Permian (251 Ma) Climate Variability: The North Panthalassic Oscillation

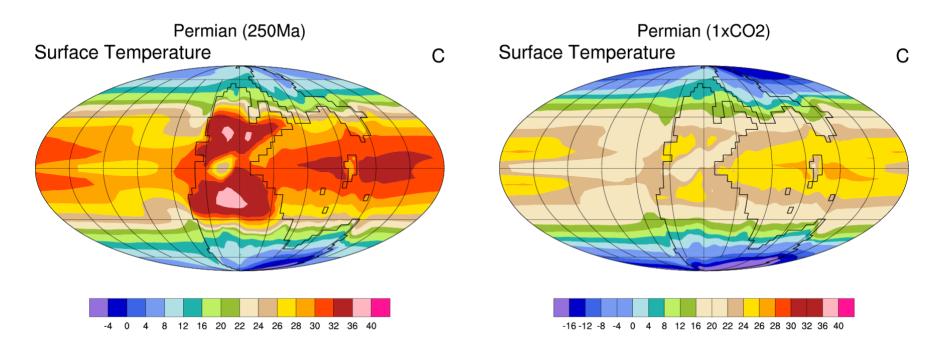
C.A. Shields and J.T. Kiehl NCAR/CGD/CCR/Paleo

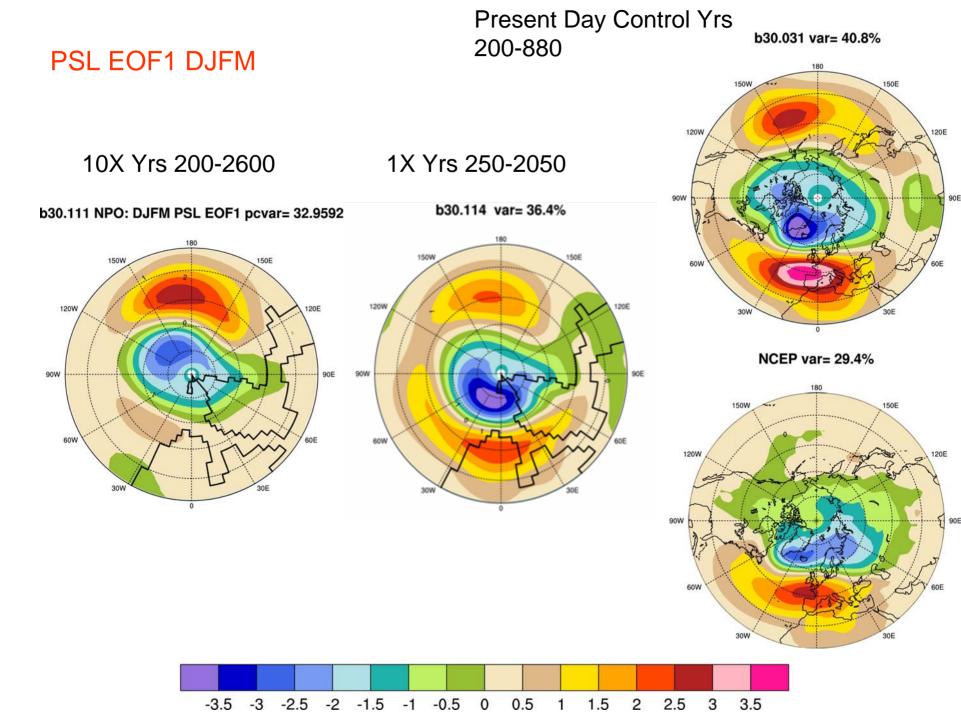
Many thanks to Clara Deser, Jeff Yin, and Bob Tomas for insightful discussions...

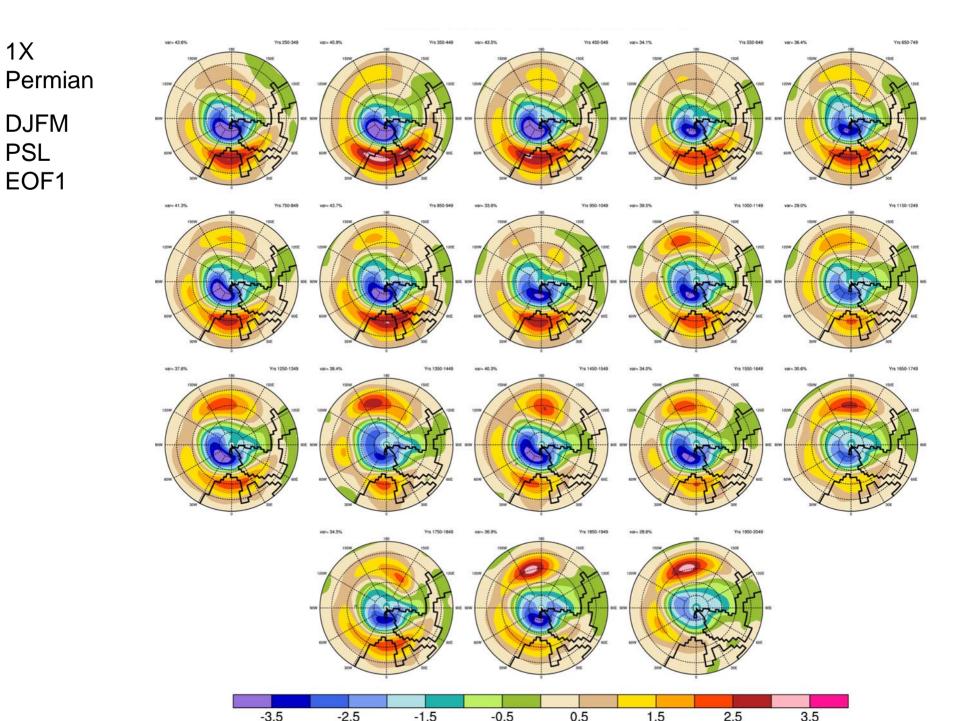
Permian Simulations

10X CO2 Fully Coupled CCSM3, T31_gx3v5, 2700 Yrs of simulation

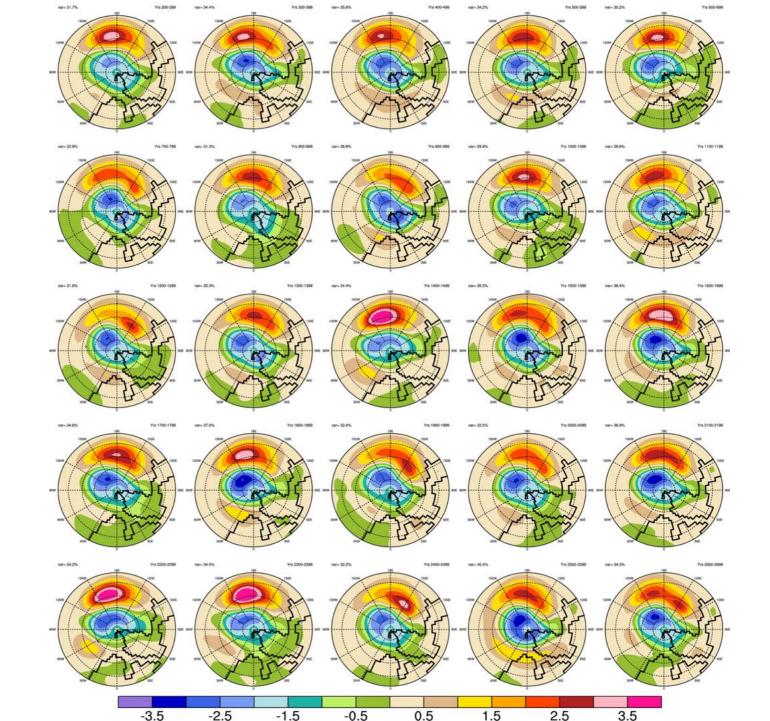
1X CO2 Fully Coupled CCSM3, T31_gx3v5, 2050 Yrs of simulation



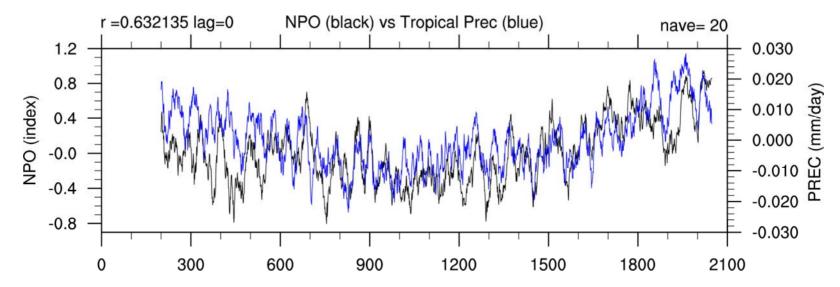




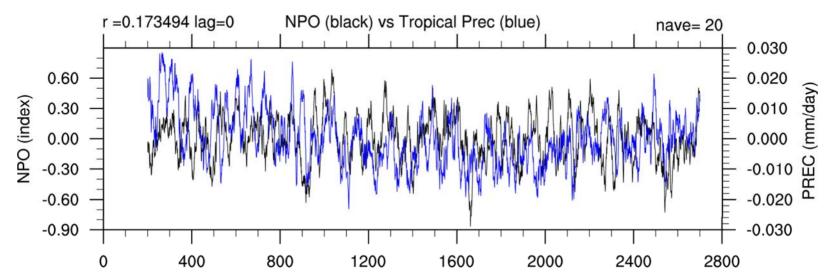
10X Permian DJFM PSL EOF1



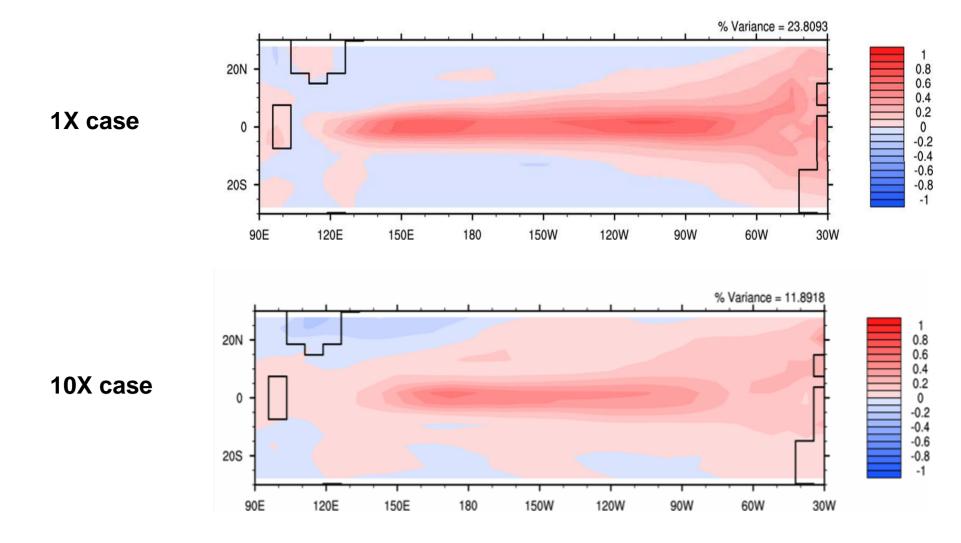
1X



10X



TS EOF1 last 100yrs of simulation



DJFM PRECIP

17 12

9

7

5

3 1

0.2

17 12

9

0.2

17 12

9

7

5

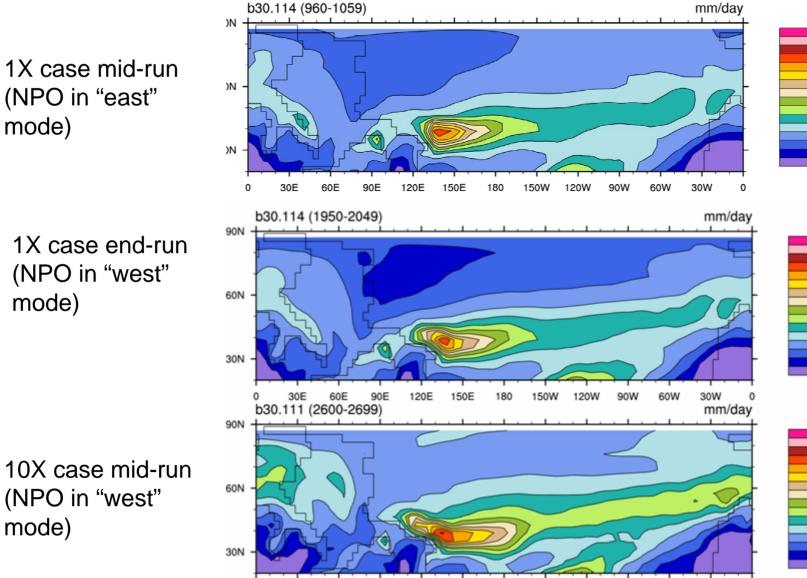
з 1

0.2

30W

0

60W



30E

0

60E

90E

120E

150E

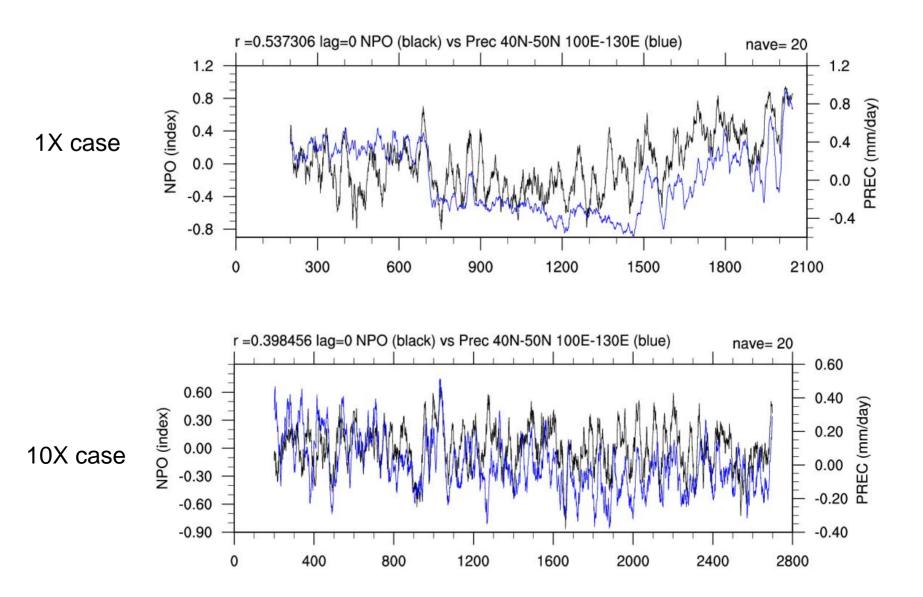
180

150W

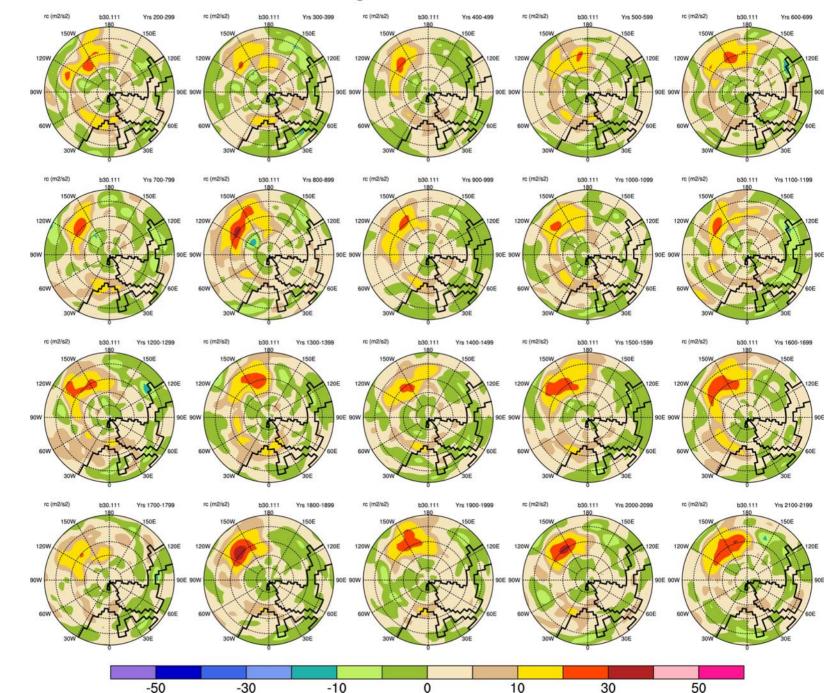
120W

90W

10X case mid-run (NPO in "west" mode)

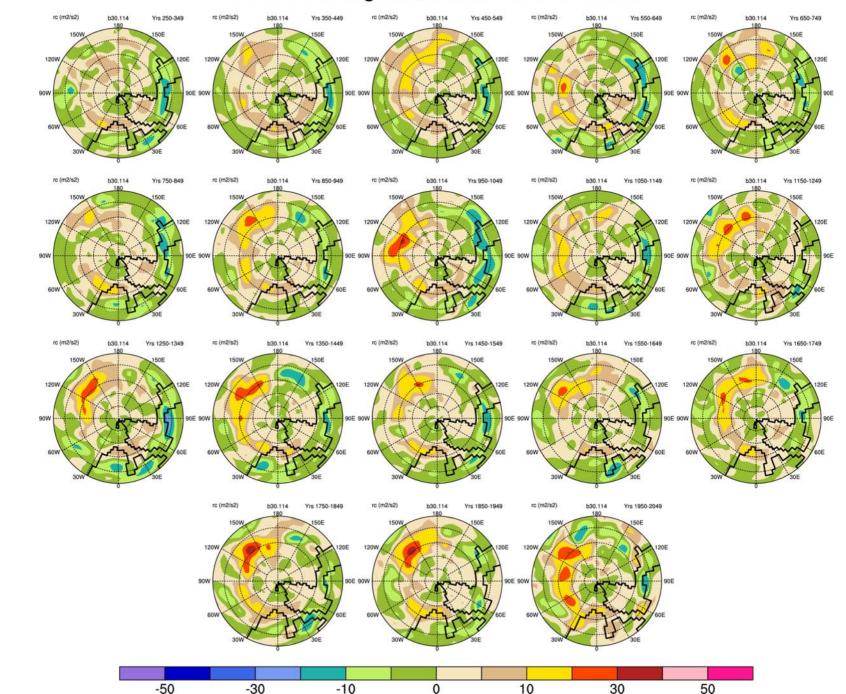


DJFM NPO regressed onto VU-V*U 250mb



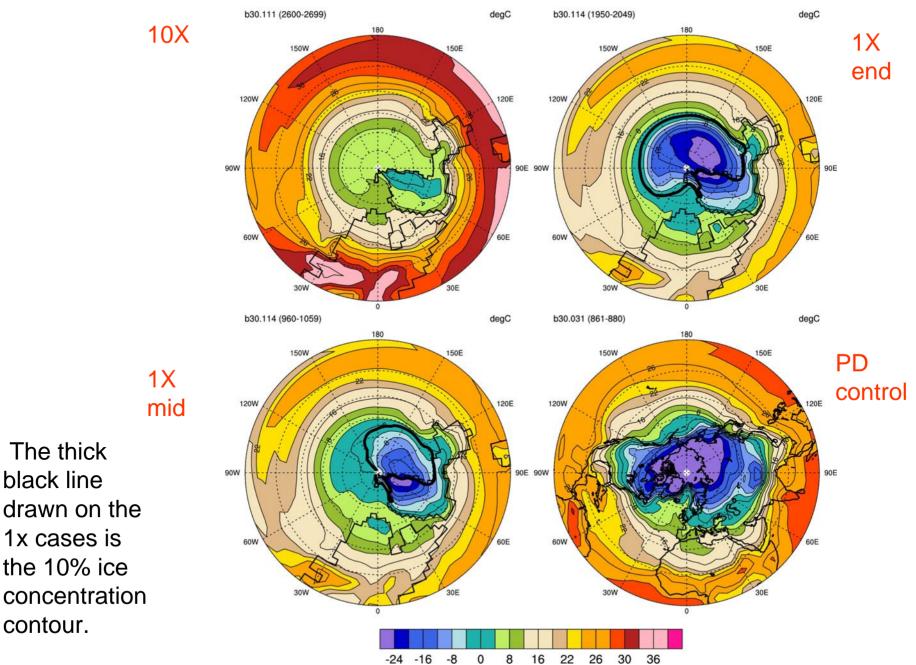
10X case

DJFM NPO regressed onto VU-V*U 250mb

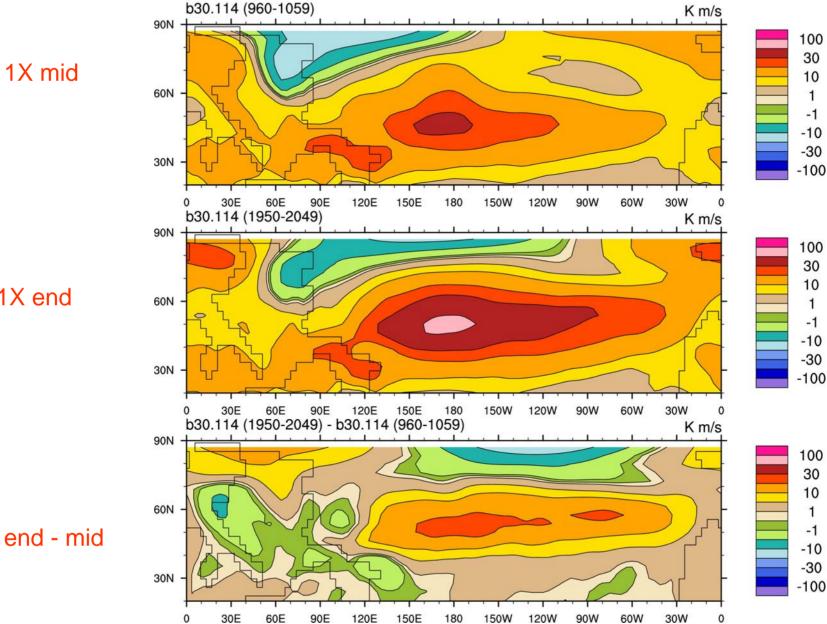


1X case

DJFM TS

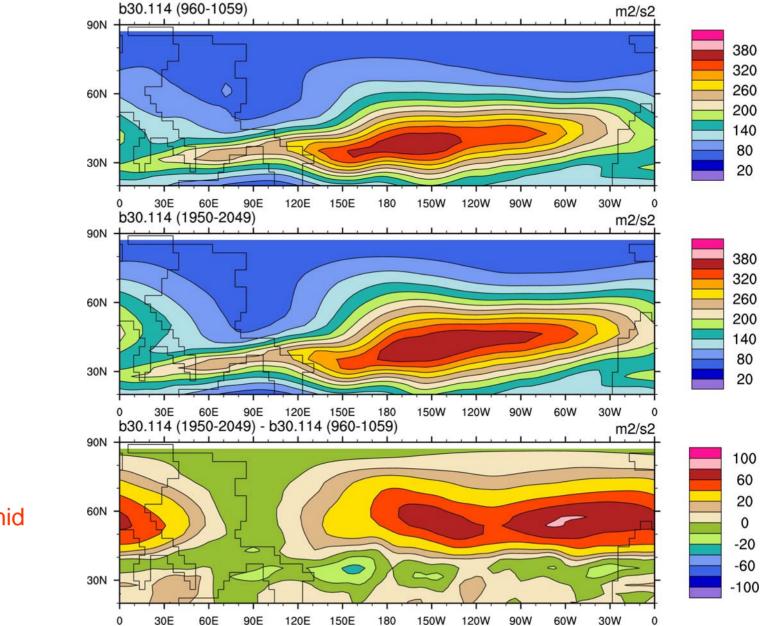


DJFM V'T' (VT-V*T) 850mb



1X end

DJFM EKE 250mb





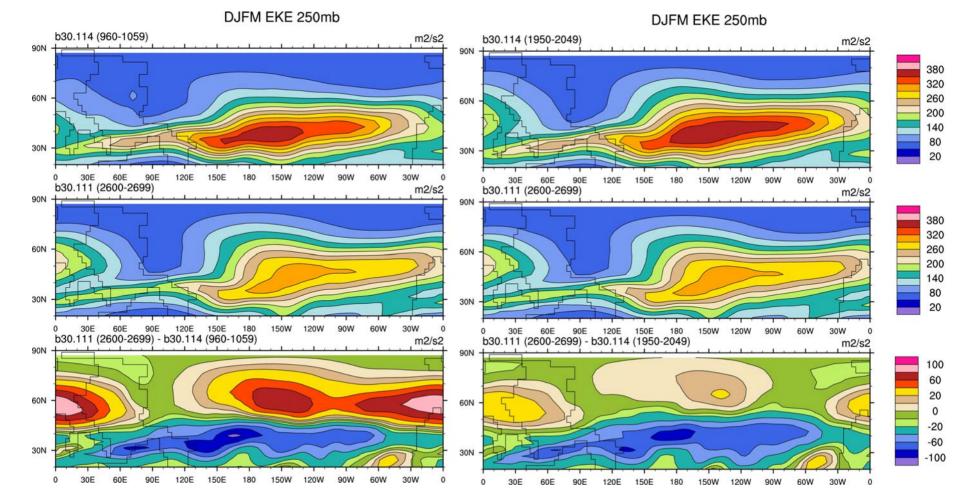




Winter Eddy Kinetic Energy 250mb

1Xmid, 10X, 10x – 1Xmid

1Xend, 10X, 10x - 1Xend



Conclusions

1) NPO/AO can shift modes, i.e. area of maximum PSL variability, in CCSM3 given:

a) changes in storm trackb) decrease in connection to tropics, (i.e, tropical rainfall)

2) For the PT CCSM3 simulations specifically

a) 10x NPO "west" mode driven by *poleward* shift in storm tracks and a weak tropical connection

b) 1X NPO mode shifts from "east" to "west" due to a *poleward and eastward migration* of maximum EKE and baroclinic zones

c) 1X shift in storm tracks primarily forced by increased temperature gradients. Increased t-grads a response from expanded ice edge into mid-latitudes, most notably across the entire Panthalassic Ocean.

THE END