

# Some Impacts of Sea Ice Micro-structure on Ice Algal Growth

Nicole Jeffery  
CNLS/CCS-2  
Los Alamos National Lab



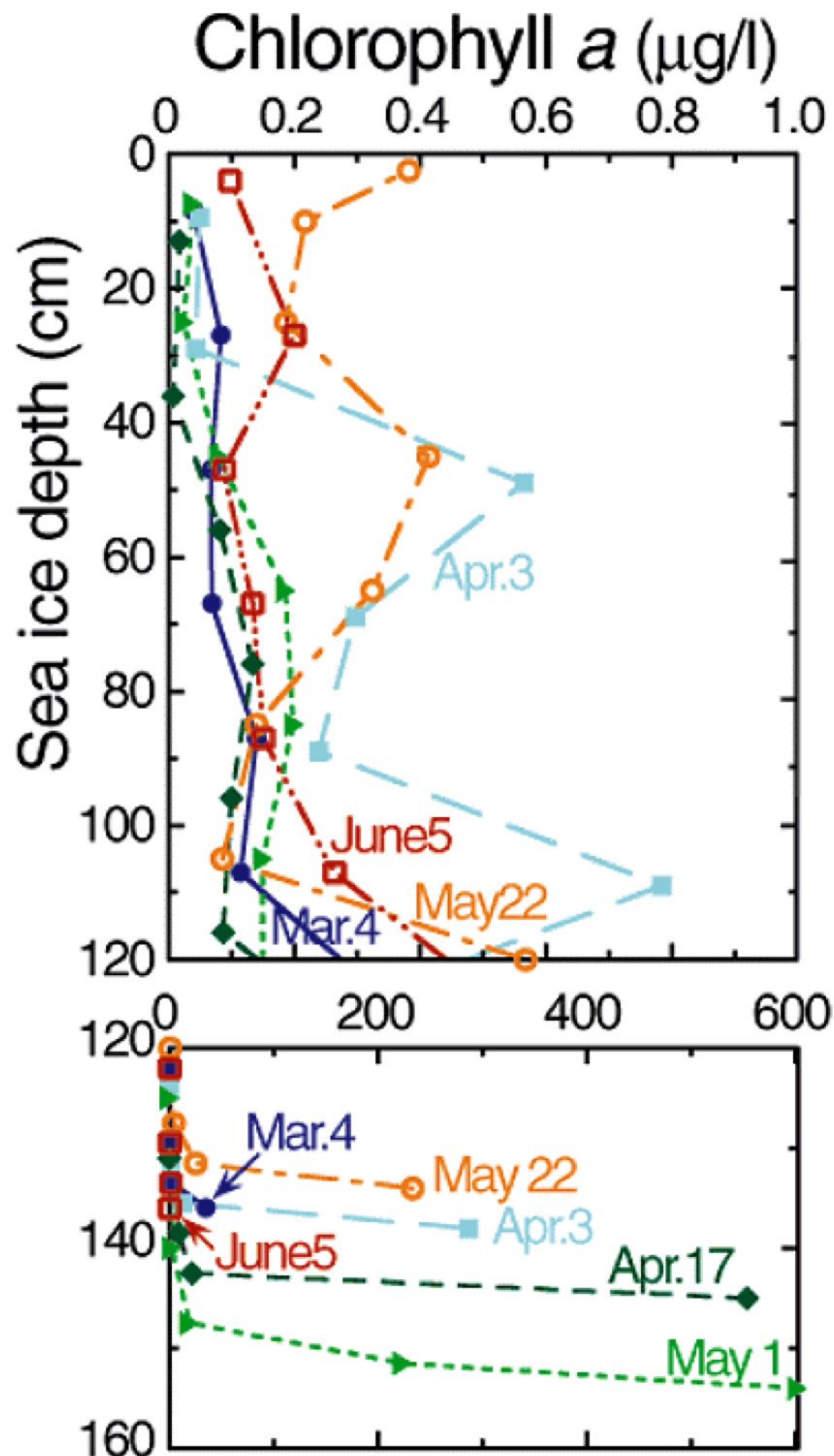
with Scott Elliott, Elizabeth Hunke, Mathew Maltrud, Clara Deal, & Meibing Jin

# Why Ice Algae?

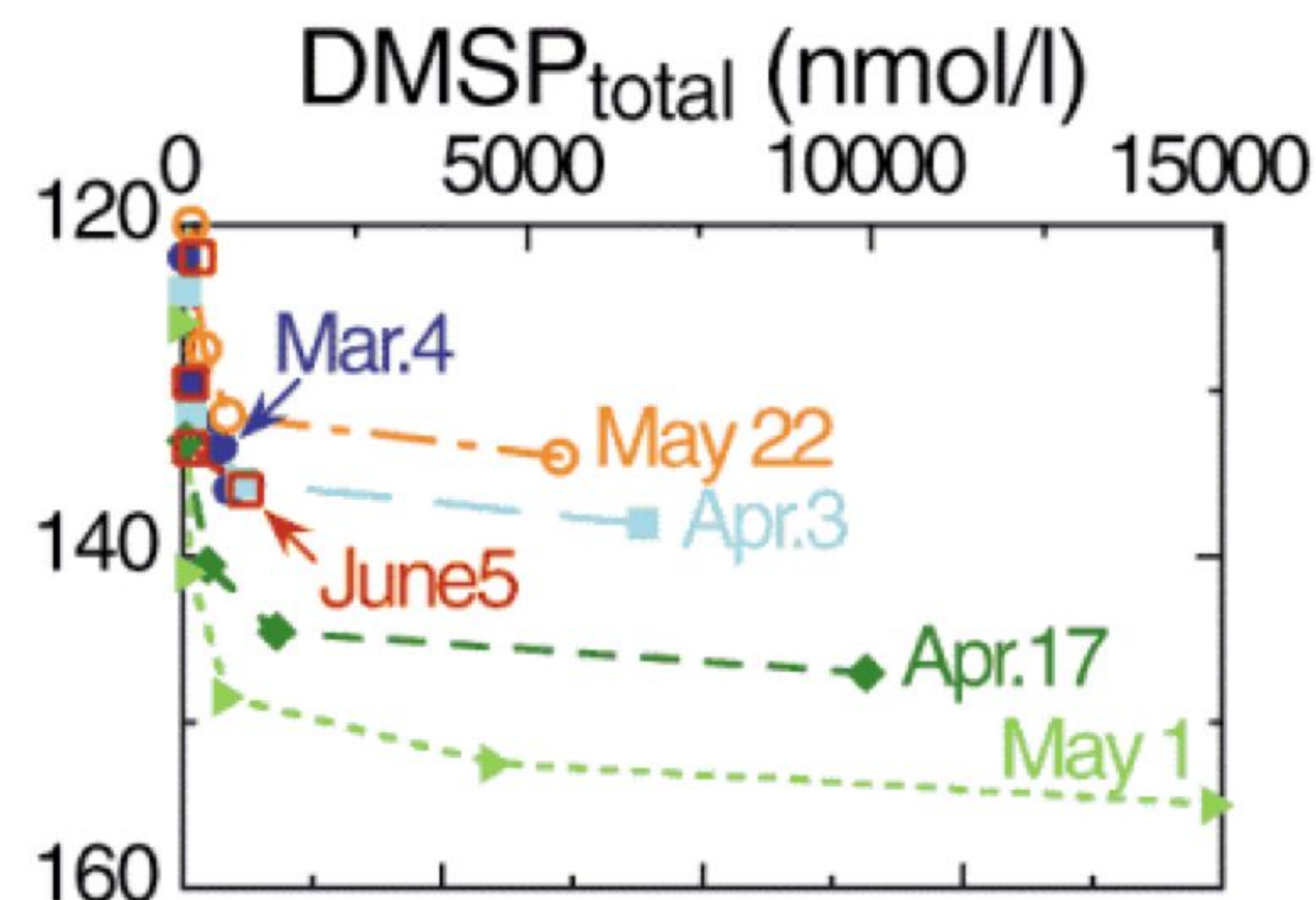
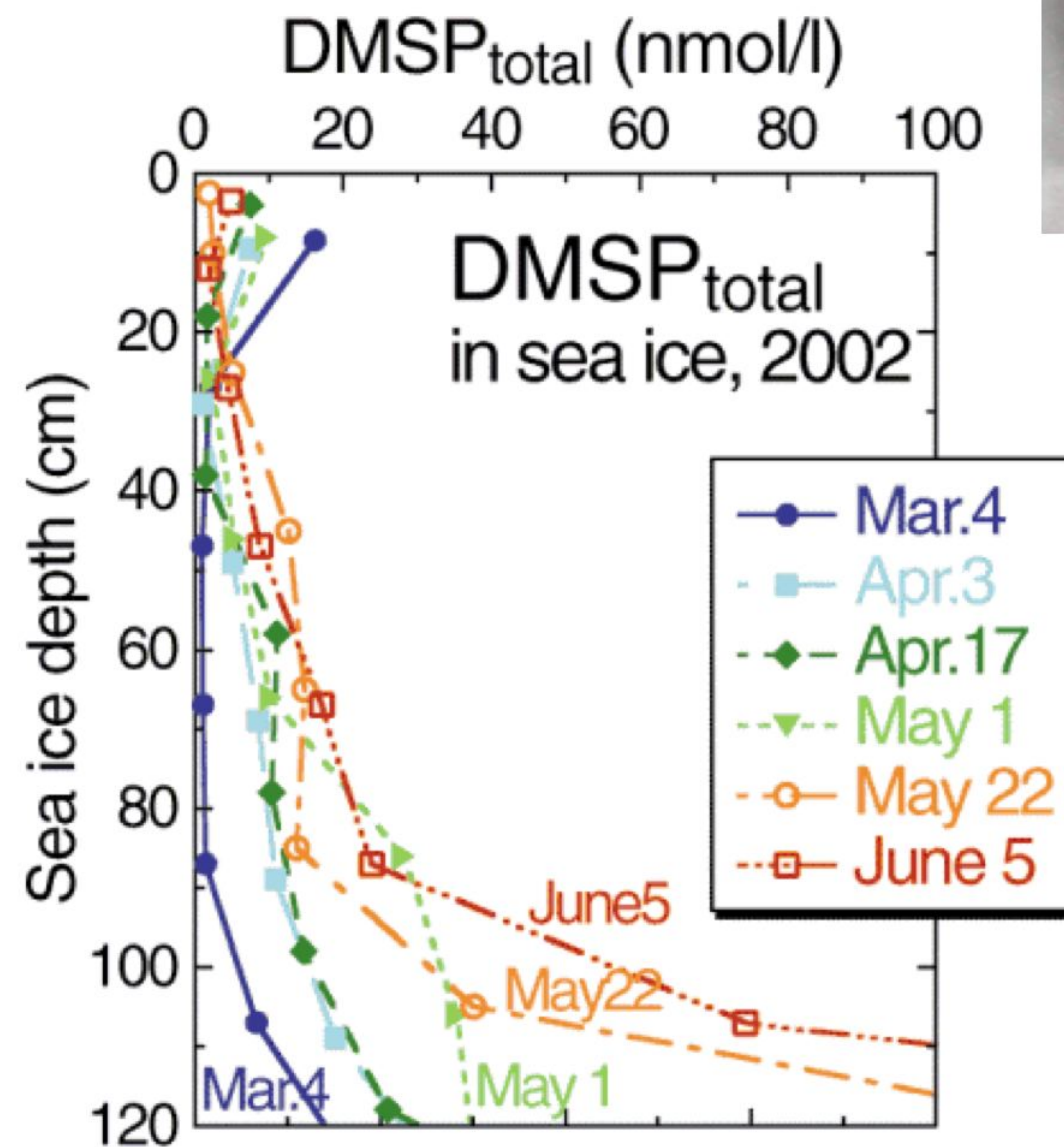
1. High concentrations of Chla and DMSP found in skeletal layer of Arctic Sea Ice



*Krembs & Deming  
University of Washington*



*Shin et al. 2003*



*Uzuka et al. 2003*

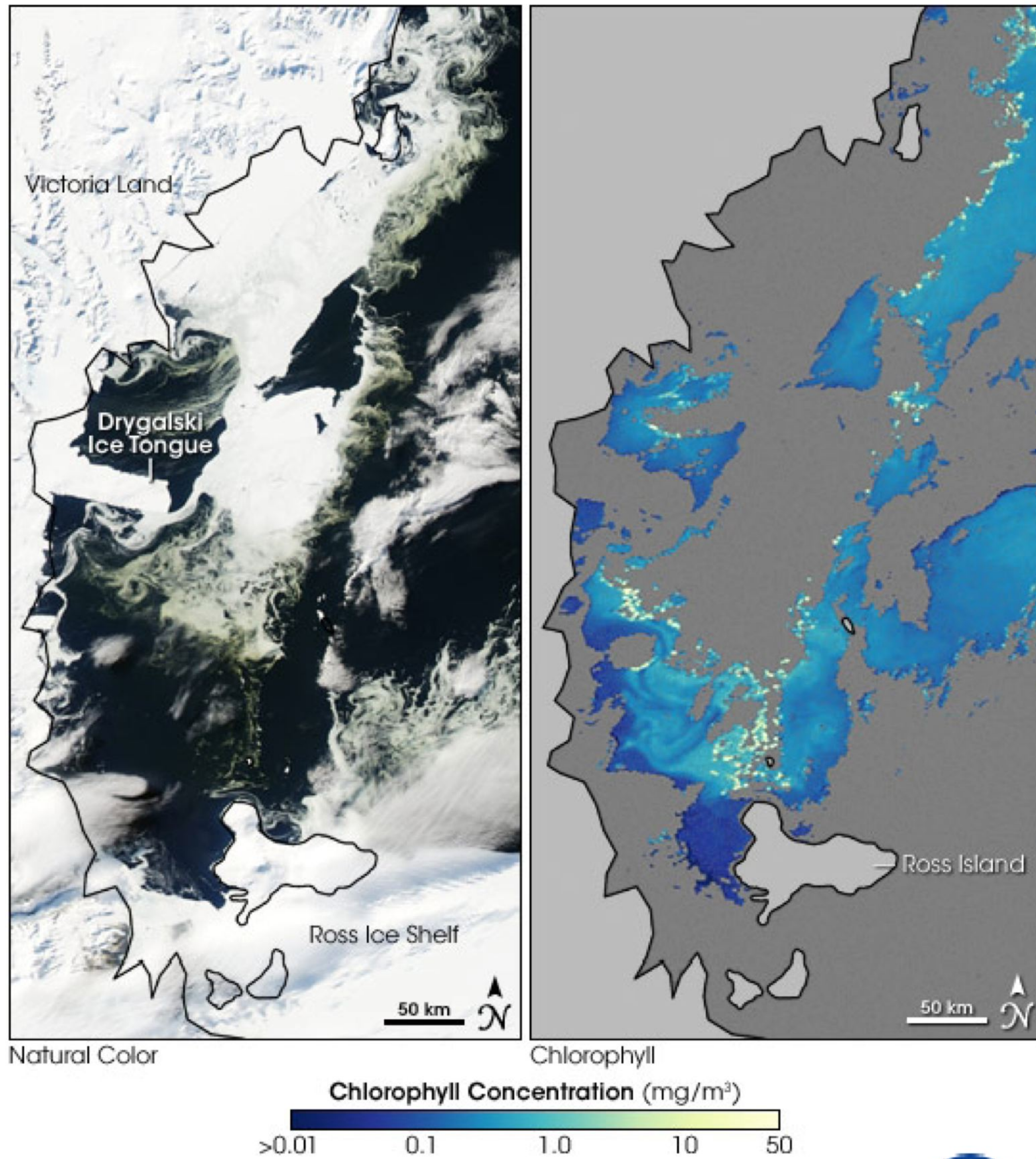
# Why Ice Algae? (cont'd)

2. Modifies sea ice optical and thermodynamic properties. ~ 85% of radiation absorbed through photosynthesis
3. Extends the growing season and provides an important niche in polar ecosystems
4. Plays a role in ice edge dynamics

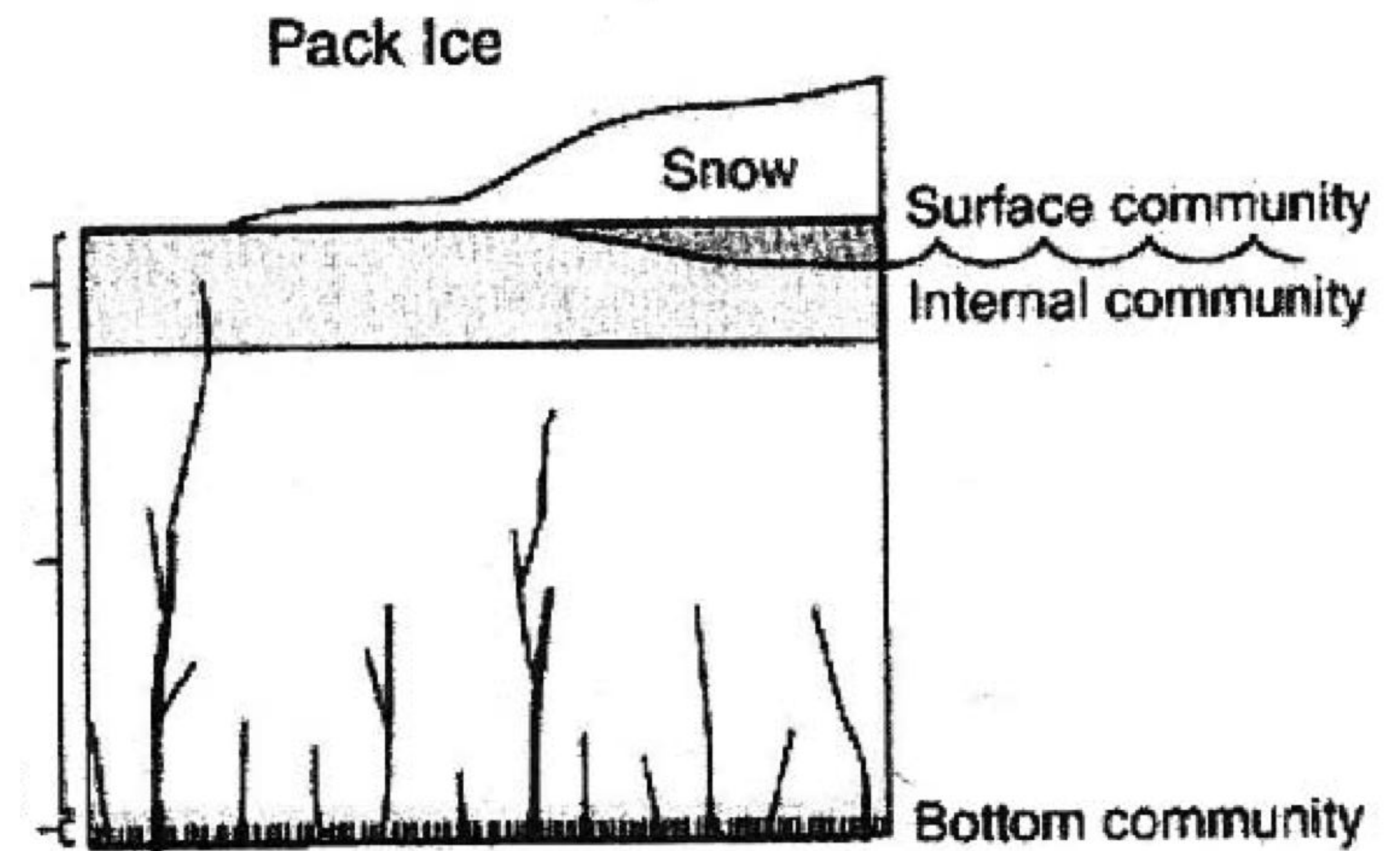
*Richard Cullather*

# Summer Bloom In the Ross Sea

MODIS image NASA's Aqua



<u>Ice Algae</u>	<u>Annual Prod.</u> (10 <sup>11</sup> g Chla y <sup>-1</sup> )
Surface	5
Freeboard	2
Interior	0.2
<u>Bottom</u>	<u>0.2-1</u>



(Arrigo et al. 2004)

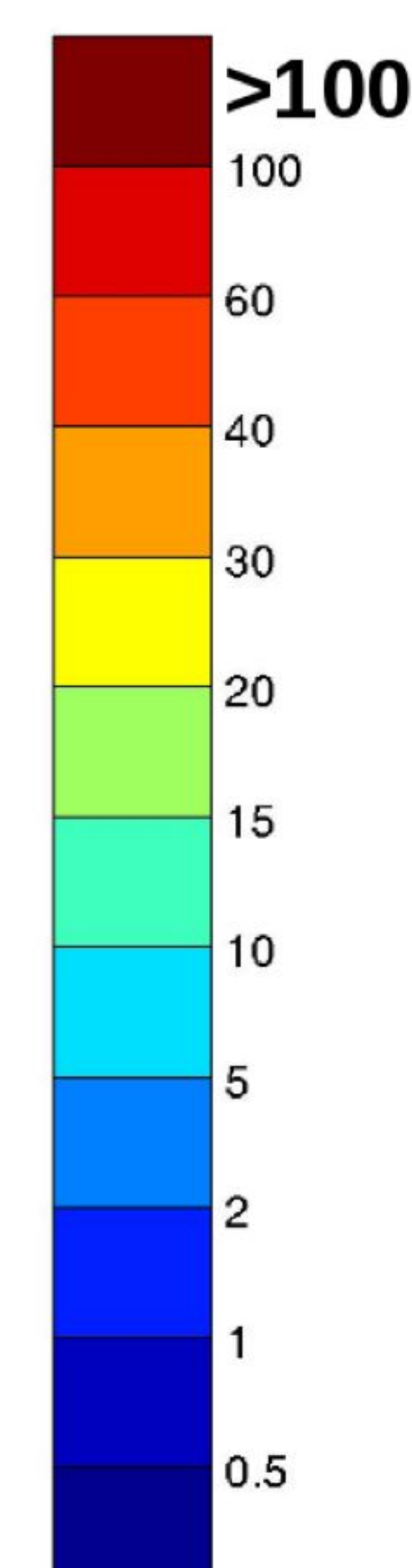
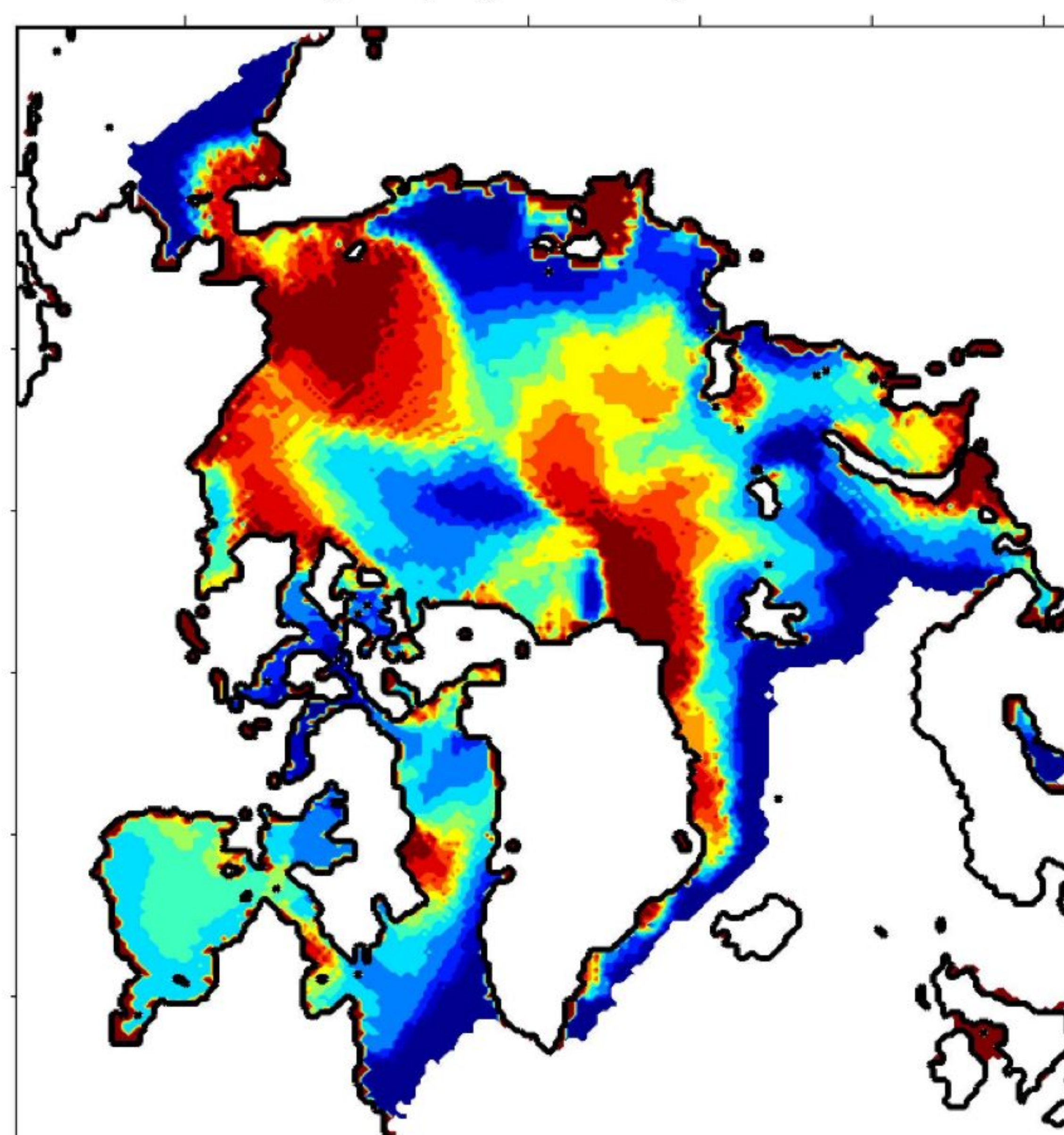
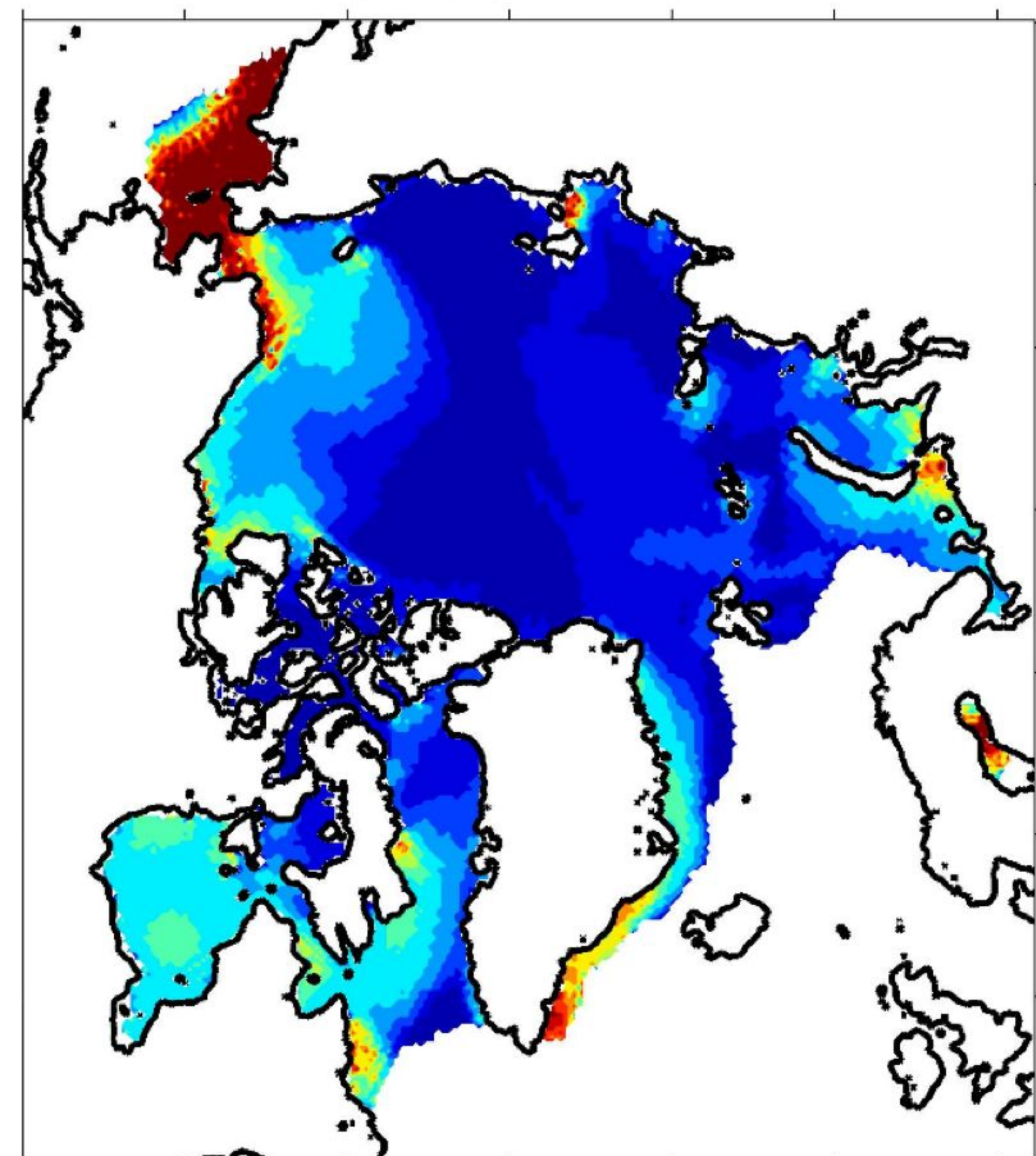
Ice Edge Dynamics



# Skeletal Layer Model

Ice Algae (mg Chl m<sup>-2</sup>) 04-13

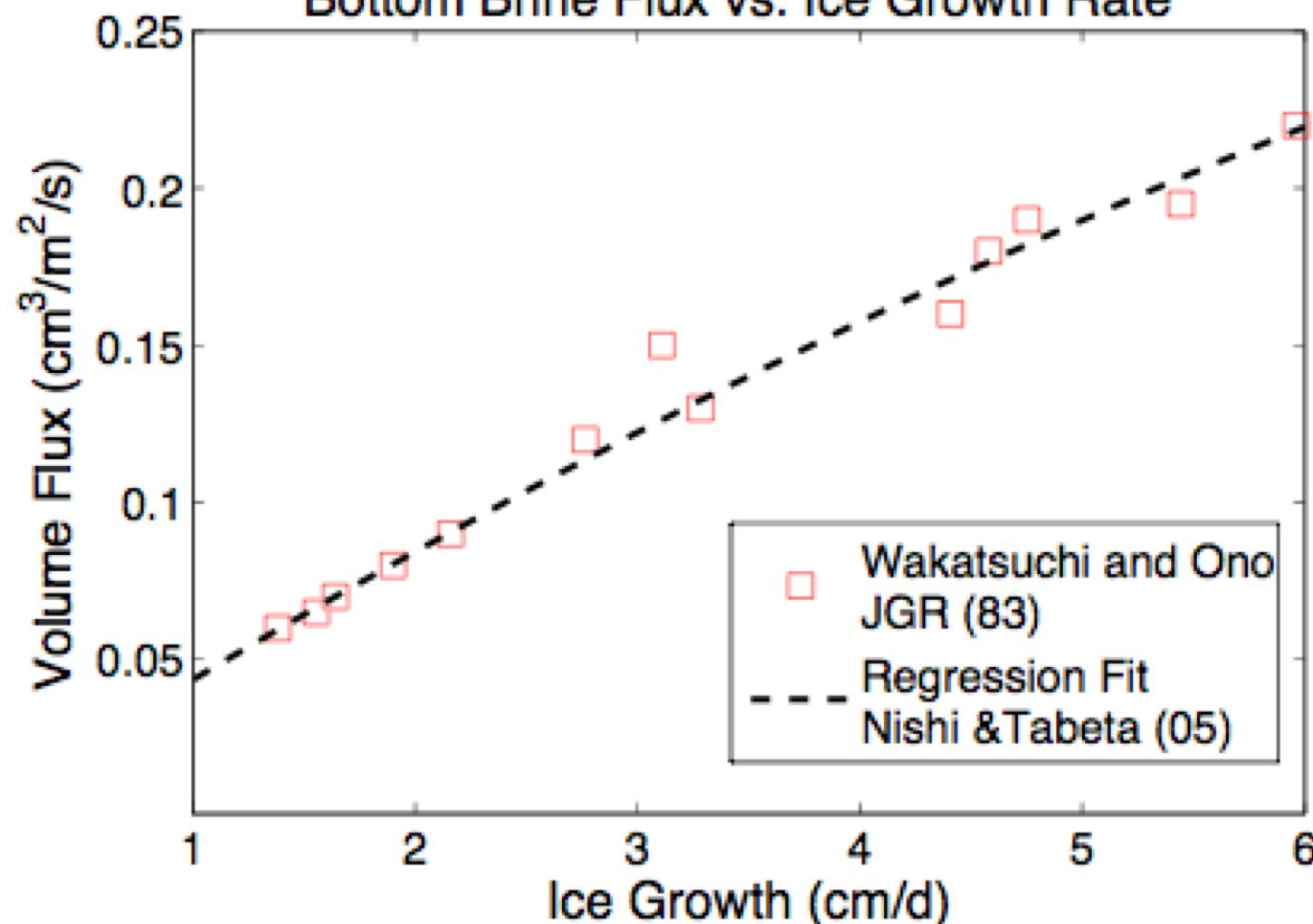
Ice Algae (mg Chl m<sup>-2</sup>) on 05-18



*Deal, Jin and Elliott*

- Ice thickness (PAR)
- Ice extent (2 cm skeletal layer)
- Ocean nutrient climatology
- Biogeochemical reactions
- Brine driven bottom flux

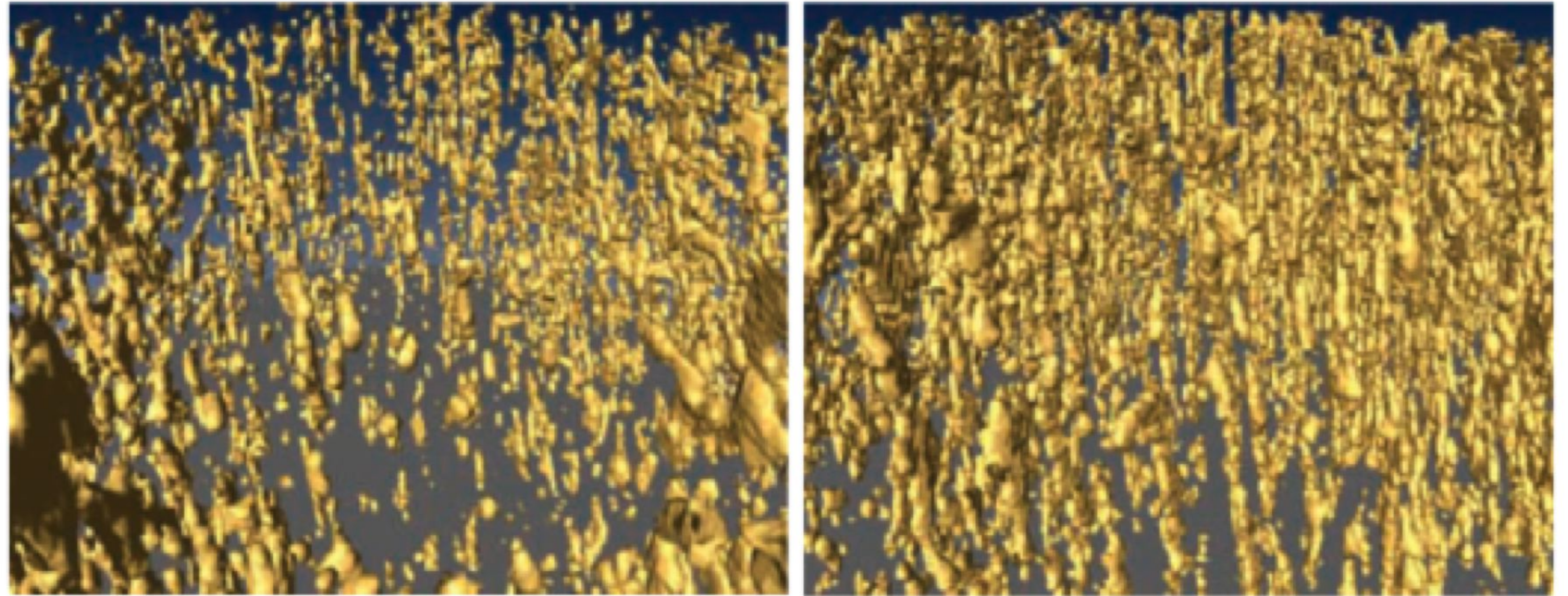
Bottom Brine Flux vs. Ice Growth Rate



# Multi-layer Ice Algal Model

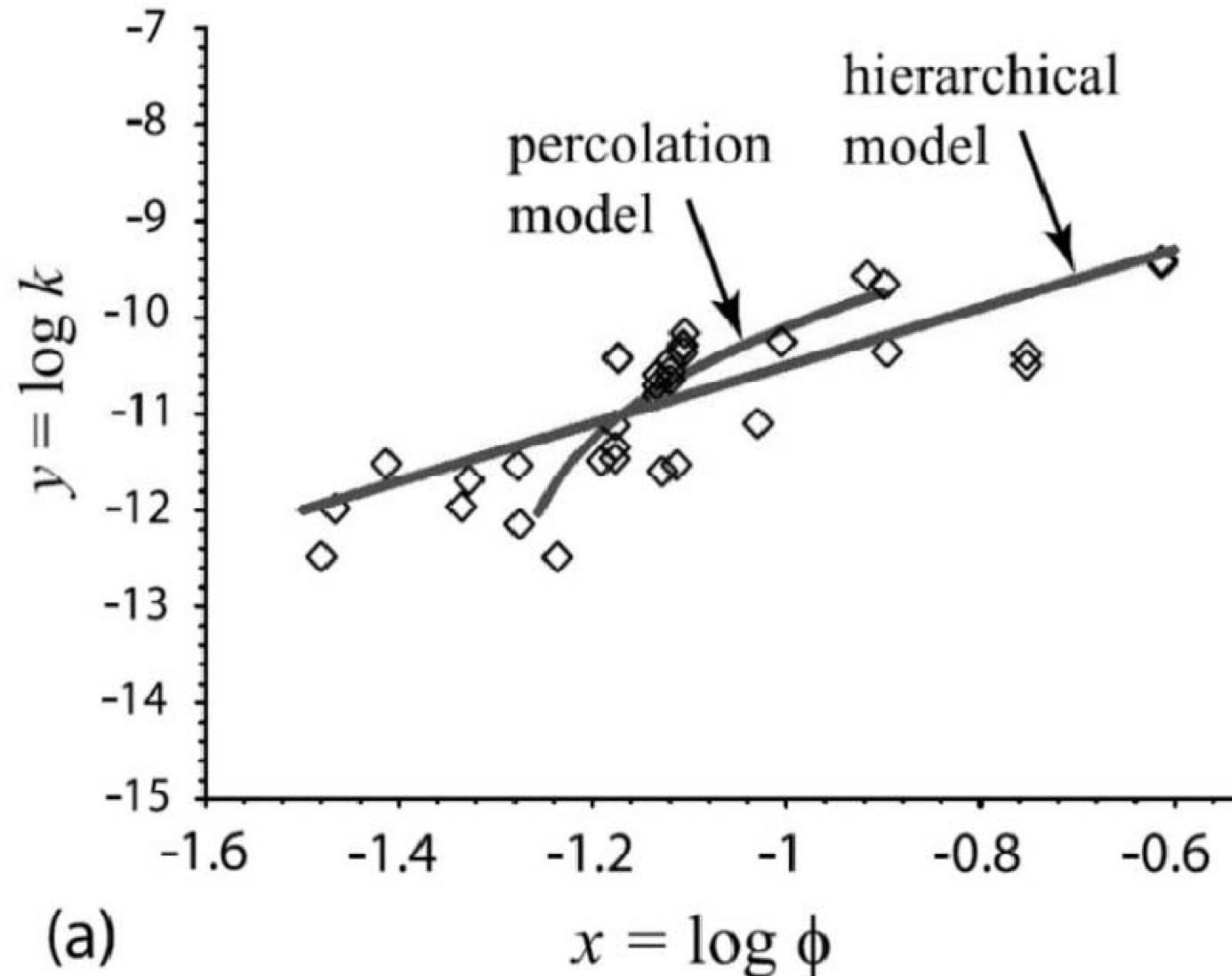
GOLDEN ET AL.: PERMEABILITY AND MICROSTRUCTURE IN SEA ICE

- Sea ice porosity ( $\phi$ ) and permeability ( $k$ ) as functions of  $T$  and  $S$



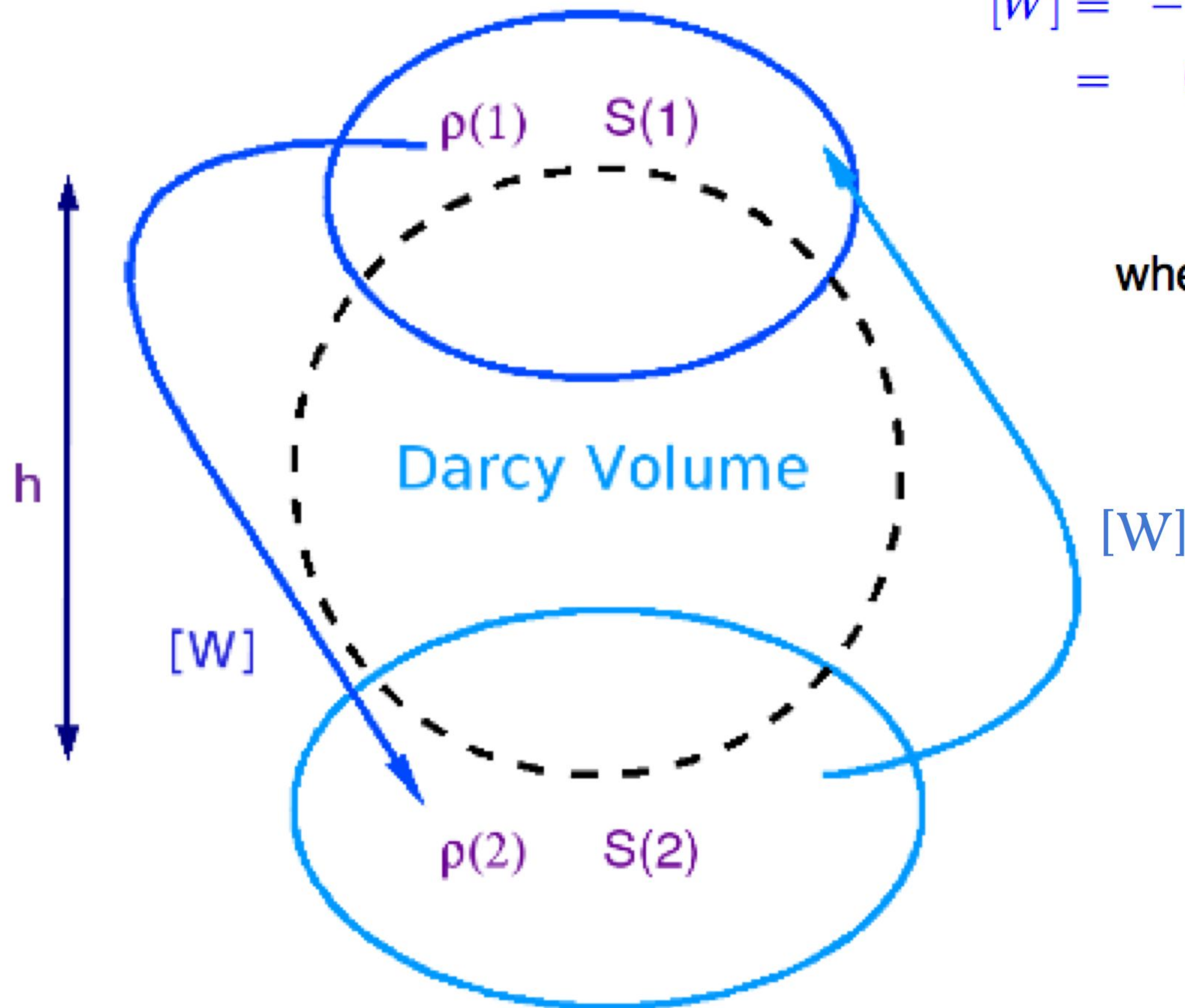
$\phi = 0.033$

$\phi = 0.075$



- Gravitationally unstable Brine motion: brine density as function of  $T$
- Nutrient transport model based are Darcy's Flow and Mixing length theory

# Mixing Length Theory



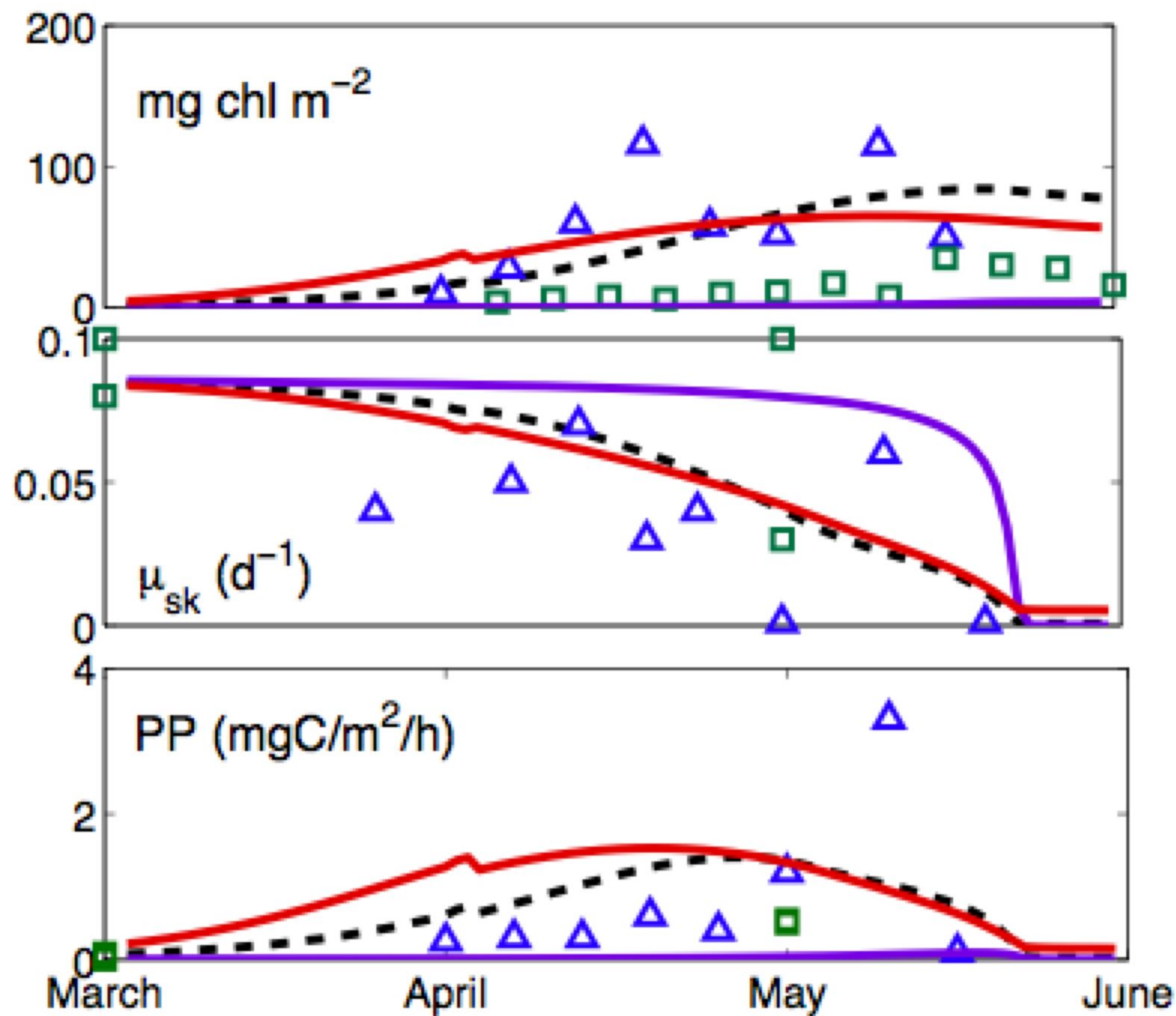
$$[W] = -\frac{kg}{\mu} [\Delta\rho]^i \quad \text{for **unstable** profiles}$$

$$= 0 \quad \text{for **stable** profiles}$$

where  $[\Delta\rho]^i = \frac{d[\rho]^i}{dz} h$

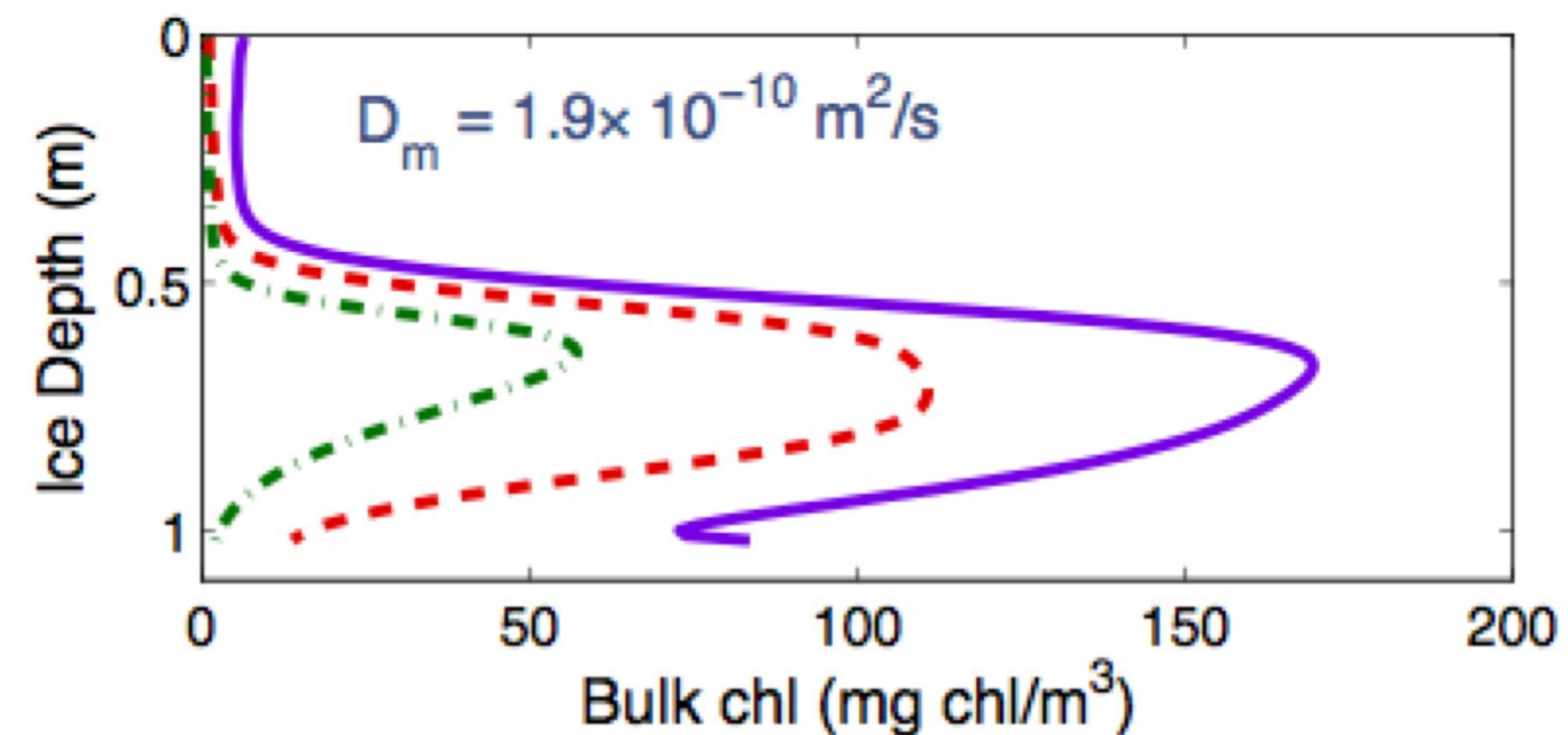
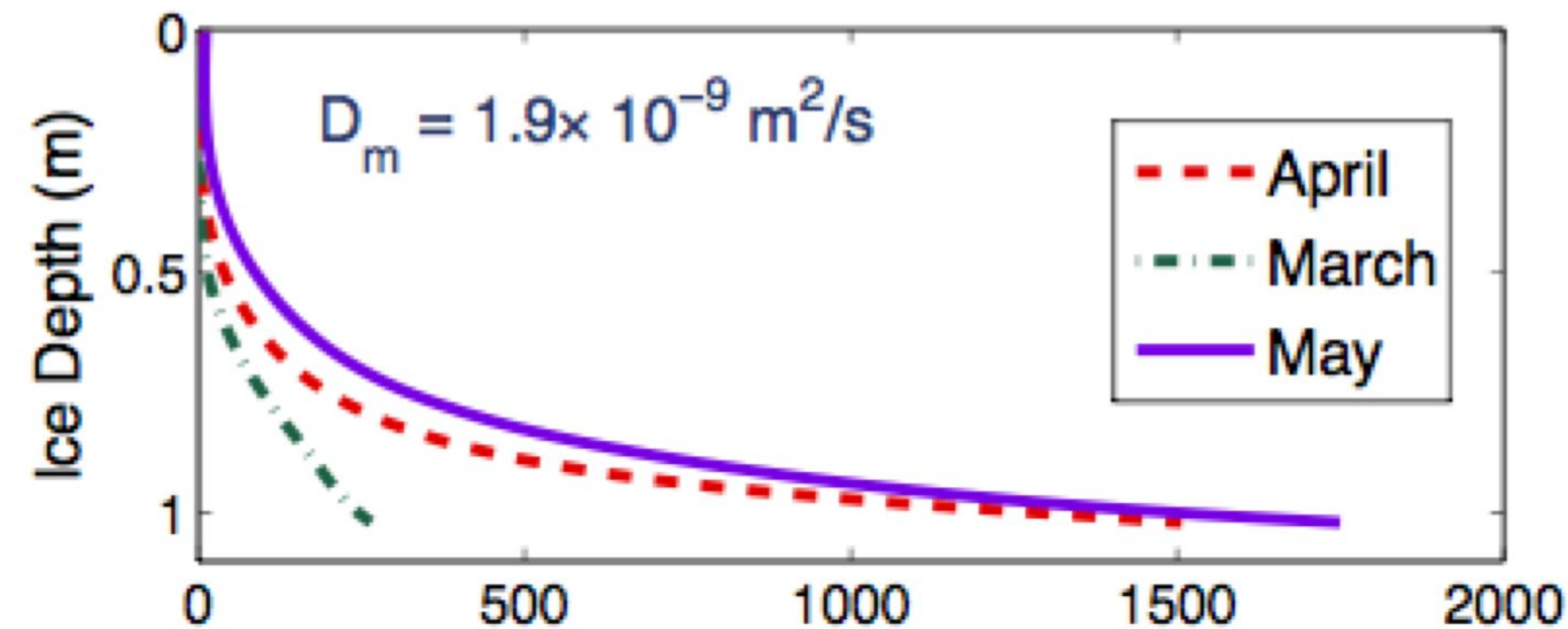
Advection  $\frac{\partial}{\partial z} (|[W]| \Delta[S]^i) \rightarrow \frac{\partial}{\partial z} \left( |[W]| h \frac{\partial [S]^i}{\partial z} \right)$  Diffusion

# Model Sensitivity: Example, $D_m$



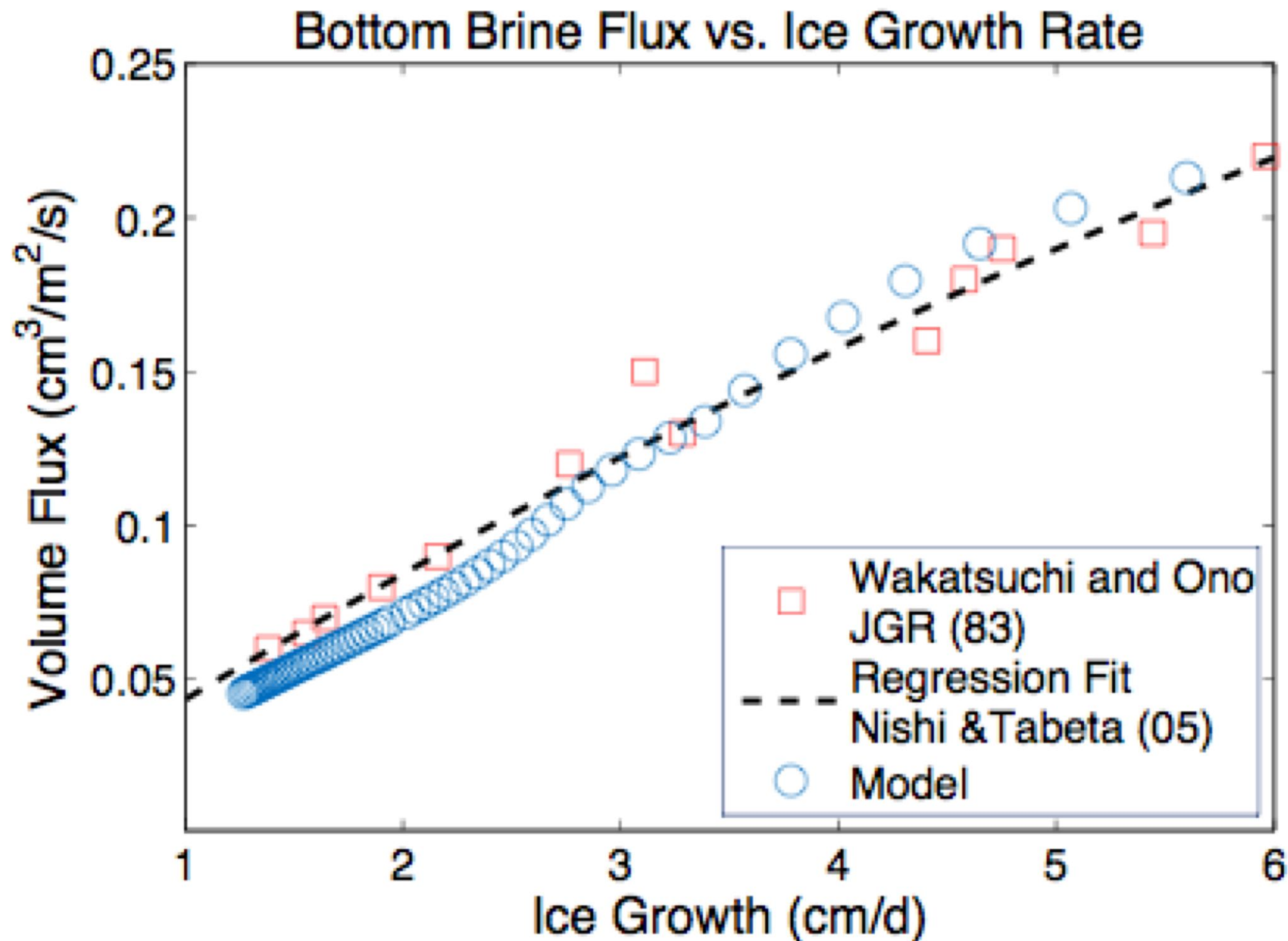
- $D_m = 1.9 \times 10^{-8} \text{ m}^2/\text{s}$
- $D_m = 1.9 \times 10^{-10} \text{ m}^2/\text{s}$
- -  $D_m = 1.9 \times 10^{-9} \text{ m}^2/\text{s}$   
Standard run

- Smooth polynomial forcing of  $T_{air}$ , and PAR
- -- Standard run:  
 $D_m = 1.9 \times 10^{-9} \text{ m}^2/\text{s}$
- Skeletal layer integration (5cm)
- $D_m$  varied 2 orders of magnitude





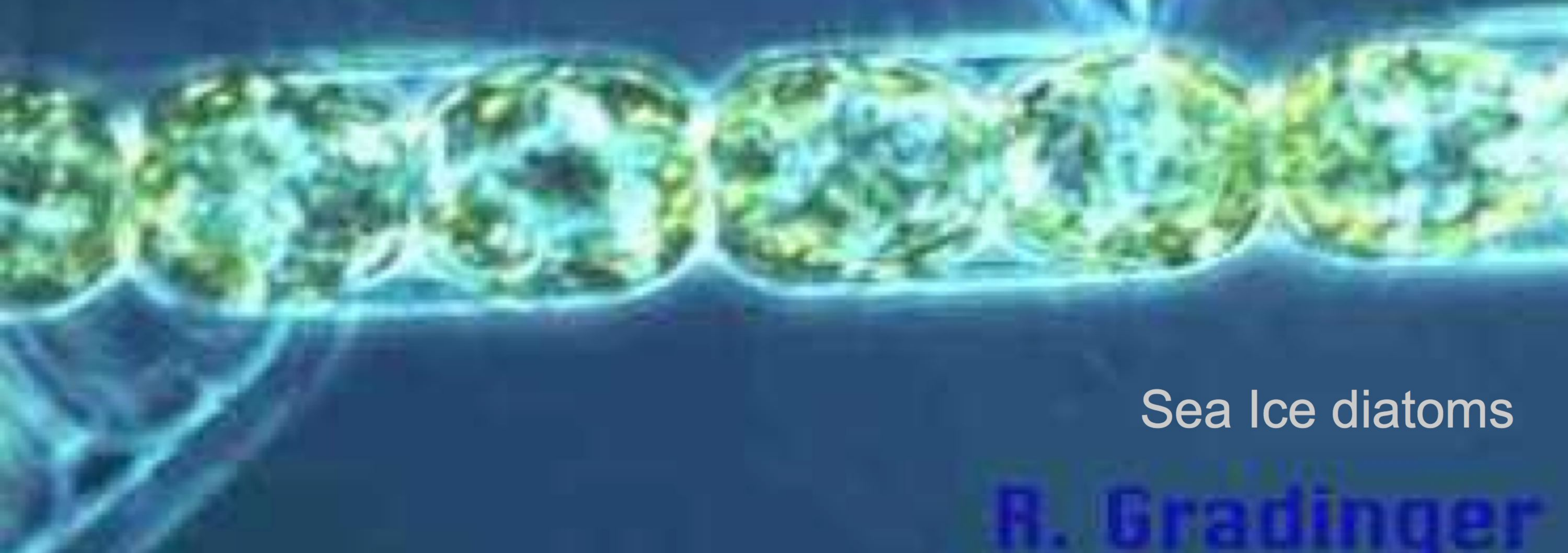
# Model Support: Is $D \propto$ Brine Flux?



- A crude comparison, but possibly
- Thickness dependence?
- Multi-layer model is consistent with skeletal layer flux assumptions
- Results are encouraging

# Continuing Research

- Sensitivity studies and data comparisons
- Include Darcy velocity driven by non-hydrostatic pressure gradients. Eg, snow loading, melt-ponds
- Prepare code for global simulations



Sea Ice diatoms

R. Gradinger