

Pushing the Limits of Pattern Scaling to Separate Tropical and Polar Climate Change Signals

Stephanie Hay and Paul Kushner

Department of Physics, University of Toronto

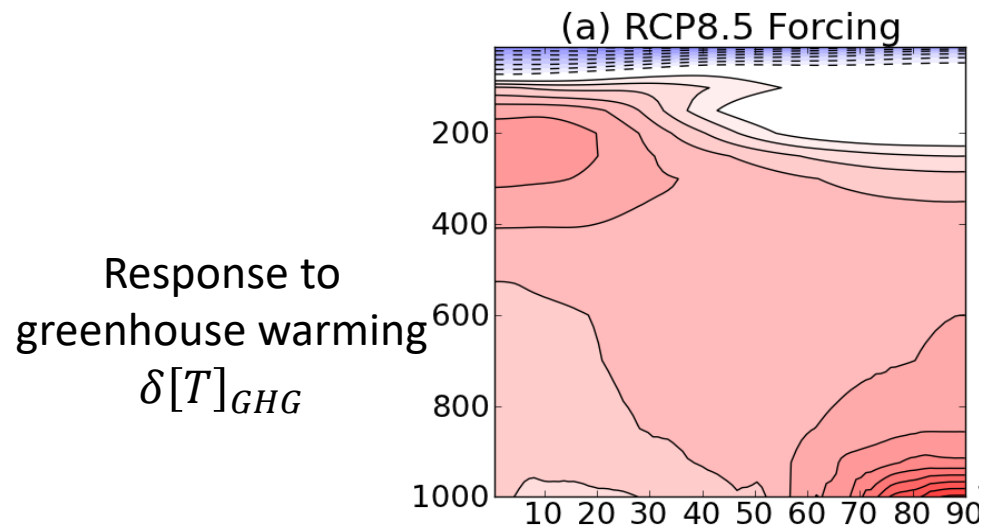
R Blackport (*U. Exeter/ECCC*), K McCusker (*Rhodium*), T. Oudar (*CNRM/U. Toulouse*), L. Sun (*CSU*), M. England (*SIO/UCSD*), J. Screen (*U. Exeter*), C. Deser (*NCAR*), L. Polvani (*Columbia U.*)

Key points:

- **We use pattern scaling to deal with different sea ice loss protocols and models.**
- **This analysis brings out both Arctic and tropical drivers.**
- **Pattern scaling fails in the North Atlantic sector given nonlinearity and experimental/model inconsistency.**



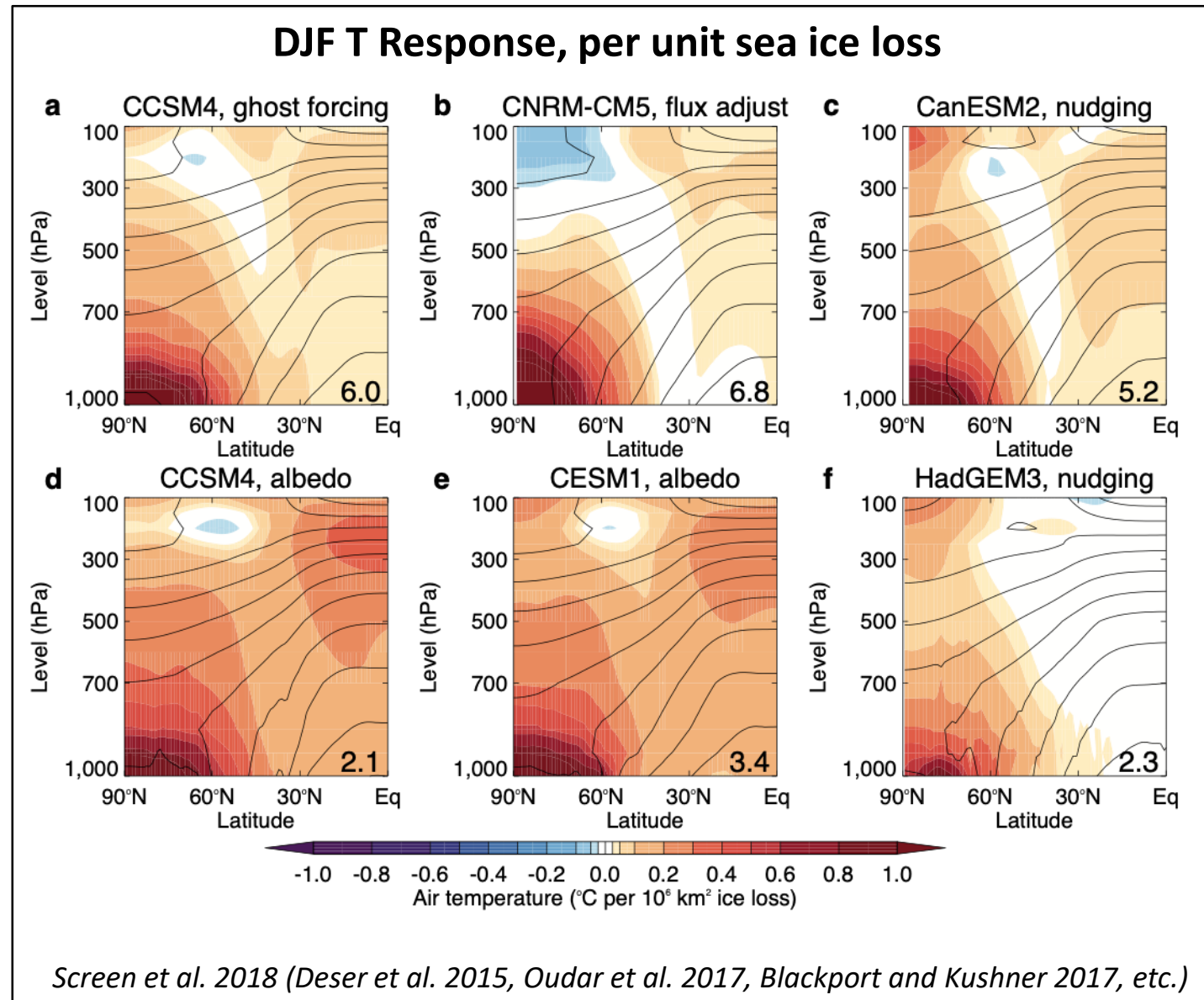
Response and Sensitivity for Zonal- and Annual-Mean Temperature, $[T]$



To account for cross coupling, Blackport and Kushner developed a multi-parameter pattern scaling approach.

Mini Global Warming in the Response to Sea Ice Loss

- Coupled ocean-atmosphere models forced by sea ice loss exhibit *cross coupling* between Arctic and tropical responses.



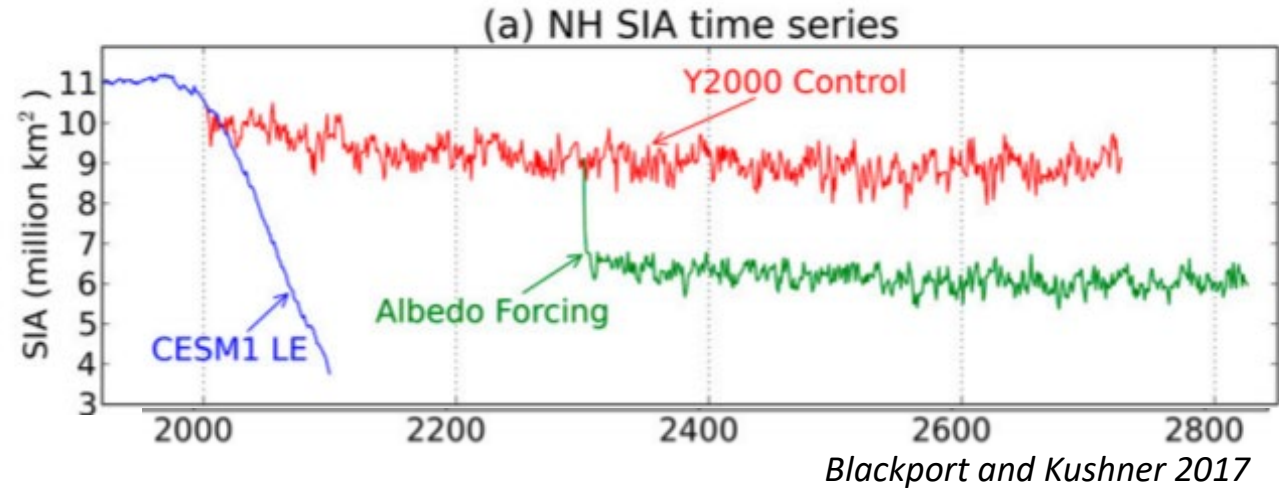
Different models and experimental protocols

⇒ *Model Experiments (MEs)*

Model: CESM1:

Experiments:

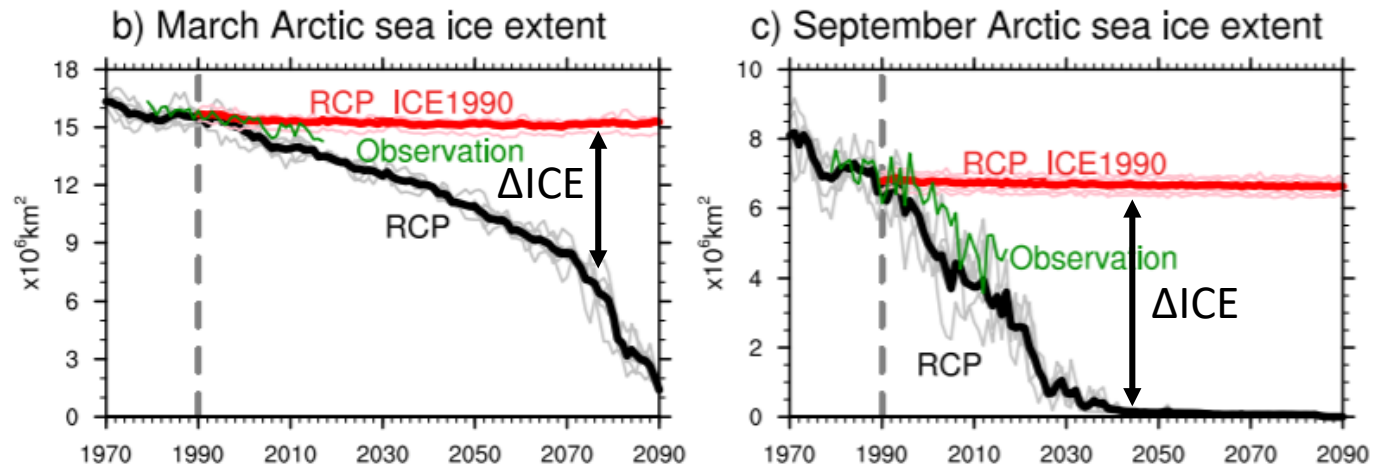
- Projected greenhouse warming (transient LENS RCP8.5)
- Albedo reduction (time slice)



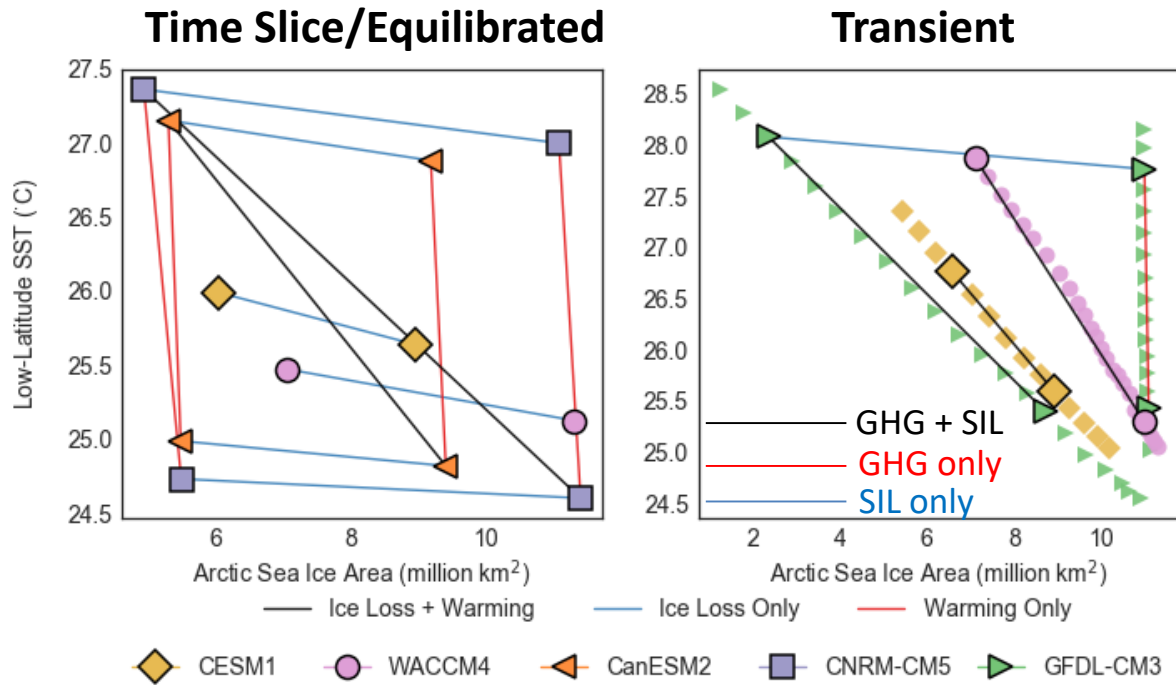
Model: GFDL-CM3

Experiments:

- RCP8.5 (transient)
- Sea ice nudging (transient)



Sun et al. 2018



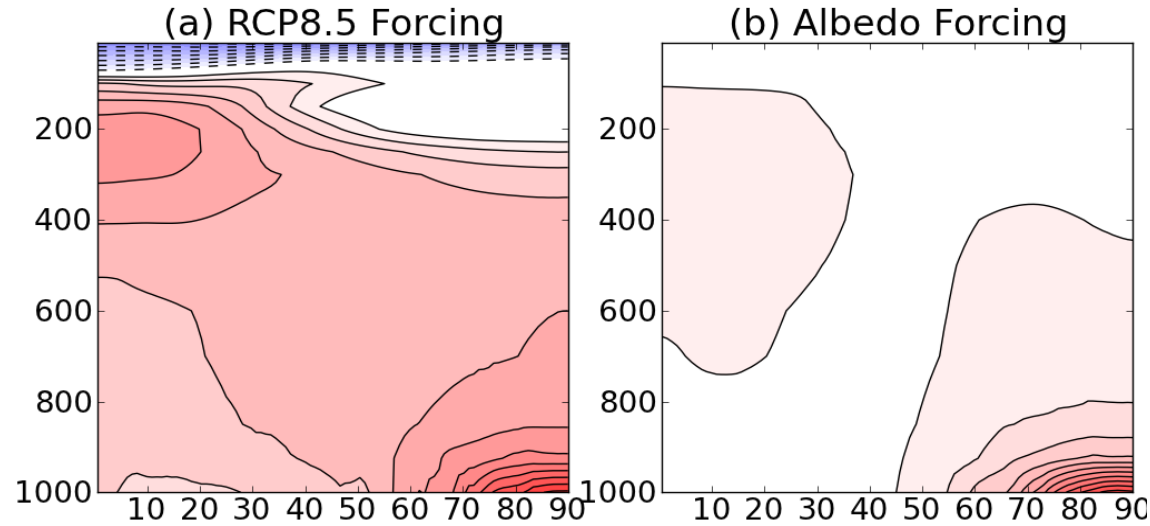
We combined five sets of model experiments (MEs), all featuring cross coupling.

We use pattern scaling (e.g *Harvey et al. 2012, Blackport and Kushner 2017*) to find *sensitivity patterns*

- Sensitivity $\left(\frac{\partial Z}{\partial T_l}\right)_I$ of field Z to low-latitude warming in absence of sea ice loss.
- Sensitivity $\left(\frac{\partial Z}{\partial I}\right)_{T_l}$ of field Z to change in sea ice in absence of low-latitude warming.

Response and Sensitivity for Zonal- and Annual-Mean Temperature, $[T]$

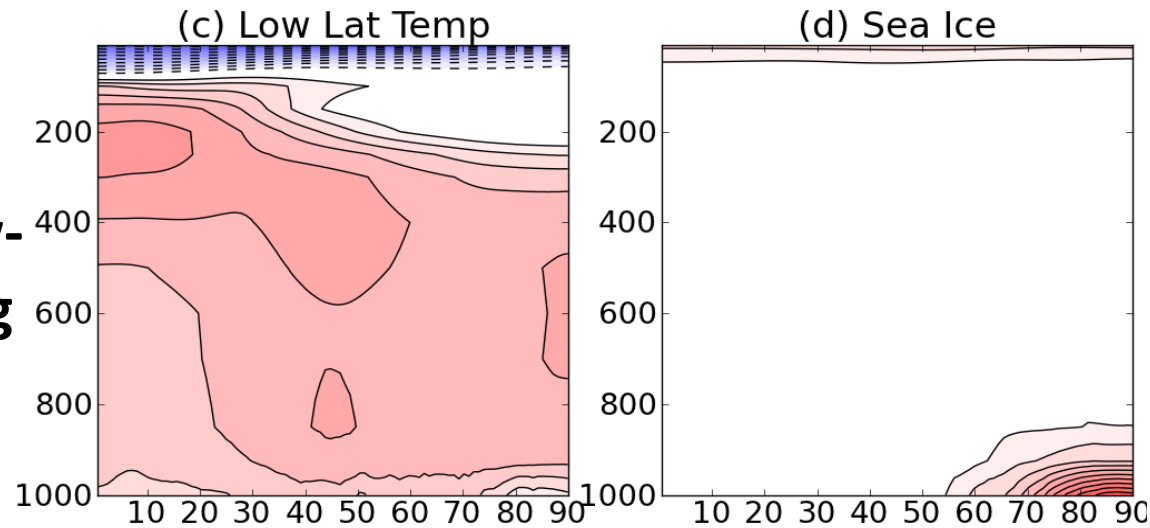
Response to
greenhouse
warming
 $\delta[T]_{GHG}$



Response to
albedo change
 $\delta[T]_{albedo}$

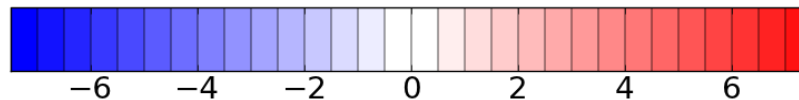
Sensitivity to low-
latitude warming

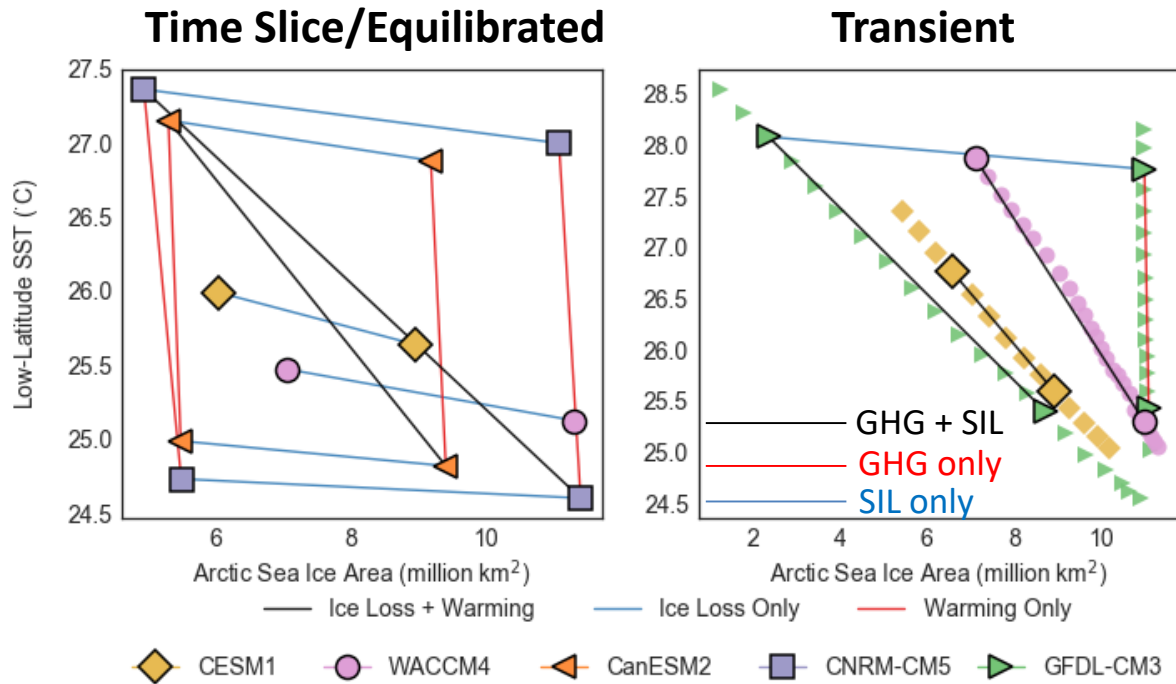
$$\left(\frac{\partial[T]}{\partial T_l}\right)_I$$



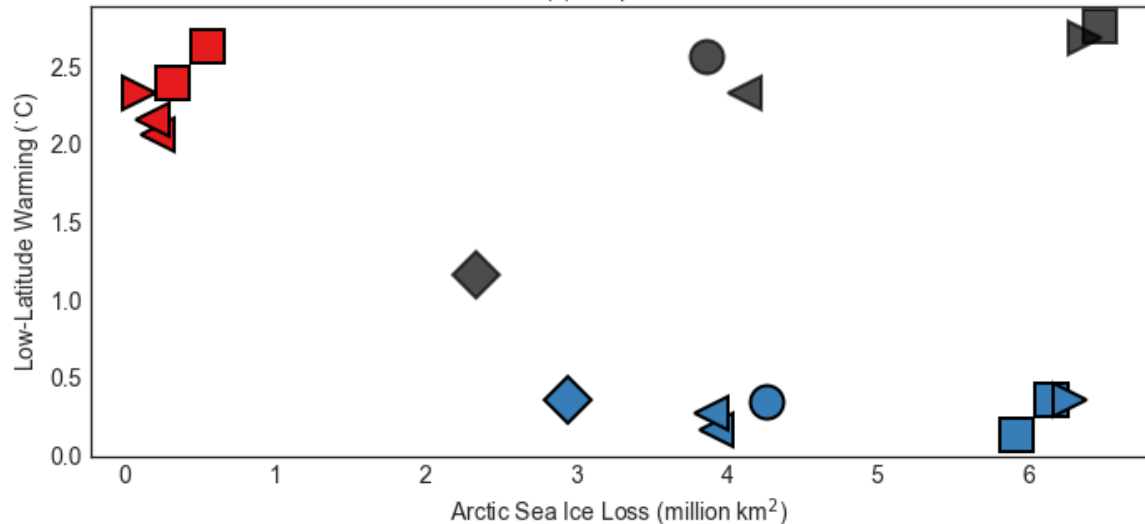
Sensitivity to
sea ice loss

$$\left(\frac{\partial[T]}{\partial I}\right)_{T_l}$$





(c) Responses



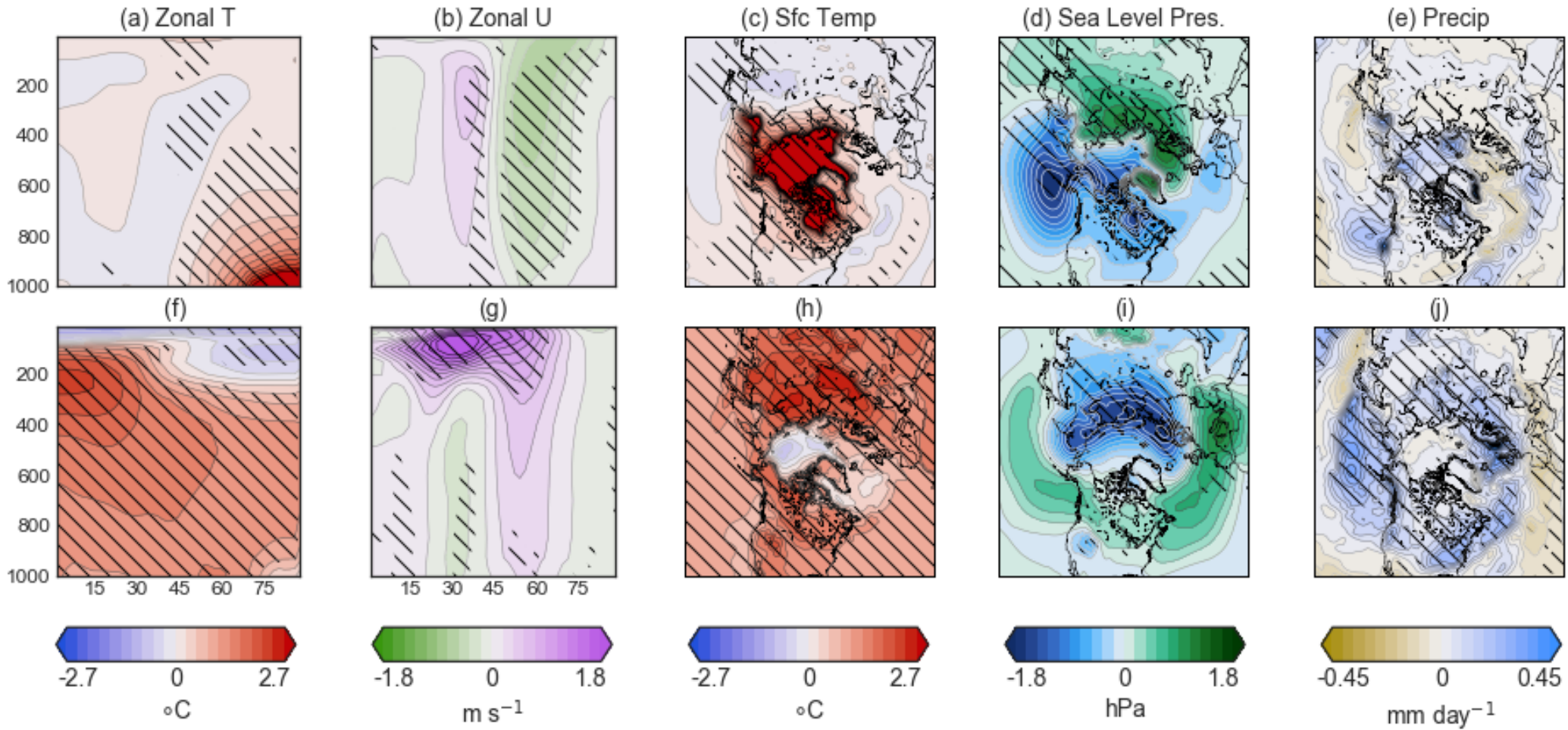
In this ensemble of opportunity, which atmosphere-ocean sensitivities are robust?

How do the sea ice loss responses relate to the rest of global warming?

And where will we hit the limits of pattern scaling approaches?

Multi-Model Mean Atmospheric Sensitivity Patterns

Sensitivity to ice loss
(per 2M km²
Arctic ice loss)

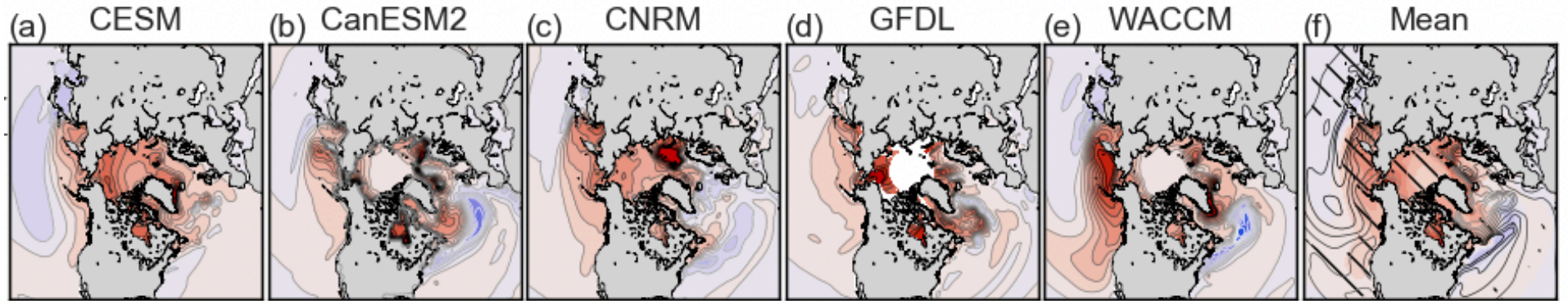


Sensitivity to low-
latitude warming
(per 1°C
low-latitude warming)

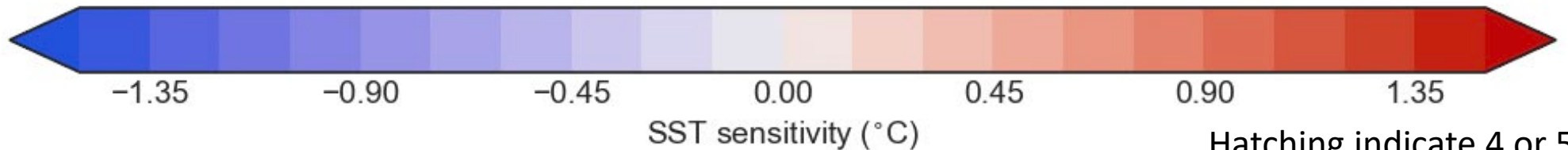
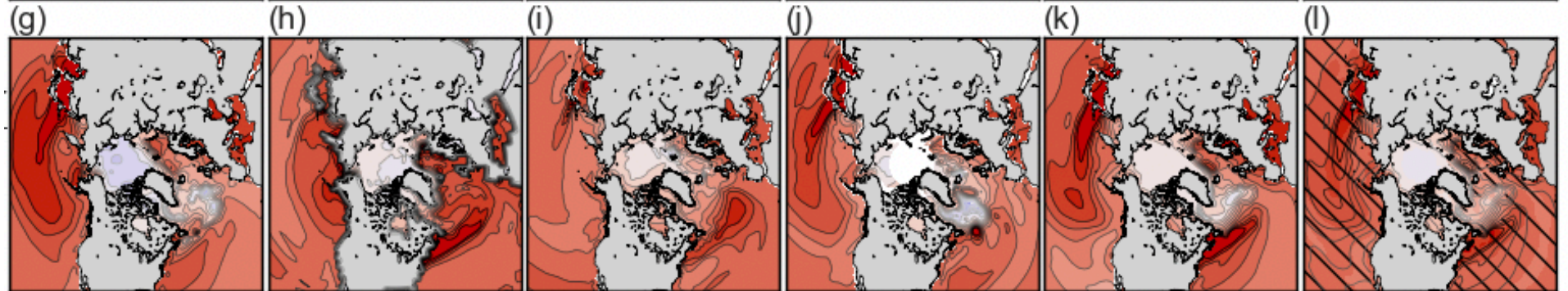
DJF climatological and multi-model mean. Hatching indicate 4 or 5 out of 5 experiments agree in sign

Sensitivity Patterns for the Ocean: SST

Sensitivity to ice loss
(per 2M km^2
Arctic ice loss)



Sensitivity to low-
latitude warming
(per 1°C
low-latitude warming)

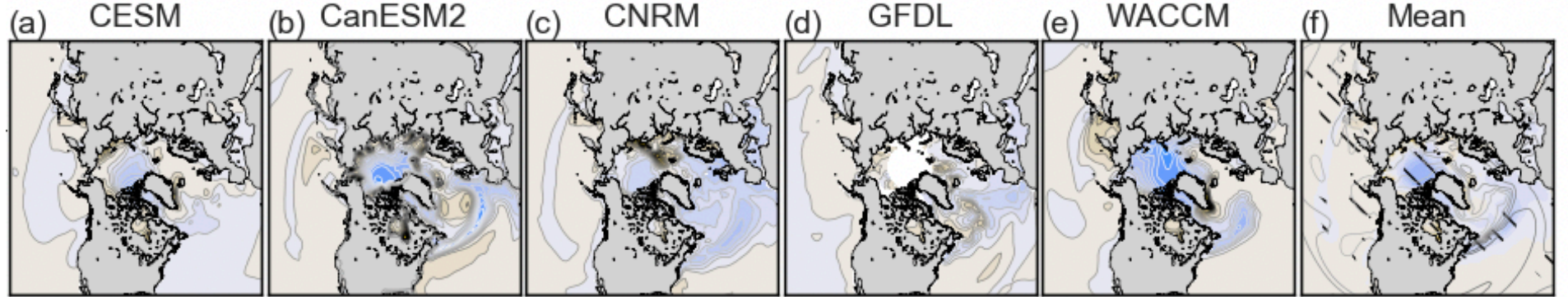


ANN climatological mean for individual and multiple models.

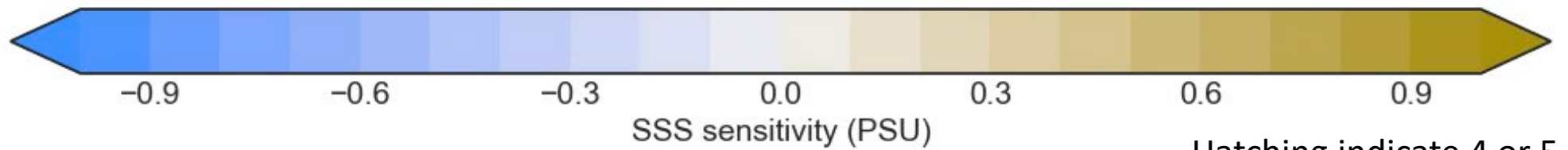
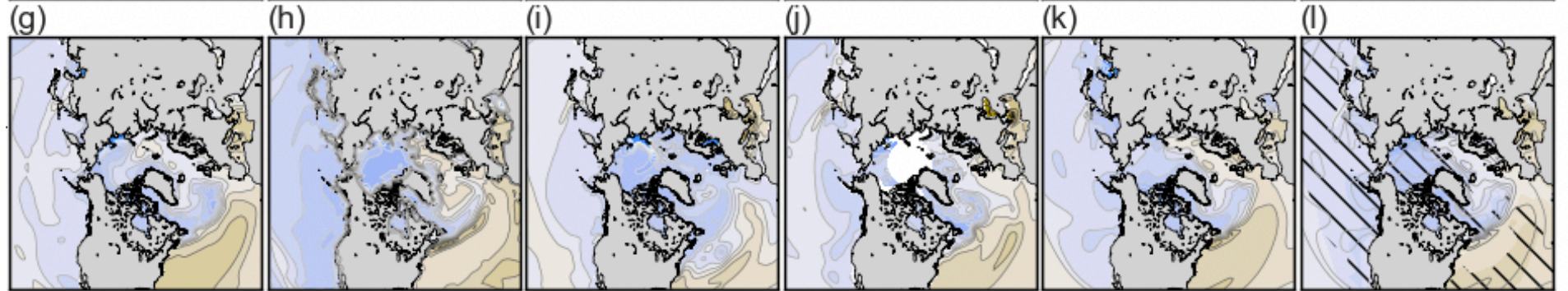
Hatching indicate 4 or 5
out of 5 experiments
agree in sign

Sensitivity Patterns for the Ocean: SSS

Sensitivity to ice loss
(per $2M \text{ km}^2$
Arctic ice loss)



Sensitivity to low-
latitude warming
(per 1°C
low-latitude warming)

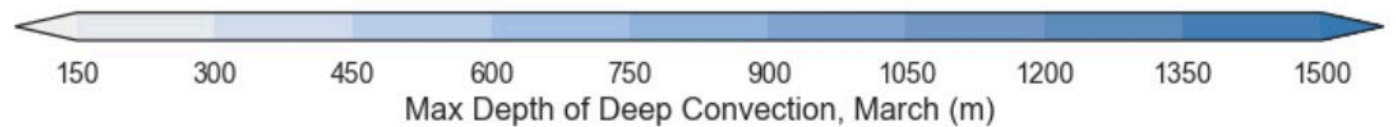
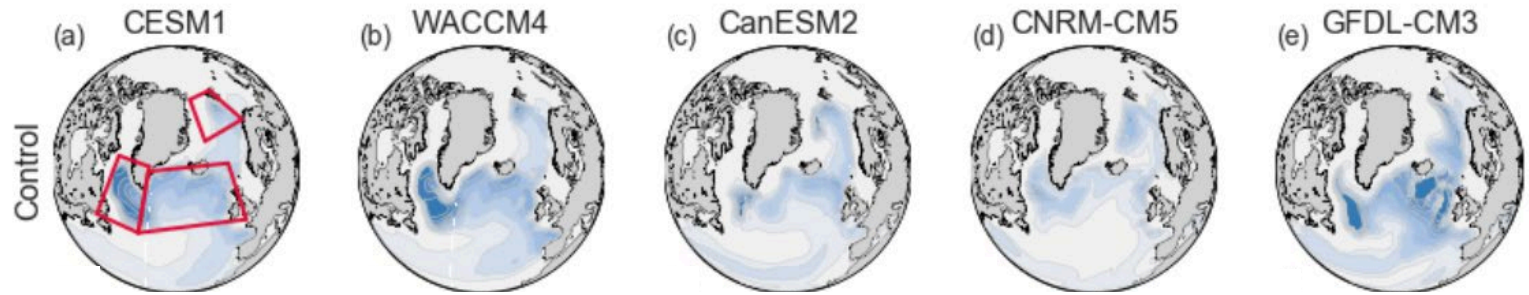


ANN climatological mean for individual and multiple models.

Hatching indicate 4 or 5
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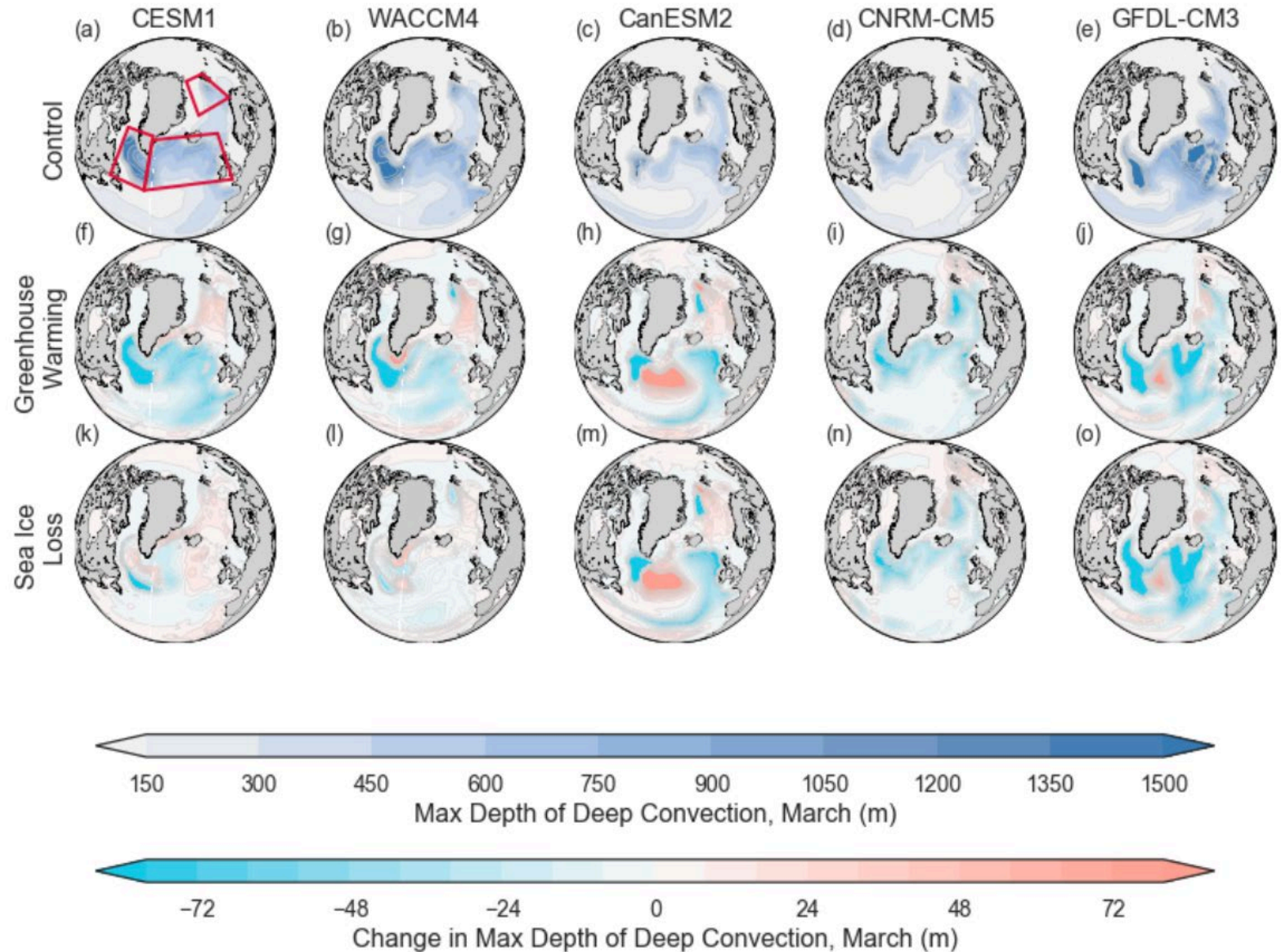
The North Atlantic: where pattern scaling breaks down ...

- Ocean convection changes are too ME dependent for pattern scaling.
 - Inconsistent control simulations
 - Different experimental protocols for sea ice loss.
- So we'll look at *responses* instead of *sensitivities*.



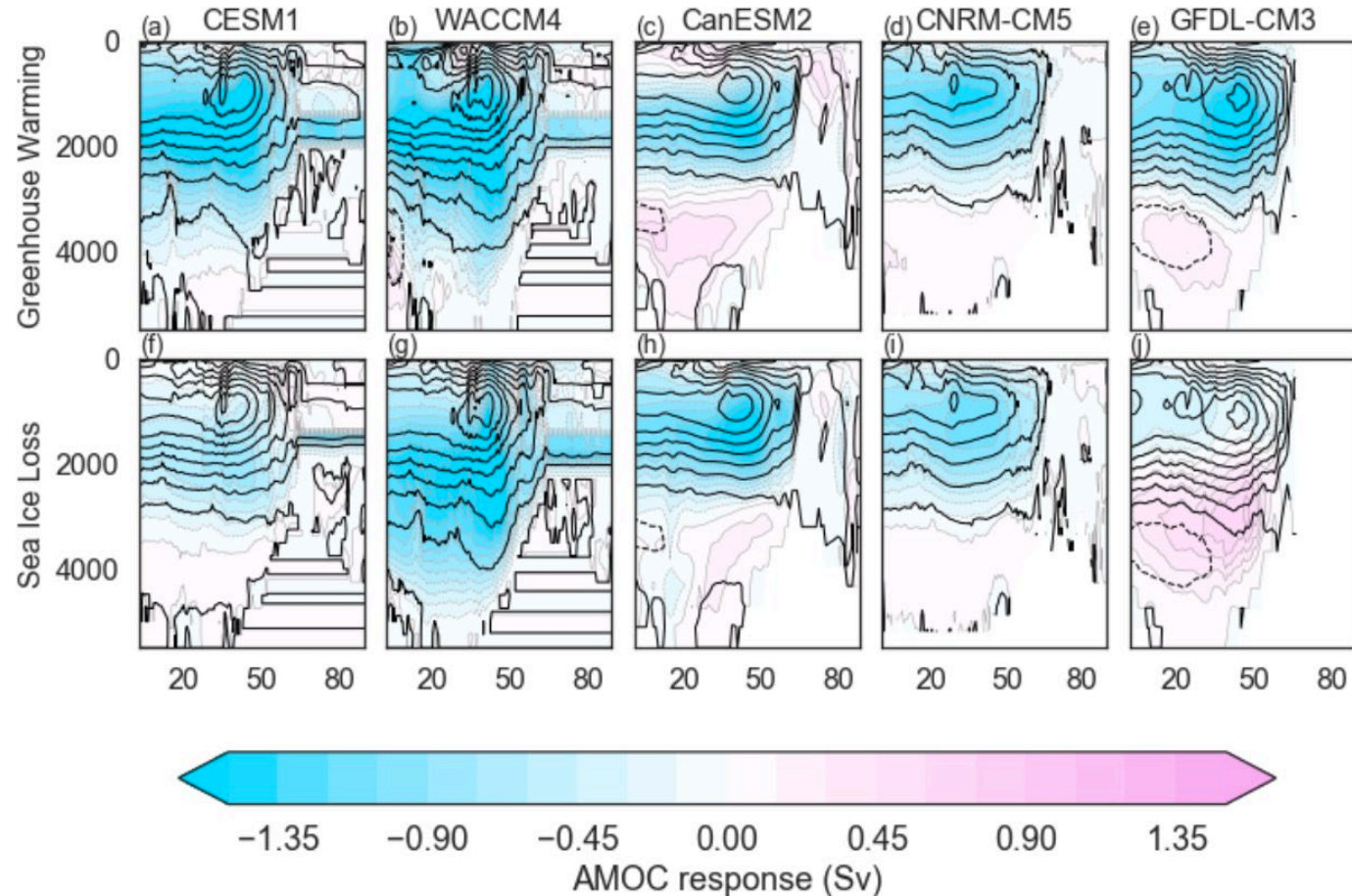
The North Atlantic: where pattern scaling breaks down ...

- Sea ice loss and greenhouse warming have mixed impacts on convection.
- Sometimes sea ice loss can be isolated as the primary driver, and sometimes not.



The North Atlantic: where pattern scaling breaks down ...

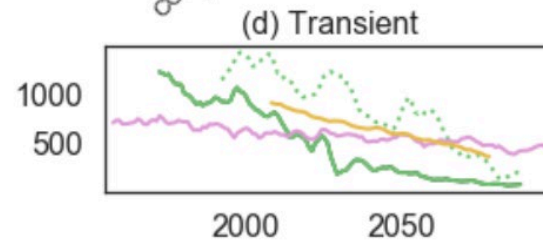
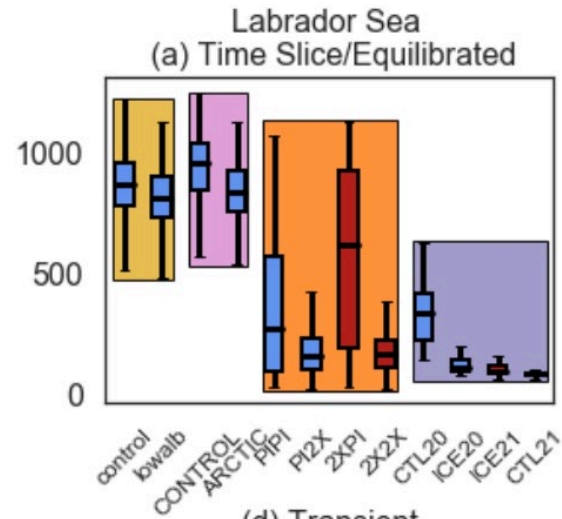
- ME dependence is also a feature of AMOC responses ...
- NCAR models with similar ocean components can have very different AMOC control simulations and responses.
- One model (CanESM2) shows a greater reduction in AMOC from sea ice loss than from greenhouse warming!



Conclusion

- Pattern scaling provides an estimate of model sensitivity kernels, to learn about the consequences of Arctic and tropical warming.
 - Sea ice loss → negative feedback on the circulation response (“tug of war”).
 - Precipitation response from low-latitude warming extends well into Arctic domain.
 - Salinity and precipitation changes from sea ice loss localized to Arctic.
- Model and experimental protocol dependence is too pronounced for pattern scaling to work in the North Atlantic.
 - Convection and AMOC responses are not robust.
- PAMIP (or similar) coordination required to obtain robust characterization of sensitivity of overturning circulation and convection.

Response of Deep Convection by Convection Zone

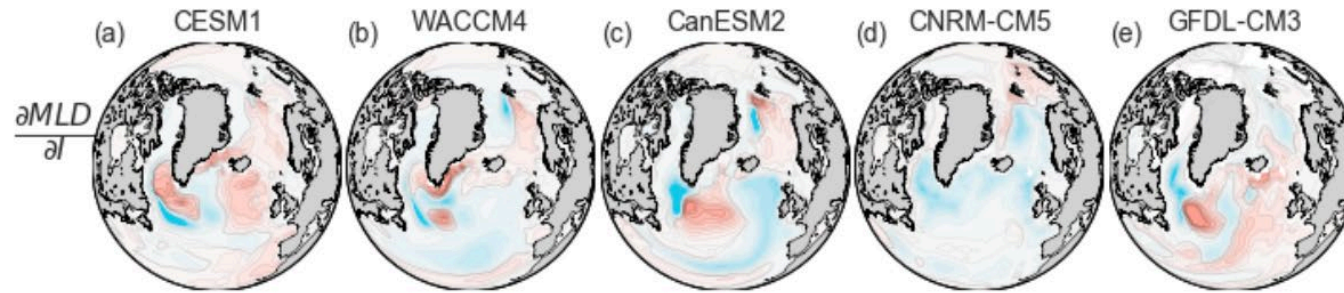


— RCP8.5 RCP8.5 w/o ice loss

■ CESM1
 ■ WACCM4
 ■ CanESM2
 ■ CNRM-CM5
 ■ GFDL-CM3

Sensitivity Patterns for the Ocean: North Atlantic Convection

Sensitivity to ice loss (per $2M \text{ km}^2$ Arctic ice loss)



Sensitivity to low-latitude warming (Per 1°C low-latitude warming)

