

# ENSO diversity in CESM2

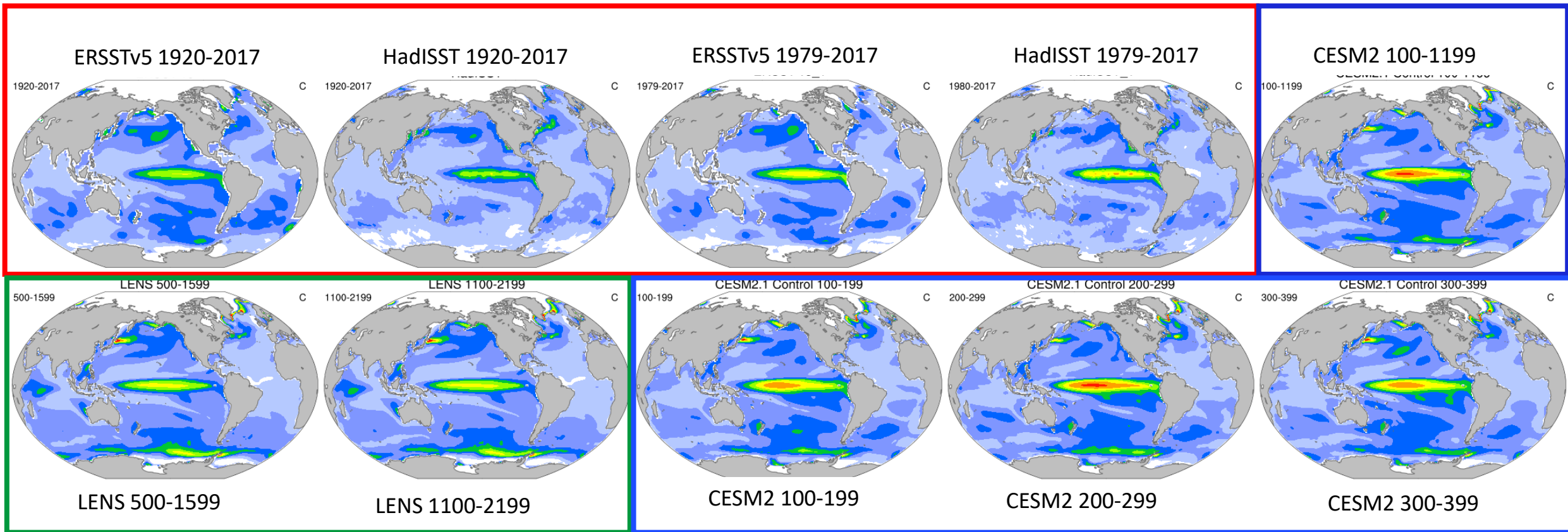
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ENSO is a fundamental mode of climate variability. It is very important for climate models to represent it realistically.

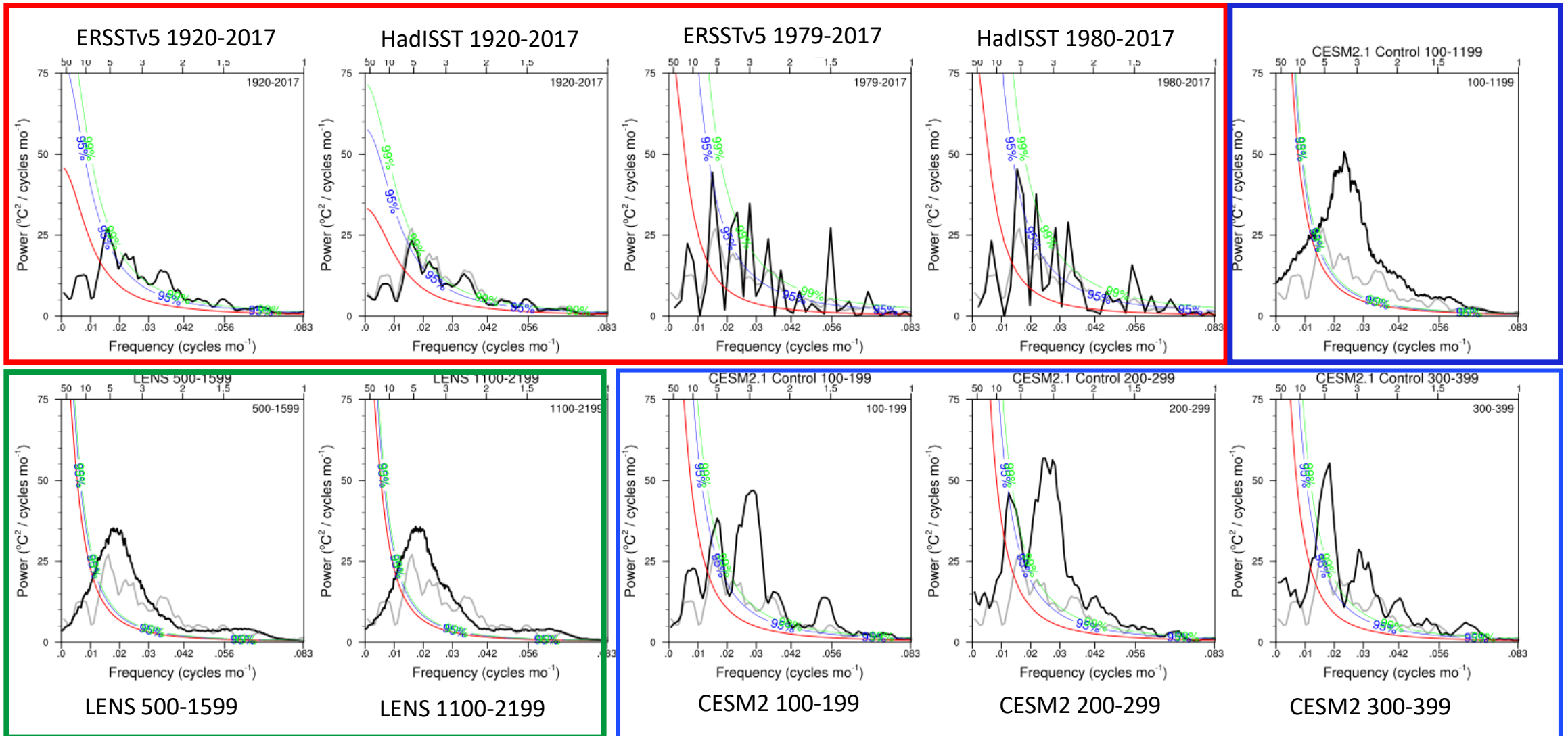
1. Basic ENSO characteristics in CESM2
2. Why do we care about ENSO diversity?
3. Examine basic characteristics of ENSO diversity using different indices

# SST standard deviation (DJF)



The westward displacement of the interannual SST anomalies is still pronounced in CESM2. CESM2 variance larger than LENS.

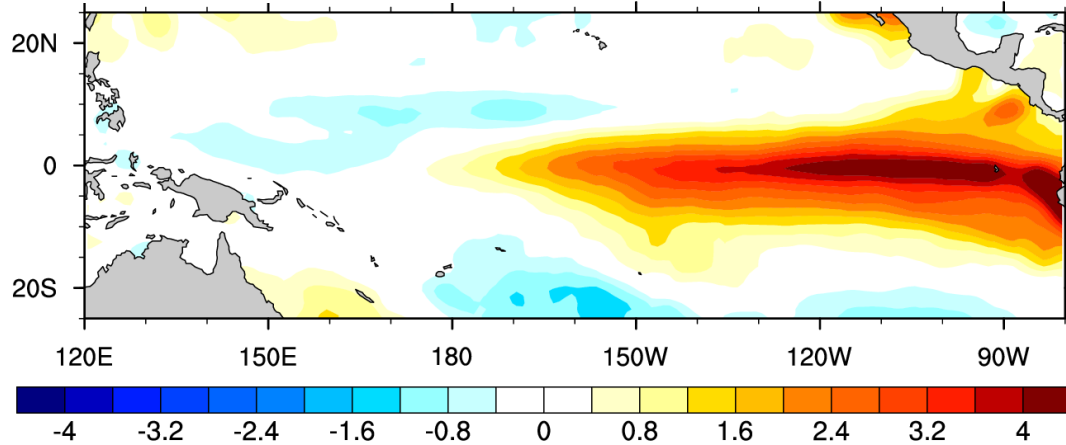
# Spectra of Nino3.4 (Monthly, detrended)



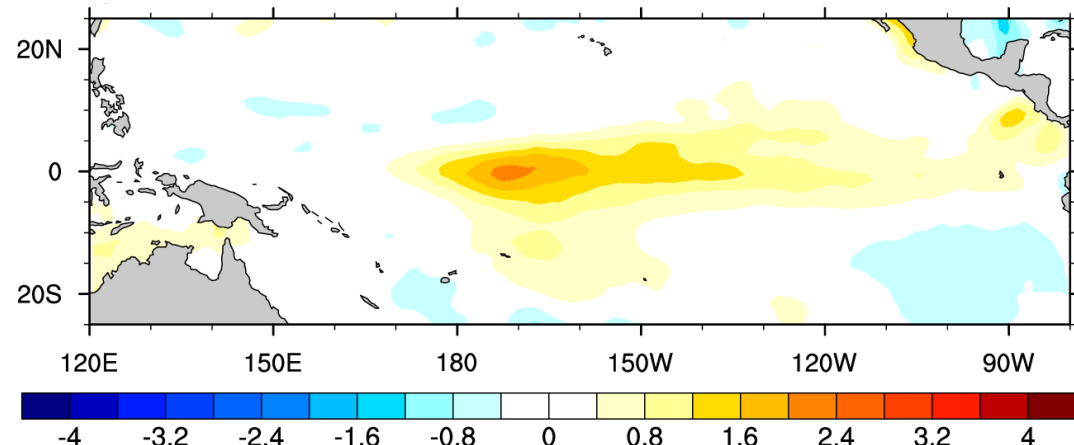
# El Nino events differ in amplitude, spatial pattern and evolution

## Examples of differences spatial pattern

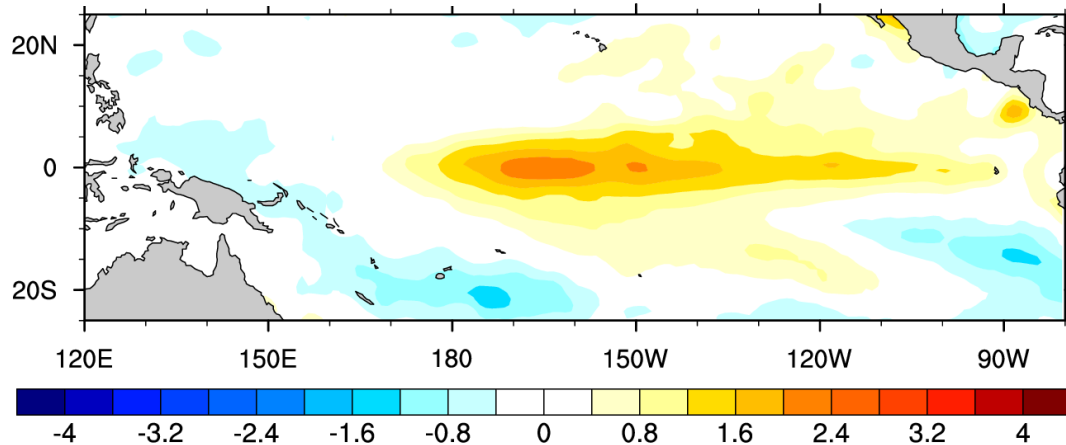
OISST DJF 1997/98



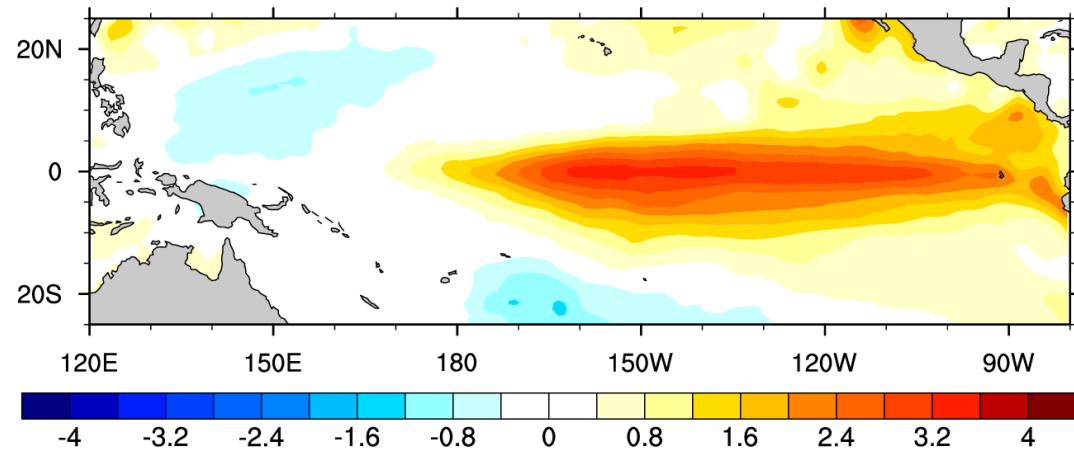
OISST DJF 2002/03



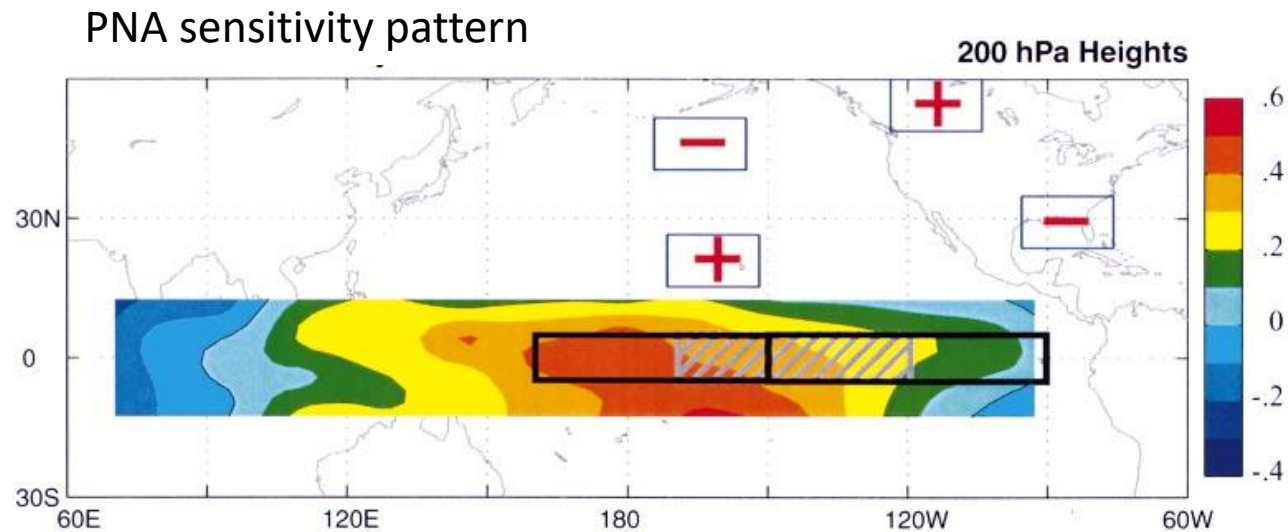
OISST DJF 2009/10



OISST DJF 2015/16



# Dependency of atmospheric teleconnections on the location of equatorial SST anomalies



Barsugli and Sardeshmukh 2002

PNA is most sensitive to SST anomalies in the central-western Pacific.

The most effective ENSO events to “excite” the PNA are those with a large projection on the sensitivity pattern

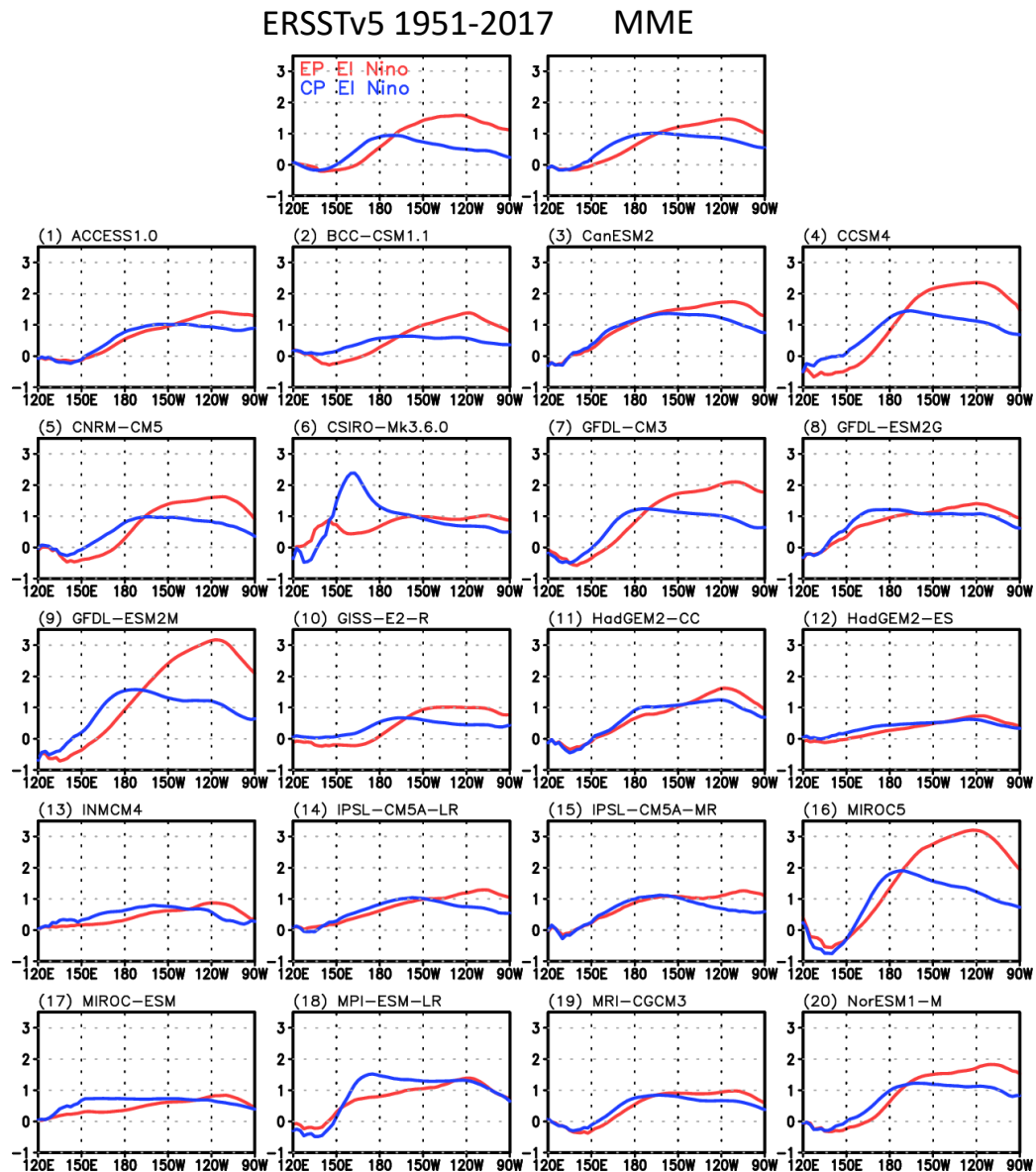
Relatively weak CP events may be as influential as strong EP events

## Metrics used for ENSO diversity classification

- El Niño Modoki index (Ashok et al. 2007)
- Niño3/Niño4 approach (Kug et al. 2009; Yeh et al., 2009)
- Niño3.4/Trans-Niño-Index (TNI, Trenberth & Stepaniak, 2001)
- EP/CP approach (Kao and Yu, 2009)
- EP/CP subsurface index method (Yu et al., 2011)
- $N_{CT}/N_{WP}$  indices (Ren and Jin, 2011)
- $EP_{new}/CP_{new}$  indices (Sullivan et al., 2016)
- E/C indices (Takahashi et al., 2011)
- Sea Surface Salinity indices (Singh et al., 2011; Qu and Yu, 2014)
- OLR-based indices (Chiodi and Harrison, 2010; Johnson and Kosaka, 2016; Williams and Patricola, 2018)
- Spatio-temporal indices (Lee et al. 2014)



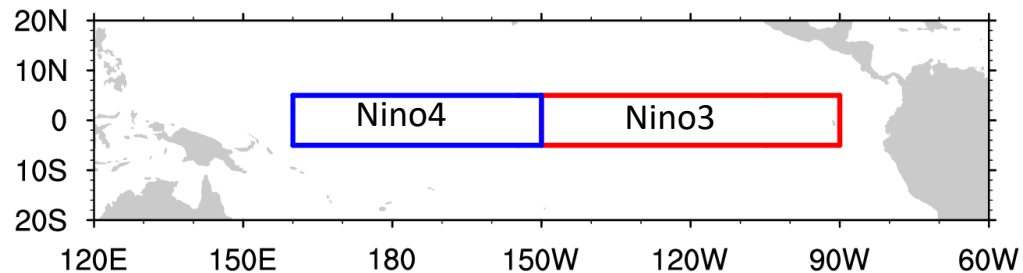
# ENSO diversity in the CMIP5 models



EP and CP events have been identified using the Nino3 and Nino4 indices.

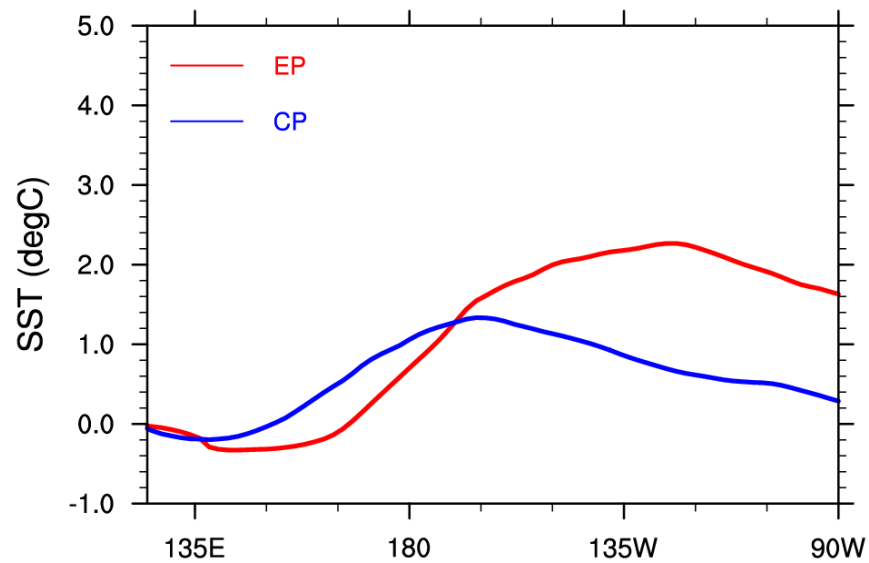
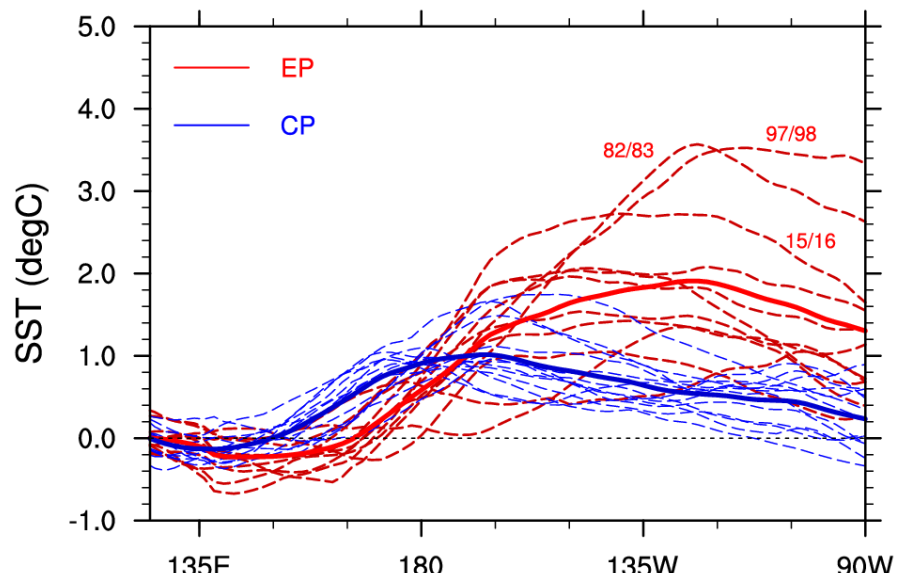
EP = Nino3 > 0.5°C and > Nino4

CP = Nino4 > 0.5°C and > Nino3

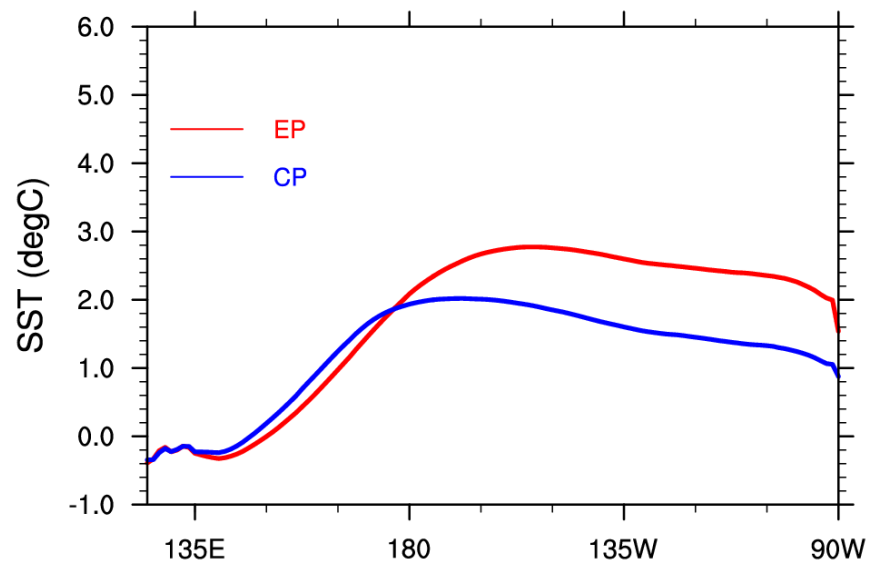
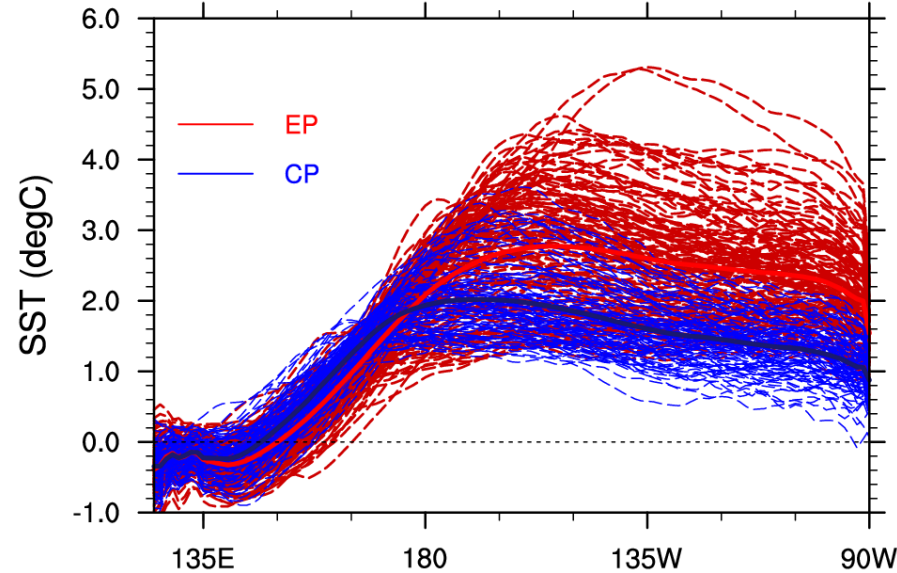


Only few models capture the diversity in the longitudinal profile of tropical SST anomalies.

# What about CESM2?



ERSSTv5  
1951-2017



CESM2  
1-1200

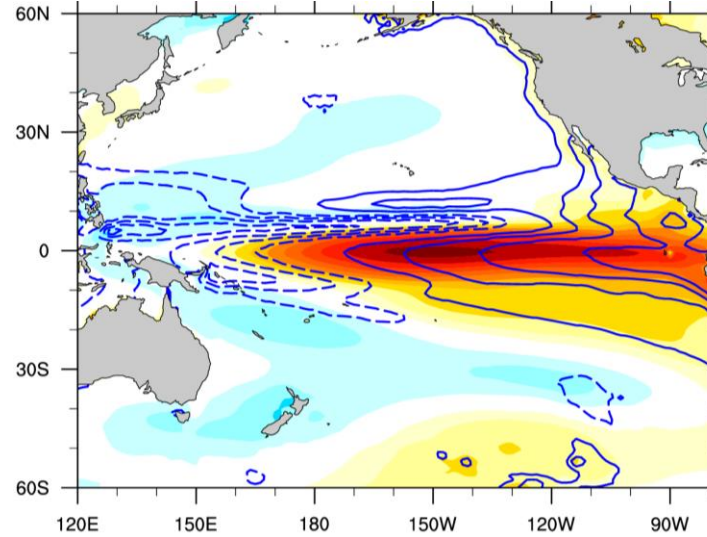
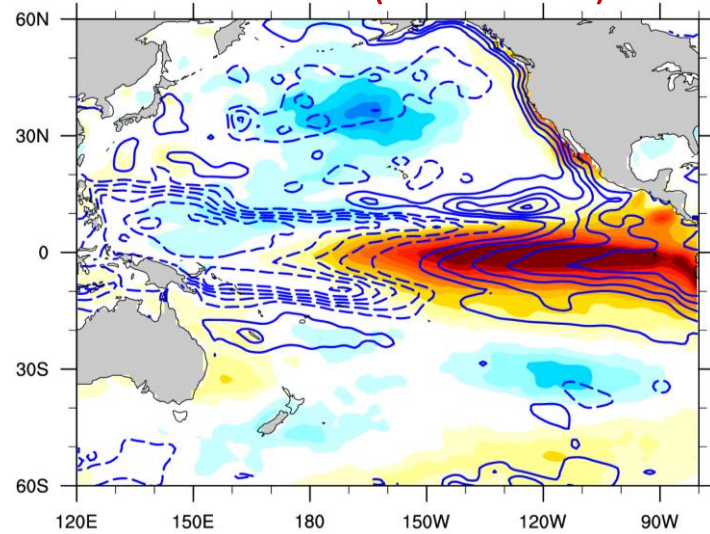


# SST and SSH composites for EP and CP events (Niño3/Niño4 approach)

ORAS4 (1958-2015)

CESM2 (1-1200)

EP

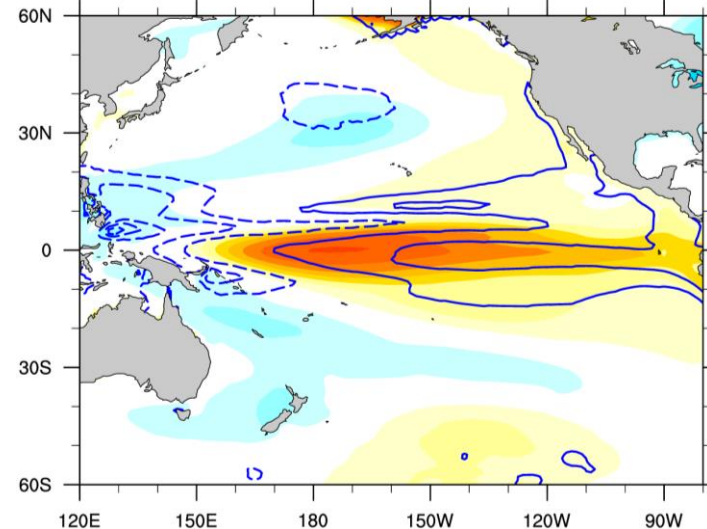
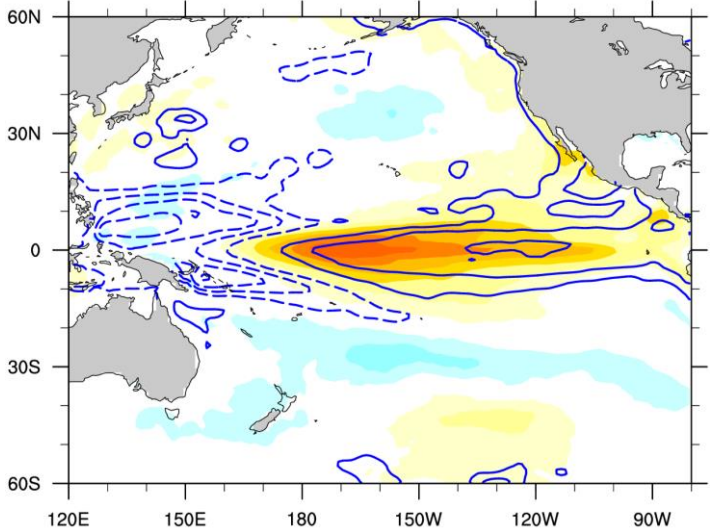


SSH (thermocline depth) anomalies larger during EP events

SSH anomalies more intense along the US West Coast during EP events

In CESM2 main difference between EP and CP is in the eastern equatorial Pacific

CP

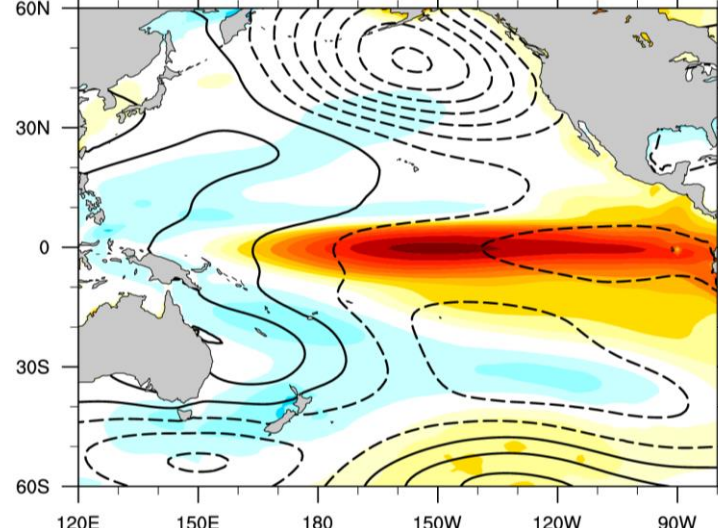
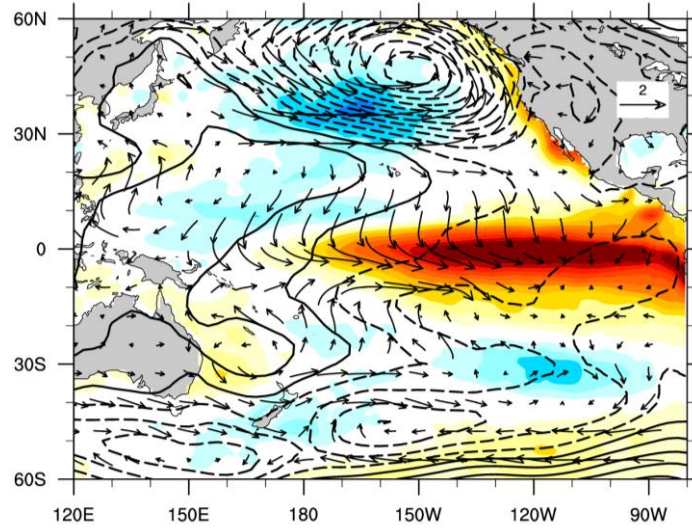


# SST and SLP composites for EP and CP events (Niño3/Niño4 approach)

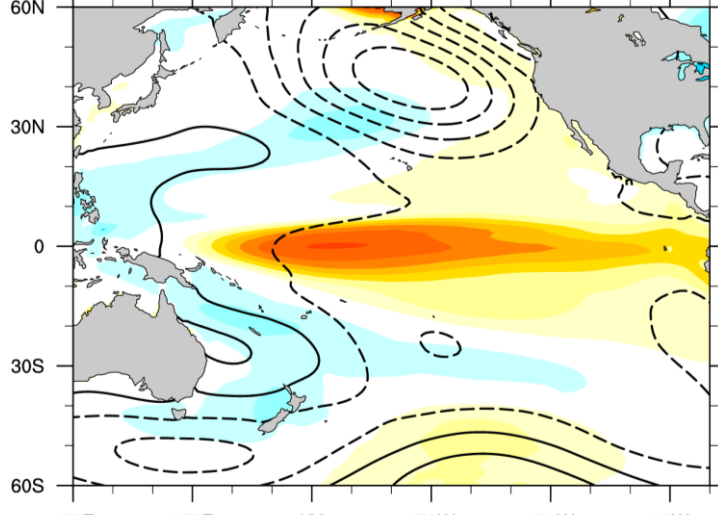
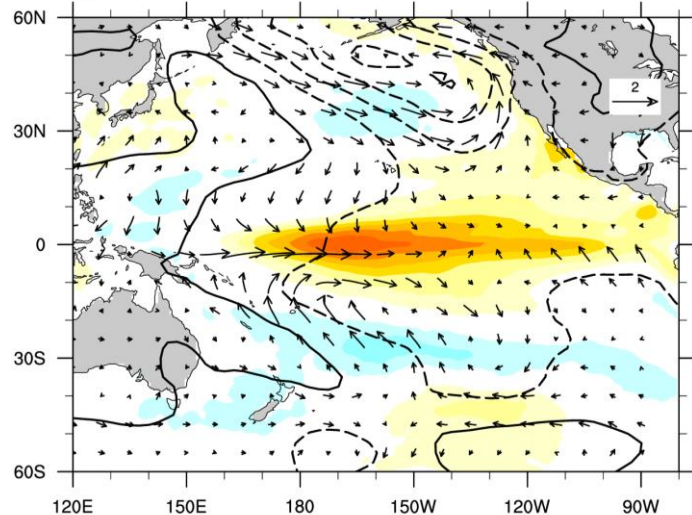
ORAS4, NCEP/NCAR

CESM2

EP



CP



Zonal gradient of SLP (and surface winds) along the equator is confined further to the west during CP events than during EP events. SLP zonal gradients are displaced to the west relative to obs for both EP and CP

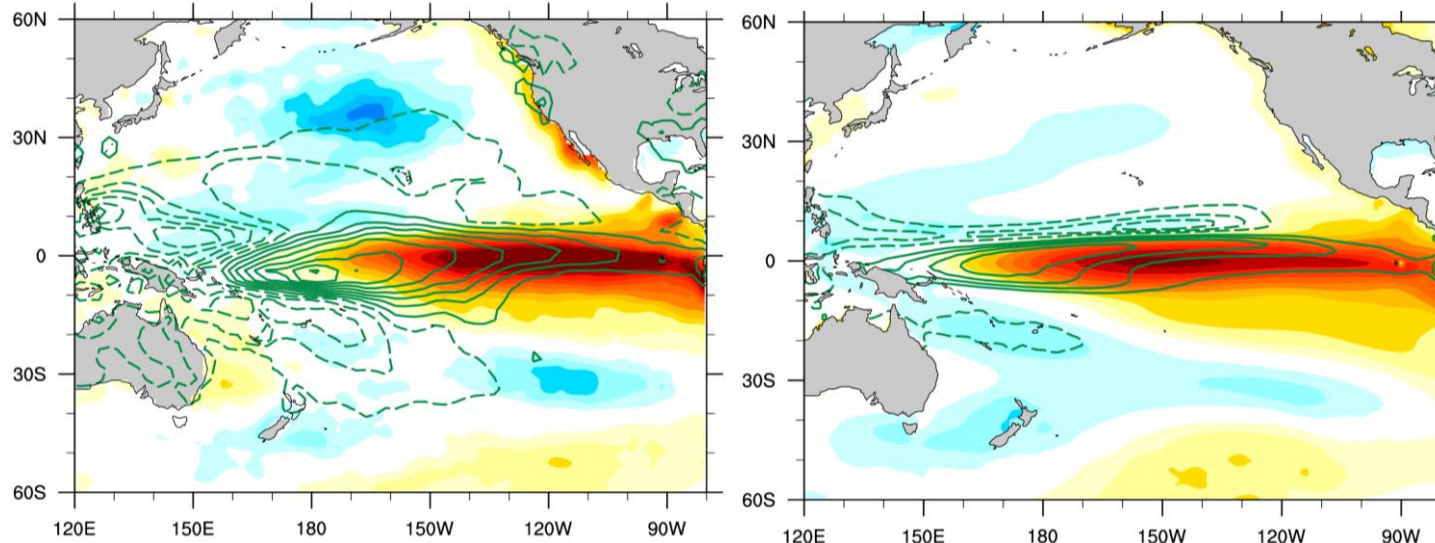


# SST and Precip composites for EP and CP events (Niño3/Niño4 approach)

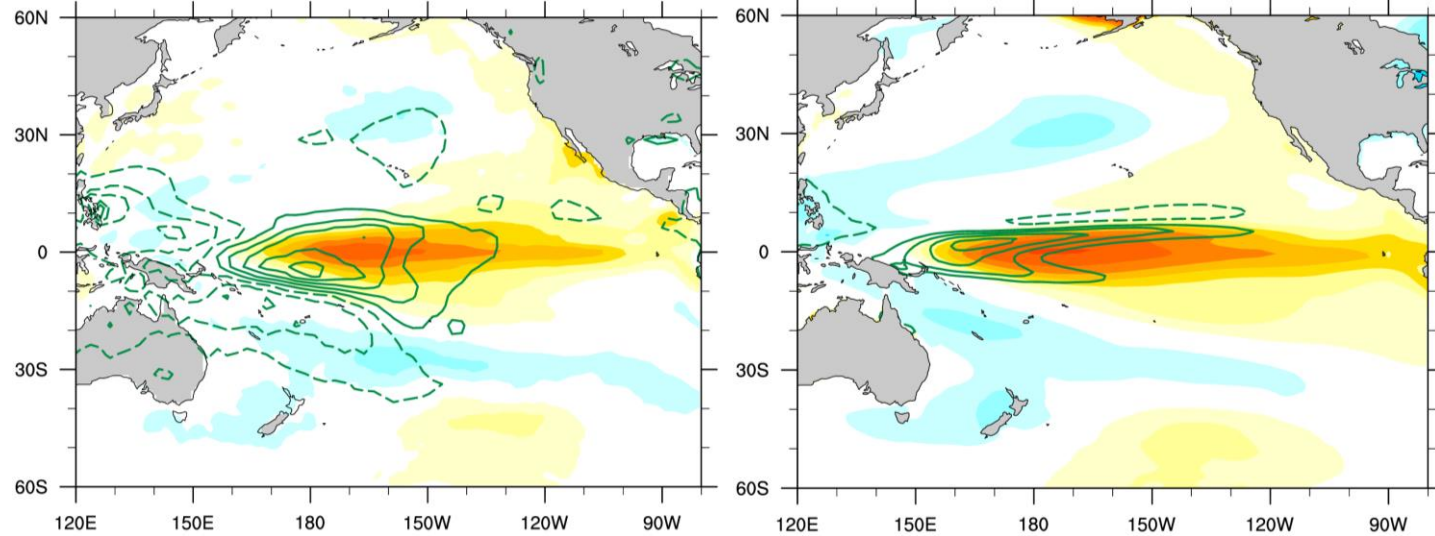
ORAS4/NOAA reconstructed

CESM2

EP



CP



Precip anomalies along the equator are confined further to the west during CP events than during EP events. SLP zonal gradients are displaced to the west relative to obs for both EP and CP

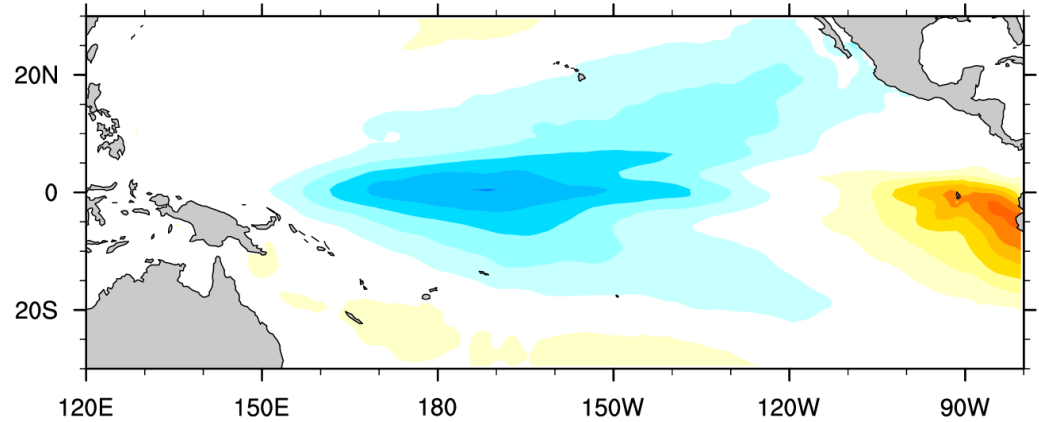
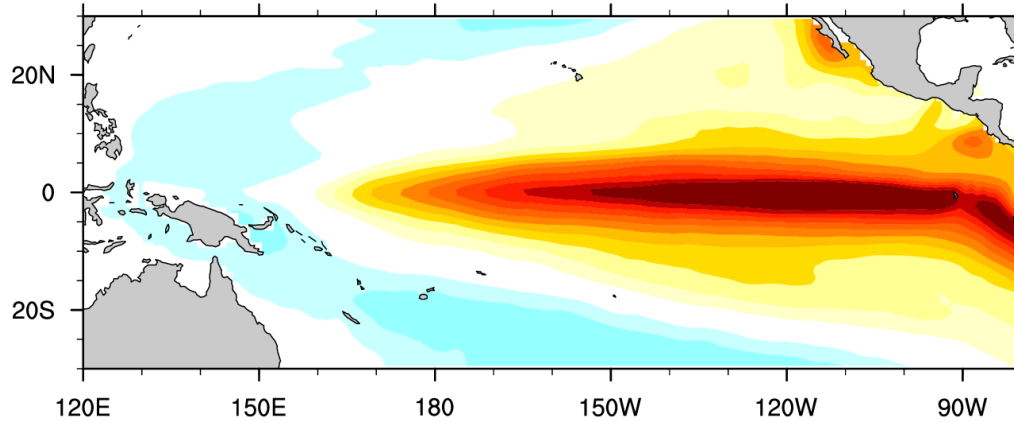
Two degrees of freedom are needed to capture the evolution of interannual SST anomalies along the equator (Trenberth and Stepaniak, 2001)

$$\text{TNI} = \text{Nino1} + 2_N - \text{Nino4}_N$$

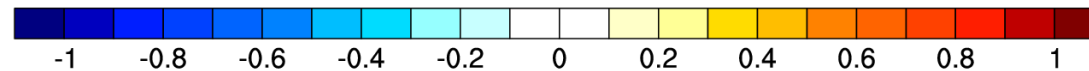
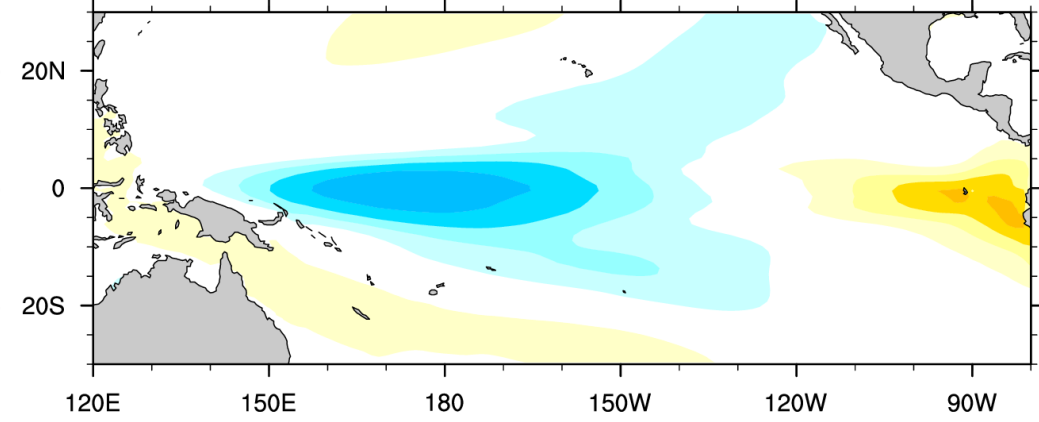
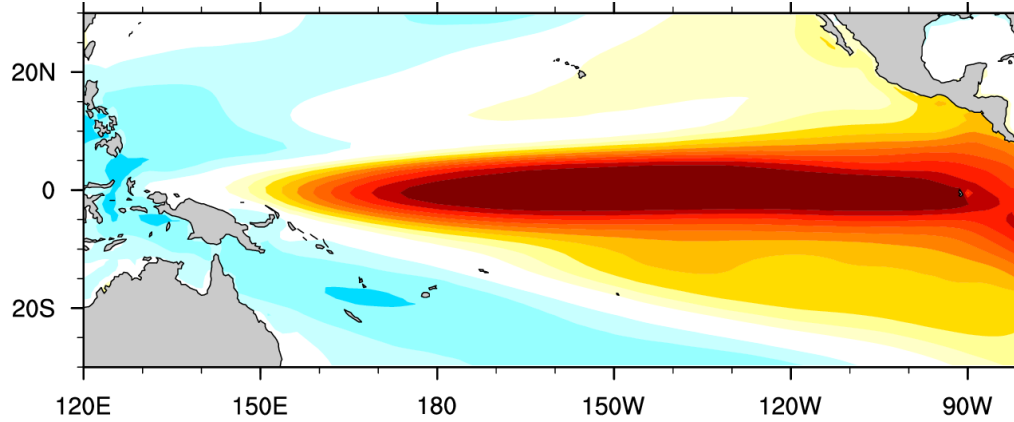
Nino3 –SST regression

TNI –SST regression

ORAS4



CESM2

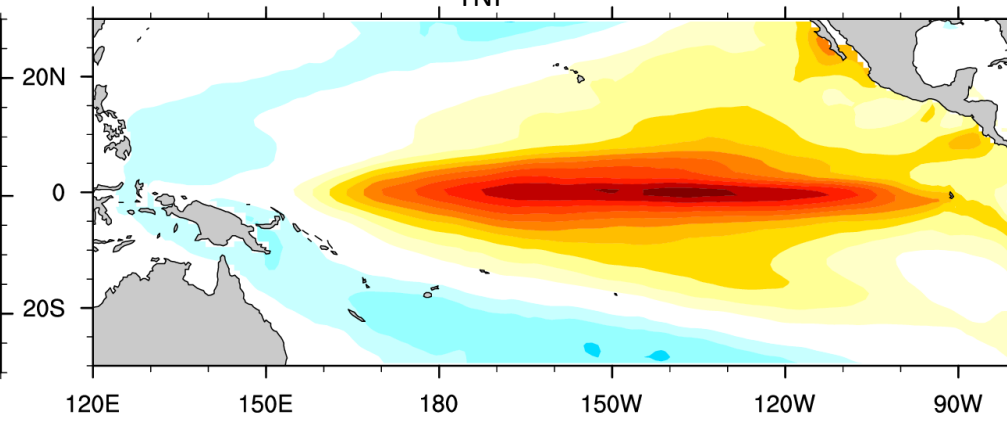
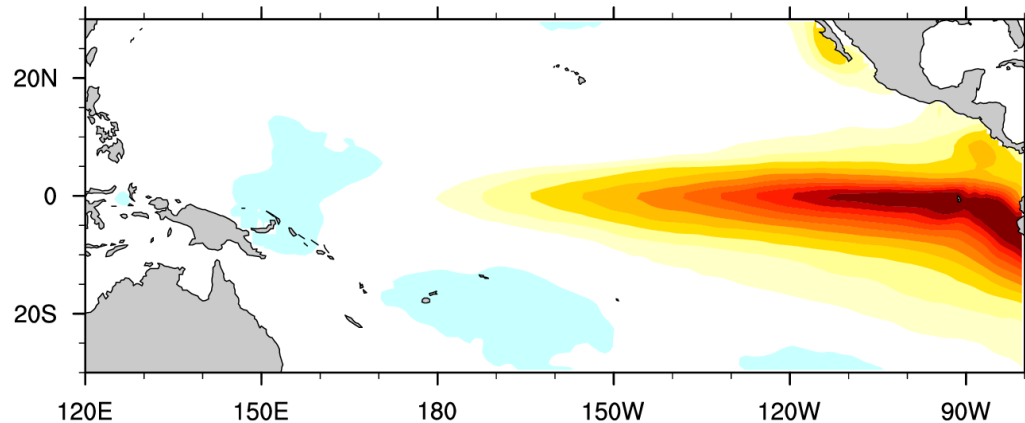


# Linear combination of Nino3 and TNI can produce EP and CP spatial patterns

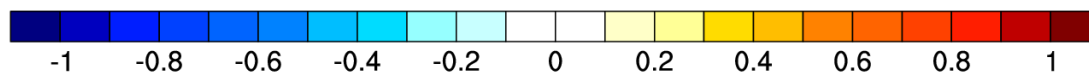
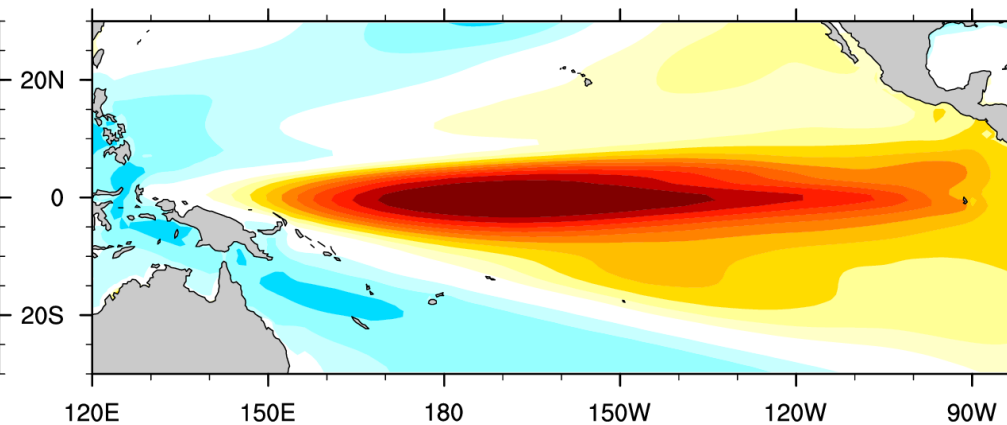
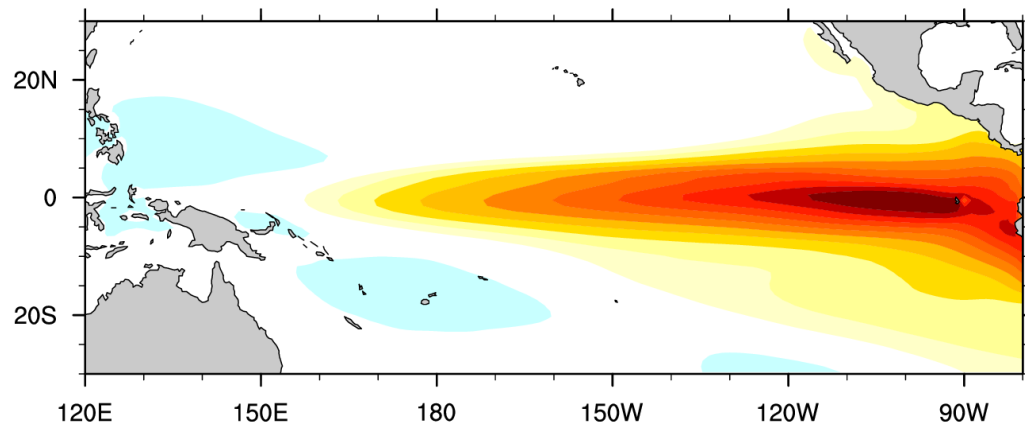
$$EP_{TNI} = Nino3 + TNI$$

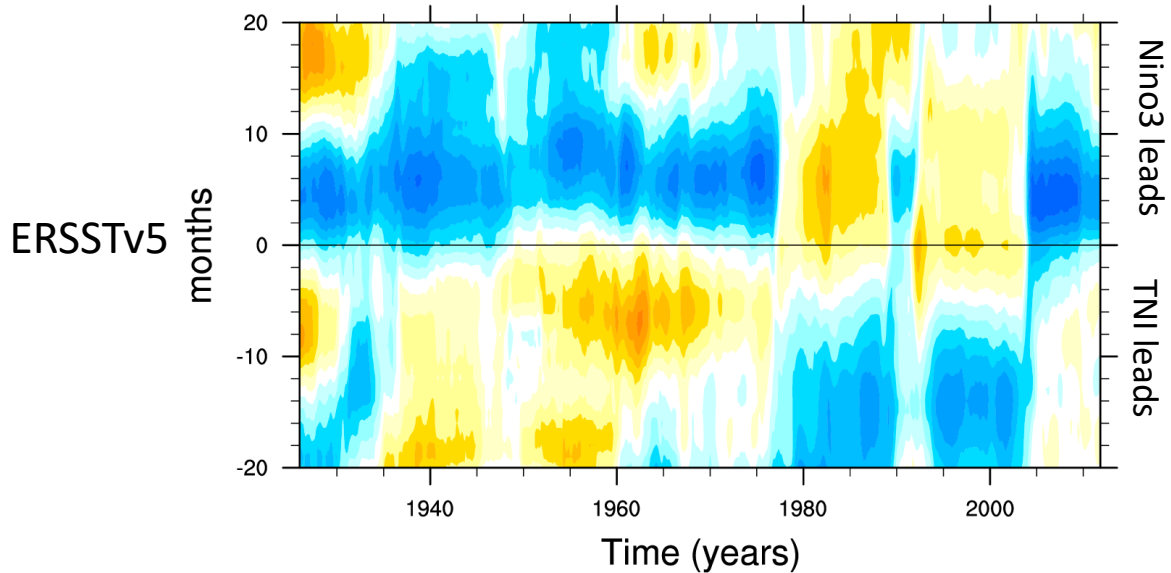
$$CP_{TNI} = Nino3 - TNI$$

ORAS4



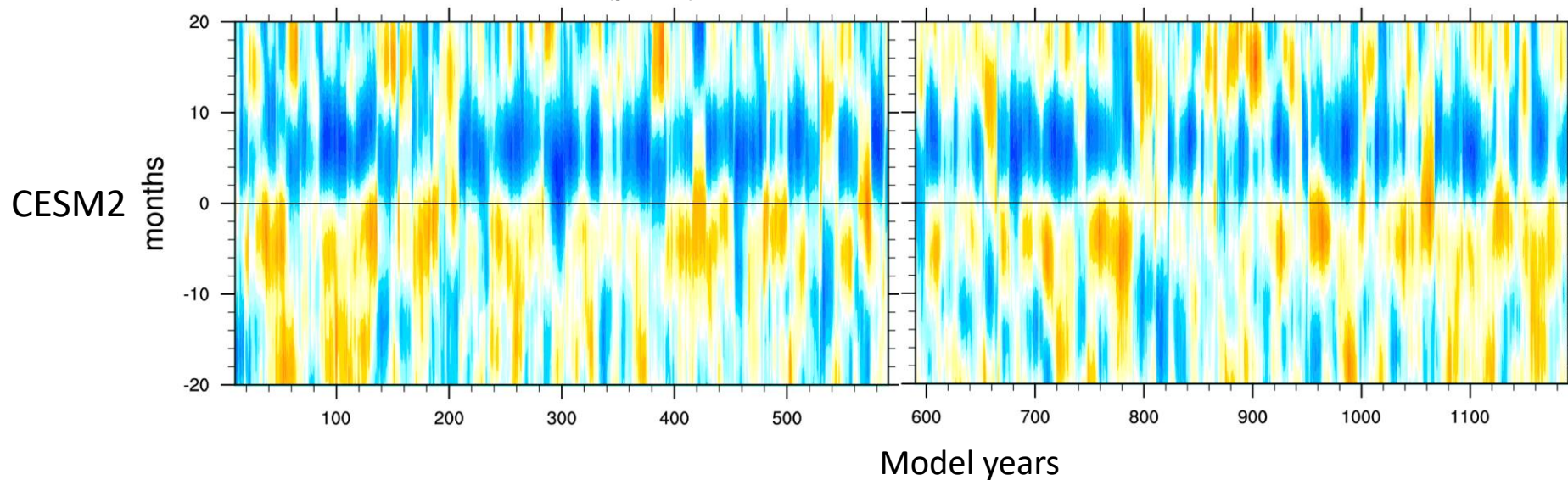
CESM2



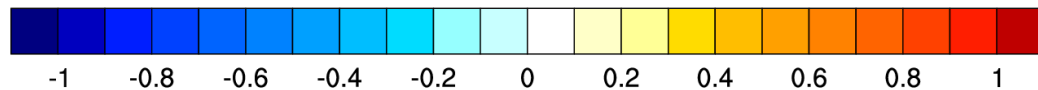


12-year sliding lag-correlation between Nino3 and TNI show anomaly propagation

Negative Corr. = Westward propagation  
 Positive Corr. = Eastward propagation

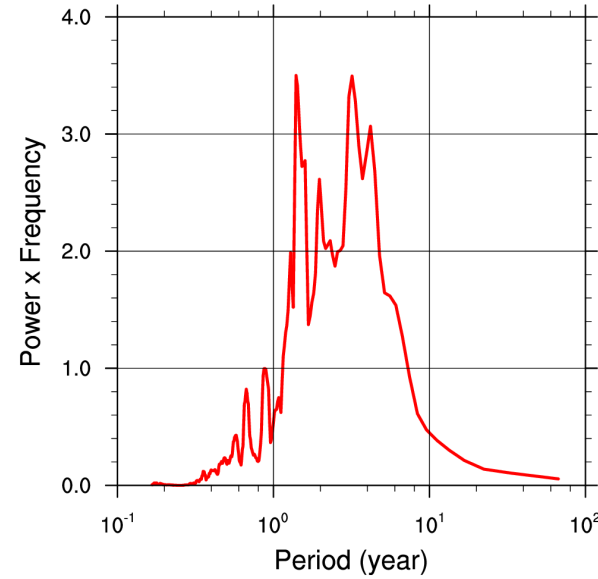
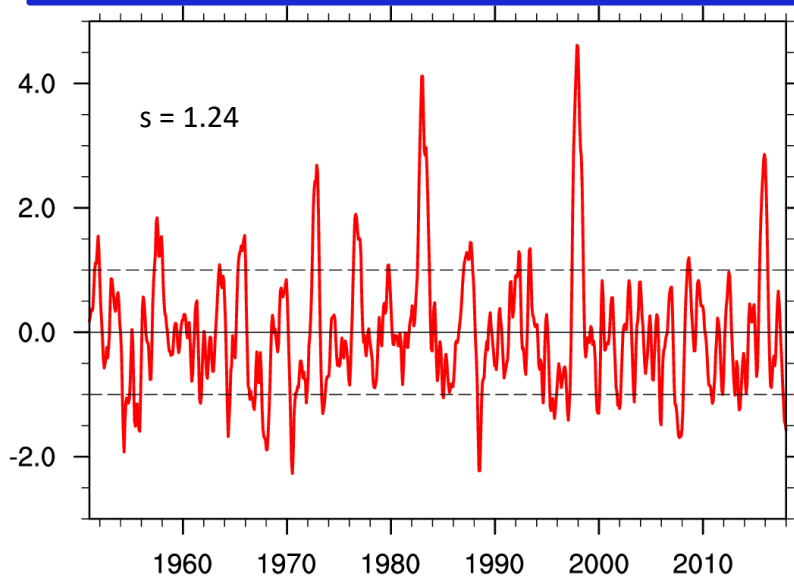


Propagation is predominantly westward in the model, similar to observations

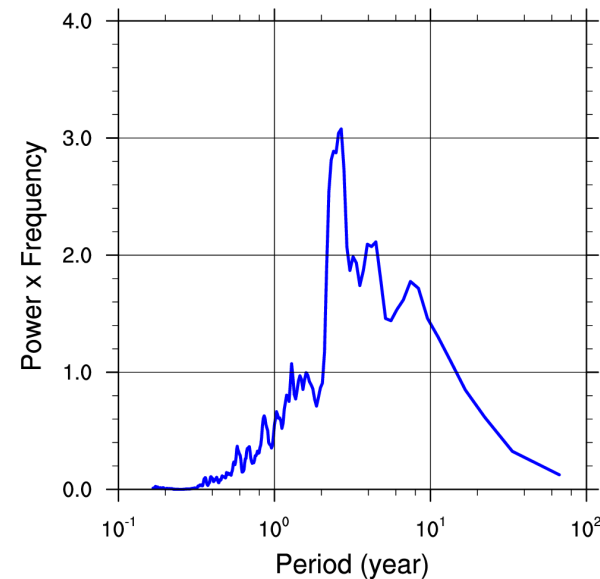
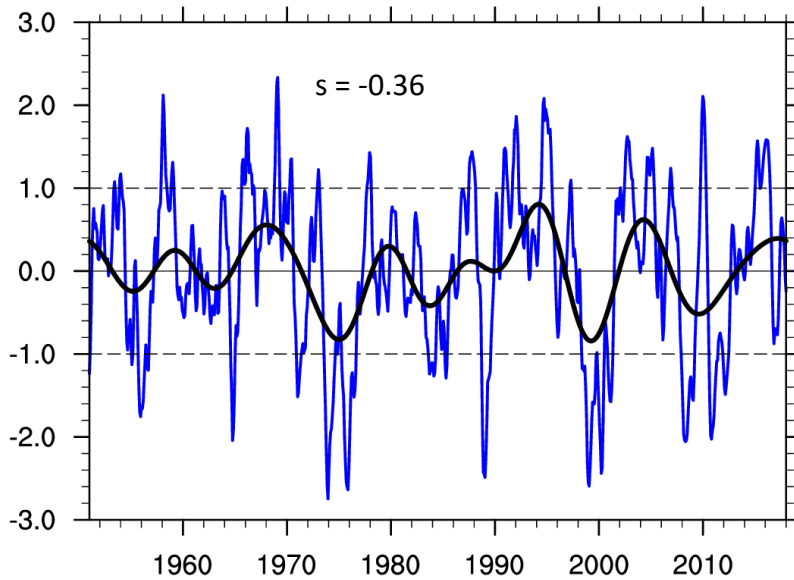




$EP_{new}$  and  $CP_{new}$  show different spectral characteristics  
 (Sullivan et al., 2016) ERSSTv5



$EP_{new} = \text{Niño3} - 0.5 \text{ Niño4}$

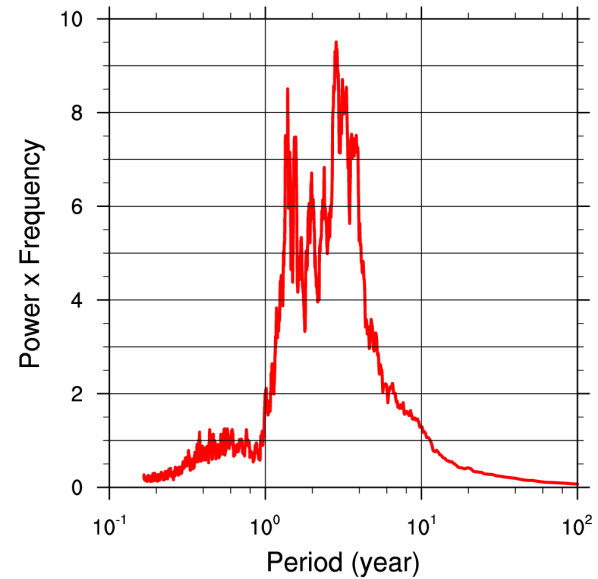
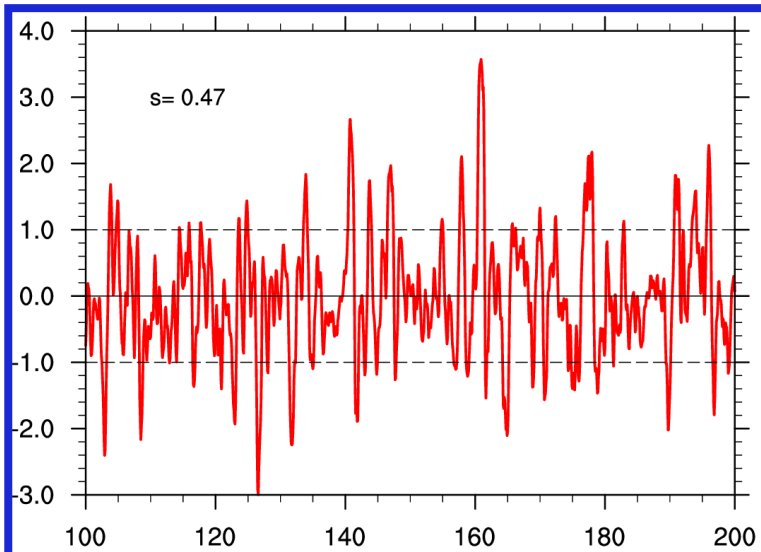


$CP_{new} = \text{Niño4} - 0.5 \text{ Niño3}$

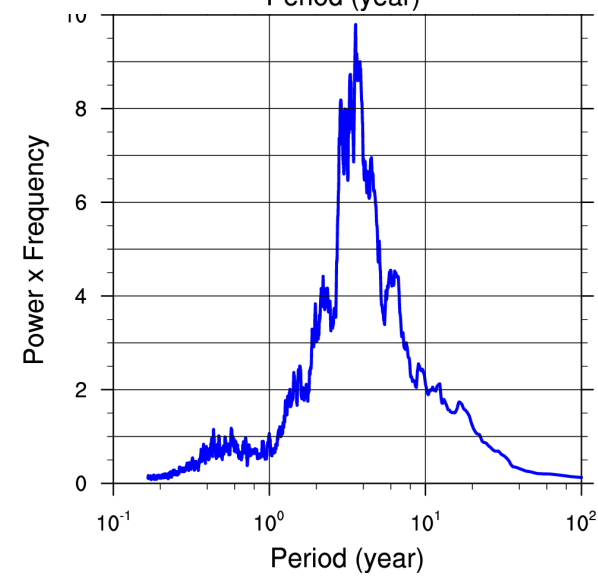
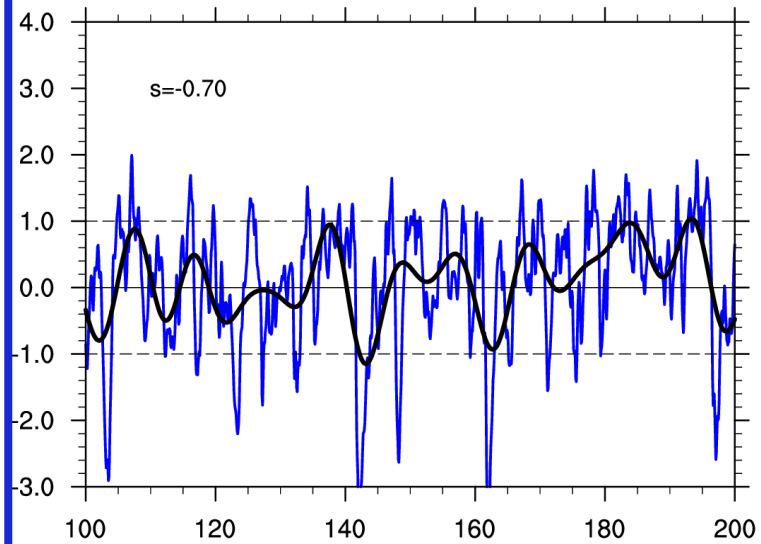
Niño3 and Niño4 normalized  
 ERSSTv5, 1951-2017

# CESM2

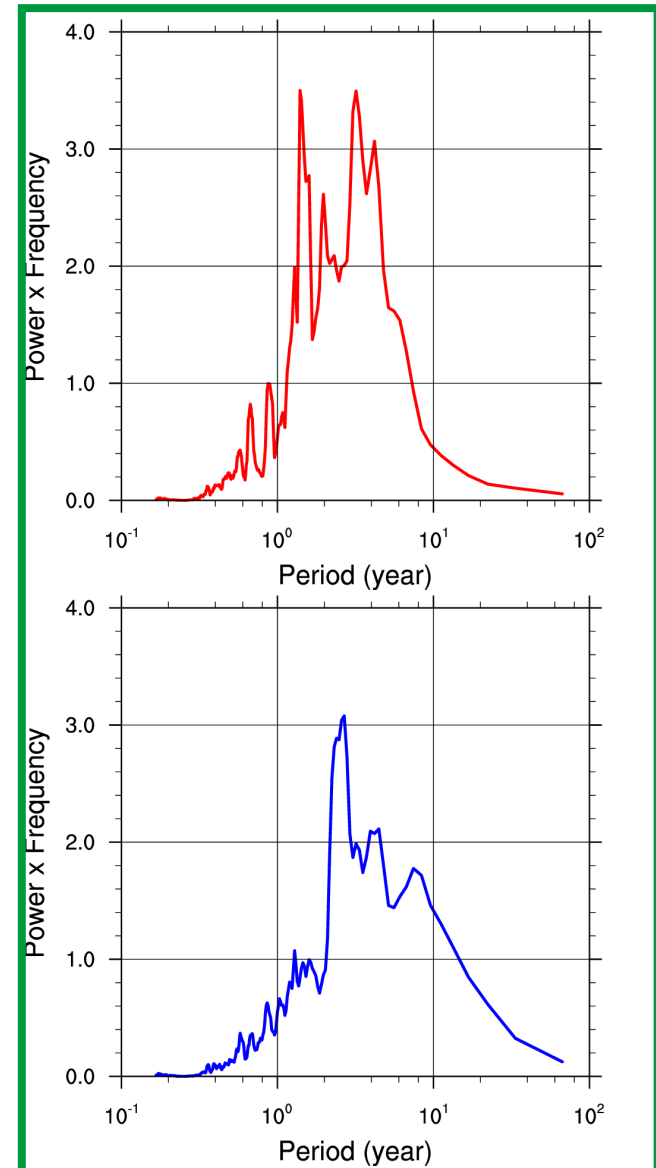
EP<sub>new</sub>



CP<sub>new</sub>



# ERSSTv5



## Conclusions

- CESM2 exhibits more overlap between EP and CP events – less diversity? This is associated with the excessive westward extension of El Nino SST anomalies
- Composite SSH, SLP, and Precipitation have qualitative similar patterns to the observed
- SST anomalies zonal propagation is consistent with observations
- El Nino/La Nina asymmetry, as indicated by the skewness, is different than observed