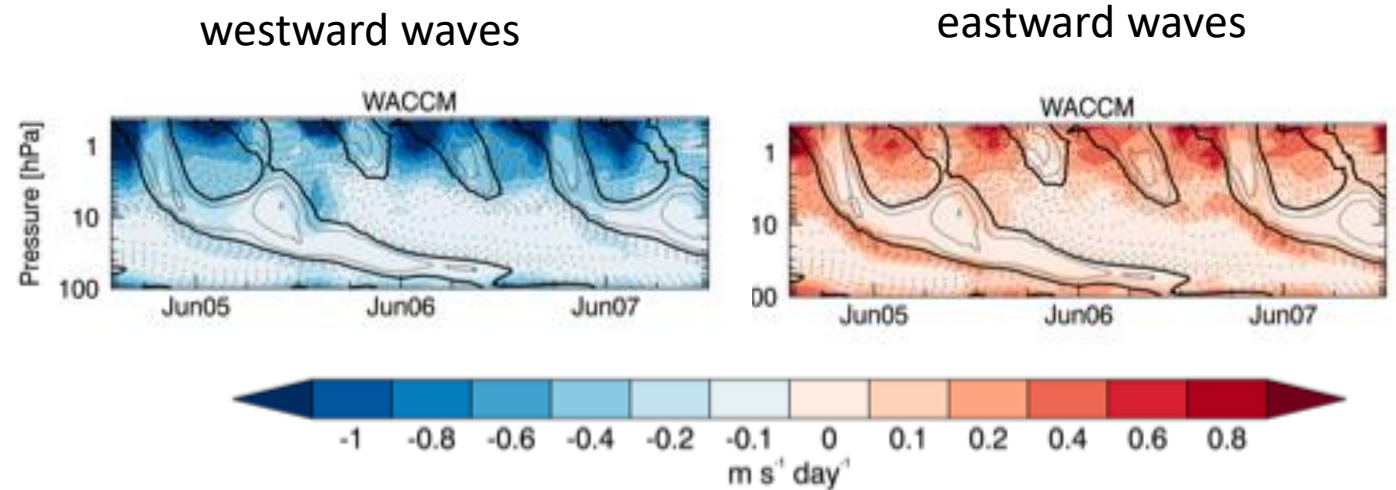
All models, including WACCM, have weak wave forcing. In WACCM, net EP flux divergence is negative.

Where is the forcing by fast Kelvin waves?



next step....

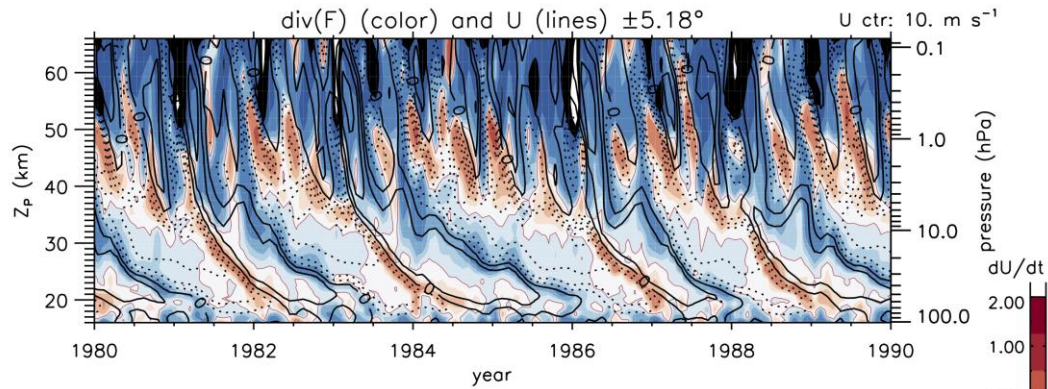
- Laura looked at wave forcing in the QBOi models using output at every 6 hrs.
- Rolando looked in detail at the wave forcing in the full WACCM output (i.e., daily; not interpolated to standard QBOi levels).



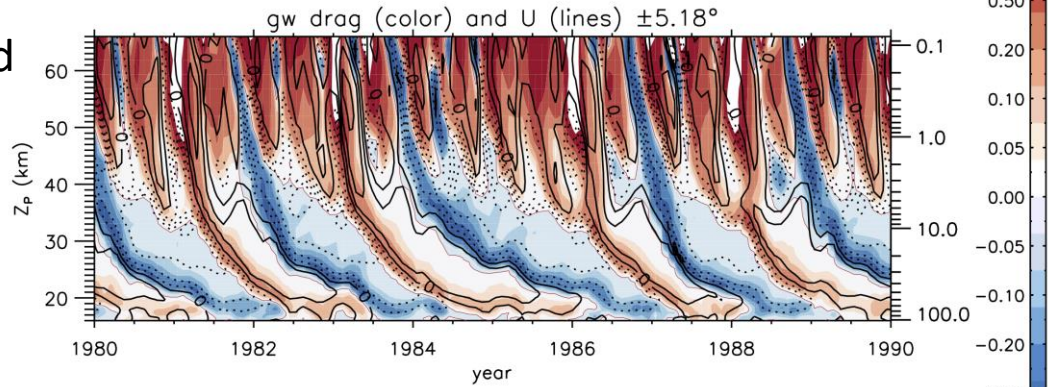
There is a high degree of cancellation since the westward & eastward waves deposit momentum at same locations and times.

WACCM wave forcing at the Equator ($\pm 5^\circ$ average)

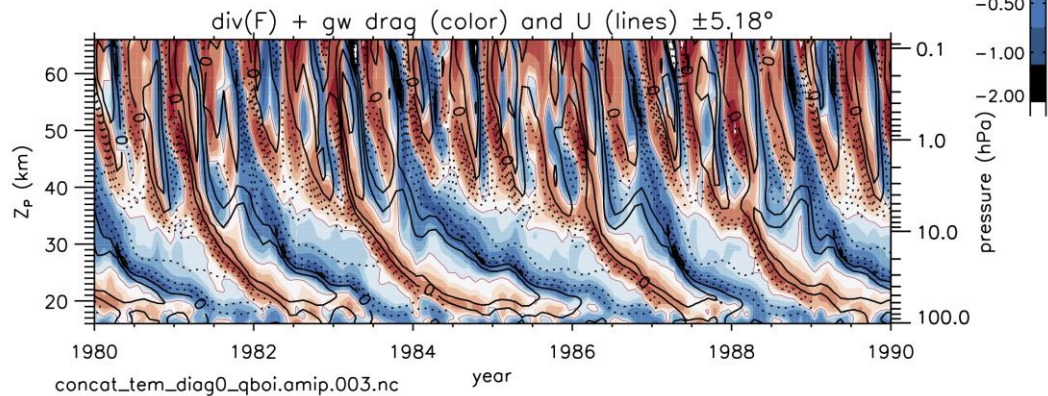
resolved waves



parameterized gravity waves



sum

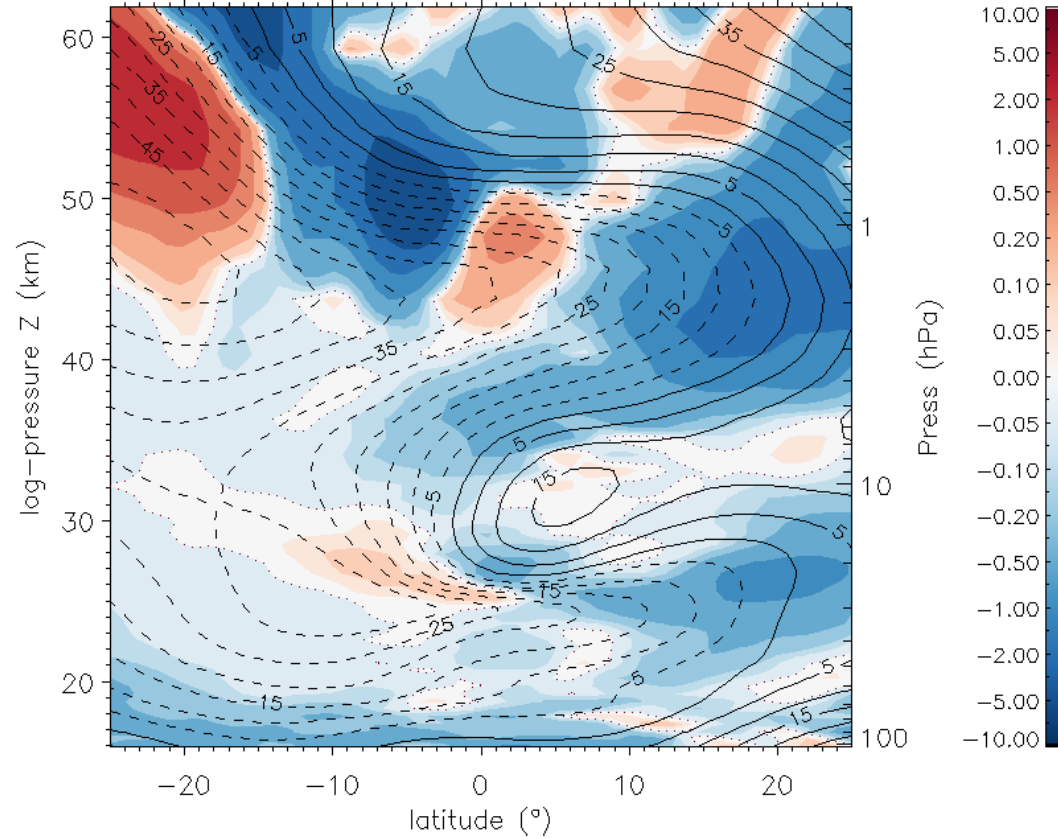


- most of the eastward acceleration in the SAO region is due to GW
- above 50 km, $\text{div}(F)$ is mainly negative
- this suggests Kelvin waves are not doing much of anything to produce the W phase of the stratopause SAO in WACCM

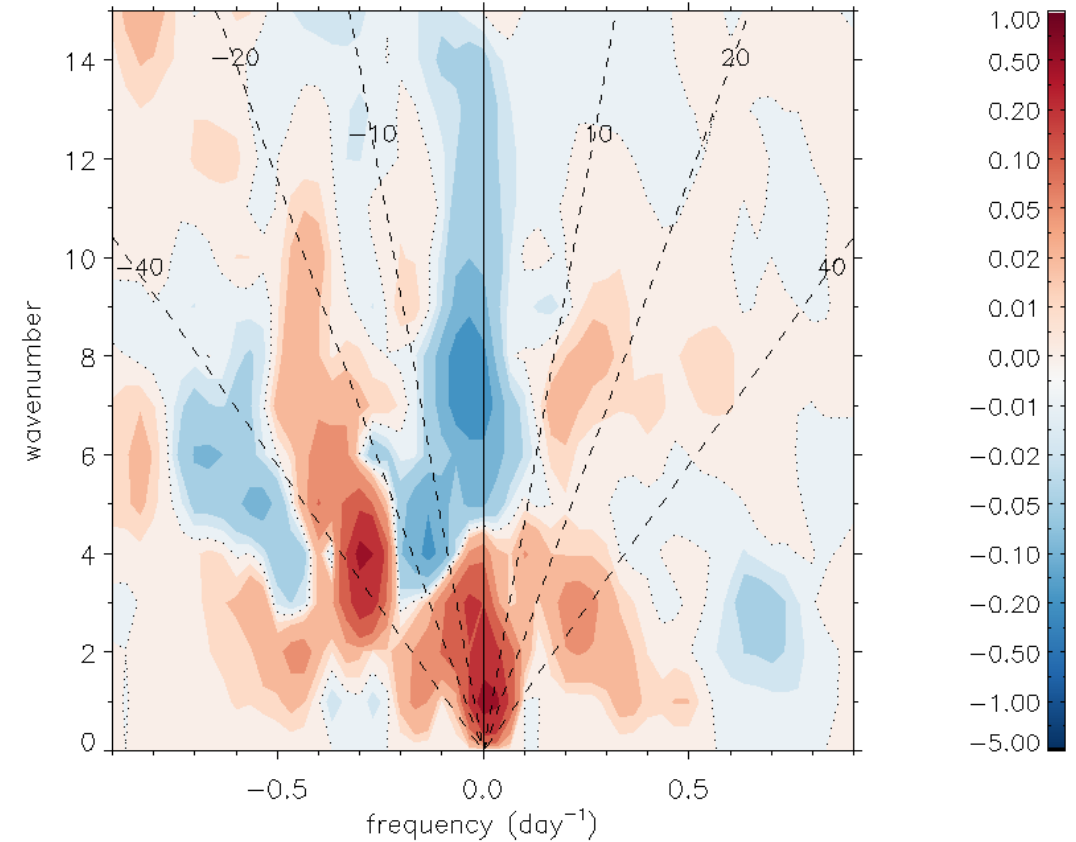
SAO in February 1993, WACCM-chem run

div(F) spectral density ($\text{ms}^{-1} \text{wno}^{-1}$) 19930215

div(F) (ms^{-2}) (-0.900, -0.033) cpd, all k, 19930215



mean(-2.4°:5.2°) 47.6 km (1.1 hPa)



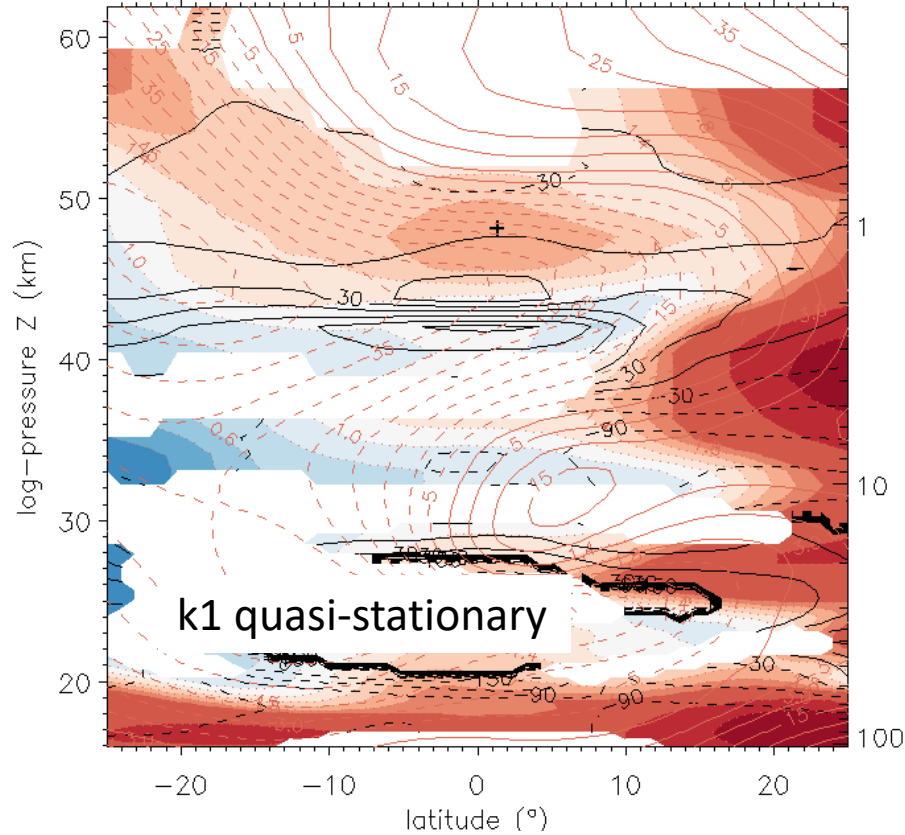
- the only near-equatorial $\text{div}(F) > 0$ occurs in the region of easterly wind
- this appears to be part of a broader +/- acceleration pattern, centered just S of Equator
- region of positive $\text{div}(F)$ is centered just N of Equator

- positive $\text{div}(F)$ just N of the Equator is due mainly by quasi-stationary $k=1$ and 3-day westward $k = 4$

T' wave structures (U contours superimposed)

T Coh² Amplitude: (-0.900, 0.900) cpd

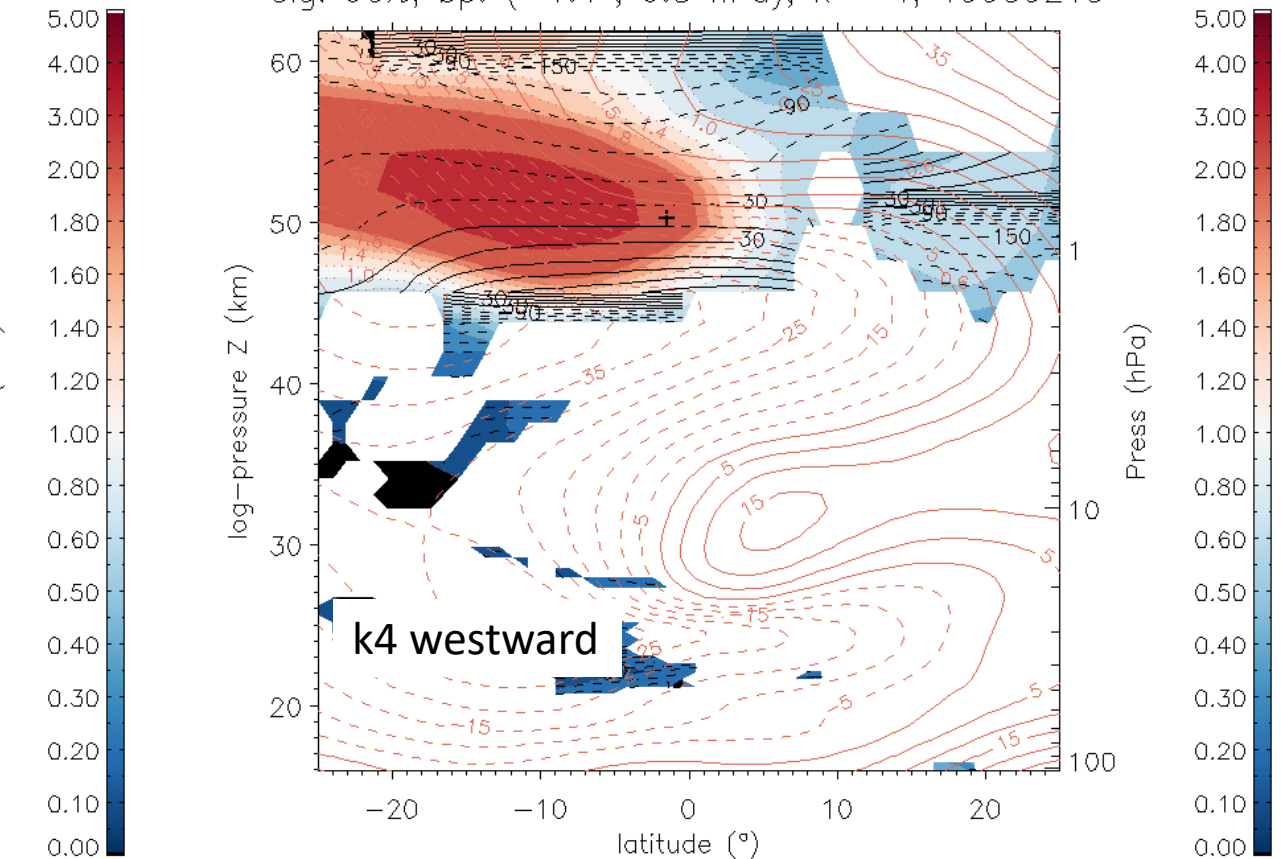
sig: 90%, bp: (1.4°, 1.1 hPa), k = 1, 19930215



- symmetric structure near Equator
- likely a Kelvin wave but not a “fast” Kelvin wave
- subtropical structure looks like Rossby wave (likely unrelated to equatorial forcing)

T Coh² Amplitude: (-0.900, -0.033) cpd

sig: 90%, bp: (-1.4°, 0.8 hPa), k = 4, 19930215



- antisymmetric structure
- Rossby-gravity wave? not coherent with anything at lower altitude
- in situ generation?

Status

- This version of WACCM successfully simulates a good QBO. Unfortunately, the standard WACCM6 has lower vertical resolution and, as a result, the QBO has shortcomings, particularly at its lowest extension just above the tropical tropopause.
- The easterly phase of the stratopause SAO is driven by the Brewer-Dobson circulation (well-simulated).
- The westerly phase of the stratopause SAO in WACCM is driven by parameterized GW.
- Kelvin waves do not contribute much in the way of forcing to the SAO or the time-mean wind. Their contribution is weak compared to GW and even to other waves of uncertain origin.

Where do we go from here?

- MISSING: observational evidence about which waves are actually driving the eastward acceleration or the time-mean westerly winds and shear.
- Would improved vertical resolution in the upper stratosphere change the representation of resolved waves and/or their dissipation?
- Are fast Kelvin waves not forced efficiently in the troposphere in WACCM? This lack might be a result of the Zhang-McFarlane deep convection parameterization (cf. Ricciardulli and Garcia, 1999), which has short-term variability that is unrealistically low.
- Can the gravity wave parameterization be “tuned” to provide the missing eastward forcing?