

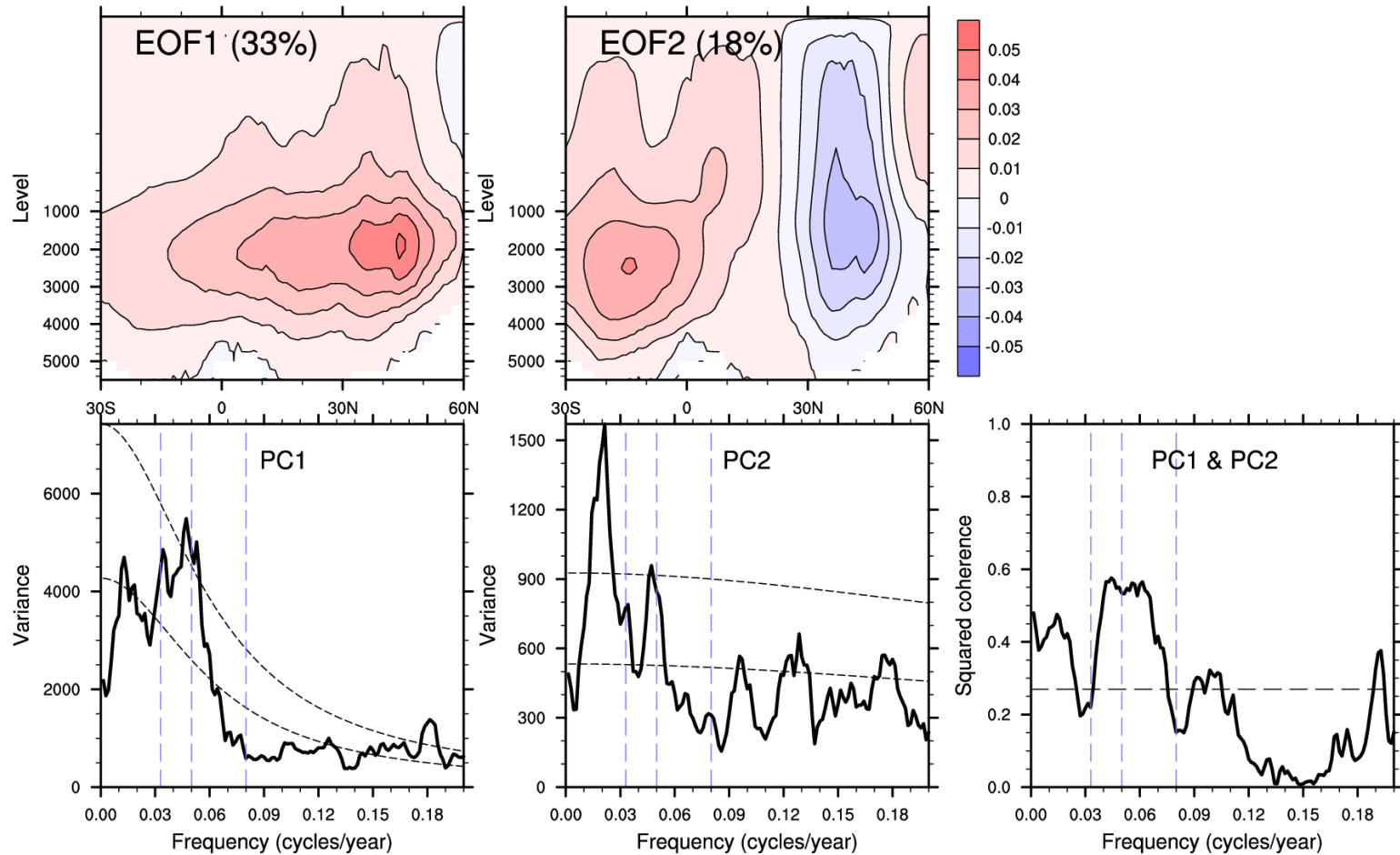
Predictability of Atlantic Overturning Circulation and Associated Surface Patterns in Two CCSM3 Large Ensemble Experiments

**Haiyan Teng
Grant Branstator
Jerry Meehl**

NCAR/CGD

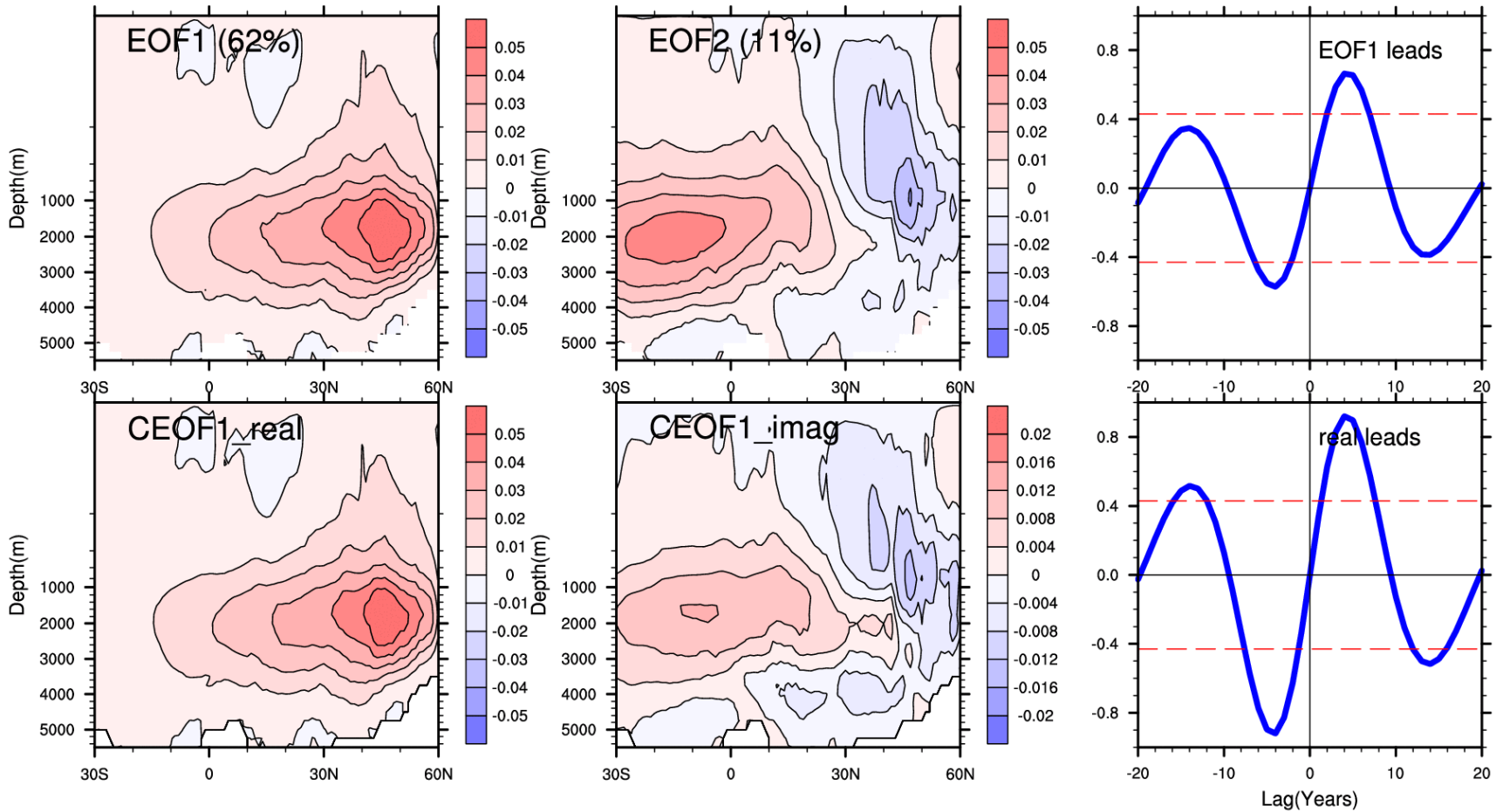
July 1st, 2010

AMOC EOF1 & EOF2

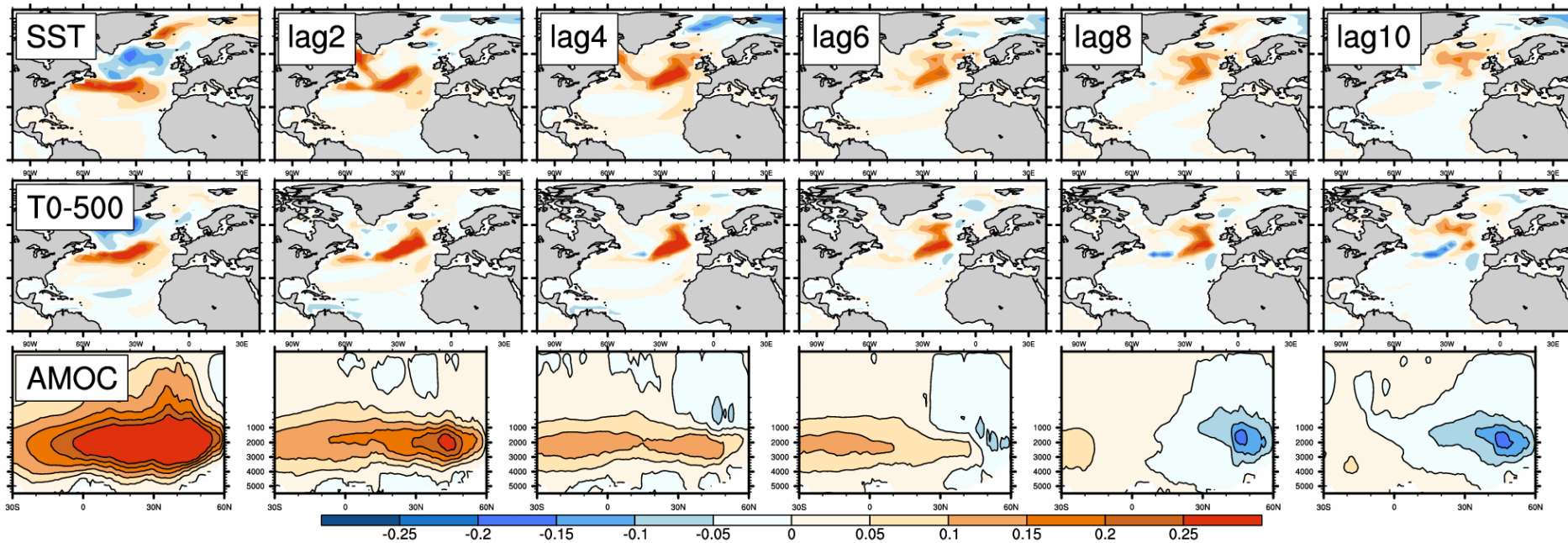


From year 300-999 of the present-day control

10-30-yr AMOC EOF1,2 & CEOF1



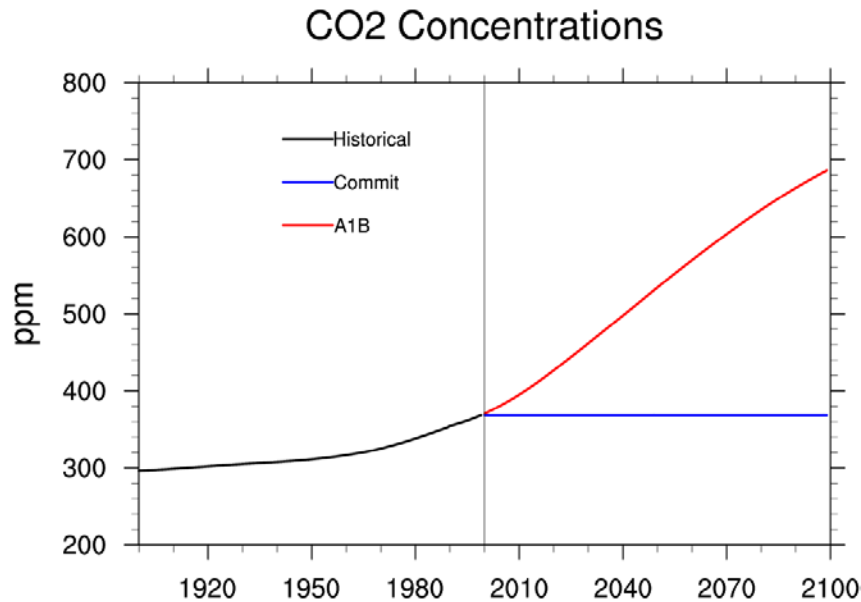
Lag regression to AMOC PC1



Tropical SST anomalies are weak!

SST 1-2-1 smoothed

Model & Experiments



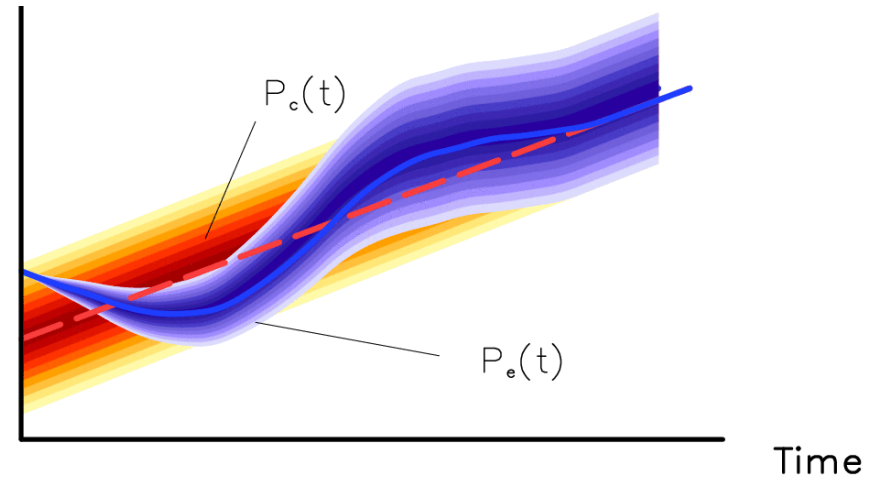
- **CCSM3.0, T42 atm, 1deg ocn**
- **40-member ensemble run**
 - **A1B, Commitment**
 - **2000-2061**
 - **atm perturbed ICs**
 - **Same ocn/Ind/ice ICs**
- **A1B (II), A1B(III), A1B (IV)**
 - **Each has unique ocn/ice/Ind ICs**
 - **20 years**
- **Last 700 yrs of the 1000-yr control**

Relative Entropy

Total predictability: $P_e(t)$ vs. $P_c(0)$

Initial-value predictability: $P_e(t)$ vs. $P_c(t)$

Forced predictability: $P_c(t)$ vs. $P_c(0)$



Kleeman (2002)

$$R = \int_S P_x(s) \log_2 \left[\frac{P_x(s)}{P_b(s)} \right] ds$$

$P_x(s)$: prediction
 $P_b(s)$: base

For normal distribution:

$$R = \frac{1}{2} \log_2(e) \left\{ \ln \left[\frac{\det(\sigma_b^2)}{\det(\sigma_x^2)} \right] + \text{trace} \left(\frac{\sigma_x^2}{\sigma_b^2} \right) + (\mu_x - \mu_b)^T (\sigma_b^2)^{-1} (\mu_x - \mu_b) - n \right\}$$

dispersion

signal

Initial-value Predictability & Forced Predictability

Assumptions:

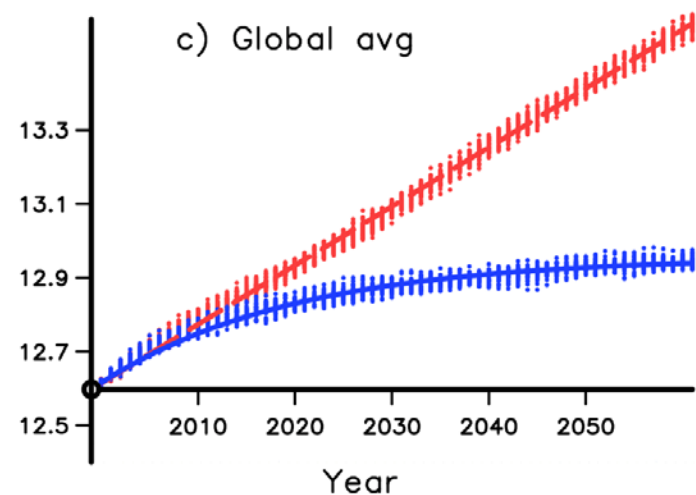
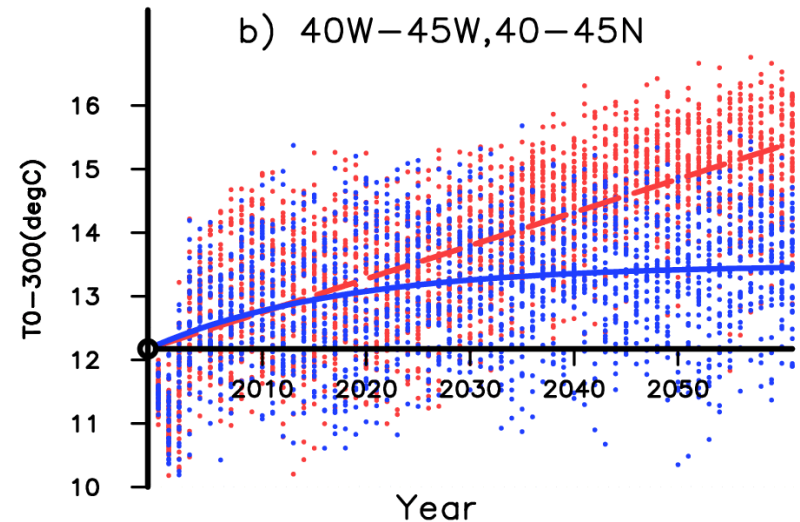
1. Gaussians
2. Climate covariance does not change with time.
3. Climate mean can be well approximated by an analytical function of time.

Time evolving climate mean:

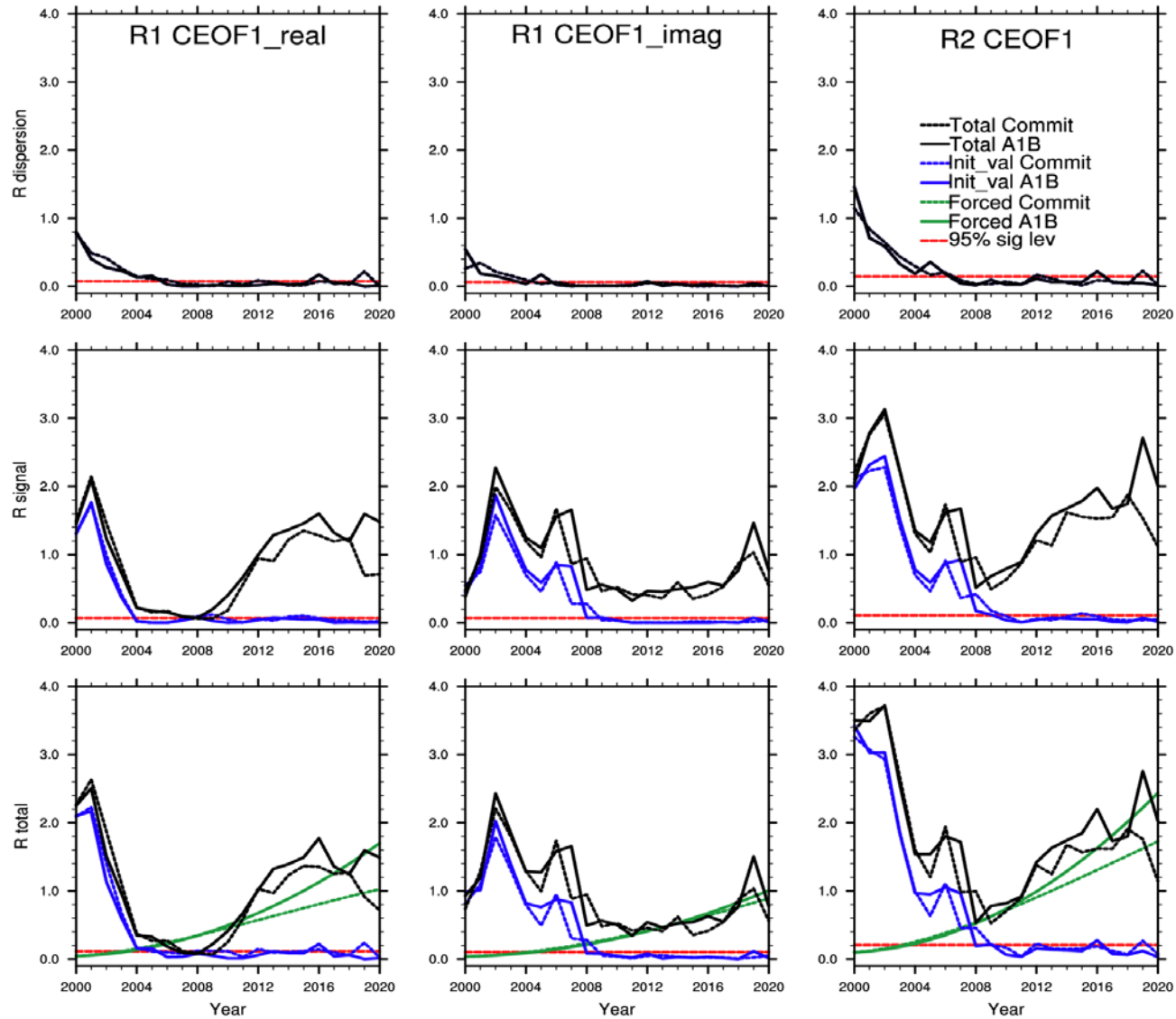
$$\bar{T}_{A1B}(t) = \bar{T}_{1999} + k(t - 1999)$$

$$\bar{T}_{commit}(t) = \bar{T}_{1999} + A(1 - e^{-t/\tau})$$

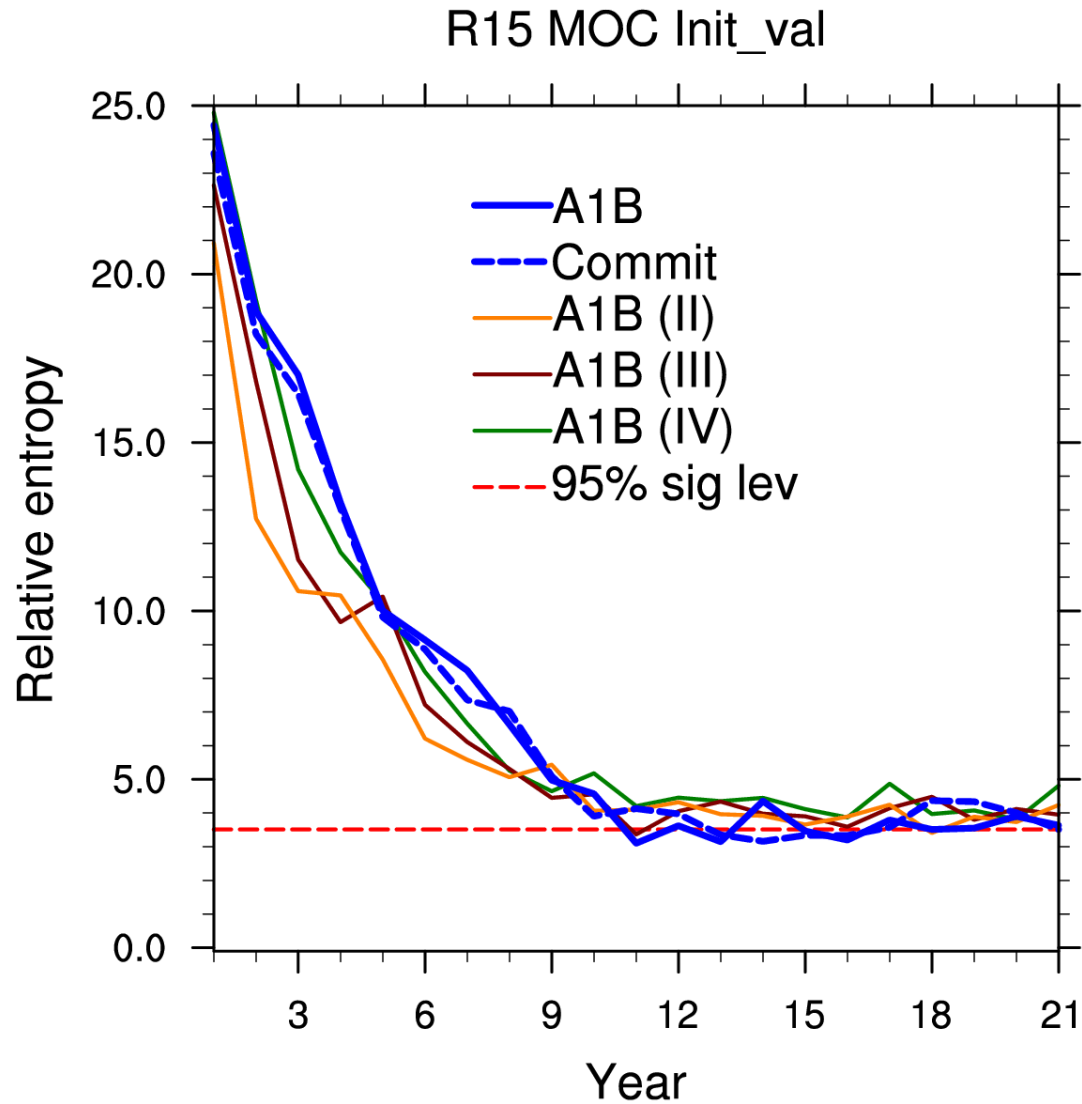
- ✓ Initial-value components
- ✓ Forced response



Predictability of the CEOF1 of AMOC



Predictability of the AMOC (15 EOFs)

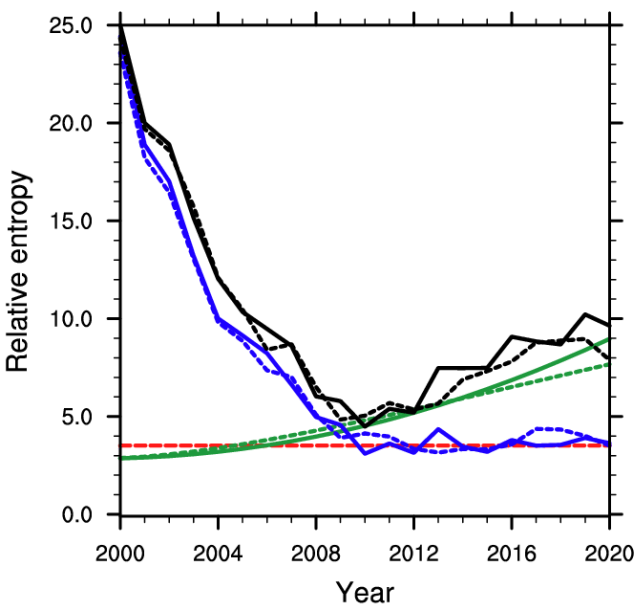


AMOC

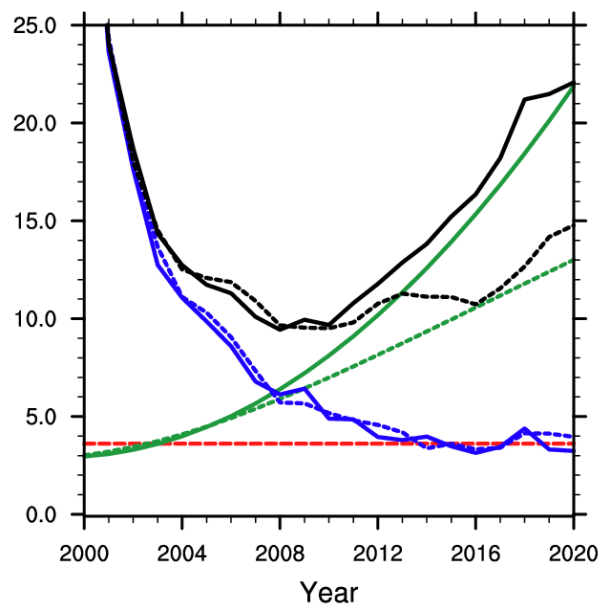
T0-500

SST

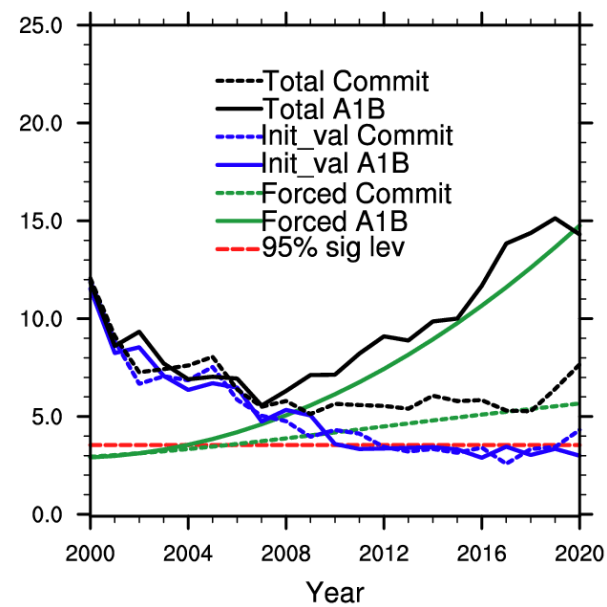
R15 MOC



R15 T0-500



R15 SST

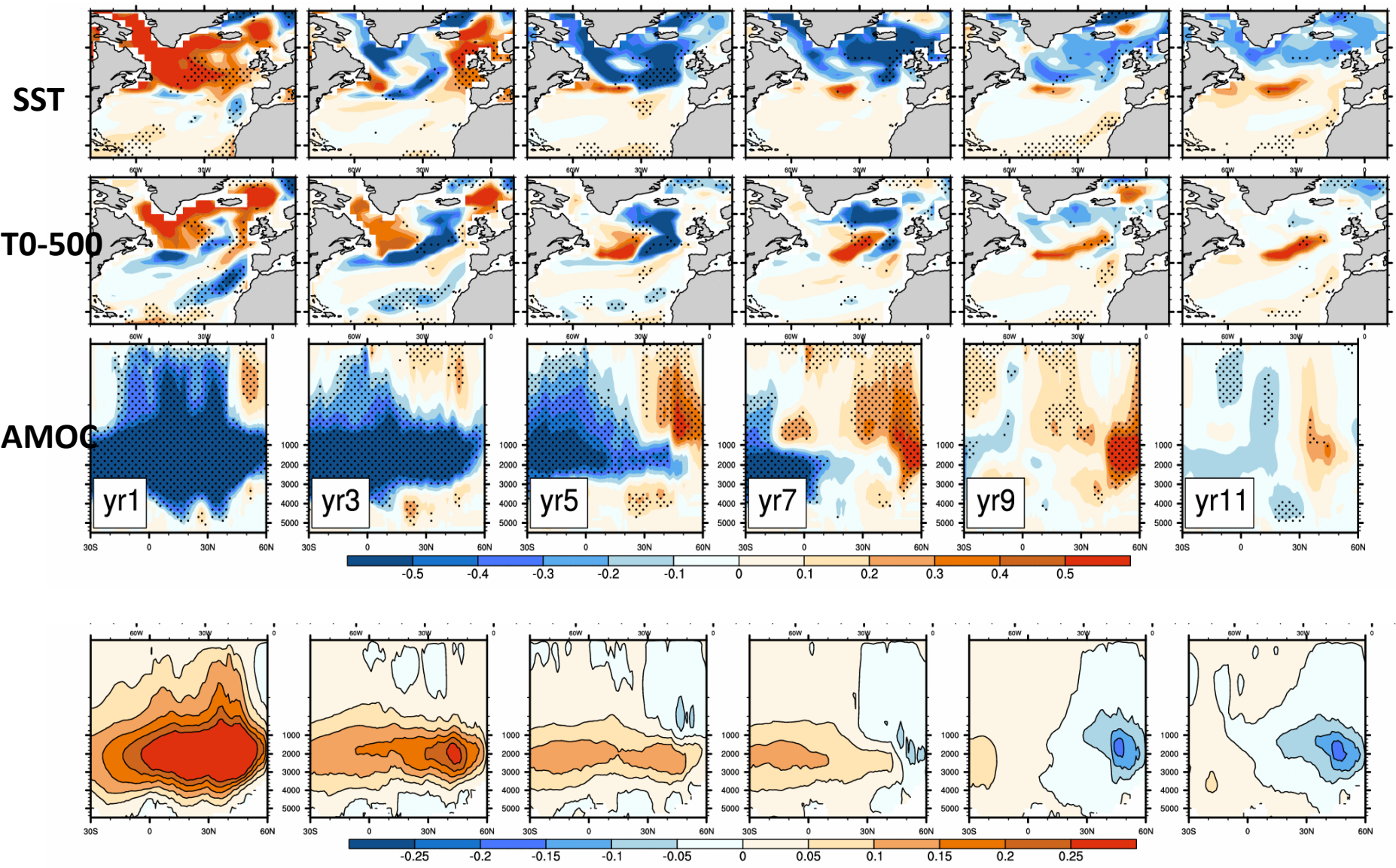


AMOC

- The initial-value component of the AMOC is predictable for about 10-15 years in the CCSM3 large ensemble experiments.
- The predictability mainly comes from the signal component.
- There is a tendency of the EOF1 to evolve to the EOF2 on the decadal time scale.
- The forced response is more predictable than the initial-value component after 10 years.
- The AMOC is no more predictable than the subsurface temperature.

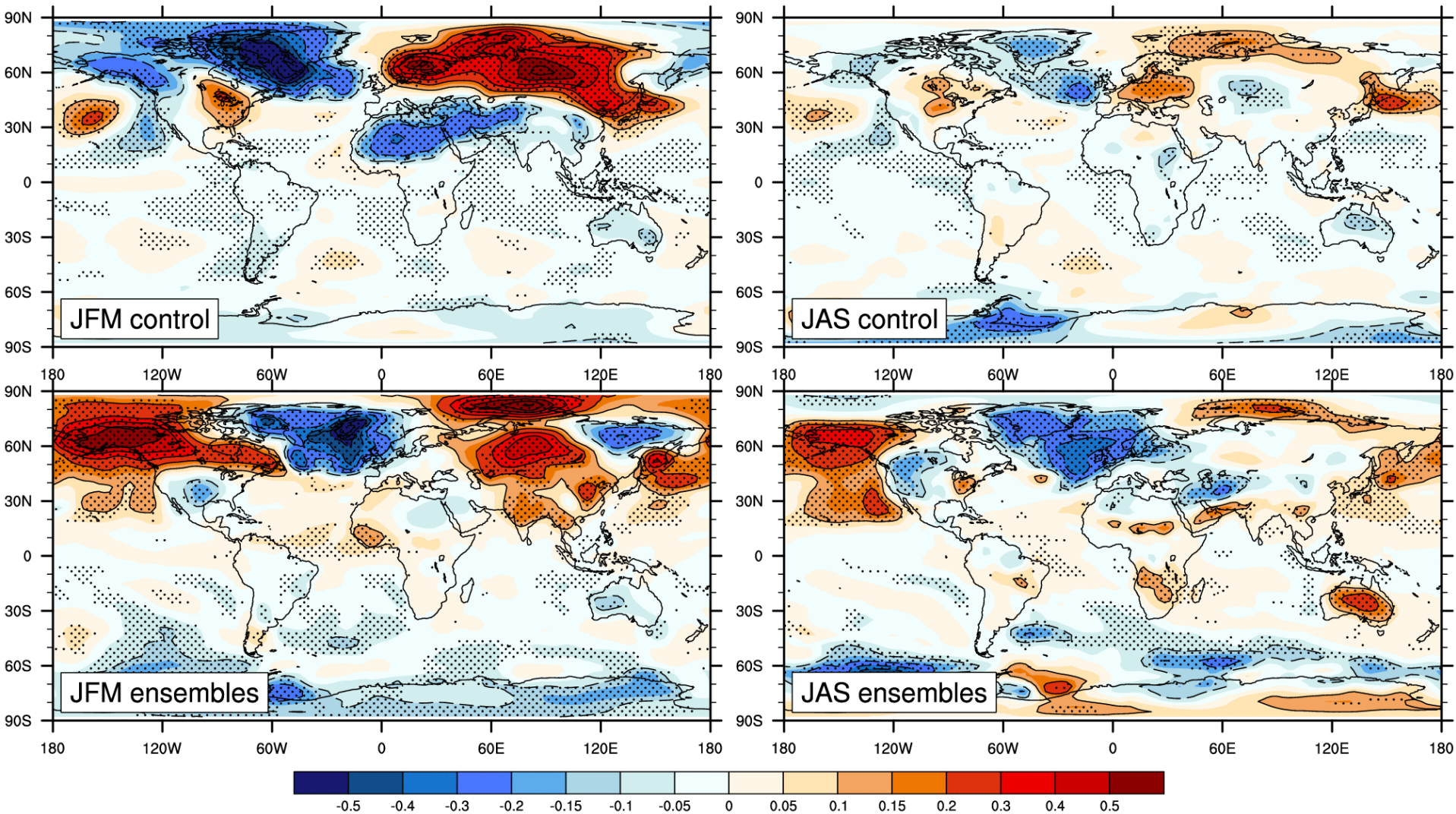
atm

- Are there signals in the atmospheric large-scale surface climate?
- Are the signals associated with the AMOC?
- Relative importance of the natural variability and the forced response

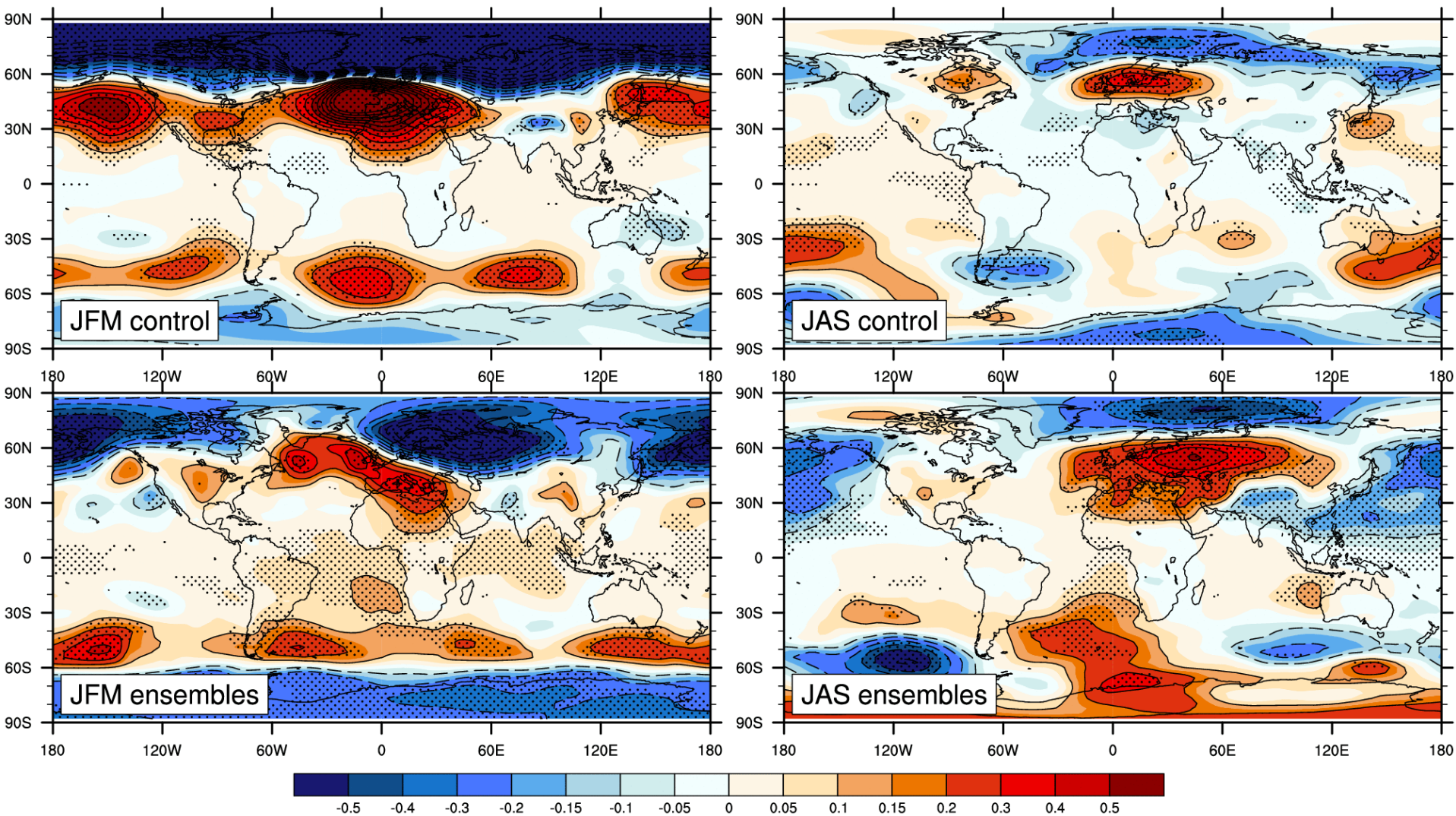


1-2-1 smoothed to remove the ENSO signals

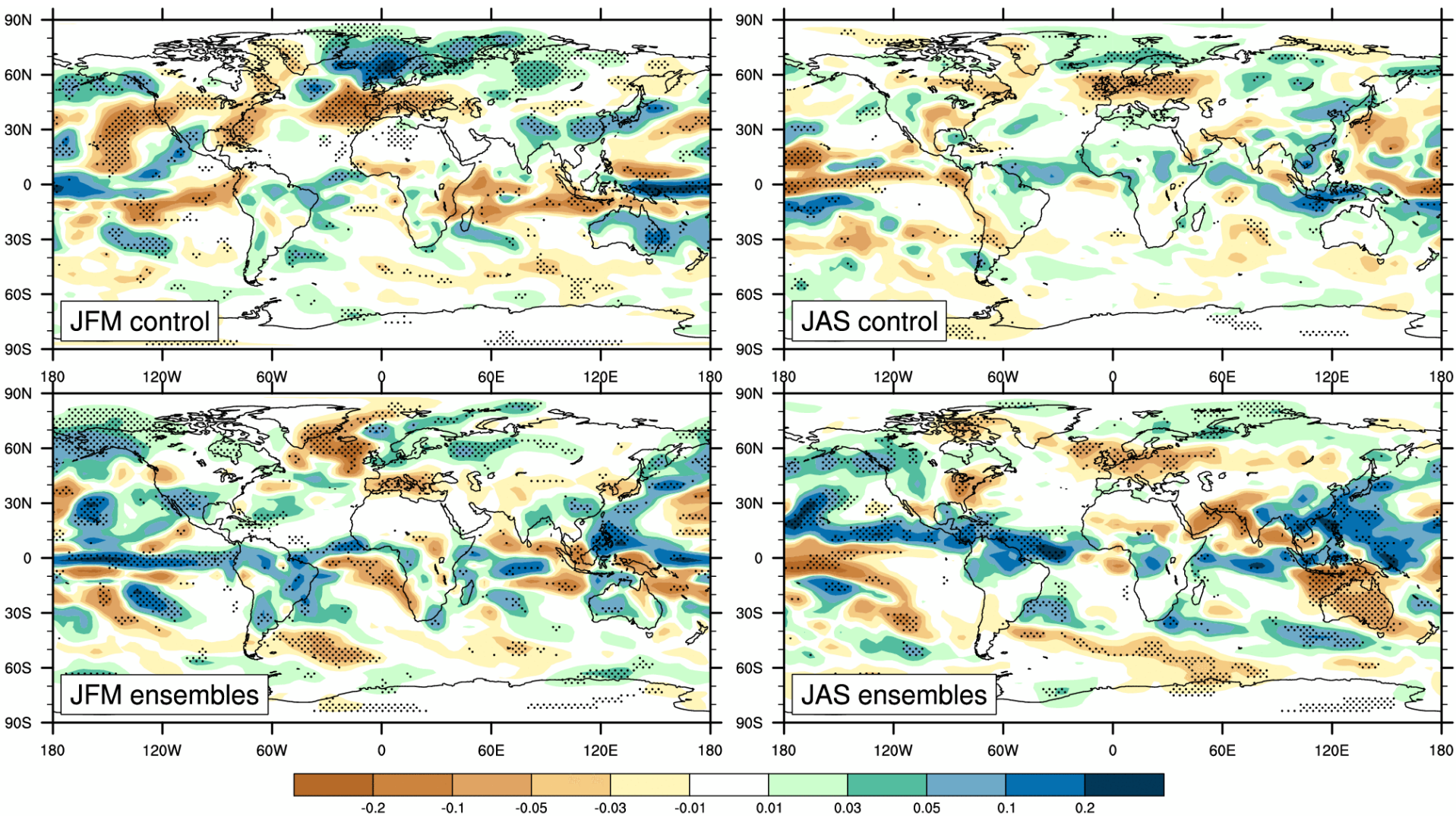
Ensemble mean TAS in yr5-10



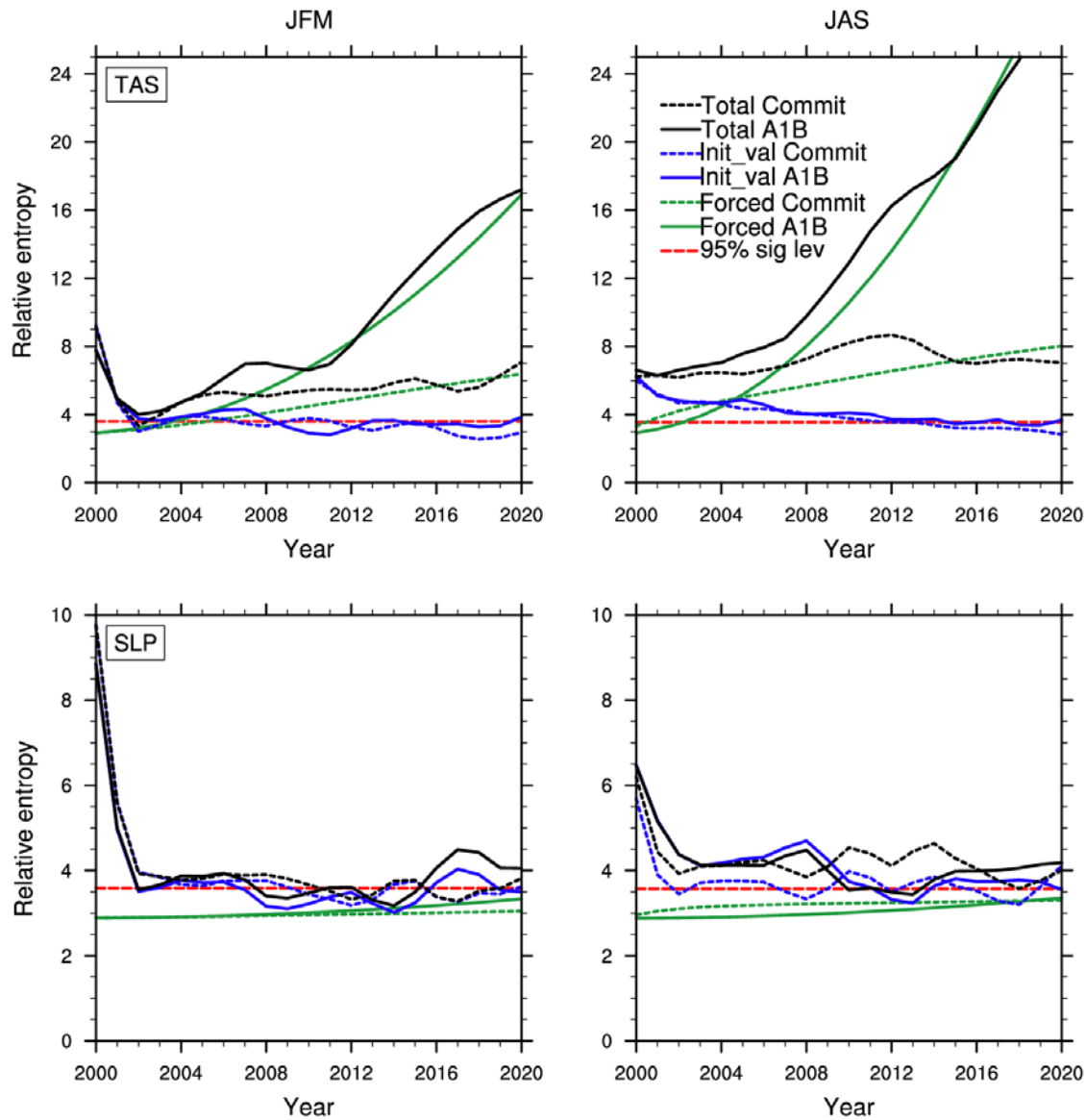
Ensemble mean SLP in yr5-10



Ensemble mean Precipitation in yr5-10



R of 15EOFs in NH (20N-90N)




Summary of the atm

- The large ensemble mean indicates that the AMOC grows from $-EOF1$ to $-EOF2$ in the first decade. Both the subsurface temperature and SST signals are consistent with the control run composite at the same AMOC phase.
- The summer atmospheric surface climate is more predictable than the winter counterpart.
- From year 5 to year 10, the large ensemble means in SLP/precipitation over the North Atlantic and Europe suggest a positive NAO, and the anomalous patterns agree with the control run at the same phase of the AMOC.

Summary of the atm

- Despite the extended initial-value predictability in the NH summer, the forced response in the TAS becomes more dominant than the initial-value predictability after about five years.
- The large ensemble mean patterns in North America are different from those in the control run composite , possibly due to the disagreement in the phase of the PDO.
- The extended predictability in the atmospheric circulation supports a two-way coupling in the midlatitudes.



Atm has much
less predictability
than ocn!

air-sea coupling?

SLP Variance (CCSM3-CAM3)/CAM3

