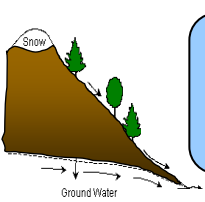


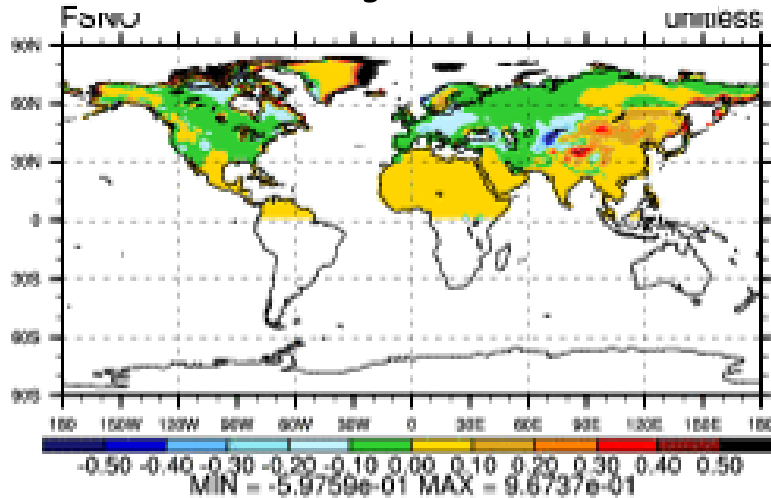
Arctic relevant LMWG changes for CLM4

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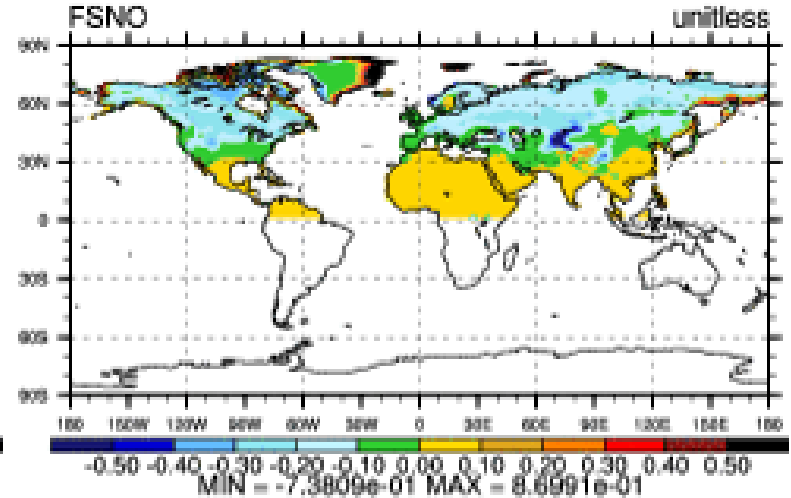


Results from Community Snow Project: Snow Cover Fraction (ANN)

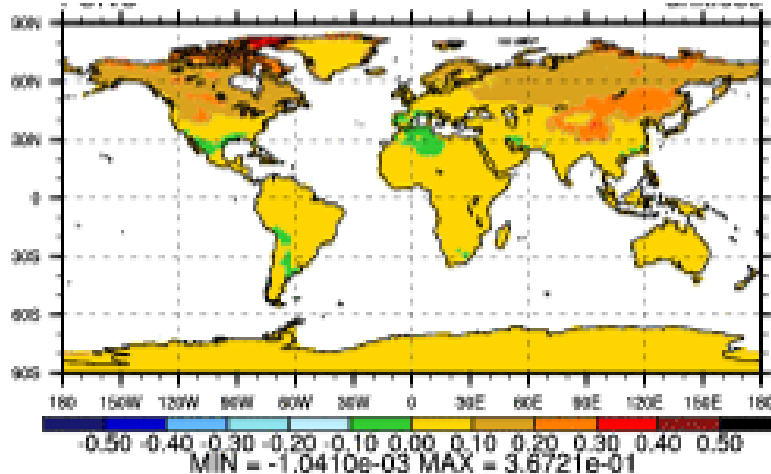
Community Snow - Obs



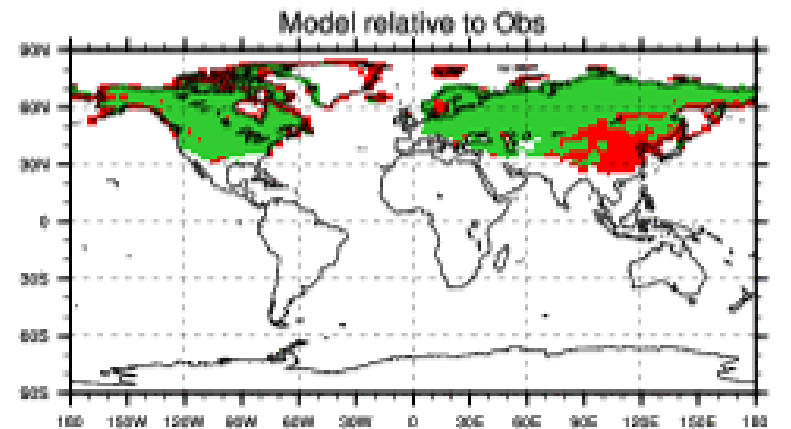
Control - Obs

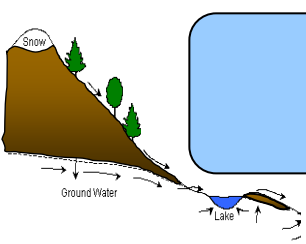


Community Snow - Control



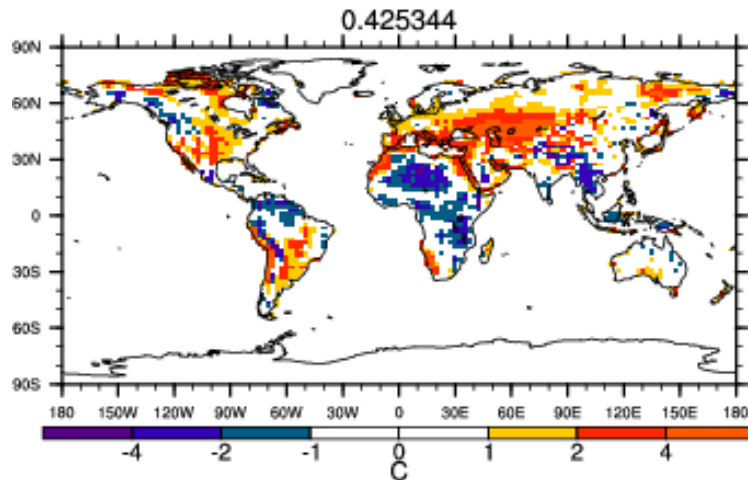
Case1+ (green) and Case2+ (red) relative to obs



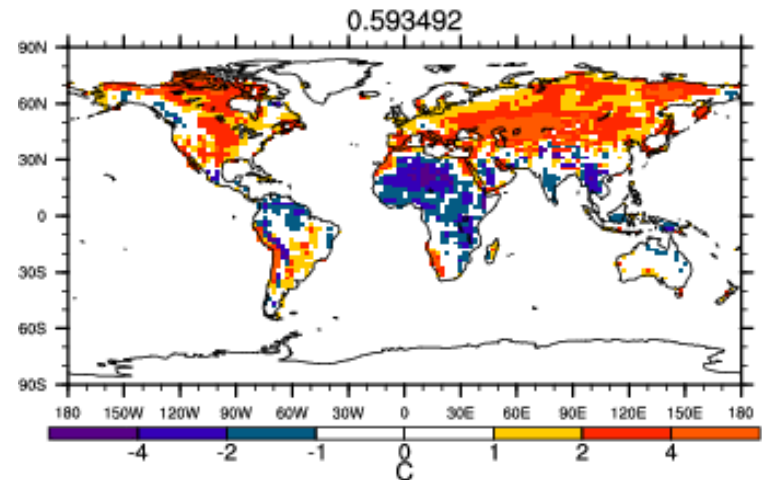


Results from Community Snow Project: Surface air temperature (ANN)

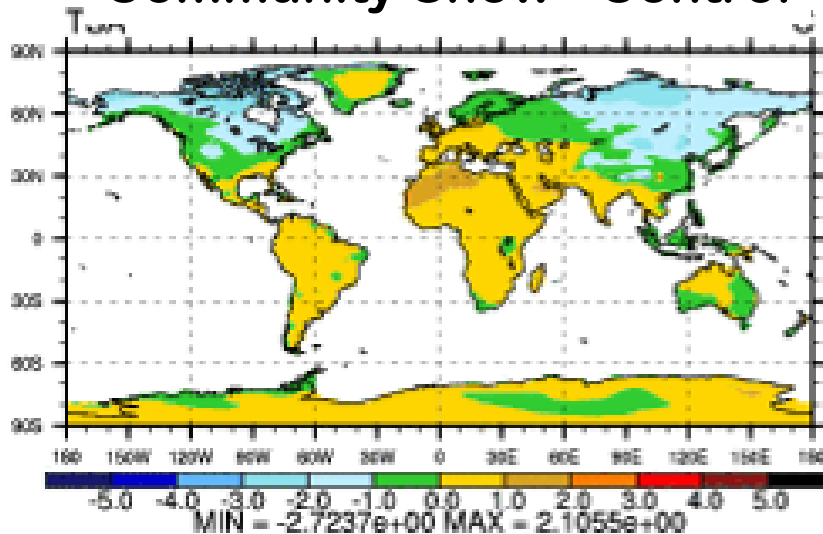
Community Snow - Obs



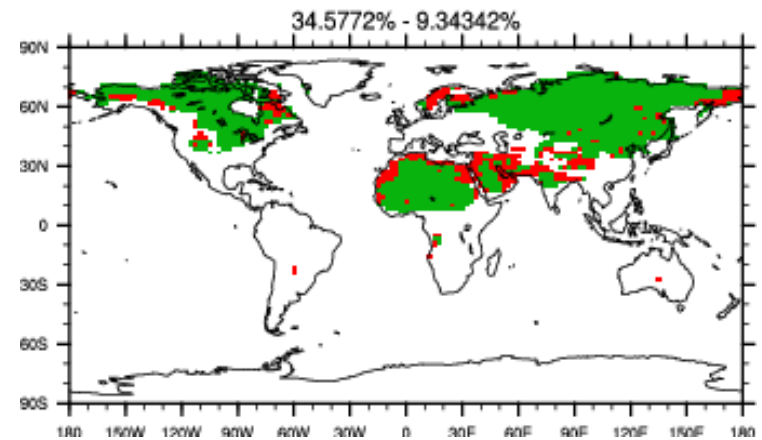
Control - Obs

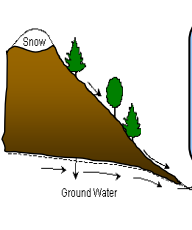


Community Snow - Control

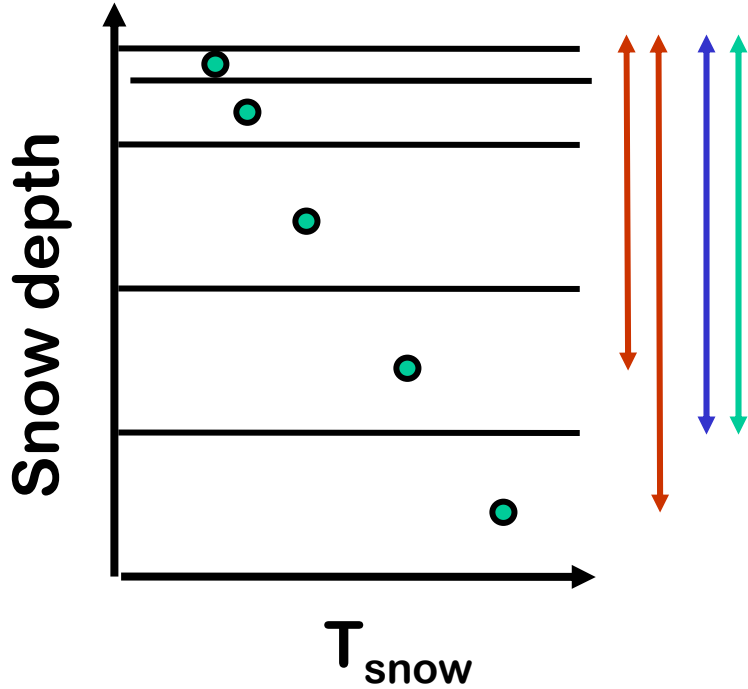
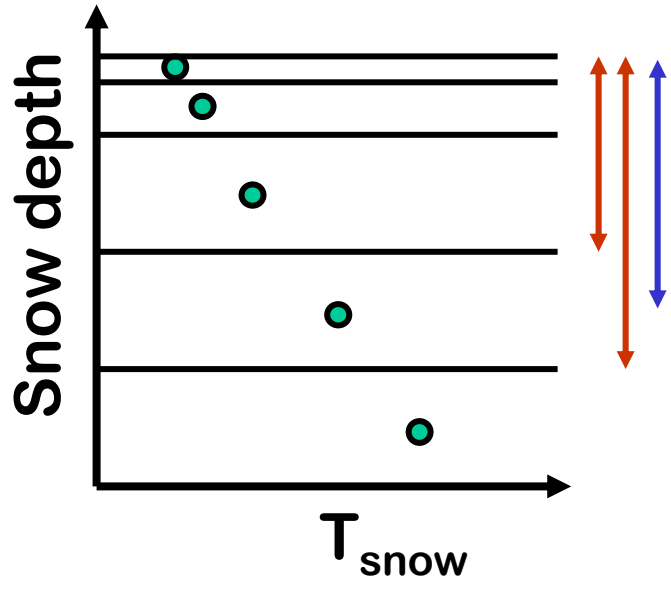
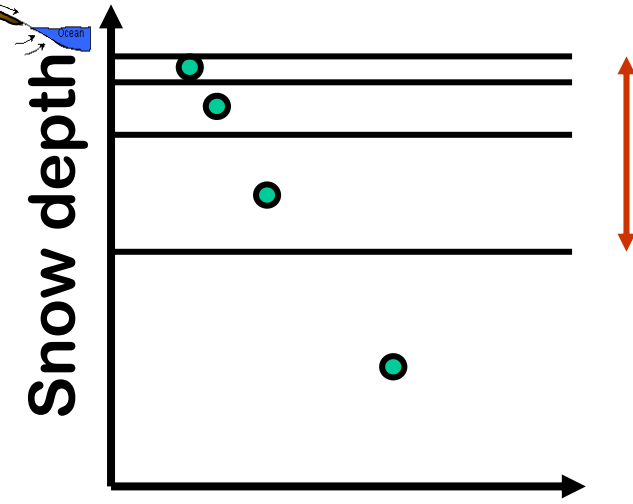


Model relative to Obs

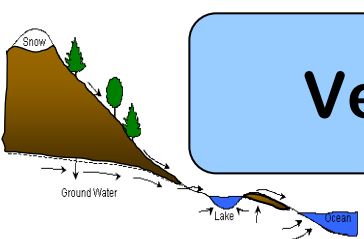




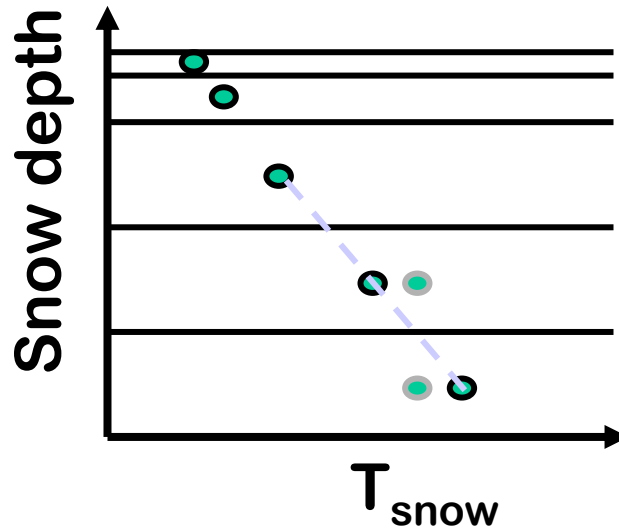
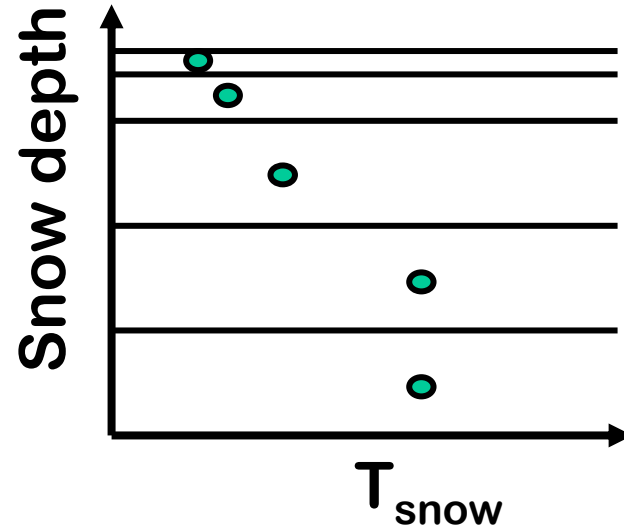
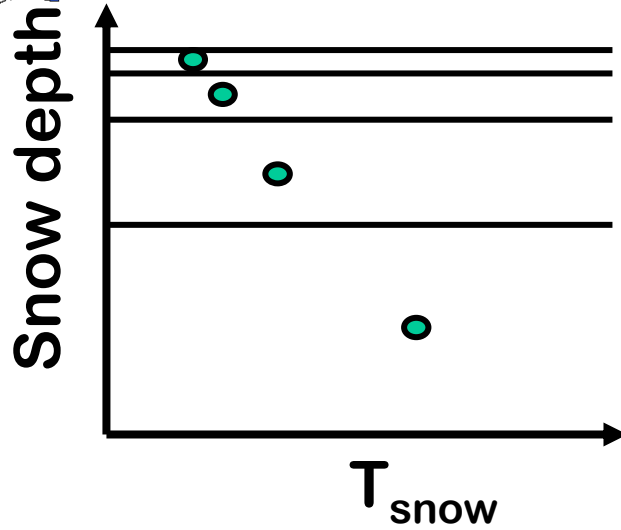
Snow compaction

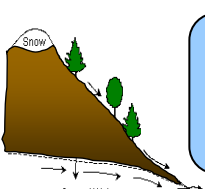


Impact: 10-20% shallower snow pack



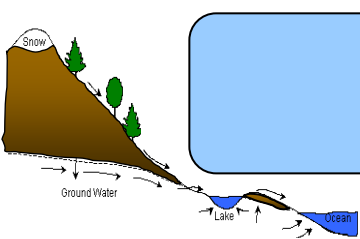
Vertical temperature profile after layer splitting





Arctic relevant LMWG changes for CLM4

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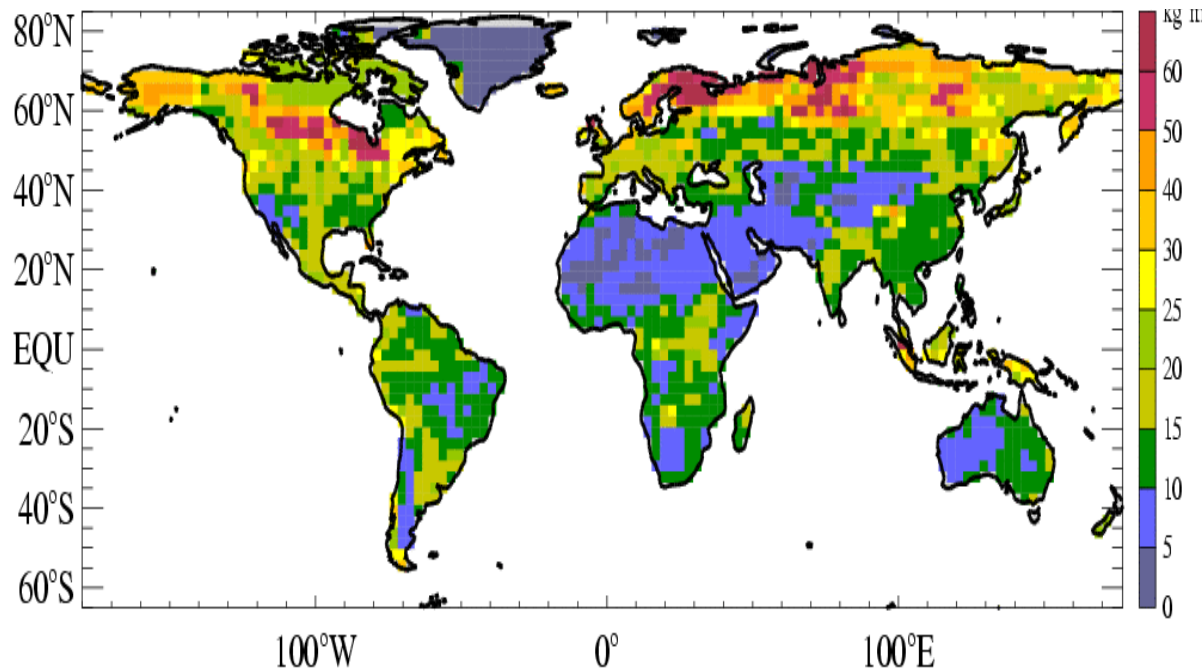


CLM soil carbon density dataset

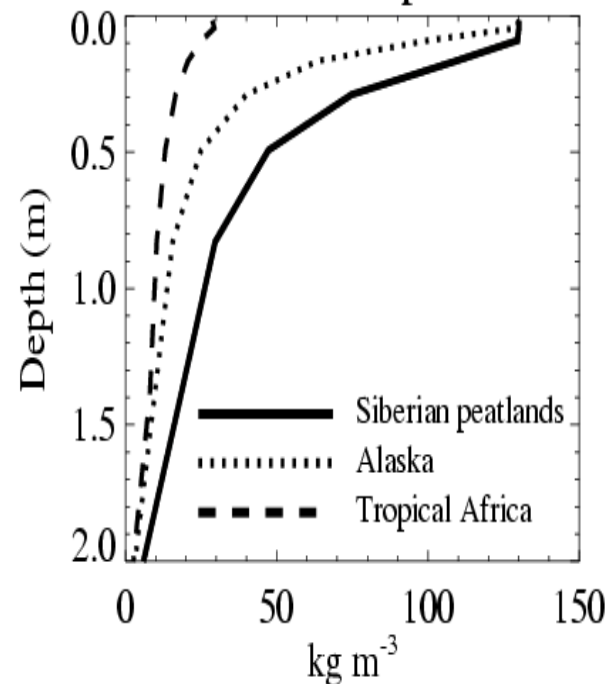
Source data from Global Soil Data Task

Soil carbon content

Kg m^{-2}

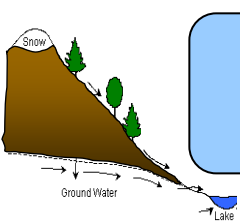


Soil carbon profile

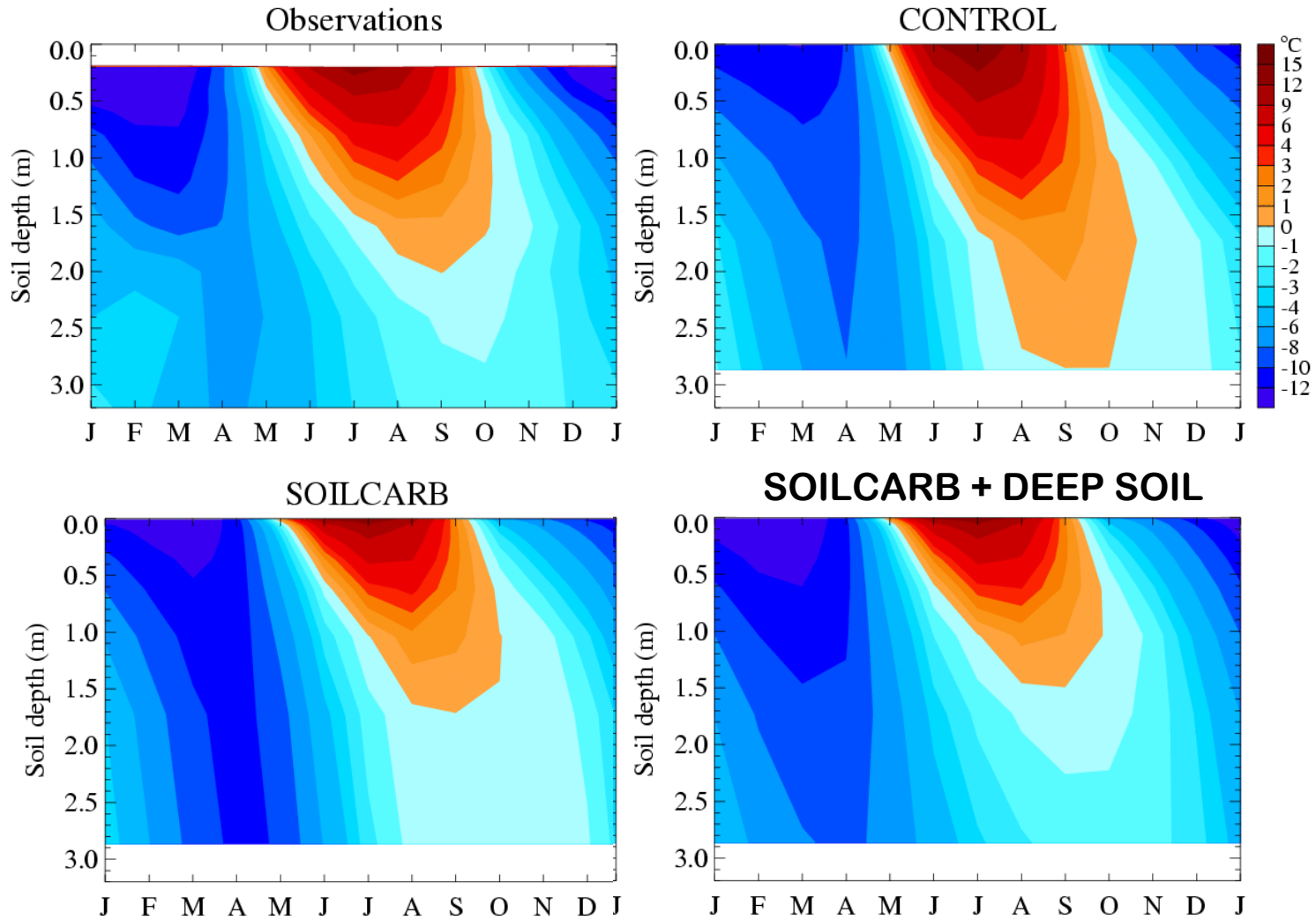


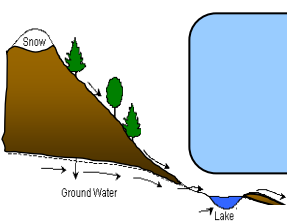
$f_{sc,i} = \rho_{sc,i} / \rho_{peat}$ fraction of layer i that is organic matter

$i = 1, 2, 3, \dots, 10$ soil layers



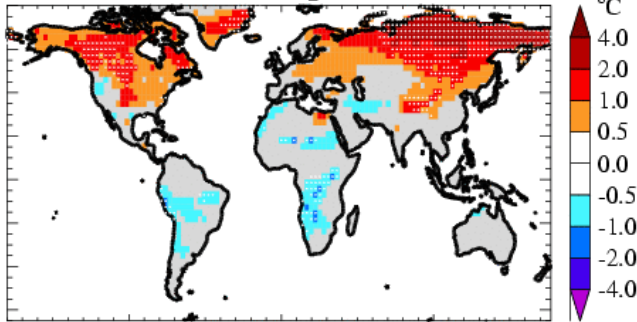
Annual cycle-depth soil temperature plots Siberia



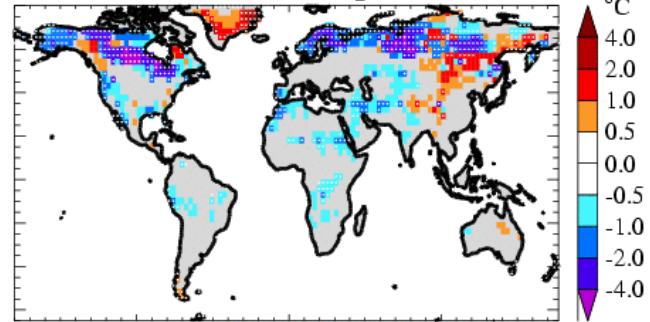


SOILCARB – CONTROL (JJA)

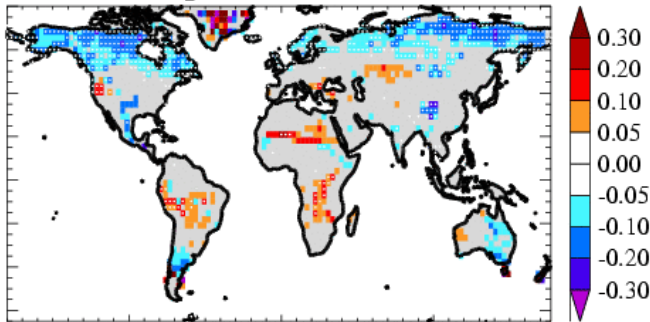
2m Air Temperature



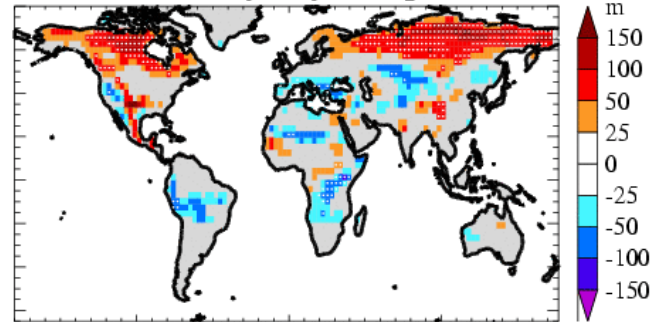
Column Soil Temperature



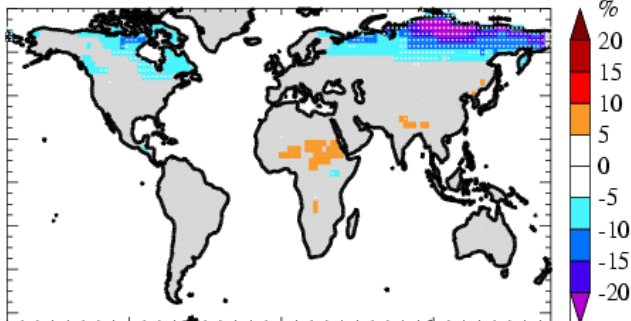
Evaporative Fraction



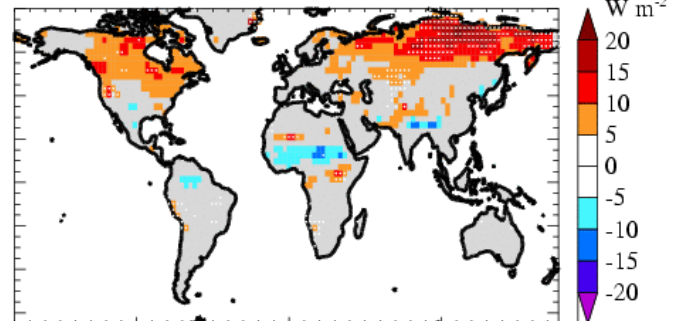
Boundary Layer Depth

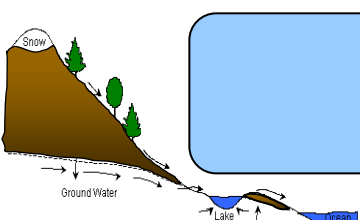


Low Cloud Fraction



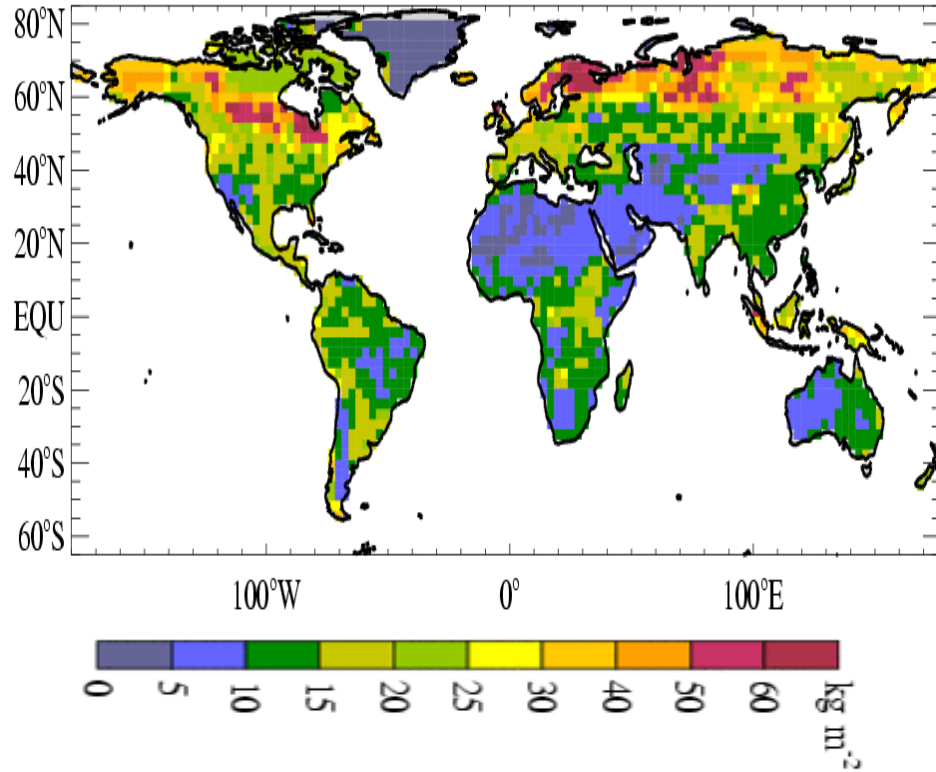
Net Surface Radiation



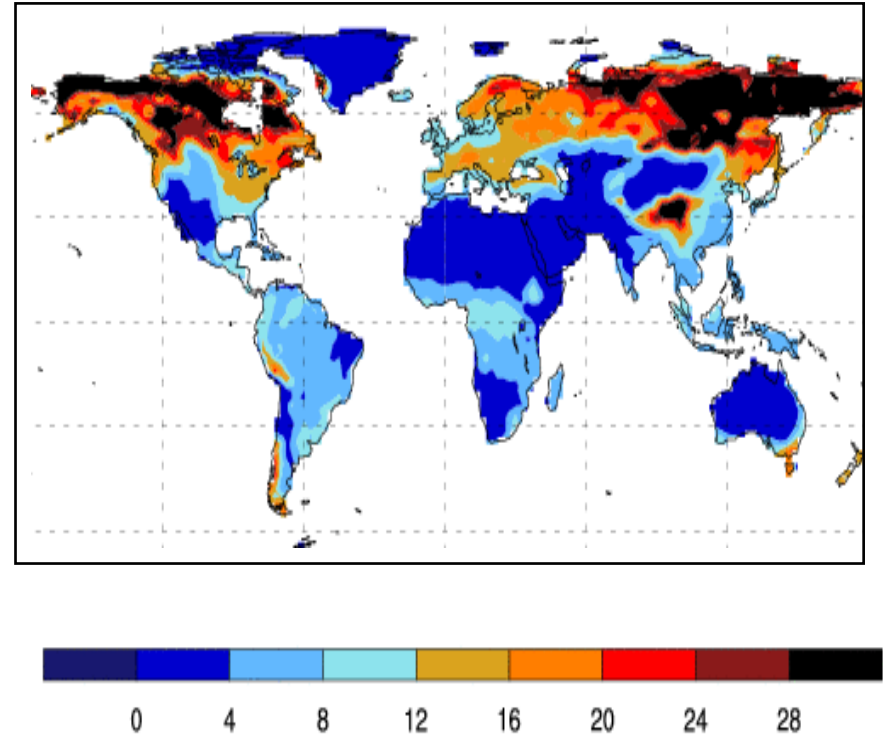


Soil carbon in CLM-CN

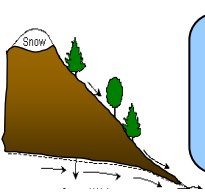
Soil carbon content: Obs



Soil carbon content: CLM-CN



Goal is to couple CLM-CN soil carbon with organic soil parameterization



Arctic relevant LMWG changes for CLM4

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Accelerated Arctic land warming and permafrost degradation during rapid sea ice loss



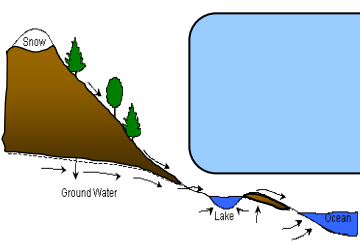
David Lawrence¹, A.G. Slater²,
B. Tomas¹, M. Holland¹, and C. Deser¹



¹ NCAR / CGD
Boulder, CO



² NSIDC / CIRES
Boulder, CO



2007 – A record year in the Arctic

2007 minimum sea ice extent (Sept)



Image courtesy NSIDC

Western Arctic (Aug to Oct 2007)

warmest on record

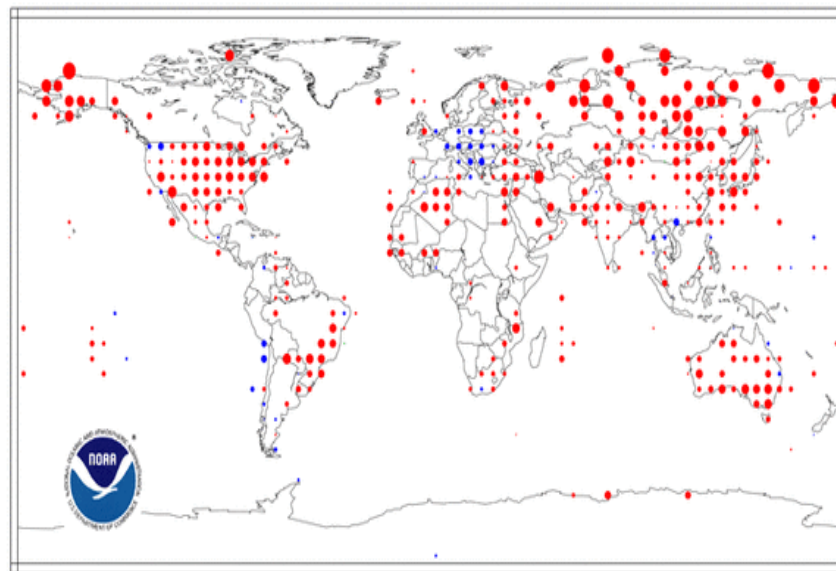
+2.3°C over 1978-2006 mean

(preliminary CRU data)

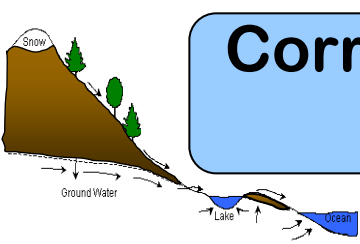
Temperature Anomalies Sep-Nov 2007

(with respect to a 1961-1990 base period)

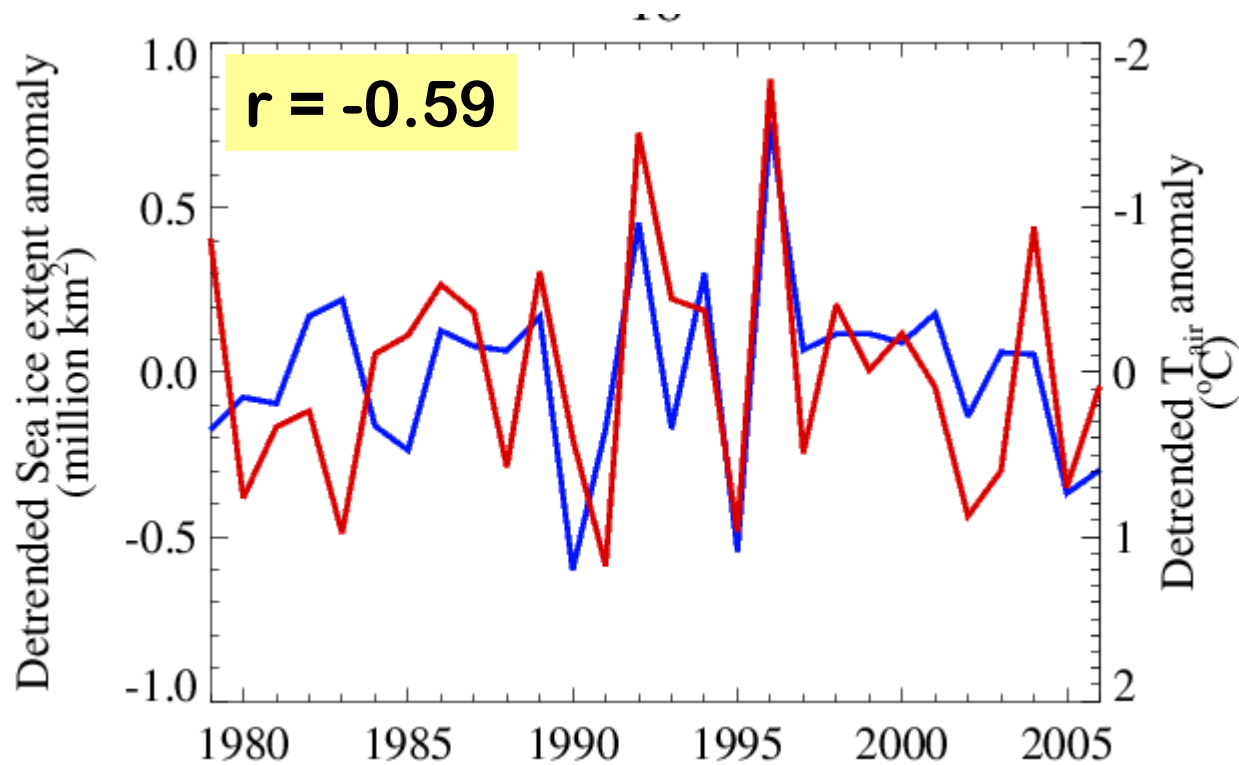
National Climatic Data Center/NESDIS/NOAA

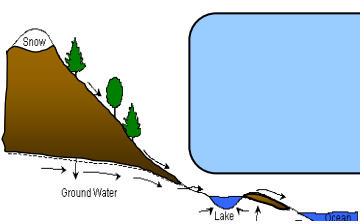


Degrees Celsius



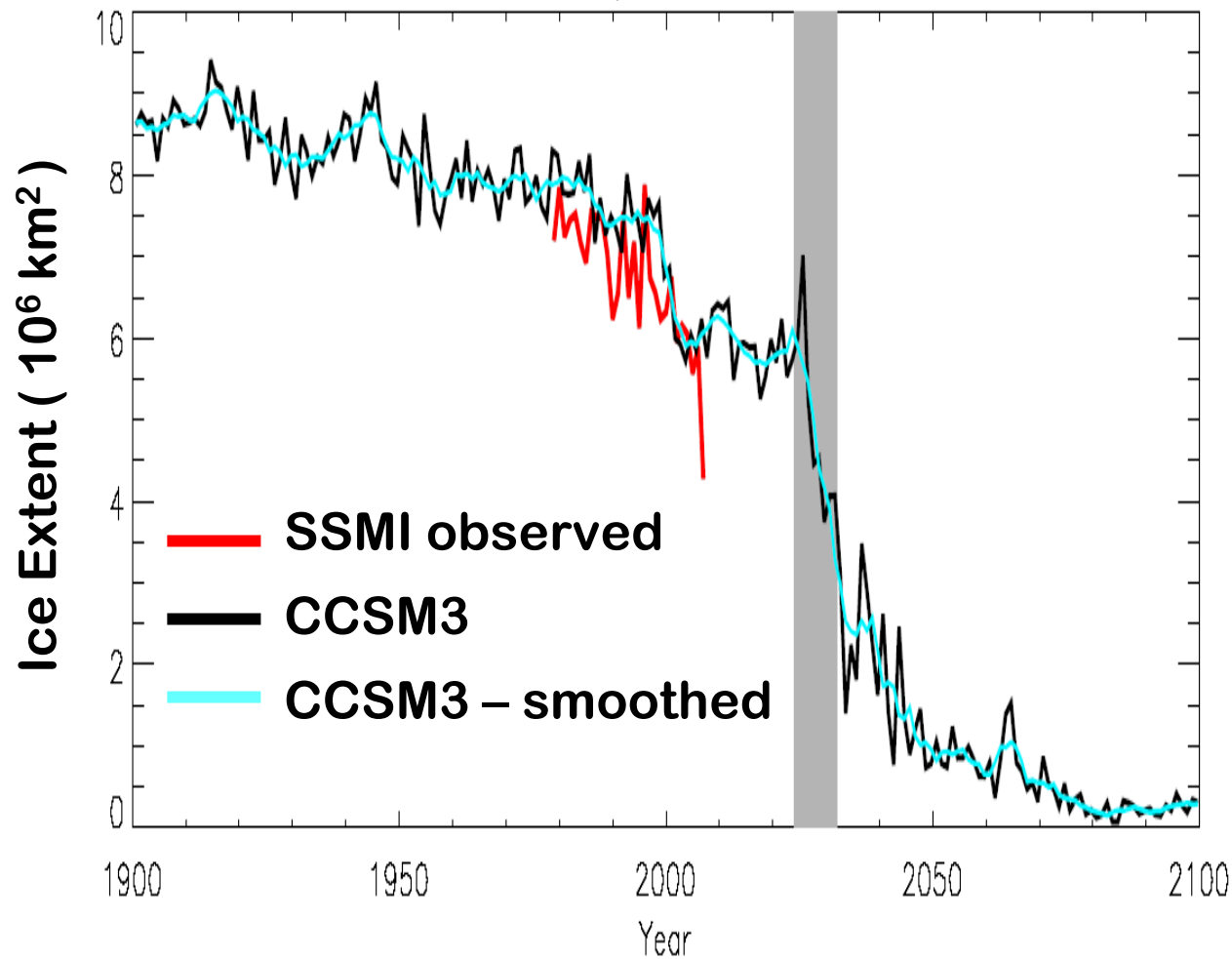
Correlation between Sept sea ice extent and western Arctic July to October T_{air} (CRU)





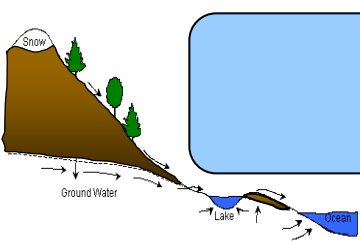
Observed and modeled sea ice extent

September sea ice extent

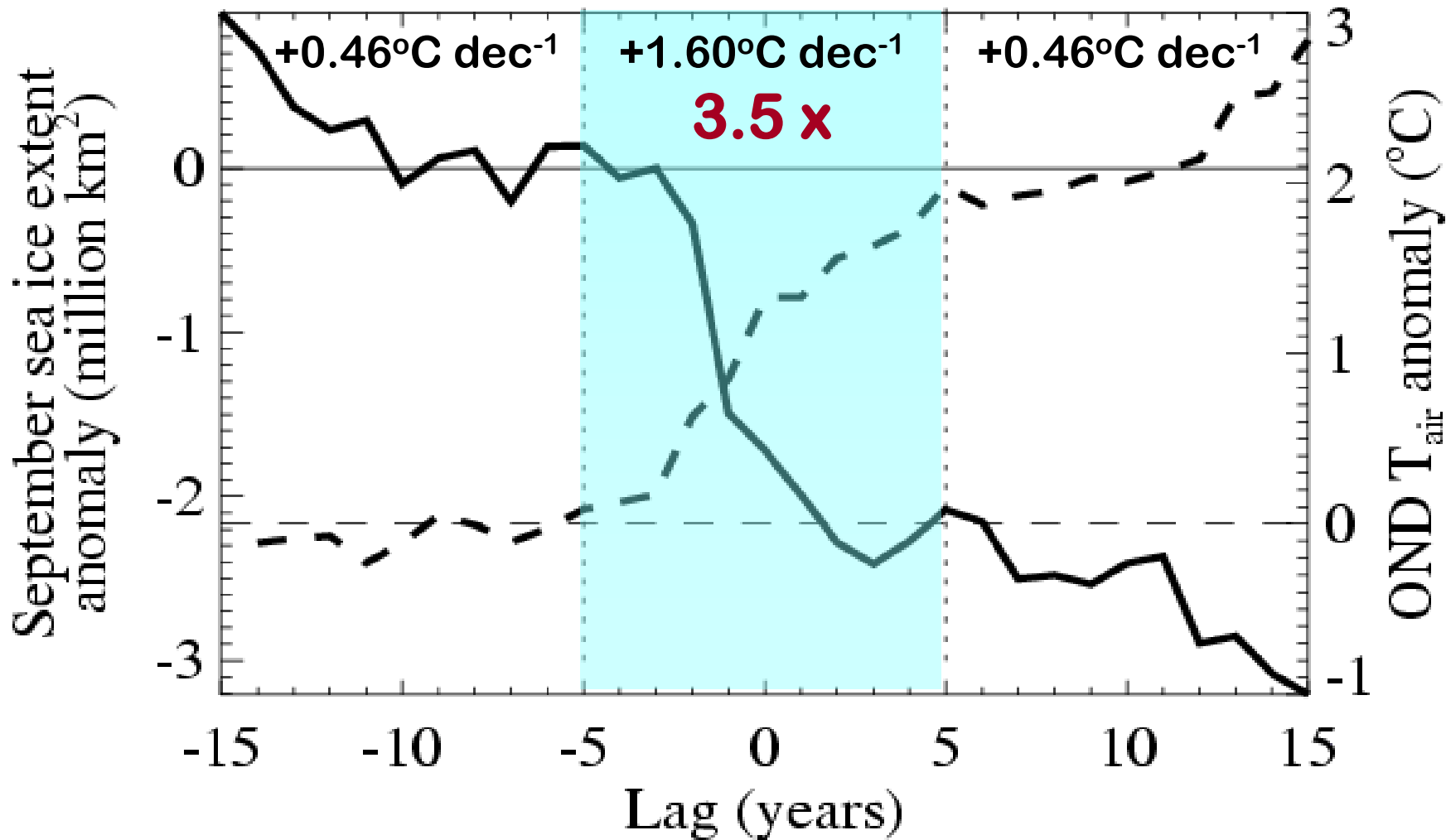


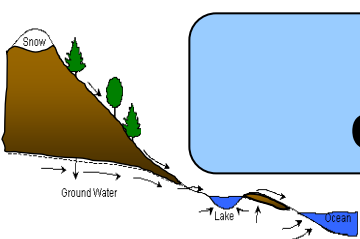
9 rapid sea ice loss events across 8-member CCSM3 A1B ensemble ranging from 2012 to 2045

(Holland et al. 2006)

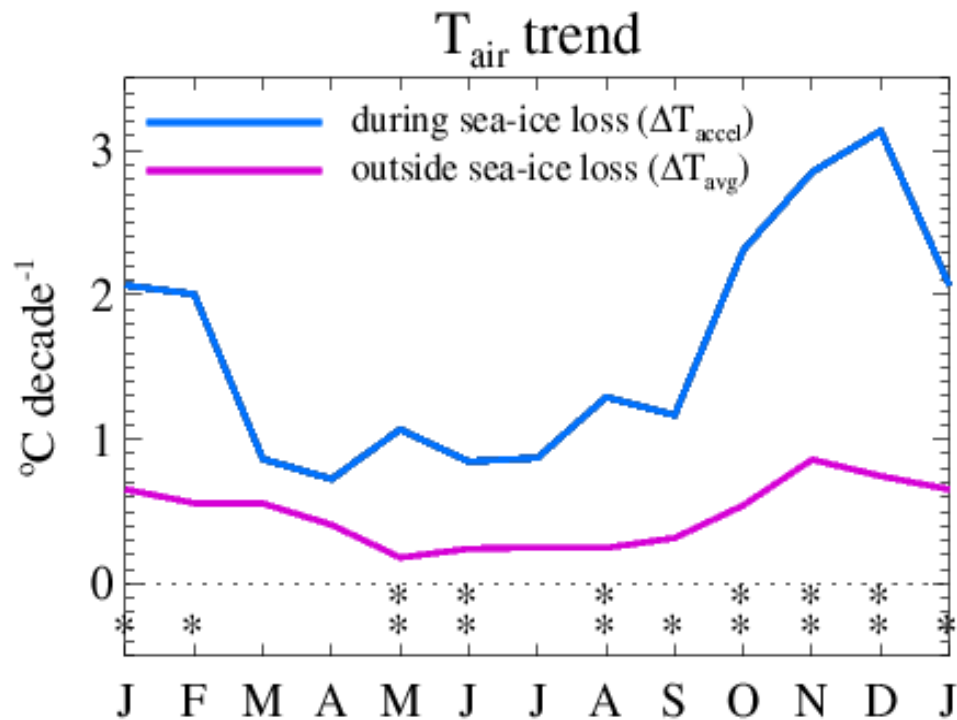
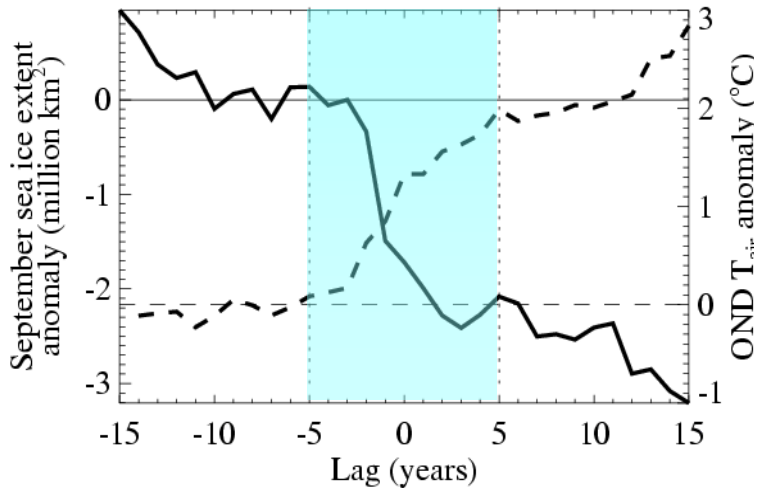


Lagged composite: sea ice extent and western Arctic land T_{air}



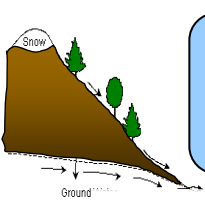


T_{air} trend during and outside rapid sea ice loss periods



* P < 0.05

** P < 0.01

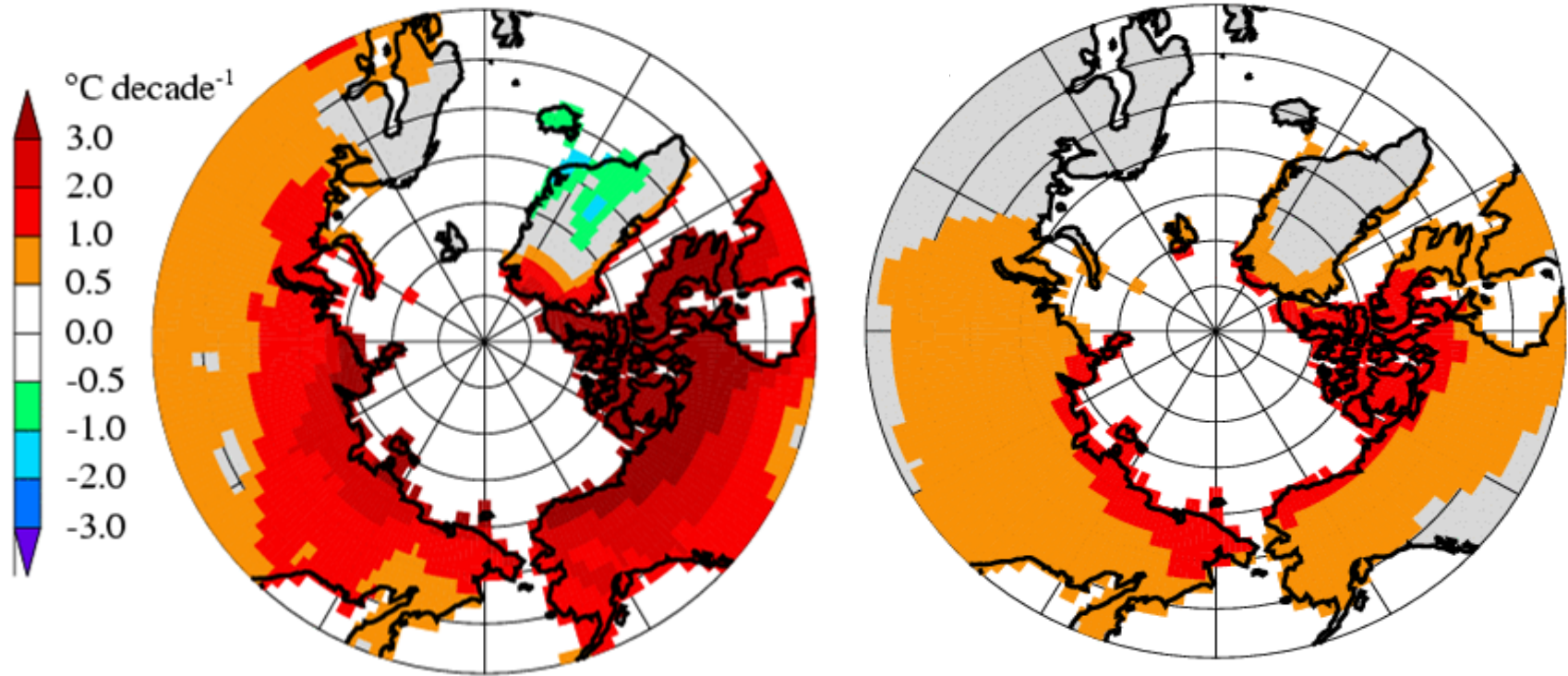


Trends during and outside rapid sea ice loss periods

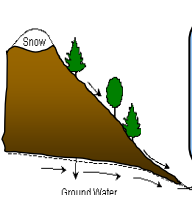
T_{air} trend:

during sea-ice loss periods

outside sea-ice loss periods



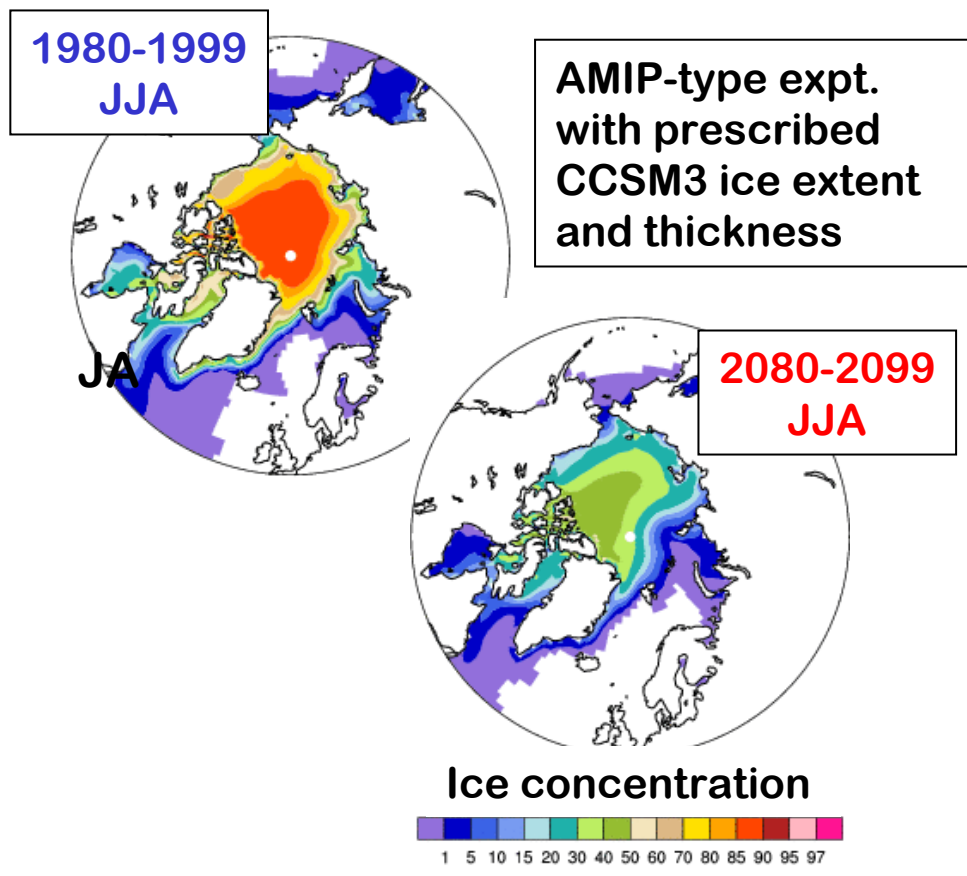
- Similar trends in q_{air} , LW_{\downarrow}
- No statistically significant or spatially coherent trends in P , snow depth, or SW_{\downarrow} , P



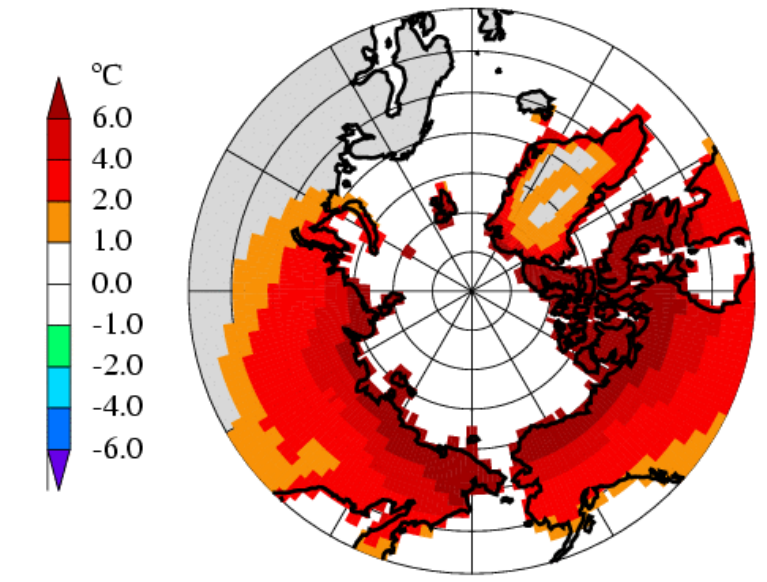
Is warming a response to or a forcing of sea ice loss?

Impact of sea ice on climate experiments

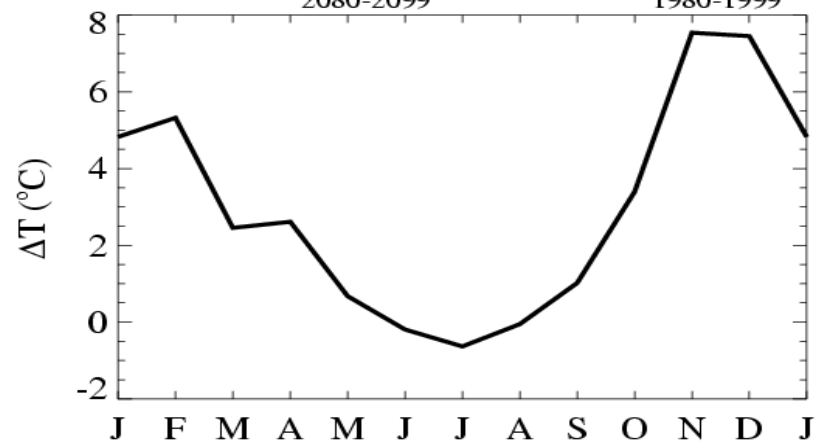
[Climate Analysis Working Group](#)

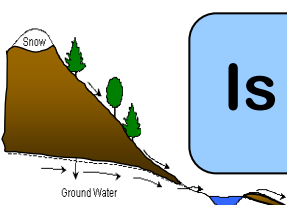


ΔT (OND)
SEAICE₂₀₈₀₋₂₀₉₉ - SEAICE₁₉₈₀₋₁₉₉₉



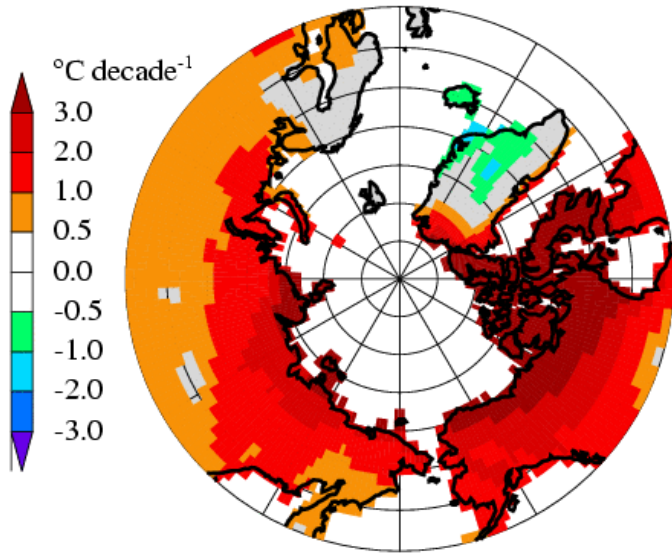
SEAICE₂₀₈₀₋₂₀₉₉ - SEAICE₁₉₈₀₋₁₉₉₉



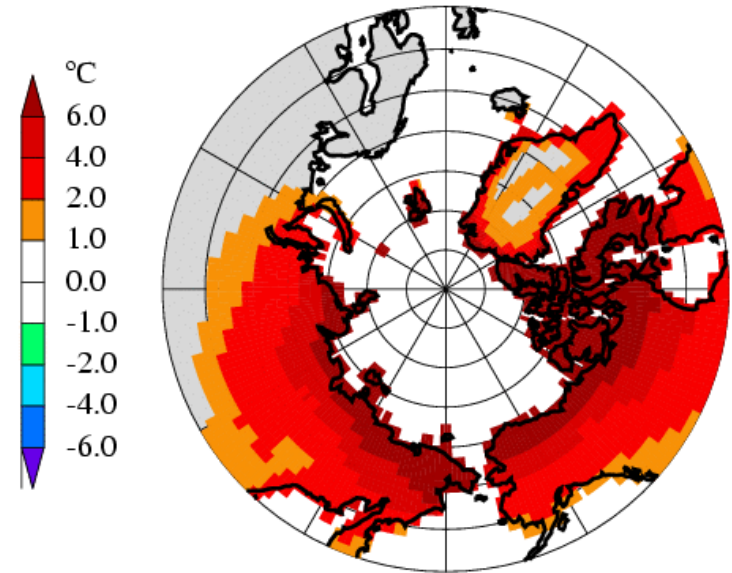


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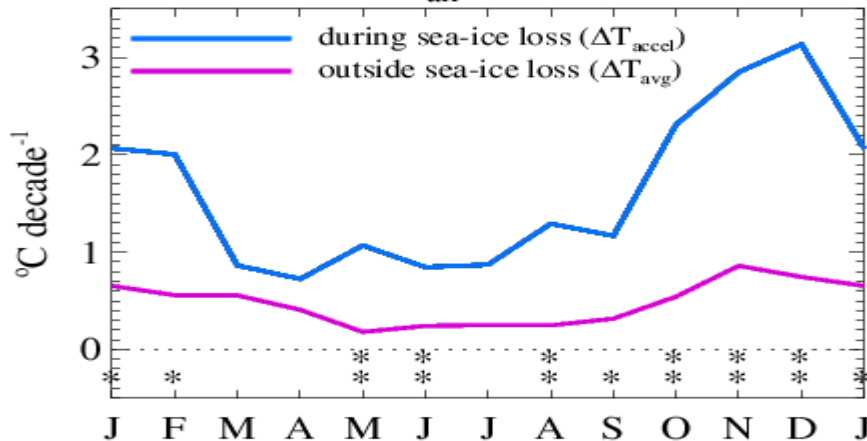
ΔT_{air} (OND)
during sea-ice loss periods



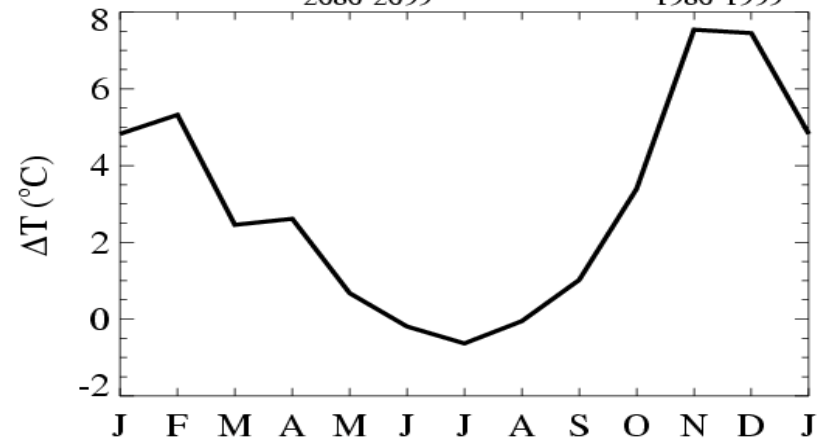
ΔT (OND)
 $\text{SEAICE}_{2080-2099} - \text{SEAICE}_{1980-1999}$

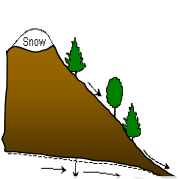


T_{air} trend



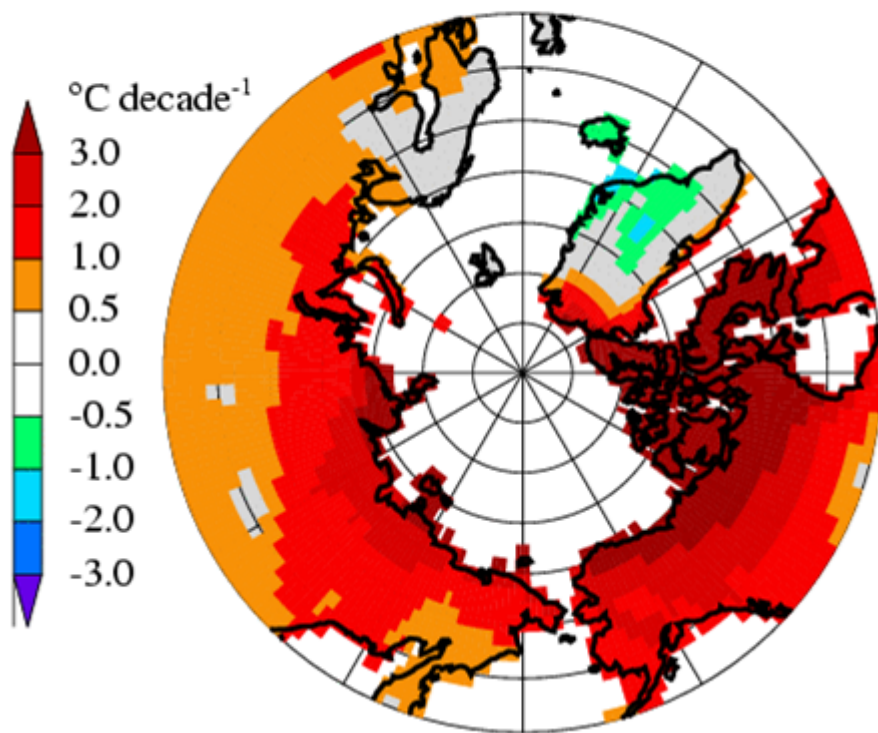
$\text{SEAICE}_{2080-2099} - \text{SEAICE}_{1980-1999}$



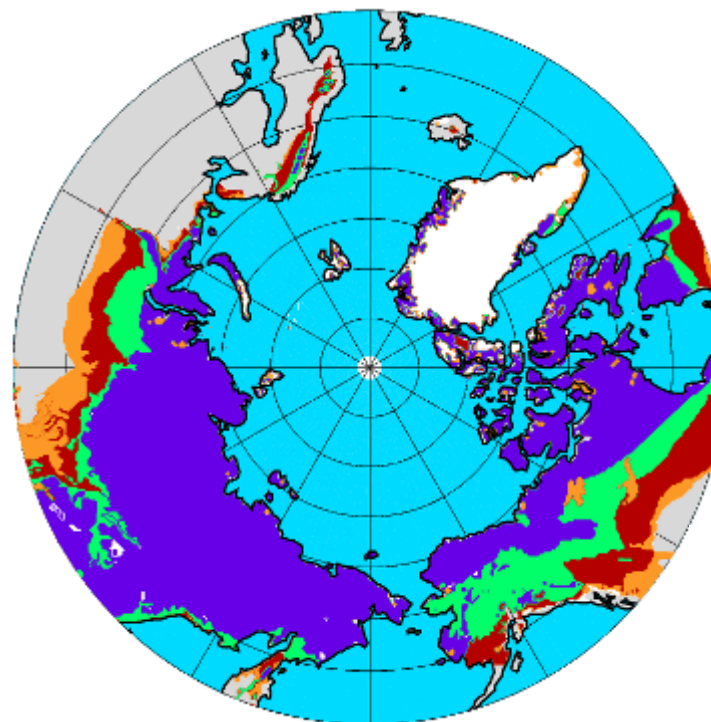


Collocation of accelerated warming with permafrost zones

T_{air} trend
during sea-ice loss periods

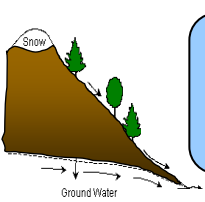


IPA Permafrost
Distribution Map

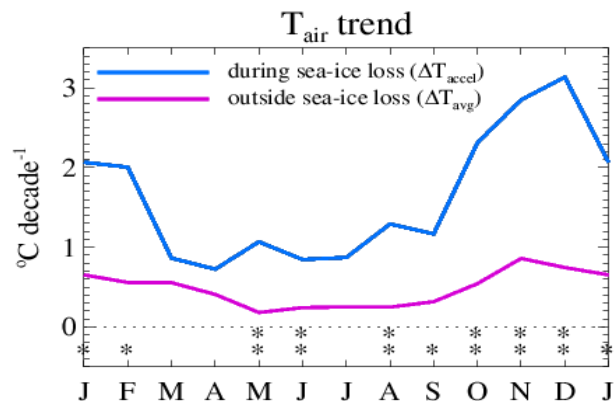
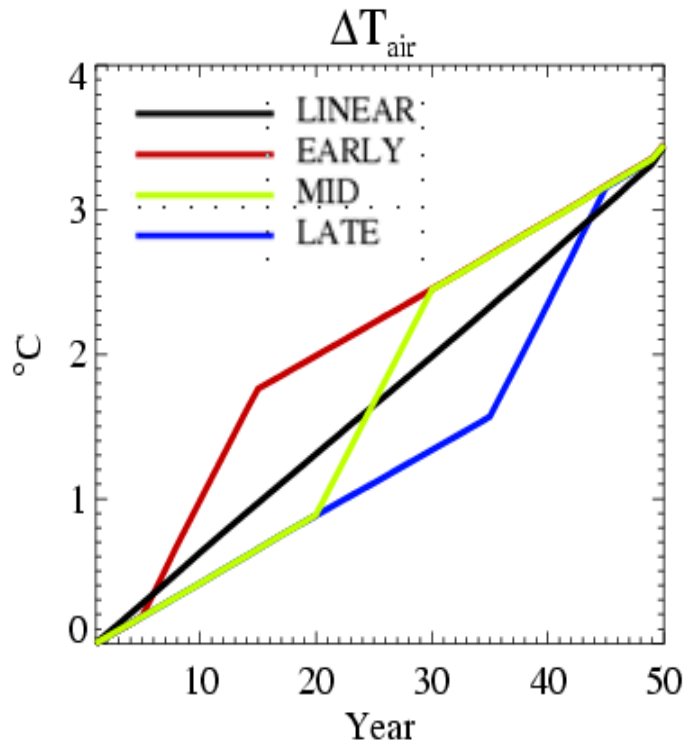


Question: How does rapid sea ice loss and associated accelerated land warming affect permafrost?

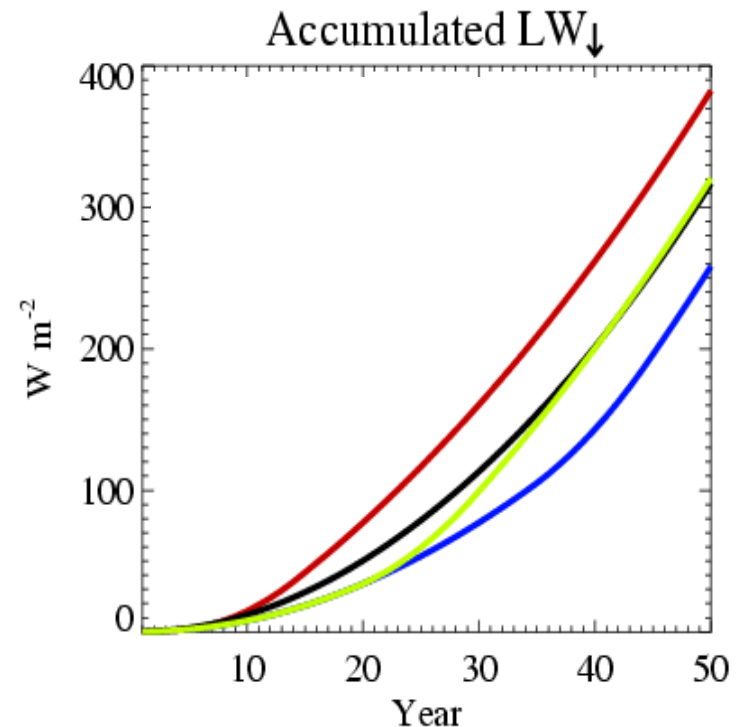
- Continuous (90-100% cover)
- Discontinuous (50-90%)
- Sporadic (10-50%)
- Isolated (0-10%)

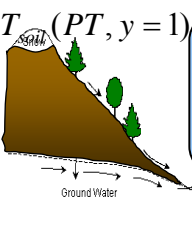


Idealized CLM forcing scenarios



- Generate four meteorological forcing data sets by adding ΔT_{air} anomaly time series to observed year 2000 forcing.
- q_{air} also adjusted to conserve RH.

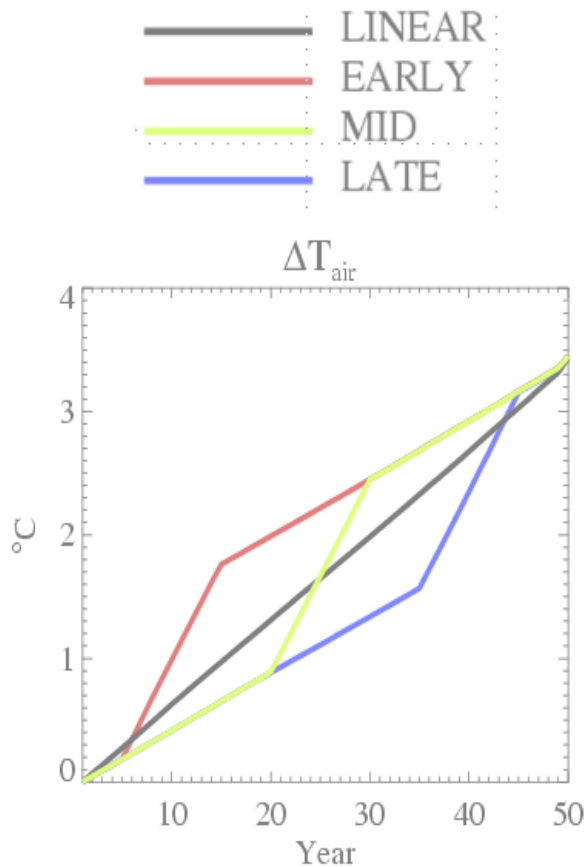


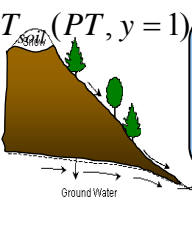


Impact of accelerated warming on permafrost

Experimental design

- Use CLM3.5 with improved permafrost dynamics
- Spin up for 400 years with repeat year 2000 forcing
- Force CLM with idealized scenario experiments
- Consider only points where $\Delta\text{snow_depth} < 10\%$





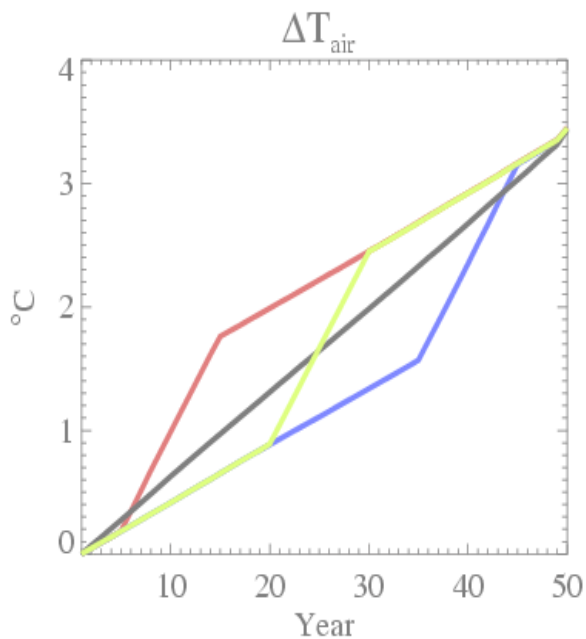
Impact of accelerated warming on permafrost



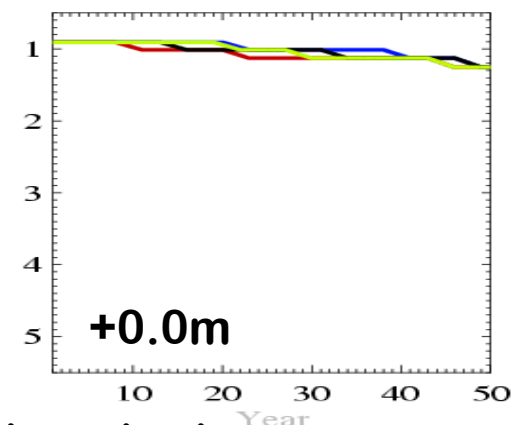
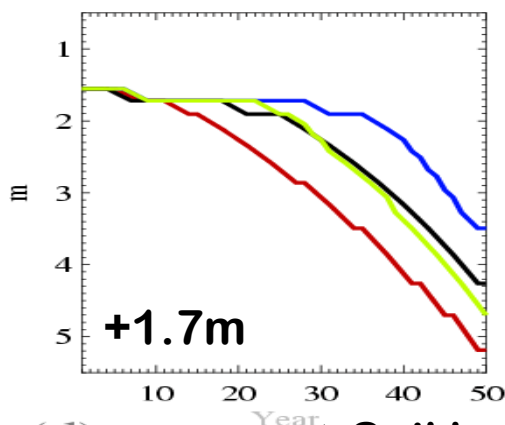
Warm
-0.3°C

$$\overline{T_{soil}(PT, y = 1)}$$

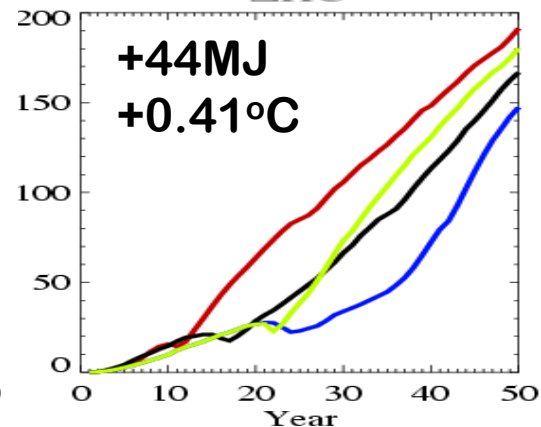
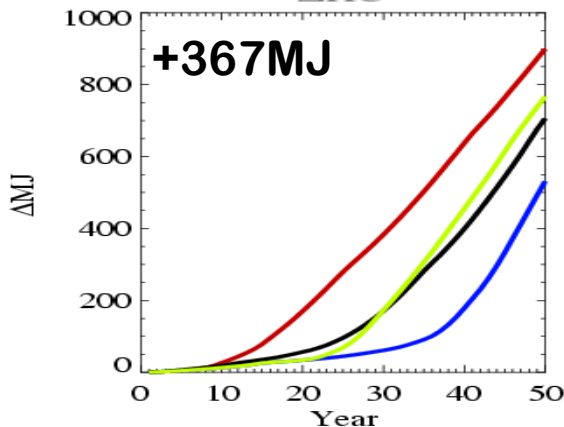
Cold
-5.8°C



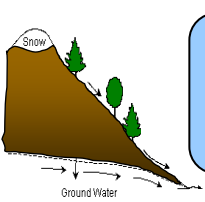
(c) Depth to permafrost table



(d) Δ Soil heat content



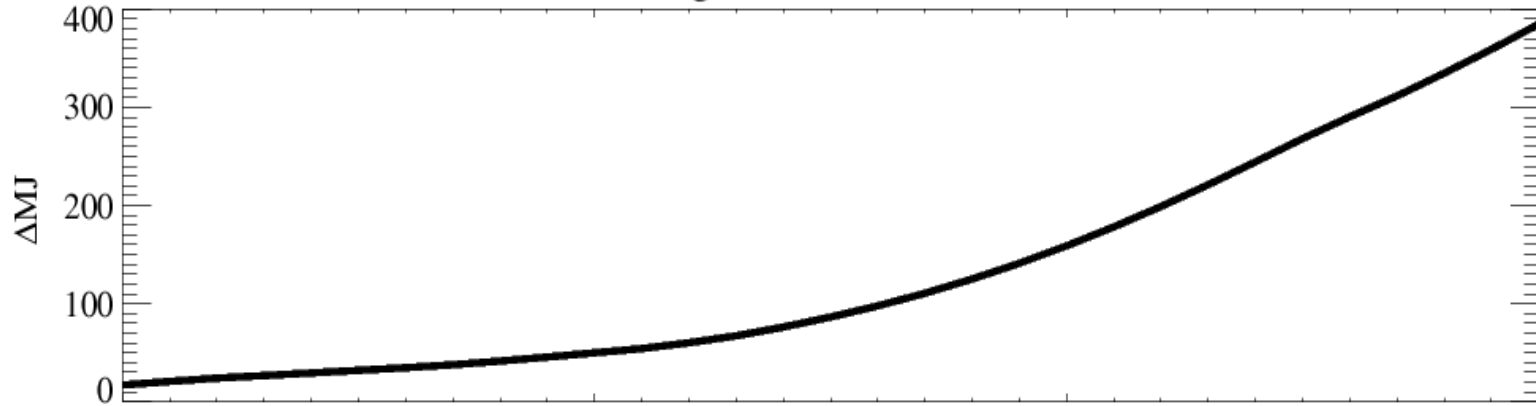
Beltrami et al. (2002)
estimate avg land heat gain
1950 to 2000 of 70MJ



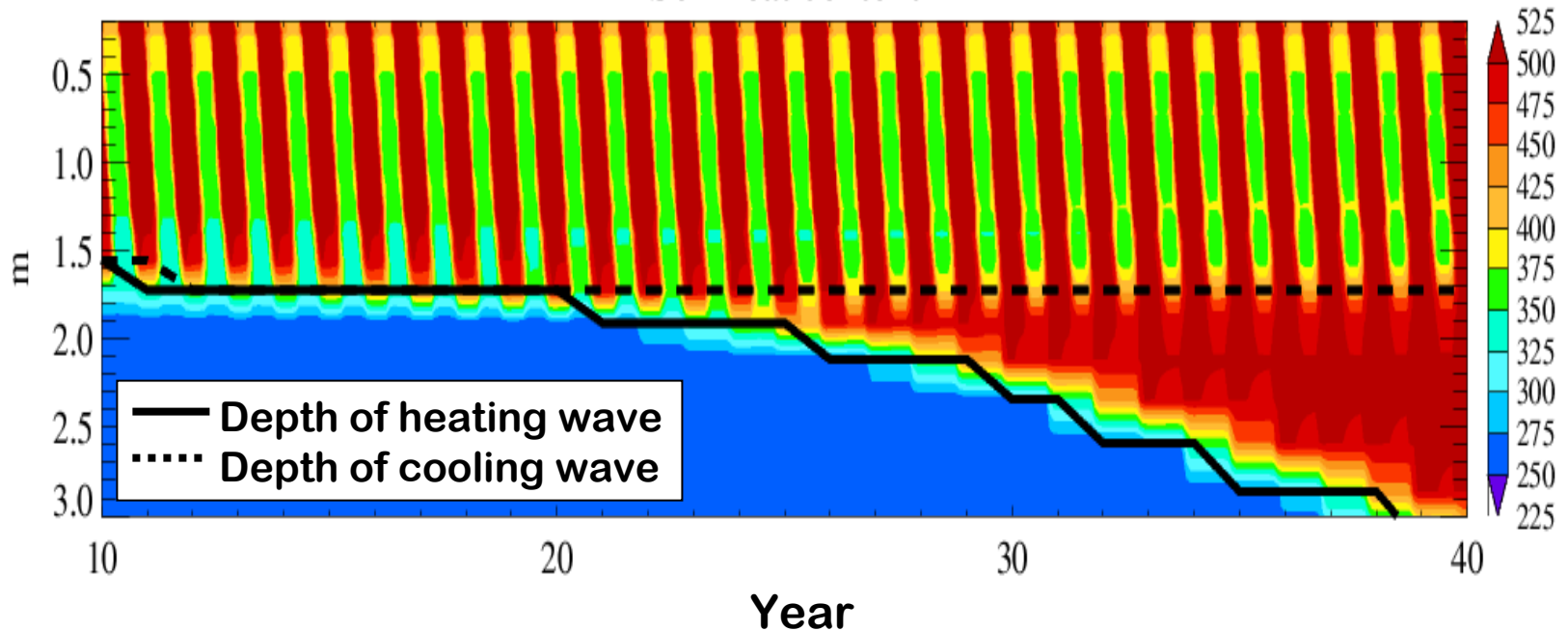
Soil heat accumulation

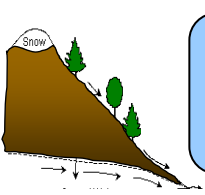
Warm permafrost case, LINEAR expt

Change in soil heat content



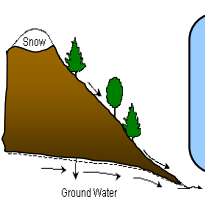
Soil heat content



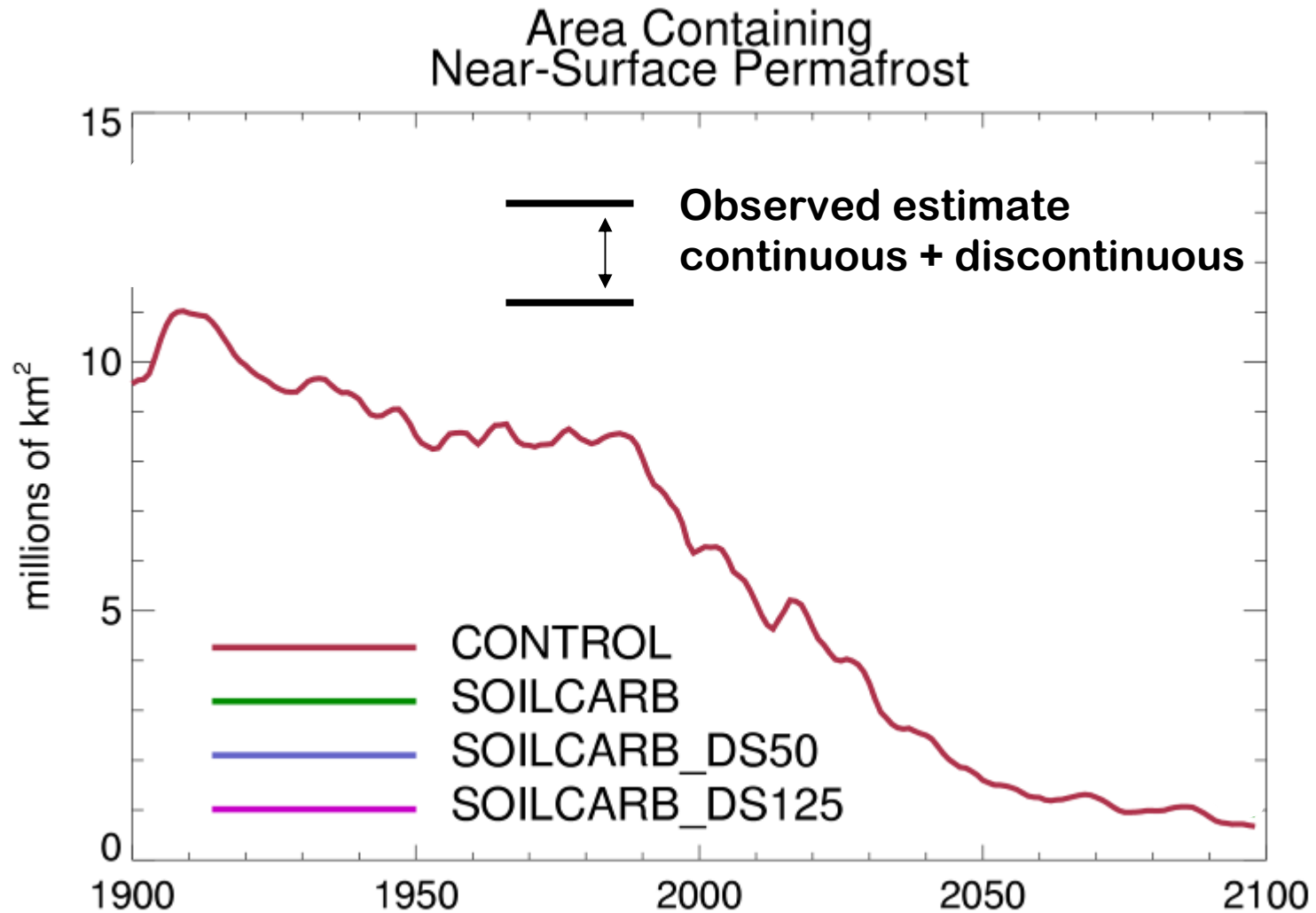


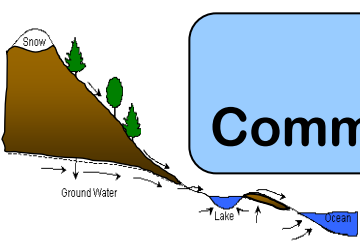
Summary

- Periods of **rapid sea ice loss** in CCSM3 induce a **3.5x increase** in warming trends compared to secular climate-change trends for western Arctic. This is consistent with 2007 events.
- Accelerated warming signal **extends up to 1500km inland** and is apparent throughout most of the year, **peaking in autumn**.
- Accelerated warming substantially increases ground heat accumulation – **the earlier the event the greater the long-term impact**.
- Enhanced heat accumulation can lead to **rapid degradation** of warm permafrost and preconditions colder permafrost for **earlier and/or more rapid degradation**.
- Accelerated warming is likely to have broader impacts on vulnerable Arctic **ecosystems, biogeochemical cycling, infrastructure**.



Near-surface permafrost degradation

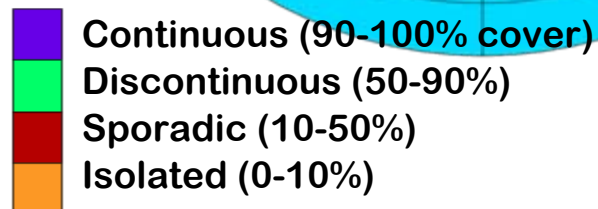
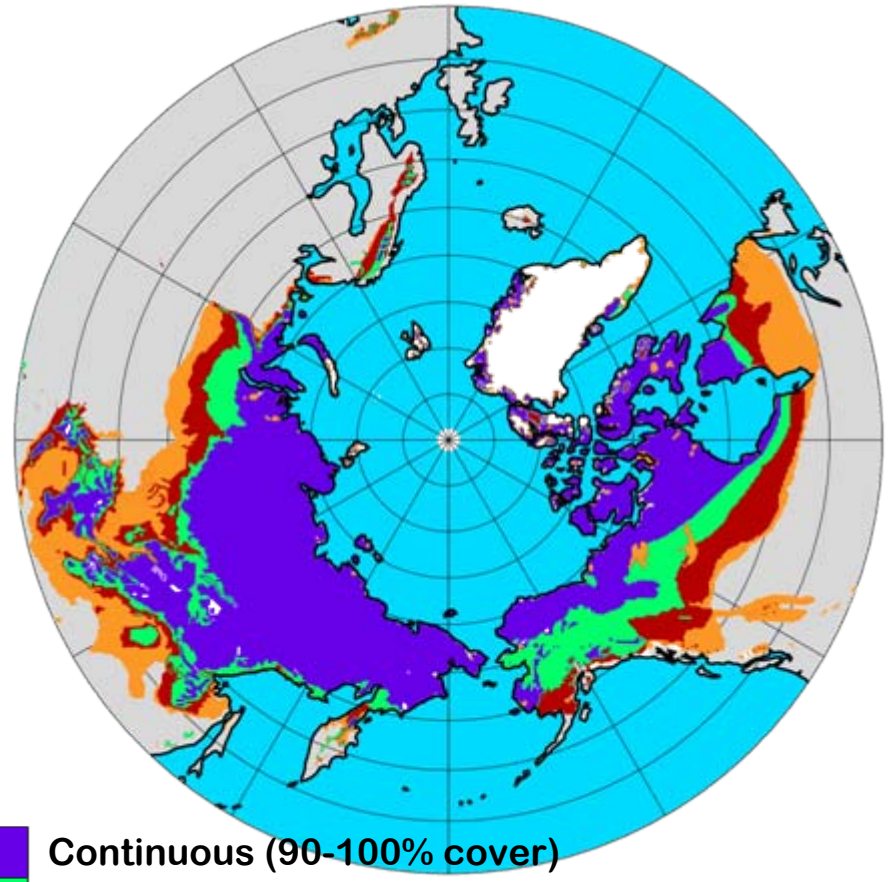
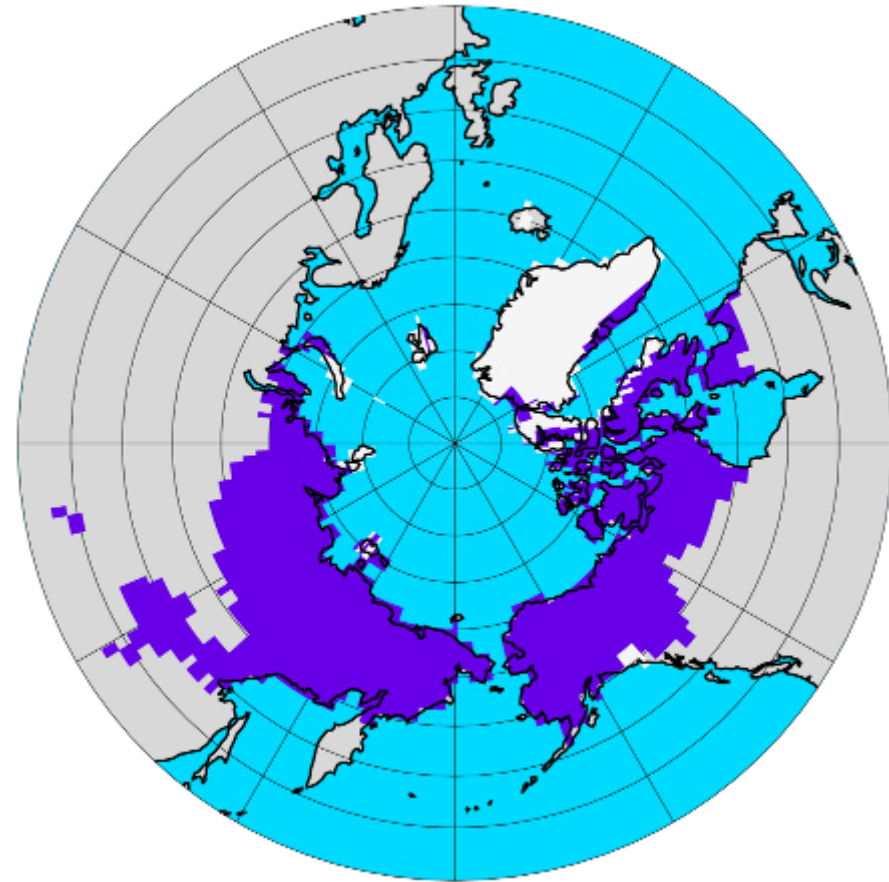


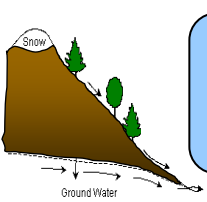


Near-surface permafrost distribution in Community Climate System Model (CCSM3) IPCC AR4 Ensemble

**CCSM3
(1980 – 1999)**

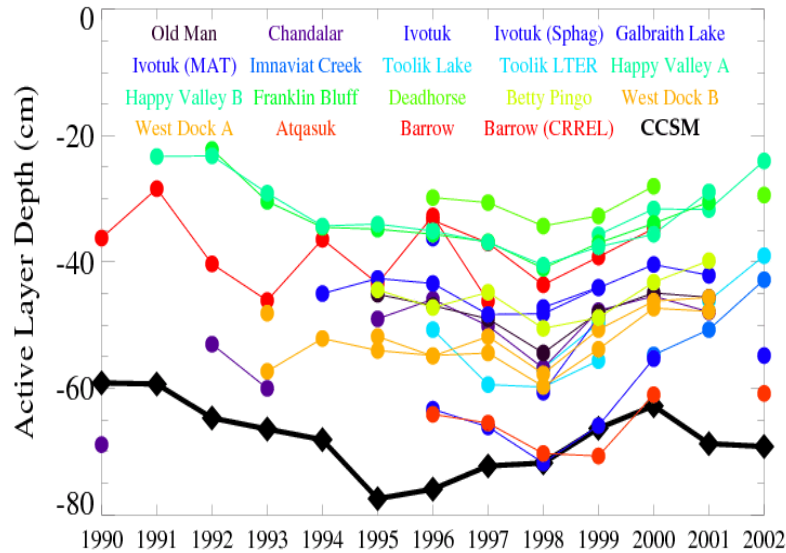
**IPA Permafrost
Distribution Map**



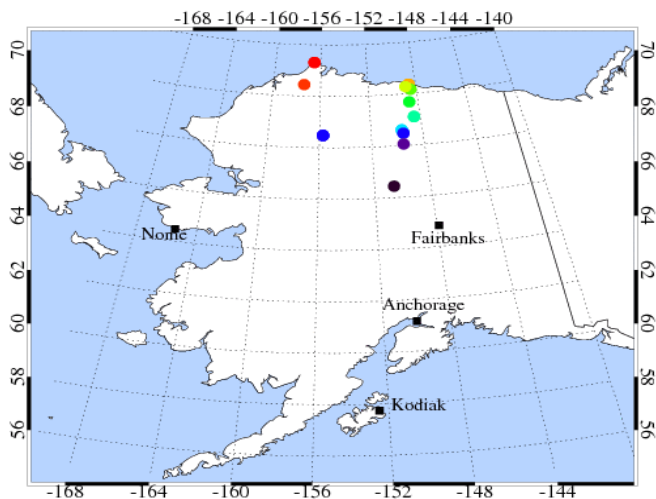


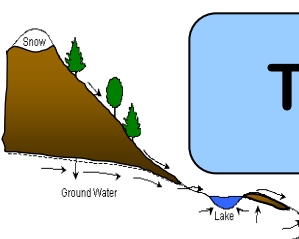
Potential sources of bias in the CCSM3 permafrost simulation

CALM Monitoring Sites



- Biases in the simulated climate
 - T_{air} , snow, variability
- Errors due to common simplifications in GCM land-surface schemes
 - soils represented as pure mineral soils (sand, silt, clay) no organic soil
 - shallow soil column (~3m), small number of soil levels
 - no excess soil ice
 - no vertical heat advection by water
 - lack of subgrid scale features, thermokarst, aspect, microtopography





Thermal and hydraulic parameters for organic soil

Soil type	λ_{sat}	λ_{dry}	Θ_{sat}	k_{sat}
Sand	3.12	0.27	0.37	0.023
Clay	1.78	0.20	0.46	0.002
Peat	0.55	0.05^a	0.9^{a,b}	0.100^b

$$\Theta_{\text{sat},i} = (1 - f_{\text{sc},i}) (0.489 - 0.00126 \% \text{sand}_i) + f_{\text{sc},i} \Theta_{\text{sat},\text{sc}}$$

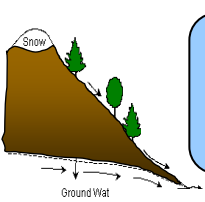
λ_{sat} sat. thermal conductivity

λ_{dry} dry thermal conductivity

Θ_{sat} volumetric water at saturation

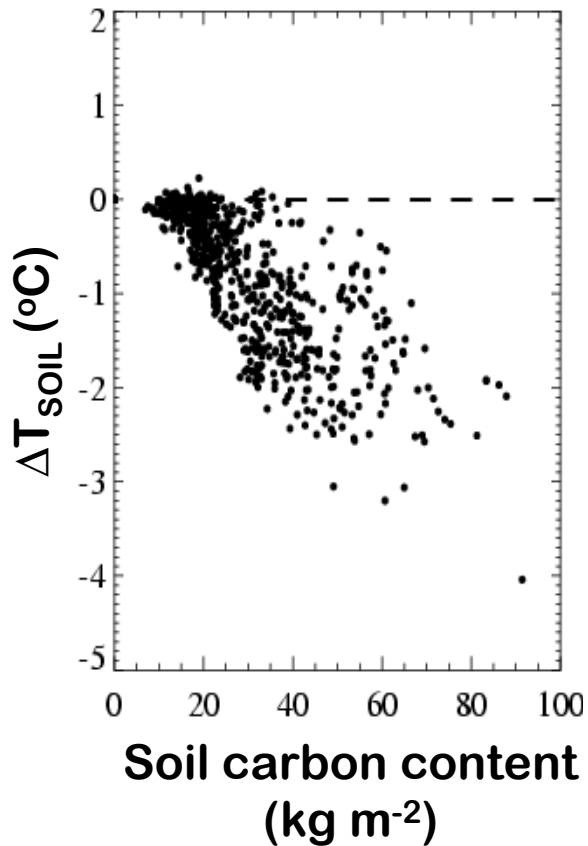
k_{sat} sat. hydraulic conductivity

^a Farouki (1981), ^b Letts et al. (2000)

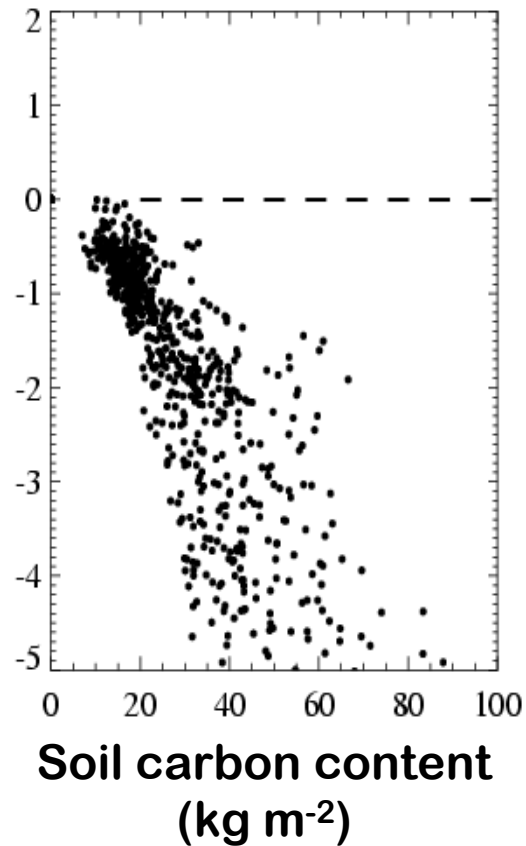


Impact on soil temperature (SOILCARB – CONTROL)

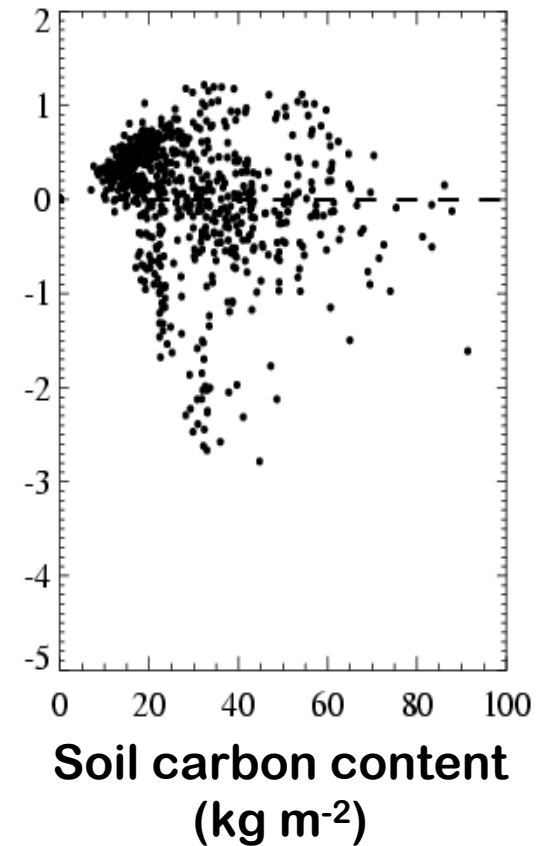
Annual Mean



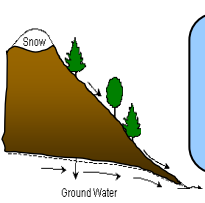
JJA Mean



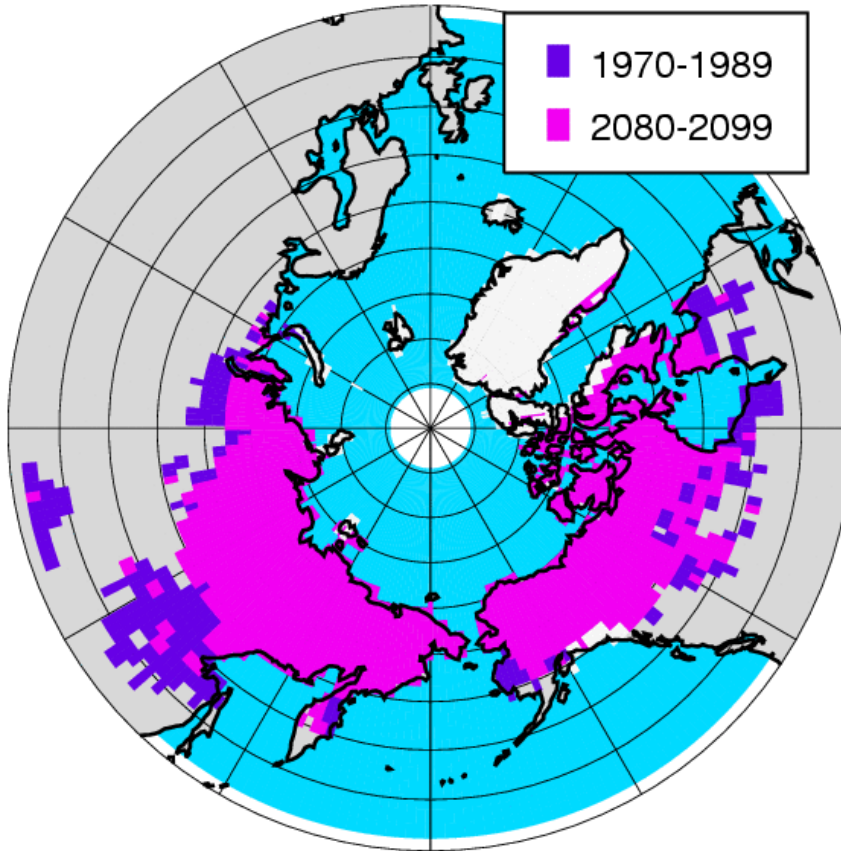
DJF Mean



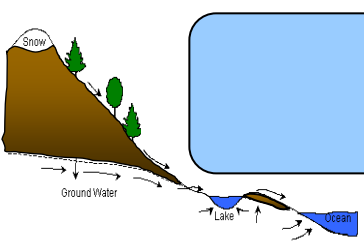
Organic soil layers cool soil temperatures by up to -3°C



Deep permafrost (10-30m)



**Most deep
permafrost still
exists at the end of
the 21st century**



Future Work: Simulating Arctic terrestrial feedbacks in CCSM



Global warming

Carbon sequester

Shrub growth

Permafrost warms and thaws

Enhanced [nitrogen]

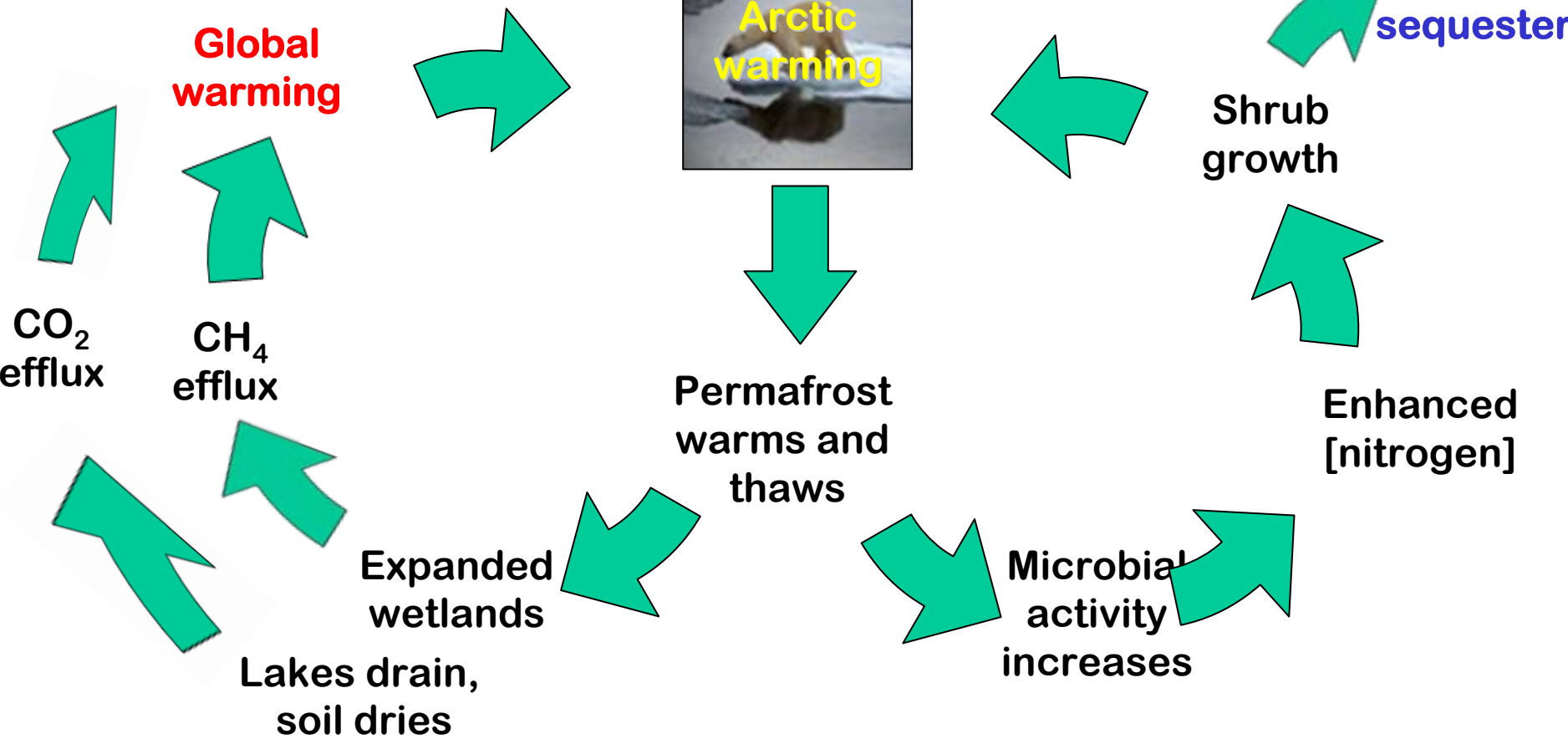
CO₂ efflux

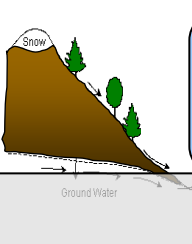
CH₄ efflux

Expanded wetlands

Microbial activity increases

Lakes drain, soil dries





Future Work: Simulating Arctic terrestrial feedbacks in CCSM



Global warming

CLM-CN (carbon cycle model): simulates large Arctic soil carbon pool

CLM-DGVM (dynamic vegetation): recently added shrub type

Carbon sequester

Shrub growth

CO₂ efflux
CH₄ production

Permafrost warms and thaws

Enhanced [nitrogen]

Wetland CH₄ emission model

Microbial

Reduce climate biases – snow model, low clouds, resolution in boundary layer, cold region hydrology

Dynamic wetlands, soil subsidence, thermokarst

wetlands

Lakes drain,