

OCE-1559153.

### Paleoceanography and Paleoclimatology

RESEARCH ARTICLE 10.1029/2019PA003644

#### Key Points:

- Modeled deep ocean tidal dissipation approximately doubled during the LGM, but the magnitude is dependent on LGM ice sheet extent
- Increase in LGM tidal mixing

#### Glacial Ice Sheet Extent Effects on Modeled Tidal Mixing and the Global Overturning Circulation

S.-B. Wilmes<sup>1,2</sup>, A. Schmittner<sup>1</sup>, and J. A. M. Green<sup>2</sup>

<sup>1</sup>College of Environmental, Atmospheric and Ocean Sciences, Oregon State University, Corvallis, OR, USA, <sup>2</sup>School of Ocean Sciences, Bangor University, Bangor, UK

# Modeling Tidal Mixing in the Glacial Ocean

#### **CESM Ocean Modeling Group Meeting, April 15, 2020**

### AGU1



Andreas Schmittner<sup>1</sup>, **Sophie-Berenice Wilmes<sup>1,2</sup>**, J. A. Mattias Green<sup>2</sup>, G. Danabasoglu<sup>3</sup>

<sup>1</sup>Oregon State University

<sup>2</sup>Bangor University

<sup>3</sup>NCAR

## Methods

### **Tide Model Simulations**

- OTIS (Oregon State Tidal Inversion Software)
- M2, S2, O1, K1
- Different resolutions (up to 1/12°)
- Different Internal Tide (IT) Drag Parameterizations
- Different LGM bathymetries (ICE5G, ICE6G)

#### **Climate Model Simulations**

- University of Victoria (UVic) model
- Tidal Mixing Parameterization (Jayne, St. Laurent, Simmons)
- Model of Ocean Biogeochemistry & Isotopes (MOBI) includes paleo tracers  $\delta^{13}$ C, and radiocarbon
- Vary Southern Ocean buoyancy fluxes



# **Tidal Dissipation**



Wilmes et al., (2019)





## **Tidal Mixing Parameterization**

Diapycnal Diffusivity:  $k_{\rm u} = k_{\rm l}$ 

### $k_{\rm bg} = 0.3 \times 10^{-4} \, {\rm m}^2/{\rm s}$

Subgrid-scale bathymetry:

 $\epsilon = \frac{1}{\rho} \sum_{z'>z}^{H} \sum_{f}^{TC} q_{TC} D_{IT,TC}(x, y, z') F(z, z'),$ Dissipation Efficiency = Fraction of locally dissipated energy

$$bg + \frac{\Gamma\epsilon}{N^2},$$

#### **Considers only locally dissipated** energy, which is only 1/3 of the total!

$$q_{\rm TC} = \begin{cases} 1, \text{for}|y| > y_{c,\rm TC} \\ 0.33, \text{otherwise.} \end{cases}$$

Schmittner & Egbert (2014) Geosc. Mod. Devel.

![](_page_3_Picture_9.jpeg)

![](_page_3_Picture_10.jpeg)

### Saltier AABW, shallower & weaker AMOC

Reduced atmospheric meridional moisture flux in Southern Hemisphere

Depth (m) 2000 3000  $\mu_{SH}$  = 0 4000 5000 1000 Depth (m) 2000  $\mu_{\mathsf{SH}}$  = 0.1 3000 4000 AABWA 5000 1000  $\mu_{\rm SH} = 0.25 \stackrel{(i)}{=} 100$ 2000 3000 4000 5000 1000 (E) 2000 Depth  $\mu_{\mathsf{SH}}$  = 0.5 300 4000 500 1000 Depth (m) 2000 3000  $\mu_{SH}$  = 1 4000 500 -40 -20 Latitude (°N)

Wilmes et al., (in prep.)

**PD tides** 

![](_page_4_Picture_7.jpeg)

![](_page_4_Picture_8.jpeg)

![](_page_4_Picture_9.jpeg)

![](_page_4_Picture_10.jpeg)

![](_page_4_Picture_11.jpeg)

0

**ICE-6G tides** 

![](_page_4_Picture_13.jpeg)

![](_page_4_Figure_14.jpeg)

![](_page_4_Figure_15.jpeg)

![](_page_4_Figure_16.jpeg)

![](_page_4_Figure_17.jpeg)

5

Streamfunction (Sv)

**ICE-5G tides** 

![](_page_4_Picture_19.jpeg)

![](_page_4_Picture_20.jpeg)

![](_page_4_Figure_21.jpeg)

![](_page_4_Figure_22.jpeg)

![](_page_4_Figure_23.jpeg)

LGM tidal mixing

- Increases k<sub>v</sub>
- Increases AMOC
- Increases AABW

![](_page_5_Figure_4.jpeg)

Wilmes et al., (in prep.)

### Radiocarbon

![](_page_6_Figure_1.jpeg)

Wilmes et al., (in prep.)

#### PD tides

**ICE-6G tides** 

![](_page_6_Figure_6.jpeg)

![](_page_6_Figure_7.jpeg)

![](_page_6_Figure_8.jpeg)

![](_page_6_Figure_9.jpeg)

![](_page_6_Figure_10.jpeg)

![](_page_6_Figure_11.jpeg)

![](_page_6_Figure_12.jpeg)

![](_page_6_Figure_13.jpeg)

![](_page_6_Figure_14.jpeg)

μ <sub>SH</sub> = 0	Depth (m)	1000 2000 3000 4000 5000	AABWA -4.1Sv
μ <sub>SH</sub> = 0.1	Depth (m)	1000 2000 3000 4000 5000	AABWA -4.1Sv
μ <sub>SH</sub> = 0.25	Depth (m)	1000 2000 3000 4000 5000	AABWA -3.7Sv
μ <sub>SH</sub> = 0.5	Depth (m)	1000 2000 3000 4000 5000	AABWA -3.6Sv
μ <sub>SH</sub> = 1	Depth (m)	1000 2000 3000 4000 5000	AABWA -2.5Sv

-60

Wilmes et al., (in prep.)

δ13**C** 

PD tides

AMOC 0,1Sv

AMOC 7,1Sv

AMOC 9,1Sv

AMOC 10.7Sv

![](_page_7_Figure_9.jpeg)

**ICE-6G tides** 

![](_page_7_Figure_11.jpeg)

AABWA

AABWA

AABWA

AABWA

-60

![](_page_7_Figure_12.jpeg)

![](_page_7_Figure_13.jpeg)

![](_page_7_Picture_14.jpeg)

![](_page_7_Picture_15.jpeg)

![](_page_7_Picture_16.jpeg)

![](_page_7_Figure_17.jpeg)

![](_page_7_Figure_18.jpeg)

![](_page_7_Figure_19.jpeg)

40

AMOC 8,5Sv

AMOC 10.1Sv

AMOC 12.1Sv

AMOC 15.1Sv

60

.

 $\delta^{13}$ C (permil)

-0.5

![](_page_7_Picture_26.jpeg)

![](_page_8_Figure_0.jpeg)

#### **Atlantic Profiles**

![](_page_8_Figure_2.jpeg)

# Conclusions

- Increased tidal mixing in LGM is robust result, but quantitatively depends on reconstructed bathymetry (basin geometry; land ice extent)
- Increased diffusivities increase AMOC & AABW flow rates
- MOC geometry (AMOC depth) strongly affects isotopes
- Effect of increased mixing is more subtle but improves model-data agreement
- All these results are conservative because they neglect changes in remotely dissipated energy
  - Would be good to include remotely dissipated energy in future simulations e.g. by using Eden & Olbers parameterization (who wants to collaborate on this?)