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Developing CISM: The Community Ice Sheet Model

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University of Montana

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| Introducing CISM | Higher-order physics | Software | |
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| Outline | | | |

1 Introducing CISM

2 Higher-order physics

3 Software

4 Verification

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Model pedigree

- Glimmer: thermomechanical shallow ice approximation (SIA) with sliding ¹
 - Excellent starting framework
 - Native NetCDF support
 - Couples to climate drivers
- First-order diagnostic model from Pattyn ²
- First-order diagnostic and prognostic model from Payne and Price
- Additional improvements: basal water, climate drivers, sparse solver packages, and more...

¹I. Rutt *et. al.*, J. Geophysical Research, 2009

²F. Pattyn, J. Geophysical Research, 2003

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Improve access

A graphical interface to the model



Assemble and redistribture modeling data

as well as climate drivers, and visualization scripts.



Model wiki pages

facilitating collaborative content generation



3. understand the outcome of ice sheet modeling experiments.

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Advantages of Pattyn model

- Three-dimensional
- Resolves transition zone between deformational flow and sliding
- Simplifying assumptions compared to full Stokes remove need to solve vertical velocity and pressure



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Why multiple models?

- Helped solve common problems sped development
- Diversity in computational methods
- For intercomparison: use a common framework
- As examples: show how other contributions can integrate (common signature)

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Conservation equations

Incompressible conservation of mass:

$$abla \cdot \mathbf{v} = \mathbf{0}$$

Conservation of momentum:

$$\rho_i \frac{d\mathbf{v}}{dt} = \nabla \cdot \mathbf{T} + \rho_i \mathbf{g}$$

Neglecting acceleration term:

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Constitutive relations

Full constituative equations (i.e. Full Stokes)

$$\mathbf{T} - \rho \mathbf{I} = 2\mu \begin{pmatrix} \frac{\partial u}{\partial x} & \frac{1}{2} (\frac{\partial u}{\partial y} + \frac{\partial v}{\partial x}) & \frac{1}{2} (\frac{\partial u}{\partial z} + \frac{\partial w}{\partial x}) \\ \frac{1}{2} (\frac{\partial v}{\partial x} + \frac{\partial u}{\partial y}) & \frac{\partial v}{\partial y} & \frac{1}{2} (\frac{\partial v}{\partial z} + \frac{\partial w}{\partial y}) \\ \frac{1}{2} (\frac{\partial w}{\partial x} + \frac{\partial u}{\partial z}) & \frac{1}{2} (\frac{\partial w}{\partial y} + \frac{\partial v}{\partial z}) & \frac{\partial w}{\partial z} \end{pmatrix}$$

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Constitutive relations

Full constituative equations (i.e. Full Stokes)



Glen-type flow law

$$\mu = \frac{A^{\frac{-1}{n}}}{2} \dot{\epsilon}^{\frac{1-n}{n}}$$

Constitutive relations

Full constituative relation

$$\mathbf{T} - p\mathbf{I} = 2\mu \begin{pmatrix} \frac{\partial u}{\partial x} & \frac{1}{2} \left(\frac{\partial u}{\partial y} + \frac{\partial v}{\partial x} \right) & \frac{1}{2} \left(\frac{\partial u}{\partial z} + \frac{\partial w}{\partial x} \right) \\ \frac{1}{2} \left(\frac{\partial v}{\partial x} + \frac{\partial u}{\partial y} \right) & \frac{\partial v}{\partial y} & \frac{1}{2} \left(\frac{\partial v}{\partial z} + \frac{\partial w}{\partial y} \right) \\ \frac{1}{2} \left(\frac{\partial w}{\partial x} + \frac{\partial u}{\partial z} \right) & \frac{1}{2} \left(\frac{\partial w}{\partial y} + \frac{\partial v}{\partial z} \right) & \frac{\partial w}{\partial z} \end{pmatrix}$$

Reduced constituative relation

$$\mathbf{T} - p\mathbf{I} = 2\mu \begin{pmatrix} \frac{\partial u}{\partial x} & \frac{1}{2} \left(\frac{\partial u}{\partial y} + \frac{\partial v}{\partial x} \right) & \frac{1}{2} \frac{\partial u}{\partial z} \\ \frac{1}{2} \left(\frac{\partial v}{\partial x} + \frac{\partial u}{\partial y} \right) & \frac{\partial v}{\partial y} & \frac{1}{2} \frac{\partial v}{\partial z} \\ \frac{1}{2} \frac{\partial u}{\partial z} & \frac{1}{2} \frac{\partial v}{\partial z} & -\frac{\partial u}{\partial x} - \frac{\partial v}{\partial y} \end{pmatrix}$$

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Boundary conditions

Stress-free surface

$$\mathbf{T} \cdot \mathbf{\hat{n}_s} = 0$$

Linear bed strength

$$\beta^2 \mathbf{\hat{t}} \cdot \mathbf{v} = \mathbf{\hat{t}} \cdot (\mathbf{T} \mathbf{\hat{n}}_{\mathbf{b}}) = \tau_b$$

Ice shelf front (vertically averaged)

$$\mathbf{T} \cdot \hat{\mathbf{n}} = \frac{1}{2} \rho_i g H \left(1 - \frac{\rho_i}{\rho_w} \right) \hat{\mathbf{n}}$$

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Challenges integrating Pattyn model

Data incompatability problems:

 Transposed coordinate ordering relative to Glimmer

$$(z, x, y) \rightarrow (y, x, z)$$

Co-located grid versus staggered grid



Adapted from Glimmer documentation

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Solution:

Build a facade around Pattyn's code that transforms data

Can run on either the ice grid or the velocity grid

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- Build a facade around Pattyn's code that transforms data
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Integration software design

Primary concern: Don't "clobber" SIA operations!

- Higher-order velocity fields placed in parallel data structure to SIA velocity fields
- Facade and HO dynamic core reside in seperate modules
- Enabled via configuration file option, off by default:

Enabling higher-order computation

```
[ho_options]
diagnostic_scheme = 1
```

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Point type mask



Shelf front descritization

Ice Front Normal Vector



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CISM as a verification platform

- Python test infrastructure
- Generates test data outside of CISM in NetCDF format
- Fully automates test setup, run, interpretation

Scripted ISMIP-HOM example

python ISMIP-HOM/verify.py -abc --20km --40km

Has been used as a partial regression suite

ISMIP-HOM A



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ISMIP-HOM C



3.1

EISMINT-Ross: Velocity Map



Discrete points show the observed velocity ³ at RIGGS stations

³Thomas *et. al.* 1984

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EISMINT-Ross: Comparison to RIGGS stations



| | | Max vel. |
|--------------|-------|----------|
| Model | X^2 | (m/a) |
| Bremerhaven1 | 3605 | 1379 |
| Bremerhaven2 | 12518 | 1663 |
| Chicago1 | 5114 | 1497 |
| Chicago2 | 5125 | 1497 |
| Grenoble | 5237 | 1508 |
| Missoula | 4962 | 1495 |

D. MacAyeal, Ann. Glaciology, 1996

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Conclusion

- Higher-order models can be integrated into CISM
- New framework, infrastructure support parallel efforts
- Intercomparison efforts promising
- Main bottleneck, elliptical equation solve, is scalable

Future Work:

- More diversity in dynamic cores
- Prognostic solver integration
- Inverse modeling and coupled shelf/stream solve