

The background of the slide is a grayscale aerial photograph of a river delta. The image shows a complex network of channels and distributaries that branch out from a single point, creating a dense, web-like pattern. The channels vary in width and depth, with some appearing as dark, narrow lines and others as broader, lighter-colored paths. The overall texture is highly detailed and organic, characteristic of natural river systems.

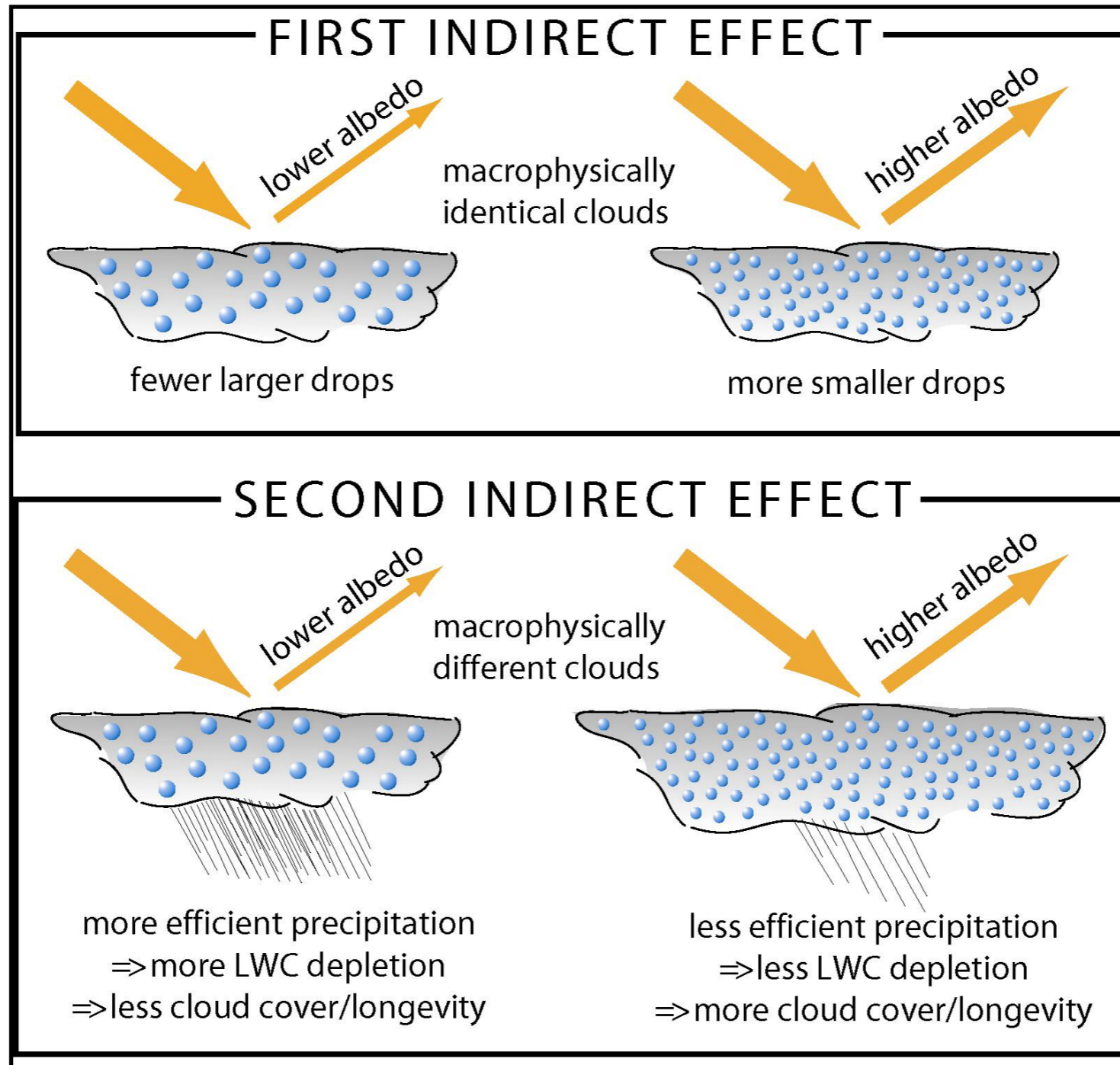
# Aerosol indirect effects in PBL clouds-LES vs. SCAM5

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Funded by NOAA Sc-Cu CPT

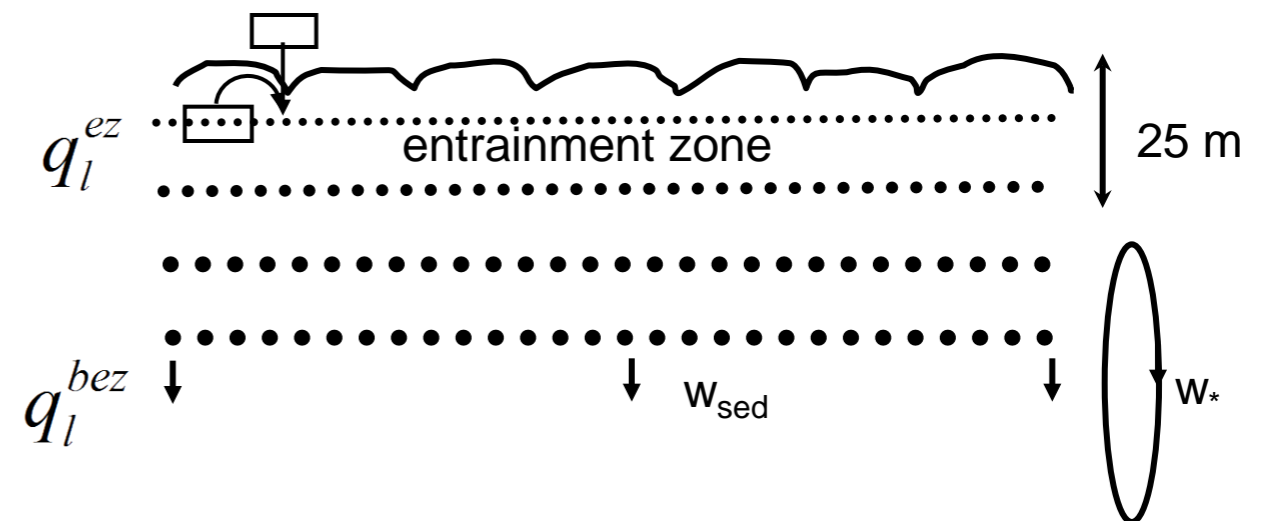
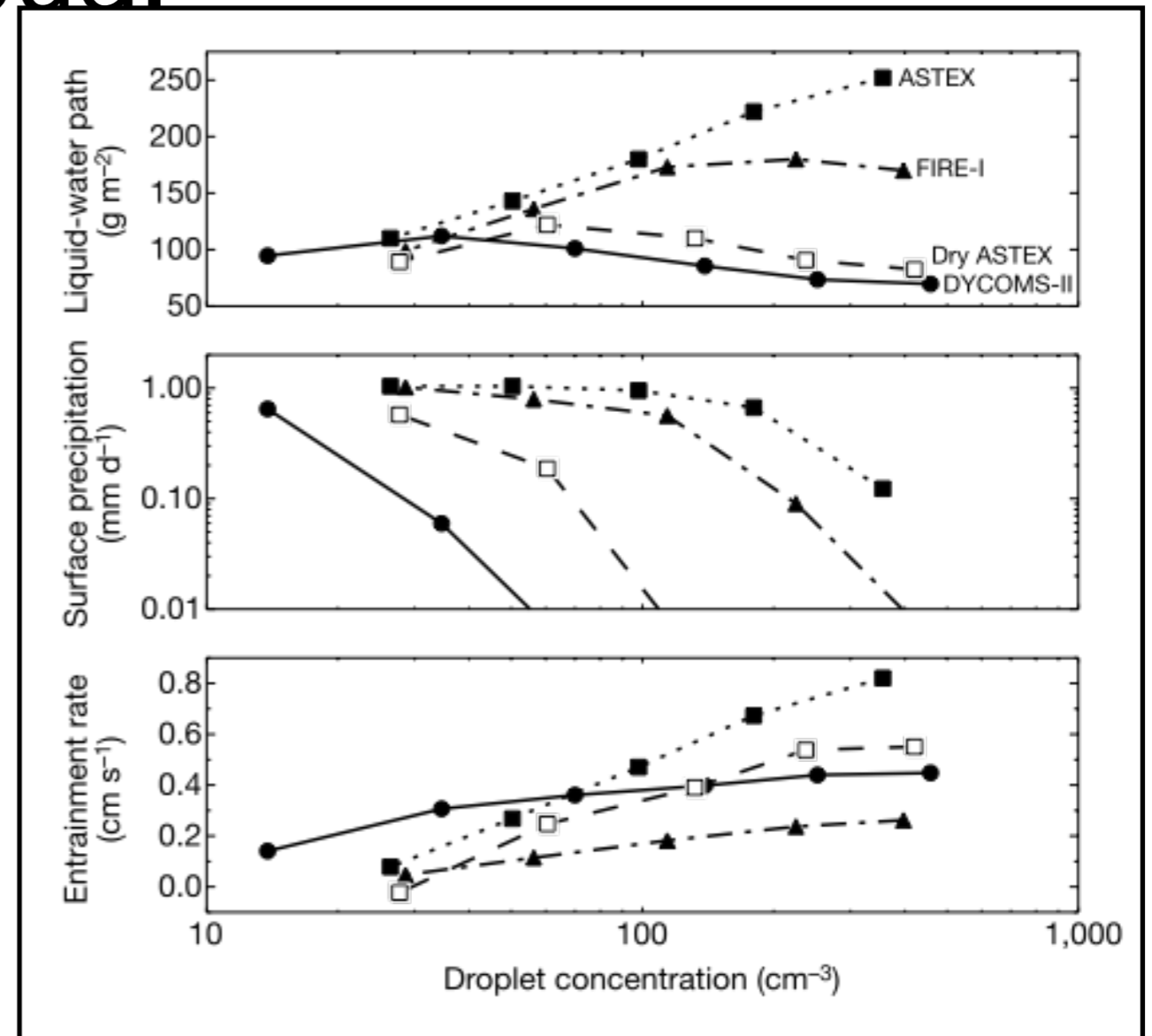
# Aerosol Impacts on Clouds

- Aerosols can impact clouds through
  - cloud brightening (1st indirect effect) or
  - modification of cloud lifetime (2nd indirect effect).



# More aerosol thins nearly nonprecipitating Sc cloud.

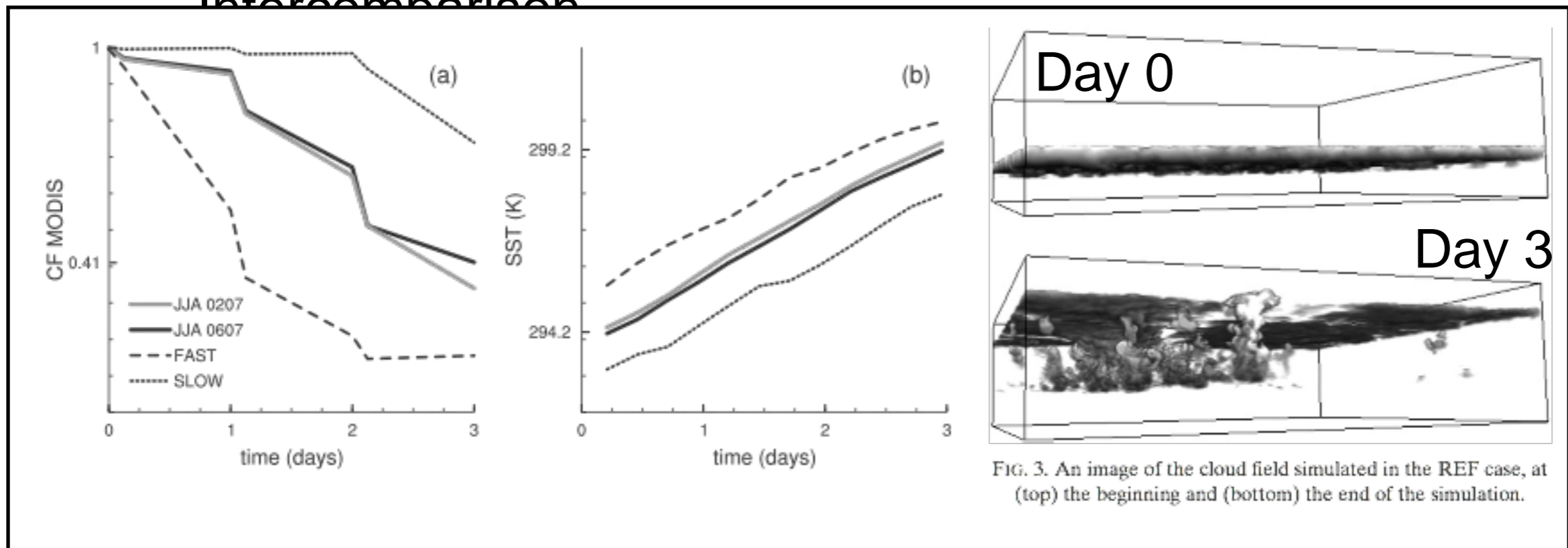
- Ackerman et al (2004) found that stratocumulus clouds only thicken with increasing cloud droplet concentration  $N_d$  until surface precipitation rate becomes small ( $<0.1 \text{ mm d}^{-1}$ ).
- Due to enhanced entrainment of dry air with higher  $N_d$ .
- Bretherton et al (2007): Higher  $N_d \rightarrow$  Less sedimentation  $\rightarrow$  more efficient entrainment, due to increased evaporation of liquid water in the entrainment zone.



# Our Study: Case Setup

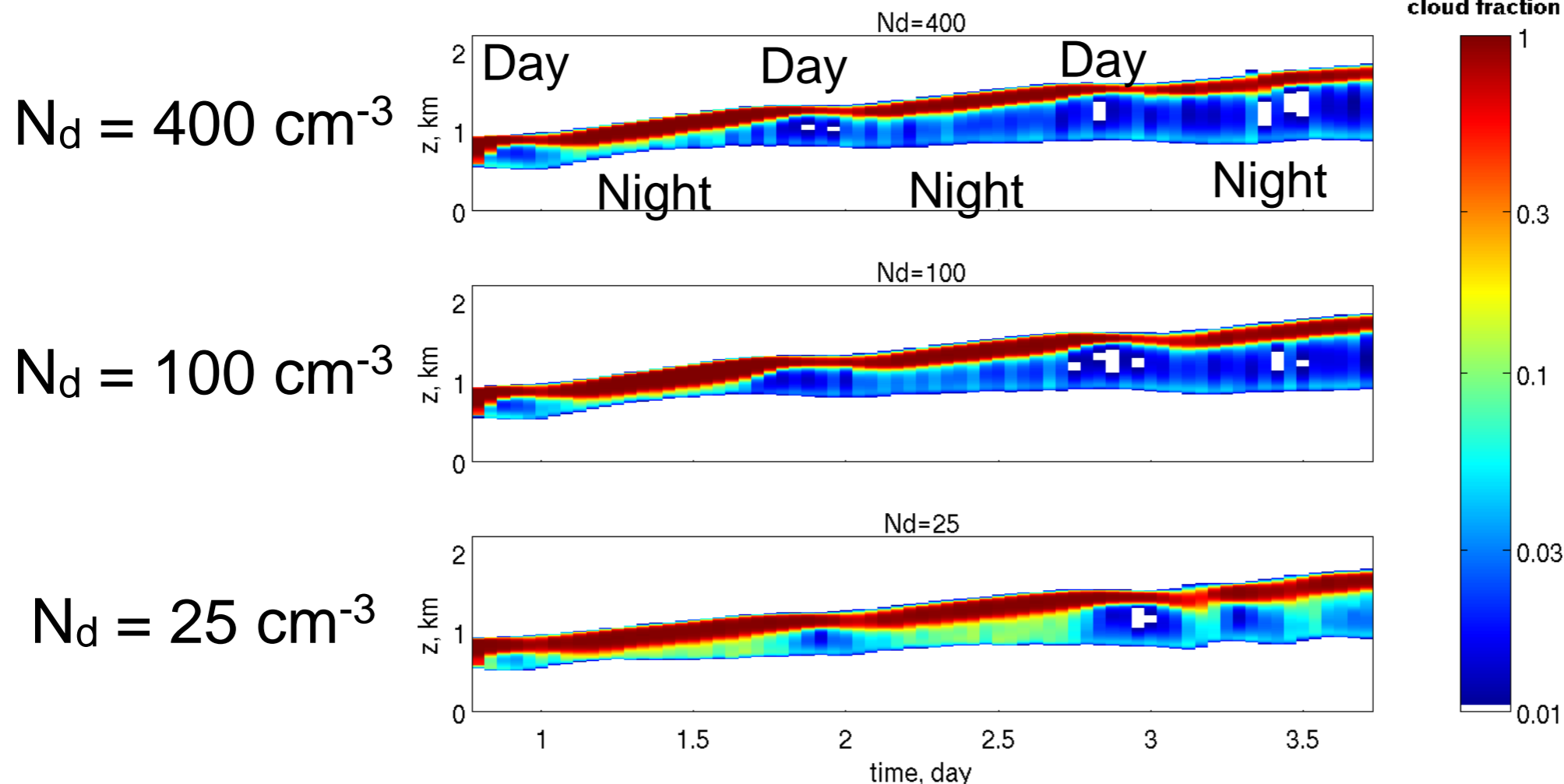
- Stratocumulus to trade cumulus transition: a composite case from the Northeast Pacific (Sandu, Stevens & Pincus, 2010; Sandu & Stevens, 2011). Summertime conditions (JJA2006-7).
- Simulation follows composite Lagrangian trajectory over progressively warmer SSTs with fixed subsidence.
- Finish after 3 days before breakup of capping Sc cloud.
- Basis for a GCSS Boundary Layer Cloud WG

Intercomparison

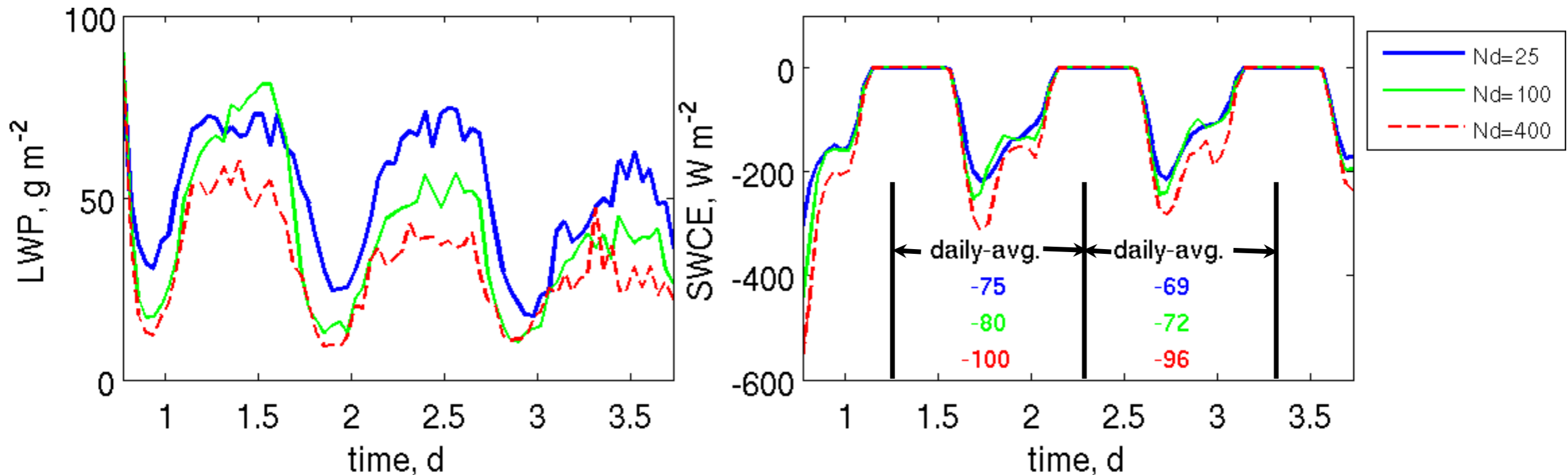


# LES results

- Large eddy simulation model: System for Atmospheric Modeling, v. 6.8 (SAM, Khairoutdinov & Randall, 2003).  $L_x=L_y\sim 4.5\text{km}$ .  $\Delta x=\Delta y=35\text{m}$ ,  $\Delta z=5\text{m}$  from  $\sim 0.5\text{-}2.5\text{km}$ .
- Microphysics: Khairoutdinov & Kogan (2000) with fixed  $N_d=25, 100, 400/\text{cm}^3$ .
- Radiation: RRTMG w/cloud droplet effective radius computed from LWC and  $N_d$ , assuming  $\sigma_g=1.2$ . Includes diurnal cycle.

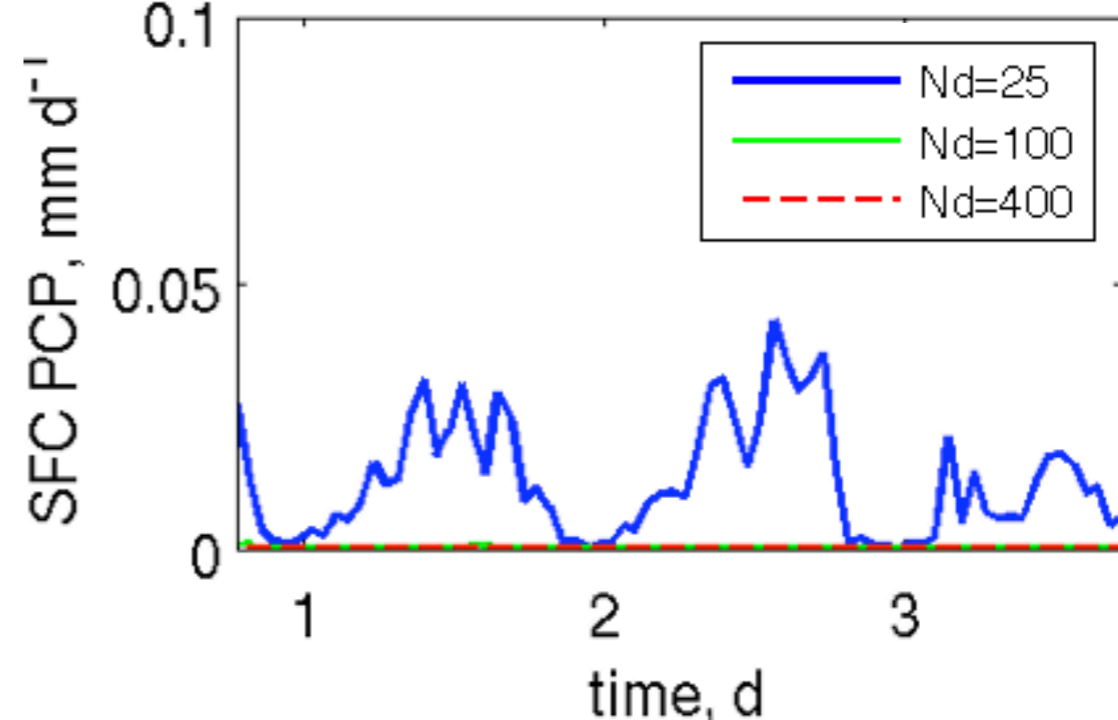
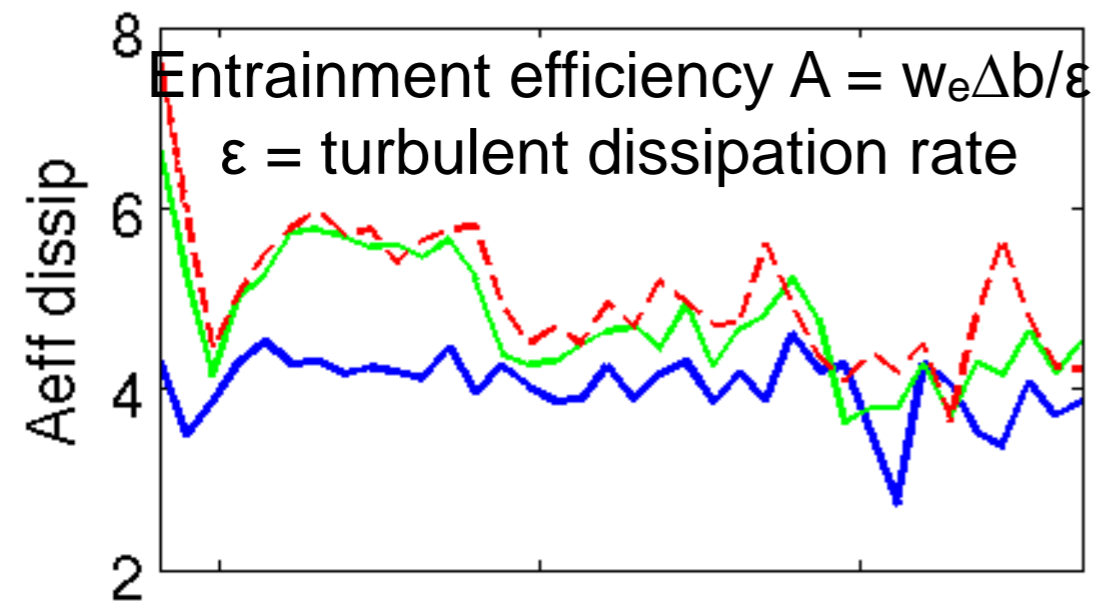
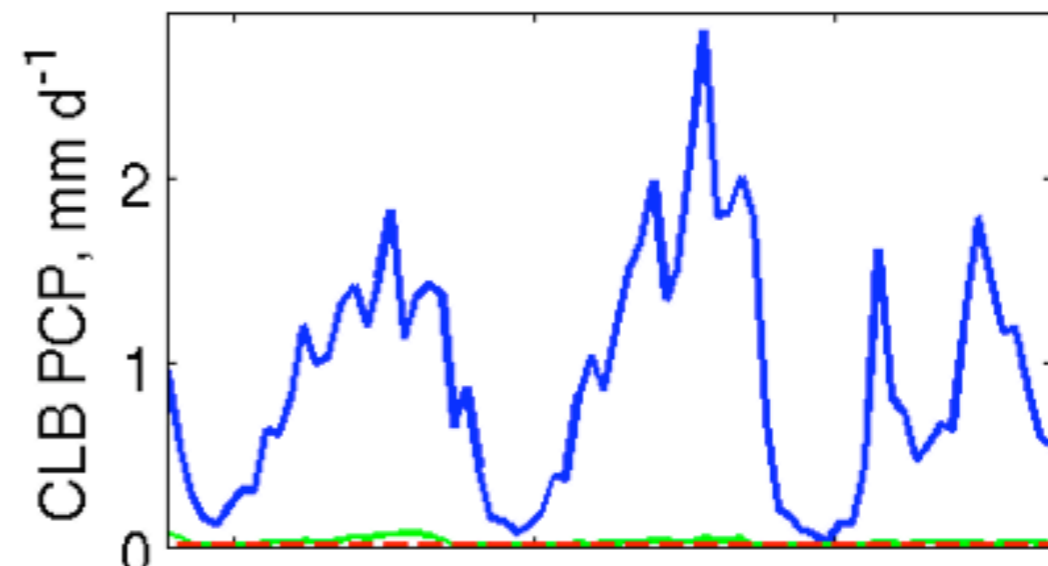
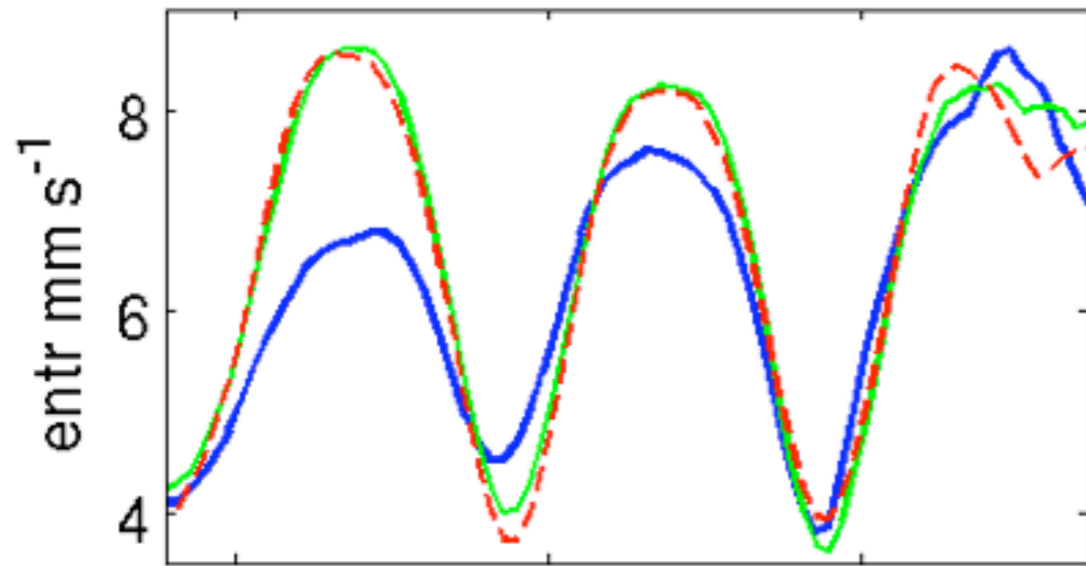


# Cloud thickness and albedo response to $N_d$



- Optical depth of an Sc layer  $\tau \sim \text{LWP}^{5/6} N_d^{1/3}$ .  
40% decrease in LWP  $\leftrightarrow$   $4 \times N_d$ .
- $N_d$  25  $\rightarrow$  100 cm<sup>-3</sup>: 35% daytime LWP decrease, little albedo increase.
- $N_d$  100  $\rightarrow$  400 cm<sup>-3</sup>: little daytime LWP decrease, Twomey effect reigns.

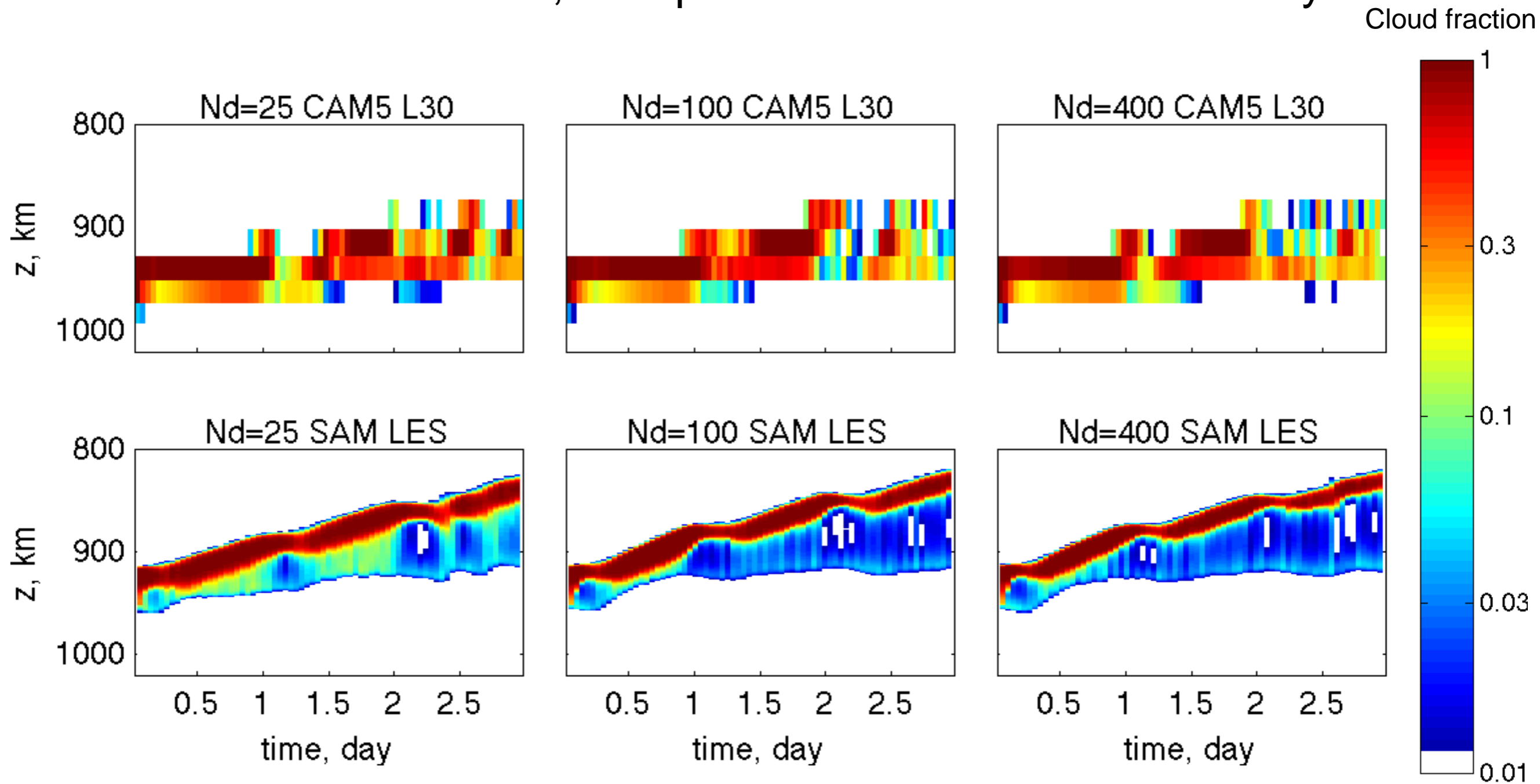
# Entrainment and Drizzle



- Entrainment efficiency increases with  $N_d$  as expected.
- Drizzle evaporating below cloud base is significant for  $N_d=25$

# Can SCAM5 reproduce this behavior?

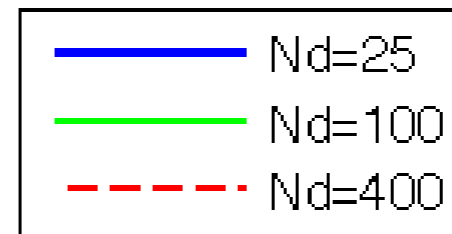
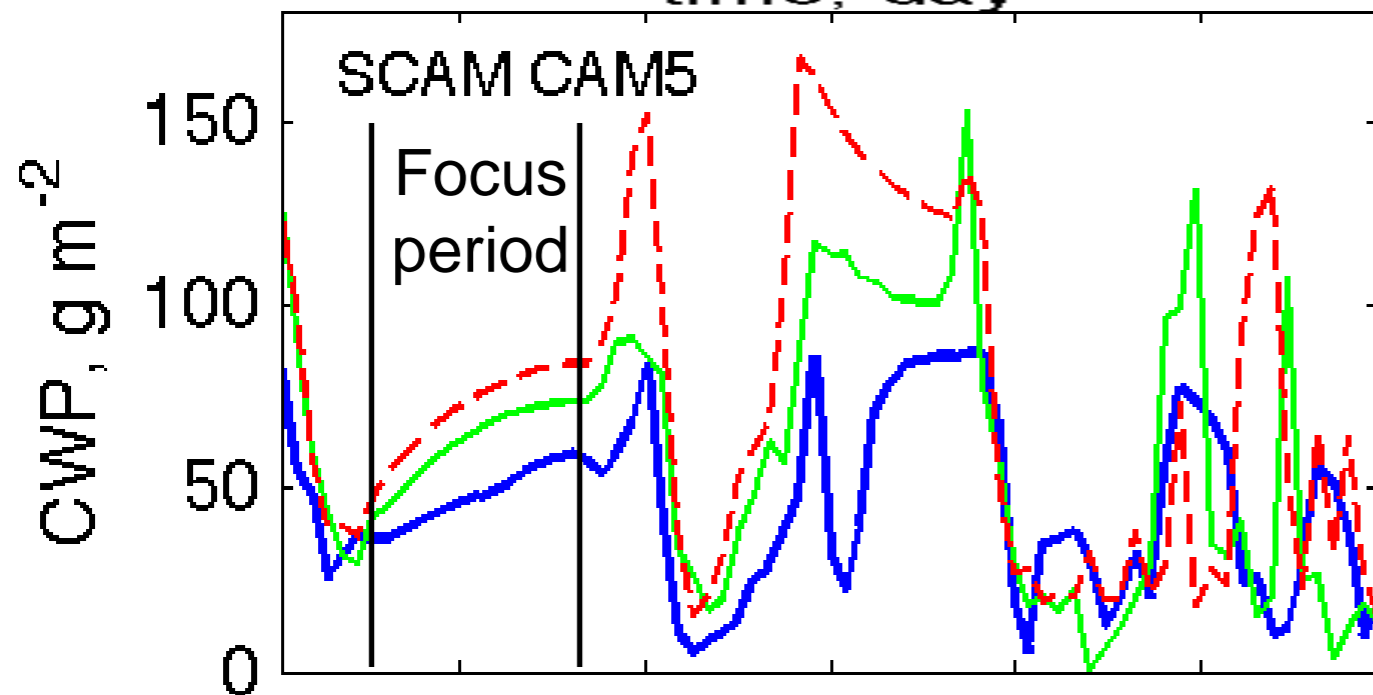
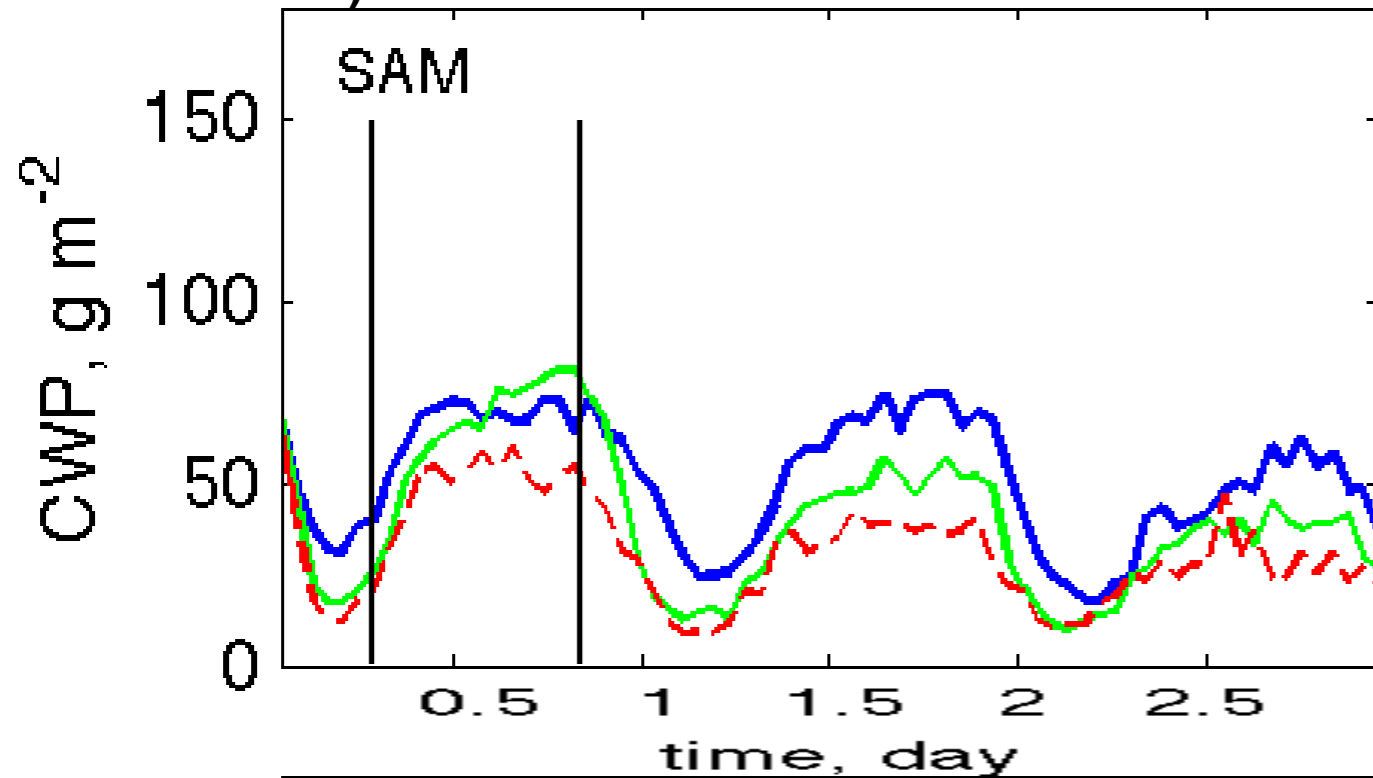
- 30-level SCAM5 not bad, except too little cloud on the last day.



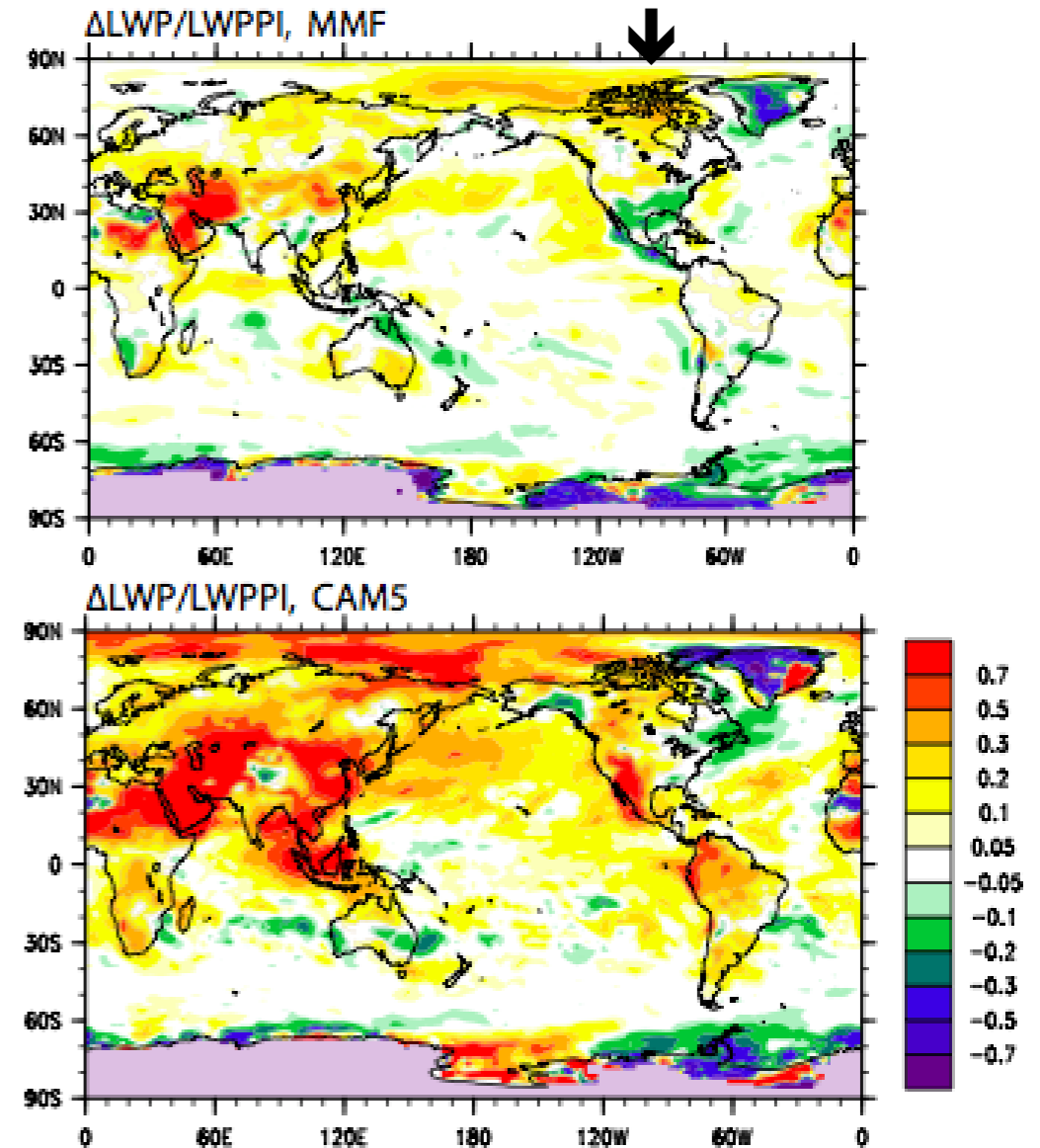


# But 2<sup>nd</sup> indirect effect opposite to LES!

- SCAM5 has thicker cloud with increasing  $N_d$ . CAM5 simulations also show more positive  $dLWP/dN_d$  than SP-CAM, contributing to their stronger aerosol indirect effect (Wang et al. 2011).



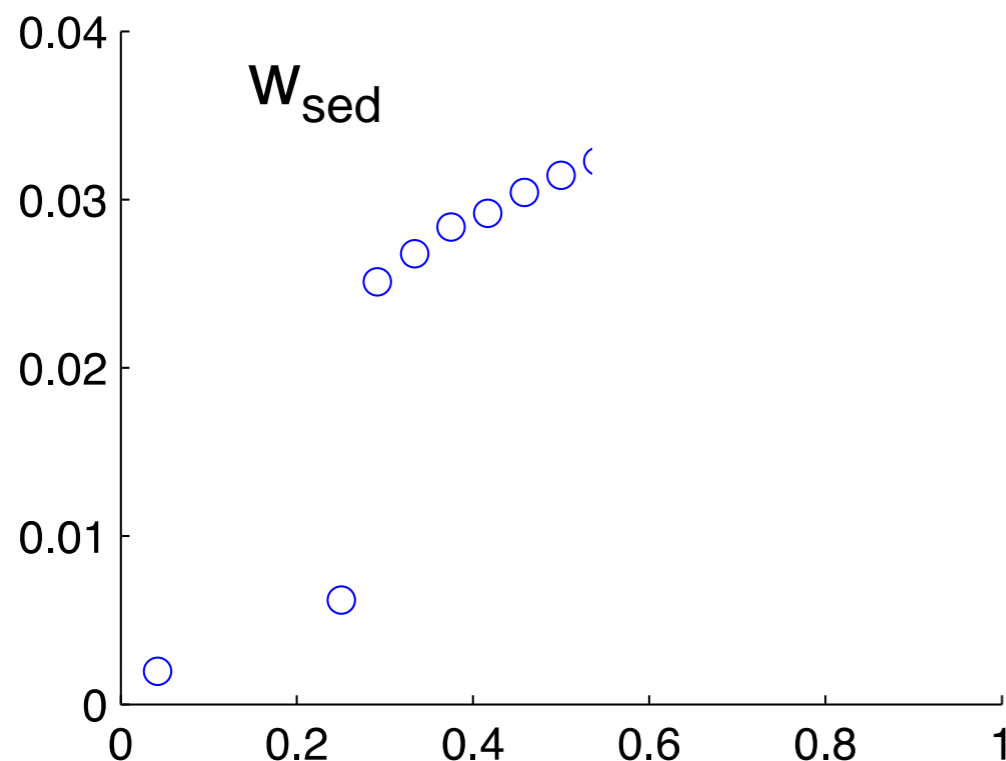
Reminiscent of  
Wang et al.  
2011



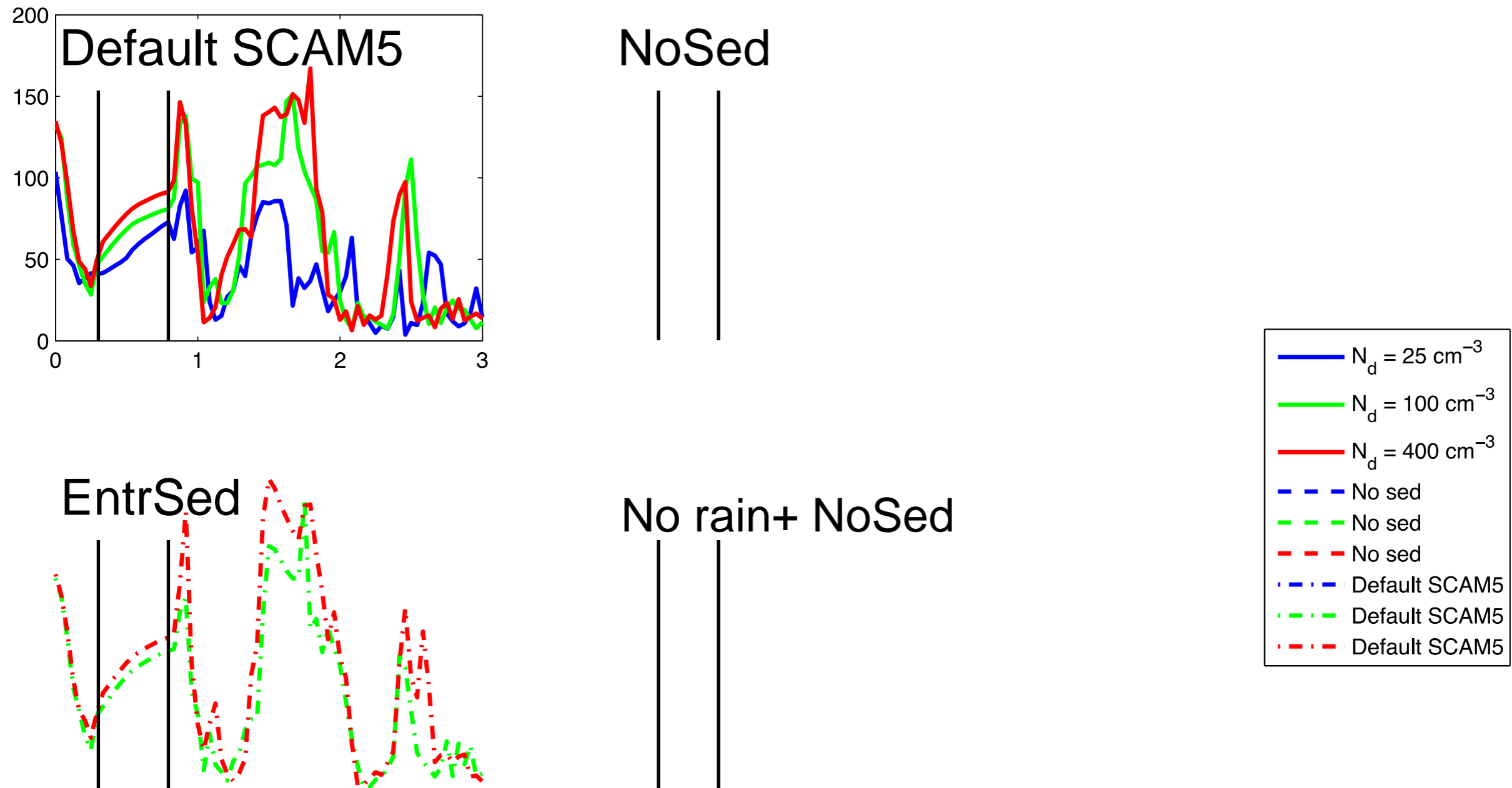
# Sensitivity studies

- Default CAM5 has cloud droplet sedimentation at a predicted rate  $w_{\text{sed}}$  in stratiform microphysics, but no other entrainment-sedimentation feedback
- NoSed: Cloud droplet sedimentation off in stratiform microphysics.
- EntrSed: Add entrainment-sedimentation feedback by multiplying evaporative enhancement factor  $evhc - 1$  in UWMT entrainment rate by (Bretherton et al 2007)

$\exp(-a_{\text{sed}} w_{\text{sed}}/w_*)$ ,  $a_{\text{sed}} = 9$  (LES-tuned),  $w_* = \text{convective velocity} \sim 1 \text{ m s}^{-1}$

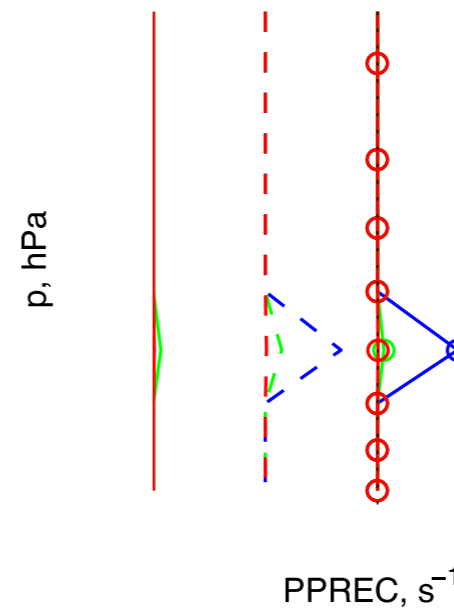
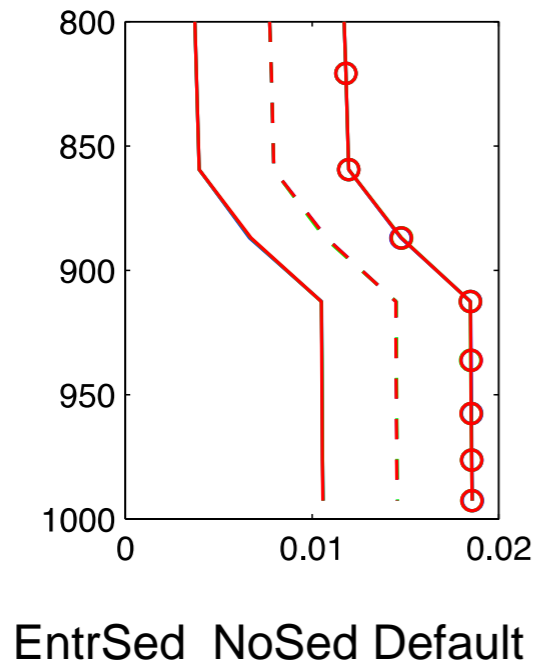


# Sedimentation not the issue



- Differences apparent in first night, when simulated PBL is well-mixed.
- Addition of stratiform sedimentation reduces LWP in all cases
- Addition of entrainment-sedimentation feedback brings some LWP back
- But  $N_d = 25$  vs. 400 LWP difference as large with no sedimentation.

# So evaporating drizzle is a likely culprit



- Look at  $t = 0.5$  day (first night). Significant evaporating drizzle for  $N_d = 25$ .
- Loss of  $q_l$  during each timestep comparable to  $400-25 \Delta q_l$

# Conclusions

- Large eddy simulations of composite Sc→Cu transition over NE Pacific.
- Aerosol sensitivity studied via prescribed cloud droplet number concentrations  $N_d = 25, 100, 400 \text{ cm}^{-3}$  in Sc→Cu transition case.
- LES simulations show cloud thins as  $N_d$  increases, due to drizzle and cloud droplet sedimentation effects on entrainment. Leads to near cancellation of first indirect (Twomey) effect on daytime cloud albedo during first full day of simulation.
- SCAM5 simulations produce reasonable transition simulations have opposite dependence on  $N_d$  to LES, even if cloud droplet sedimentation is turned off. We don't know why.
- This case may shed light on  $dLWP/dN_d$  and 2<sup>nd</sup> AIE in global CAM5.