Constraining the aerosol indirect effect

Jón Egill Kristjánsson (Univ. Oslo)

Many thanks to: Alf Kirkevåg (met.no), Corinna Hoose (Univ. Oslo), Leo Donner (GFDL)

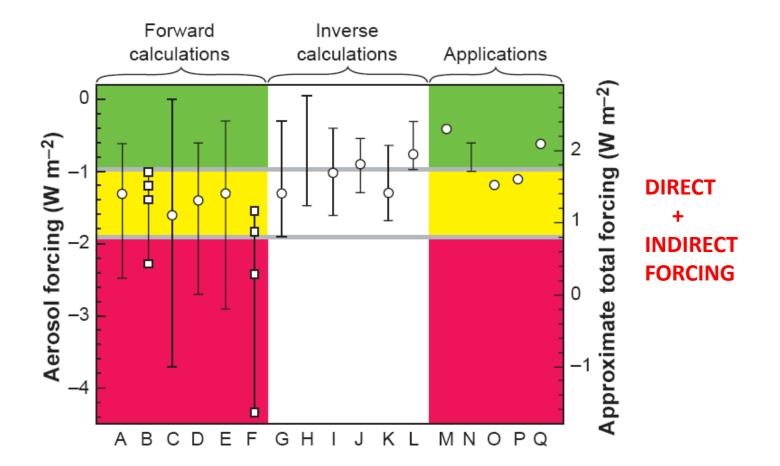
Problem Definition

- Most GCMs give an aerosol indirect effect which is too high compared to results from residual calculations – Why?
- Many models have built in constraints on parameter values that keep the indirect effect within reasonable bounds – Is this justifiable?
- What can be done?

Aerosol Indirect Effect

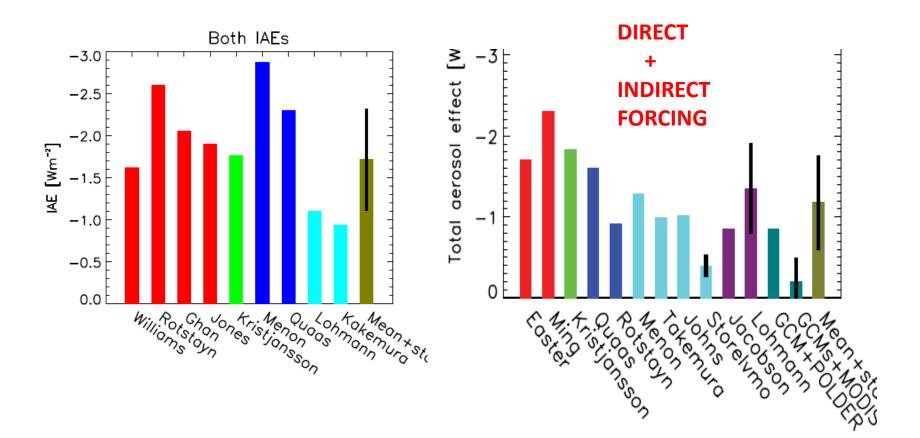
- <u>Definition</u>: Change in Cloud Radiative Forcing due to Anthropogenic Aerosols
- Model estimates of AIE are sensitive to:
- The Aerosol Scheme, in particular the Treatment of Natural Aerosols
- Parameterizations of Aerosol-Cloud Interactions
- The Atmospheric State in the host model, in particular the Cloud Properties

Models tend to overestimate the aerosol indirect effect



Anderson et al. (2003: Science)

Models tend to overestimate the aerosol indirect effect?



Lohmann and Feichter (2005: ACP)

IPCC (2007)

Sensitivity to background aerosols

TABLE 4. Globally averaged annual means of the AIE evaluated by the model from the difference in cloud radiative forcing. Also included are the NH and SH, and land and ocean averages. NR-DE refers to the AIE calculated as the difference between net radiation and the direct aerosol effect. Global annual Δ LCC and Δ LWP between PD and PI emissions are also given.

	AIE (W m^{-2})				NR-DE (W m^{-2})			
Experiment	Land	Ocean	NH	SH	Global	global	$\Delta LCC (\%)$	Δ LWP (g m ⁻²)
CTRL-R	-3.13	-1.31	-2.56	-1.09	-1.82	-1.95	0.15	-1.10
NEWCLD-R	-2.39	-1.22	-1.82	-1.27	-1.55	-1.72		-0.30
NEWCLD-M-7.5	-7.83	-2.99	-6.16	-2.56	-4.36	-4.68	1.18	7.80
NEWCLD-M-5.0	-2.91	-1.42	-2.39	-1.29	-1.84	-1.81	0.33	0.90
NEWCLD-M-5.0-P	-4.08	-1.75	-3.41	-1.41	-2.41	-2.53	0.56	1.90

Menon et al. (2002: J.Atmos.Sci.)

Cloud Susceptibility

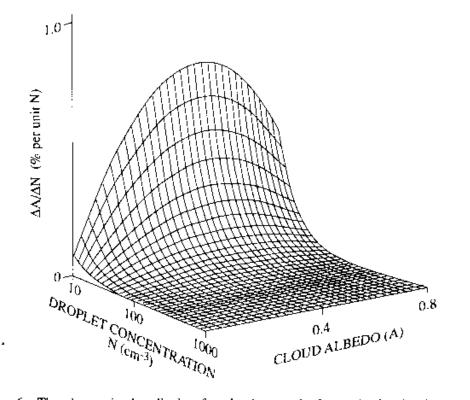
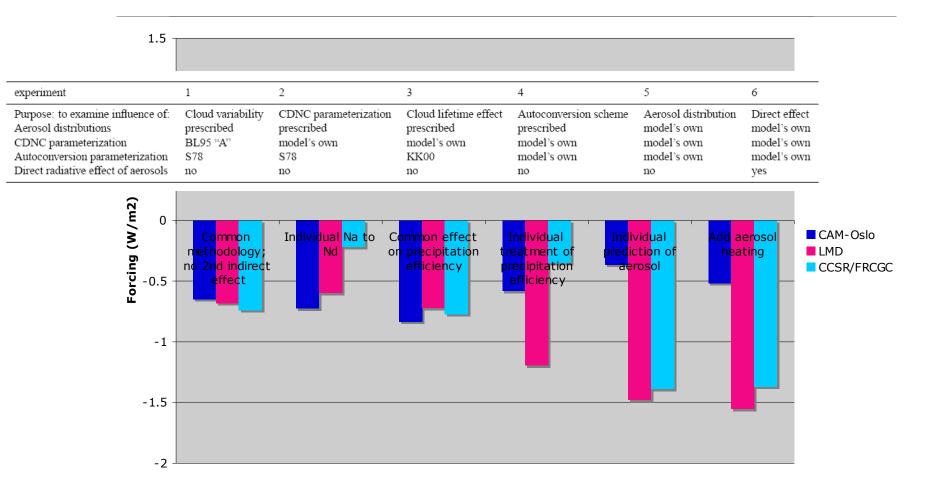


Figure 6 The change in the albedo of a cloud per unit change in the droplet concentration $(\Delta A/\Delta N)$ as a function of the cloud albedo (A) and the droplet concentration (N), for a cloud with a constant liquid-water content. (Adapted from Twomey, 1991.)

Hobbs (1993: Academic Press)

Indirect forcing in 3 GCMs: Model estimates differ mainly due to different parameterizations and different emissions



Penner et al. (2006: ACP)

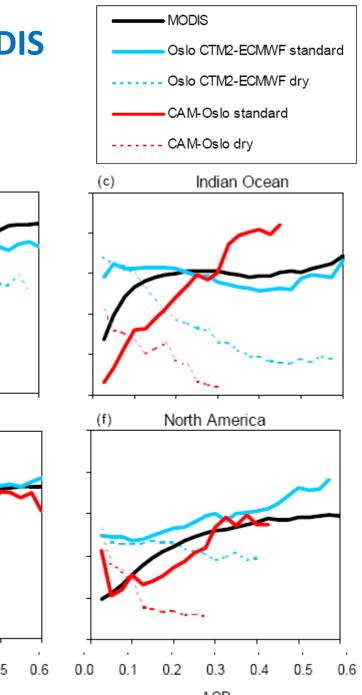
Cloud Fraction vs. AOD: GCM and MODIS

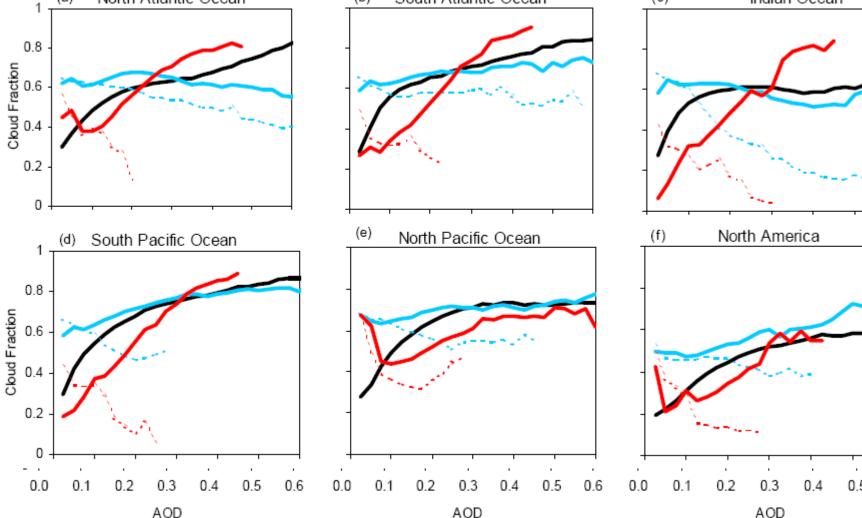
North Atlantic Ocean

(a)

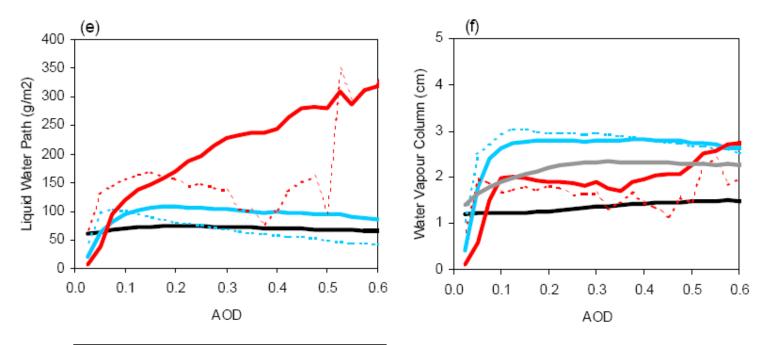
Myhre et al. (2007: ACP)

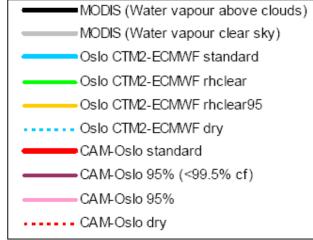
(b)





South Atlantic Ocean



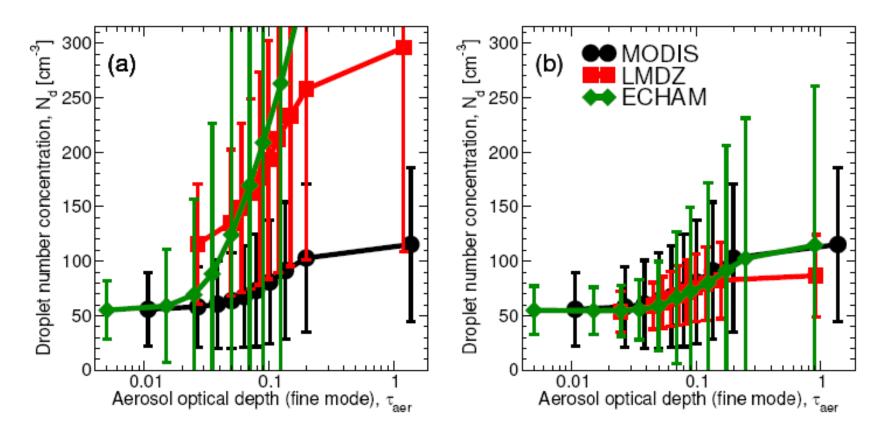


Myhre et al. (2007: ACP)

Constraining the indirect effect with observations

ORIGINAL

ADJUSTED



Quaas et al. (2006: ACP)

Constraining the indirect effect with observations

Table 1. Global annual mean radiative forcings by the total aerosol indirect effect.

Experiment	Standard (Wm^{-2})	Modified (Wm^{-2})
LMDZ	-0.84	-0.53
ECHAM4	-1.54	-0.29

Quaas et al. (2006: ACP)

Constraining CDNC reduces the indirect effect

- ECHAM does not allow CDNC < 40 cm⁻³
- How realistic is this constraint?
- What is the implication of it?

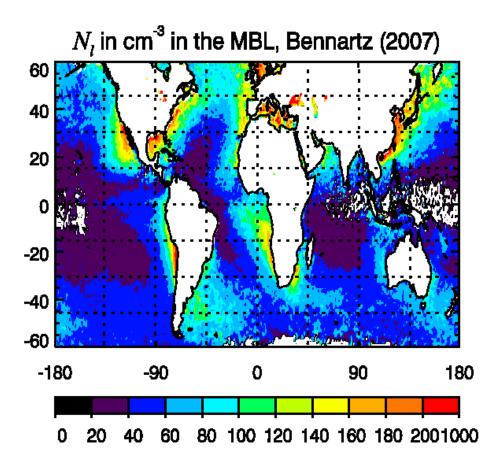


 Table 2. Statistics of Cloud Droplet Number Concentration for Remote Ocean Areas With at Least 1500 km Distance to the Next Major Landmass^a

Area	N All, cm ⁻³	N No Drizzle, cm ⁻³	N Drizzle, cm ⁻³	Fraction Drizzle, %
North Atlantic	89(99) ± 27	$118(120) \pm 23$	$50(56) \pm 10$	$56(48) \pm 20$
South Atlantic	$67(77) \pm 29$	$93(96) \pm 17$	$34(39) \pm 7$	$64(59) \pm 30$
North Pacific	$64(74) \pm 22$	$84(88) \pm 19$	$38(44) \pm 9$	$57(49) \pm 27$
South Pacific	$40(49) \pm 16$	$69(74) \pm 15$	$32(38) \pm 7$	$86(82) \pm 23$
South Indian	$42(51) \pm 18$	$76(80) \pm 14$	$32(38) \pm 7$	$79(72) \pm 21$

^aThe values given are two and a half year mean value with one standard deviation. Results are presented for all stratiform boundary layer clouds and separated in clouds with high/low likelihood of drizzle. The values in parentheses give the estimates that are derived using the parameterization of k derived by Lu and Seinfeld [2006]. Standard deviations for the estimates using Lu and Seinfeld [2006] are almost identical to the standard deviations for k = 0.8 and are not reported.

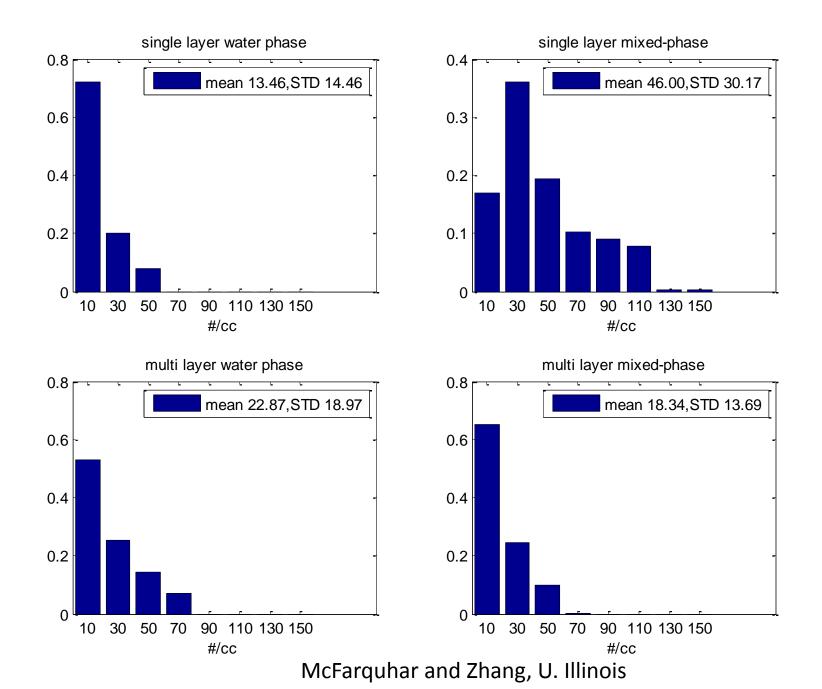
MCFARQUHAR ET AL.: MIXED-PHASE ARCTIC CLOUD OBSERVATIONS

Table 2. N_i, N_w, r_{ei}, r_{ew}, LWC, and TWC Averaged Over All Spirals Flown Through Single-Layer Mixed-Phase Clouds on 9 October, 10 October, and 12 October^a

Date	LWC, g m ^{-3}	IWC, gm ⁻³	r _{ew} , μm	r _{ei} , μm	$N_w, x10^3 L^{-1}$	N_i, L^{-1}
9 Oct	0.193 ± 0.131	0.025 ± 0.060	9.37 ± 2.23	25.48 ± 1.30	72.21 ± 34.37	5.62 ± 12.10
10 Oct (a)	0.174 ± 0.120	0.015 ± 0.032	9.04 ± 2.41	24.61 ± 2.35	25.74 ± 13.43	1.60 ± 2.40
10 Oct (b)	0.154 ± 0.116	0.006 ± 0.006	10.93 ± 2.57	25.76 ± 5.72	23.00 ± 9.97	2.04 ± 2.06
12 Oct	0.193 ± 0.116	0.006 ± 0.018	9.07 ± 2.29	25.15 ± 7.28	51.73 ± 16.60	2.07 ± 4.97

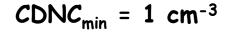
^aStandard deviations correspond to deviations of the average value of each spiral from the average value integrated over all of the spirals.

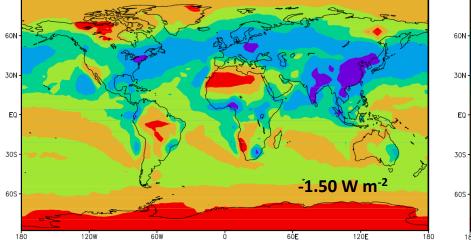
McFarquhar et al. (2007: JGR)

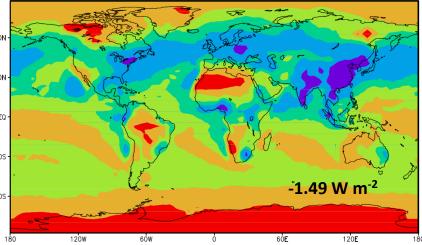


Tests with CAM3-Oslo (1-year runs)

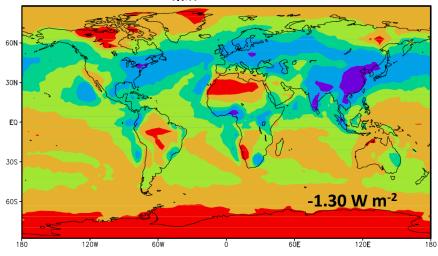
no cut-off on CDNC



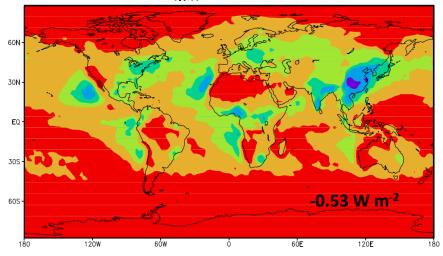




 $CDNC_{min} = 10 \text{ cm}^{-3}$



 $CDNC_{min} = 40 \text{ cm}^{-3}$



Anthropogenic Ice Nuclei

Ехр. →	LIQ- UID	CON- TROL	KAO- LINITE	LESS DUST	COAT- ING
∆LWP (g m⁻²)	+ 0.87	- 0.07	- 0.32	+ 0.68	+ 0.64
∆IWP (g m⁻²)	- 0.04	+ 0.20	+ 0.36	+ 0.52	- 0.46
∆R _{eff} (µm)	- 0.44	- 0.33	- 0.32	- 0.41	- 0.43
INDIR (W m ⁻²)	- 0.49	- 0.07	- 0.10	- 0.18	- 0.27

Storelvmo et al. (2008: J.Atm.Sci., in press)

Summary and Conclusions

PROBLEM

- Most GCMs struggle to keep the aerosol indirect forcing low enough to yield realistic climate simulations (~ -1 W m⁻²)
- The AIE is very sensitive to background (pre-industrial) aerosols
- Comparisons to observations indicate too large sensitivity of cloud parameterizations to aerosol burdens

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POSSIBLE SOLUTIONS

- Constraining with satellite data yields significantly suppressed indirect effect
- Constraining with prescribed bounds on CDNC reduces AIE, but violates observations in remote regions
- Indications that AIE of mixed-phase clouds may reduce the overall indirect forcing
- Enhanced evaporation may lead to a positive indirect effect in trade wind cumuli (Feingold)
- Competition effect reduces AIE (Ghan)

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 Comparisons to observations indicate too large sensitivity of cloud parameterizations to aerosol burdens

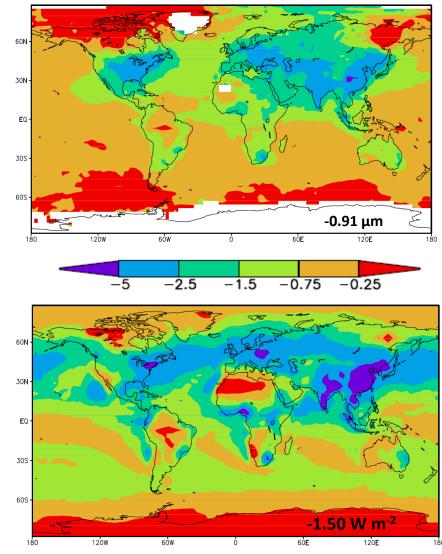
POSSIBLE SOLUTIONS

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- Indications that AIE of mixed-phase clouds may reduce the overall indirect forcing
 - Inhanced evaporation may lead to a positive indirect effect in trade wind cumuli (Feingold)
 - **Competition effect reduces AIE (Ghan**

Reports of low CDNC measurements

- Bower et al. (2006: Atm. Res.): In-situ ship measurement from remote area SH ocean: 8 cm⁻³
- Yum and Hudson (2004: JGR): Southern Hemisphere Oceans: 20-40 cm⁻³
- Bennartz (2007: JGR): MODIS-based retrievals: Average values of 41±17 cm⁻³ in PBL clouds in South Pacific and South Indian Oceans
- McFarquhar et al. (2007: JGR): Arctic measurements in mixed-phase clouds (M-PACE) of between 23±10 cm⁻³ and 72±34 cm⁻³

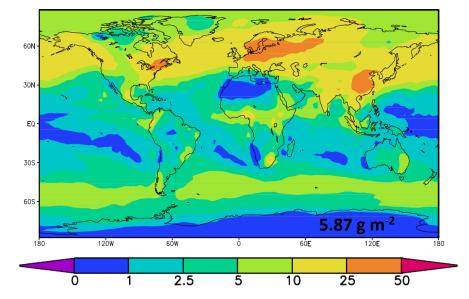
standard CDNC treatment (no lower cut-off)



Change in effective radius as seen from satellite

1+2. indirect radiative forcing

Change in cloud liquid water path



Present day:

LWP = 133.1 g m⁻² Reff = 12.93 μ m CDNCint = 3.95e6 cm⁻²

Pre-industrial:

$$CDNC_{min} = 1 \text{ cm}^{-3}$$

60N 30N -EQ -30S -60S -0.91 µm 60E 120E 120W 60W -1.5 -0.75 -0.25 -5 -2.5 60N 30N-EQ 30S-60S --1.49 W m⁻² 60E 120E 180 120W 60W ò

1+2. indirect radiative forcing

Change in cloud liquid water path

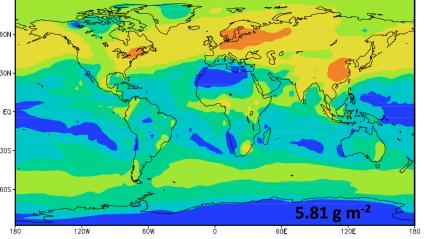
60N -30N -EQ -30S -60S-5.81 g m⁻² 60W 60E 120E 120W ó 2.5 25 50 0 5 10 1

Present day:

= 133.6 g m⁻² LWP = 12.95 μm Reff CDNCint = 3.95e6 cm⁻²

Pre-industrial:

= 127.8 g m⁻² LWP Reff = 13.87 μm CDNCint = 2.57e6 cm⁻²



Change in effective radius as seen from satellite

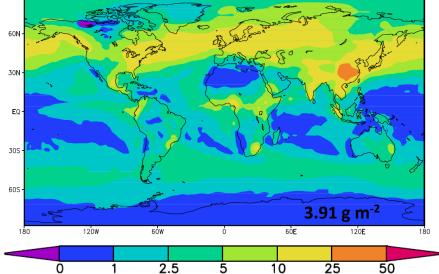
$CDNC_{min} = 10 \text{ cm}^{-3}$

60N 30N EQ 30S -60S -0.79 µm 60E 120E 120W 60W -1.5 -0.75 -0.25 -5 -2.560N 30N EQ 30S · 60S · -1.30 W m⁻² 60E 120E 180 120W 60W

Change in effective radius as seen from satellite

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Present day:

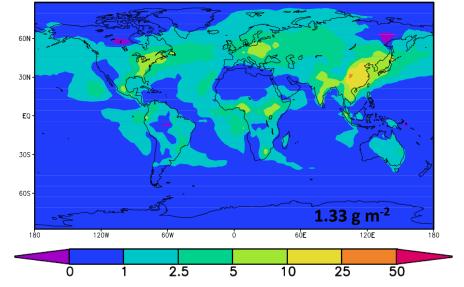
Pre-industrial:

$$CDNC_{min} = 40 \text{ cm}^{-3}$$

60N-30N -EQ -30S -60S --0.44 µm 60E 120E 120W 60W -1.5 -0.75 -0.25 -5 -2.5 60N 30N EQ 30S 60S --0.53 W m⁻² 60E 120E 180 120W 60W ó 180

1+2. indirect radiative forcing

Change in cloud liquid water path

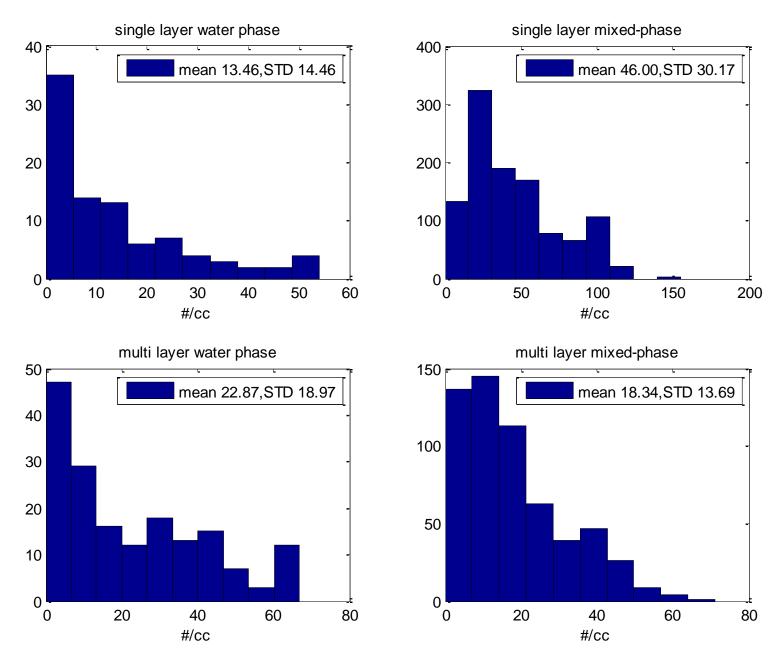


Present day:

Pre-industrial:

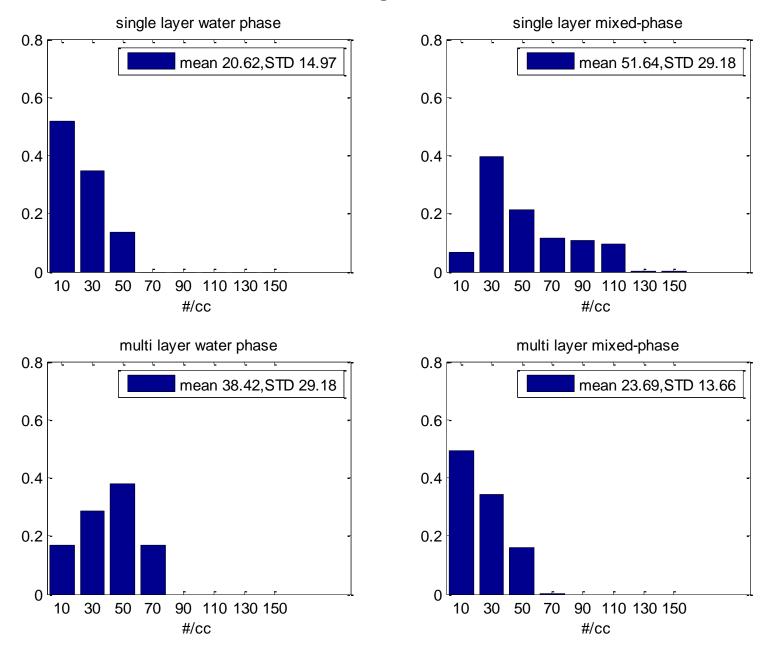
LWP = 148.1 g m⁻² Reff = 11.43 μ m CDNCint = 2.57e6 cm⁻²

Change in effective radius as seen from satellite



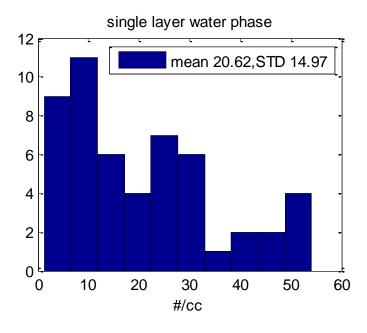
McFarquhar and Zhang, U. Illinois

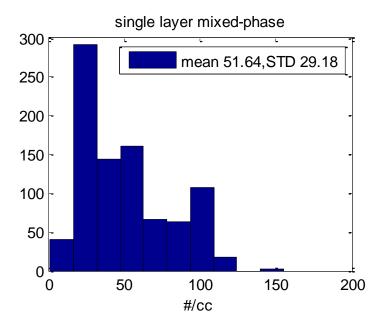
LWC>0.05g/m³

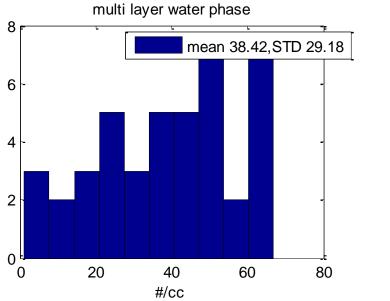


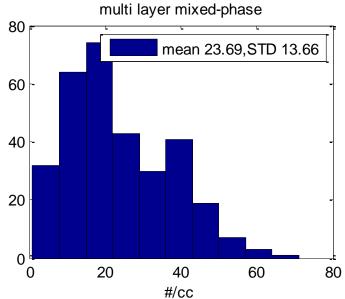
McFarquhar and Zhang, U. Illinois

LWC>0.05g/m³









McFarquhar and Zhang, U. Illinois

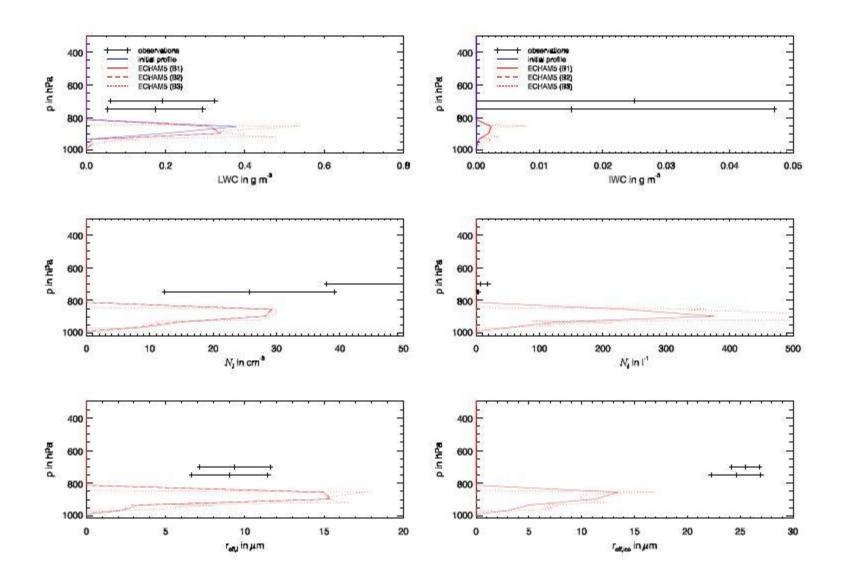
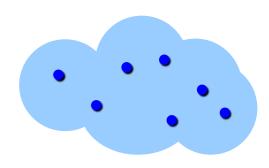


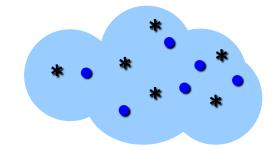
Figure 4.7: Model results for M-PACE period B, averaged over the 12h simulation period. Simulations B1 and B2 lie close together. Observations are aircraft data from McFarquhar et al. (2007), averaged over several vertical spirals for two flights during period B. The standard deviations correspond to deviations of the vertically averaged value of each spiral to the average value over all spirals per flight.

C. Hoose (2008: Ph.D. thesis)

Warm and cold clouds

Warm clouds \longrightarrow clouds with T > 0°C

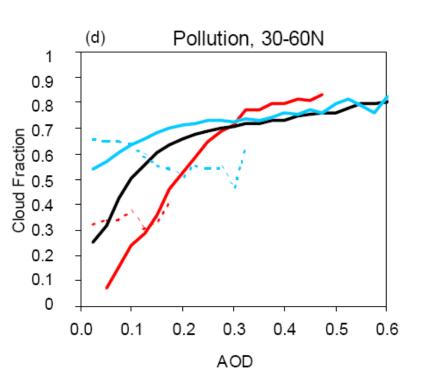


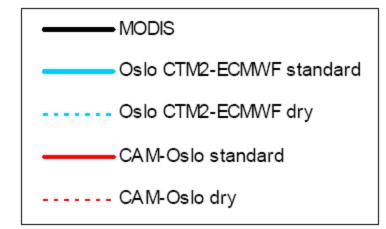


mixed-phase clouds (~ -35°C < T < 0°C)

Cold clouds

ice clouds (cirrus) (T < ~ -35°C)</p>





Aerosol Indirect Forcing in CAM3-Oslo

1. indirect effect standard simulation (-1.78 W m^{-2}) 1. indirect effect Prognostic CDNC (-1.13 W m⁻²) 90N 90N · 60N 60N -5 -2 -5 -5 30N -30N -<u>___</u> EQ EQ -2 30S -30S -60S 60S 90S 90S 120E 180 120E 120W 60E 120W 60W 60E 180 60W 180 180 0 $\left(\right)$ -1 -0.5 -1 -5 -2 0 -2 -0.5 -1

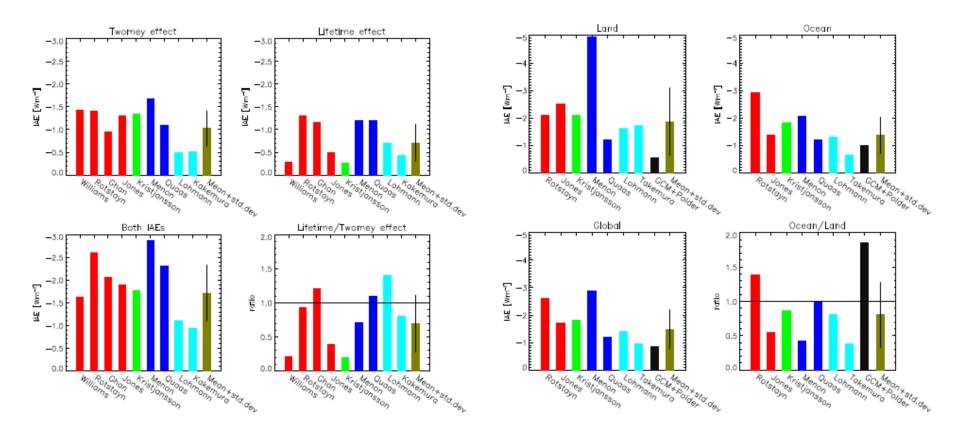
Indirect forcing reduced by 35%, largely due to competition effect!

Diagnostic CDNC

Seland et al. (2008: Tellus A)

Prognostic CDNC

Model Estimates of the Aerosol Indirect Effect



Lohmann & Feichter (2005: ACP)

Cloud Feedback

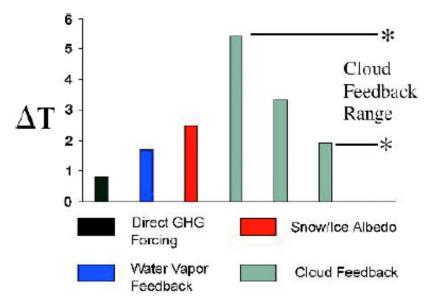


FIG. 13. The response of a single climate model to an imposed doubling of CO_2 as different feedbacks are systematically added in the model (adapted from Senior and Mitchell 1993). Different treatments of cloud processes in the model produce a large spread in predicted surface temperature due to CO_2 doubling. Sensitivity to the treatment of clouds and cloud-radiative processes

Stephens (2005: J. Climate)