

# Surface mass balance of the Greenland Ice Sheet simulated with CESM

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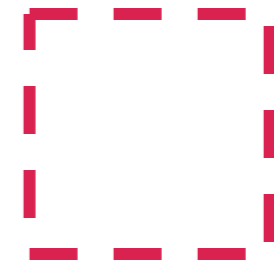
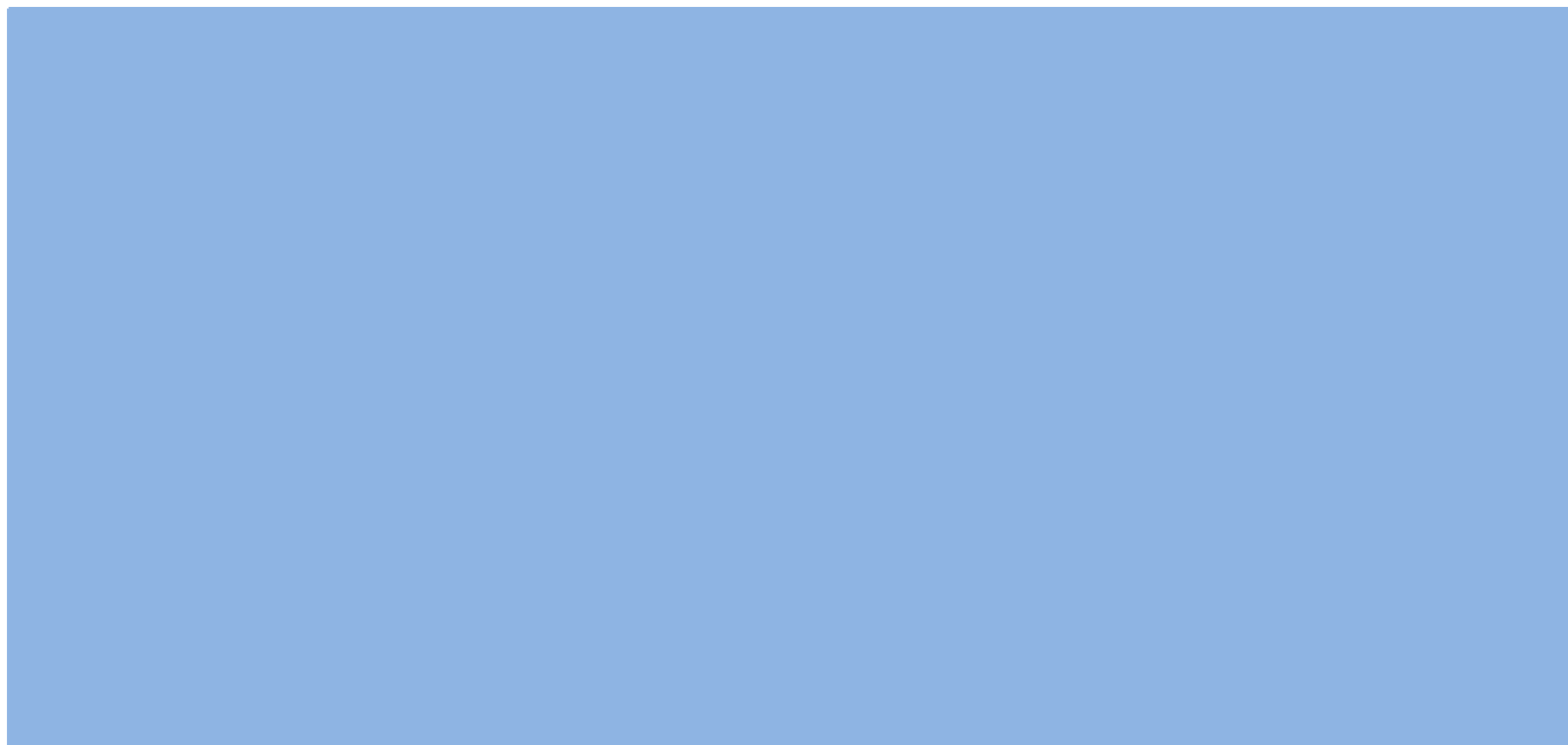
# overview

1. CLM simulations with MOAR forcing
  - a) 20<sup>th</sup> century SMB
  - b) 21<sup>st</sup> century projections (RCP8.5)
2. Sensitivity tests with the fully coupled model (“year 2000” climate)
  - a) Ice albedo
  - b) Resolution
3. Conclusions and future plans

# Model and set-up: IG-MOAR runs

Set-up: land model is run at  $1^\circ$  with data atmosphere from a previous 20<sup>th</sup> century run (MOAR). SMB is calculated at 10 elevation classes and downscaled to finer grid (10 km). Ice dynamics off.

Validation: regional model RACMO (forced by ERA-40/ECMWF operational 1958-2008; *Ettema et al. GRL, 2008*). 11 km resolution.

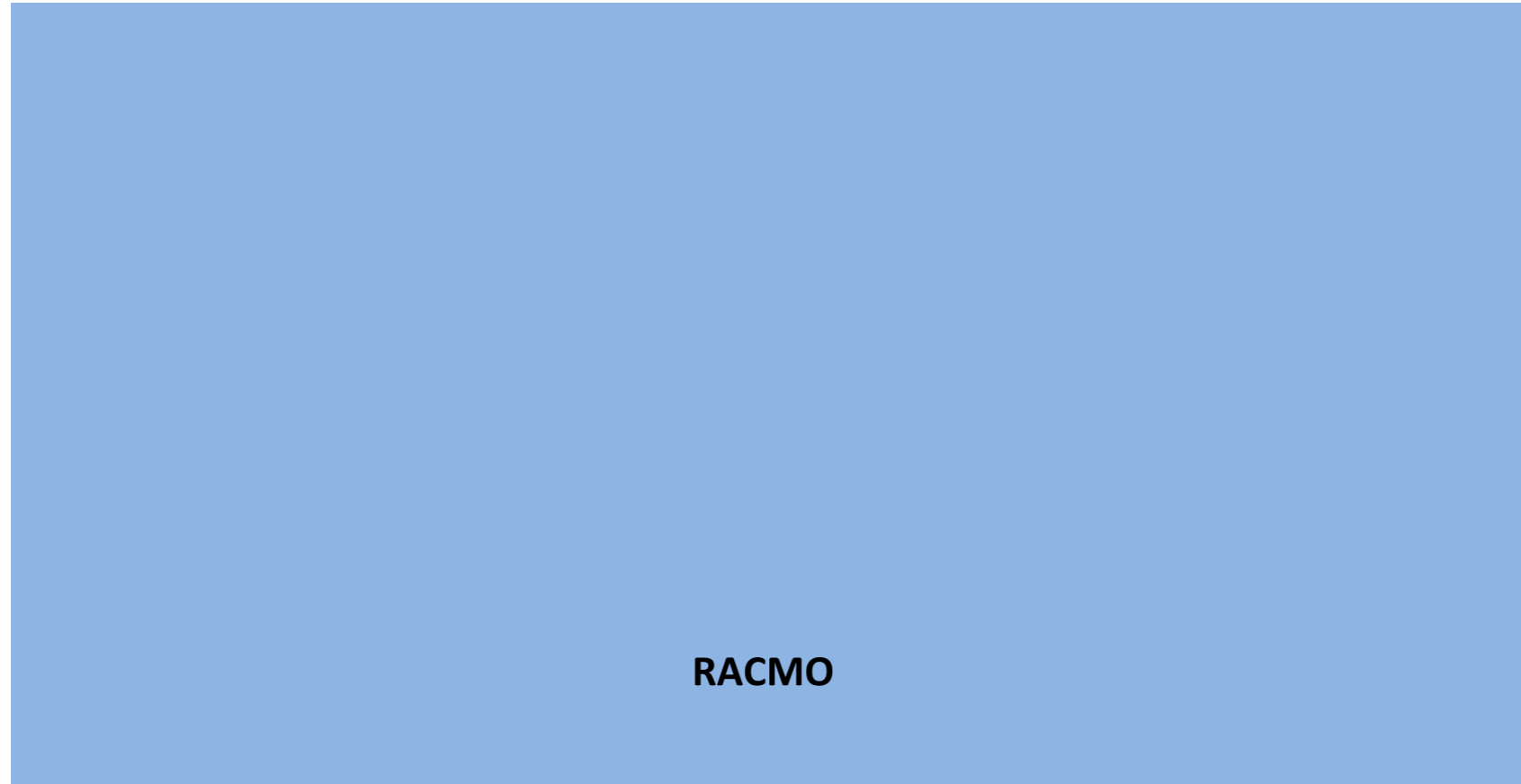


# Simulated climate 1958-2005





# Simulated climate 1958-2005



- Two bands of max. precip along SW and E margins are well captured
- Precipitation is overestimated in N interior and underestimated in SE
- Near surface climate agrees well with RACMO, with some discrepancies
  - July albedo is too low in W margin and N tundra regions. The second is probably due to overestimation of prescribed glacier extent
  - Incoming LW is higher in winter (RACMO underestimates it), possibly improving LW forcing

# Simulated surface mass balance

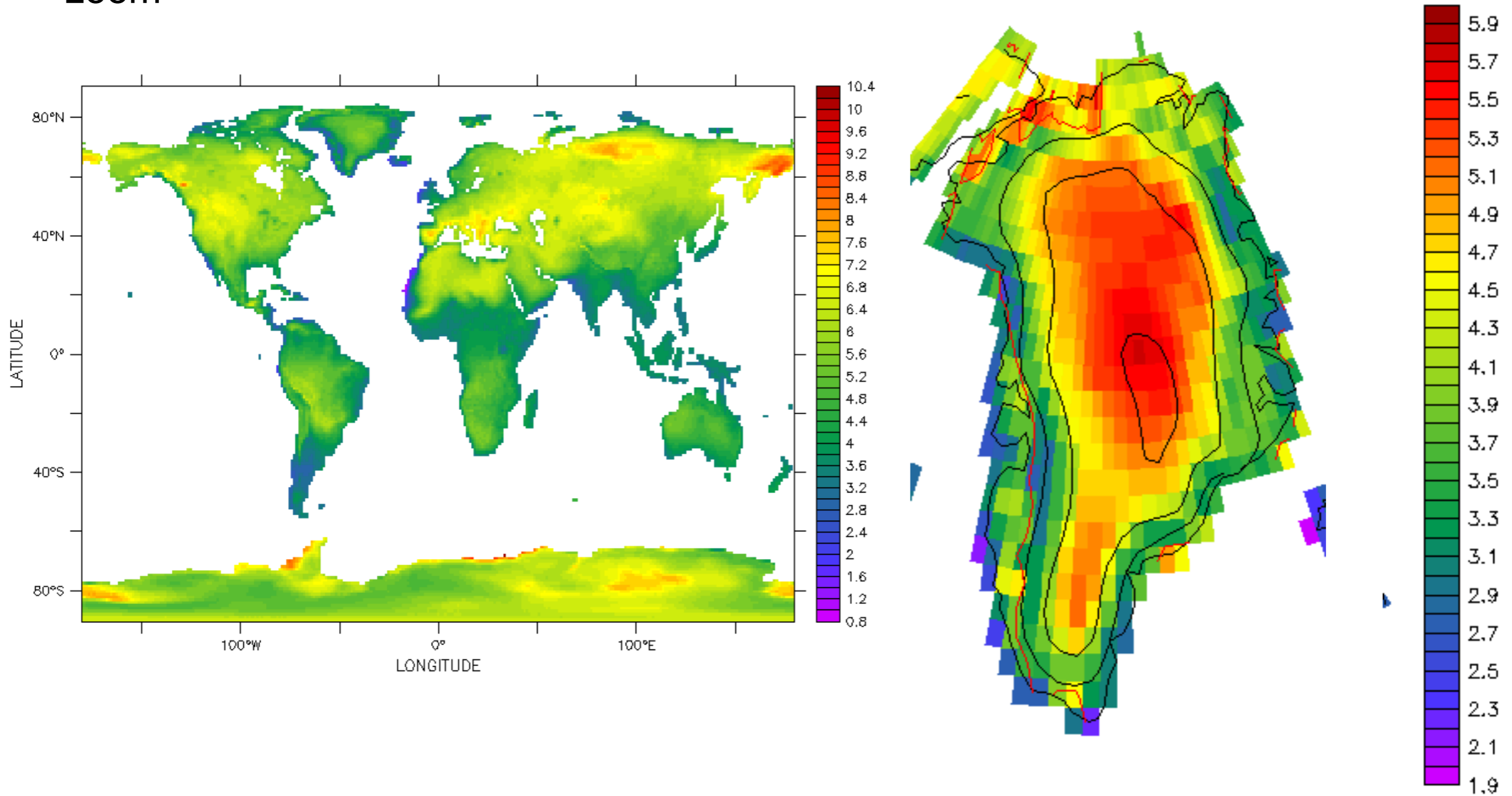


# Simulated surface mass balance



# 21<sup>st</sup> century projections (RCP8.5)

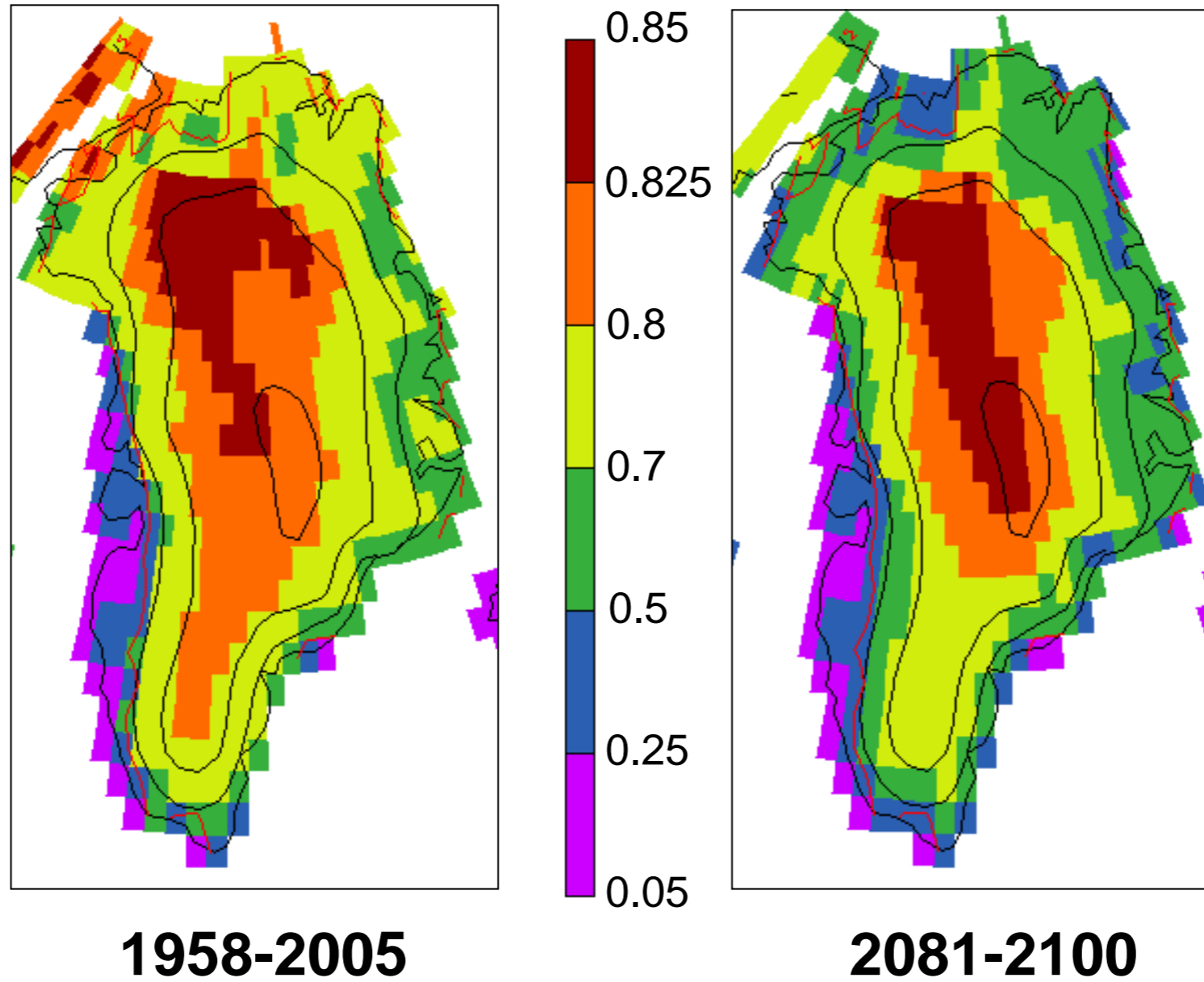
Summer near-surface temperature increase 2081-2100 wrt 1958-2005 [K], global and zoom



# 21<sup>st</sup> century runs: note of caution

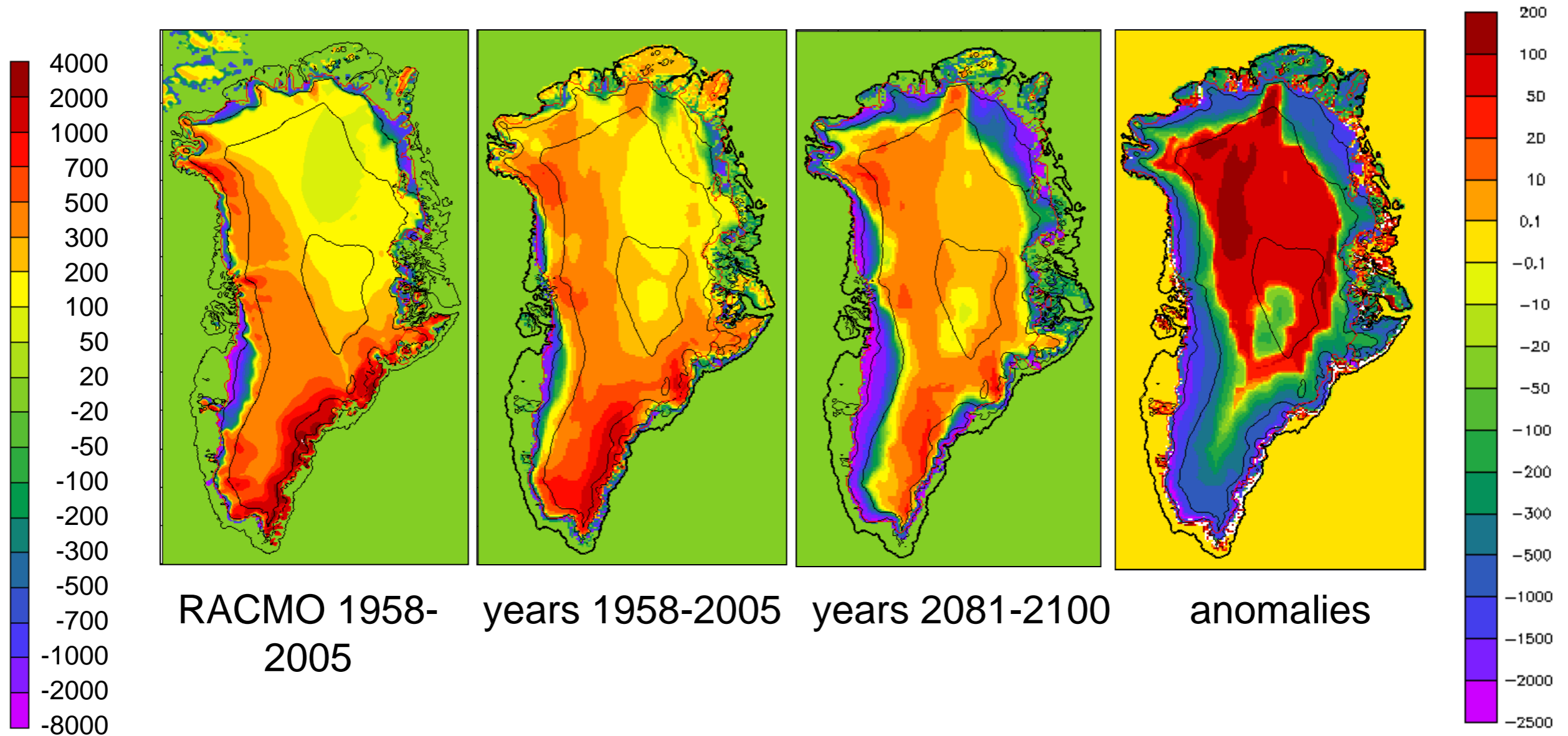
- Ice dynamics are off.
- Here we focus only on surface mass balance change (currently half of mass loss is due to glacier calving)
- **no SMB-height feedback** (increase of melt with decreasing elevation)!

# July albedo (RCP8.5)



# Surface mass balance RCP8.5 (kg m<sup>-2</sup>)

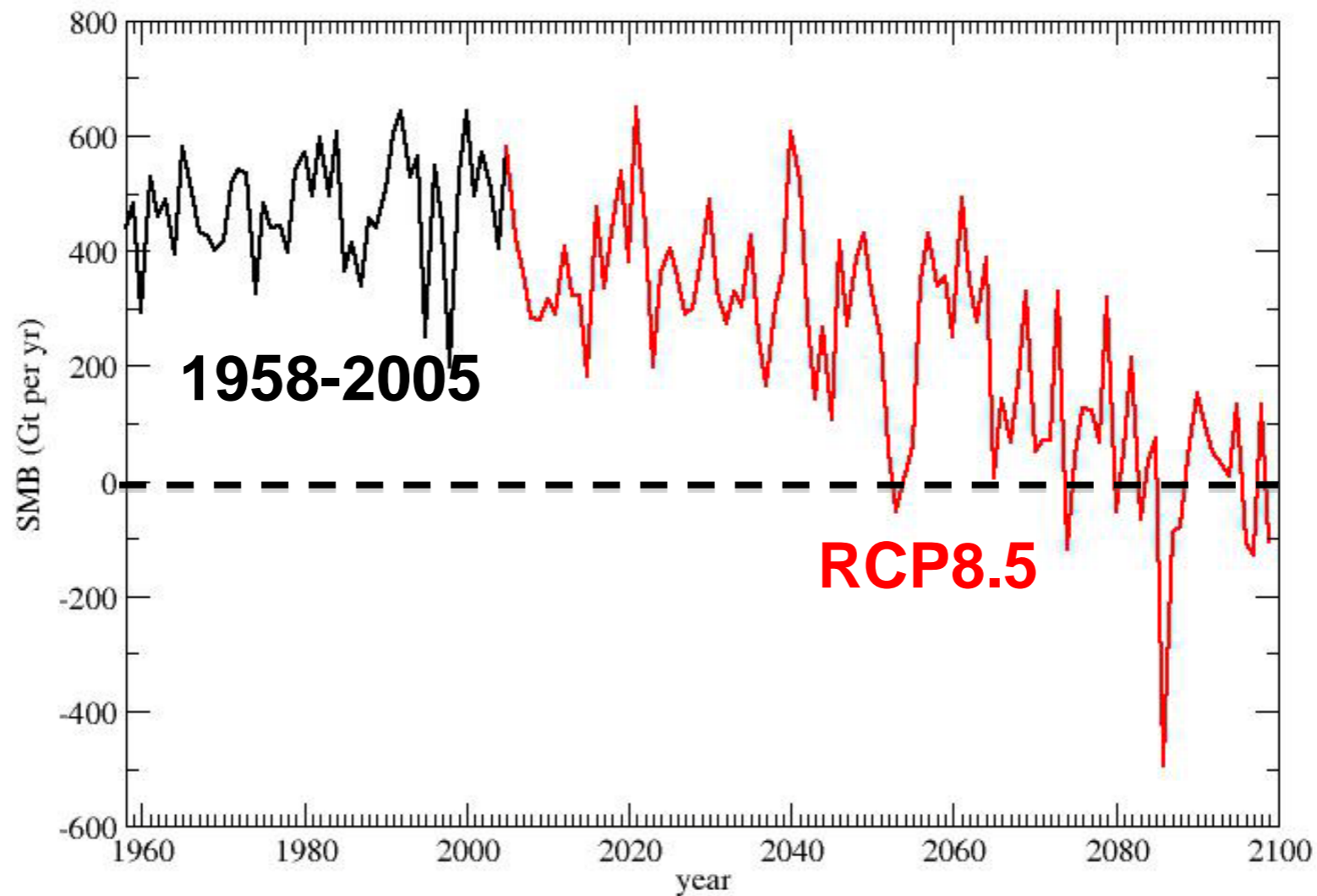
Net surface mass balance at 10 km resolution



1. CLM with MOAR forcing → b) 21<sup>st</sup> century SMB

# SMB time evolution

Time series of downscaled (10 km) net surface mass balance [Gt yr<sup>-1</sup>] integrated over Greenland ice sheet





# Total SMB, FV1 grid (ice sheet + glaciers)

Surface mass balance terms integrated over ice sheet and ice caps [Gt yr<sup>-1</sup>]. Total area is  $2.019 \times 10^6$  km<sup>2</sup>. RACMO values (*Ettema et al., GRL, 2009*) are listed for the sake of validation. Stds are given in parenthesis

Variable	RACMO 1958-2005	1958-2005	2081-2100
Precip	743 (78)	974 (105)	1275 (81)
Rain & rain frac	46 0.06	136 (26) 0.14	321 (41) 0.25
Sublim	26 (3)	66 (4)	81 (4)
SMB	469 (107)	380 (125)	-424 (225)

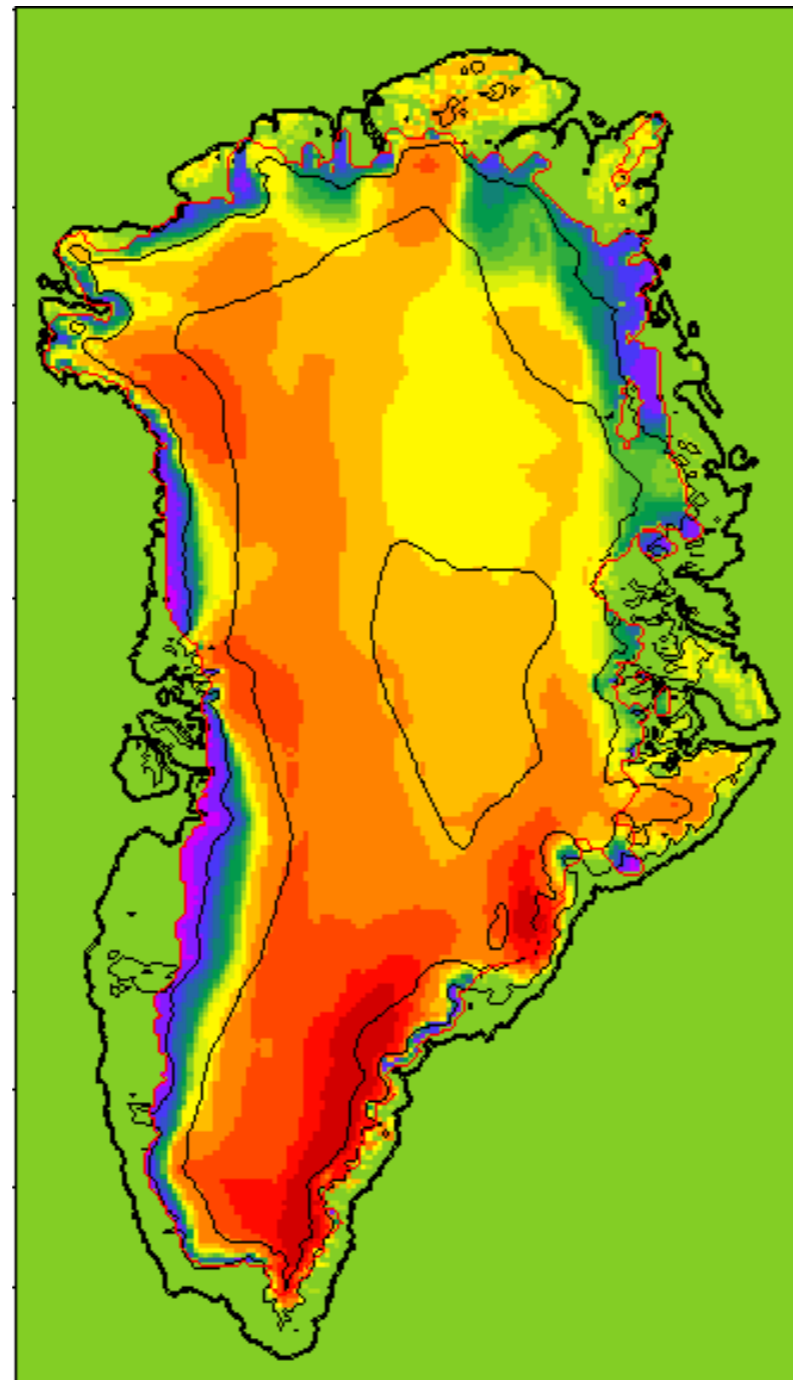
# Summary

- We present results from simulations with the land model of the Community Earth System Model forced with a data atmosphere from a previous CESM run for the 21<sup>st</sup> century (2005-2100, RCP8.5). The Greenland surface mass balance is downscaled to 10 km resolution at the heights given by *Bamber et al. (2001)*. A simulation with the same setup but forced with 20<sup>th</sup> century data from CESM is used as reference for 21<sup>st</sup> century climate change.
- By the end of the 21<sup>st</sup> century, ablation has increased with similar magnitude **at all margins**. The equilibrium line has moved **to 2000 m** at all margins except the NW, where its height is close to 1500 m.
- Precipitation rates over glaciated areas of Greenland increase **by 31%** from 1958-2005 to 2081-2100. The annual rain fraction increases from 14% to 25%.
- The increase in near-surface temperatures is highest in the interior of the ice sheet (~6 K above 2000 m). Areas at elevations between 1000 and 2000 m are warmer by ~4.5 K, and below 1000 m, by ~3 K.
- Outlook: Simulations will be repeated with the coupled atmosphere-ocean-land model

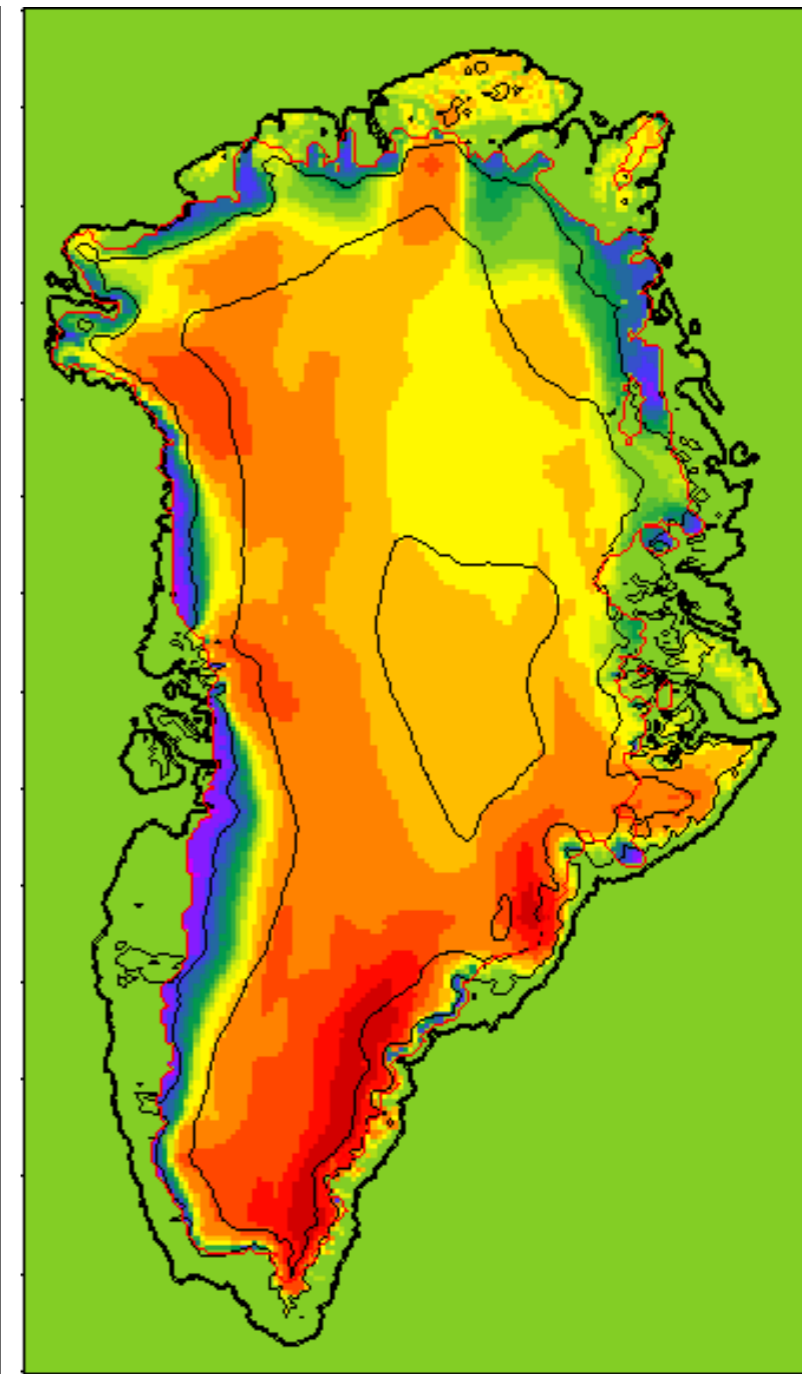
# Ice albedo sensitivity

- Increasing ice albedo does not change much the width of ablation areas, because it depends on whether the snowpack is melted (depends primarily on snow albedo)
- It reduces melting at the highest melting areas, because they have long ablation season

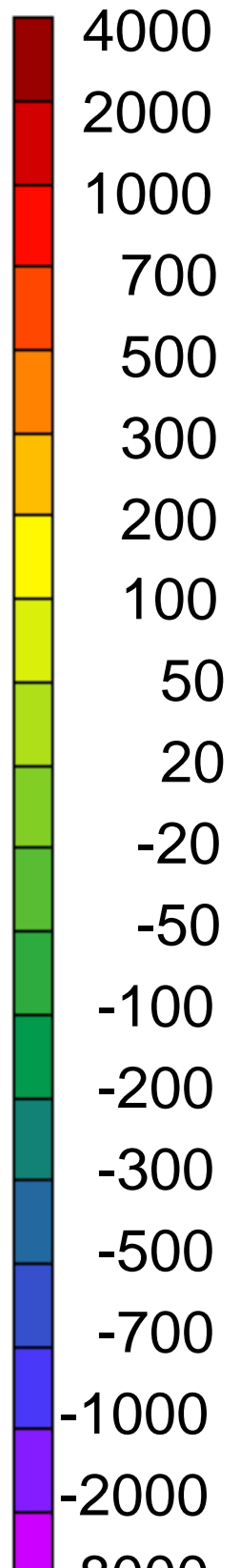
Contours:  
*Ice sheet margin,*  
1000, 2000, 3000 m topo



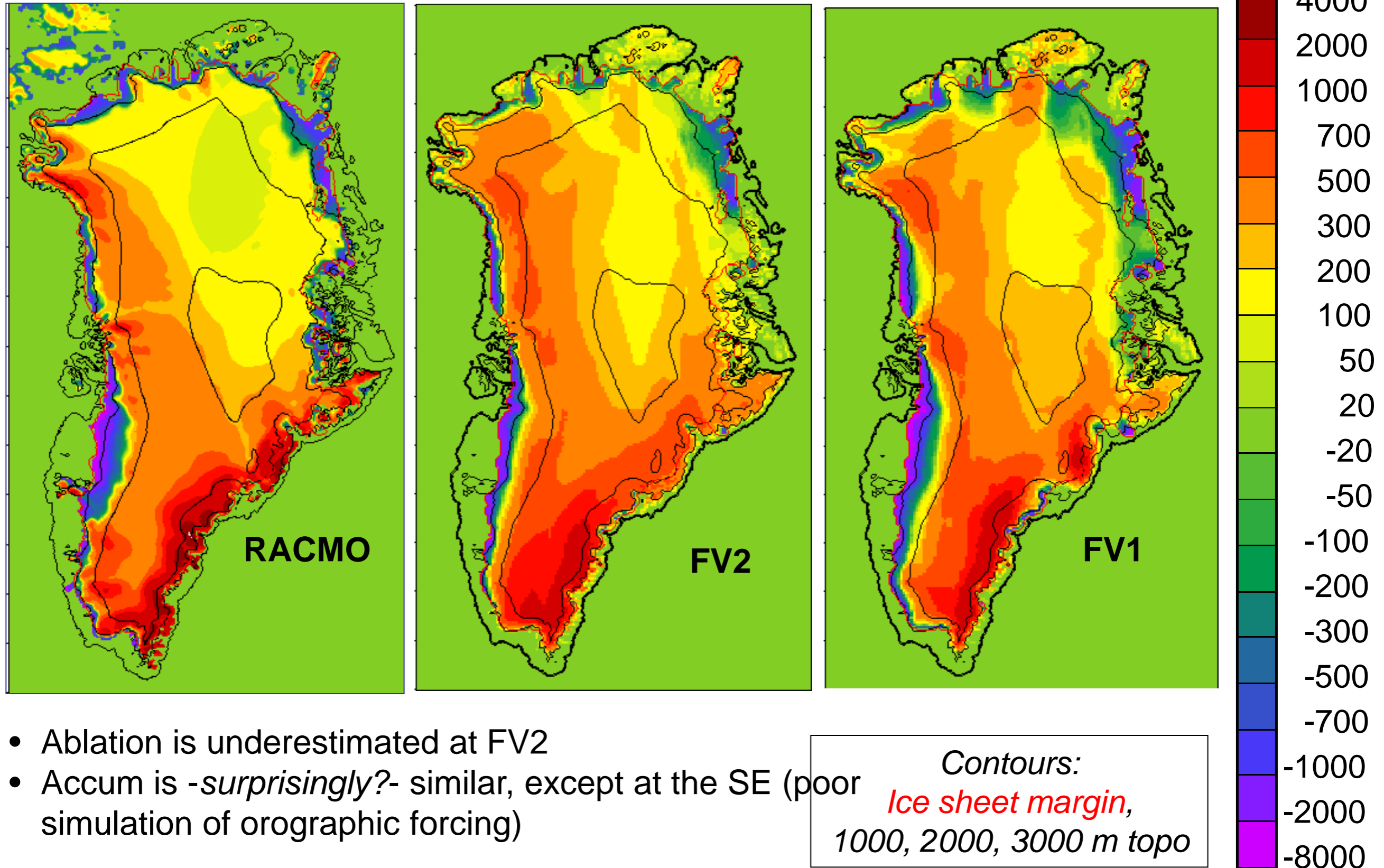
$alb_{vis}=0.60$   
 $alb_{nearIR}=0.40$



$alb_{vis}=0.75$   
 $alb_{nearIR}=0.50$



# Sensitivity to resolution: BG2000 FV1 vs FV2



- Ablation is underestimated at FV2
- Accum is *-surprisingly?-* similar, except at the SE (poor simulation of orographic forcing)

# Net SMB at 10 km resolution (kg m<sup>-2</sup>)

	2000-climate
FV1	<b>361 (79)</b>
FV2	<b>509 (67)</b>
FV1 high albedo (0.75/0.50)	<b>386 (98)</b>

For comparison, these numbers for pre-industrial SMB

Variable	10 km grid	Pre-ind CCSM4 FV1/FV1	Pre-ind CCSM4 FV1/FV2	RACMO	Other reg models (*)
SMB	<b>380 (97)</b>	429 (121)	315 (132)	469 (107)	288/356/287
Area (10 <sup>6</sup> km <sup>2</sup> )	<b>1.685</b>	2.131/2.01 9			

# Conclusions

- Climate and SMB compares well to state-of-the-art regional model RACMO
- Main biases possibly due to biased glacier mask
- First 21<sup>st</sup> century runs (only SMB) show increase of ELA height to 2000 m & SMB < 0 by end of century
- Relatively high sensitivity to model resolution
- Relatively low sensitivity to choice of ice albedo

# Ongoing work

- Improved glacier mask from Bamber et al.
- Spin-up of BG1850CN run (reference CMIP run)
- Next: CMIP runs (whole model with glacier elevation classes and carbon cycle)
  - 20<sup>th</sup> century
  - 21<sup>st</sup> century