

Decoupling of Late Paleozoic epicontinental sea and ocean $\delta^{18}\text{O}$ in iCESM

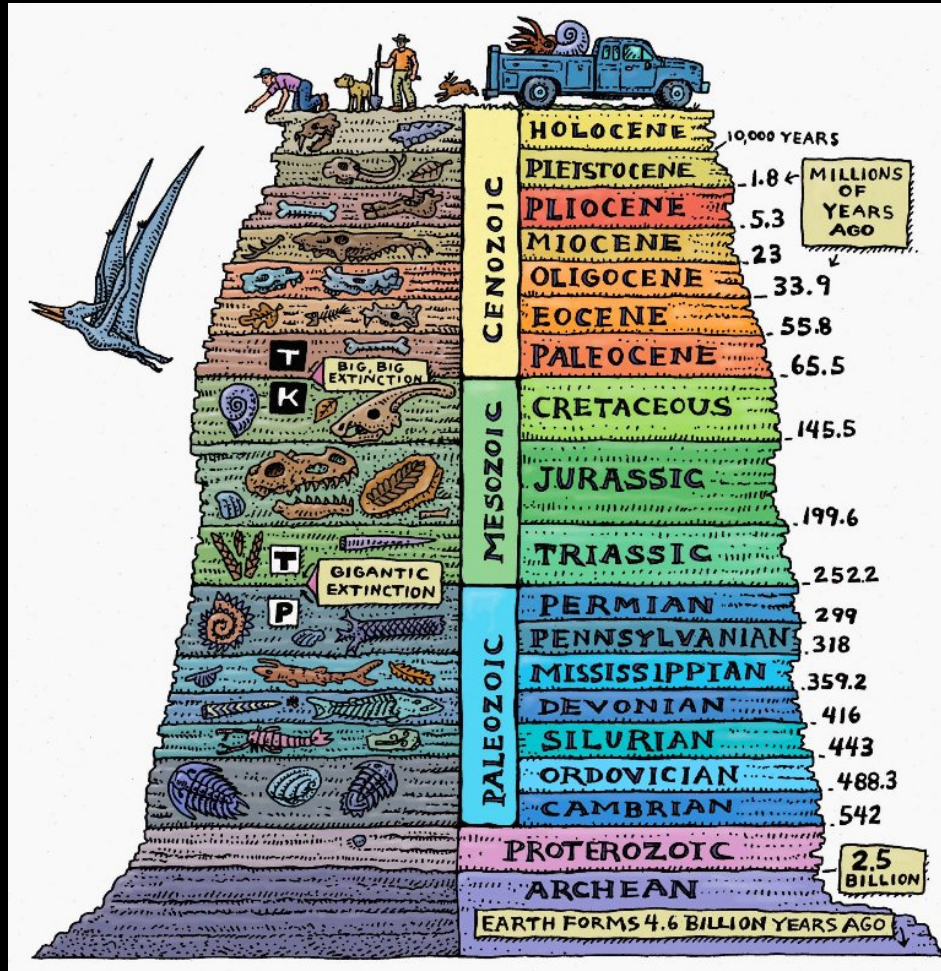
NCAR Paleoclimate Working Group Meeting 2019

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²University of California, Davis

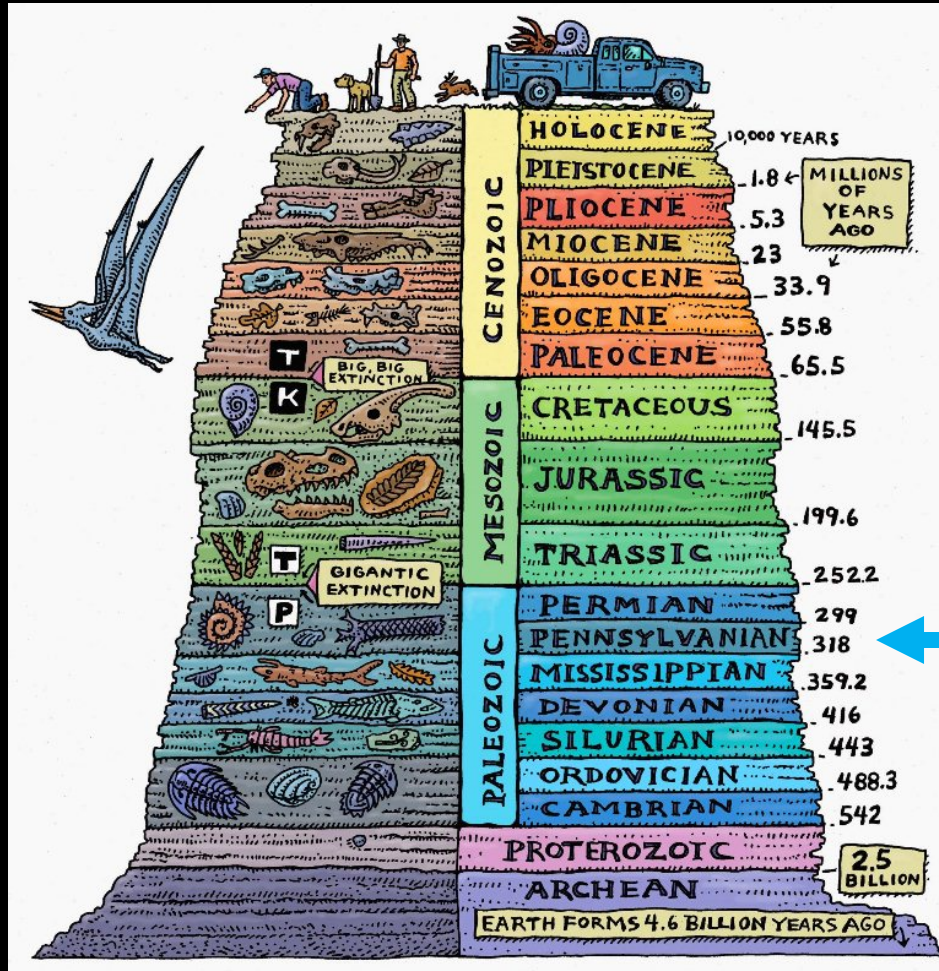
Introduction: Late Paleozoic Ice Age



- Icehouse periods make up <25% of the past billion years, yet glacial-interglacial cycles persisted for ~70 Ma of the Paleozoic (Montañez & Poulsen, 2013)

} LPIA

Introduction: Late Paleozoic Ice Age



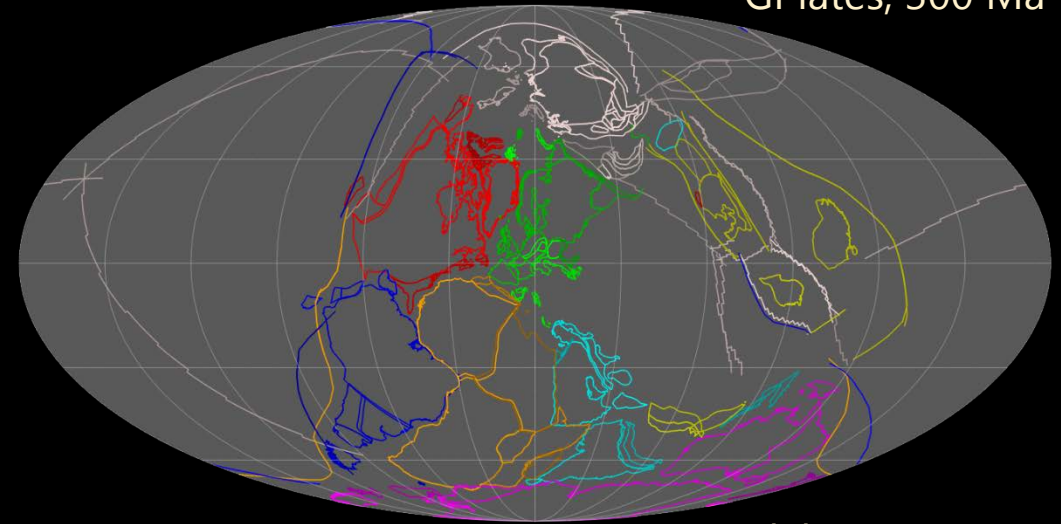
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Pennsylvanian

Introduction: Seawater reconstructions

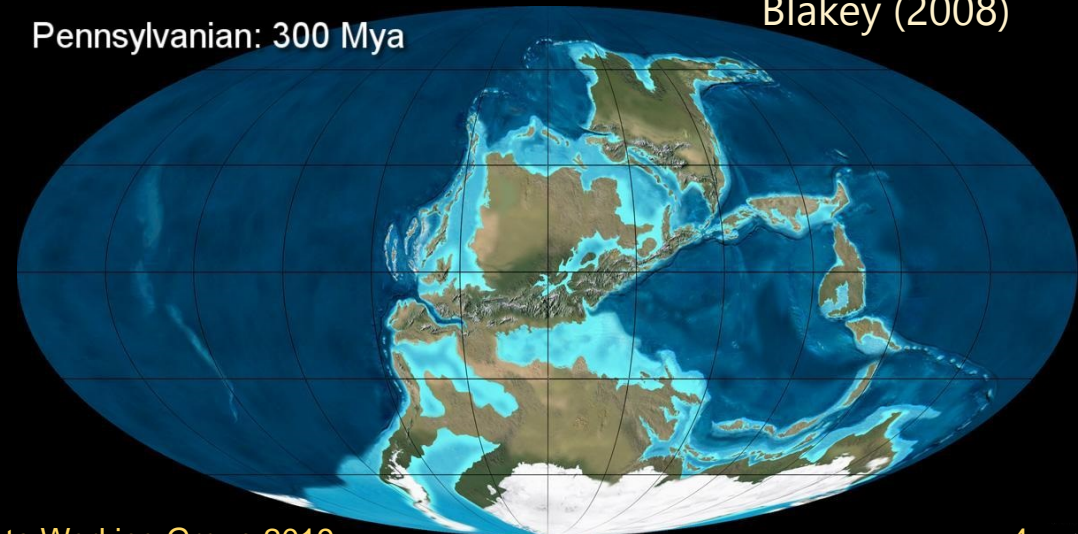
- Mean paleo-seawater conditions, especially pre-Mesozoic, largely inferred from epicontinental seas
 - Shallow carbonate platforms and interior sea successions
- Assume that ancient ocean chemistry is coupled to epicontinental seawater isotope records ($\delta^{18}\text{O}_{\text{sw}} = \sim 0\text{‰}$)

GPlates, 300 Ma



Pennsylvanian: 300 Mya

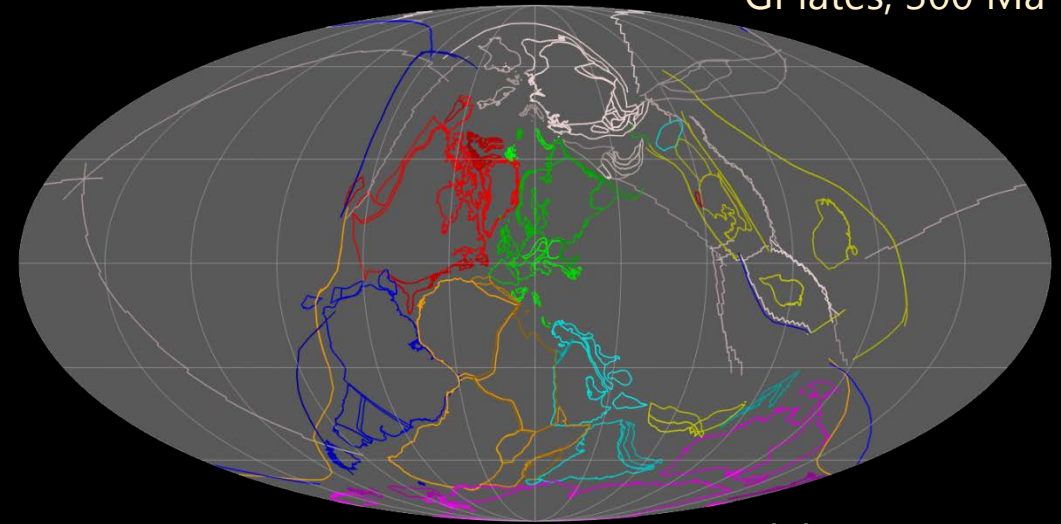
Blakey (2008)



Introduction: Seawater reconstructions

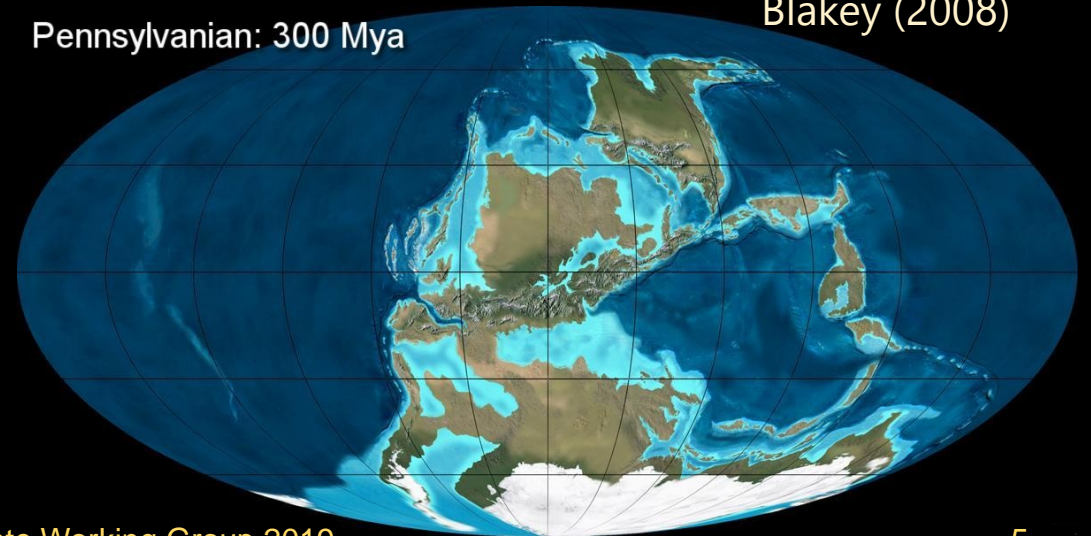
- Mounting geologic evidence for the overprinting of local processes on global trends in epicontinental records
 - Shallow depths and limited communication with the open ocean cause high variability in seawater conditions (Montañez et al., 2018; Joachimski & Lambert, 2015; Roark et al., 2017; Brand et al., 2009)

GPlates, 300 Ma

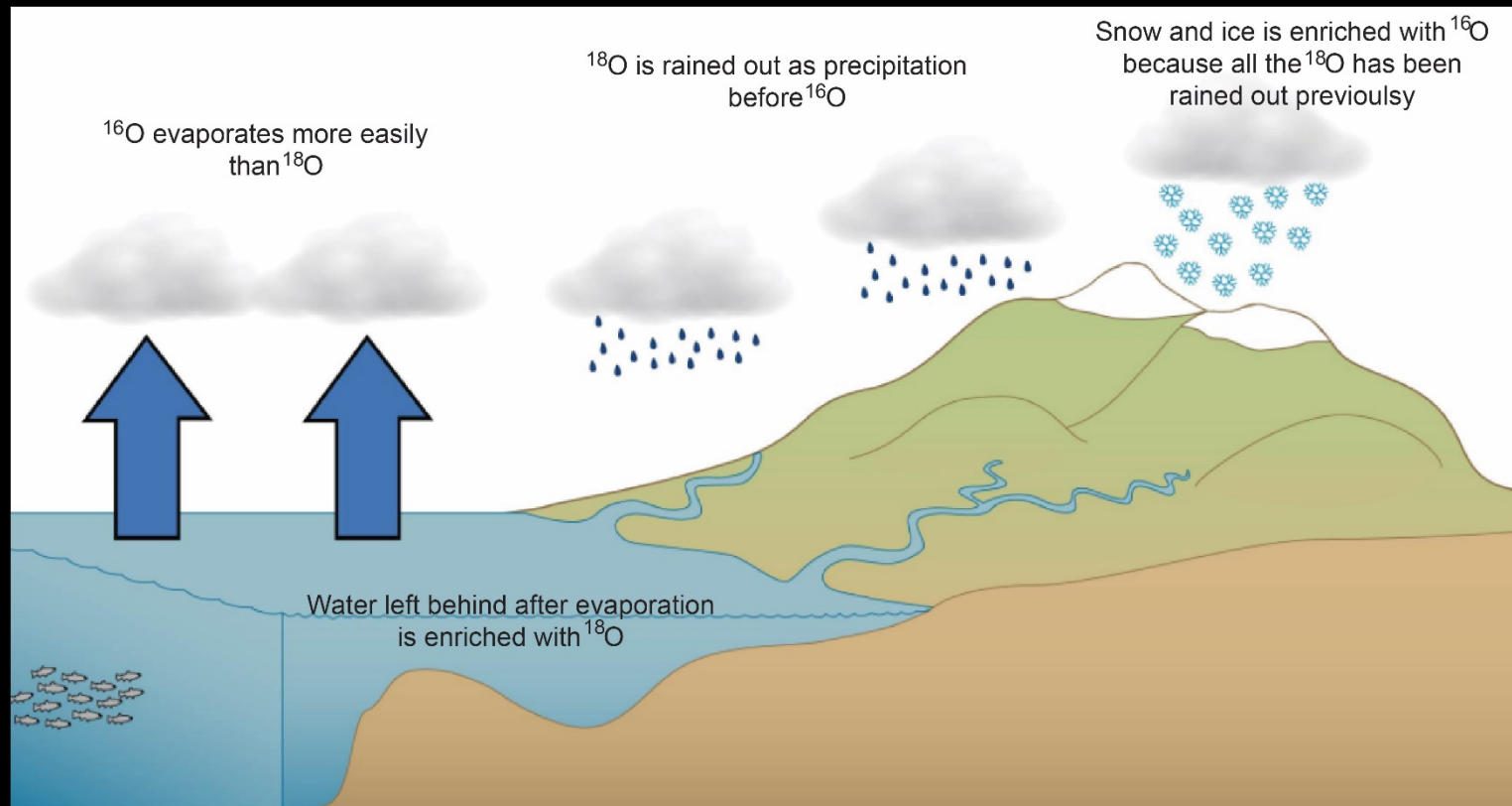


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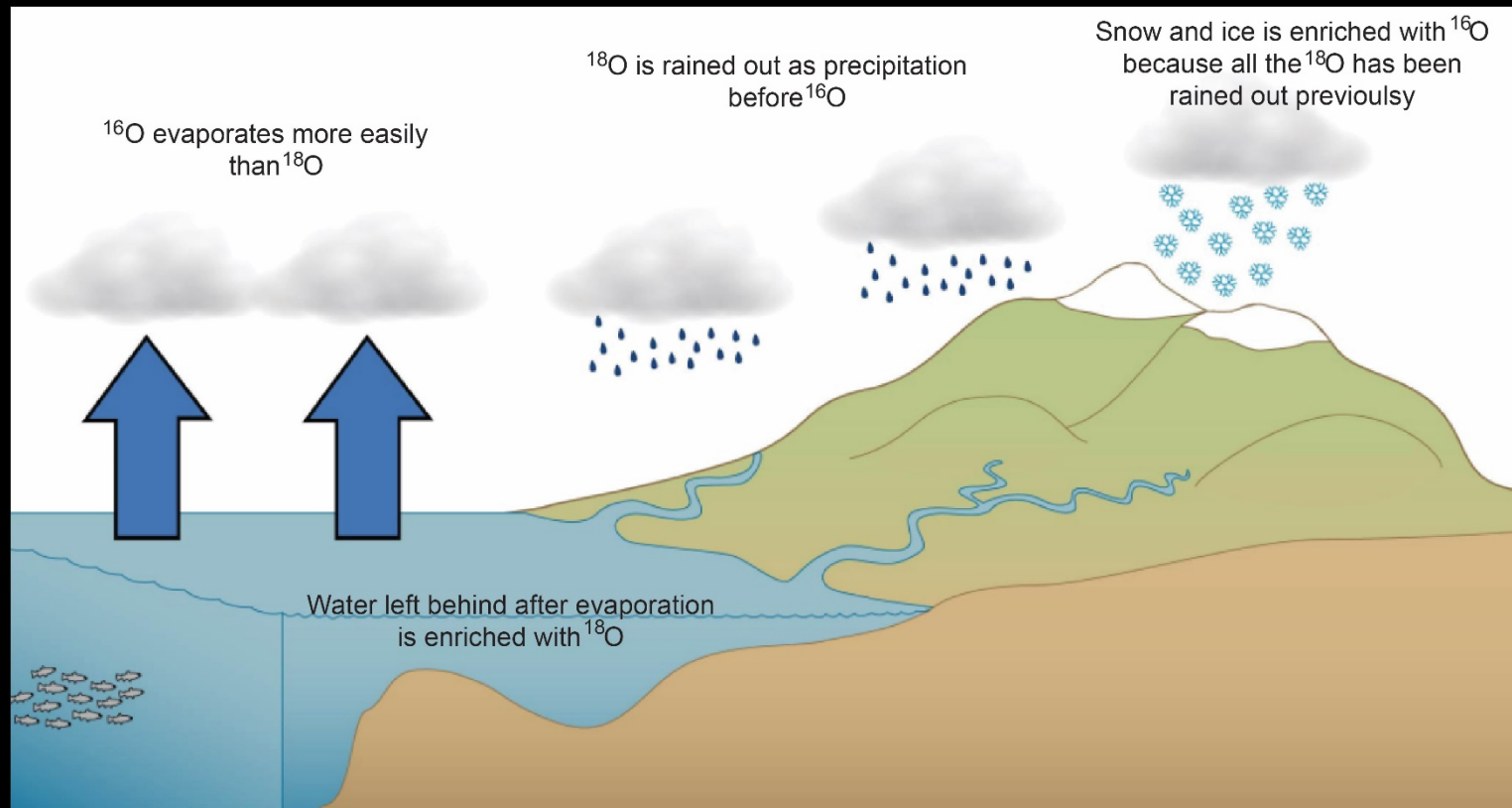


Introduction: $\delta^{18}\text{O}$ of seawater



- Evaporation:
 $\uparrow\delta^{18}\text{O}$
- Precipitation
and runoff:
 $\downarrow\delta^{18}\text{O}$

Introduction: $\delta^{18}\text{O}$ of seawater



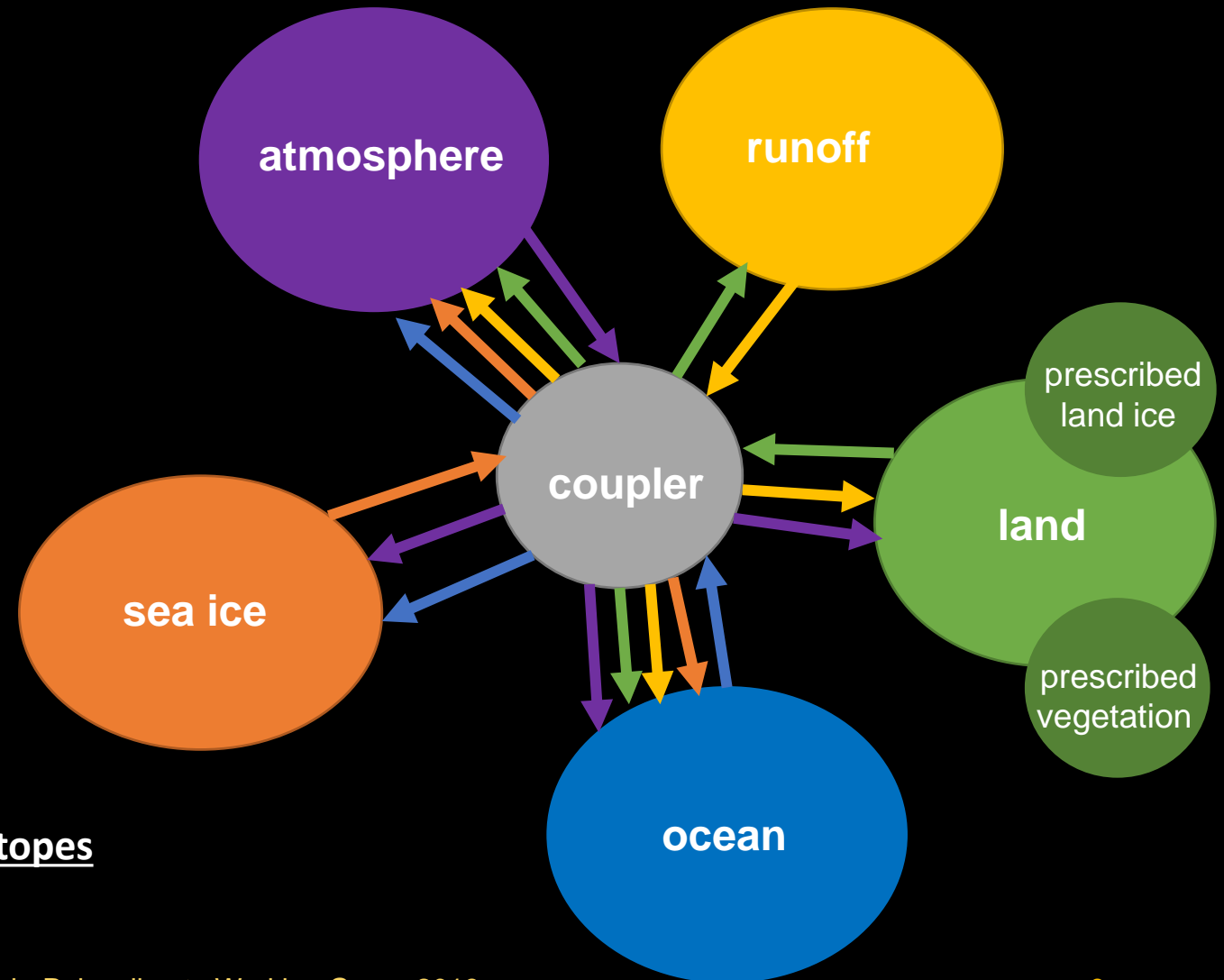
- Evaporation: $\uparrow\delta^{18}\text{O}$
- Precipitation and runoff: $\downarrow\delta^{18}\text{O}$
- Ice sheet growth: $\uparrow\delta^{18}\text{O}$
- Ice sheet melt: $\downarrow\delta^{18}\text{O}$

Goals of the present study

1. Present the first simulations of Late Paleozoic glacial and interglacial states that resolve epicontinental sea dynamics
2. Investigate what processes influence seawater $\delta^{18}\text{O}$ in epicontinental seas, and whether these environments are coupled to mean ocean conditions
3. Compare simulations with published geochemical proxy data to constrain seawater $\delta^{18}\text{O}$ in the Midcontinent sea

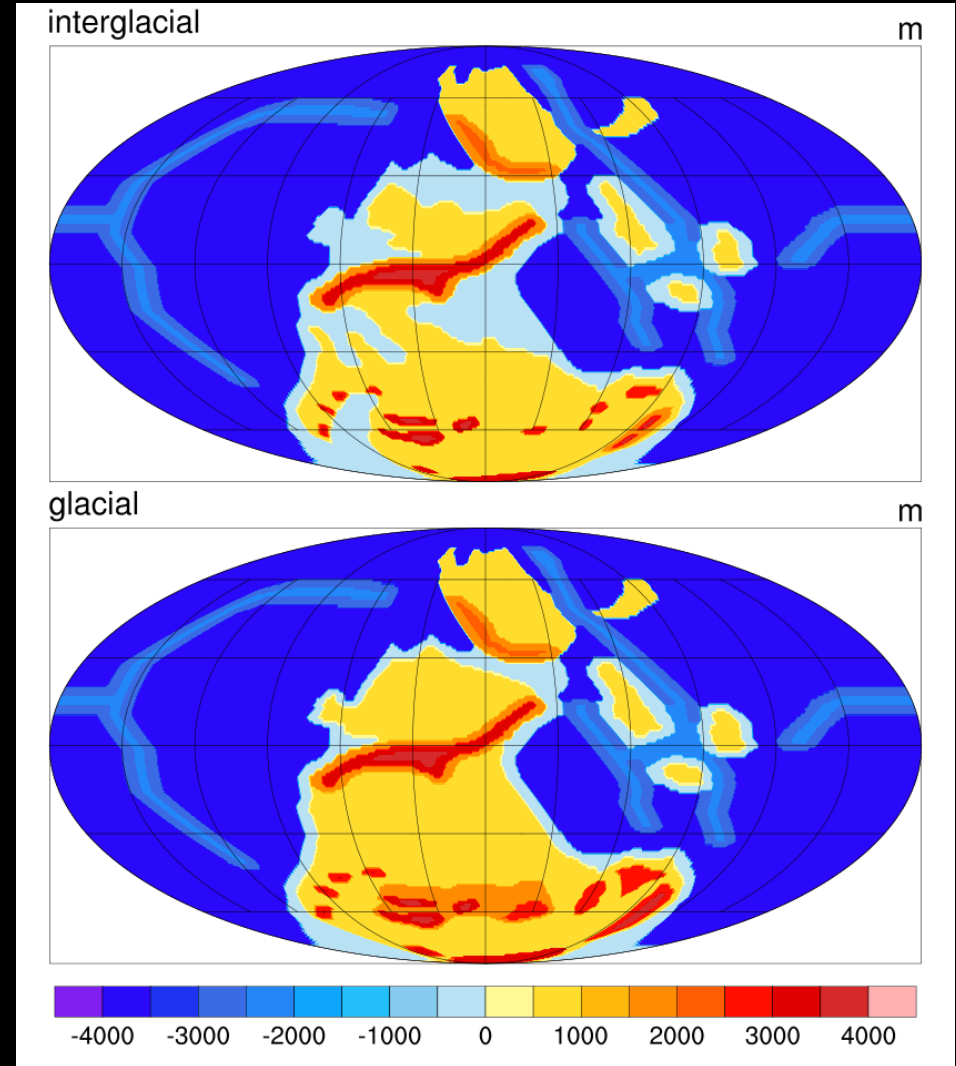
Methods: iCESM1 setup

- CAM5 and CLM4 (1.9°x2.5°)
- POP2 and CISM4 (~1°)



Methods: Base simulations

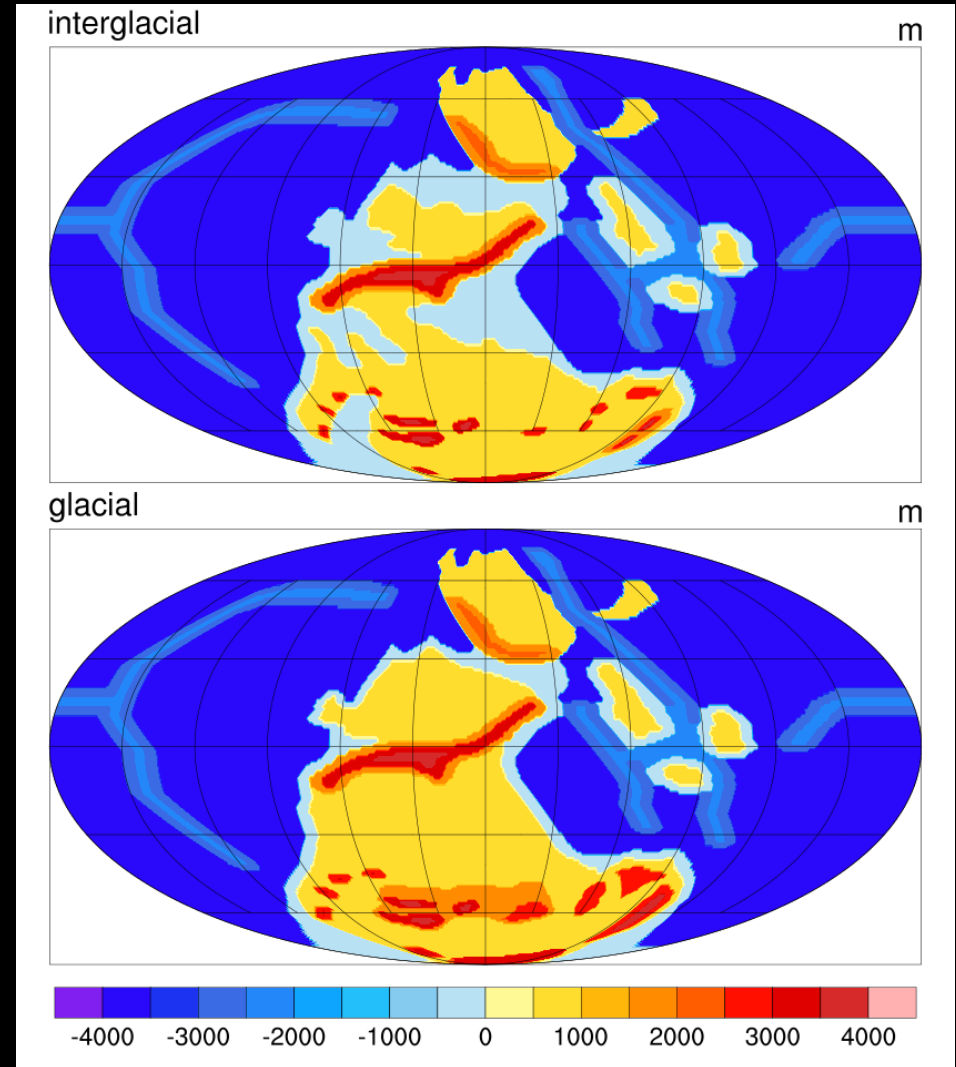
Case	Glacial	Interglacial
Runtime	2000 years	2000 years
Paleogeography	300 Ma Low sea level/ High land ice	300 Ma High sea level/ Low land ice
Atmospheric [CO ₂]	280 ppmv	560 ppmv
Solar luminosity	97.5% modern	97.5% modern
Initial $\delta^{18}\text{O}$ ocean	0.5‰	0‰



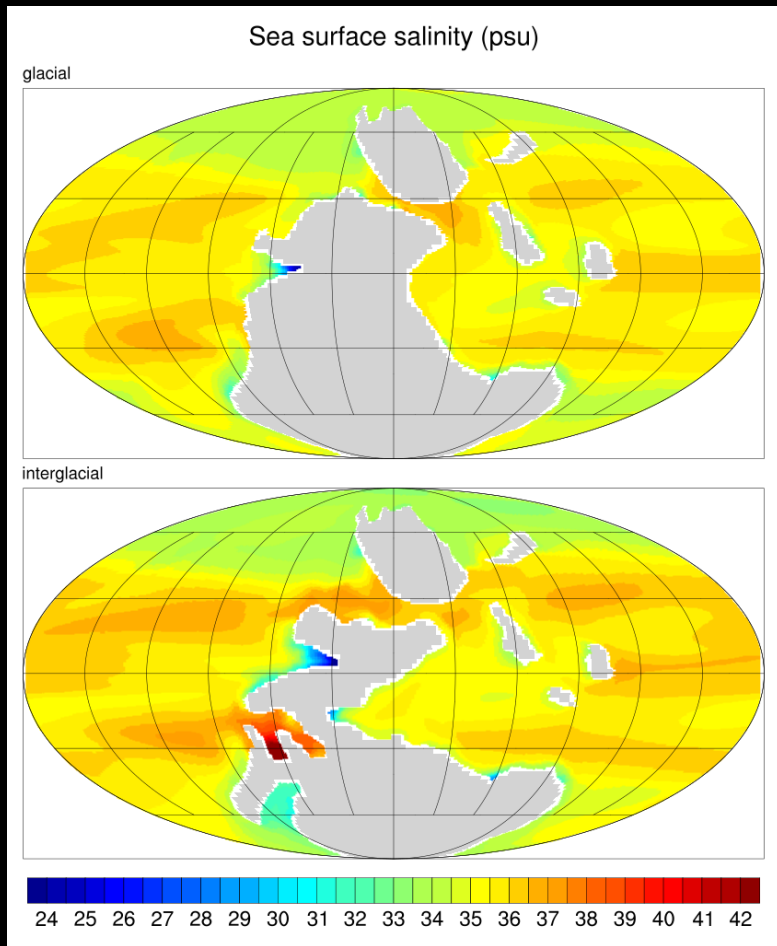
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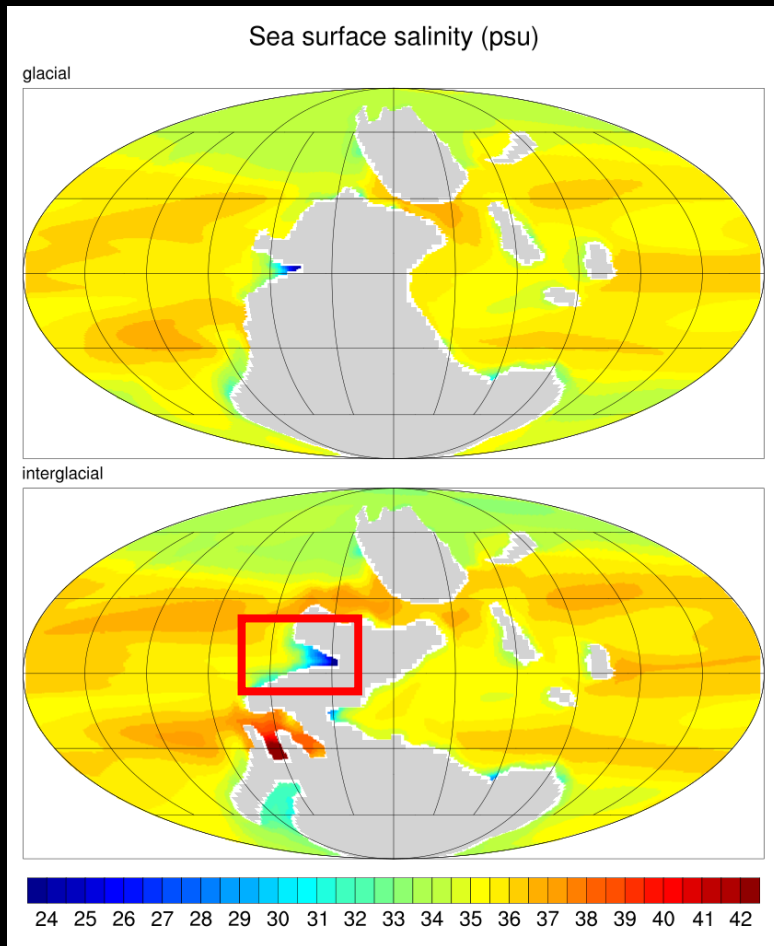
*Differences between two simulations



Results: Mean annual SSS

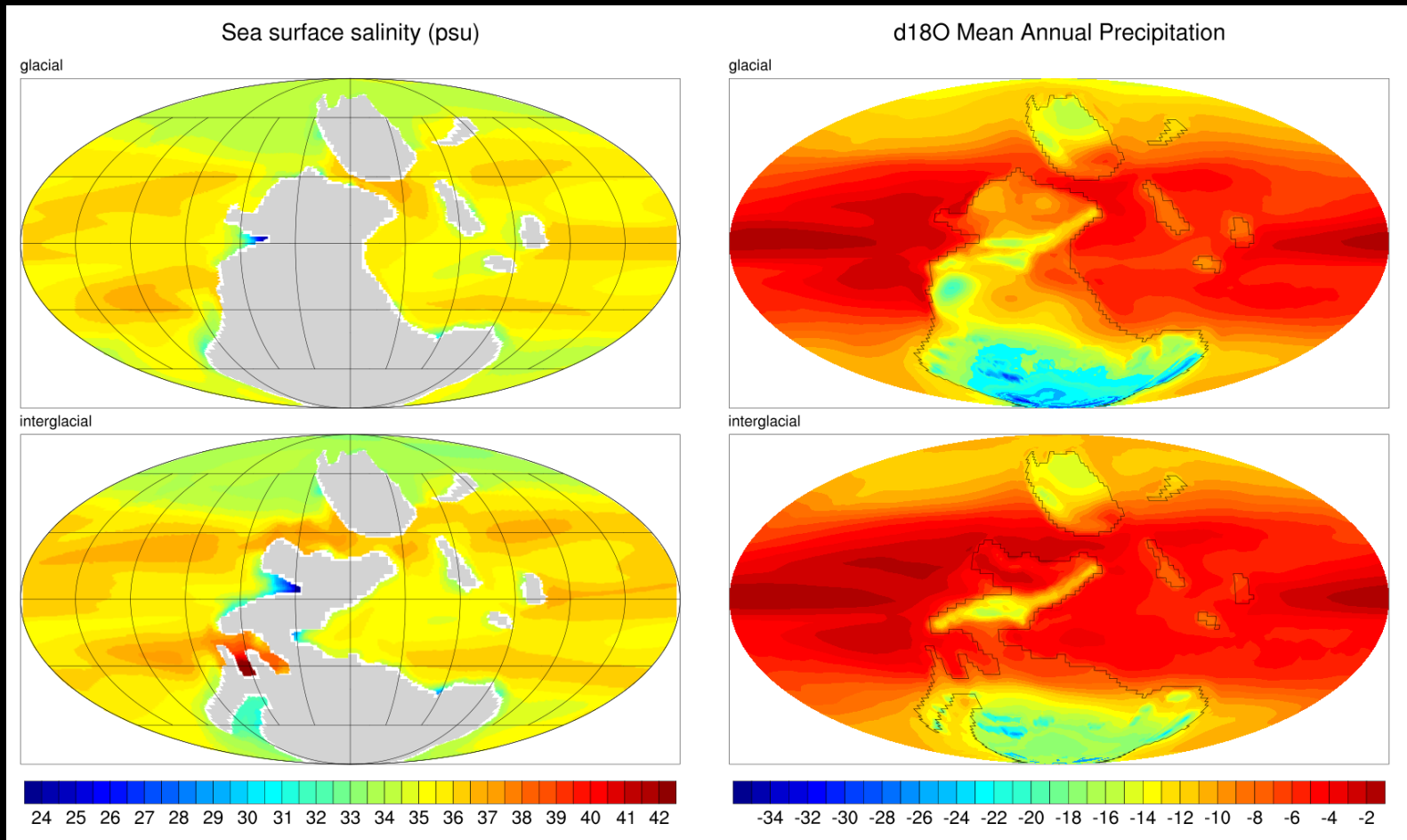


Results: Mean annual SSS



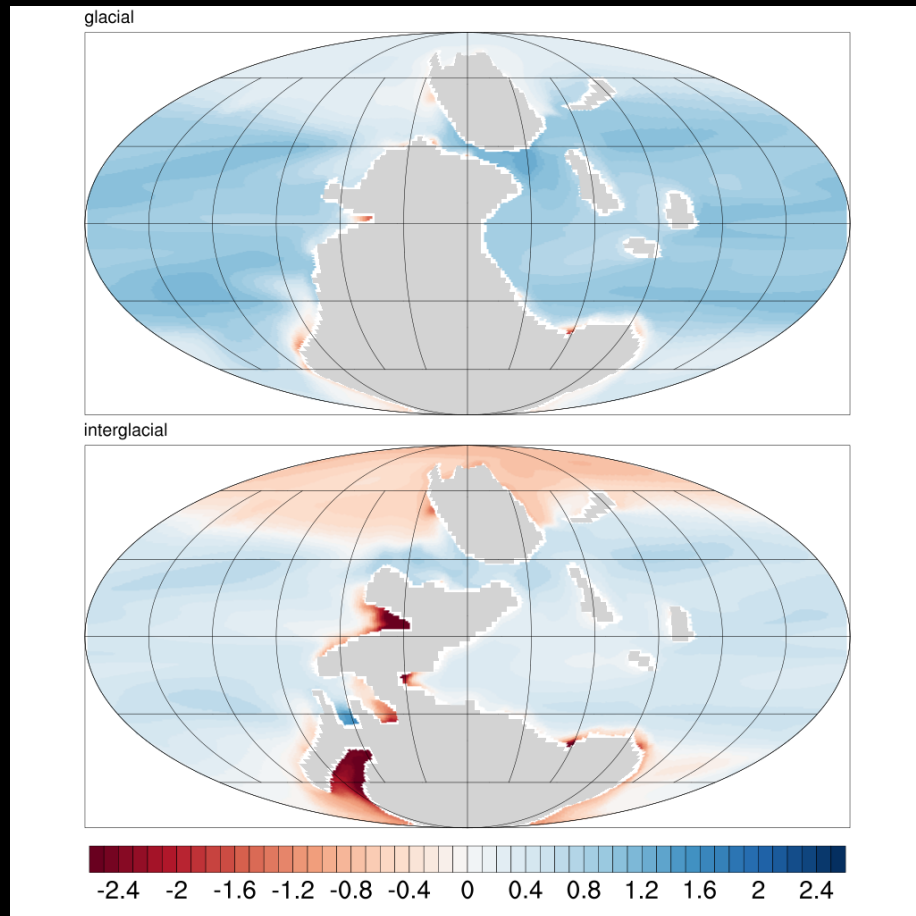
- Interglacial salinity in Midcontinent sea is low, particularly during interglacial high stand

Results: Mean annual SSS & precipitation



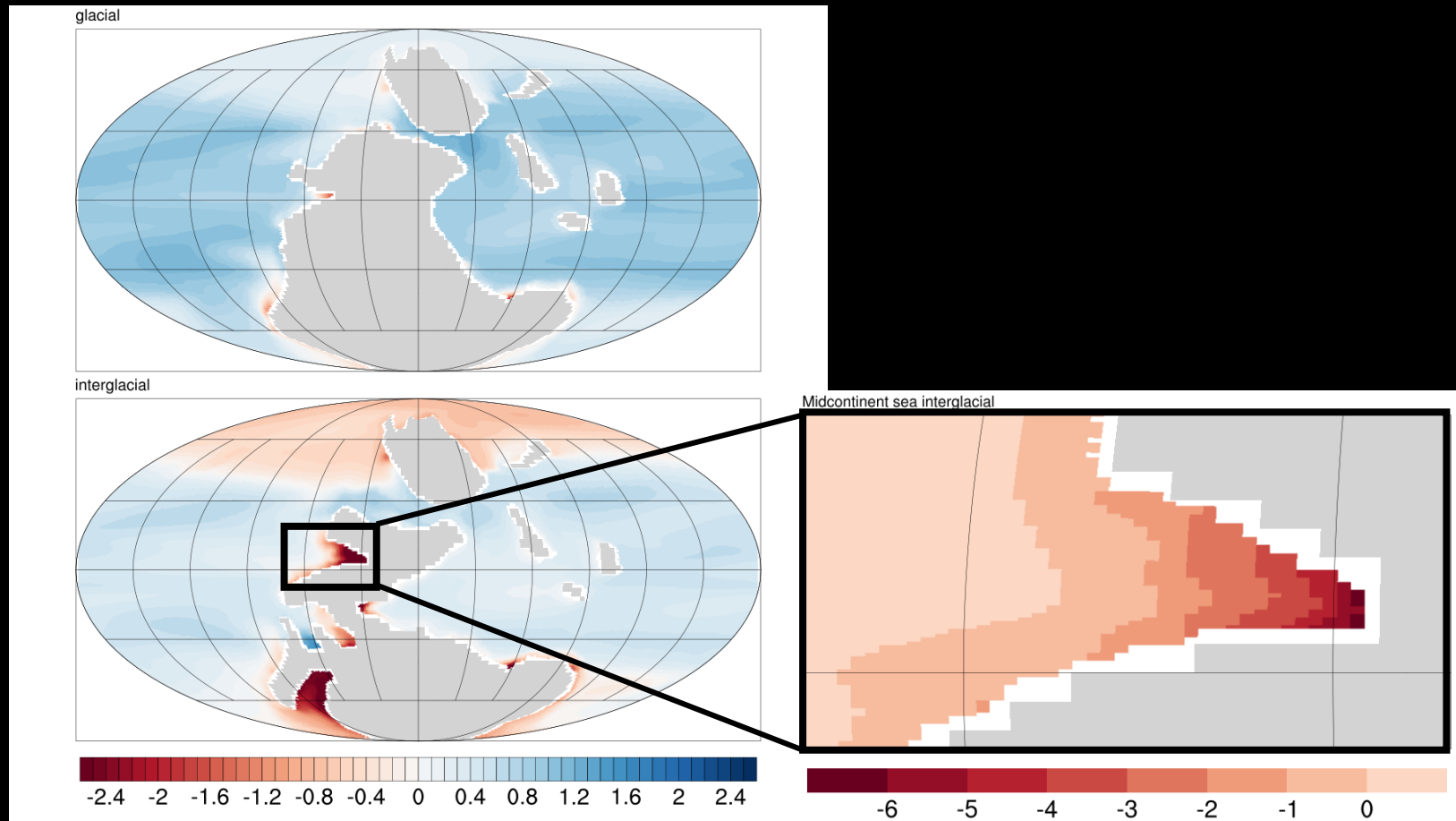
- Interglacial salinity in Midcontinent sea is low, particularly during interglacial high stand
- Depleted $\delta^{18}\text{O}$ of precipitation and runoff from Central Pangean Mountains into Midcontinent sea

Results: Mean annual seawater $\delta^{18}\text{O}$ (‰)



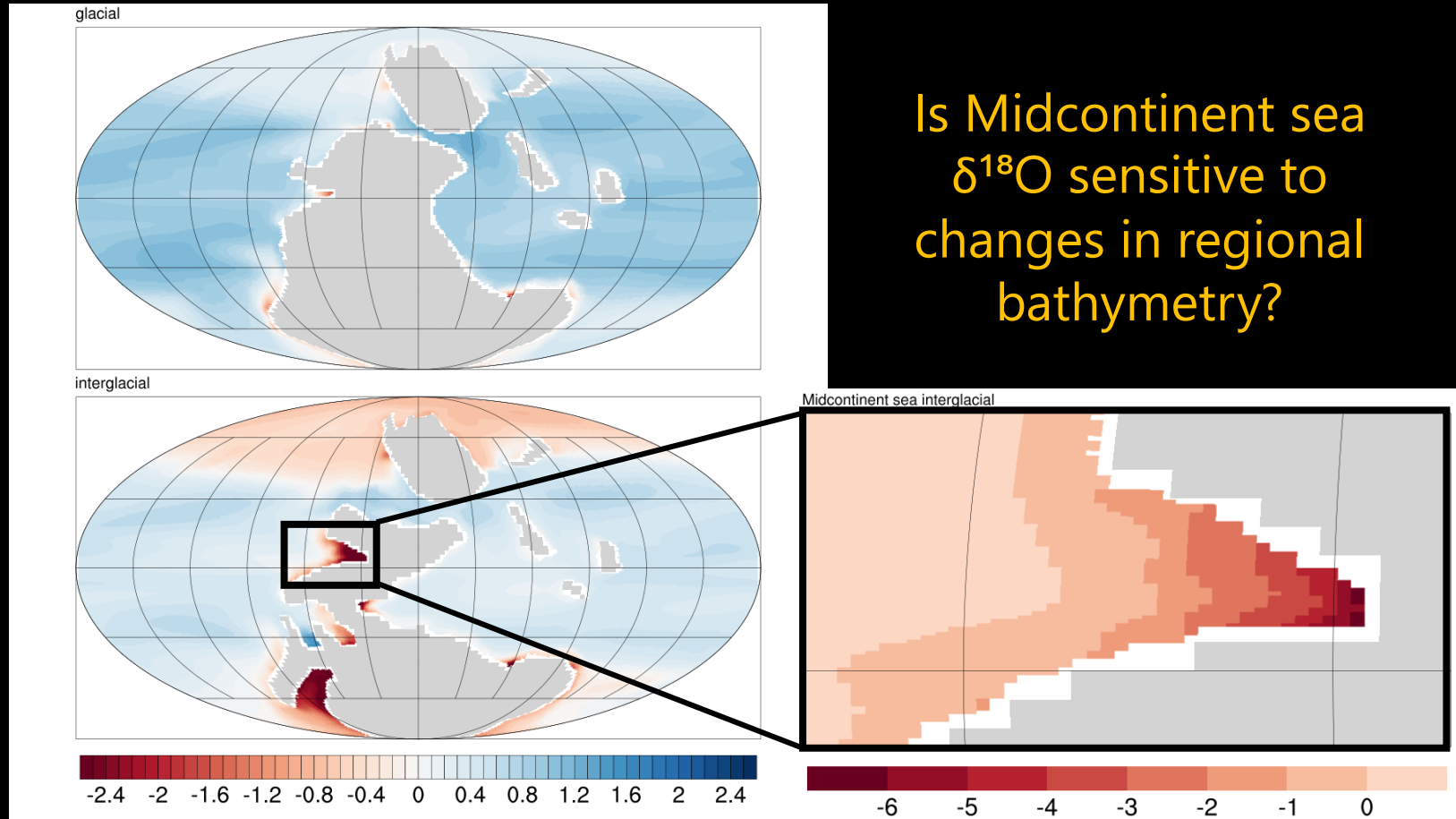
- Relatively more enriched $\delta^{18}\text{O}_{\text{sw}}$ in glacial
- Gradient of decreasing $\delta^{18}\text{O}_{\text{sw}}$ from equator to the poles

Results: Mean annual seawater $\delta^{18}\text{O}$ (‰)



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- Gradient of decreasing $\delta^{18}\text{O}_{\text{sw}}$ from equator to the poles
- Interglacial $\delta^{18}\text{O}_{\text{sw}}$ gradient of $>5\text{‰}$ over Midcontinent sea from high riverine influx

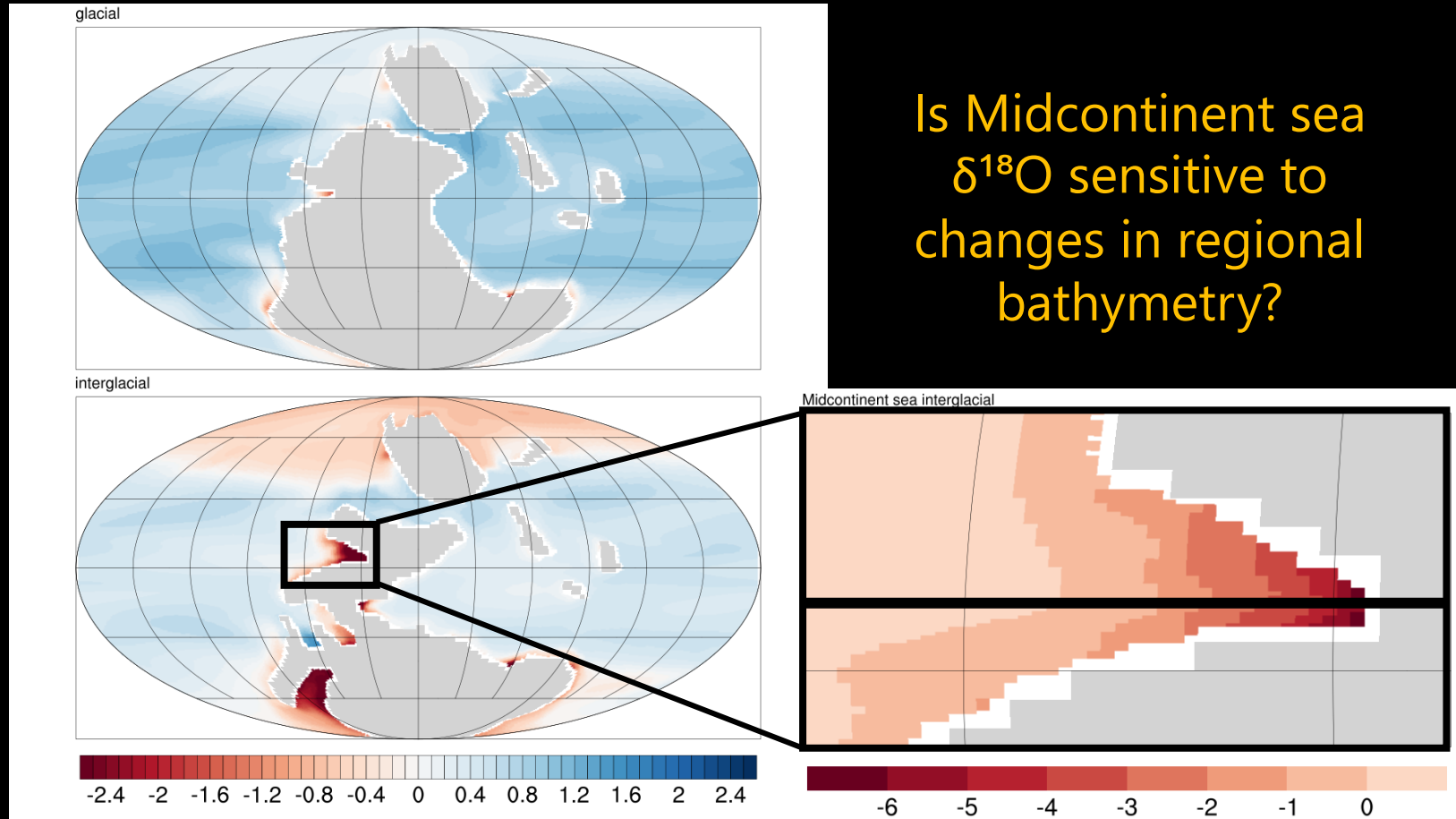
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Is Midcontinent sea $\delta^{18}\text{O}$ sensitive to changes in regional bathymetry?

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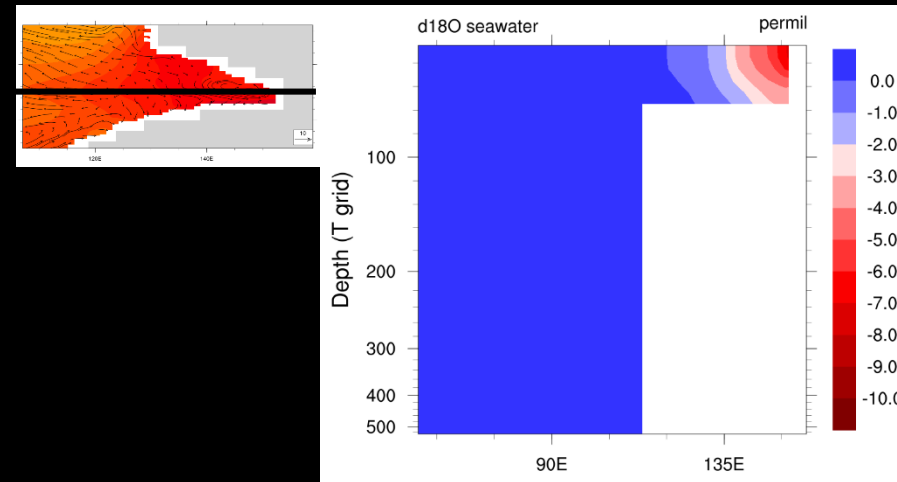


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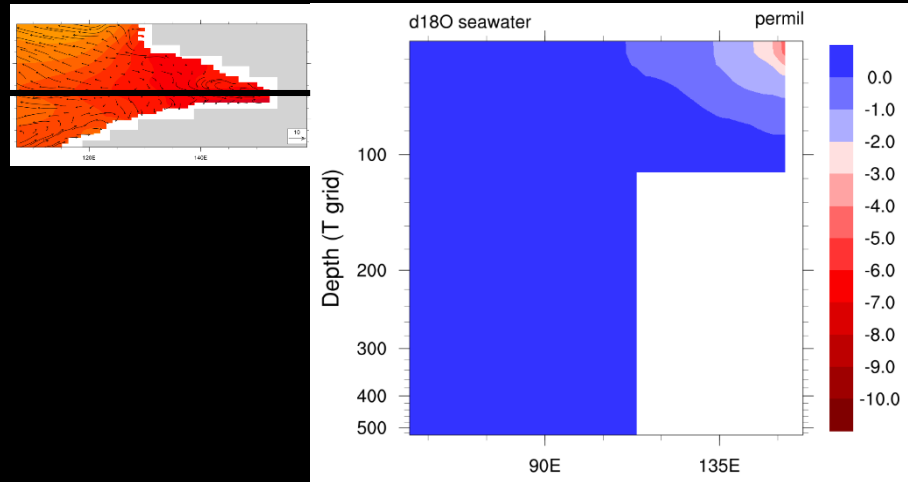
Results: Sensitivity of Midcontinent $\delta^{18}\text{O}$

60 m depth

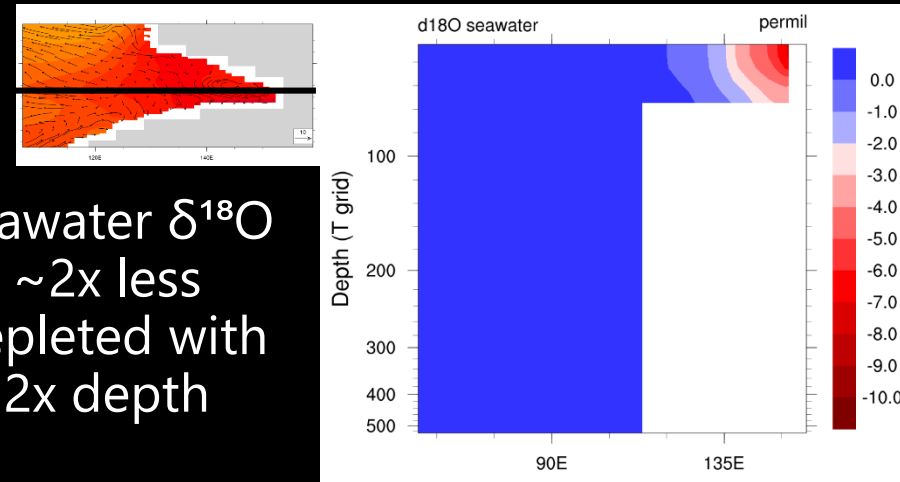


Results: Sensitivity of Midcontinent $\delta^{18}\text{O}$

120 m depth



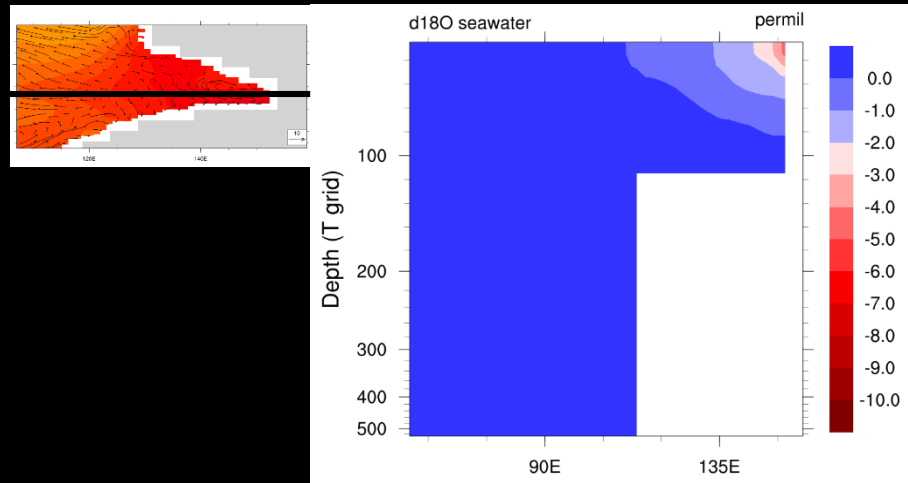
60 m depth



Seawater $\delta^{18}\text{O}$
~2x less
depleted with
2x depth

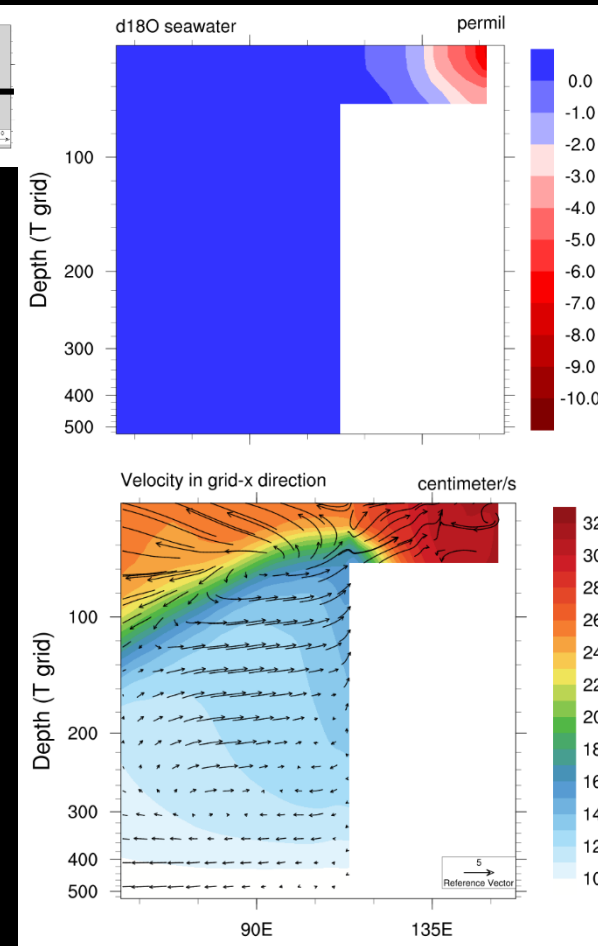
Results: Sensitivity of Midcontinent $\delta^{18}\text{O}$

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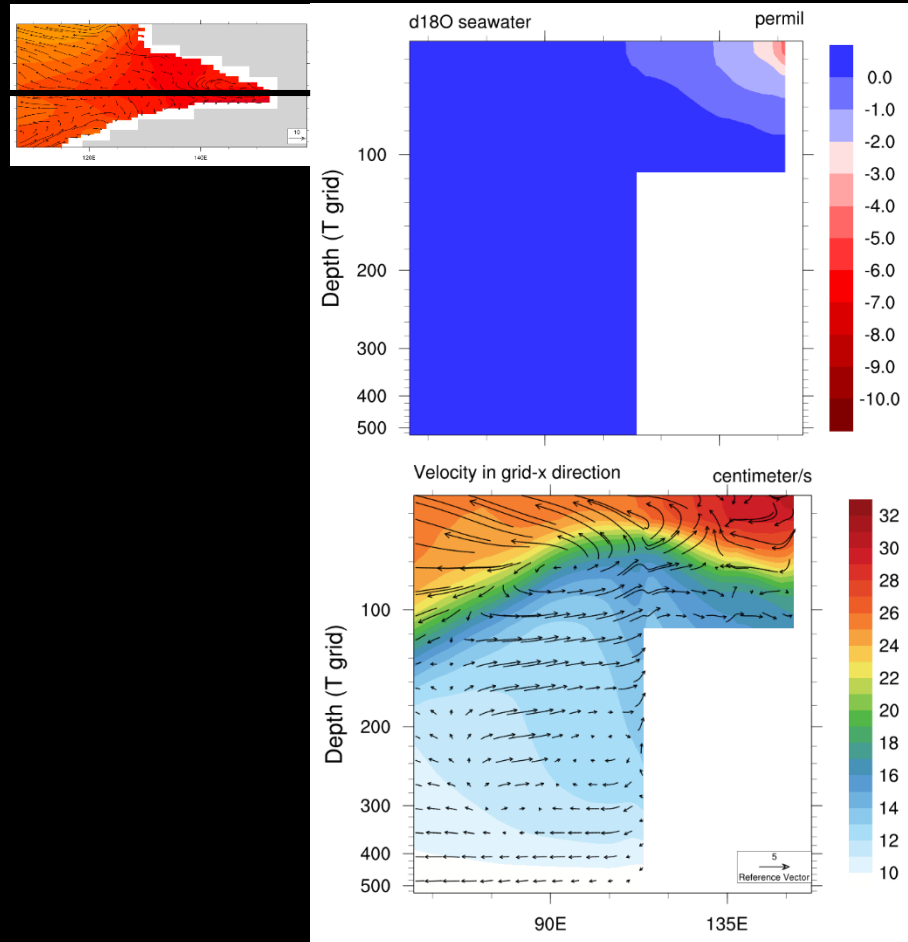
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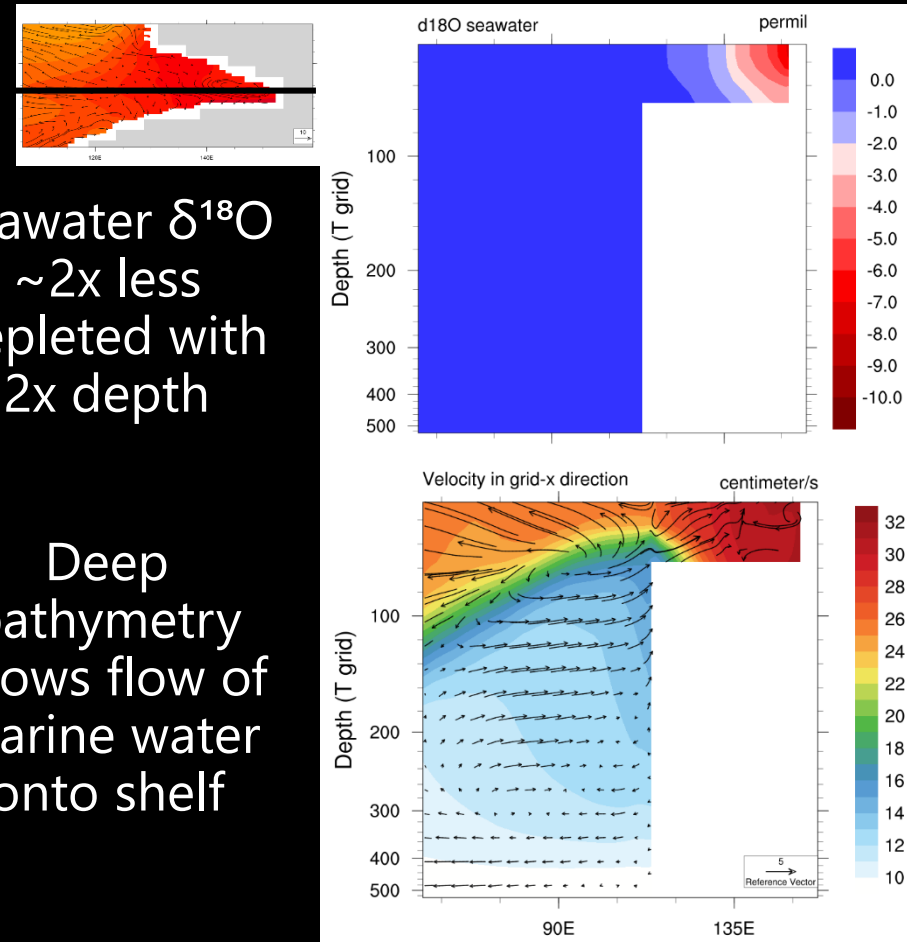
120 m depth

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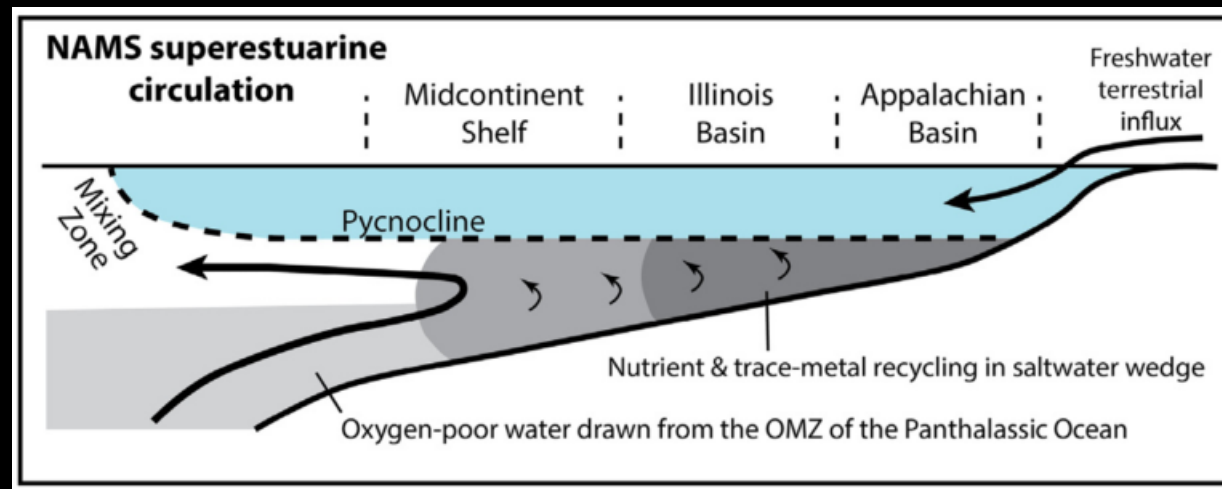
Seawater $\delta^{18}\text{O}$
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Deep
bathymetry
allows flow of
marine water
onto shelf



Results: Sensitivity of Midcontinent $\delta^{18}\text{O}$

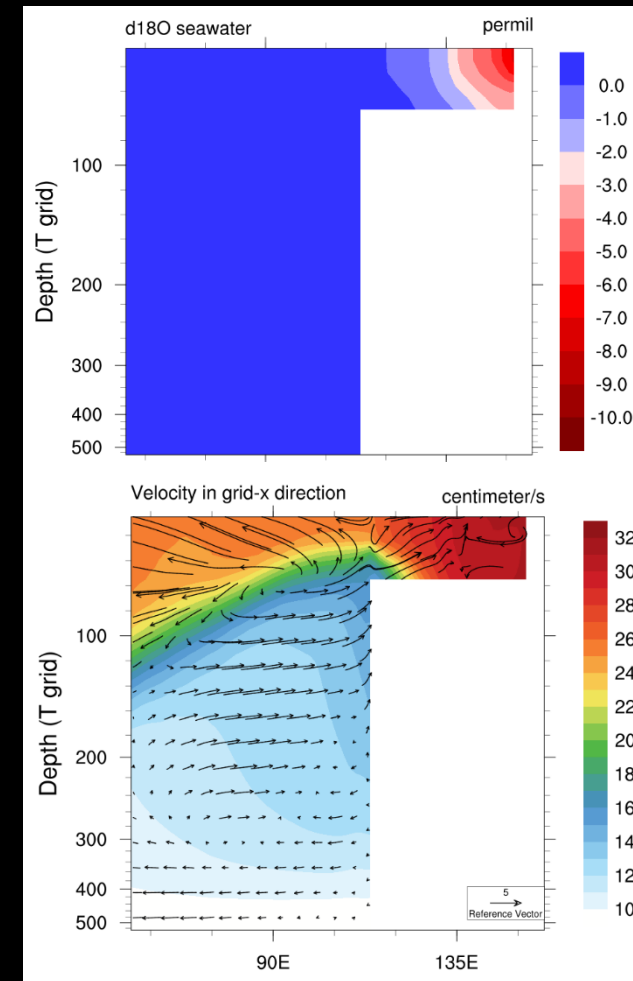
- North American Midcontinent Sea (NAMS) has a mean depth of 50 m (Algeo & Heckel, 2008)



(Turner et al., 2018)

Results: Sensitivity of Midcontinent $\delta^{18}\text{O}$

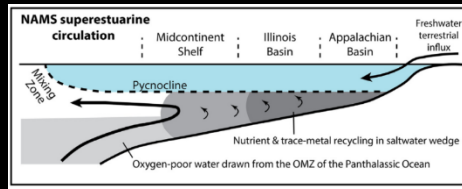
Uniform 60 m



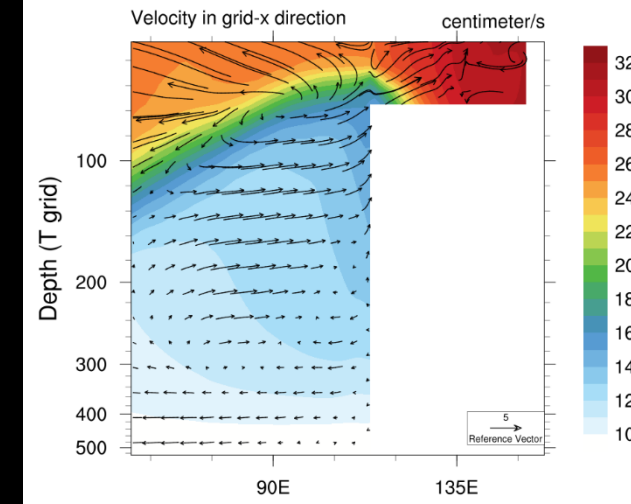
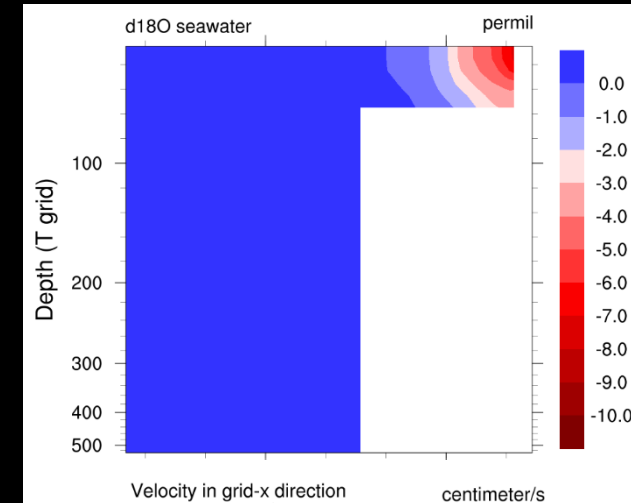
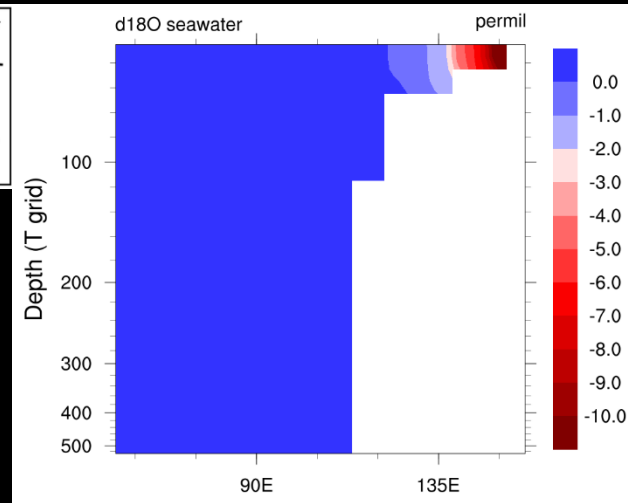
Results: Sensitivity of Midcontinent $\delta^{18}\text{O}$

Mean 50 m

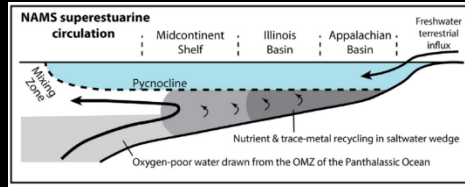
Uniform 60 m



Despite open coastline, shallow depth creates $> 10\text{‰}$ gradient in seawater $\delta^{18}\text{O}$



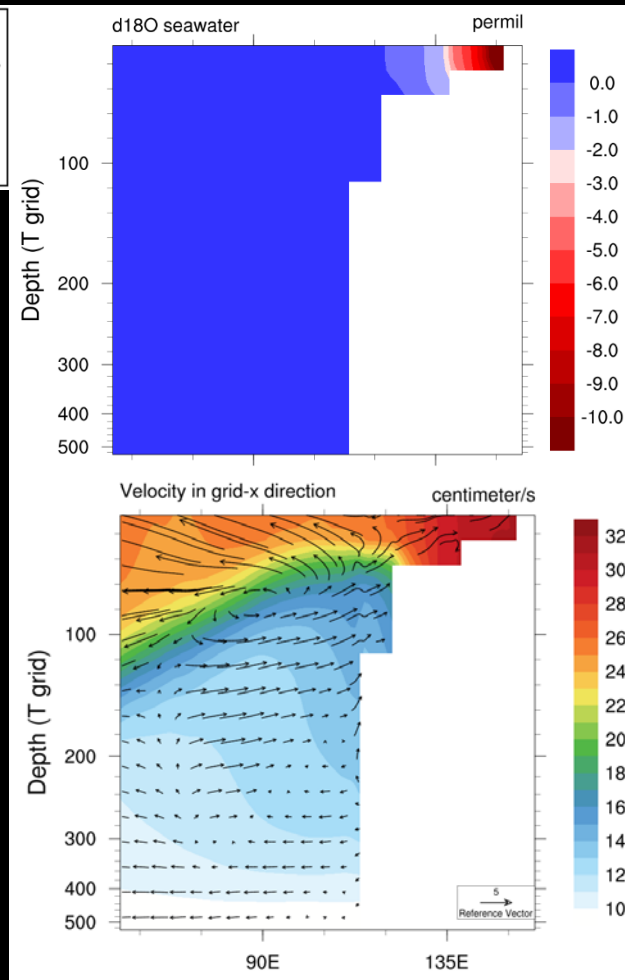
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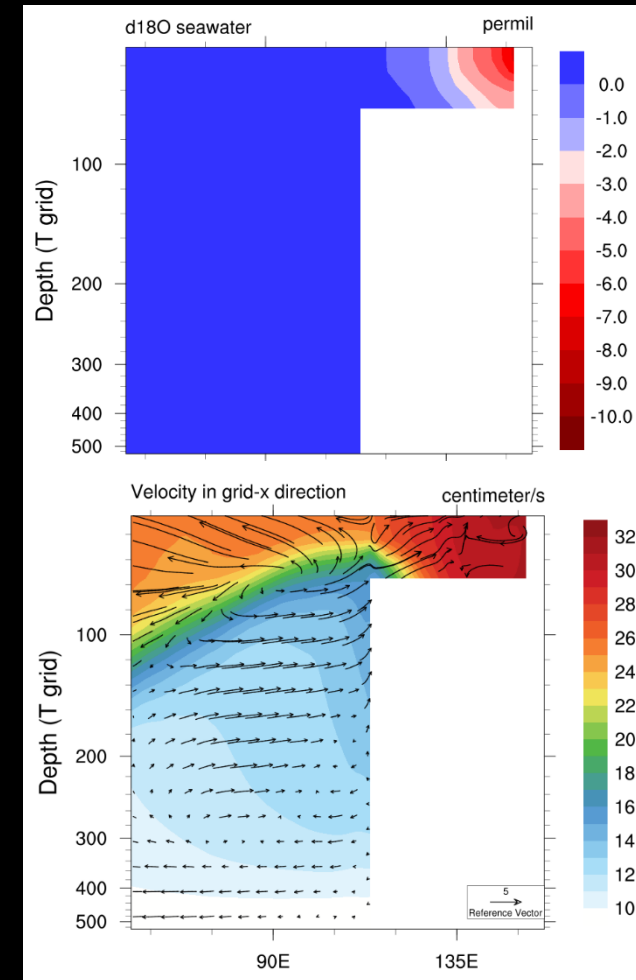
Despite open coastline, shallow depth creates $>10\text{‰}$ gradient in seawater $\delta^{18}\text{O}$

Mean depth >50 m needed for flow of marine deepwater into eastern sea

Mean 50 m



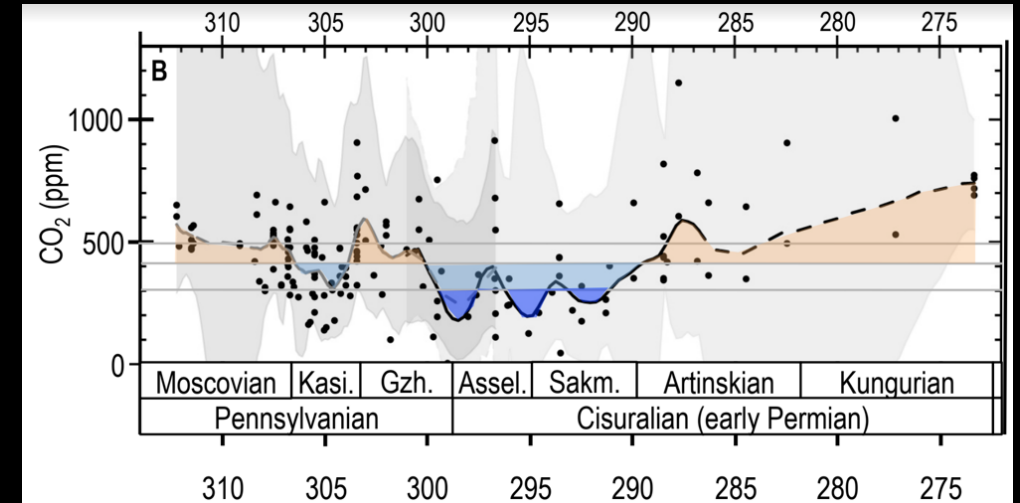
Uniform 60 m



Results: Data-model comparison

(Isabel Montañez, personal comm.)

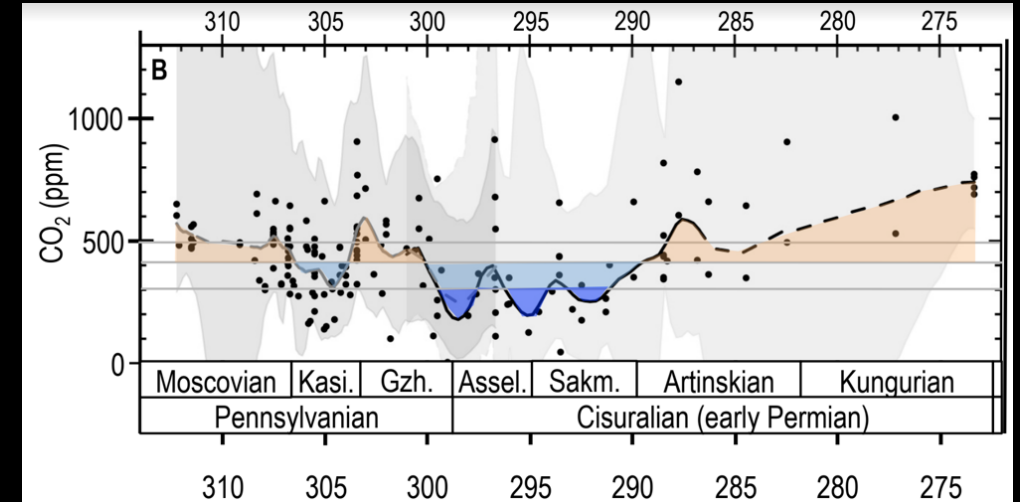
- Compiled proxy data from an interglacial (304 – 302 Ma) interval
- Calculated paleotemperatures using simulated and assumed $\delta^{18}\text{O}_{\text{sw}}$



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- Brachiopod carbonate $\delta^{18}\text{O}$

$$T(^{\circ}\text{C}) = 15.7 - 4.36(\delta^{18}\text{O}_{\text{calcite}} - \delta^{18}\text{O}_{\text{sw}}) + 0.12(\delta^{18}\text{O}_{\text{calcite}} - \delta^{18}\text{O}_{\text{sw}})^2$$

Hays & Grossman (1991)

- Conodonts phosphate $\delta^{18}\text{O}$

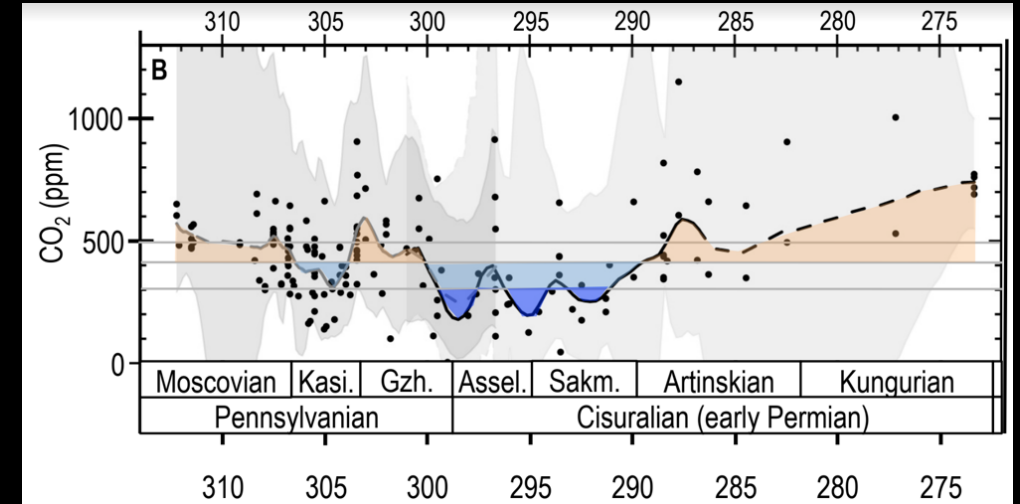
$$T(^{\circ}\text{C}) = 118.7 - 4.22(\delta^{18}\text{O}_{\text{phosphate}} + (22.6 - \delta^{18}\text{O}_{\text{NBS120c}}) - \delta^{18}\text{O}_{\text{sw}})$$

Puceat et al. (2010)

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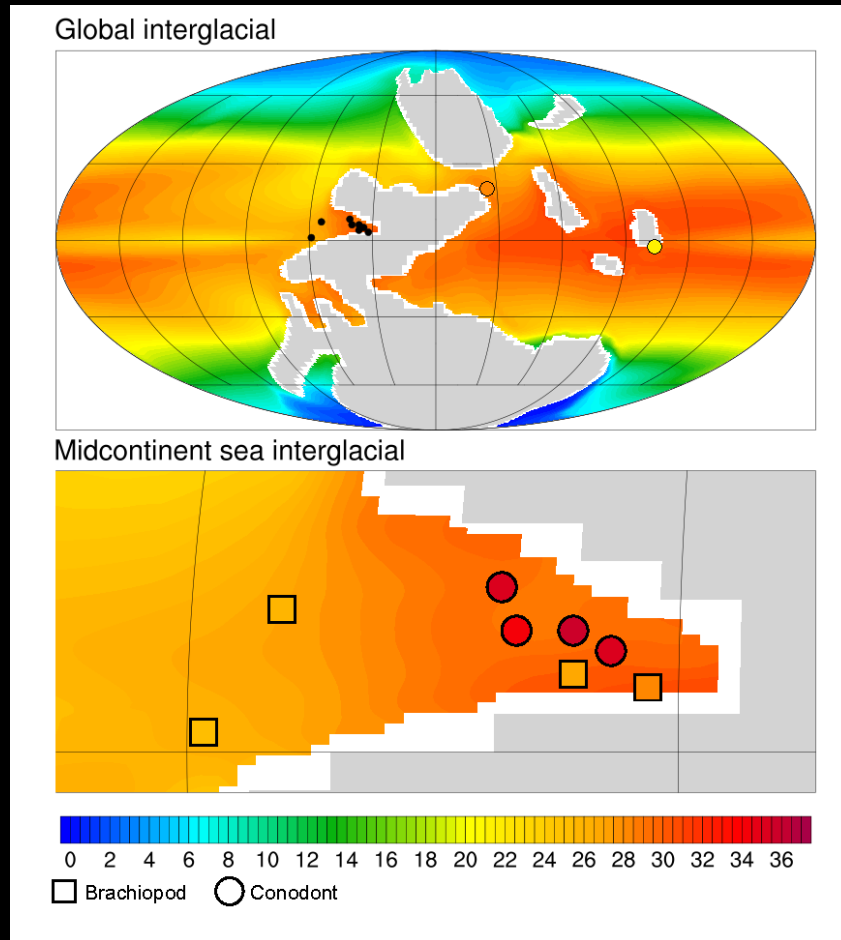
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Usually assumed
to be $\sim 0\text{‰}$

Results: Data-model comparison

$\delta^{18}\text{O}_{\text{sw}} = 0\text{‰}$



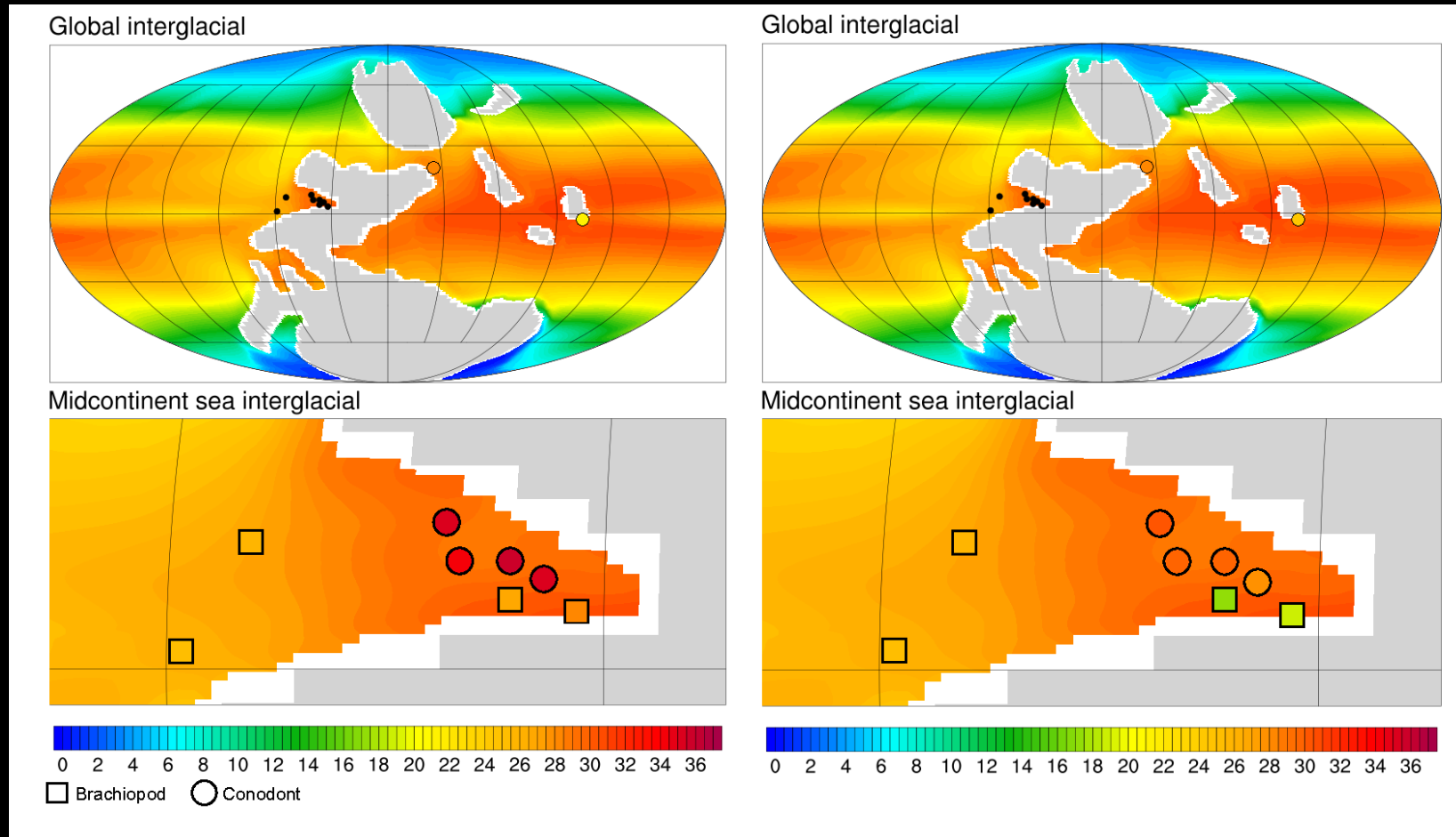
Mean annual SSTs and
proxy-derived
temperatures ($^{\circ}\text{C}$)

Proxy-derived
temperatures are too
high compared to
simulated SSTs
(Joachimski & Lambert,
2015; Roark et al., 2017)

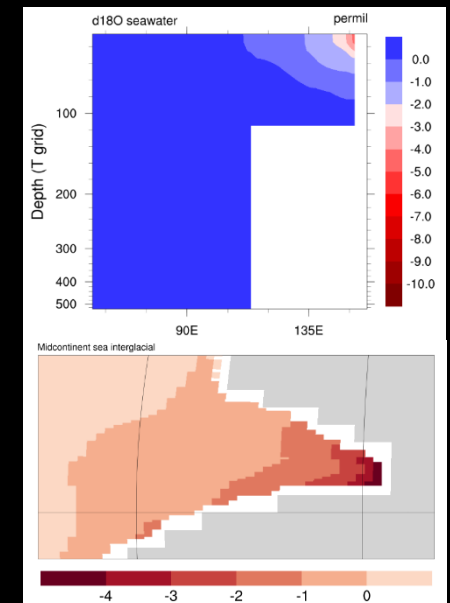
Results: Data-model comparison

$\delta^{18}\text{O}_{\text{sw}} = 0\text{‰}$

Model $\delta^{18}\text{O}_{\text{sw}}$



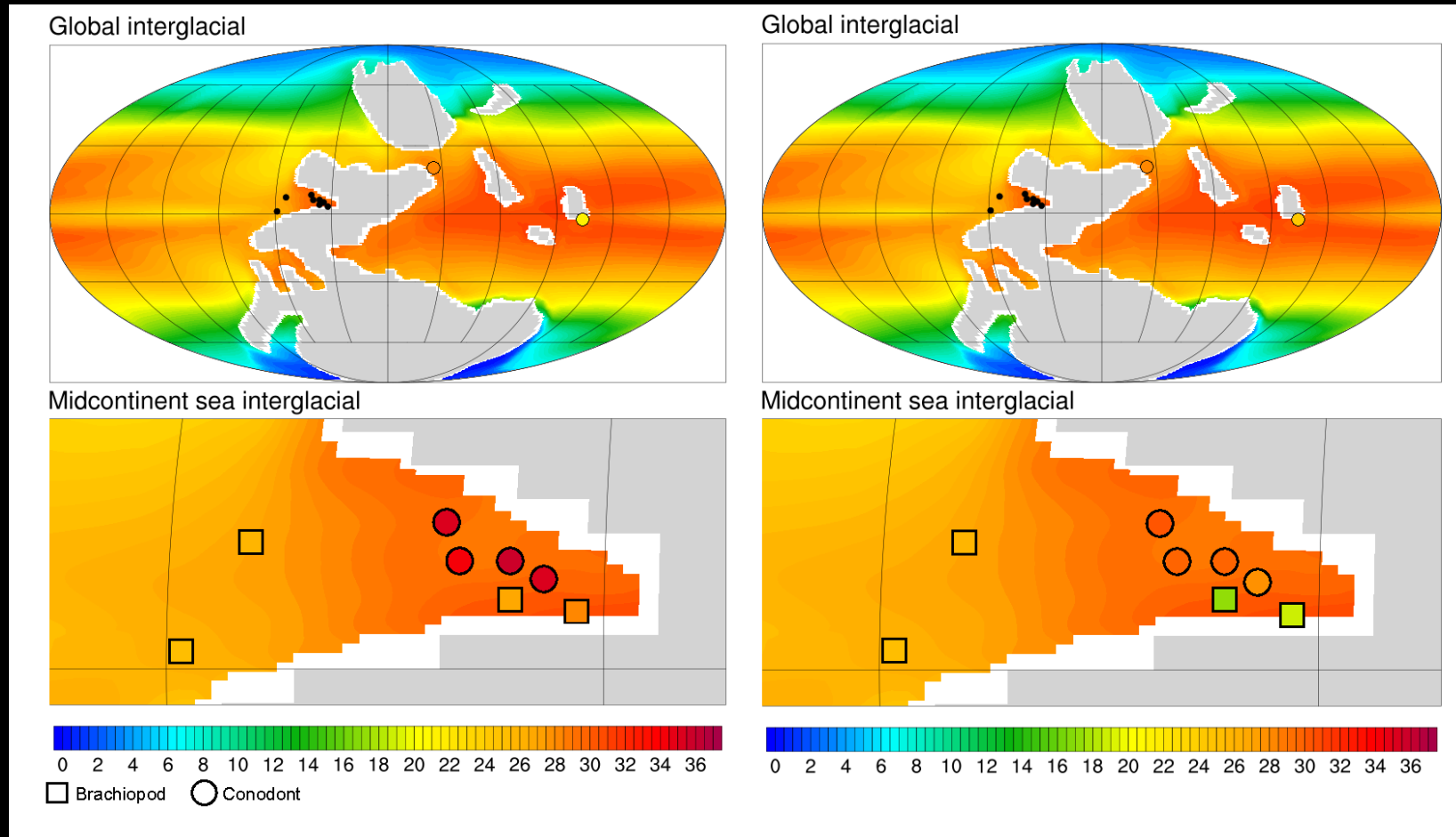
Conservative $\delta^{18}\text{O}_{\text{sw}}$



Results: Data-model comparison

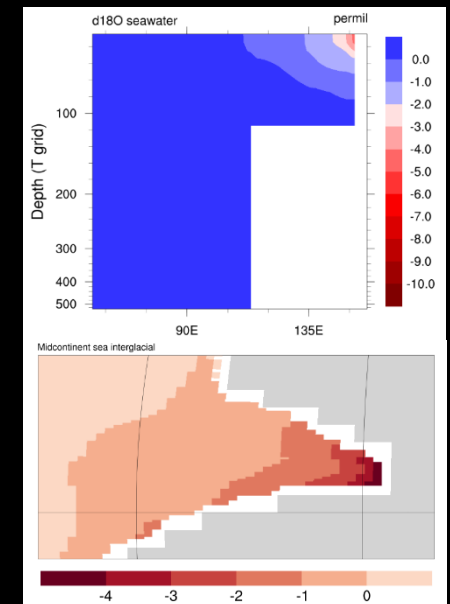
$\delta^{18}\text{O}_{\text{sw}} = 0\text{‰}$

Model $\delta^{18}\text{O}_{\text{sw}}$



Proxy-derived temperatures more closely resemble simulated SSTs

Conservative $\delta^{18}\text{O}_{\text{sw}}$



Conclusions

- Decrease in seawater $\delta^{18}\text{O}$ across the Midcontinent sea during interglacial high stands due to excess runoff and precipitation
 - Magnitude of seawater $\delta^{18}\text{O}$ depletion is dependent on regional bathymetry

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 - Magnitude of seawater $\delta^{18}\text{O}$ depletion is dependent on regional bathymetry
- Knowledge of the regional hydrologic processes that influence epicontinental seas may improve $\delta^{18}\text{O}$ paleotemperature estimates in deep time
 - Such processes can be constrained with isotope-enabled Earth system models (i.e. iCESM)



Thank you!