Progress on U.W. PBL and Macrophysics reformulations Interactions among cloud, radiation, and PBL turbulence

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the same and with some

Sungsu Park, Christopher S. Bretherton, Phil J. Rasch Dept. of Atmospheric Sciences, Univ. of Washington, Seattle NCAR

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Outline

- I. Inconsistency between cloud fraction and in-cloud LWC
- II. Offline computation of stratiform macrophysics
- III. Introduction of a new macrophysics scheme
- IV. UW PBL + UW ShCu + RK Microphysics + PBR Macrophysics
- V. UW PBL + UW ShCu + MG Microphysics + PBR Macrophysics

Inconsistency between stratus fraction and in-stratus LWC

Interplay among various processes in stratocumulus



Interplay among various processes in stratocumulus



Dcean

 \rightarrow Too strong aerosol indirect effect in current CAM ?





Isothermal, cloud in the PBL top layer is a black body, transparent clear air

$$\longrightarrow \left(\frac{dT}{dt}\right)_{LW} = -\left[\frac{g}{C_p \cdot \Delta p}\right] (a \cdot \varepsilon \cdot \sigma T^4)$$
$$\varepsilon = 1 - \exp(-1.66 \cdot k \cdot \hat{q}_l) \cdot \frac{\Delta p}{\rho \cdot g}) \quad k = 90.36$$

[K hr-1]	CAM	Nature
Ambiguous L.	-1.2	-4.9
PBL Top L.	~ -4.0	~ 0

Ambiguous Layer

PBL Top Layer



 θ_v

As a result, current CAM suffers from

- \rightarrow Strong inversion at the PBL top
- \rightarrow Too weak entrainment
- \rightarrow Too shallow, cold, moist PBL
 - Too much (less) subtropical stratocumulus in downstream (upstream)
 - Suppression of nocturnal deepening of PBL
 - Too strong ENSO amplitude due to too weak SST damping by weak upward LHF

Summary Inconsistency between cloud fraction and in-cloud LWC can exert large Influences on global climate system through complex feedbacks among cloud, radiation, and moist turbulence

 \rightarrow What caused inconsistency between cloud fraction and in-cloud LWC ?

What caused inconsistency between cloud fraction and in-cloud LWC ? \rightarrow Stratiform Macrophysics Scheme

- Isolate stratiform macrophysics from the CAM and perform off-line computation
- Force the ambiguous layer at p = 900 [hPa], T = 280 [K], q_v = 6.84 [g kg⁻¹], q_l = 0.16 [g kg⁻¹], a = 0.6, Δ p = 20 [hPa] with various external forcings of temperature (A_T) and water vapor (A_{qv})
- Neglect cumulus, cloud ice, and precipitation
- Examine $\Delta a vs \Delta q_{I,cloud}$
- Test for CAM30, CAM35, Equilibrium CAM35, and Triangular PDF with a half width of total specific humidity = 0.1*q_s(T,p)

How stratiform net condensation rate Q is computed in CAM













Equilibrium CAM35 produces consistent cloud fraction and in-cloud LWC. \rightarrow Realistic LW cooling rate and reasonable response of moist turbulence within PBL

A New Stratiform Macrophysics Scheme

- 1. Uses equilibrium cloud fraction and equilibrium state variables for computing Q
- 2. Formulation based on conservative scalars
 - Consistent with the assumption of uniform T within the grid
- 3. Incorporation of fusion heat in computing Q
 - Treatment of ice and mixed-phase clouds
- 4. Explicit treatment of cumulus cloud
 - Cumulus and stratus are non-overlapped in each layer and have their own in-cloud LWC and cloud fraction

Original Macrophysics



- Overlap
- In-cumulus CWC = In-stratus CWC

New Macrophysics



- Non-overlap
- In-cumulus CWC ≠ In-stratus CWC

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 - Cumulus and stratus are non-overlapped in each layer and have their own in-cloud LWC and cloud fraction
- 5. Has the following functionalities:
 - Stratus fraction formula based on either RH or triangular PDF (CAMstfrac)
 - For PDF-based cloud, $U_{clr} \rightarrow 1$ as $a_{st} \rightarrow 1$ consistent with the real world
 - Specify in-cloud CWC (LWC+IWC) of newly formed or dissipated stratus from zero (cc=0) to the CWC of pre-existing stratus (cc=1)
 - □ Force in-stratus CWC to be bounded by externally-specified limiting values (qcst_min>0, qcst_max>0) by performing pseudo condensation-evaporation in each layer.
 - Natural removal of 'empty' (a>0, q_{I,cloud}=0) and 'dense'(a=0, q_{I,cloud}>0) cloud
 - Potential replacement of Vavrus polar cloud fix
 - Explicit or implicit computation of Q by choosing different iteration number (niter)
 - Potential to allow super-saturation within the stratus in any phases

CAMUW + RK Micro + New Macro

- 5 years AMIP runs using version CAM3_5_42
- Several refinements are made to UW PBL and UW shallow convection
 - Increase turbulent master length scale in the convective regime
 - Refined computation of TKE at the entrainment interfaces
 - Maximum cumulus updraft core fractional area of 10 % instead of 5 %
 - Refined identification of penetrative entrainment zone
- Cumulus fraction and in-cumulus CWC are not included in computing radiation and grid-mean CWC.
- Switches in the macrophysics scheme
 - CAMstfrac = RH cloud, cc = 1, niter = 3, qcst_min = 0.01, qcst_max = 3 [g kg⁻¹]



CAMUW + MG Micro + New Macro

- Condensation (Q) into cloud liquid $[Q_w = (1-f)^*Q]$ and cloud ice $[Q_i = f^*Q]$ are explicitly treated within the macrophysics scheme by setting
 - $f = q_{i,cloud} / (q_{i,cloud} + q_{i,cloud})$
 - Bergeron-Findeisen process within the MG microphysics scheme is treated as a separate process independent of Q
- In the MG microphysics, droplet activation occurs
 - only when Q > 0 instead of when 'q_{l,cloud}, q_{i,cloud} > 0'
 - only one time at each time step, that is, mtime = 1 instead of mtime = $\Delta t / t_o$ where t_o is a mixing time scale of aerosol within cloud



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- Cumulus fraction and in-cumulus CWC are explicitly included in computation of radiation and grid-mean CWC with appropriate tunings (dp1 = 0.03, $c_0 = 0.02$)
- Switches in the macrophysics scheme
 - CAMstfrac = RH cloud, cc = 0, niter = 3, qcst_min = 0.01, qcst_max = 3 [g/kg]

LWP. Annual Mean





Increase of LWP in the trade cumulus & deep convection regimes due to explicit treatment of in-cumulus LWC

SWCF. Annual Mean

Old Macro

New Macro

TOM SW cloud forcing mean= -59.87

W/m~S~2~N~

TOA SW cloud forcing mean= -49.21

₩/m~S~2~N~

45 30 15 -15 -30 -45 -60 -75 -90 -105 -120 -135 -150 -170

> 120 100 80 60 40 20 10 -10 -20 -40 -60 -80 -100 -120





CERES Observation

TOA SW cloud forcing mean= -48.59

W/m~S~2~N~

mean = 10.63

rmse = 16.83

ΔSWCF. New – Old

W/m~S~2~N~



- Huge improvement of SWCF
- Enhanced SWCF in the trade due to explicit contribution of Cu CWC to radiation





'Qi= f^*Q' + 'Forcing in-stratus CWC to be bounded by two limiting values' \rightarrow Increase of in-cloud IWC in the storm track

UW PBL + UW ShCu + MG Micro + New Macro

	RESTOM	SWCF	LWCF	SWCF	TGCLDLWP	PRECT	PREH2O	LHFLX
	[W m ⁻²]	[W m ⁻²]	[W m ⁻²]	at 60°S. DJF	[g m ⁻²]	[mm day ⁻¹]	[mm]	[Wm ⁻²]
Base ⁰	-6.3	-63.3	32.6	-170	105	2.91	26.5	85.1
Enhance entrainment ¹ (a2l = 30 ← 15)	-4.5	-61.4	32.4	-163	100	2.91	26.5	85.1
Zero-CWC of newly formed stratus ²	-7.4	-64.2	32.3	-168	90	2.99	26.3	87.5
Cloud drop activation only when Q > 0 ³	-1.7	-56.2	30.1	-147	90	2.96	26.2	86.4
Increase activation time scale of CCN ⁴ (mtime = 1 ← 1.5)	-5.4	-62.0	32.2	-170	101	2.92	26.3	85.5
Enhance conversion of deep Cu LWC to precip. ⁵ ($c_o = 0.02 \leftarrow 0.01$)	-4.9	-61.7	32.1	-170	97	2.92	26.4	85.4
Reduce deep Cu fraction ⁶ (dp1 = 0.03 ← 0.05)	-5.6	-62.7	32.4	-170	100	2.92	26.5	85.3
Neglect Cu contribution in computing radiation	-2.6	-58.8	32.0	-157	85	2.91	26.6	85.2
0 +1(a2l=20)+ 2 + 3 + 4 + 5	0.3	-51.7	27.9	-140	75	3.07	25.8	89.8
Observation	0	-48.6~-54.2 (CERES, ERBE)	27.2 ~30.4 (CERES, ERBE)	-148 (CERES)	78 ~ 113 (NVAP, MODIS)	2.61 (GPCP)	25.0 (ERA40)	82.4 (ERA40)

Summary of Global Model Performance

Variable	CAM30 rms error	CAM3.5 (Revise ZM deep conv.)	UW PBL UW ShCu RK Micro ZE Macro	UW PBL* UW ShCu* MG Micro ZE Macro	UW PBL UW ShCu MG Micro* PBR Macro	
SLP	3.5 hPa	0.82 0.86		0.88	0.85	
Surface wind stress	0.05 N m ⁻²	0.82	0.81	0.76	0.84	
Zonal wind at 300 hPa	4.5 m s ⁻¹	0.77 0.80 0.73		0.73	0.74	
Surface rainfall	1.7 mm day ⁻¹	0.93	0.92	0.93	0.99	
Air temperature at 2m	3.5 K	0.94	0.95	0.91	0.93	
SWCF	22.8 W m ⁻²	1.02	0.91	1.14	0.88	
LWCF	11.7 W m ⁻²	0.97	0.93	0.86	0.88	
т	2.1 K	0.84	0.81	0.97	0.88	
RH	11.0 %	0.78	0.76	0.79	0.76	
Climate Bias Index (CBI)		0.88	0.86	0.88	0.86	

Conclusion

- We developed a new stratiform macrophysics scheme which
 - ensures consistency between cloud fraction and in-cloud LWC by using equilibrium variables for computing Q,
 - removes many conceptual and mathematical inconsistencies in CAM's cloud system model by
 - using conservative scalars in computing Q,
 - taking into account of fusion heat,
 - treating cumulus (CWC as well as fraction) separately from stratus,
 - mimicking a PDF-approach in computing stratus fraction and in-stratus CWC,
 - forcing in-stratus CWC to be bounded by two limiting values,
 - allowing the possibility of super-saturation within stratus
- MG microphysics is modified, so that droplet activation occurs only when Q > 0 instead of whenever stratus exists → substantially reduced the bias of SWCF → may help to reduce too strong aerosol indirect effect (-2.3 → -1.1 Wm⁻²) ?
- Overall, our new macrophysics scheme deepens PBL in the stratocumulus deck and increases LWP in the trade cumulus regime, resulting in improved skill scores of SWCF and T. However, simulation of precipitation and surface wind stress is worsen due to enhancement of already strong hydrological cycle.
- Future and on-going works:
 - Find a tuning set to weaken hydrological cycle
 - Analysis of diurnal cycle and ENSO amplitude

Stochastic Ocean Mixed Layer Model

[Frankignoul and Hasselman 1977; Deser et al. 2003; Park et al. 2006]

$$\rho \cdot C_p \cdot \overline{H} \cdot \frac{\partial}{\partial t} SST' = -\lambda \cdot SST' + Others$$



Q': Downward surface heat flux anomaly

 $\lambda < 0$: SST anomaly is amplified \longrightarrow Positive feedback

aerosol

$$= \lambda_{LHF} + SHF + \lambda_{(SW + LW)cloud} + \lambda_{(SW + LW)clear}$$
PBL turbulence,
convective deflation stratocumulus,
anvil cirrus water vapor,
aerosol

convective deflation