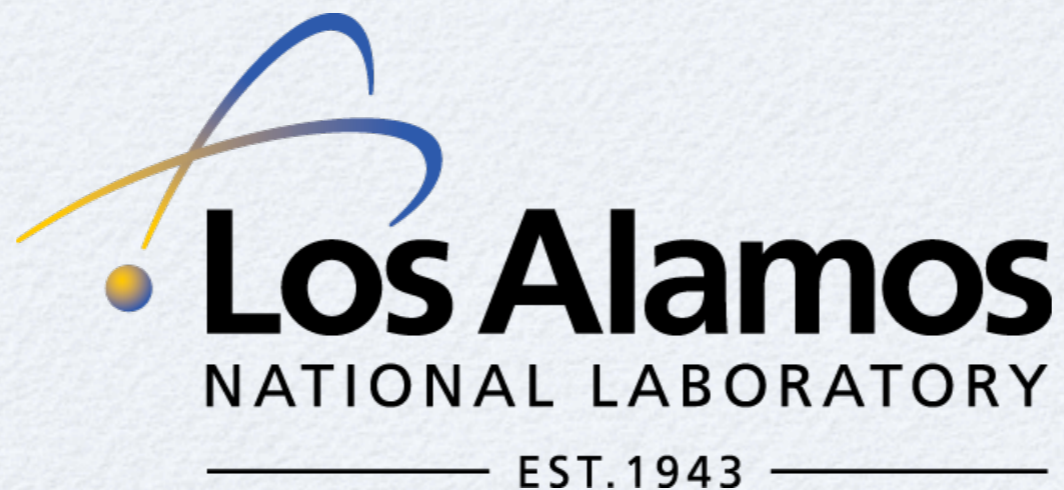


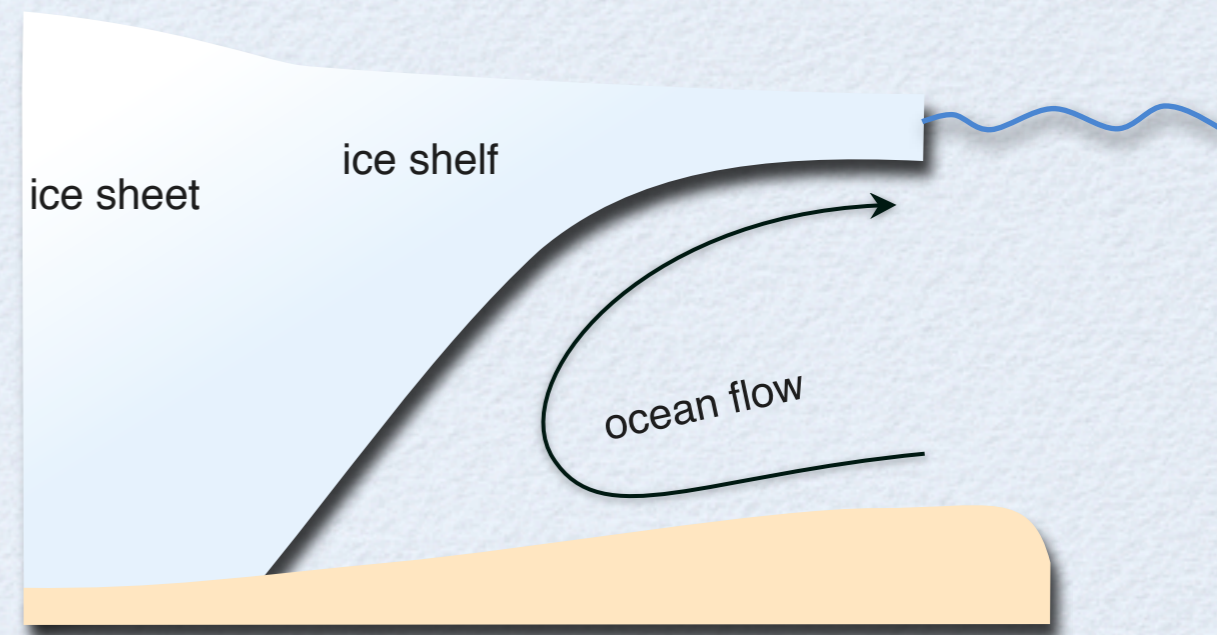
AN UPDATE ON MODELING LAND-ICE/
OCEAN INTERACTIONS IN CESM

Xylar Asay-Davis



OUTLINE

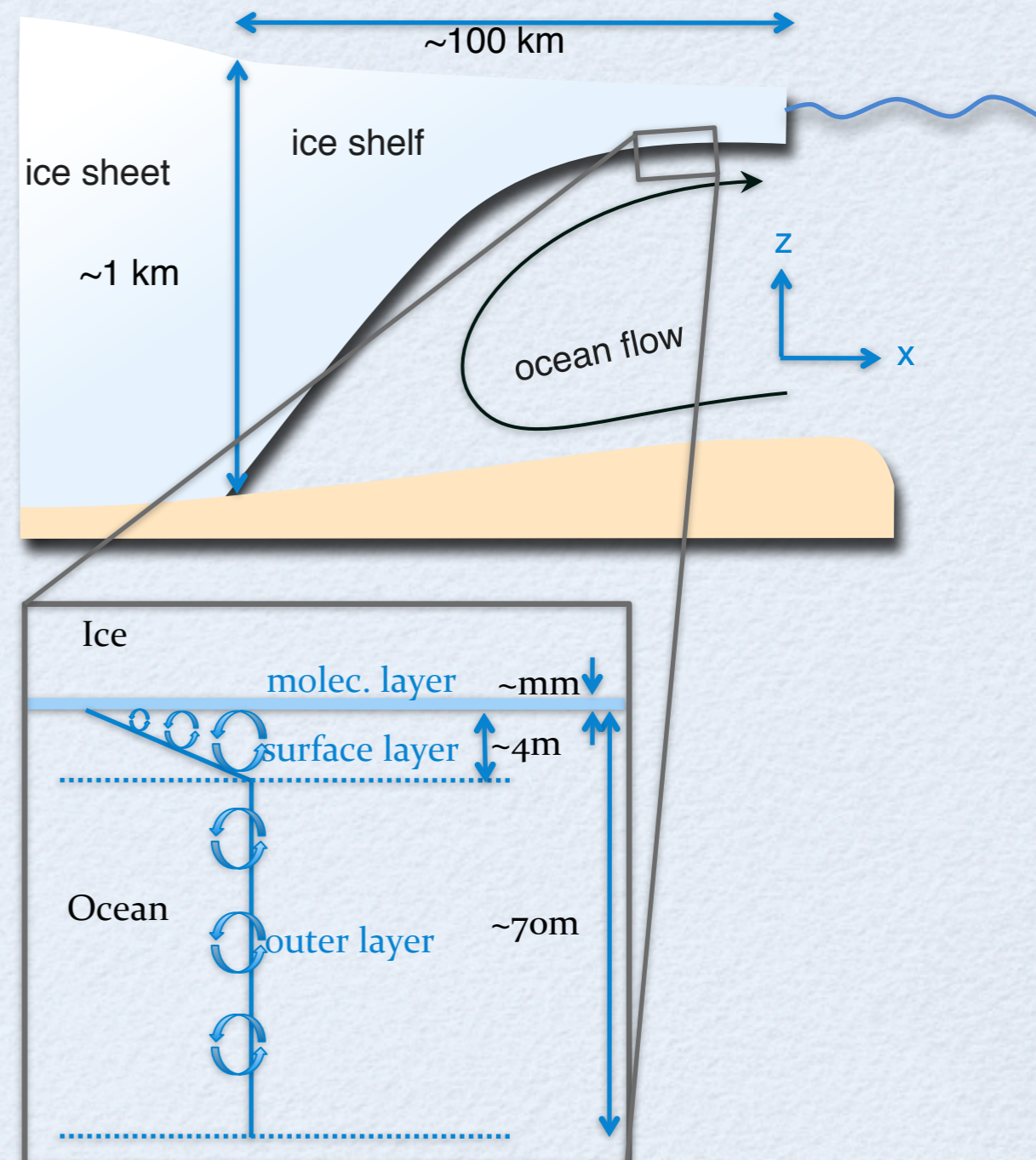
- Ice-shelf / Ocean Coupling
- Boundary Conditions
- Challenges with the Immersed Boundary Method
- A New Partial Cells Approach



ICE-SHELF/OCEAN COUPLING

Challenging Physics:

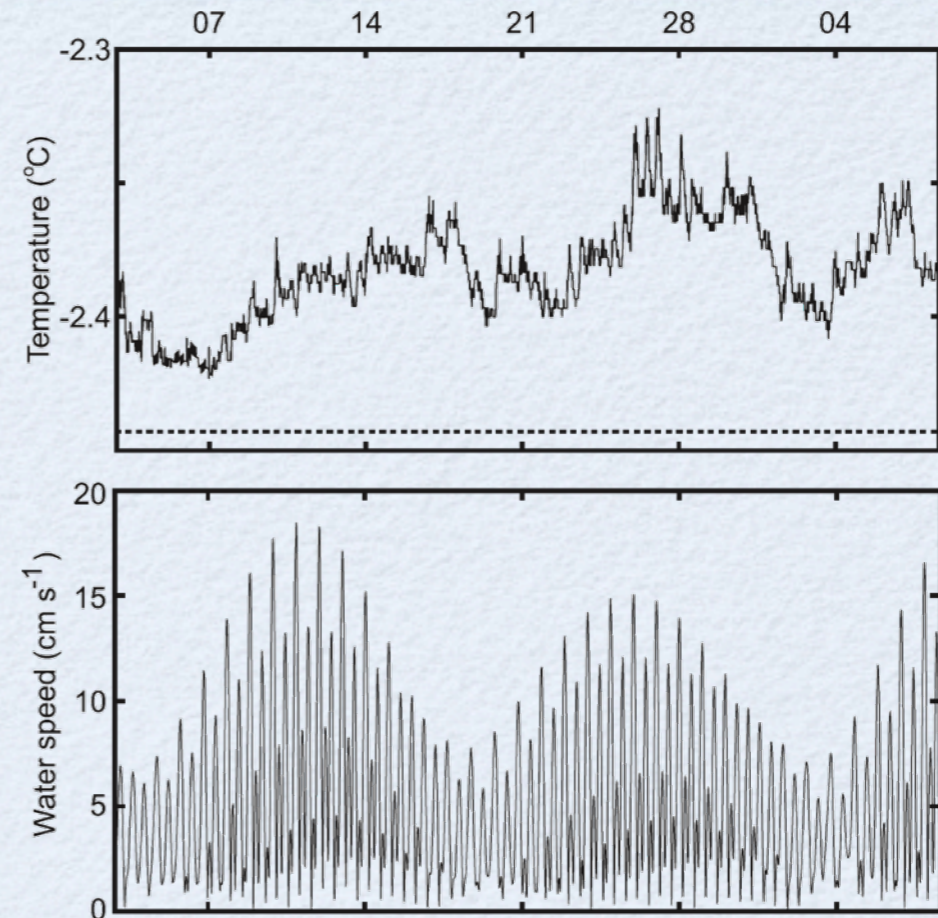
- Many length scales



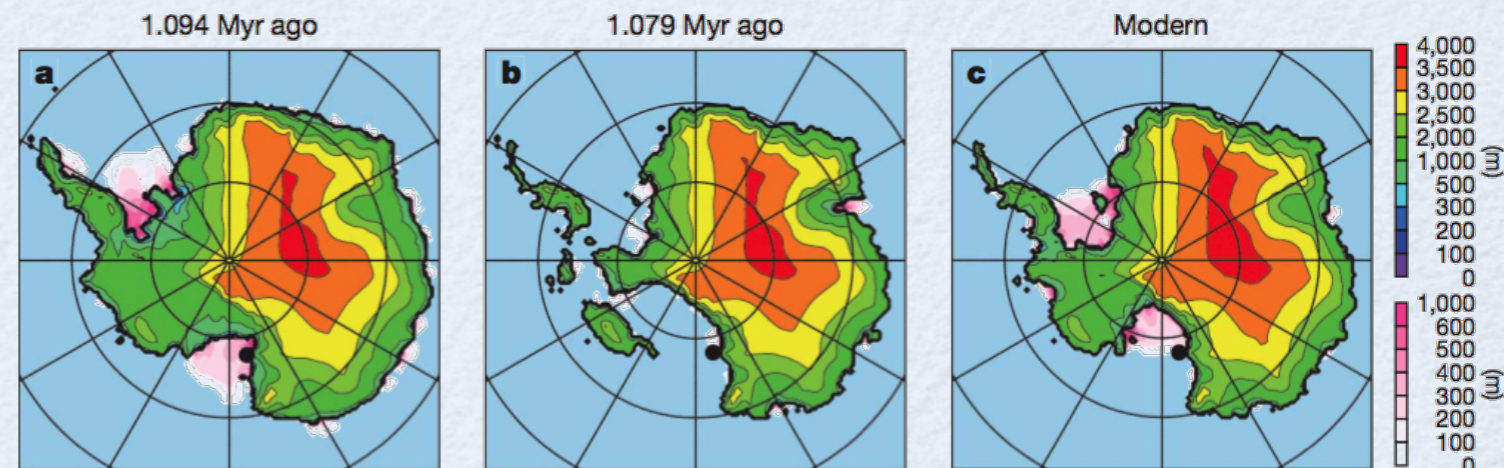
ICE - SHELF/OCEAN COUPLING

Challenging Physics:

- Many length scales
- Many time scales



Jenkins et al. 2010

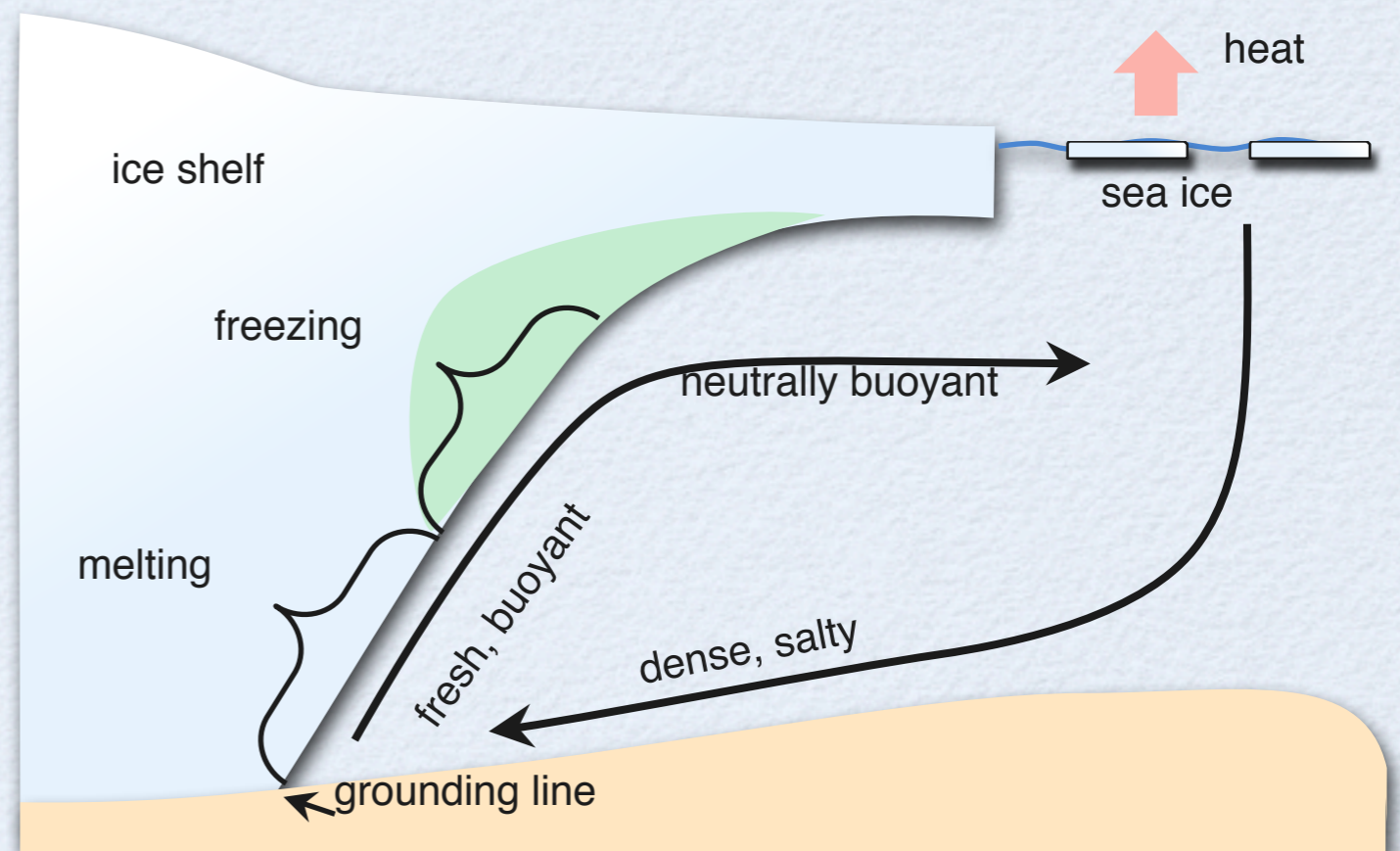


Pollard and DeConto 2009

ICE/OCEAN COUPLING

Challenging Physics:

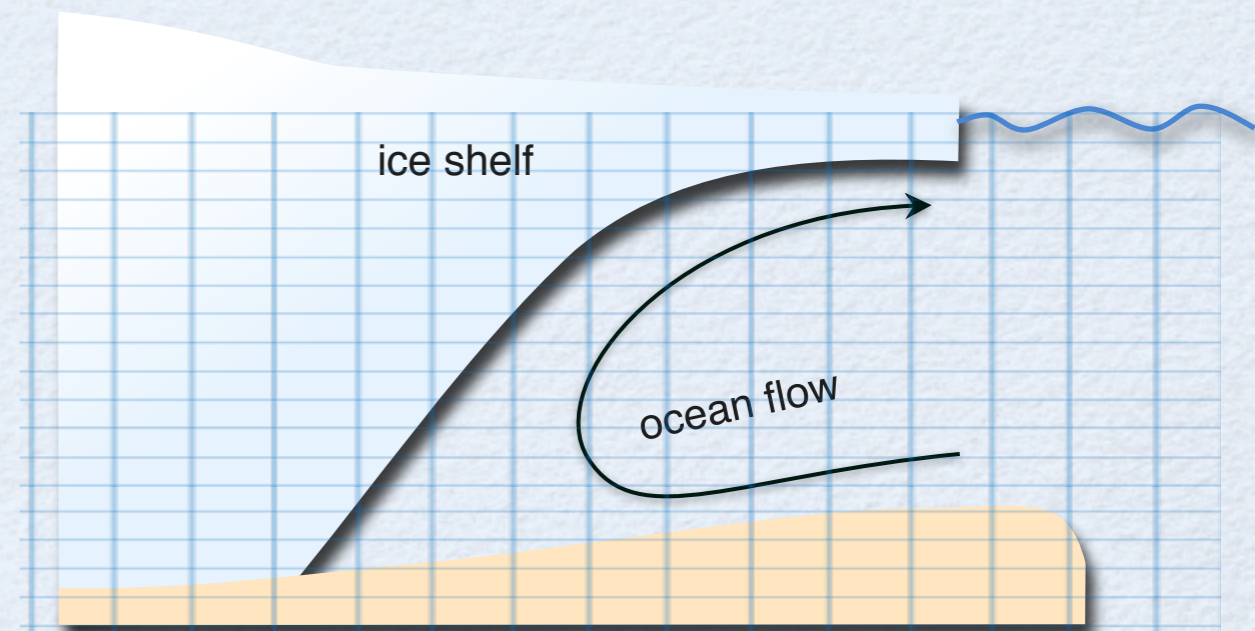
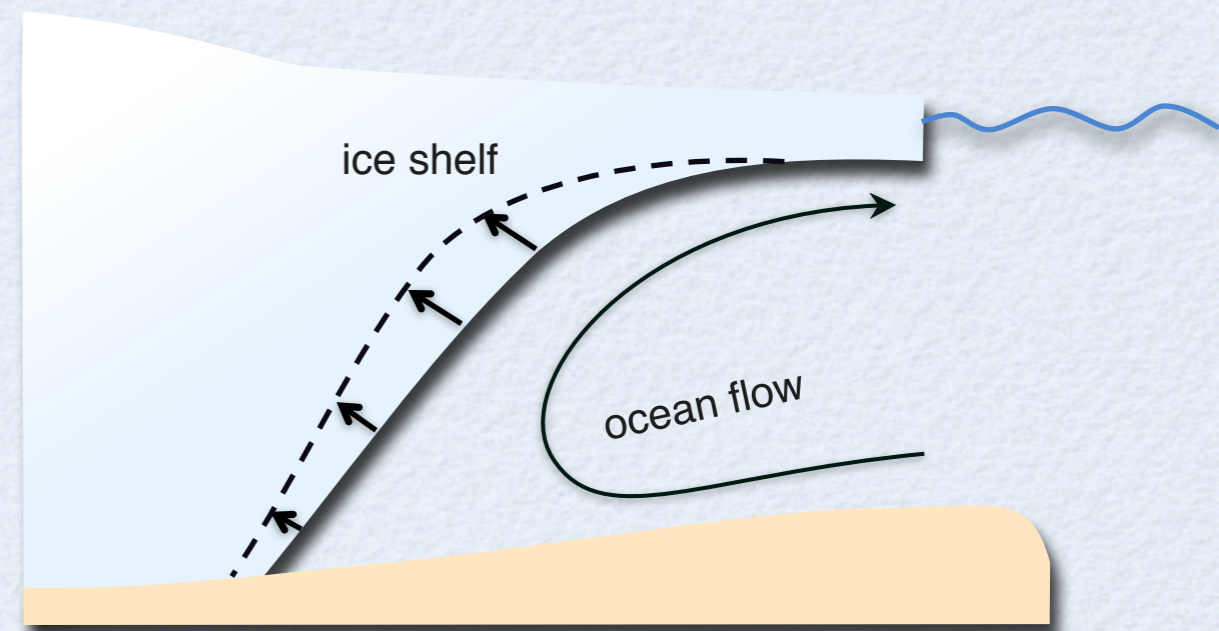
- Many length scales
- Many time scales
- Many processes



ICE/OCEAN COUPLING

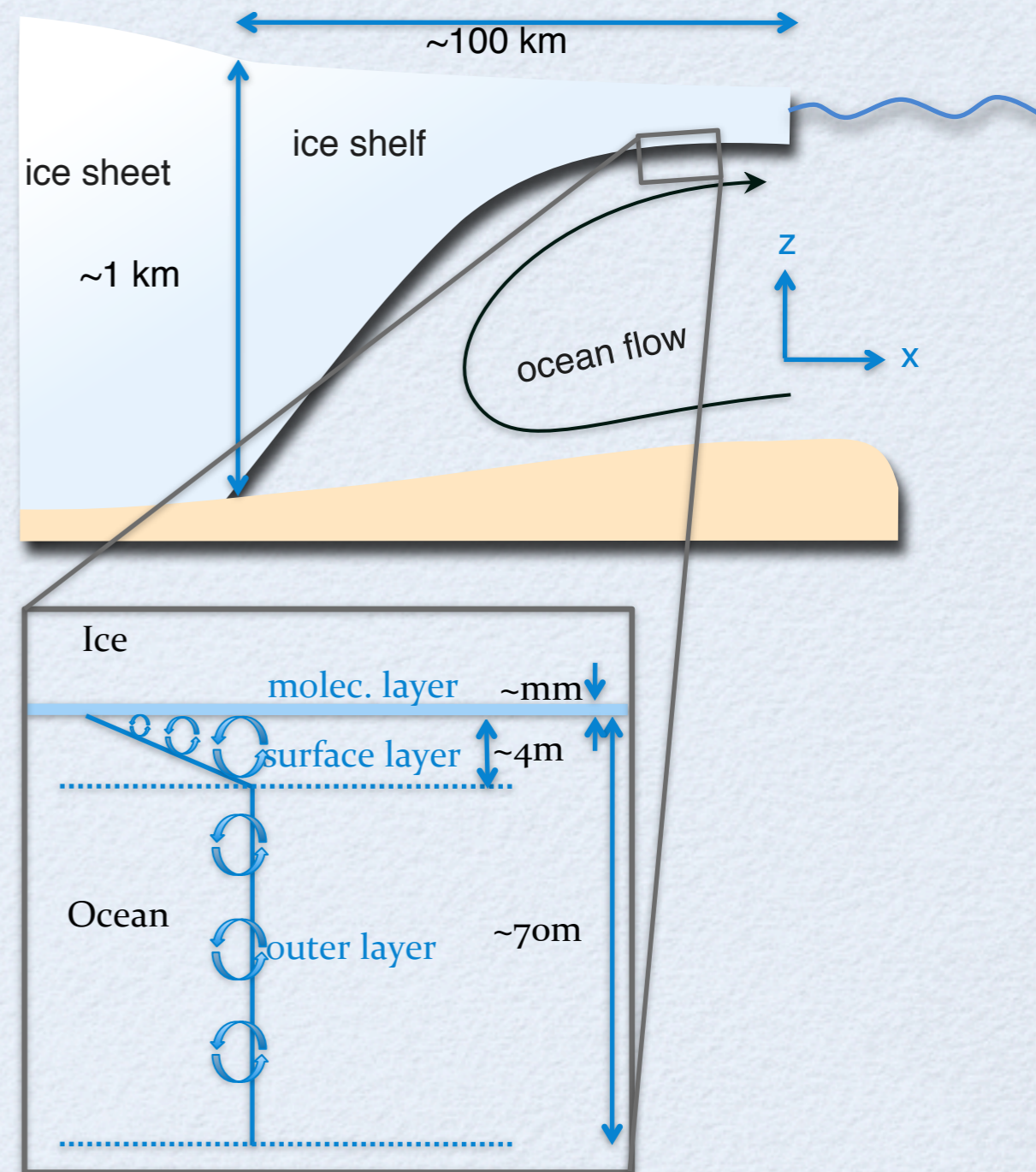
Challenging Numerics:

- Moving boundaries
- Anisotropic grids ($\Delta x \gg \Delta z$)
- Under-resolved physics



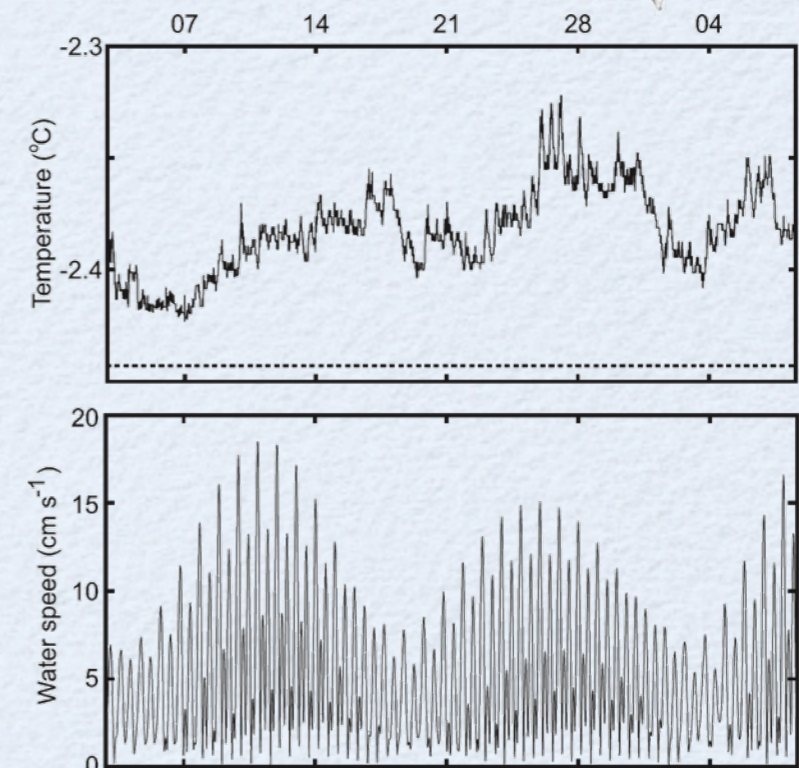
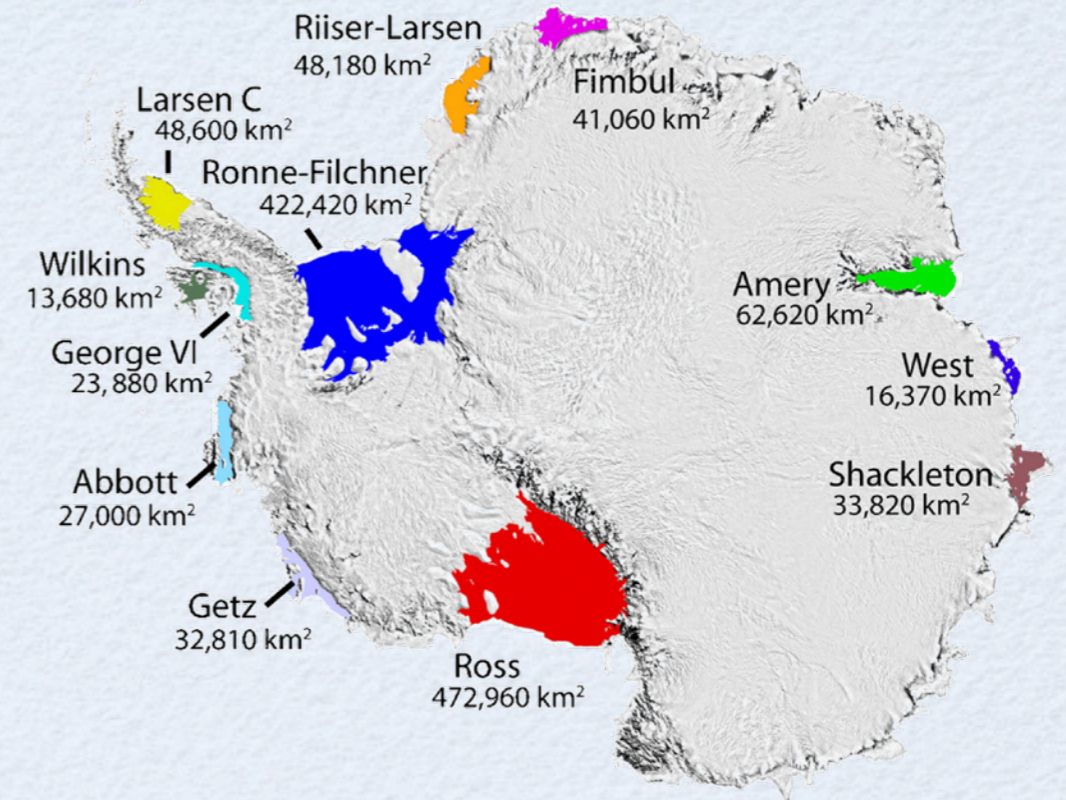
BOUNDARY LAYER PHYSICS

- Simplified version of McPhee 2008 boundary layer model for sea ice
- Gives **heat, salt, momentum** and **mass fluxes** at the interface
- Includes **stratification**, very important for rapid melting



BOUNDARY LAYER PHYSICS

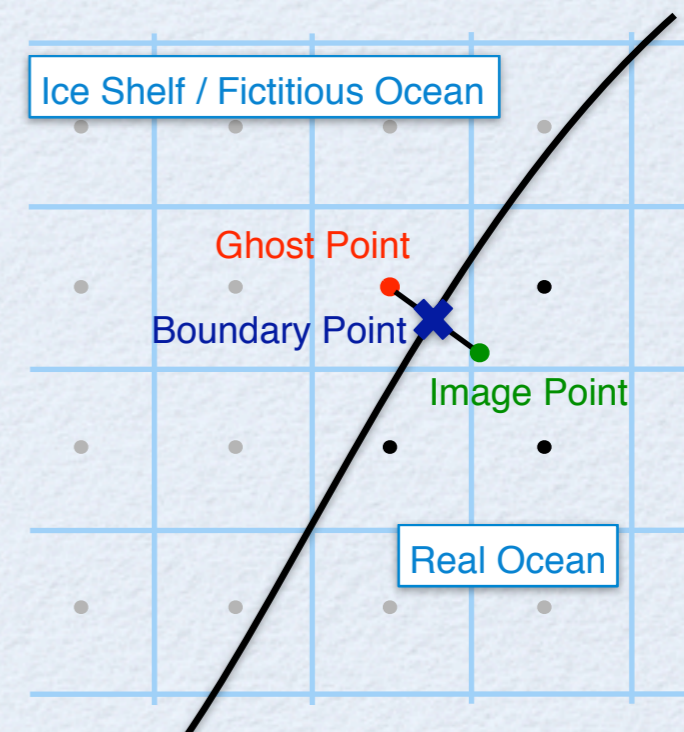
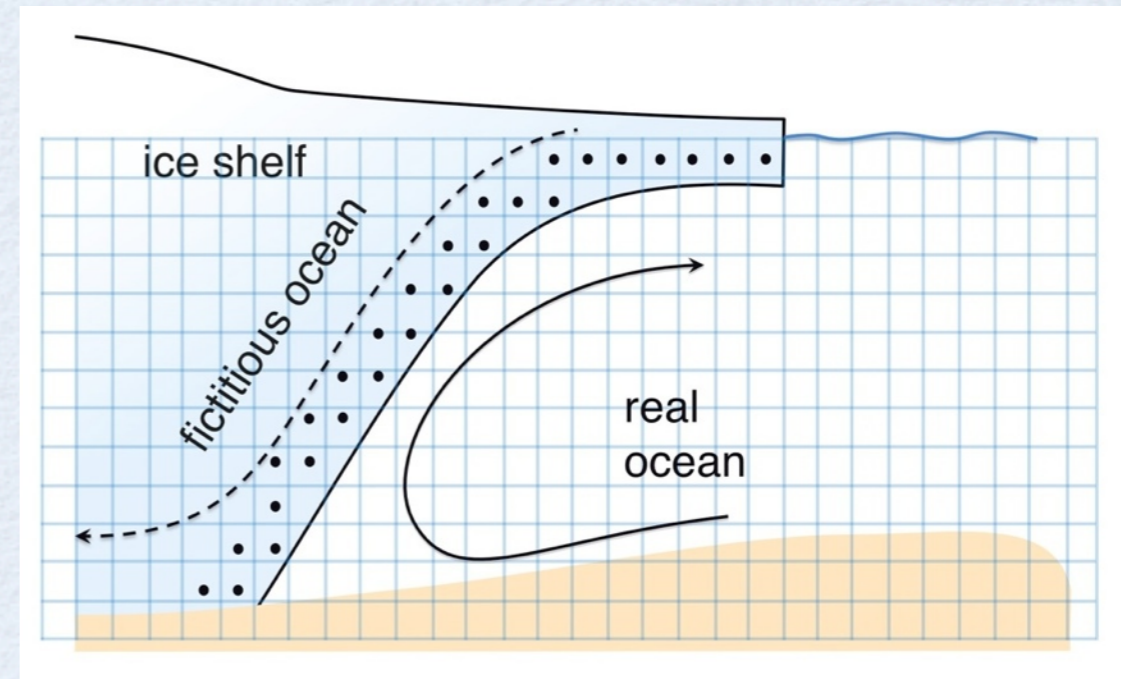
- Unknown coefficients are calibrated based on measurements under Ronne Ice Shelf (Jenkins et al. 2010)
- More calibration data expected in coming years (Fimbul, Larsen C and George VI Ice Shelves)



Jenkins et al. 2010

IMMERSED BOUNDARY METHOD

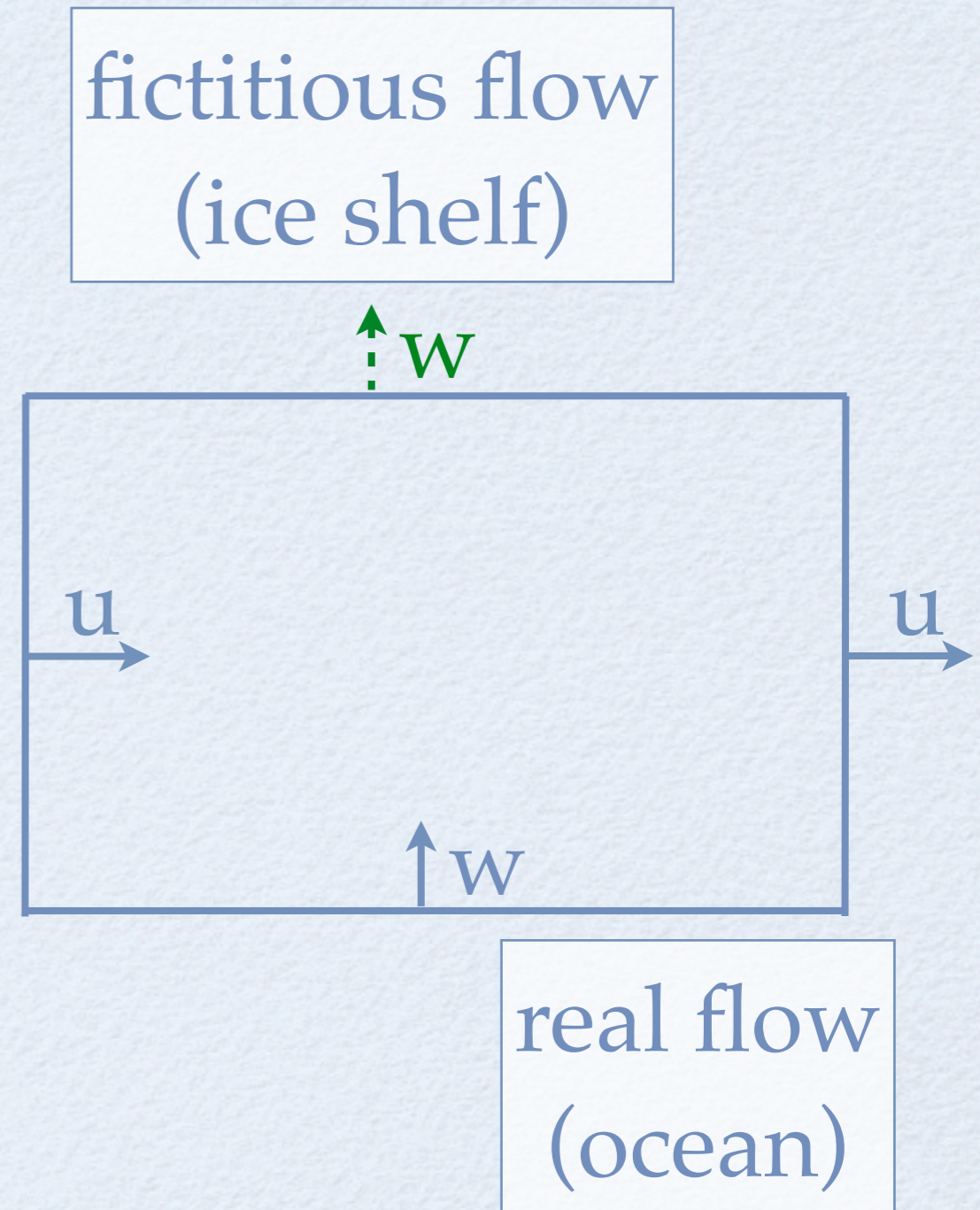
- Handle complex, moving boundaries on constant grids (including calving front)
- Boundary conditions by forcing a fictitious flow
- Extrapolate fluid values from **image points** to **ghost points** using boundary conditions (mass, heat and salt fluxes)



CHALLENGES

Continuity:

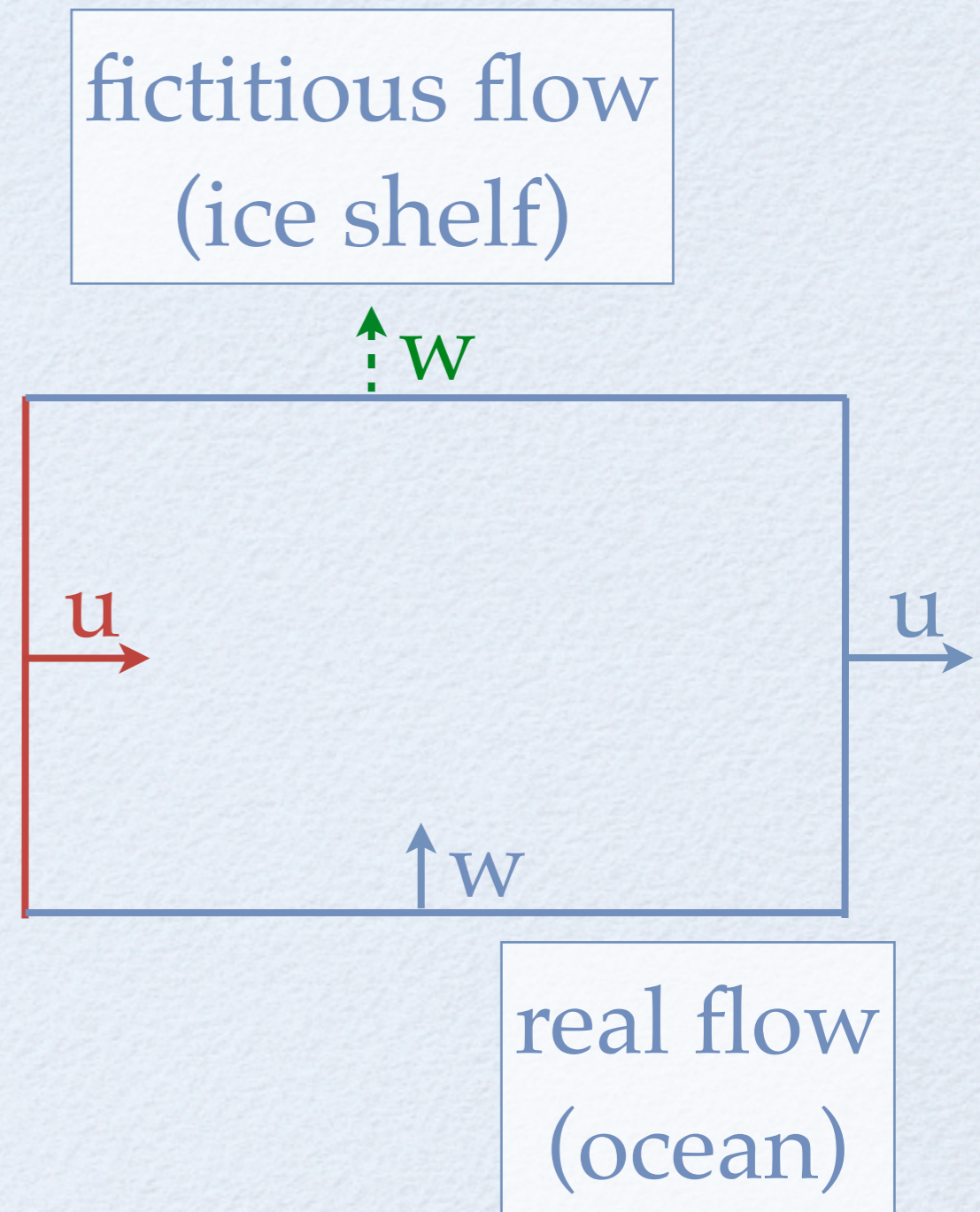
- w at the top of the cell found by continuity (sum of **fluxes** is zero)



CHALLENGES

Continuity:

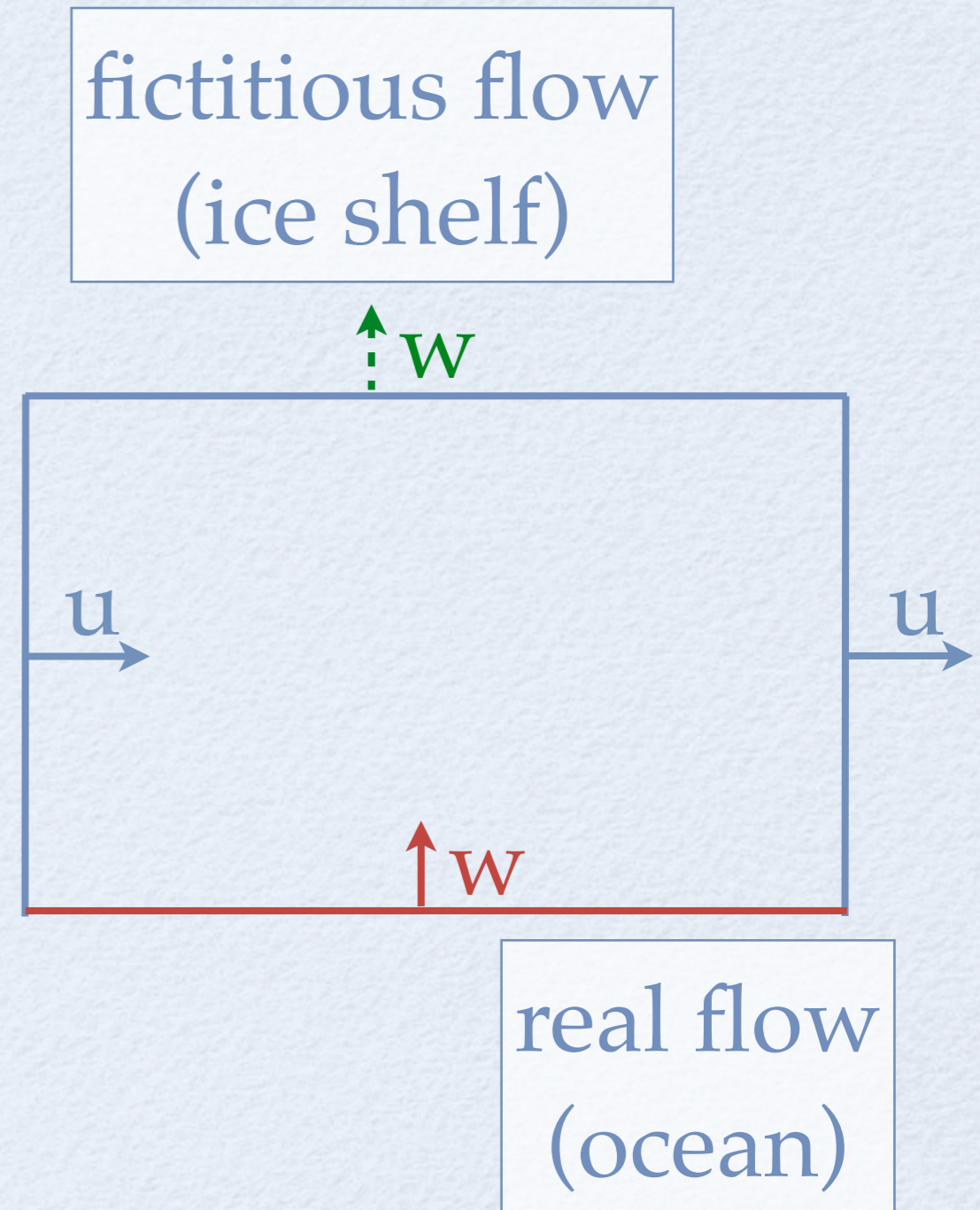
- w at the top of the cell found by continuity (sum of **fluxes** is zero)



CHALLENGES

Continuity:

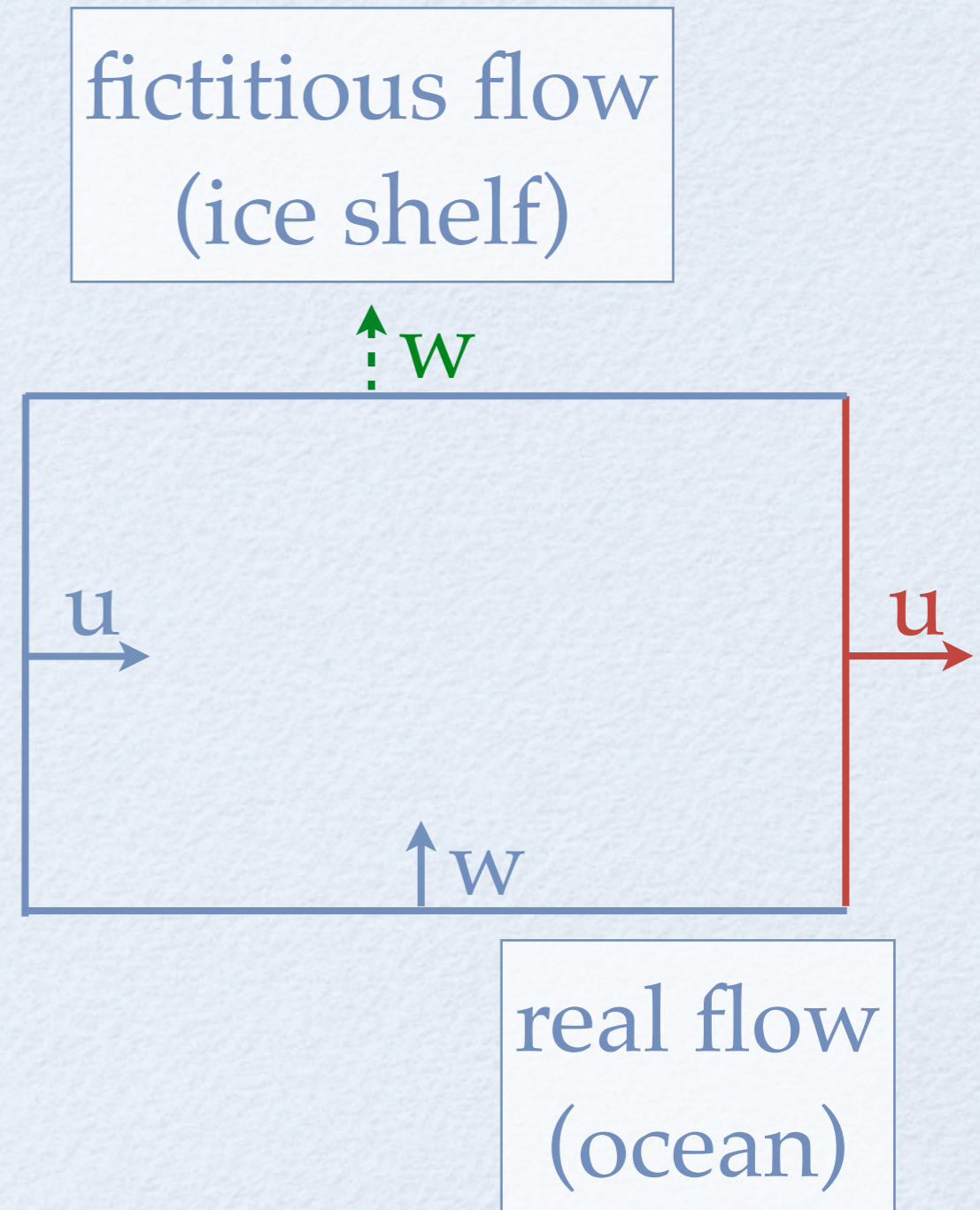
- w at the top of the cell found by continuity (sum of **fluxes** is zero)



CHALLENGES

Continuity:

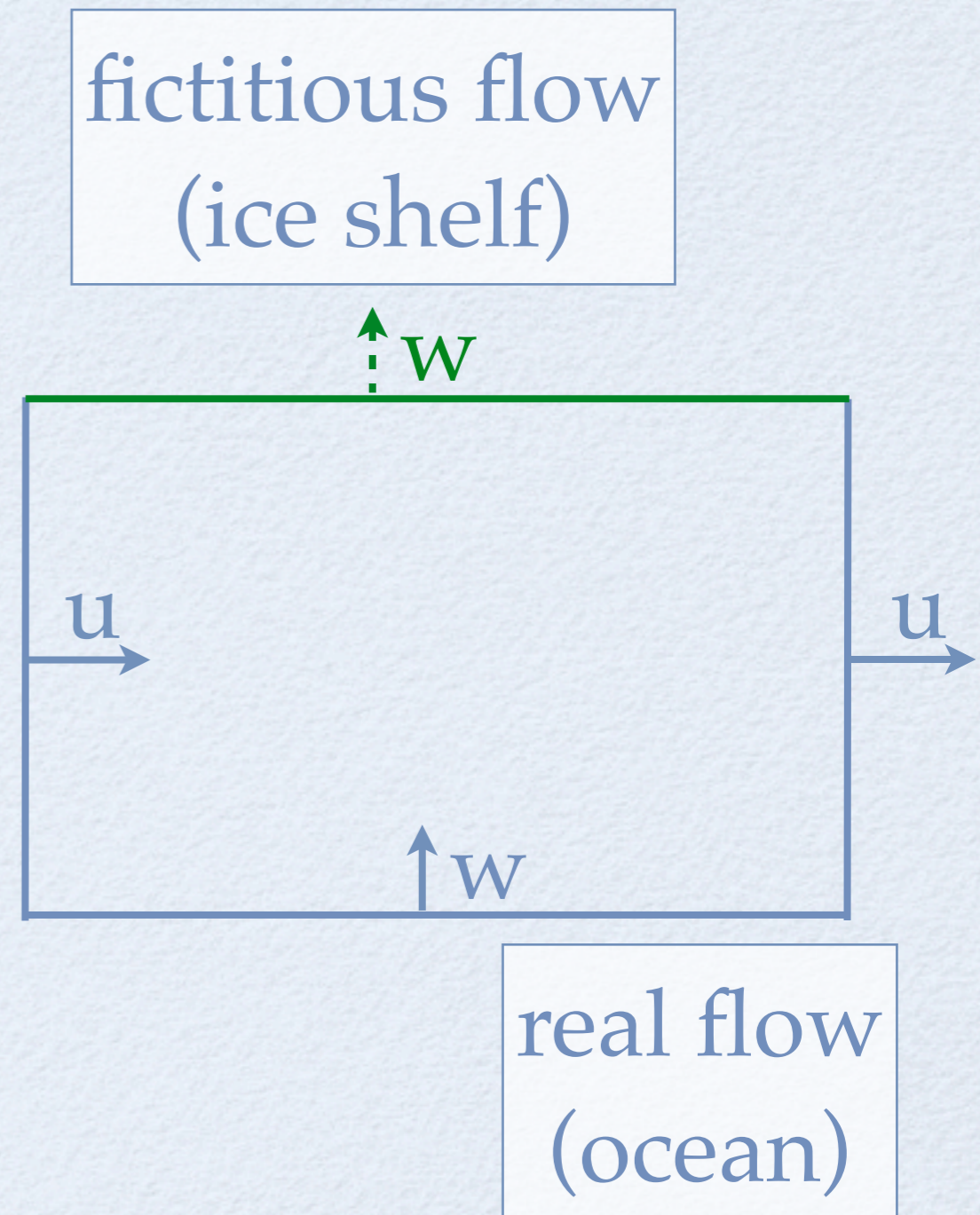
- w at the top of the cell found by continuity (sum of **fluxes** is zero)



CHALLENGES

Continuity:

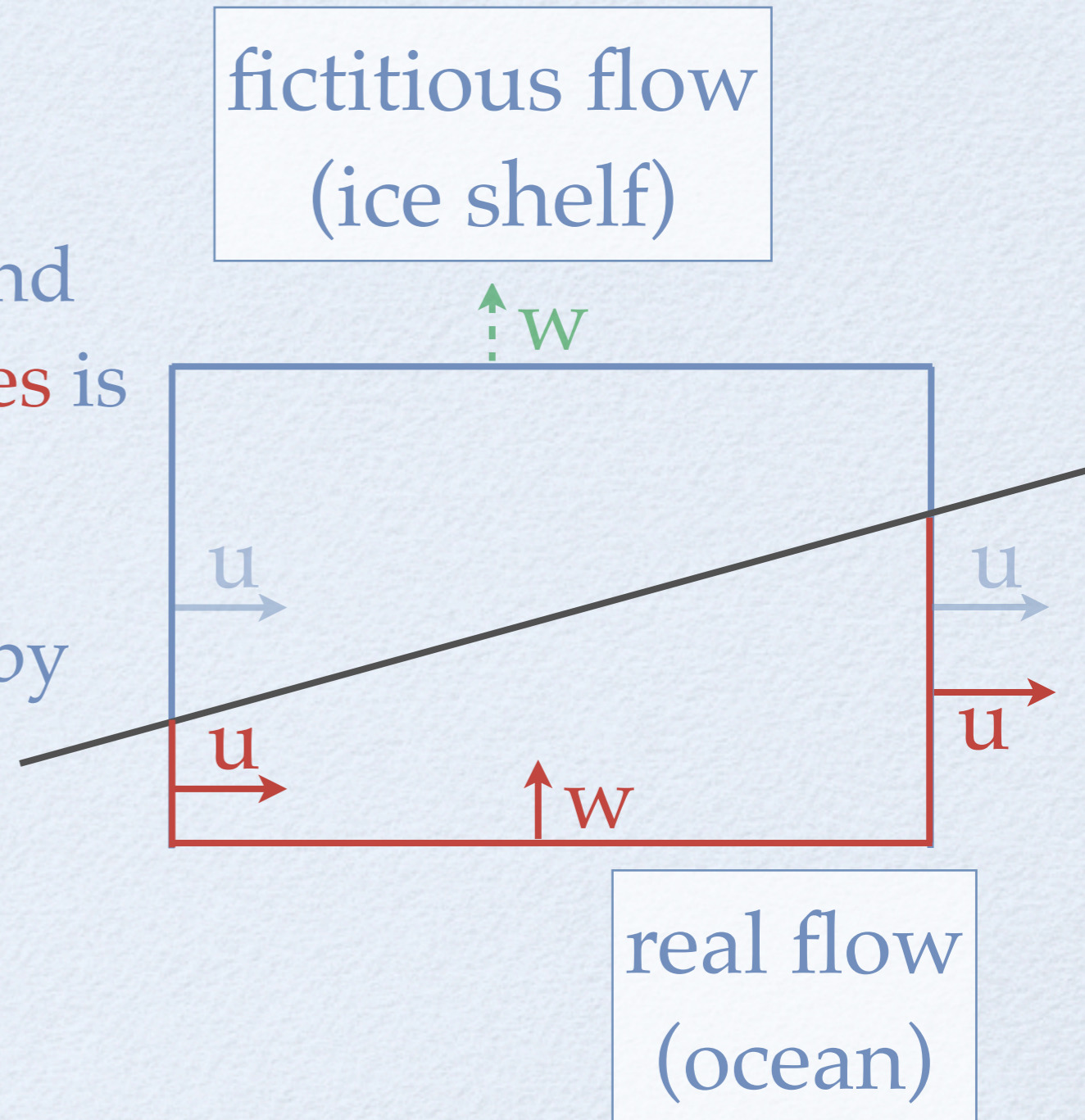
- w at the top of the cell found by continuity (sum of **fluxes** is zero)



CHALLENGES

Continuity:

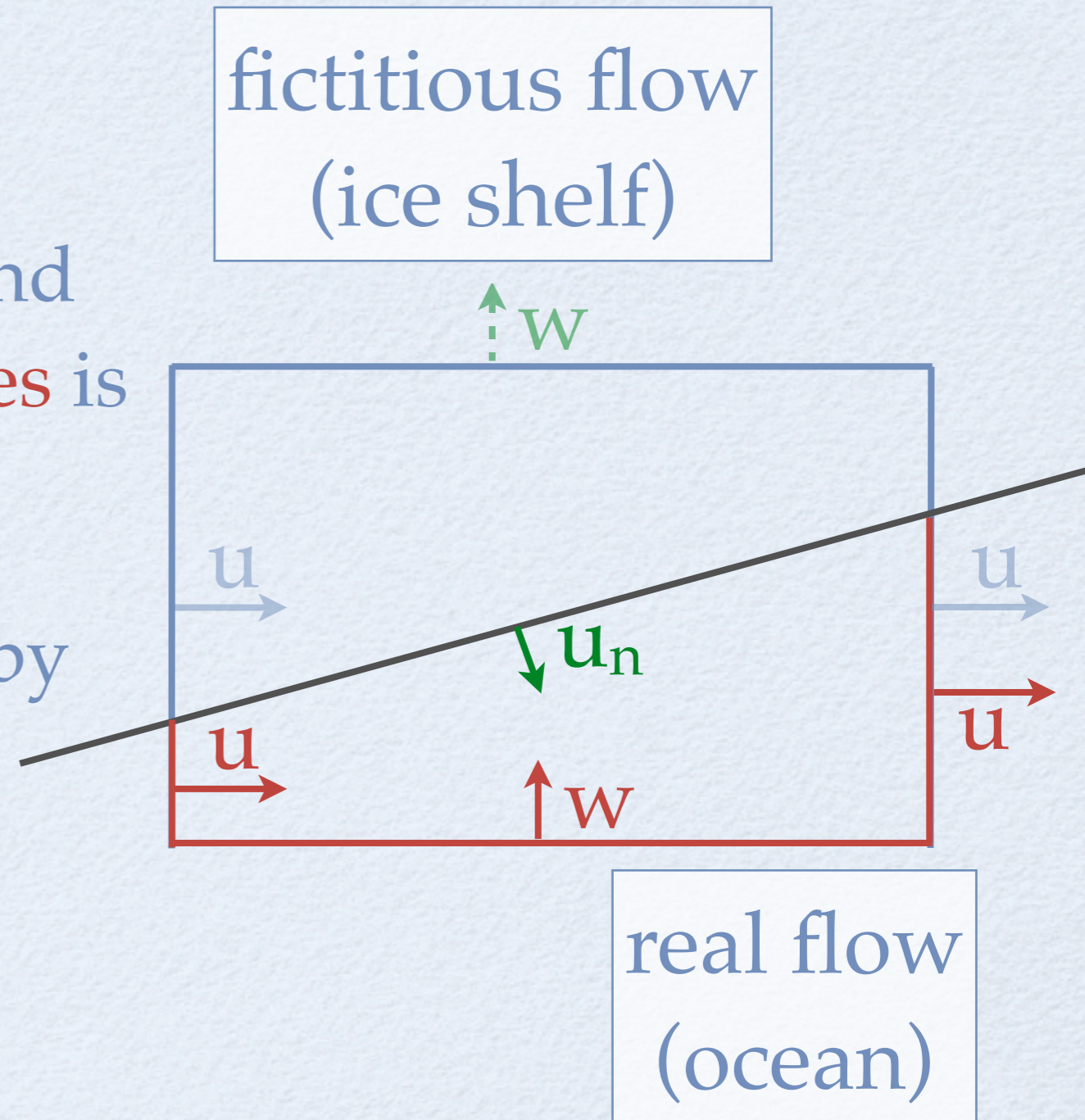
- w at the top of the cell found by continuity (sum of **fluxes** is zero)
- Normal velocity u_n found by fluxes in real flow only



CHALLENGES

Continuity:

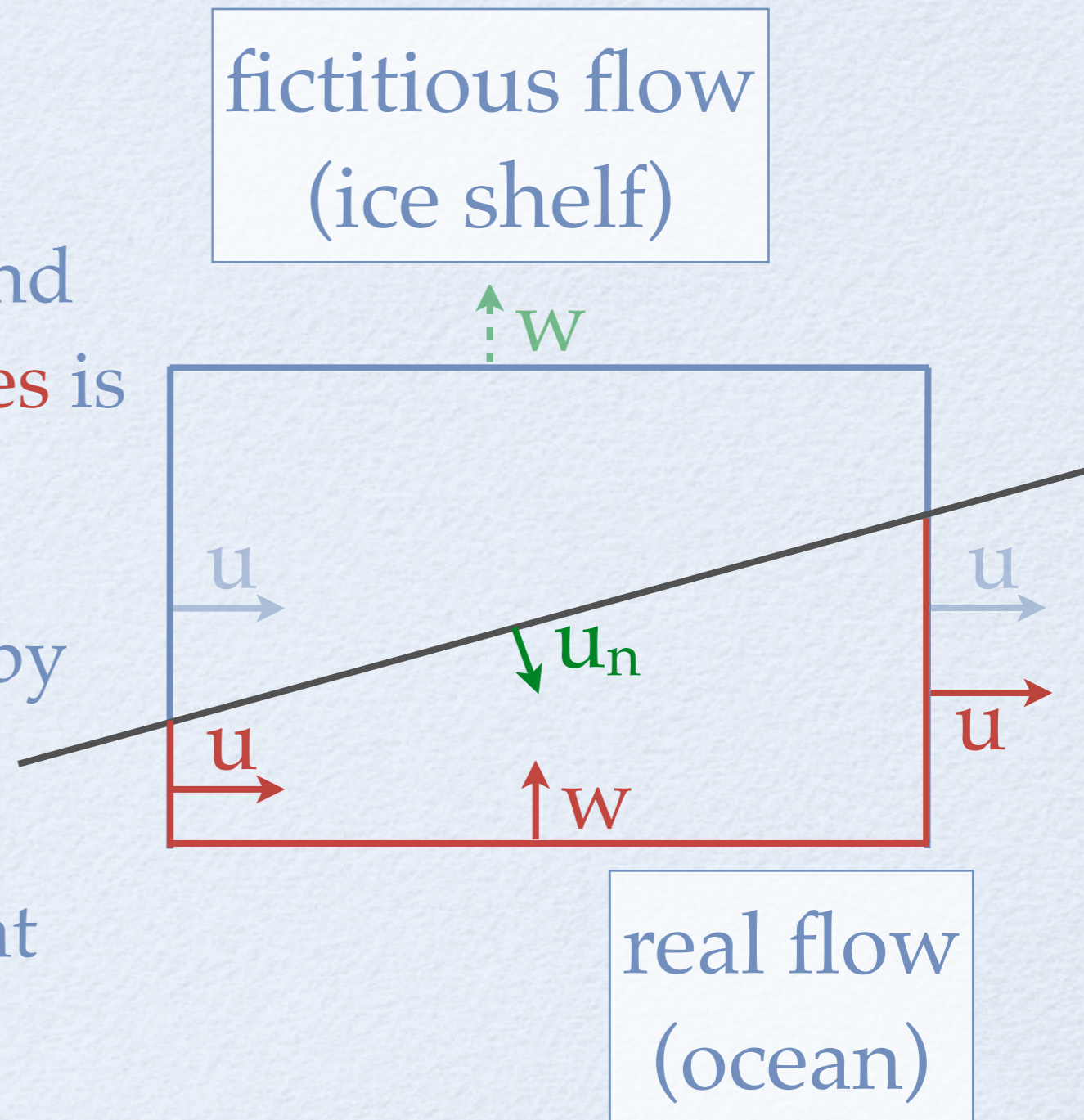
- w at the top of the cell found by continuity (sum of **fluxes** is zero)
- Normal velocity u_n found by fluxes in real flow only



CHALLENGES

Continuity:

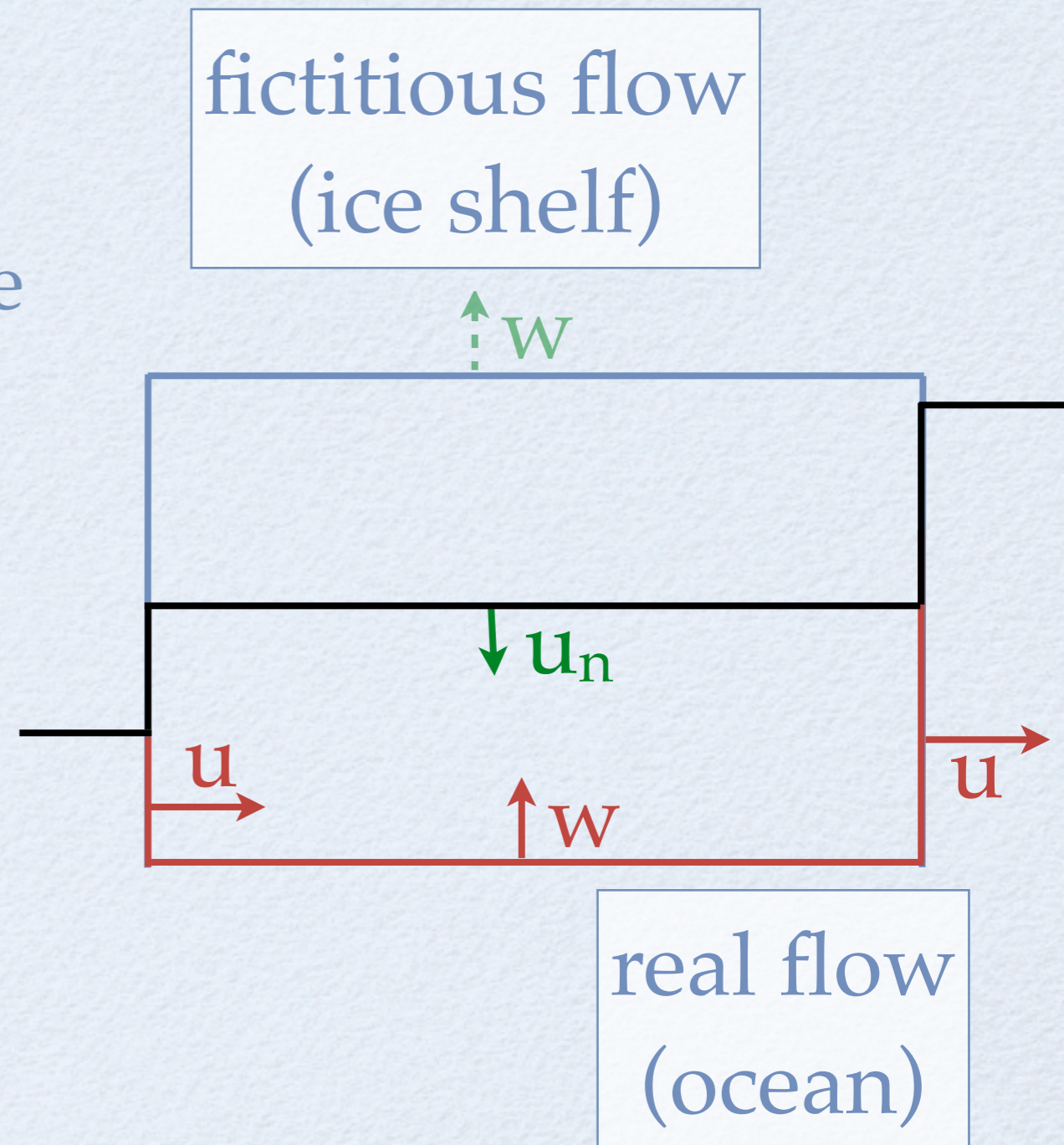
- w at the top of the cell found by continuity (sum of **fluxes** is zero)
- Normal velocity u_n found by fluxes in real flow only
- Different **flux areas** on right and left
- u_n **not** an interpolation of w



CHALLENGES

Continuity:

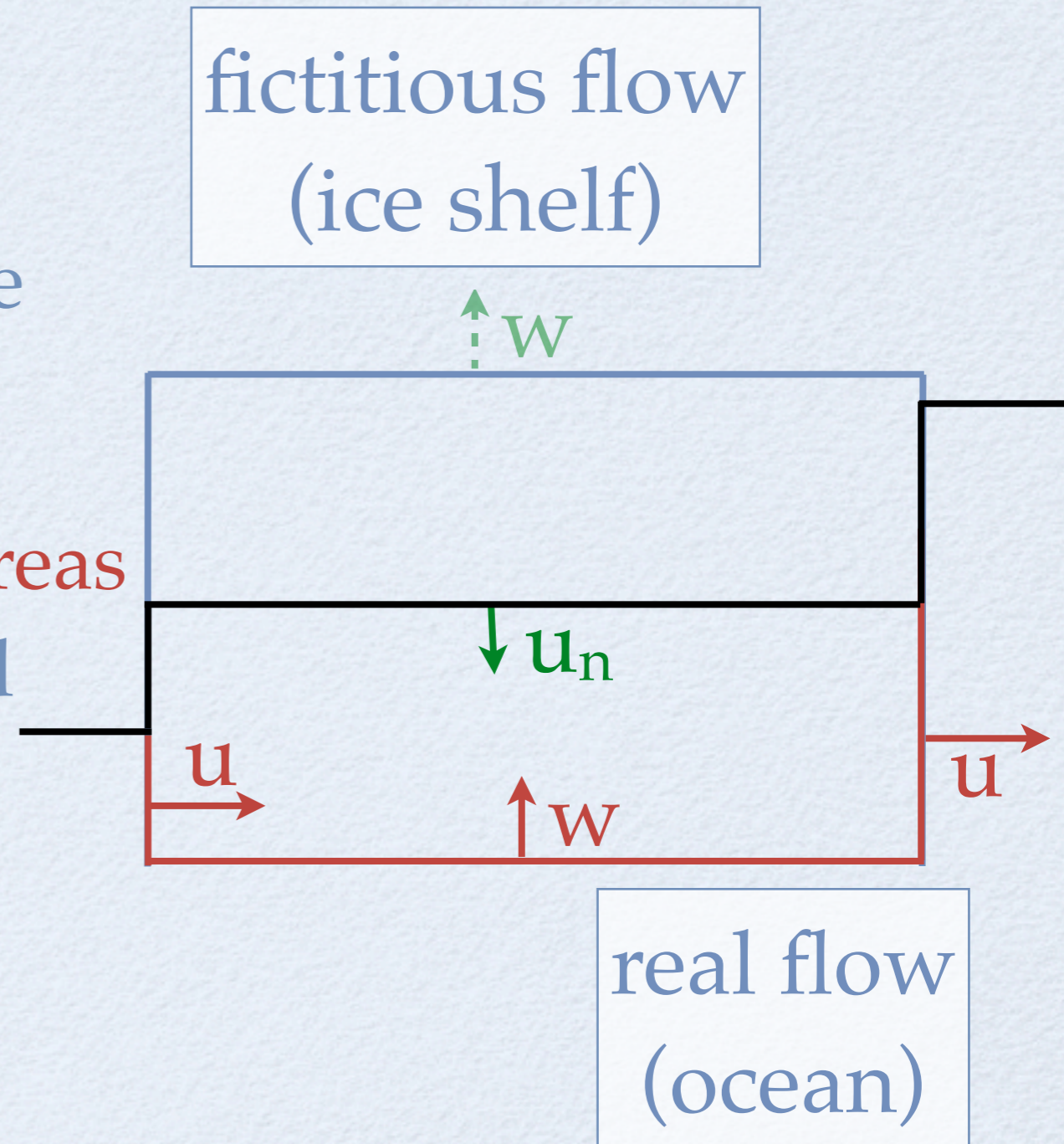
- Partial cells do not solve the problem



CHALLENGES

Continuity:

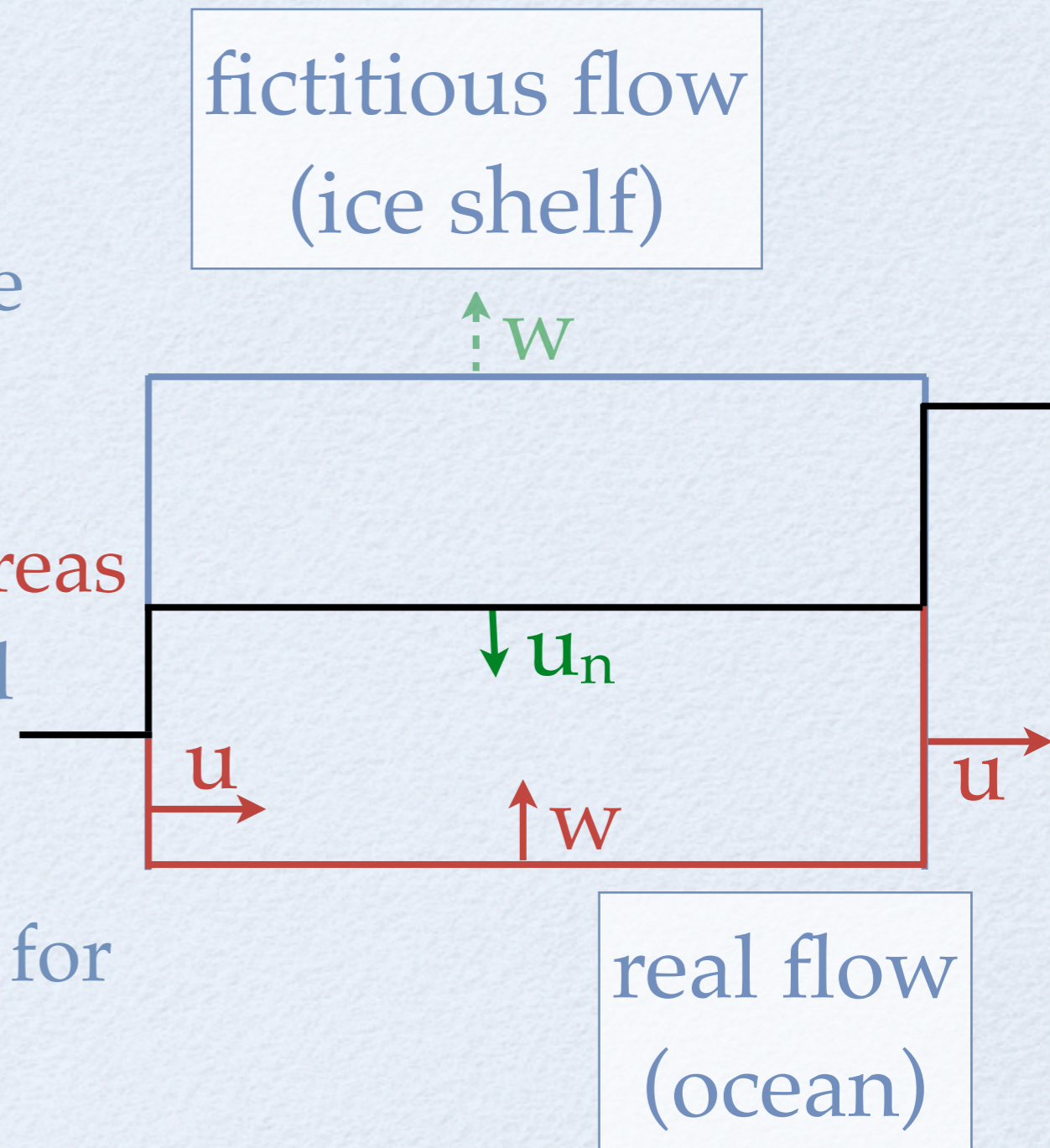
- Partial cells do not solve the problem
- Interface is flat but **fluxes areas** are still different on left and right



CHALLENGES

Continuity:

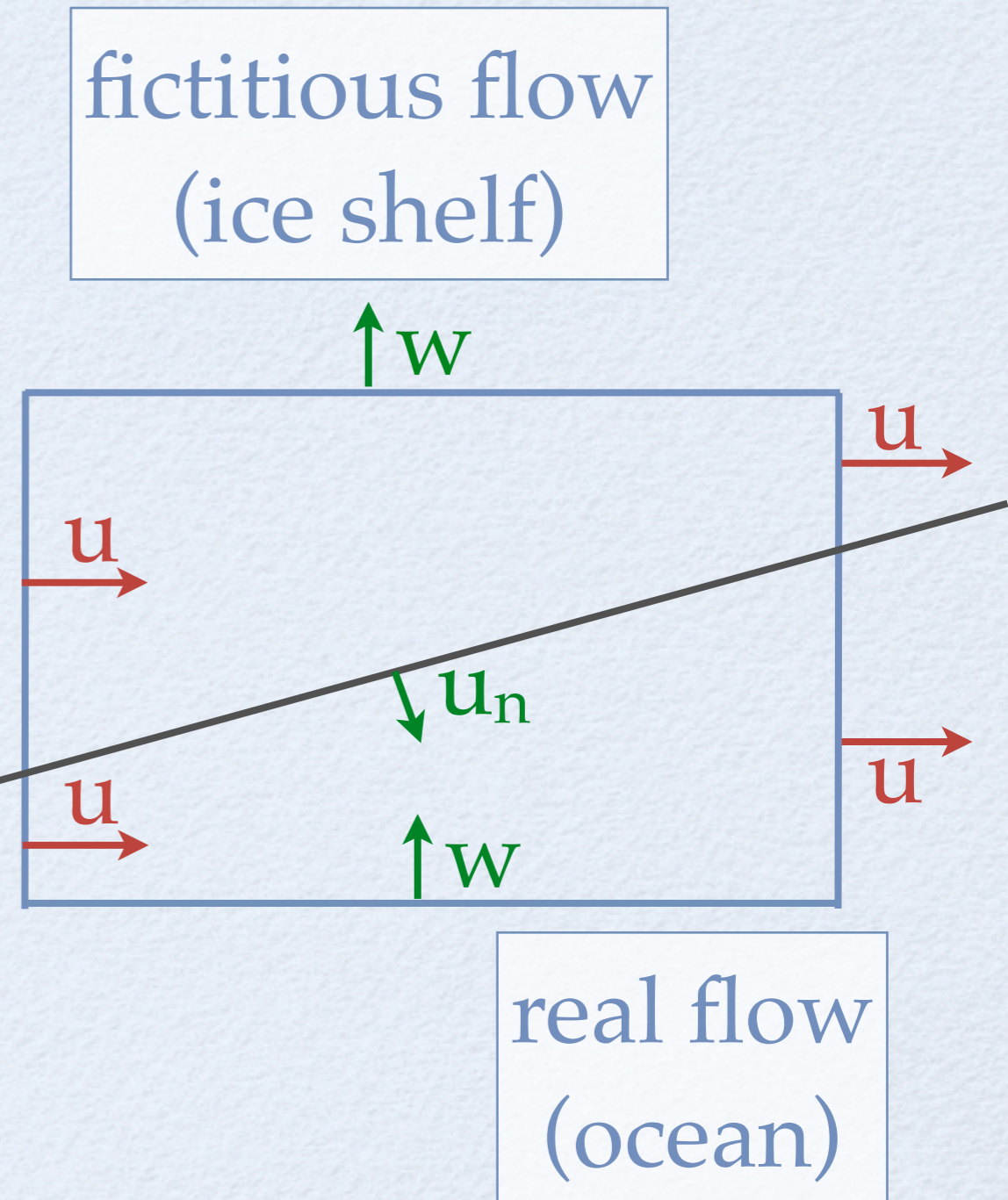
- Partial cells do not solve the problem
- Interface is flat but **fluxes areas** are still different on left and right
- The flux areas are the same for the full box: u_n is not an interpolation of w and w



CHALLENGES

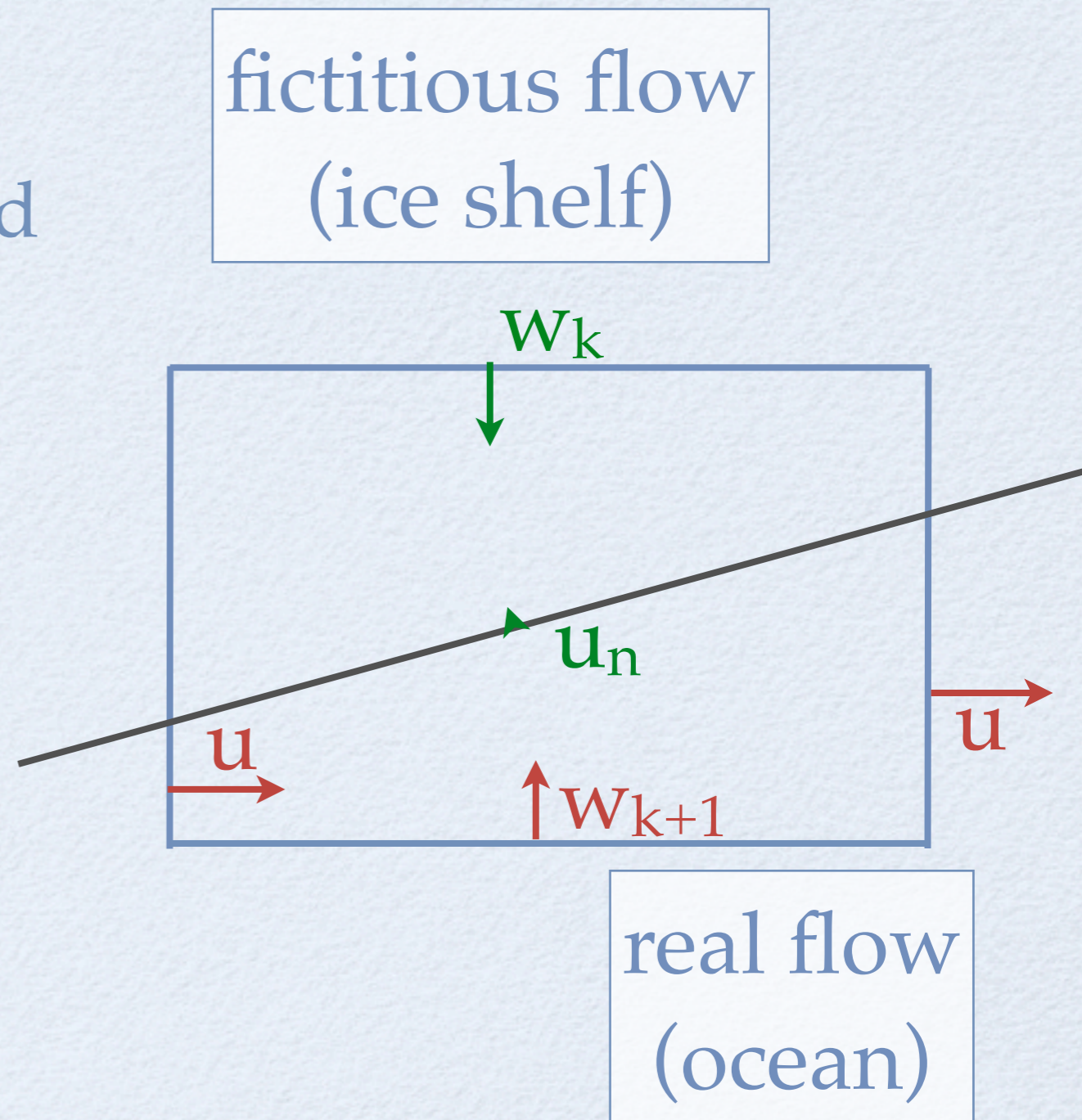
Continuity:

- Separate u 's for real and fictitious parts of cut cells would solve *this* problem
- But this would require a **complete redesign** of POP to handle the fictitious flow



CHALLENGES

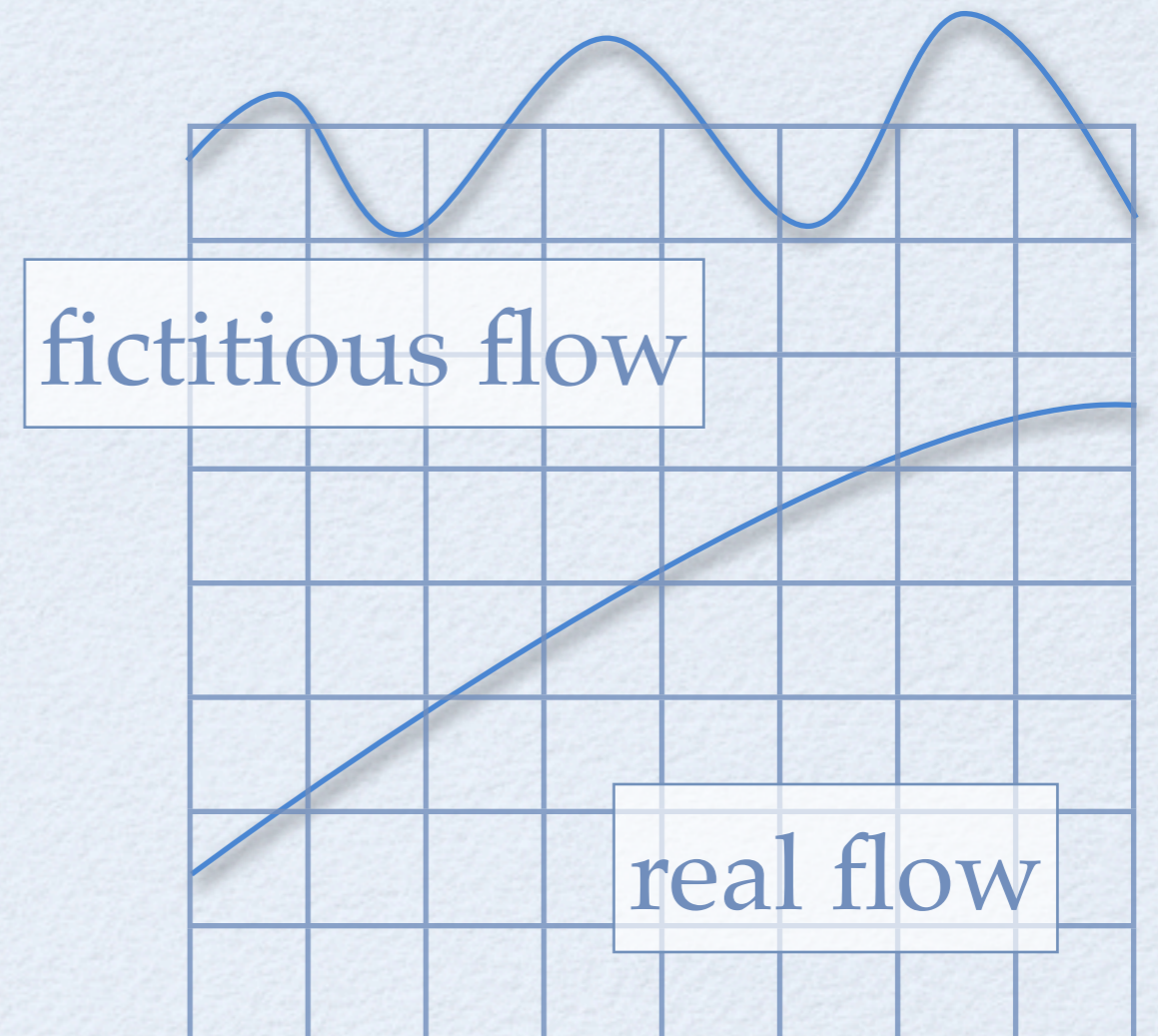
- w_k could be computed so that interpolation of w_k and w_{k+1} gives $u_n \approx 0$
- Allow divergence (ignore continuity) in fictitious flow
- But leads to *very* noisy w
- Wildly unphysical tracer advection in fictitious flow



CHALLENGES

Fictitious surface height:

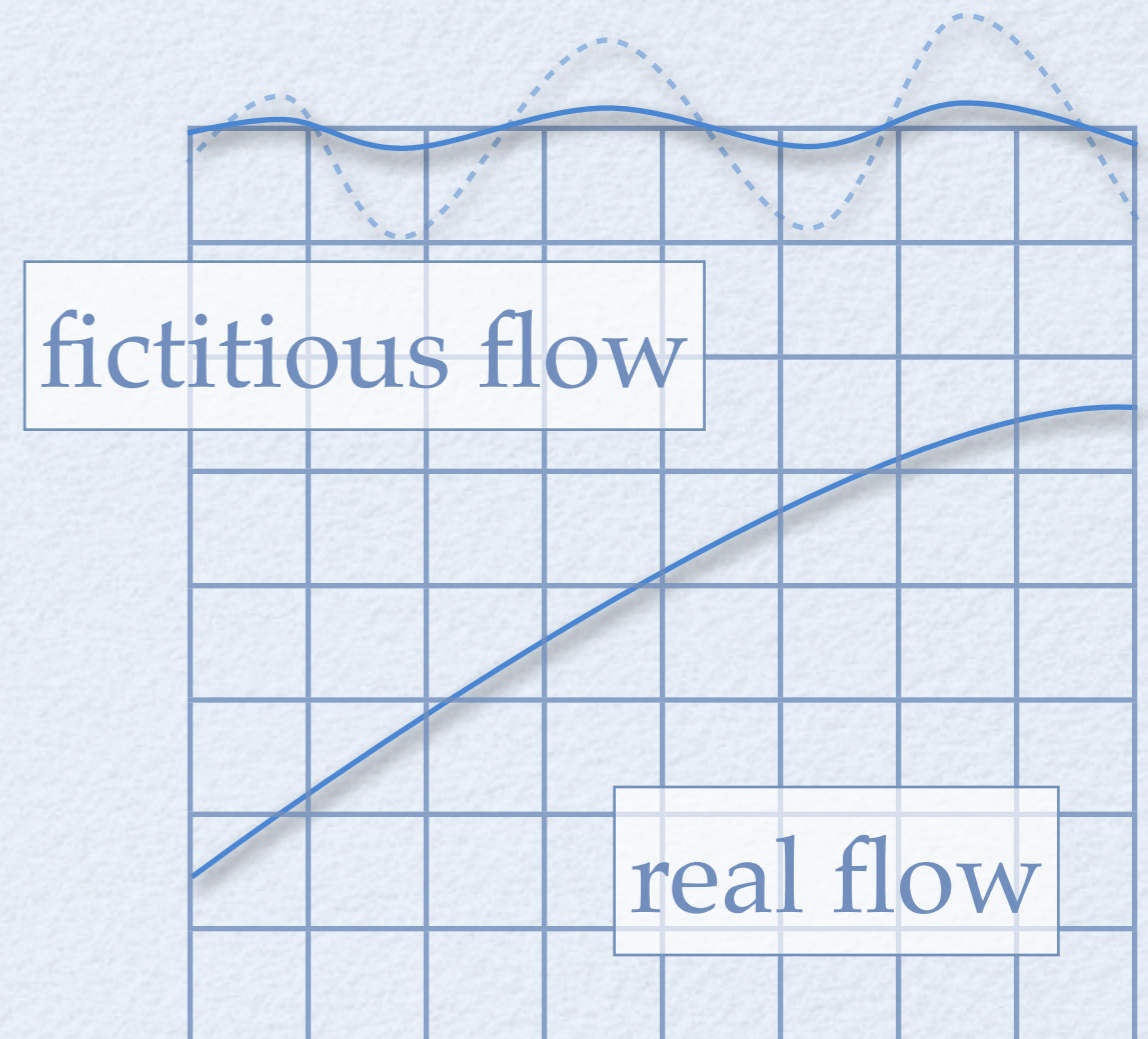
- **not** constrained by barotropic solver



CHALLENGES

Fictitious surface height:

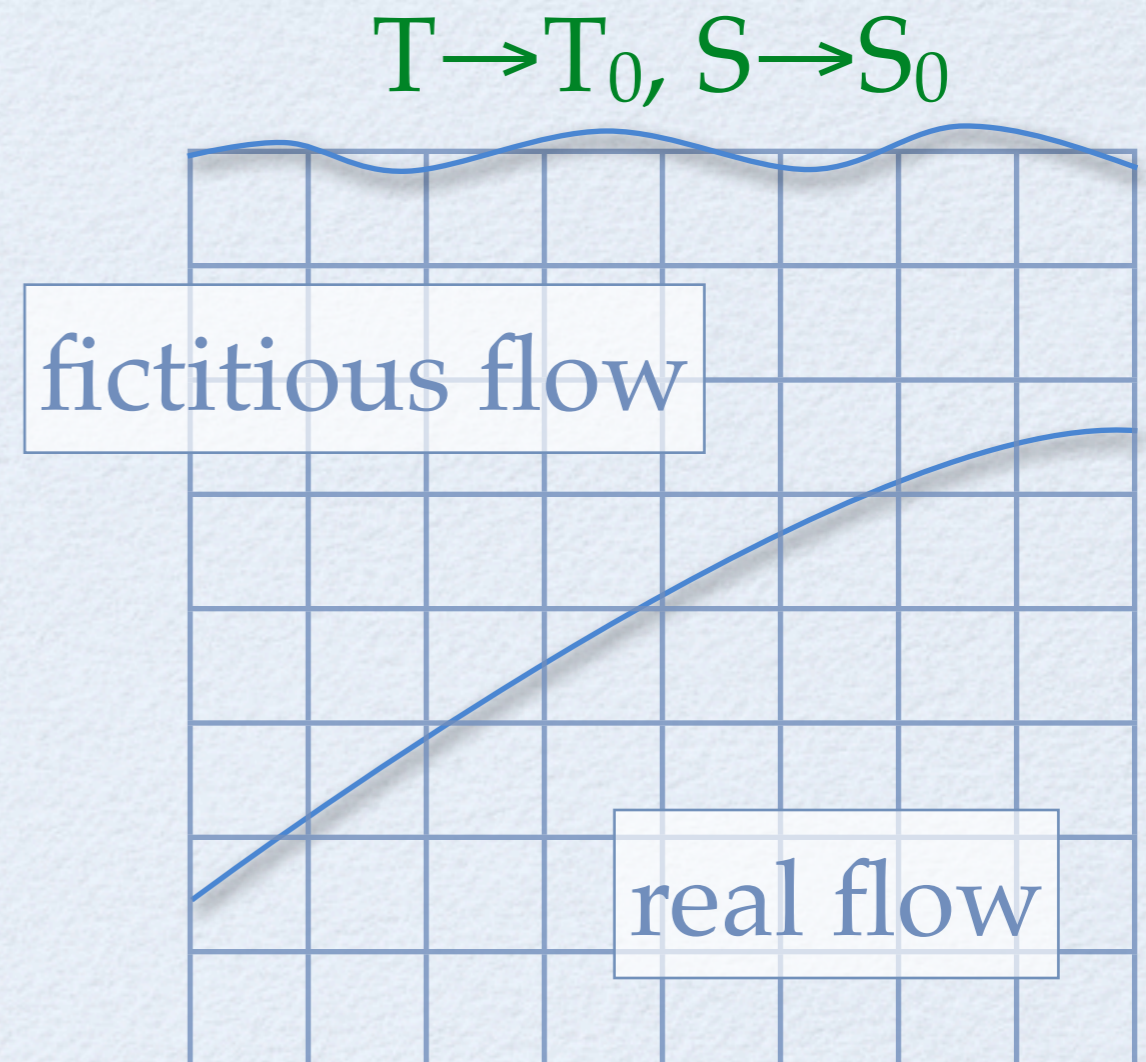
- **not** constrained by barotropic solver
- can be constrained by modifying fictitious horizontal velocities



CHALLENGES

Fictitious surface height:

- **not** constrained by barotropic solver
- can be constrained by modifying fictitious horizontal velocities
- Instability unless T/S are restored at the surface

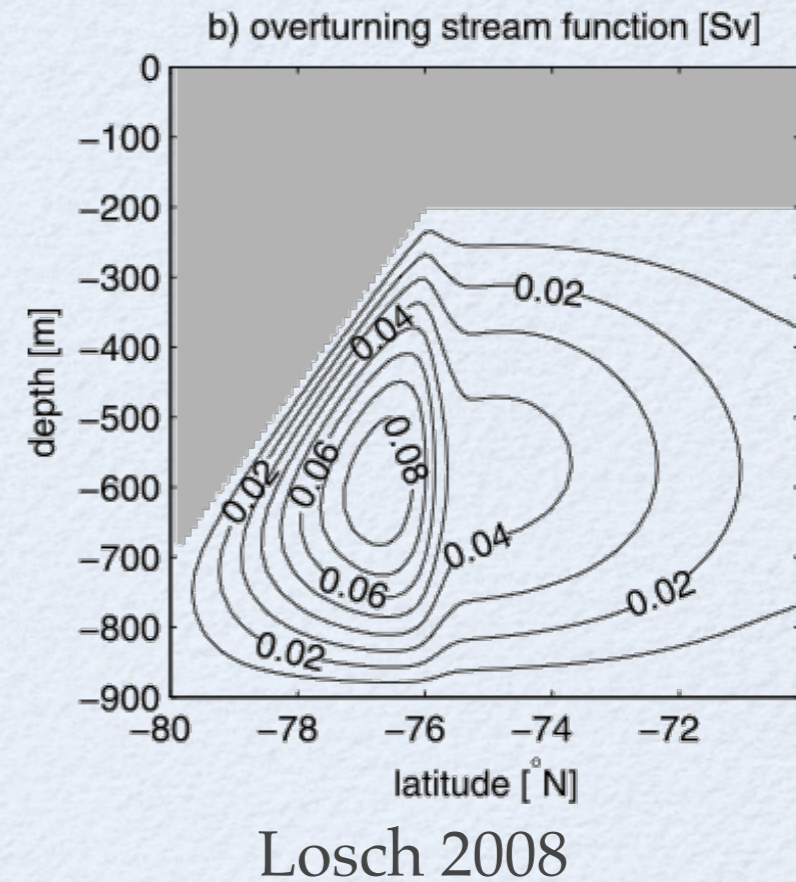


CHALLENGES

It seems that the fictitious flow
isn't worth the trouble!

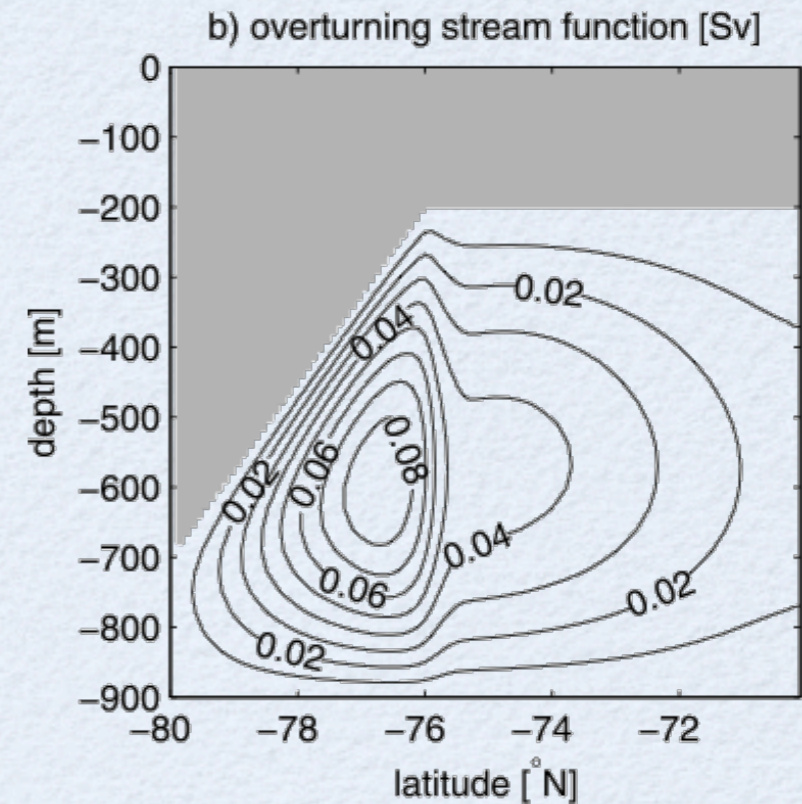
A NEW APPROACH

- Interface by partial cells
- No fictitious flow
- Based on Losch 2008: static ice shelves in MITgcm

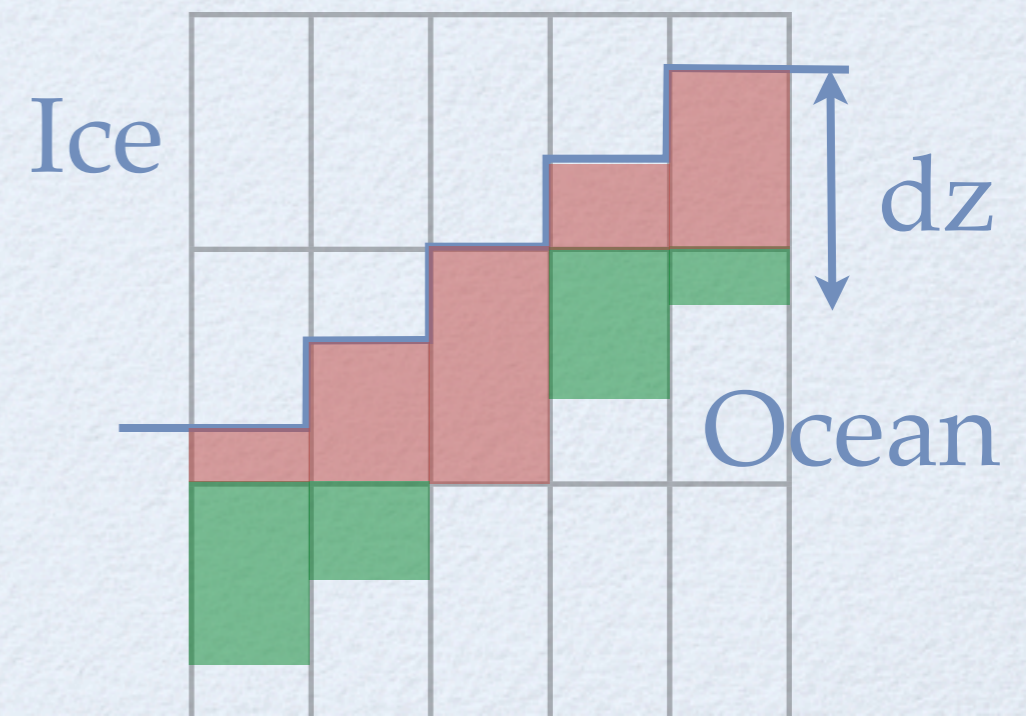


A NEW APPROACH

- Interface by partial cells
- No fictitious flow
- Based on Losch 2008: static ice shelves in MITgcm
- Salt/heat from melting/freezing mixes into both **partial cell** and **next cell below** (reduces noise)

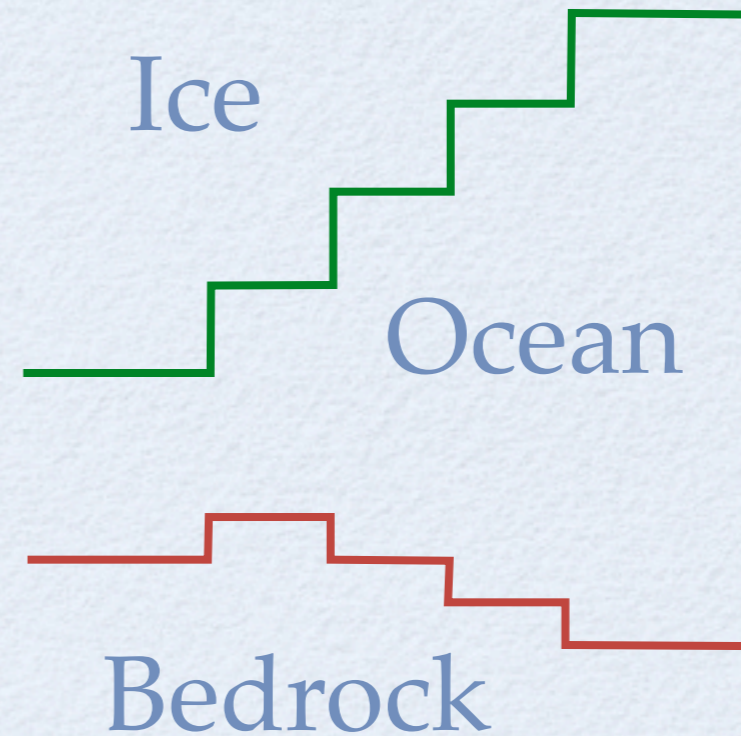


Losch 2008



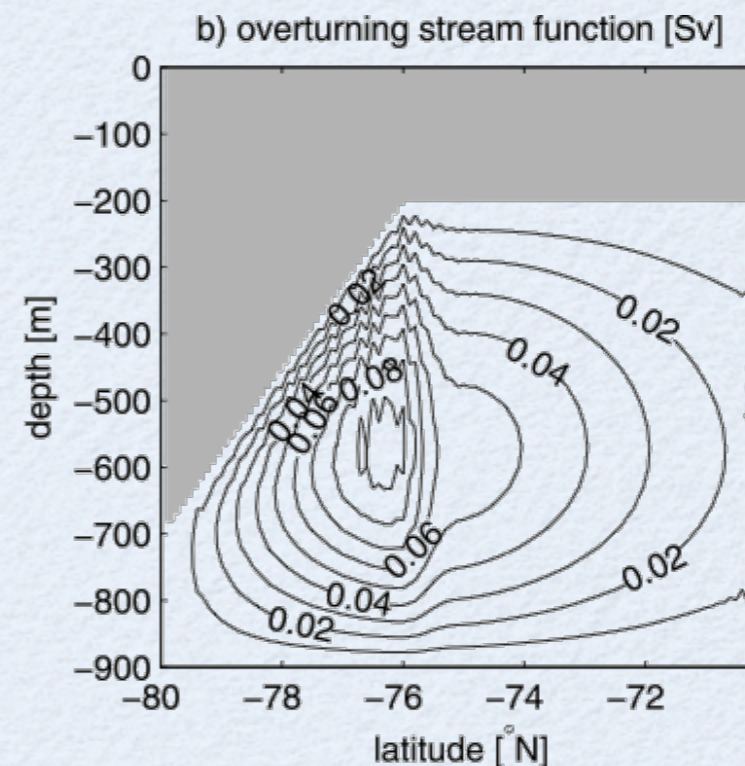
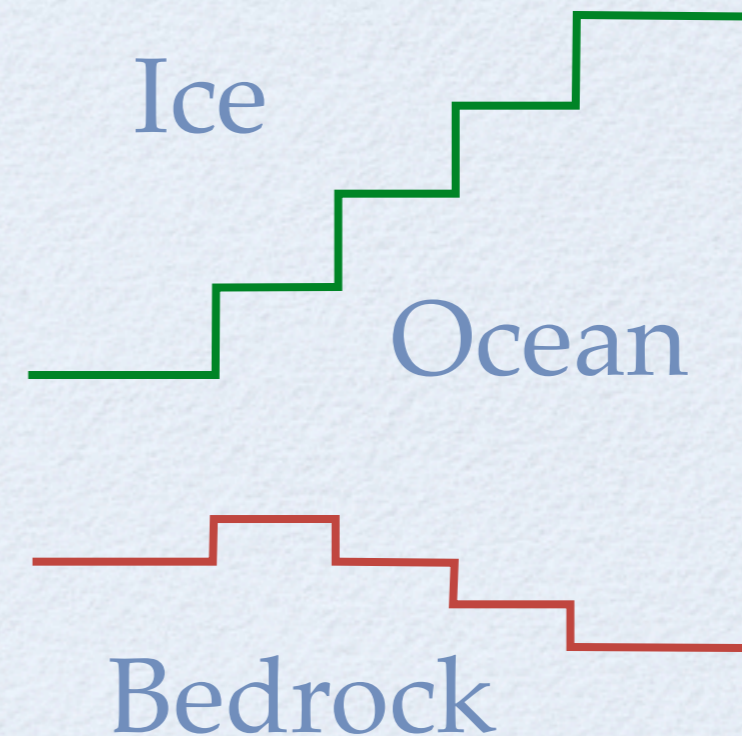
A NEW APPROACH

- Pros:
 - Static **interface** tested with other ocean models
 - Similar to **bathymetry**
 - Same boundary conditions as IBM



A NEW APPROACH

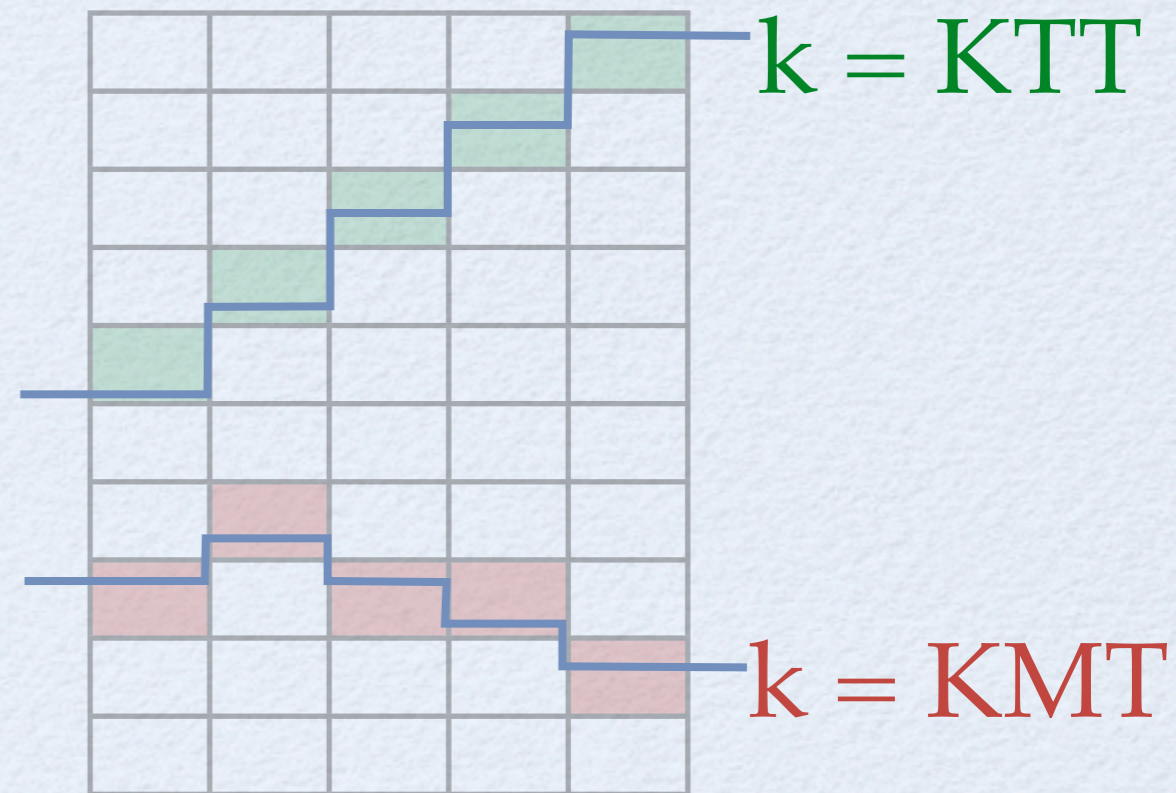
- Pros:
 - Static **interface** tested with other ocean models
 - Similar to **bathymetry**
 - Same boundary conditions as IBM
- Cons:
 - Tested only for static ice shelves
 - Stair-step geometry can lead to noisy fields
 - Pervasive modifications



Losch 2008

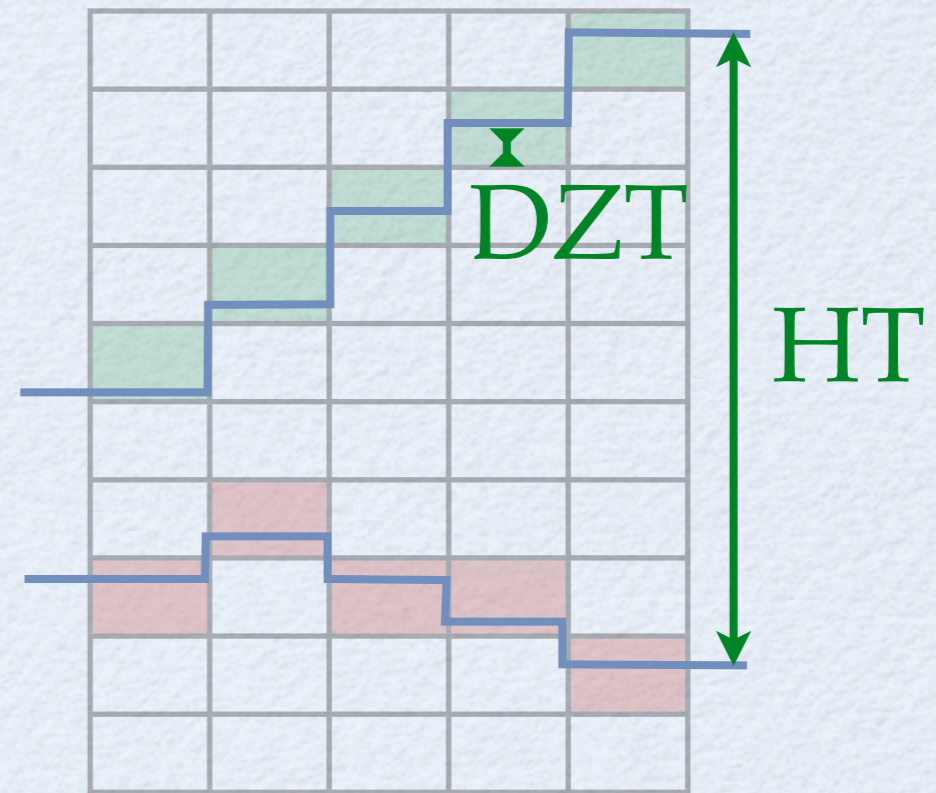
NEW APPROACH: IMPLEMENTATION

- Vertical indexing in POP from **KTT** to **KMT** (rather than 1 to KMT) on T grid, similar on U grid



NEW APPROACH: IMPLEMENTATION

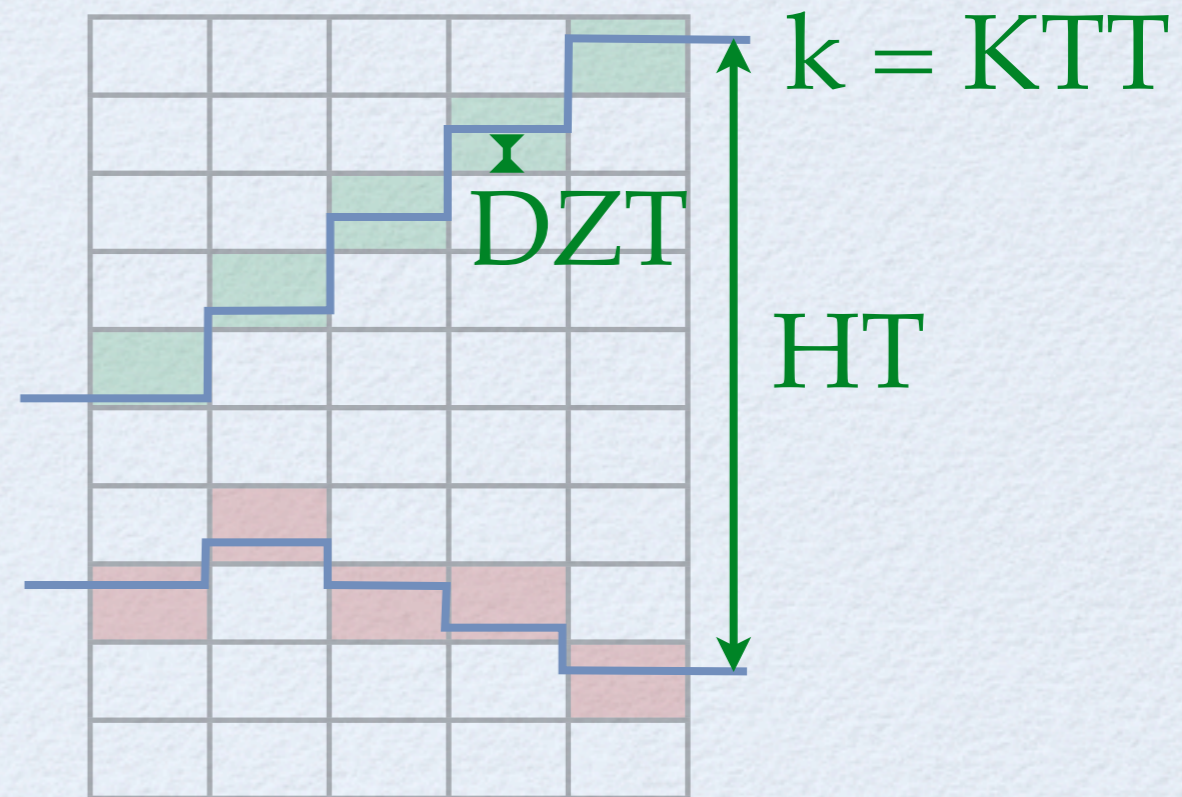
- Vertical indexing in POP from **KTT** to **KMT** (rather than 1 to KMT) on T grid, similar on U grid
- Modifications to grid cell thickness **DZT/DZU** and total thickness **HT/HU**



NEW APPROACH: IMPLEMENTATION

As interface moves:

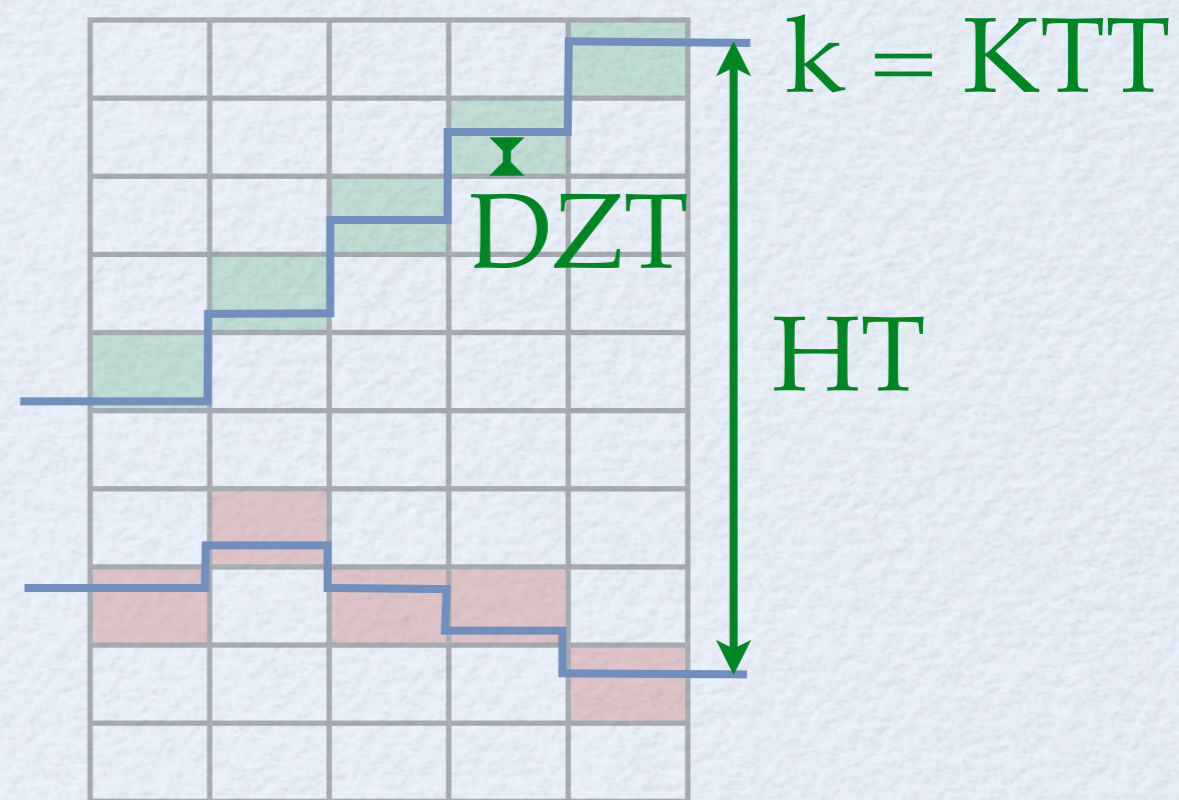
- Recompute KTT/KTU , DZT/DZU and HT/HU



NEW APPROACH: IMPLEMENTATION

As interface moves:

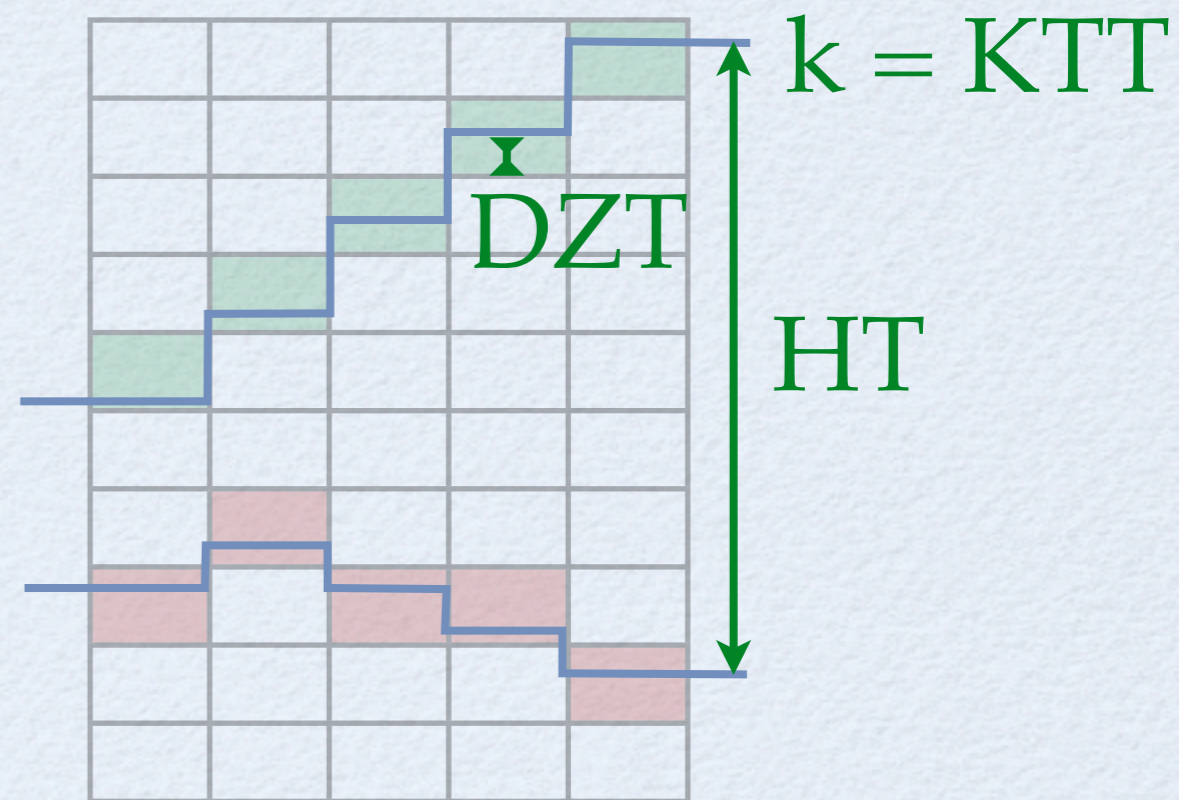
- Recompute KTT/KTU , DZT/DZU and HT/HU
- Reinitialize barotropic solver, flux-limiting advection coefficients



NEW APPROACH: IMPLEMENTATION

As interface moves:

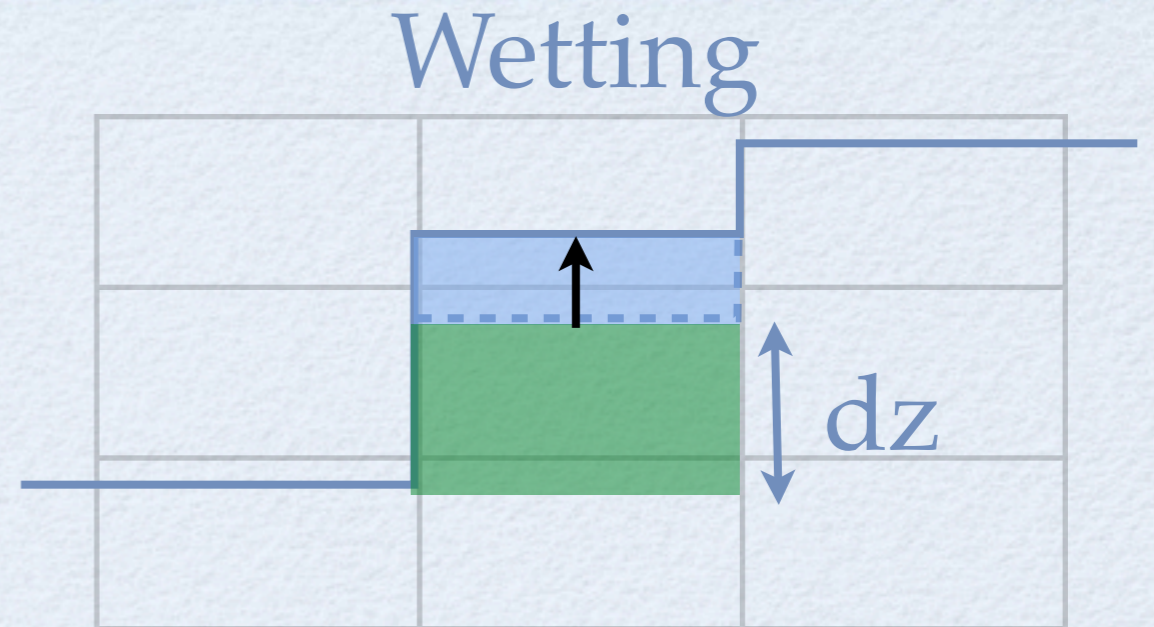
- Recompute KTT/KTU , DZT/DZU and HT/HU
- Reinitialize barotropic solver, flux-limiting advection coefficients
- Account for mass fluxes



NEW APPROACH: IMPLEMENTATION

“Wetting” and “drying” of cells:

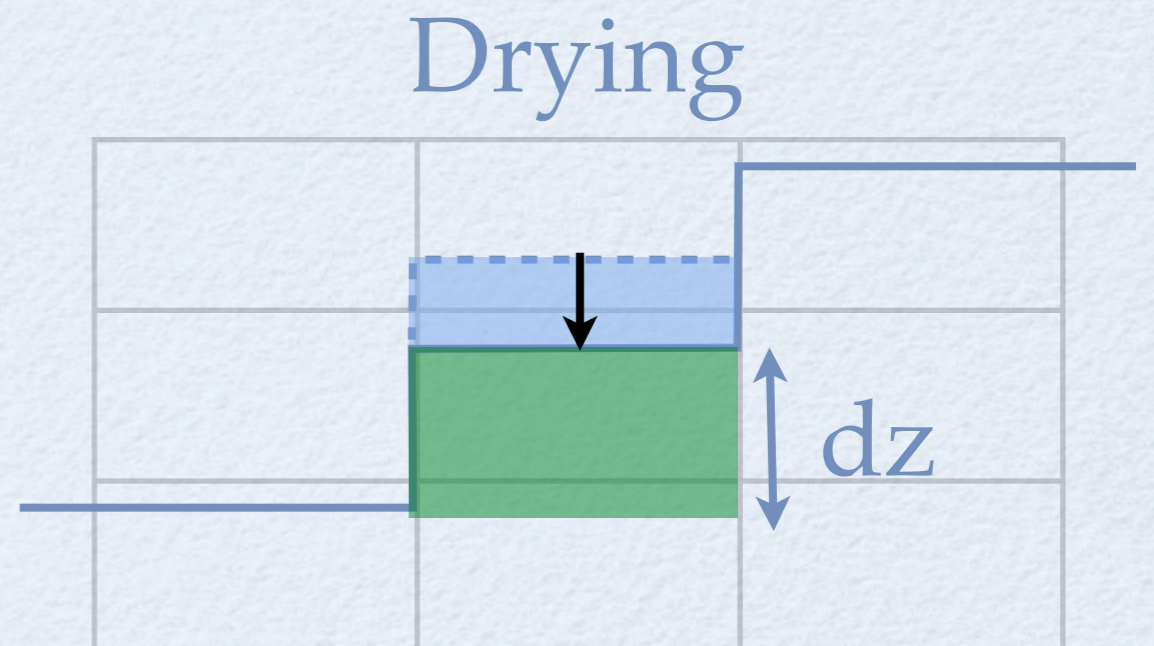
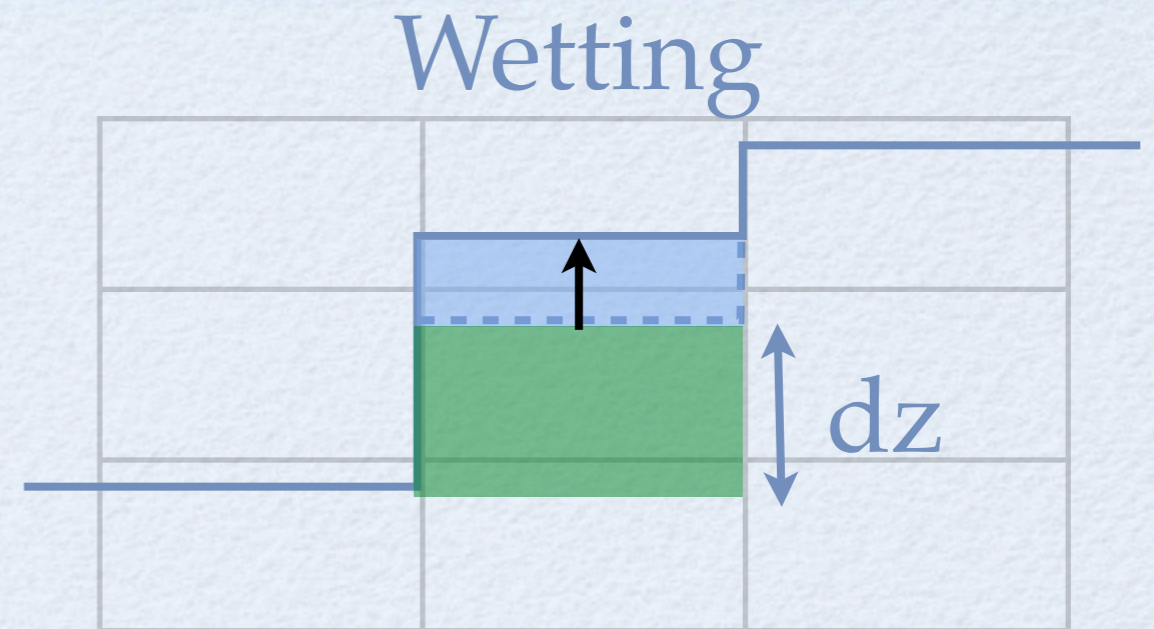
- Tracers in new “wetted” cells conservatively distributed *from* neighboring cell(s)



NEW APPROACH: IMPLEMENTATION

“Wetting” and “drying” of cells:

- Tracers in new “wetted” cells conservatively distributed *from* neighboring cell(s)
- Tracers in old “dried” cells conservatively distributed *to* neighbor(s)



SUMMARY

- Unsolved problems with Immersed Boundary Method in POP (and probably ocean models generally)
- New approach to land-ice / ocean interface similar to partial bottom cells (no fictitious flow)
- Requires substantial changes to indexing in POP
- New methods required to handle “wetting” and “drying” of grid cells as boundary moves