

Improving Ice Nucleation Parameterization for Mixed-Phase Clouds in CAM5

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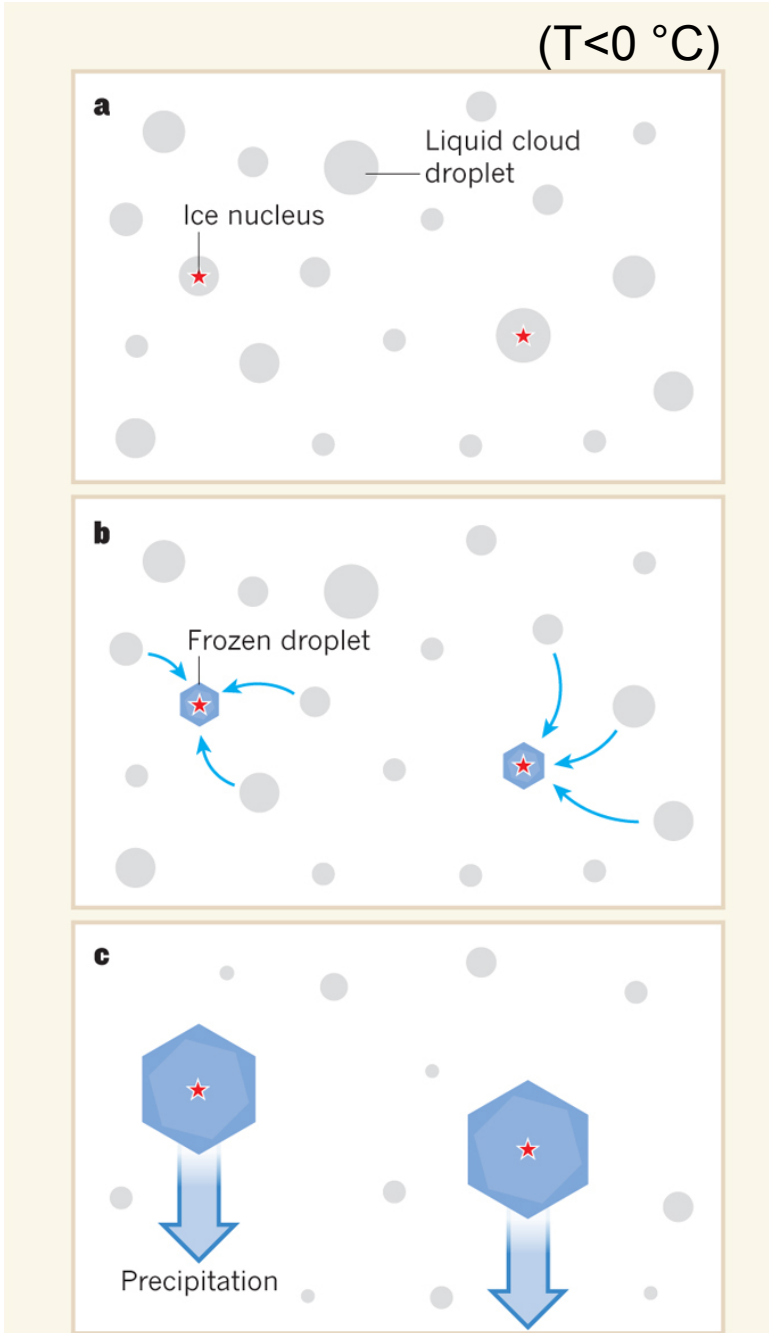
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Ice nucleation important for radiation and precipitation formation in mixed-phase clouds

Ice Nucleation

Bergeron-F. Process

Precipitation Initiation



Koop, Nature (2013)

Classical nucleation theory

Nucleation rate $J = A' r_N^2 \sqrt{f} \exp\left(\frac{-\Delta g^\# - f \Delta g_g^o}{kT}\right)$
 $f(\alpha)$, α is contact angle

Immersion/condensation

Pruppacher and Klett (1997)

Hoose et al. (2010)

$$r_{g,imm} = \frac{2v_w \sigma_{i/w}}{kT \ln(a_w e_{sw} / e_{si})}$$

$$\Delta N_{i,imm} = \sum_x \text{Min}\{f_{l,x} N_{aer,x} f_{i,max,x}, f_{l,x} N_{aer,x} [1 - \exp(-J_{imm,x} \Delta t)]\}$$

Activated/cloud-borne aerosol

Deposition

$$r_{g,dep} = \frac{2v_w \sigma_{i/v}}{kT \ln(e / e_{si})}$$

Interstitial & uncoated aerosol

$$\Delta N_{i,dep} = \sum_x \text{Min}\{(1 - f_{l,x}) (1 - f_{x,coated}) N_{aer,x} f_{i,max,x}, (1 - f_{l,x}) (1 - f_{x,coated}) N_{aer,x} \times [1 - \exp(-J_{dep,x, RH_w=0.98} \Delta t)]\}$$

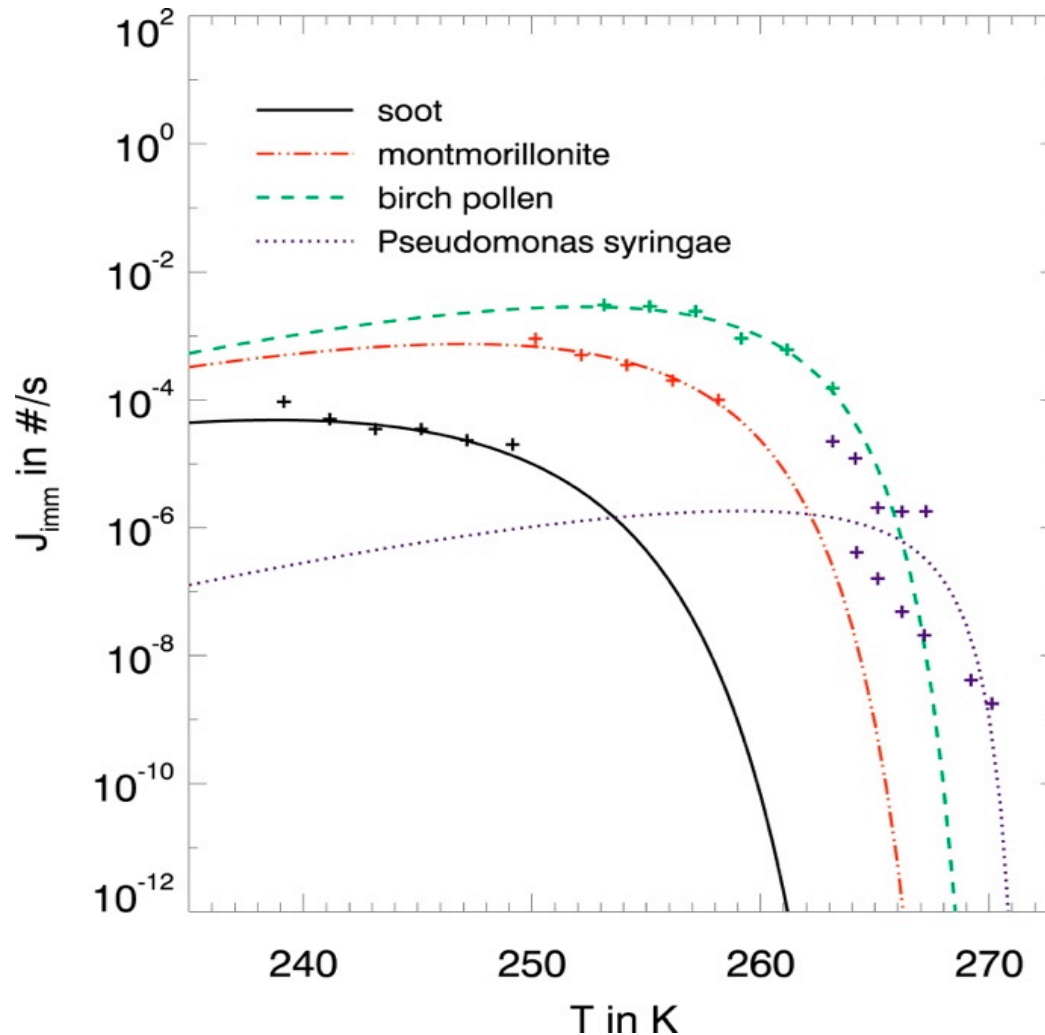
Contact

$$N_{g,contact} \approx 4\pi r_N^2 \frac{e}{v_s \sqrt{2\pi m_w kT}} \times \exp\left[-\frac{\Delta g_{dep}^\# + f \Delta g_{g,dep}^o(r_{g,imm})}{kT}\right]$$

$$\Delta N_{i,contact} = \sum_x \text{Min}\{(1 - f_{l,x}) (1 - f_{x,coated}) N_{aer,x} f_{i,max,x}, (1 - f_{l,x}) (1 - f_{x,coated}) N_{aer,x} \times [1 - \exp[-K_{coll}(r_{N,x}, r_l) N_l \text{Max}(N_{g,contact,x}, 1) \Delta t]]\}$$

PDF- α model: integrate over the PDF of contact angle α

Classical theory links ice nucleation rate to aerosol properties, constrained by experiments



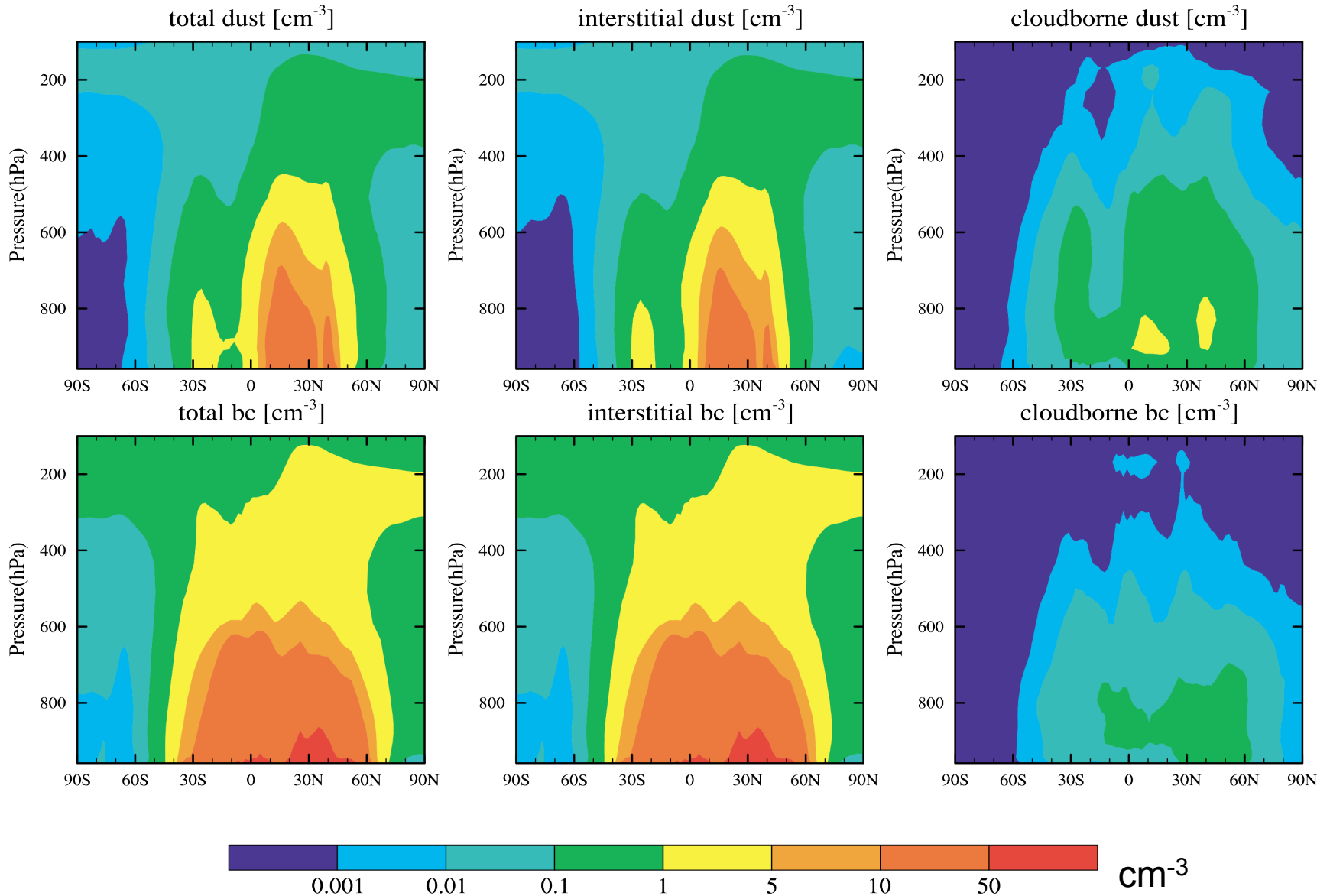
Hoose et al. (2010)

$\alpha=40.2^\circ$ (soot, immersion); $\alpha=31.0^\circ$ (dust, immersion)
 $\alpha=28.0^\circ$ (soot, deposition); $\alpha=12.7^\circ$ (dust, deposition)

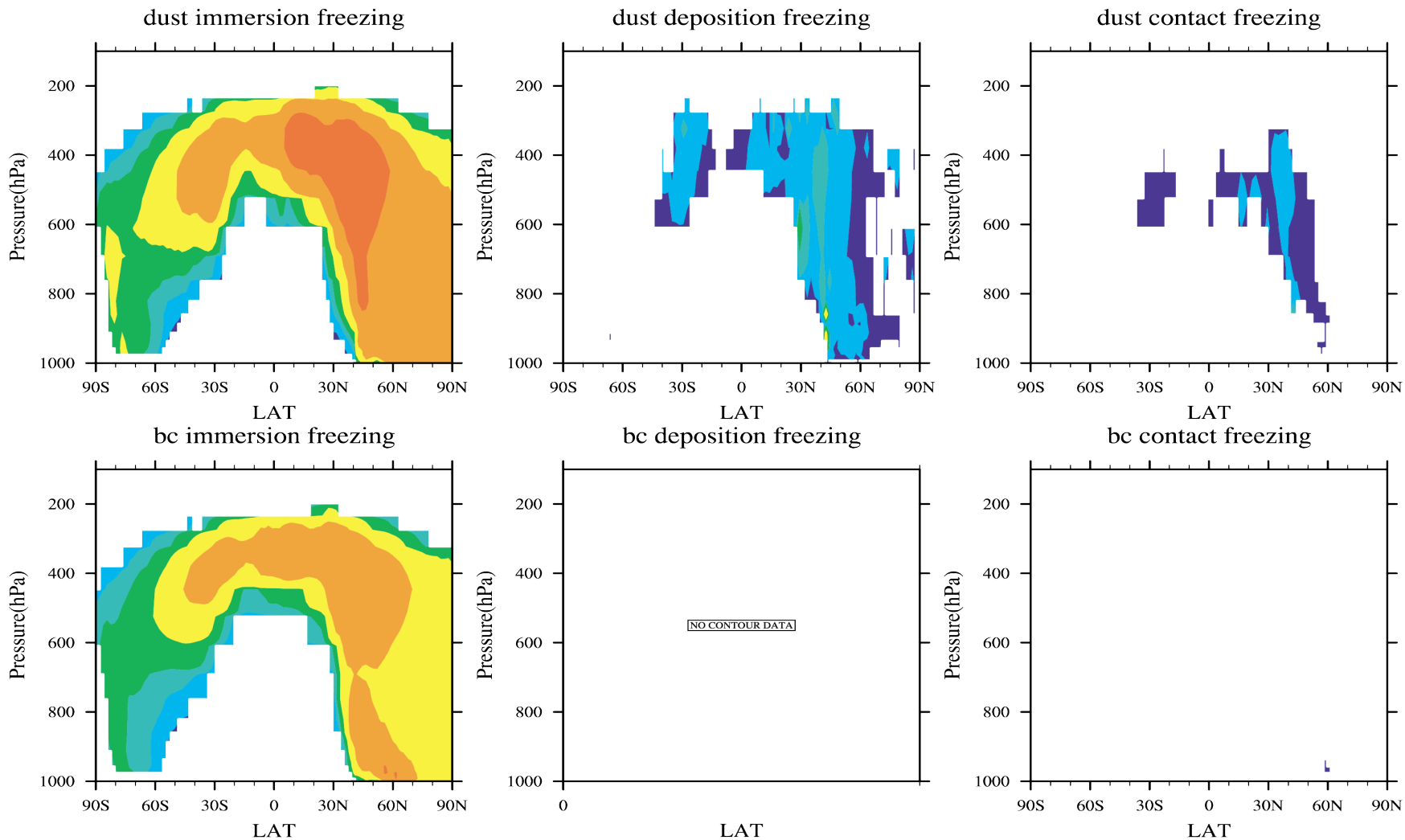
Model Experiments

- ▶ CAM5.1 with FV dynamic core, $1.9^\circ \times 2.5^\circ$, 30 levels
- ▶ 6-yr climate runs with prescribed SST and sea ice (AMIP II type of run)
 - **Default** : Meyers et al. (1992) for deposition/condensation/immersion in mixed-phase clouds, with no link to aerosol
 - **single- α** : Classical nucleation theory with single contact angle (Hoose et al. 2010)
 - **PDF- α** : Classical nucleation theory with PDF-contact angle, mean of PDF same as single α with standard deviation 0.1

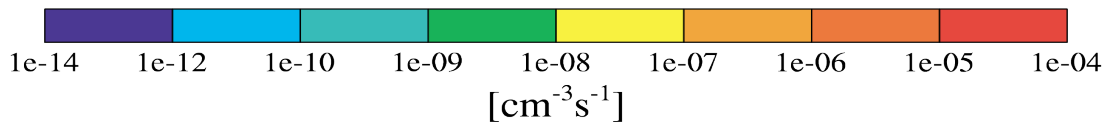
Dust and BC number (interstitial vs. cloud-borne)



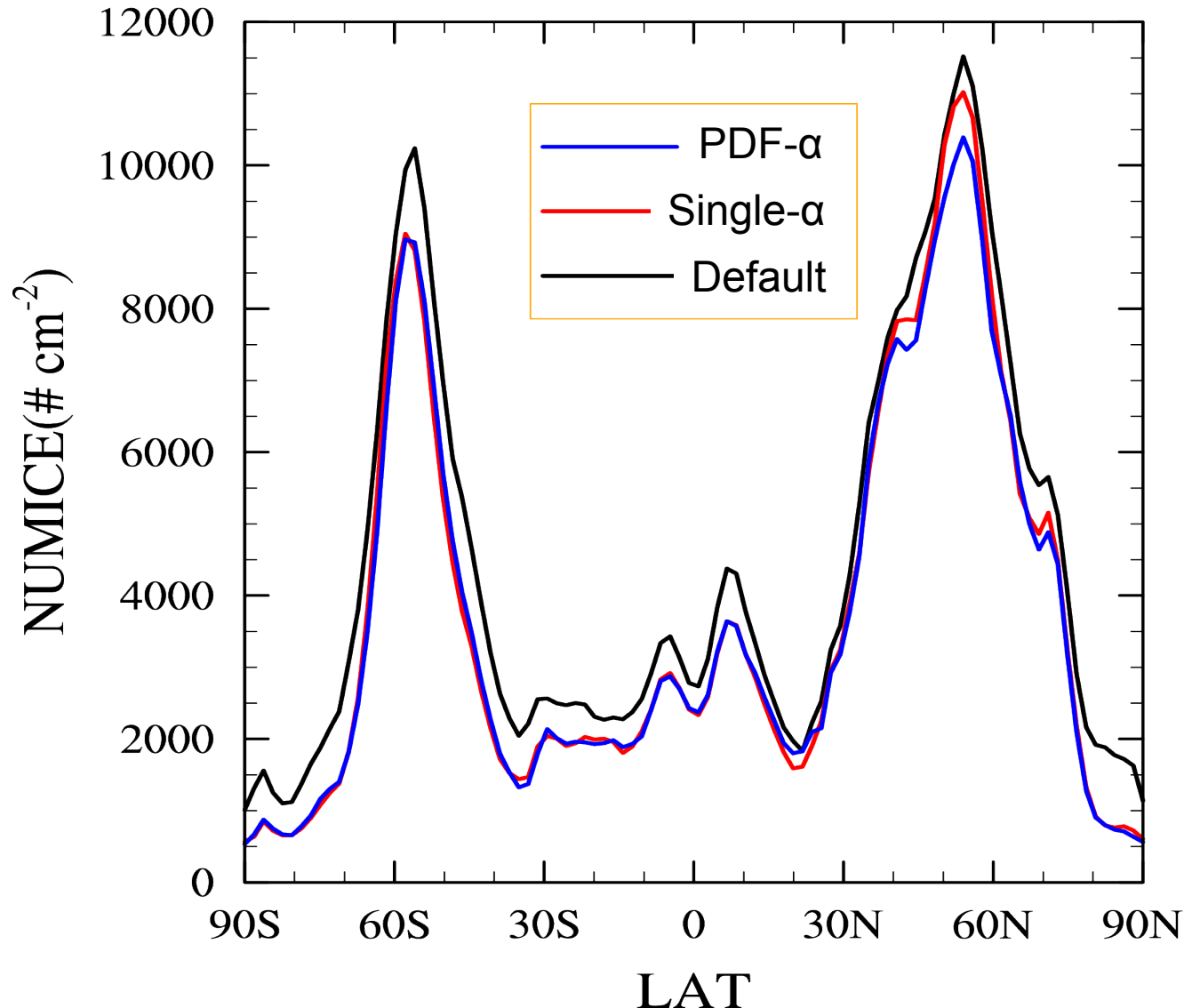
Ice nucleation in mixed-phase clouds: immersion vs. deposition vs. contact mode



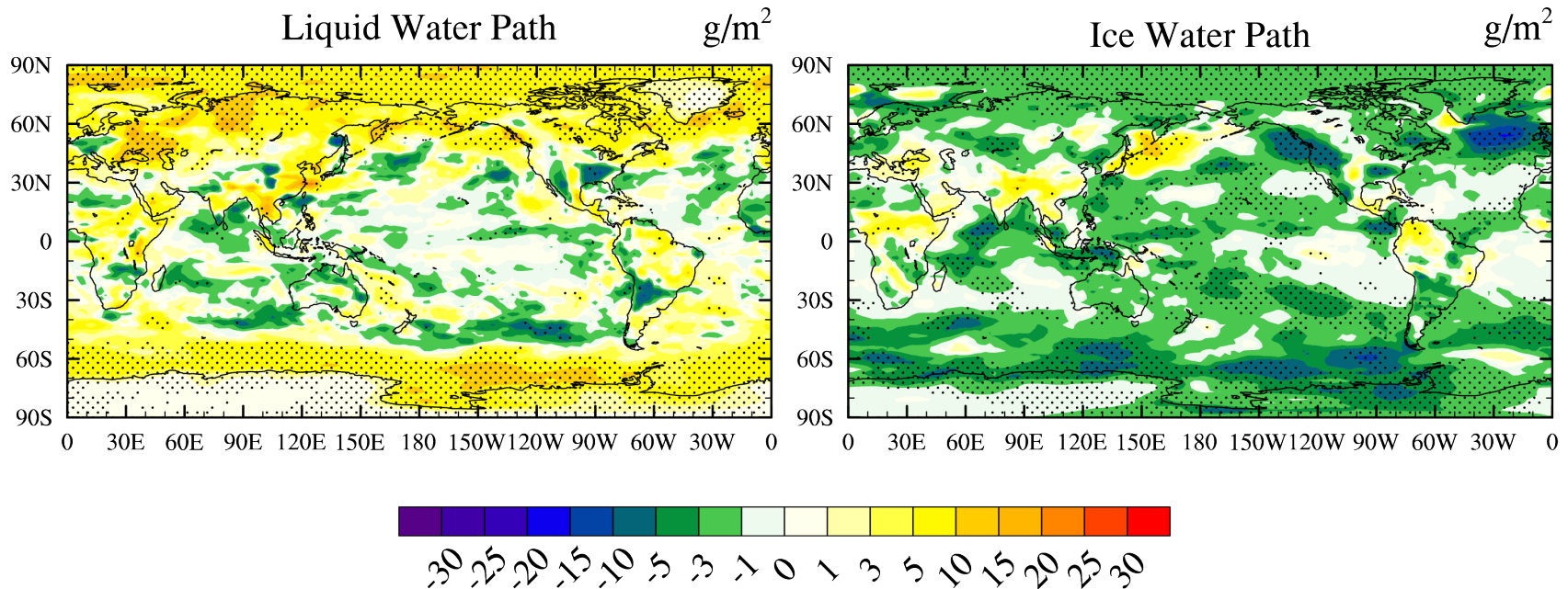
Single- α model



Column Ice Crystal Number Concentration in Mixed-Phase Clouds



Difference in Cloud Water (single α - default)



Fewer Ni \rightarrow More Liquid, Less Ice

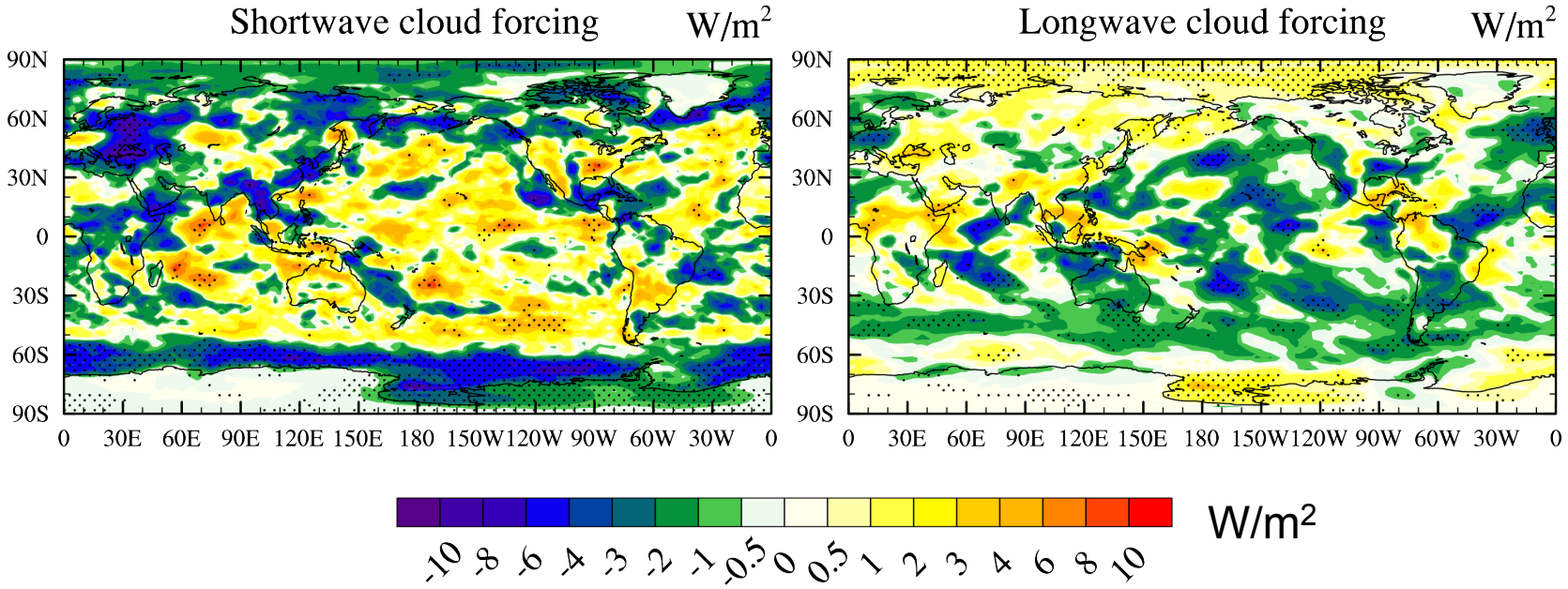
through slowing down the Bergeron process



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Difference in Cloud Forcing (single α - default)



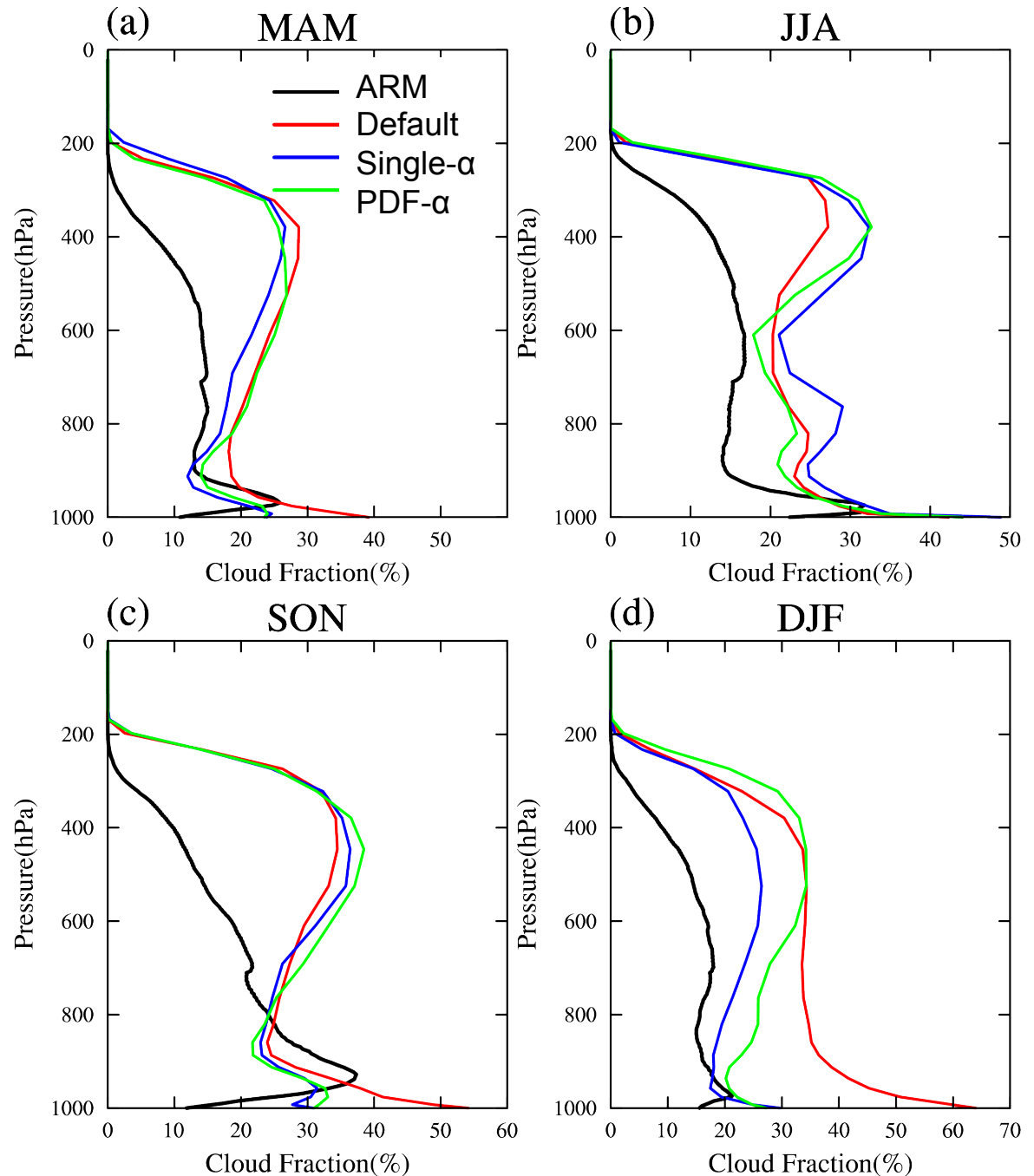
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Cloud cover at ARM Barrow

LWP -> Ts -> LTS -> LLC
-> RH -> MLC & HLC

Xie et al. (2013)



Summary

- ❑ A framework of ice nucleation based on classical nucleation theory has been implemented in CAM5, which links ice nucleation to aerosol (e.g., dust, soot, biological aerosol) properties
- ❑ Compared to Meyers et al. (default), new parameterization has a significant effect on cloud water and cloud cover, especially in high latitudes:
 - Improved comparison with long-term ARM observations

