

# Quasi-Biennial Oscillation in WACCM: Parameterization and Evaluation

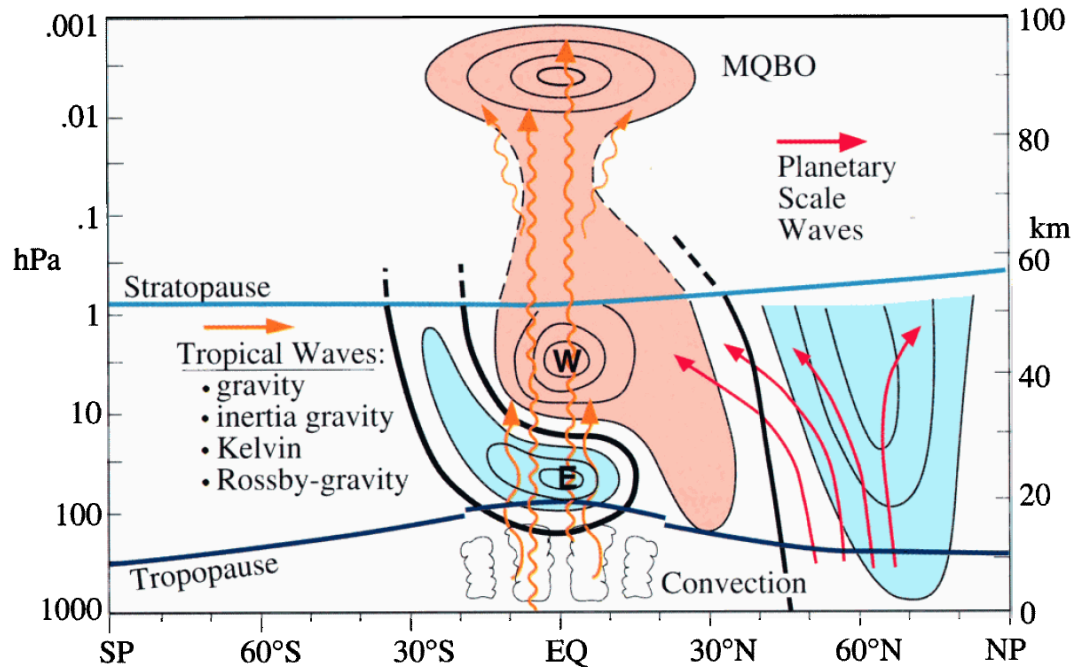
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# Overview

- Absence of QBO in WACCM and possible causes.
- Development of an inertio-gravity wave (IGW) parameterization scheme.
- Evaluation of WACCM simulations with the IGW.
  - Zonal mean wind, temperature, ozone at the equator.
  - Extratropical effects.
  - Surface signature.

# Possible Driving Forces of QBO



Baldwin et al, 2001

- $F(\text{GW/IGW})$  likely much larger than  $F(\text{PW})$
- PW (Kelvin waves, Rossby-gravity waves) resolved by WACCM (albeit weak).
- Mesoscale GW parameterized, breaking mainly in mesosphere.
- **IGW poorly resolved, and not parameterized.**

# Requirement for QBO Forcing

- QBO Acceleration rate:

- $50\text{m/s}/14\text{months} \sim 10^{-6}\text{m/s}^2$

- $Q \frac{\partial u}{\partial t} = -\frac{1}{\rho} \frac{\partial \tau}{\partial z} \approx -\frac{1}{\rho} \frac{\Delta \tau}{\Delta z}$

- $\rho_{\text{strat}} \sim 0.1\text{kg/m}^3, \Delta z \sim 10\text{km}$

- $\Delta \tau \sim 10^{-3}\text{Pa}$

- For GW with such momentum flux to break in the stratosphere, the horizontal wavelength is  $\sim 1000\text{ km}$  according to linear saturation theory.
- Also possible: intermittent mesoscale GW with large momentum flux (e.g.  $t \sim 10^{-2}\text{ Pa}$ , occurring 10% of the time). Not considered in this study.

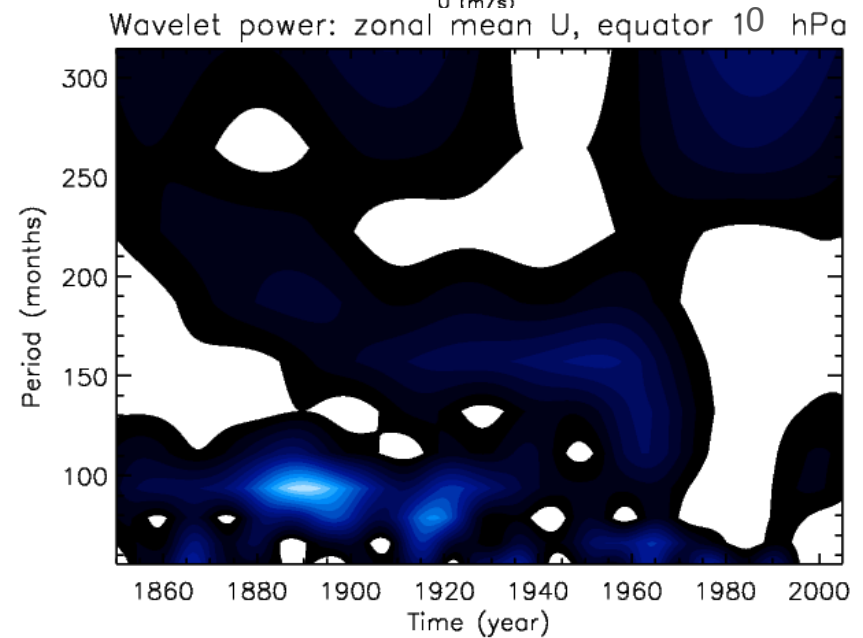
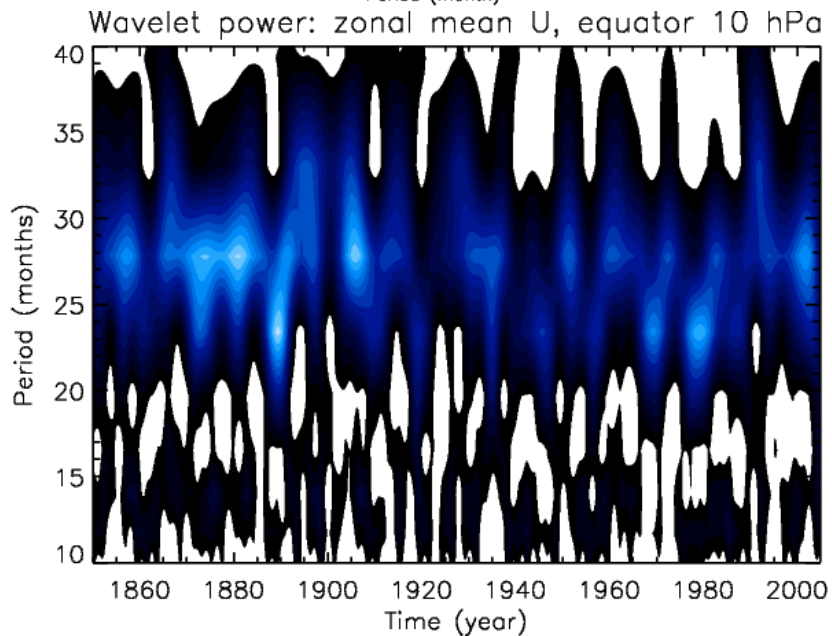
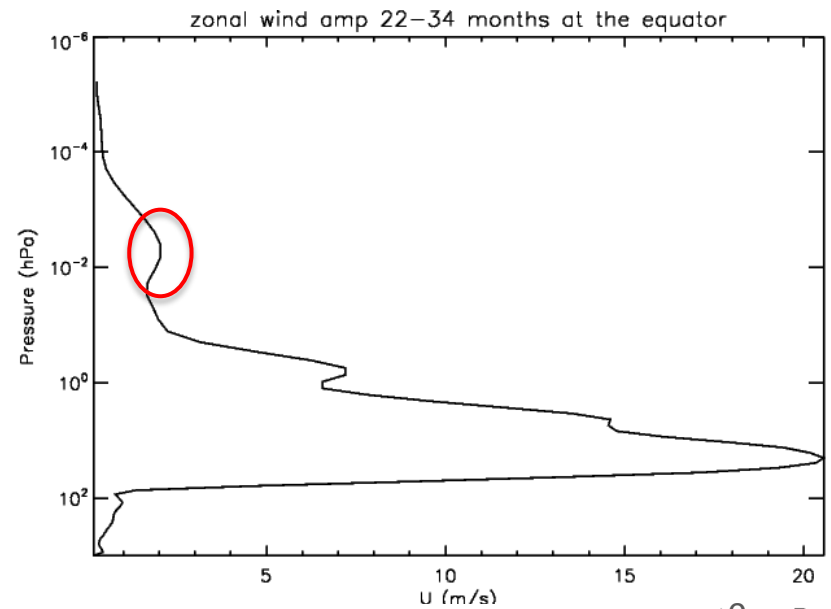
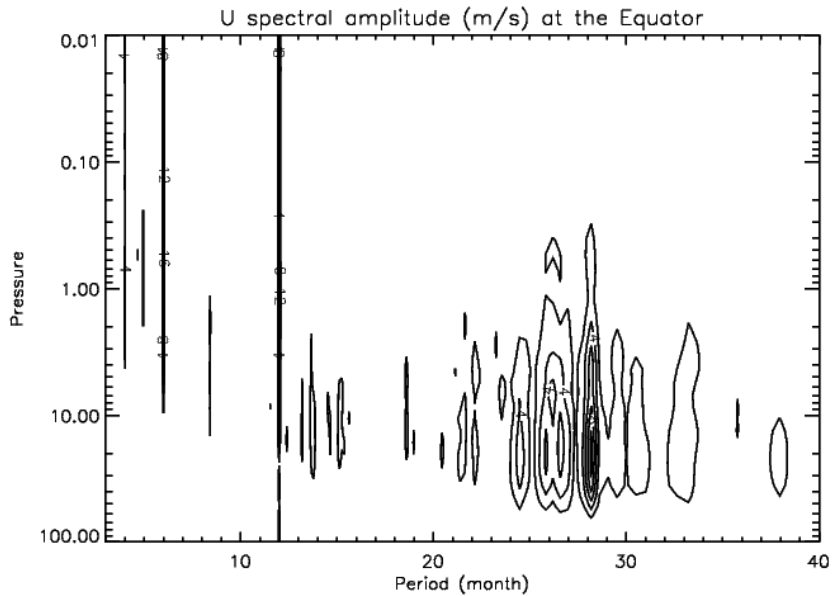
# Parameterizing Inertia-GW

- Similar to Lindzen (1981), though considering Coriolis effect (Xue et al., 2011).

$$\frac{\partial u}{\partial t} = -\frac{1}{\rho_0} \frac{\partial \tau^*}{\partial z} = \frac{k[(c-u)^2 - f^2/k^2]^{1/2}(c-u)^2}{2NH}$$

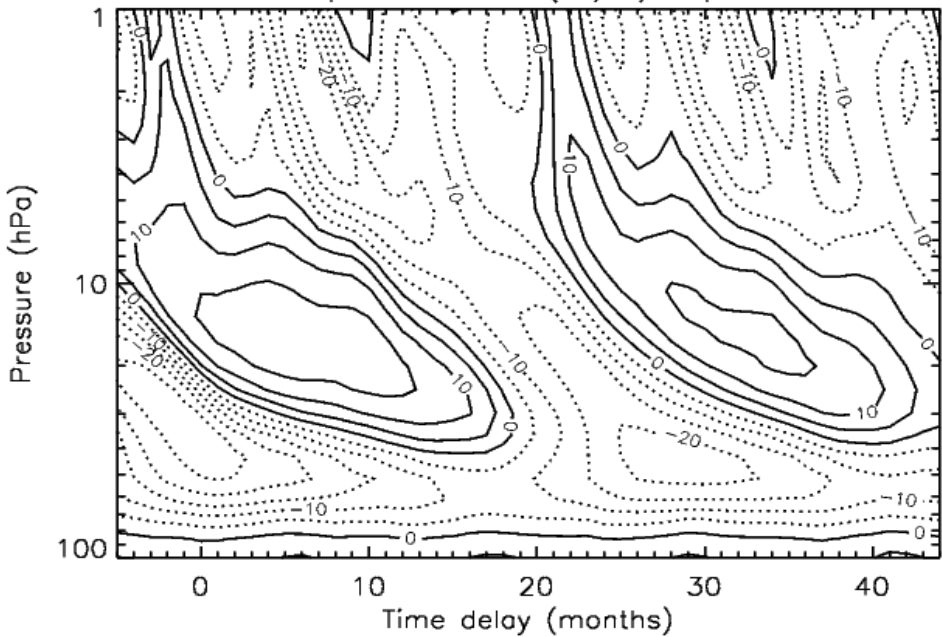
- A discrete spectrum of IGWs is launched at each grid point from tropopause between 30S-30N. Uniform longitudinal distribution.
- Implemented in CESM/WACCM4 and made a simulation 1850-2004.

# Zonal Wind Spectrum: Equator

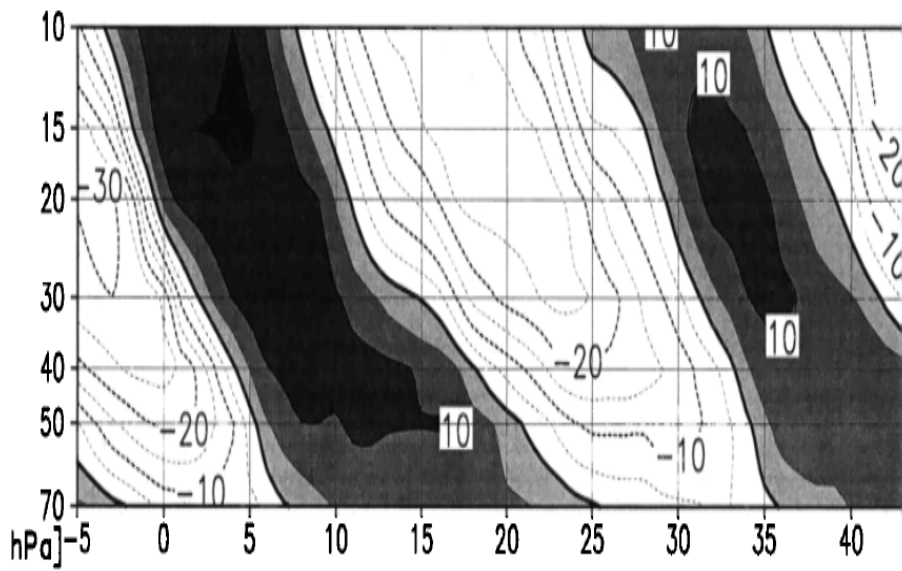


# Composite Zonal Mean U: Equator

composite ubar (m/s) equator

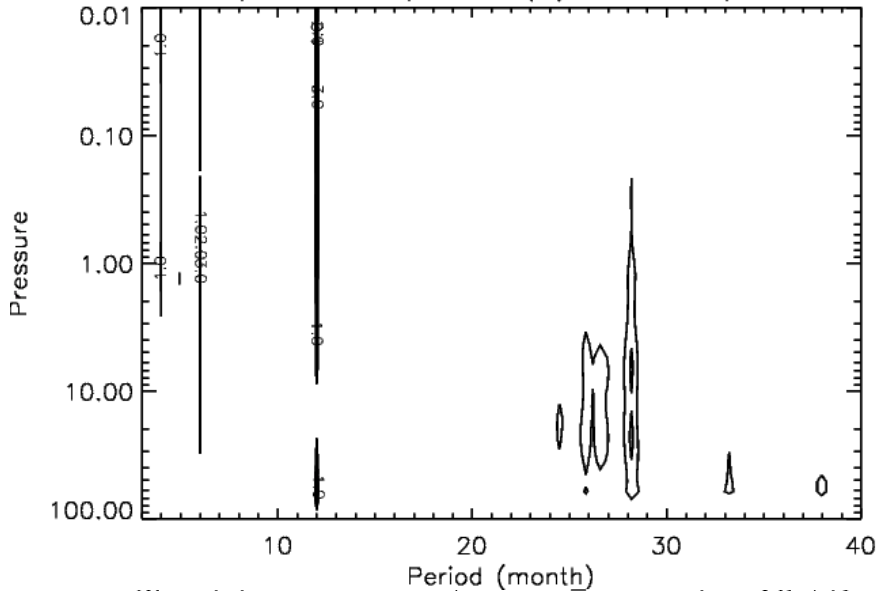


- 10 hPa: Westerly phase 18 mon. in WACCM, compared with 10 mon. in reanalysis.
- Westerly phase becomes shorter at lower altitude, opposite to the reanalysis.
- Westerly phase stops at 40 hPa.

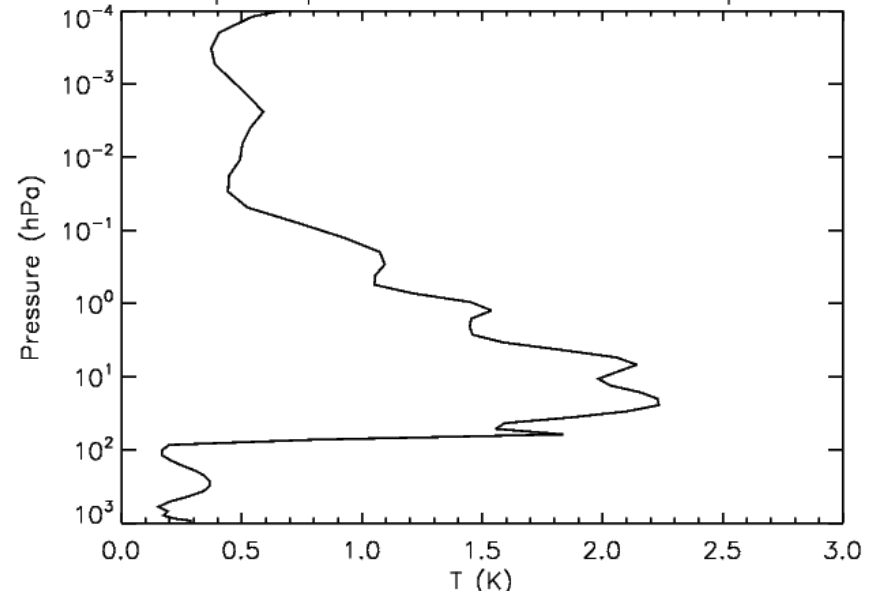


# Temperature Spectrum: Equator

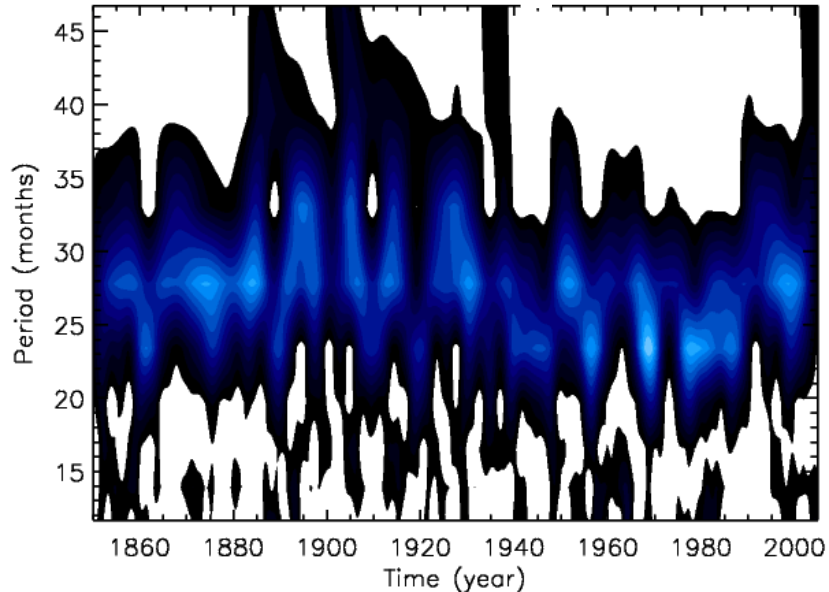
T spectral amplitude (K) at the Equator



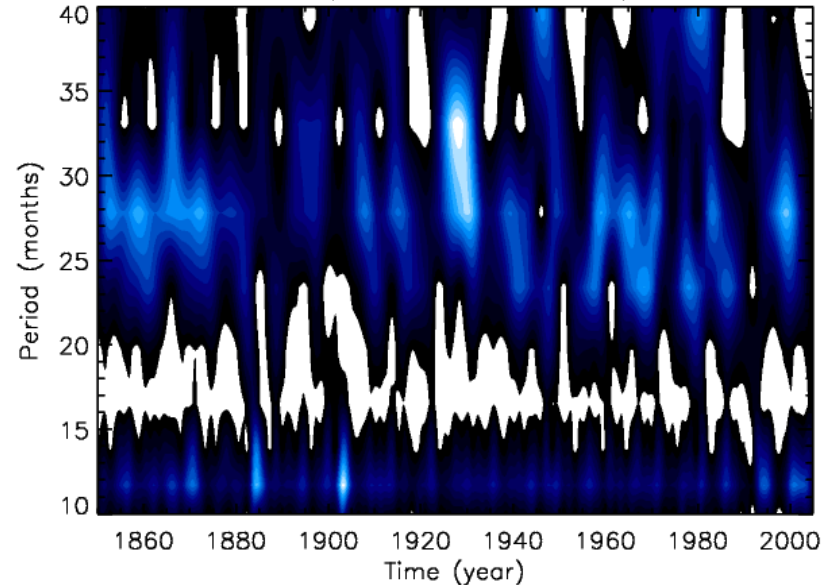
Temp. amp 22–34 months at the equator



Wavelet power: zonal mean T, equator 10 hPa

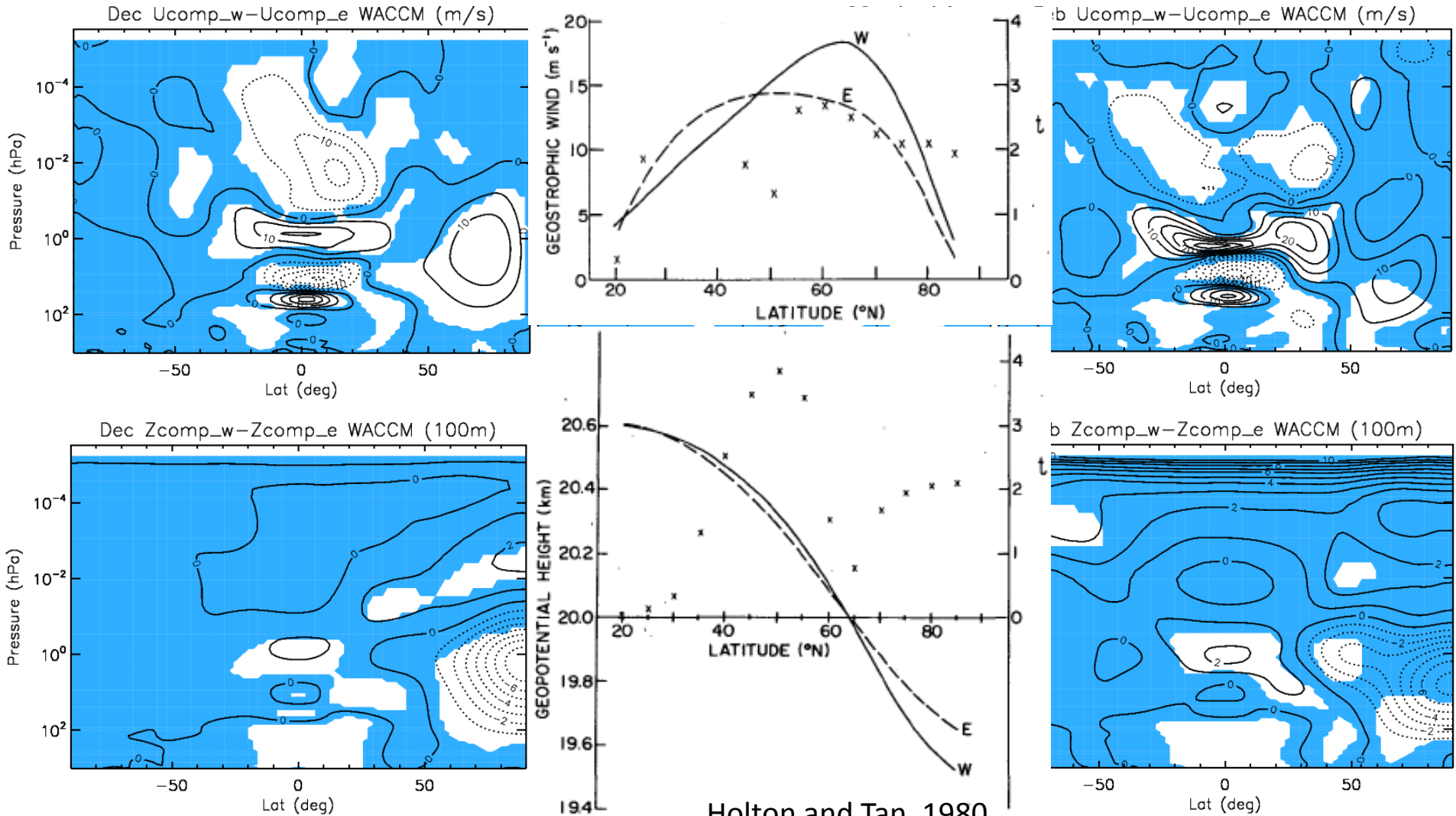


Wavelet power: total O3, equator



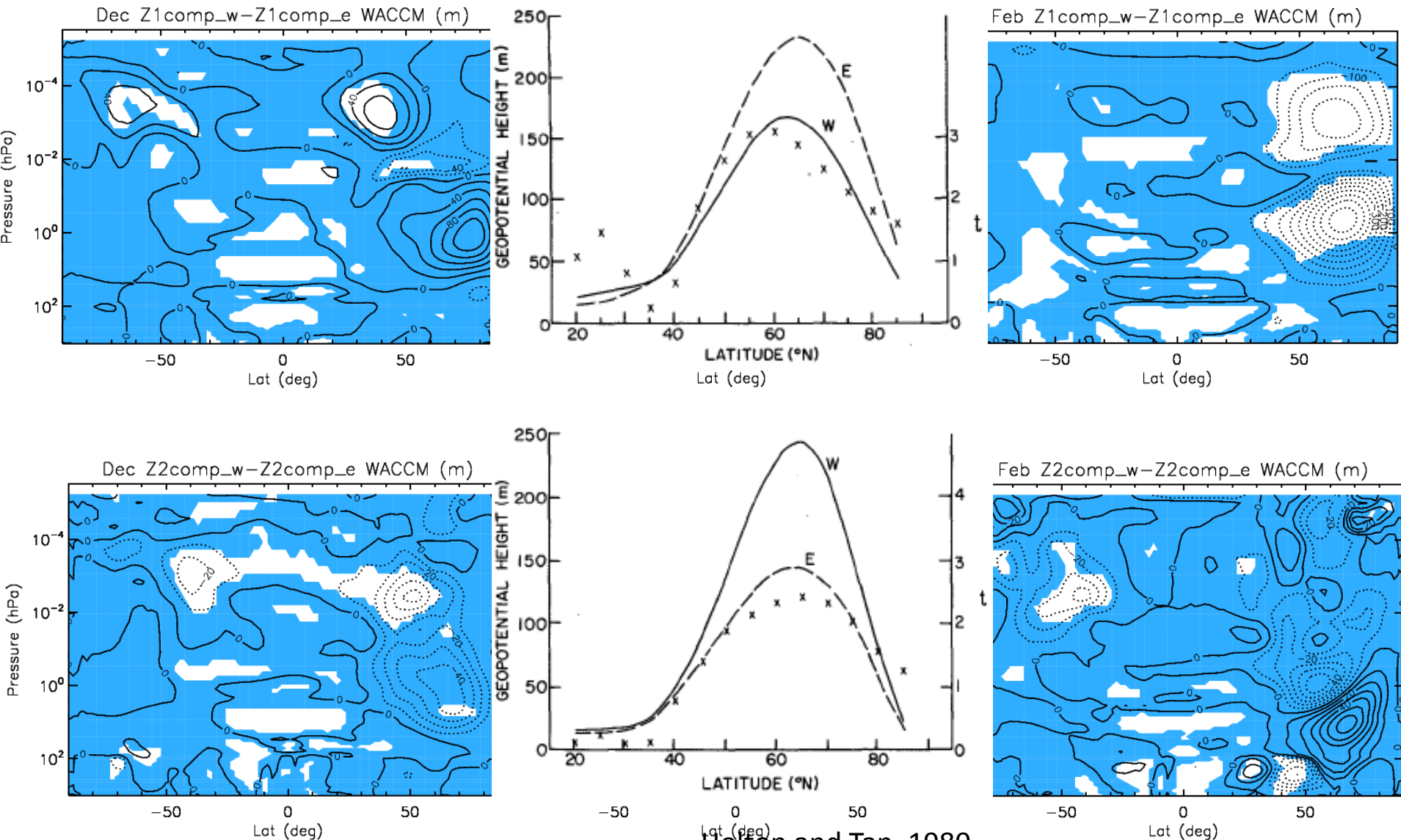


# DJF Composite Diffs: W-E



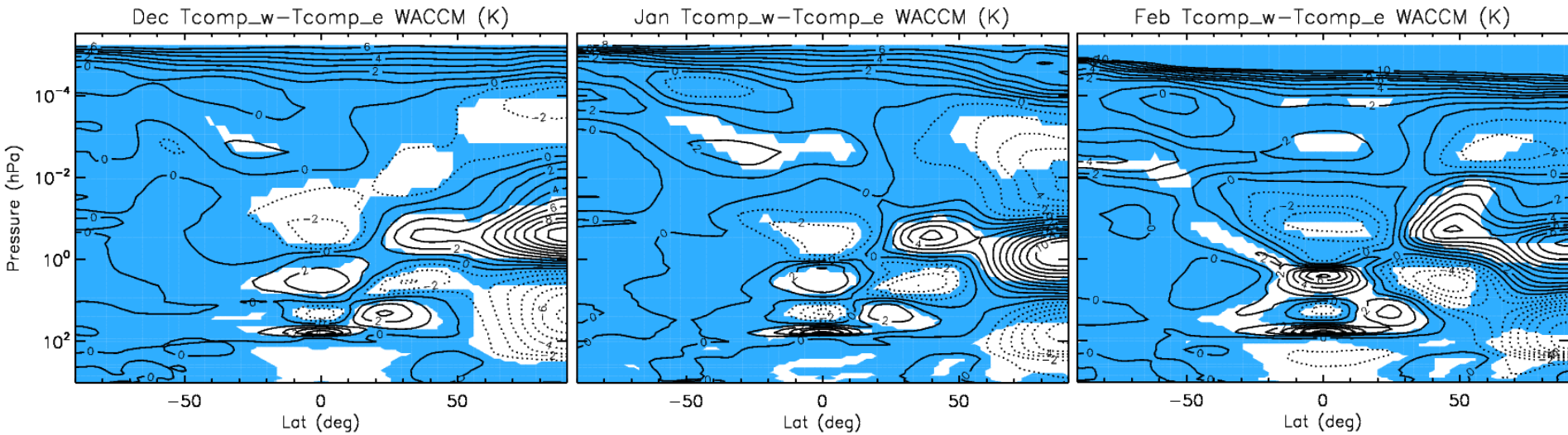
Holton and Tan, 1980

# Planetary waves (1-2): W-E



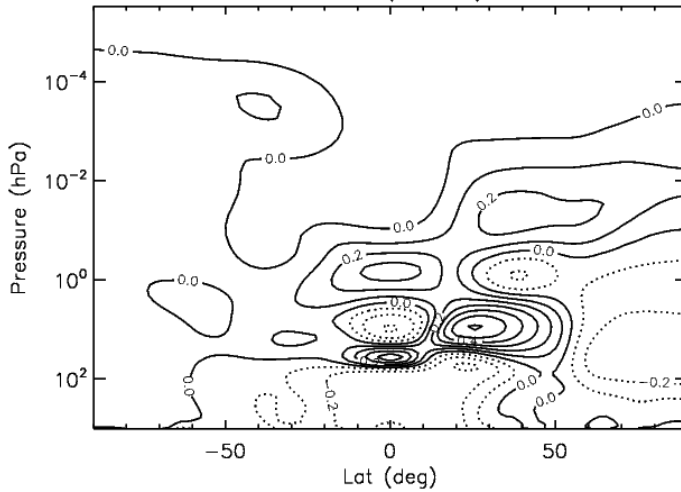
Holton and Tan, 1980

# DJF Composite T Diff: W-E

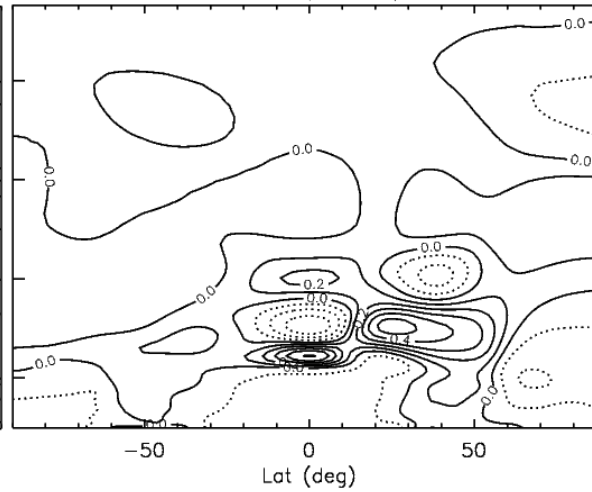


# Correlation Between Z/Z1 with U(43hPa)

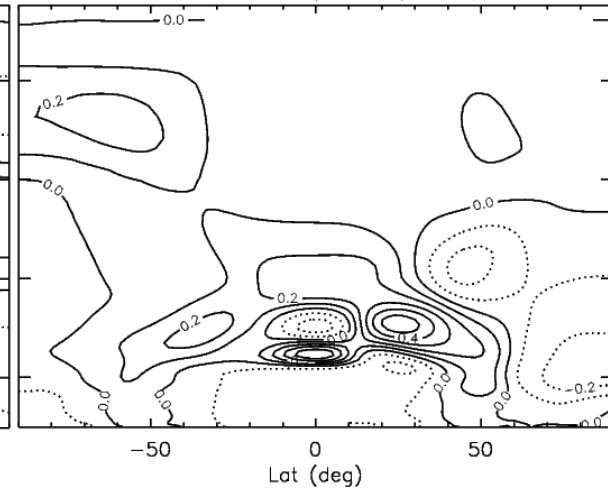
Dec Z and U(43hPa) Corr.



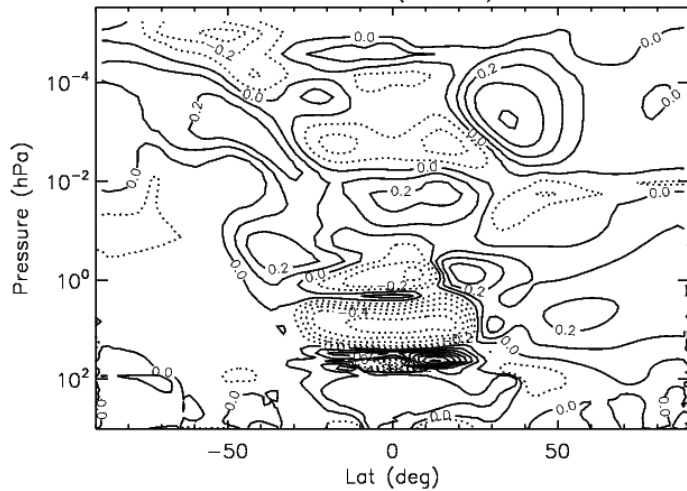
Jan Z and U(43hPa) Corr.



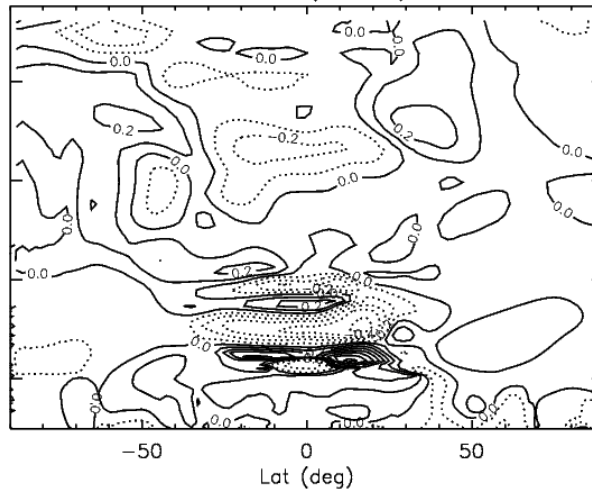
Feb Z and U(43hPa) Corr.



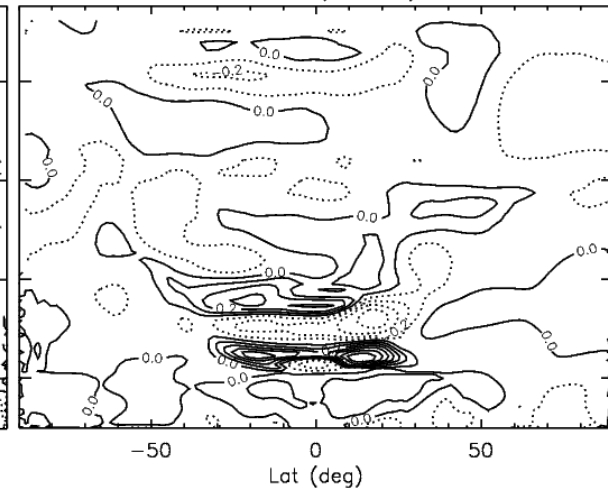
Dec Z1 and U(43hPa) Corr.



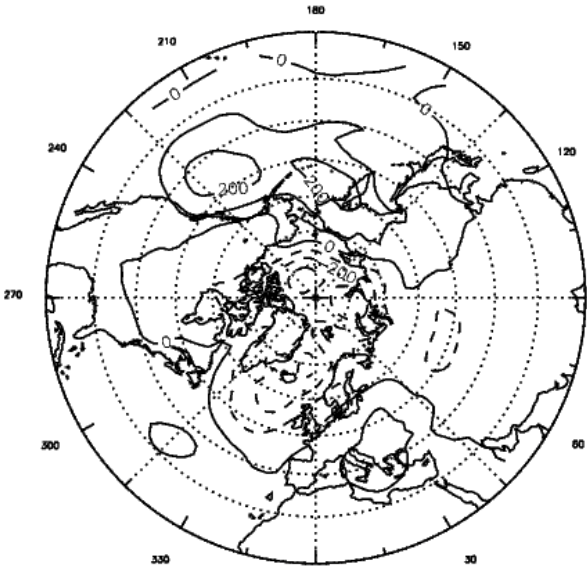
Jan Z1 and U(43hPa) Corr.



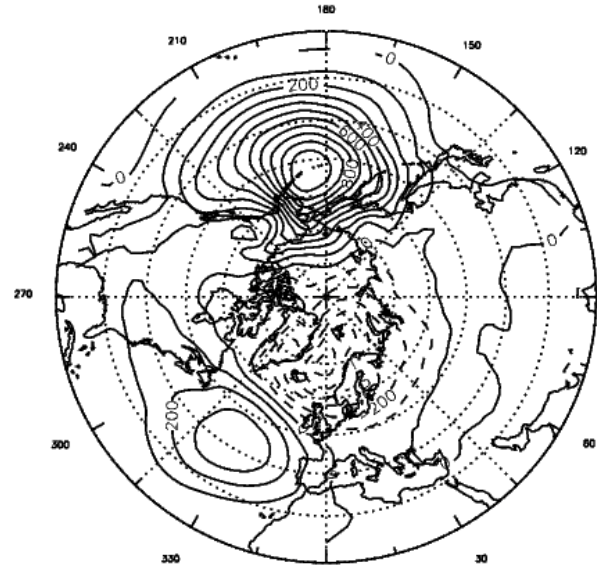
Feb Z1 and U(43hPa) Corr.



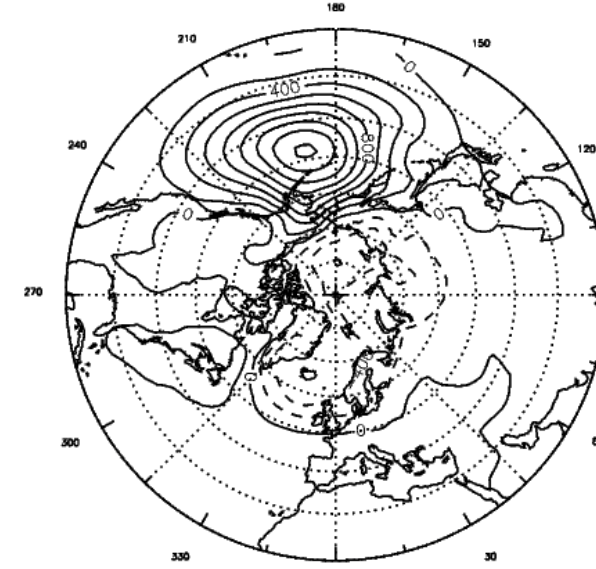
# Surface Pressure: W-E



Longitude  
Dec PScomp\_w-PScomp\_e (Pa)



Longitude  
Jan PScomp\_w-PScomp\_e (Pa)



Longitude  
Feb PScomp\_w-PScomp\_e (Pa)



Longitude  
Mar PScomp\_w-PScomp\_e (Pa)

# Summary

- The new IGW parameterization scheme produce QBO-like oscillations in CESM/WACCM4 simulation (1850-2004).
- Stratosphere zonal mean zonal wind around equator oscillates with a mean period of 28 months. Wavelet analysis shows that the period and strength of the oscillation vary with time.
- Lengths of QBO W/E phase differ from observations.
- Mesosphere QBO weak compared with observations.
- Holton-Tan relation reproduced: W phase -> lower geopotential/temperature, stronger jet at high latitudes of winter hemisphere.
- Surface pressure change consistent with reanalysis. W phase -> low pressure anomaly over winter pole and high pressure anomaly over northern Pacific (50N) and Atlantic (40N) oceans.