# Improvements in CAM4 : Moist Turbulence, Shallow Convection, and Cloud Macrophysics

CSM Meeting at Breckenridge

Jun. 16. 2009

Sungsu Park, Chris Bretherton, and Phil Rasch

# CGD.NCAR

University of Washington, Seattle, Washington Pacific Northwest National Laboratory, Richland, Washington

# DEFINITIONS

## Model Names

- CAM3.5 : CAM3.0 + Revised Deep Convection + etc. ( = Track I )
- CAM4 : CAM3.5 + All New Atmospheric Physics ( = Track V )

## Variables

- a : Cloud Fraction
- LCA : Low Cloud Amount
- TCA : Total Cloud Amount
- **S** : Lower-Tropospheric Stability,  $S \equiv \theta_{v}(700) \theta_{v}(1000)$

# **Physical Processes in CAM4**



**MOIST TURBULENCE SCHEME in CAM4** 

Ri

SI

STI

ΕI

ΤI

CL

S

е

W<sub>e</sub>



$$\frac{\partial \overline{A}}{\partial t} = -\frac{\partial}{\partial z} \overline{w'A'} = \frac{\partial}{\partial z} \left( K \frac{\partial \overline{A}}{\partial z} \right)$$

### *K* : eddy diffusivity

: Moist Richardson Number : Stable Interface : Stably Turbulent Interface : Entrainment Interface : Turbulent Interface : Stably Turbulent Layer STL : Convective Layer : Turbulent length scale : Stability function (fcn of Ri) : TKE : Entrainment rate

# Moist Turbulence Scheme in CAM4

- Diagnostic TKE-based 1<sup>st</sup> order K diffusion scheme with entrainment param.
  - Numerically stable, physically realistic, conceptually clear
  - TKE is fed into 'shallow convection' and 'cloud microphysics', and regulates the onset of cumulus updraft and cloud droplet activation

## Stratus-Top LW Cooling and In-Stratus Condensation Heating into TKE

- Sensitive to 'cloud macro-microphysics' and 'radiation' schemes
- Treatment of Stratus-Radiation-Turbulence Interactions
- Now, stratus is a dynamic (as well as radiative) driver of the climate
- Handling of the 2<sup>nd</sup> aerosol indirect effect
- Removal of the stability-based KH stratus fraction
- Activate in any layers above as well as within PBL
  - Simulate turbulences in the mid- and upper-level clouds
- Compared to CAM35 PBL scheme,
  - Much better performance in cloud-topped regime
  - Similar or superior performance in dry stable and convective regimes

# **Cloud-Radiation-Turbulence Interactions**



# Low Cloud Amount. JJA.



KH's stratus fraction

# Fog Amount. JJA.



# SHALLOW CONVECTION SCHEME in CAM4



 $w'A' = \rho \cdot M_u \cdot (A_u - \overline{A})$ 

 $M_u$ : updraft mass flux  $A_u$ : updraft scalar

IN	: Convective INhibition
CL	: Lifting Condensation Level
FC	: Level of Free Convection
NB	: Level of Neutral Buoyancy
/ <sub>u</sub>	: Updraft vertical velocity
u	: Updraft fractional area

C

พ a

# Shallow Convection Scheme in CAM4

- An entraining-detraining buoyancy-sorting updraft plume with a penetrative entrainment parameterization
  - Mass flux closure based on TKE and Convective Inhibition (CIN)
  - Close interactions with moist turbulence scheme
  - Transports momentum and aerosols as well as thermodynamic conservative scalars
  - Computes cumulus fraction and LWC, vertical velocity, updraft mass flux
  - Direct influence on the global radiation budget
- Much less sensitive to vertical resolution than CAM35
- Can simulate deep as well as shallow convective activity
- Simulate the '*real*' convective activity

#### Shallow Convective Mass Flux at Cloud Base. Annual.

### CAM4

### CAM35



### Inconsistency between 'Stratus Fraction' and 'In-Stratus LWC' in CAM35



- $\rightarrow$  distorts LW cooling profile
- $\rightarrow$  too strong inversion at the PBL top
- $\rightarrow$  too weak entrainment rate
- $\rightarrow$  too shallow and moist PBL

# Macrophysics Scheme in CAM4

- Enhance consistency between stratus fraction and in-stratus LWC
- Remove 'empty' (a>0, q<sub>l,cloud</sub>=0) and 'dense' (a=0, q<sub>l,cloud</sub>>0) stratus
- Uses a single equilibrium stratus fraction at each time step
- Liquid stratus fraction based on triangular PDF of q<sub>t</sub>
- Removal of KH's stability based stratus fraction
- Separate treatment of liquid condensation and ice sublimation
- Separate diagnose of liquid and ice stratus fractions
- Liquid condensation formula based on conservative scalars
- Cumulus is non-overlapped with stratus in each layer.
- Cumulus has its own in-cumulus LWC.
- Cumulus is radiatively active.

# Horizontal Geometry of Clouds in CAM



## Improvements of Cloud Treatment in CAM4

- Simulation of 'Interactive Cloud Droplet Number' as well as 'LWC/IWC'

   New 2-Moment Microphysics and Modal Aerosol Model
- More Realistic Radiative Properties of Clouds

   New Cloud Optics

# Simulation Results: Observation vs CAM35 vs CAM4

Observation : 42-yrs (1956-1997) EECRA ship-observations, NCEP/NCAR Reanalysis 17-yrs (1984-200) ISCCP satellite-derived radiation at surface

CAM35 : 92-yrs coupled simulation using pre-industrial GHG and aerosols

CAM4 : 87-yrs coupled simulation using pre-industrial GHG and aerosols

Interannual Correlation between  $S \equiv \theta_{\nu}(700) - \theta_{\nu}(1000)$  and Low Cloud Amount. JJA.



Line: Ship-observed LCA

### ENSO Regression Anomalies of Total Cloud Amount [%]. JAS.



# Normalized Covariance of the 1<sup>st</sup> Coupled Mode from the SVD Analysis over the North Pacific



# SVD Heterogeneous Map. SST vs TCA. JAS.

SST

TCA













# SW Surface Heat Flux Feedback $\lambda_{SW} \equiv -\partial Q_{SW}^{\downarrow} / \partial SST$ . JJA.

60N





## LW Surface Heat Flux Feedback $\lambda_{LW} \equiv -\partial Q_{LW}^{\downarrow} / \partial SST$ . JJA.



Line: Ship-observed Large Cumulus Frequency

[Wm<sup>-2</sup>K<sup>-1</sup>]

# Surface Heat Flux Feedback over the North Pacific Ocean



# **SUMMARY**

- CAM4 has much better physics and interactions among the physics than CAM35, without arbitrary kludges (e.g., stability based LCA etc.).
- Our analysis also showed that the overall practical performance of CAM4 is similar or better than CAM35.
- CAM4 can simulate many important features in a physically reasonable way, especially the ones associated with cloud processes themselves and cloudclimate interactions (e.g., marine stratocumulus clouds, cumulus, cloud-SST interaction, cloud-sea ice interaction, 1<sup>st</sup> and 2<sup>nd</sup> aerosol indirect effects, etc.).
- Some important biases in CAM4:
  - Biases common both in CAM35 and CAM4 : moist atmosphere, weak LW
     CRF (?) and LW radiative feedback
  - Biases in CAM4 : small sea-ice fraction over the Arctic in summer

## Sensitivity to Vertical Cloud-Overlapping Structure



## SW Surface Heat Flux Feedback $\lambda_{SW} \equiv -\partial Q_{SW}^{\downarrow} / \partial SST$ . JJA.

CAM35



Weaker SW feedback in summer Arctic in CAM35, probably due to the built-in negative feedback between sea ice and stability-based stratus fraction,

may explain more sea ice extent in CAM35 than in CAM4.

-40 -32 -24 -16 -8 -2 2 8 16 24 32 40

150W 120W

60W

90W

30W

30E

[Wm<sup>2</sup>K<sup>1</sup>]

30E

120E

150E

180

# The End of Presentation



Evaporation of Precipitation



LW Radiative Heating Rate - ⊟ - CAM4 





# Macrophysics Scheme in CAM4

- Uses a single equilibrium cloud fraction at each time step.
- Condensation formulation based on conservative scalars
- Remove 'empty' (a>0, q<sub>l,cloud</sub>=0) and 'dense'(a=0, q<sub>l,cloud</sub>>0) stratus
- Explicit treatment of in-cumulus LWC



# Interplay among Various Processes in Stratocumulus

Large-Scale Subsidence





Evaporation of Precipitation



LW Radiative Heating Rate - ⊟ - CAM4 





**MOIST TURBULENCE SCHEME in CAM4** 

Ri

SI

STI

ΕI

ΤI

CL

S

е

W<sub>e</sub>



$$\frac{\partial \overline{A}}{\partial t} = -\frac{\partial}{\partial z} \overline{w'A'} = \frac{\partial}{\partial z} \left( K \frac{\partial \overline{A}}{\partial z} \right)$$

### *K* : eddy diffusivity

: Moist Richardson Number : Stable Interface : Stably Turbulent Interface : Entrainment Interface : Turbulent Interface : Stably Turbulent Layer STL : Convective Layer : Turbulent length scale : Stability function (fcn of Ri) : TKE : Entrainment rate

### Interannual Correlation between $S \equiv \theta_{\nu}(700) - \theta_{\nu}(1000)$ and Low Cloud Amount. DJF.



### ENSO Regression Anomalies of Total Cloud Amount [%]. DJF.



# 3 Cloud Types in CAM3.5

## Cumulus

 $a_c = f(M)$ , M: Convective Updraft Mass Flux

## • RH (Relative Humidity) Stratus

 $a_{s,RH} = f(RH)$ , RH: Grid-Mean Relative Humidity

(Klein-Hartmann) Stratus

 $a_{s,KH} = f(S)$ ,  $S \equiv \theta_{v}(700) - \theta_{v}(1000)$ 

**Computation of Liquid Stratus Fraction** 

### PDF of q<sub>t</sub> for liquid cloud only

### Stratus Fraction as a function of RH



